



INPE/CPTEC – NOAA Meeting on Weather and Climate Modeling

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CPTEC/INPE

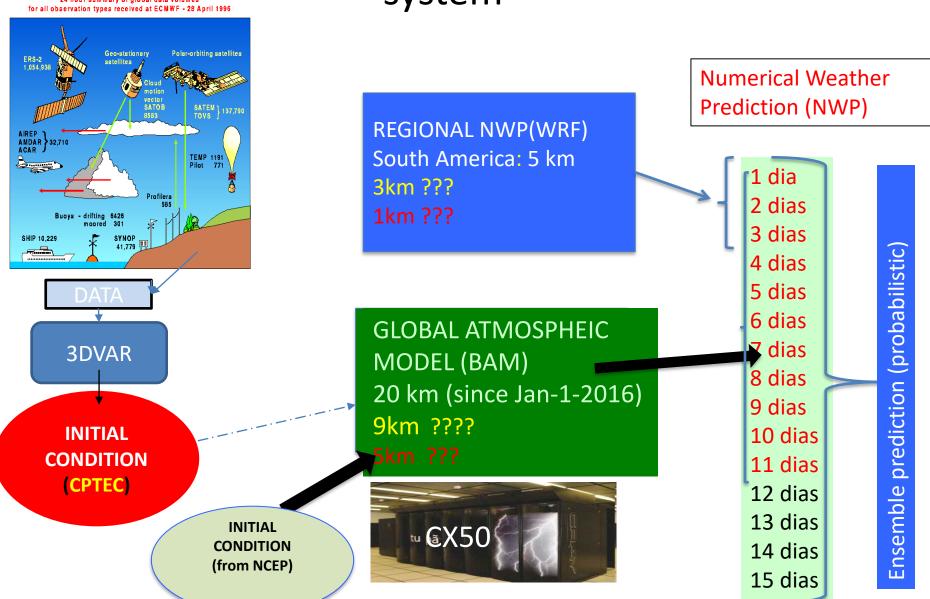
Part1. An overview on the CPTEC Global Weather and Climate Modeling System

Part2. CPTEC Global Atmospheric Modeling and frontier research topics in Tropical climate Modeling (Teleconnections, Tropical Convection, Andes, etc.)

We have three serious issues in our global weather and climate modeling

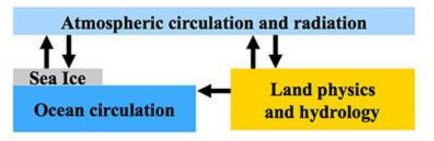
- We don't have our DAS operational yet
- Our global model has serious errors over Southern South America (dry bias)
- 3) We don't have our ESM yet.

CPTEC operational Regional and Global NWP system



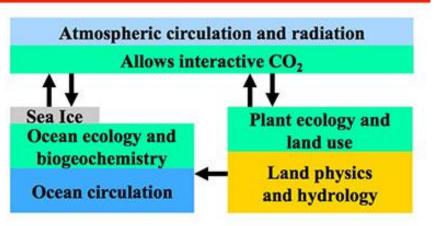
An Earth System Model (ESM) closes the carbon cycle

