# TPG4190 Seismic data acquisition and processing Lecture 20: Deconvolution

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## Geophysical Analysis - lecture 10

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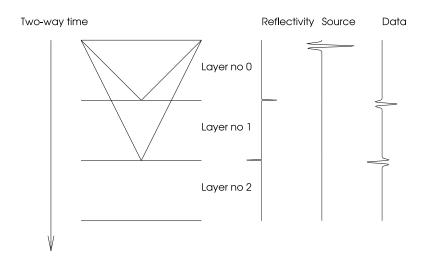


Figure: Layered earth model

vertical propagation

Reflected signal

$$a(t) = r_0 \frac{s(t - \phi_0)}{2d_0}. (1)$$

 $ightharpoonup \phi_0 = 2d_0/c_0$ : two-way traveltime

 $ightharpoonup r_0$ : reflection coefficient

 $ightharpoonup d_0$ : layer thickness

Reflection coefficient defined by

$$r_0 = \frac{c_1 \rho_1 - c_0 \rho_0}{c_1 \rho_1 + c_0 \rho_0}. (2)$$

- $ightharpoonup c_0, c_1$ : Wave velocity
- $ightharpoonup 
  ho_0, 
  ho_1$ : Density

► Signal recorded at the surface:

$$a(t) = r_0 \frac{s(t - \phi_0)}{2d_0} + r_1 \frac{s[t - (\phi_0 + \phi_1)]}{2d_0 + 2d_1},$$
(3)

- ► Neglect spherical spreading
- General model

$$a(t) = \sum_{k=0}^{N} r_k s(t - \tau_k), \tag{4}$$

 $au_k = \sum_{I=0}^k \phi_I$  two way traveltime to interface no k.

$$ightharpoonup au_k = k\Delta t.$$

$$a_k = \sum_{l=0}^{N} r_l s_{k-l}, (5)$$

$$a_k = a(t = k\Delta t),$$
  
 $s_k = s(t = k\Delta t).$  (6)

Continuous case:

$$a(t) = \int_{\tau=0}^{+\infty} r(\tau)s(t-\tau), \tag{7}$$

- Vertical wave propagation
- ► Neglect vertical spreading
- Only primary reflections
- ► Earth is a linear filter
- ightharpoonup r(t) Earth's impulse response

Convolutional model for

seismic waves

$$y(t) = r(t) * s(t), \tag{8}$$

- $\triangleright$  s(t): seismic source
- ightharpoonup r(t): reflectivity

- ightharpoonup Recover r(t) when
- ightharpoonup s(t) given
- ightharpoonup y(t) given

## Spiking deconvolution

► Fourier transformation of (8)

$$Y(f) = R(f)S(f) \tag{9}$$

Solution

$$R(f) = \frac{Y(f)}{S(f)}. (10)$$

► Equation (10) slightly reformulated

$$R(f) = S^{-1}(f)Y(f),$$
 (11)

- ►  $S^{-1} = 1/S(f)$ : inverse filter.
- $ightharpoonup S^{-1}$  removes the source
- recovers reflection coefficients

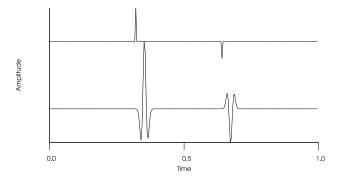


Figure: Spiking decon

The spectrum of a Ricker pulse and it's inverse

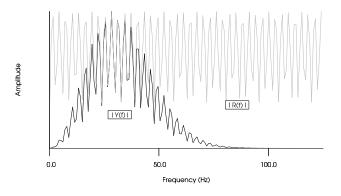


Figure: Deconvolution spectrum

The spectrum of the input data |G(f)| and the deconvolved output |R(f)|.

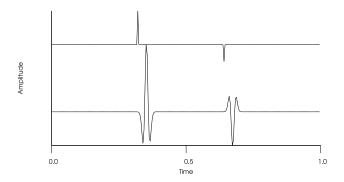


Figure: Deconvolution

► Real seismic data contains noise

$$Y(f) = R(f)S(f) + N(f),$$
 (12)

 $\triangleright$  N(f): White (random) noise

$$R(f) = S^{-1}(f)Y(f) - S^{-1}(f)N(f).$$
 (13)

 $S^{-1}(f)$  modified:

$$S^{-1}(f) = \frac{S^*(f)}{|S(f)|^2 + \epsilon},\tag{14}$$

where  $\epsilon$  is a small constant.

