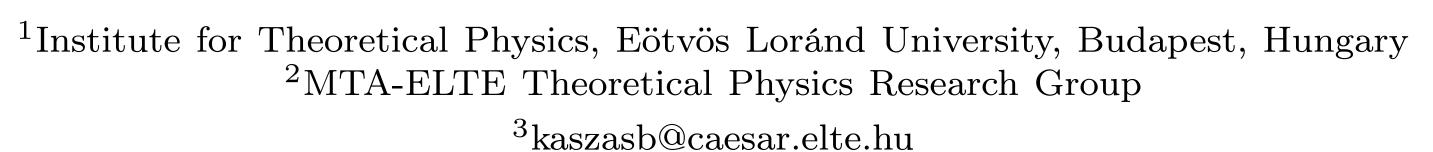
# Global cooling: the Snowball Earth transition in a climate model with drifting parameters

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## Motivation

The current configuration of our planet allows two stable climates

- A warm state, resembling today's climate
- An ice-covered Snowball Earth.

Drastic changes in solar irradiance (e. g. volcanic ash in the atmosphere) may trigger a transition to the Snowball state.

Conversely, the sudden melting can be induced by greenhouse gases.

The driving mechanism behind both the freezing and melting processes is the ice-albedo feedback.

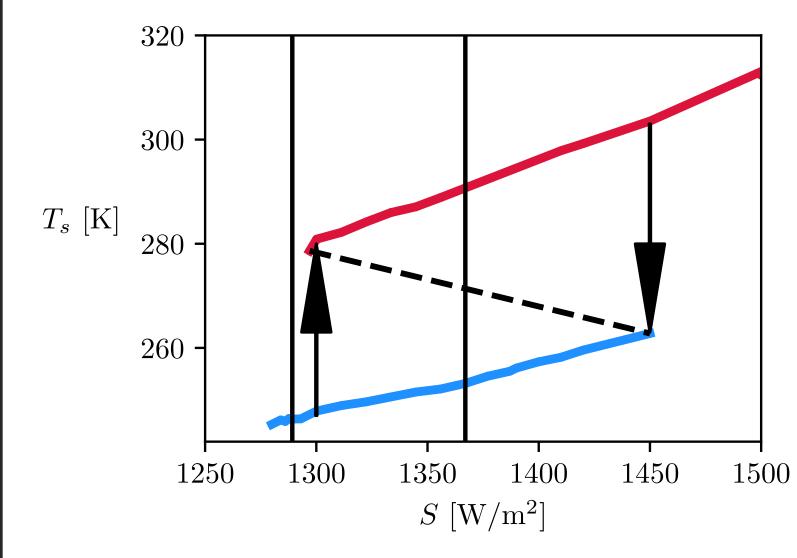
#### Plasim

The Snowball Earth transition is investigated in an intermediate-complexity climate model, PlaSim [1].

- T21 resolution, grid of  $5^{\circ} \times 5^{\circ}$
- Atmosphere is coupled to a mixed layer ocean
- A dynamical system with 10<sup>5</sup> degrees of freedom

#### Introduction

Our goal is to investigate the transition between the two stable states (attractors, from a dynamical point of view), that coexist for a wide range of the solar constant S.



- Two dynamically different stable states exist. These are displayed on the **bifurcation** diagram [2, 3].
- Between the attractors, an unstable edge state is embedded in the basin boundary (dashed lines)
- At  $S = S_{bif}$ , the warm state loses stability
- Assume a parameter drift scenario S(t)[see middle panel] for the solar constant