Climate Model

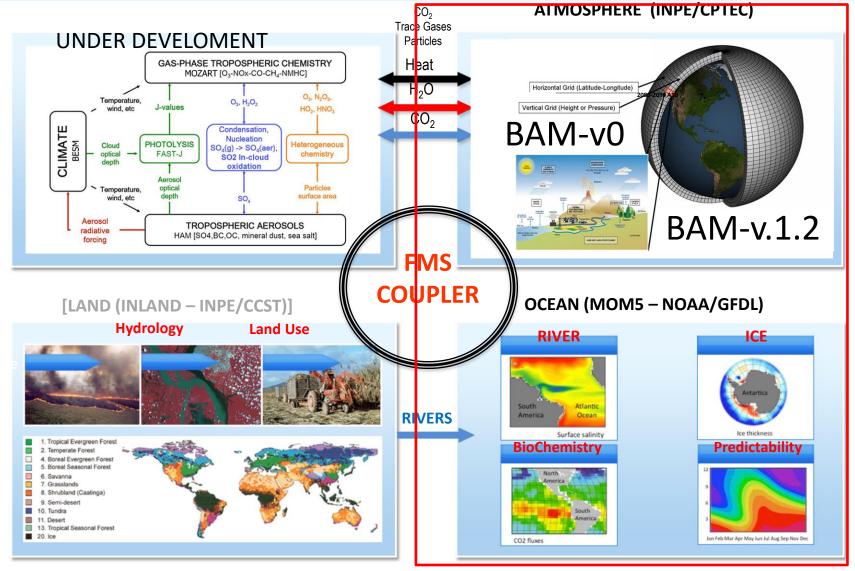


Earth System Model

ESM



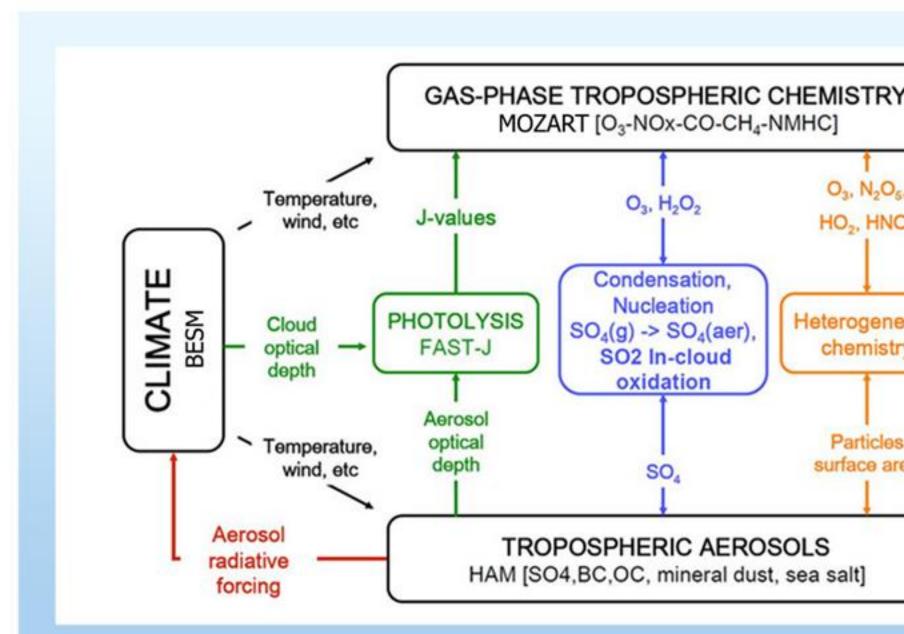
CPTEC COUPLE OCEAN-ATMOSPHERE MODEL Brazilian Earth System Model (BESM)now it is BESM_OA



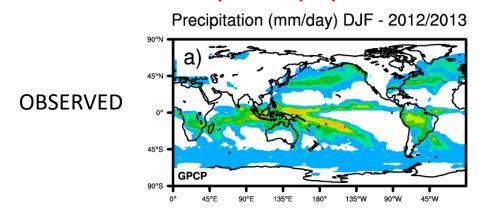
BESM_OA_2.5= BAM-vo+MOM5

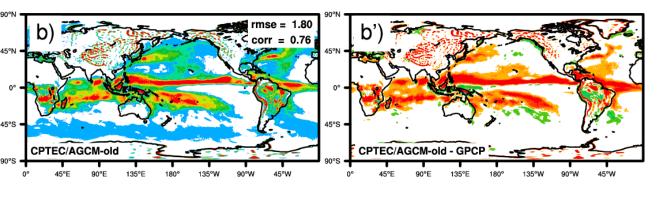
BESM_OA_2.8= BAM-v1+MOM5

Courtesy: Paulo Nobre

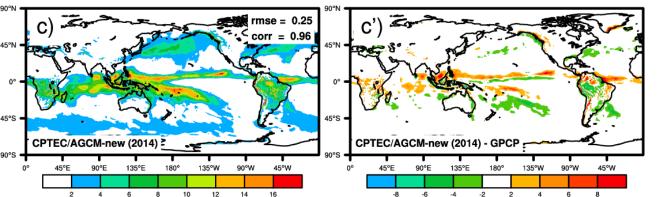


DEVELOPMENT OF THE BRAZILIAN GLOBAL ATMOSPHERIC MODEL (BAM) (details in Figueroa et. al 2016)