Inserting equation (14) into equation (13):

$$R(f) = \frac{S^*(f)}{|S(f)|^2 + \epsilon} Y(f) - \frac{S^*(f)}{|S(f)|^2 + \epsilon} N(f).$$
 (15)

Increase  $\epsilon$  to become much larger than  $|s(f)|^2$ 

$$R(f) \approx \left(\frac{1}{\epsilon}\right) S^*(f)[Y(f) - N(f)].$$
 (16)

match-filtering is a stable approach for deconvolution.

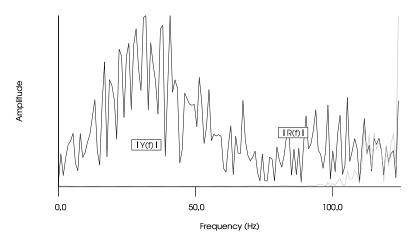


Figure: The spectrum of the input data with added noise |Y(f) + N(f)| and the deconvolved output |R(f)| (gray line).

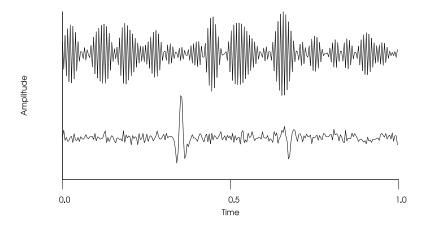


Figure: The input seismic data with added noise is shown in the lower trace, while the ouput of a spiking deconvolution filter applied to the input data is shown in the upper trace.

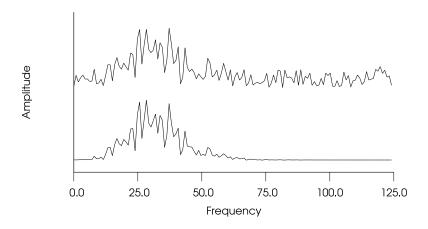


Figure: The spectrum of the input data with added noise |Y(f) + N(f)| and the stabilized deconvolved output |R(f)| (gray line).

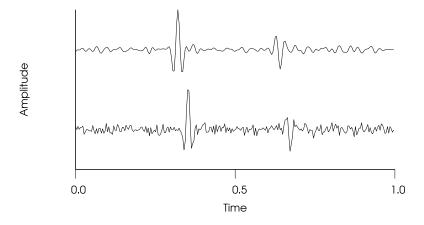


Figure: The input seismic data is shown in the lower trace, while the ouput of a match filter applied to the input data is shown in the upper trace.

The convolutional model for seismic data

$$y(t) = r(t) * s(t), \tag{17}$$

- $\triangleright$  y(t): Simulated Seismic data
- ightharpoonup r(t): Reflectivity
- ightharpoonup s(t): Seismic source pulse

y(t): Observed seismic data

We want to estimate r(t) by minimizing

$$e = \frac{1}{T} \int_{-\infty}^{+\infty} dt \, [\hat{y}(t) - y(t)]^2, \tag{18}$$

Sampled functions

$$e = \frac{1}{N} \sum_{k=0}^{N} [\hat{y}_k - y_k]^2.$$
 (19)

Inserting equation (5) into equation (19)

$$e = \frac{1}{N} \sum_{k=0}^{N} [\hat{y}_k - \sum_{l=0}^{N} r_l s_{k-l}]^2.$$
 (20)

Minimize error by

$$\frac{\partial e}{\partial r_m} = 0, \tag{21}$$

for m = 0, 1, 2, ..., N.