# References, Acknowledgements

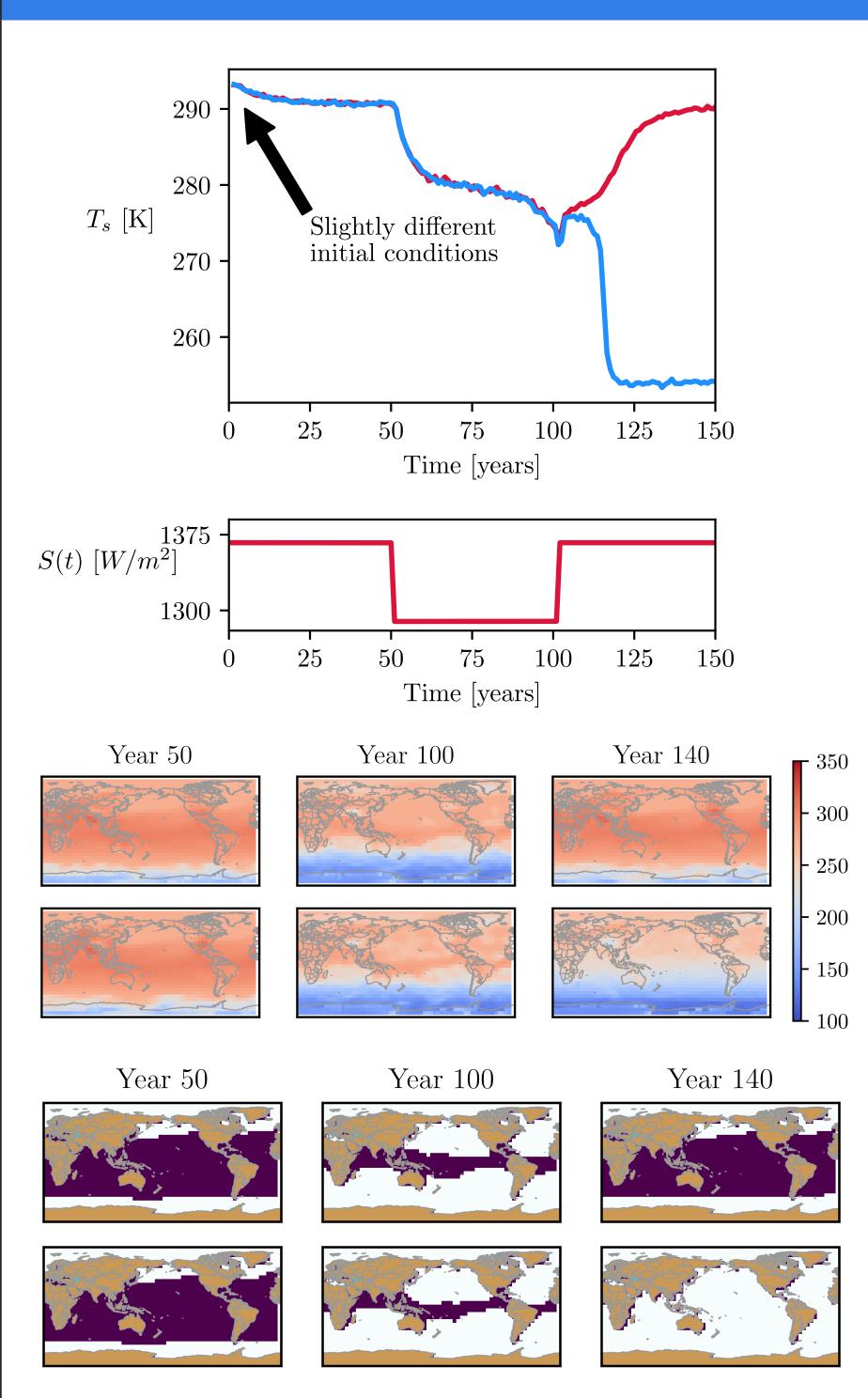
- [1] K. Fraedrich et al., Meteorol. Z. **14**:299–304 (2005)
- V. Lucarini et al., Q. J. Royal Meteorol. Soc. **136:**2–11 (2010).
- A. Tantet et al., Nonlinearity **31**:2221–2251 (2018)
- M. Herein et al., Sci. Rep. 7:44529 (2017).
- M. D. Chekroun et al., Physica D **240**: 1685–1700 (2011).