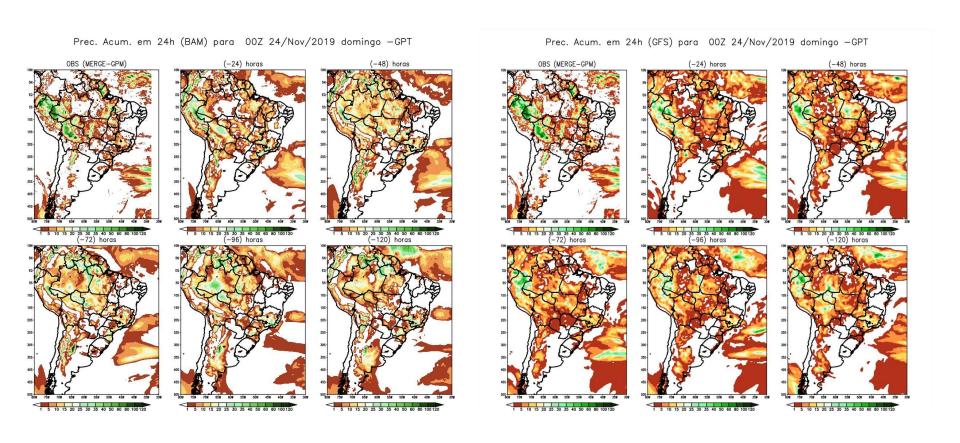


OLD CPTEC AGCM (originally came from COLA)



NEW CPTEC AGCM (BAM)
New Dynamic
New physical processes

Comparison between BAM-v1.2 (CPTEC) and GFS (NCEP) operational models



Is it possible to predict the South Atlantic Convergence Zone?

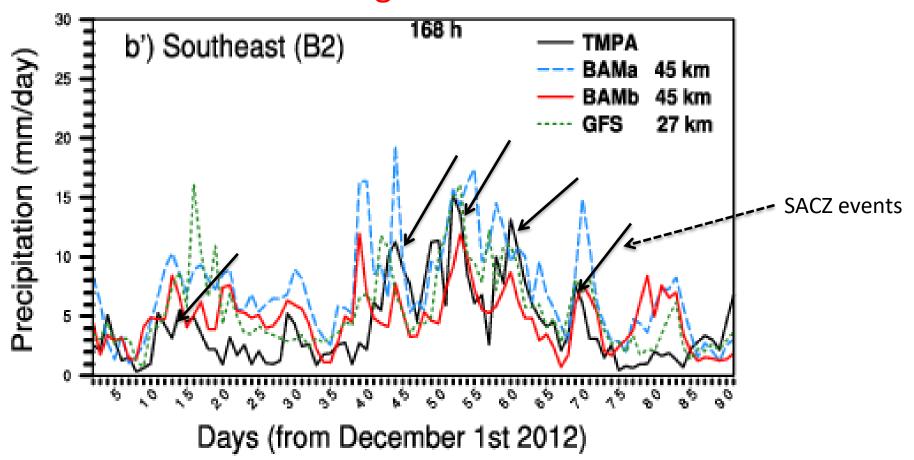


Fig. Daily mean precipitation for the period 01 December 2012 to 28 February 2013 from 168 hours (7 days) forecasts over the Southeast from TMPA and three NWP models indicated in the panel.

SOUTH ATLANTIC CONVERGENCE ZONE (SACZ)

1580

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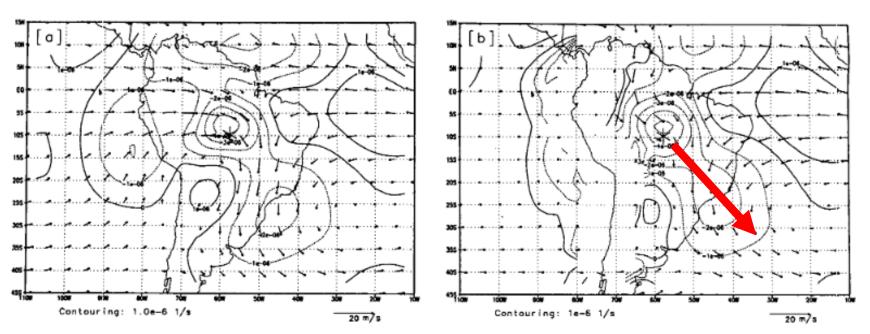
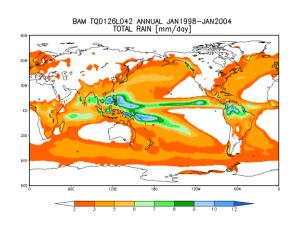
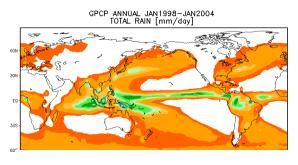


Fig. 10. Vector wind and divergence at 850 hPa in the cases with basic zonal flow: no-mountain case (experiment 3) (a) and mountain case (experiment 4) (b). Asterisk indicates the center of heat source.