Applying to equation (20)

$$0 = \frac{-2}{N} \sum_{k=0}^{N} \left[ \hat{y}_k - \sum_{l=0}^{N} s_{k-l} r_l \right] s_{k-m}.$$
 (22)

After reorganizing and multiplication with -N/2

$$0 = \sum_{k=0}^{N} \hat{y}_k s_{k-m} - \sum_{k=0}^{N} \sum_{l=0}^{N} s_{k-l} s_{k-m} r_l,$$
 (23)

To simplify introduce the cross correlation

$$\phi_{\hat{y}s}(m) = \sum_{k=0}^{N} \hat{y}_k s_{k-m}, \tag{24}$$

and the autocorrelation

$$\phi_{ss}(m) = \sum_{k=0}^{N} s_k s_{k-m}, \tag{25}$$

Equation (23) then becomes

$$\phi_{\hat{y}s}(m) = \sum_{l=0}^{N} \sum_{k=0}^{N} s_{k-l} s_{k-m} r_{l}.$$
 (26)

Make a change of the summation variable k' = k - l which gives k = k' + l

$$\phi_{\hat{y}s}(m) = \sum_{l=0}^{N} \sum_{k'=-l}^{N-l} s_{k'} s_{k'-(m-l)} r_l.$$
 (27)

To get

$$\phi_{\hat{y}s}(m) = \sum_{l=0}^{N} \phi_{ss}(m-l)r_{l}, \qquad (28)$$

for m = 0, 1, ..., N. Equation (28) is known as Levinson's *normal* equations.

Solution of equation (28) corresponds to deconvolution in the frequency domain given by equation (14) with  $\epsilon=0$ .

$$\phi_{ab}(t) = \int_{-\infty}^{+\infty} d\tau \, a(\tau)b(\tau - t), \tag{29}$$

which is the same as

$$\phi_{ab}(t) = \int_{-\infty}^{+\infty} d\tau \, a(\tau + t) b(\tau). \tag{30}$$

The last equation is proved by substitution of variables by  $u = \tau - t$ .

The auto-correlation becomes

$$\phi_{aa}(t) = \int_{-\infty}^{+\infty} d\tau \, a(\tau) a(\tau + t). \tag{31}$$

The Fourier transform of equation (30) is

$$\Phi_{ab}(f) = \int_{-\infty}^{+\infty} dt \int_{-\infty}^{+\infty} d\tau \, \exp(-2\pi i f t) a(\tau + t) b(\tau), \quad (32)$$

which becomes after a change of variable  $u = \tau + t$ :

$$\Phi_{aa}(f) \int_{-\infty}^{+\infty} d\tau \left[ \exp(2\pi i f \tau) b(\tau) \right]^* \int_{-\infty}^{+\infty} du \, \exp(-2\pi i f u) a(u), (33)$$

which is the same as

$$\Phi_{ab}(f) = A(f)B^*(f). \tag{34}$$

If a = b we get

$$\Phi_{aa}(f) = A(f)A^*(f) = |A(f)|^2.$$
 (35)

The Fourier transform of the auto-correlation is the square of the amplitude spectrum, which is also called the power spectrum.

Fourier-transform equation (28) and use equations (34) and (35):

$$\hat{Y}S^*(f) = |S(f)|^2 R(f),$$
 (36)

Solving equation (36) for R(f) we get

$$R(f) = \frac{Y(f)S^{*}(f)}{|S(f)|^{2}}$$
 (37)

which is exactly the same solution as given by equation (19) but with  $\epsilon = 0$ .

To stabilize the solution add a damping-constant  $\epsilon$  in the error function given by equation (19)

$$e = \frac{1}{N} \sum_{k=0}^{N} [\hat{y}_k - y_k]^2 + \epsilon \sum_{k=0}^{N} r_k^2.$$
 (38)

This gives the normal equations

$$\phi_{\hat{y}s}(m) = \sum_{l=0}^{N} [\phi_{ss}(m-l) + \epsilon \delta_{m,l}] r_l, \qquad (39)$$

where the Kronecker symbol  $\delta_{m,l}$  is defined as

$$\delta_{m,l} = \begin{cases} 1 & m = l \\ 0 & \text{otherwise} \end{cases} \tag{40}$$

A Fourier-transform of equation (39) gives exactly the stabilized deconvolution solution in equation (37).