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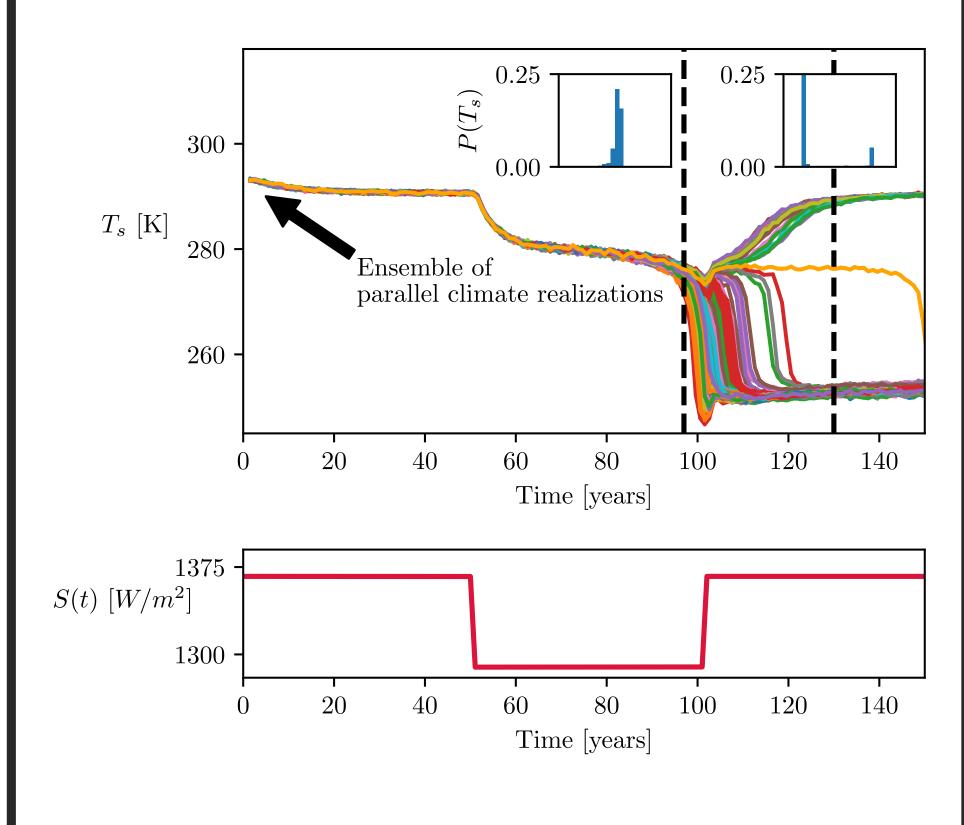
## Individual simulations



- Small difference in initial conditions  $\Rightarrow$ completely different outcomes
- Individual simulations are **not represen**tative of the transition
- There is no way to determine which attractor the system will end up on

#### Ensemble of climate realizations

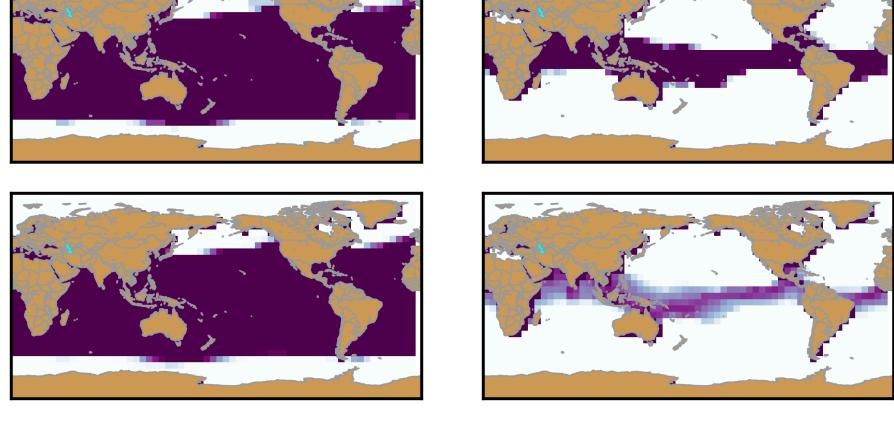
- Instead, follow an **ensemble** of parallel climate realizations [4]
- The 125 ensemble members converge to the snapshot attractor (or pullback attractor
- Statistical quantities can be computed: average, standard deviation, with respect to its natural measure