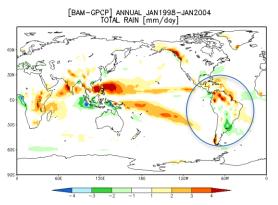
AMIP style simulation **CPTEC and UKMET Global Models**



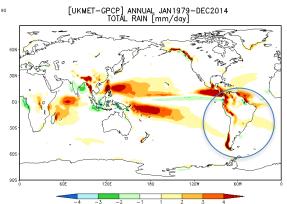
ANNUAL PRECIPITATION



UKMET ANNUAL JAN1998-JAN2004 TOTAL RAIN [mm/day]



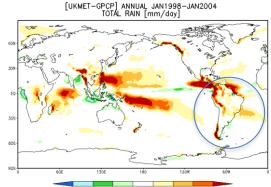
CPTEC: AMIP Style simulation (1998-2004)



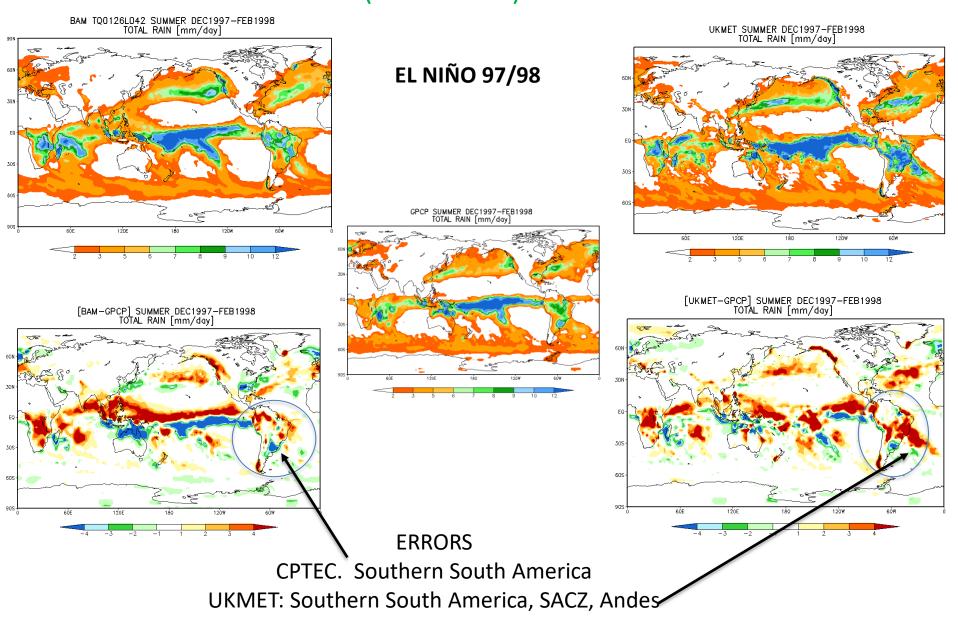
UKMET: AMIP simulation



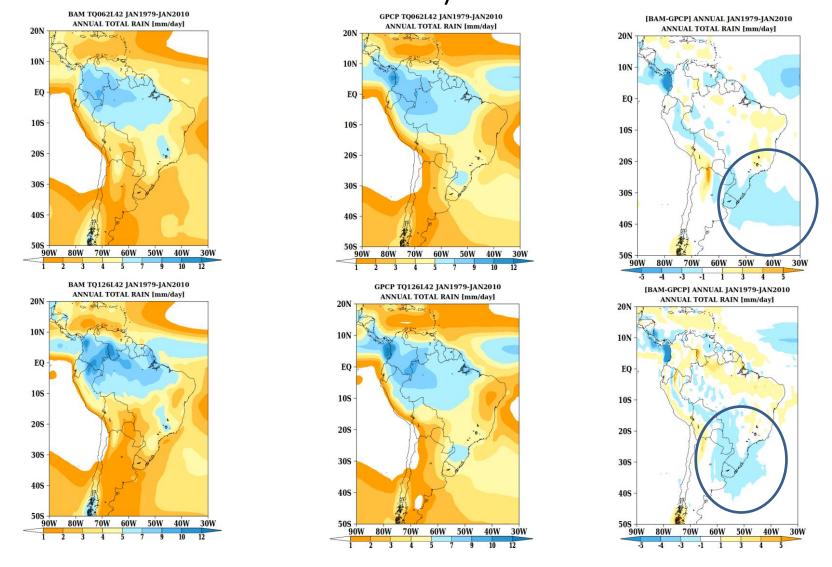
CPTEC. Southern South America **UKMET: SACZ, Andes**