▶ Predict the input  $y_k$  at future time  $y_{k+\alpha}$ .

$$y_{k+\alpha} = \sum_{l=0}^{M} p_l y_{k-l}.$$
 (41)

- ▶ Use Wiener filtering
- ► Input:  $y_k$
- ▶ Desired output:  $y_{k+\alpha}$ .

▶ Right hand side crosscorrelation:

$$\phi_{y_{k+\alpha}y_k}(m) = \sum_{l=m}^{M} y_{l+\alpha}y_{l-m}.$$
 (42)

▶ Change of variable  $I' = I + \alpha$ :

$$\phi_{y_{k+\alpha}y_k}(m) = \sum_{l'=\alpha+m}^{M} y_{l'} y_{l'-(m+\alpha)} = \phi_{yy}(m+\alpha).$$
 (43)

► The normal equations then become

$$\sum_{l=0}^{M} \phi_{yy}(l-m)p_l = \phi_{yy}(m+\alpha), \tag{44}$$

for m = 0, 1, ..., M.

Prediction error

$$\epsilon_k = y_{k+\alpha} - \hat{y}_{k+\alpha}. \tag{45}$$

- Prediction error represents non-predictable part of input data
- If input data is stationary, white and random with zero mean

$$\phi_{xx}(m) = 0$$
, for  $m \neq 0$   
 $\phi_{xx}(m) = \sigma^2$ . for  $m = 0$  (46)

- Right hand side is zero for nonzero valuesof the prediction distance
- ► Filter coefficients zero
- Predicted value zero, which is mean of input data

## Inverse filters for multiples

▶ Replace the predicted value with equation (41)

$$\epsilon_{k+\alpha} = y_{k+\alpha} - \sum_{l=0}^{N} p_l y_{k-l}. \tag{47}$$

> z-transform of equation (45)

$$A(z) = \sum_{k=0}^{N} a_k z^k \tag{48}$$

$$\sum_{k=0}^{N} z^{k} \epsilon_{k+\alpha} = \sum_{k=0}^{N} z^{k} y_{k+\alpha} - \sum_{l=0}^{N} z^{k} \sum_{l=0}^{N} p_{l} y_{k-l}, \tag{49}$$

$$z^{-\alpha} \sum_{m=k+\alpha}^{N+\alpha} z^m \epsilon_m = z^{-\alpha} \sum_{m=k+\alpha}^{N+\alpha} z^m y_m - \sum_{j=0}^{N} z^k \sum_{l=0}^{N} p_l y_{k-l}.$$
 (50)

$$z^{-\alpha}E(z) = z^{-\alpha}Y(z) - P(z)Y(z). \tag{51}$$

$$E(z) = [1 - z^{\alpha} P(z)] Y(z).$$
 (52)

► The filter

$$[1-z^{\alpha}P(z)], \qquad (53)$$

► Z-transform of the time-domain filter

$$p_k = 1, 0, 0, 0, 0, \dots, 0, -p_0, -p_1, \dots, -p_N,$$
 (54)

- ▶ The filter contains  $\alpha 1$  zeros.
- For  $\alpha = 1$

$$p_k = 1, -p_0, -p_1, \dots, -p_N,$$
 (55)

► Statistical Wiener spiking deconvolution filter.

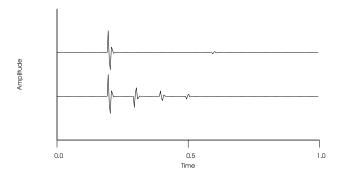


Figure: Input seismic data with multiples (lower trace) and output from multiple inverse filter (top)

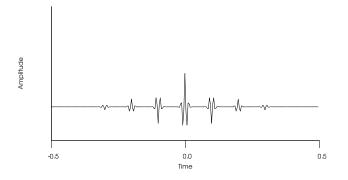


Figure: Autocorrelation of the input data shown in figure 9

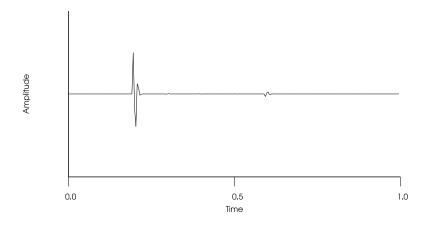


Figure: Prediction error filter applied to the input data shown in figure 9