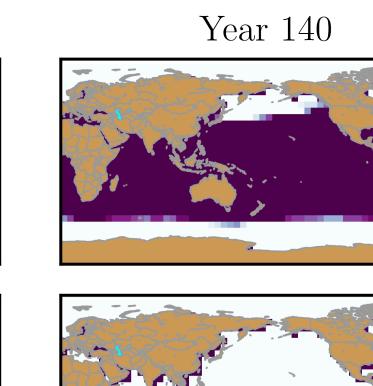


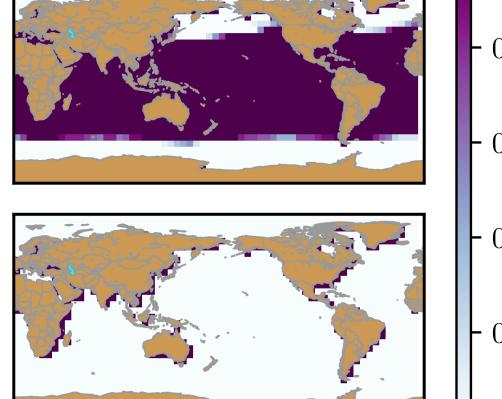
- The transition to Snowball Earth is **proba**bilistic, no prediction is possible based on a single simulation
- The snapshot attractor splits into two, resulting in two types of typical behavior
- Before the splitting, the standard deviation of the ensemble grows: precursor
- The orange curve runs near the time dependent edge state

### Typical sea ice cover ending in the two snapshot attractors

Year 100

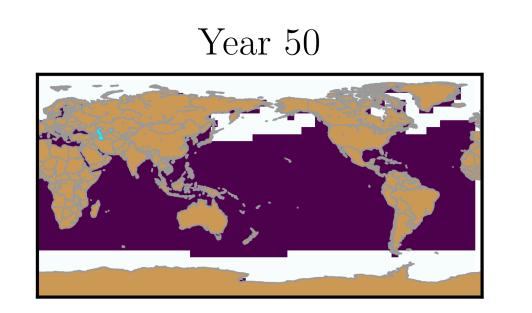




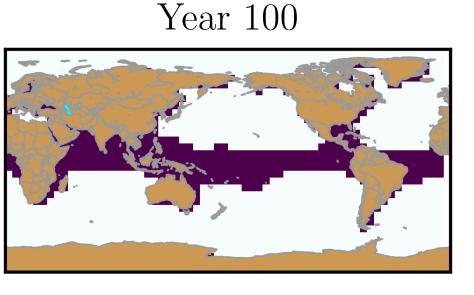


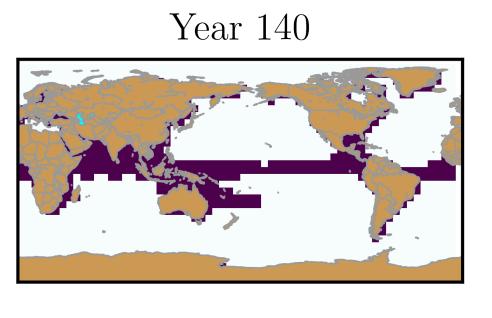
- The expected behavior (average) is computed over the two snapshot attractors
- Local probability of sea ice formation

# Sea ice cover in the edge state



Year 50





 Narrow band unfrozen sea remaining at the Equator

# Summary

- Individual trajectories are not meaningful
- Constructing the snapshot attractor with ensemble methods
- The snapshot attractor splits, which is preceded by increased standard deviation
- All three equilibria (two attractors and the edge state) are found

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