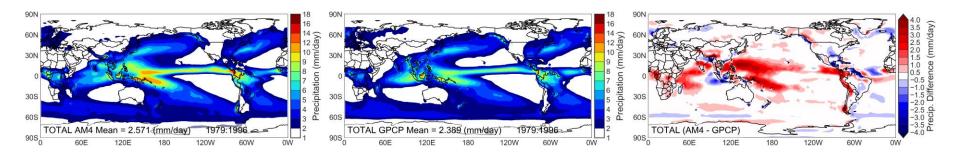


CPTEC:10-year long AMIP style (1995-2004), preliminary results. UKMET: AMIP simulation (1979-2014)

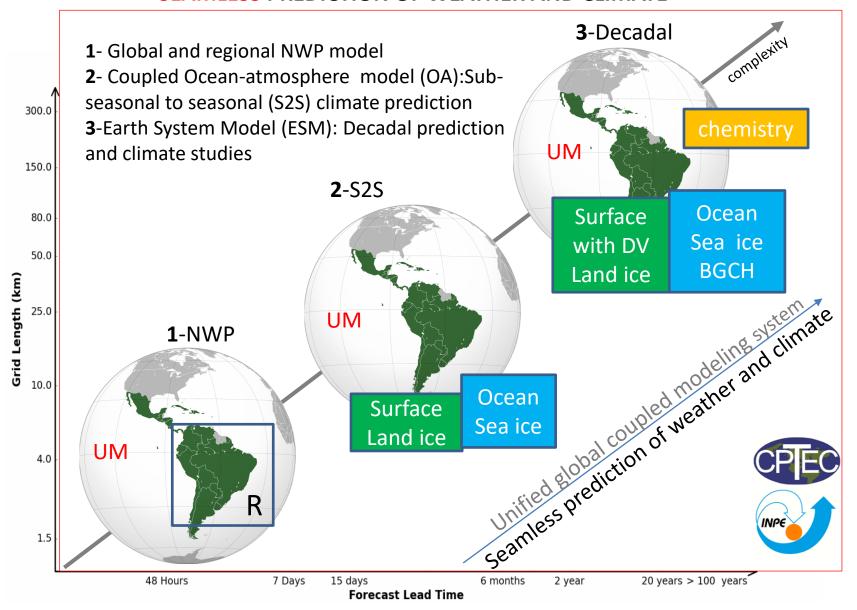


Our global model has serious errors over Southern South America (dry bias)





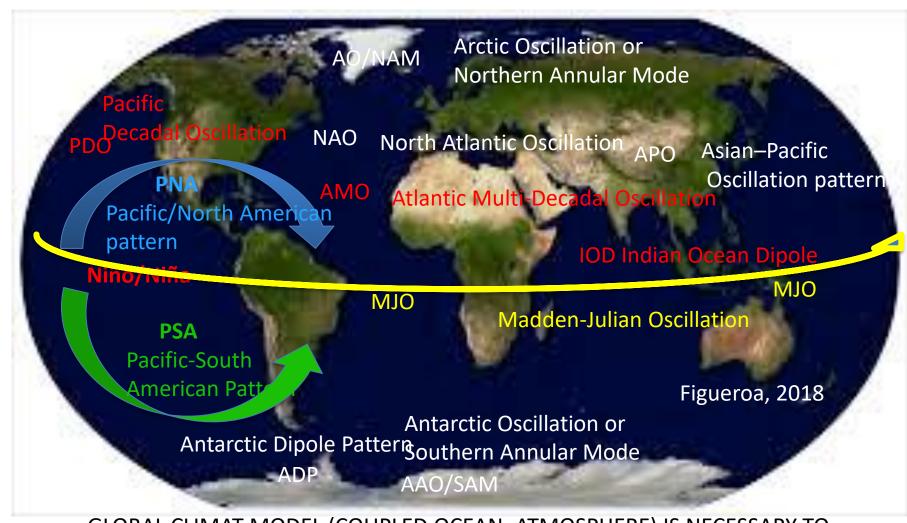
SEAMLESS PREDICTION OF WEATHER AND CLIMATE



Unified Model Global Atmosphere (UM-A)

DIFFERENT TELECONNECTIONS IN THE WORLD (spatial overview)

The most common atmospheric/oceanic oscillations and dipoles in the global system.



GLOBAL CLIMAT MODEL (COUPLED OCEAN_ATMOSPHERE) IS NECESSARY TO UNDERSTAND THE GLOBAL TELECONNECTIONS.

Benefits of the seamless approach Efficiency

Developing one system for multiple uses reduces development effort and allows improvements made for climate science to be applied in our weather forecasting systems, and vice-versa.

Understanding

Short-range forecasts can be used for learning about error growth and help study the performance of long term climate simulations. Likewise, studying the long-range climatology of the model helps constrain and understand the physical processes used in short-range NWP forecasts.

Robustness

Using the same model for regional and global modelling gives confidence that the driving mechanisms are consistent.

(from Ukmet).

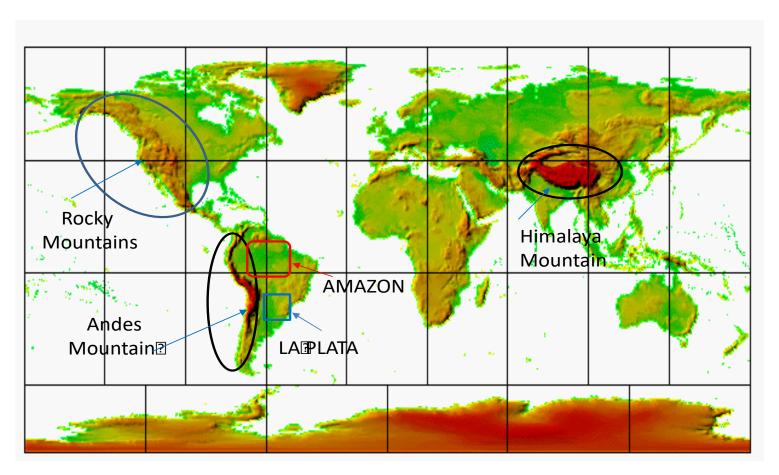


How can we represent realistically the Andes in the Climate Models?.

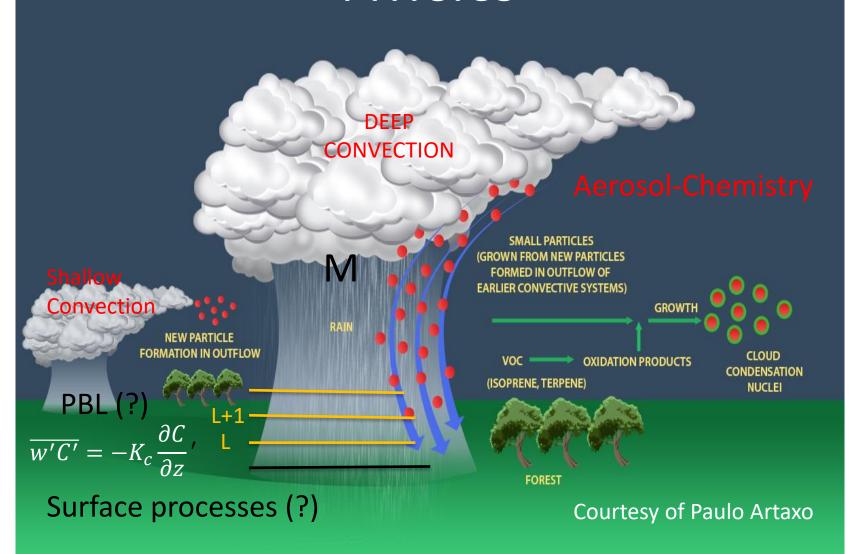
Horizontal dynamic Core Vertical Coordinate

DYNAMIC CORE FOR COMPLEX MOUNTAINS (eg. ANDES)

GLOBALTOPOGRAPHY



PHYSICS



Mass-flux parameterization approach

$$\varrho \frac{\partial a \chi_c}{\partial t} = -\frac{\partial (M_c \chi_c)}{\partial z} + E \chi_e - D \chi_c - \frac{\partial a \varrho(z) \overline{w_i \chi_i}^c}{\partial z} - \varrho \frac{\partial (1-a) \chi_e}{\partial t} = -\frac{\partial (M_c \chi_e)}{\partial z} - E \chi_e + \frac{D \chi_c - \frac{\partial (1-a) \varrho(z) \overline{w_i \chi_i}^e}{\partial z}}{\partial z}$$

$$\text{Mass Flux}$$

$$\varrho \frac{\partial a}{\partial t} = -\frac{\partial M_c}{\partial z} + E - D$$

 $\chi \epsilon \{s_t, q_t\}$

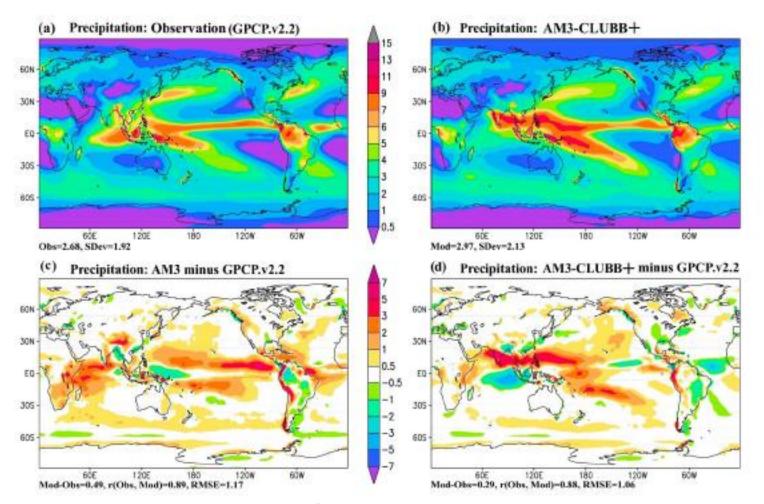
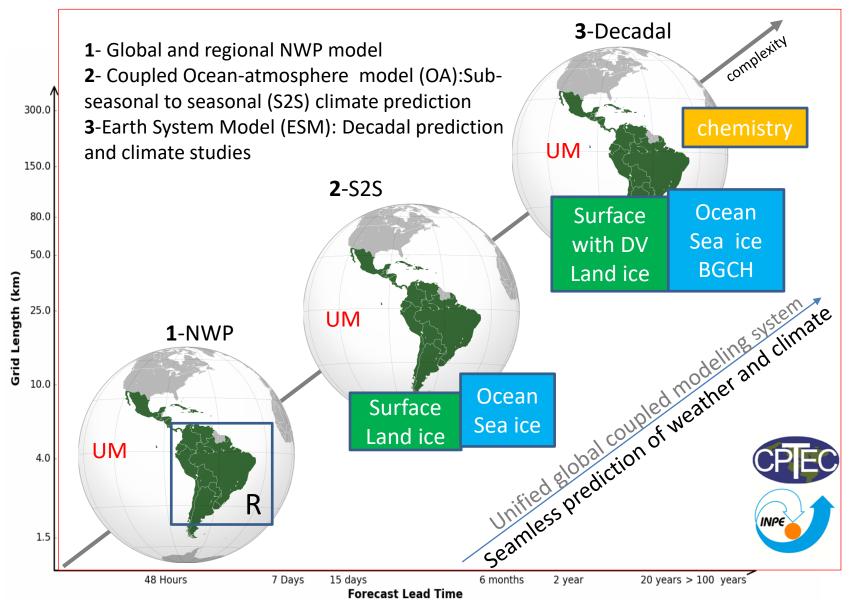


Figure 1. Annual mean precipitation rate (mm d⁻¹) from (a) the version 2 of the Global Precipitation Climatology Project (GPCP.v2.2) [*Adler et al.*, 2003], (b) AM3-CLUBB+, (c) AM3 model bias, and (d) AM3-CLUBB+ model bias, annual mean shortwave cloud forcing (SWCF) model bias from (e) AM3 and (f) AM3-CLUBB+, and longwave cloud forcing (LWCF) model bias from (g) AM3 and (h) AM3-CLUBB+.

SEAMLESS PREDICTION OF WEATHER AND CLIMATE



Unified Model Global Atmosphere (UM-A)

THANKS