Bulletin of Geography - physical geography series No 3/2010: 121-171

KRYSTYNA BRYŚ, TADEUSZ BRYŚ

Department of Agro- and Hydrometeorology, Wrocław University of Environmental and Life Sciences, Plac Grunwaldzki 24, 50–363 Wrocław, Poland krystyna.brys@up.wroc.pl

RECONSTRUCTION OF THE 217-YEAR (1791–2007) WROCŁAW AIR TEMPERATURE AND PRECIPITATION SERIES

Abstract: The authors present the results of a climatologic reconstruction of the 217-year series (1791–2007) of monthly average values of air temperature and monthly totals of precipitation in Wrocław (Breslau). The reconstruction is based on verified measurement (or observation) data from the Breslau-Sternwarte Observatory (1791–1930), which were completed (1931–2007) by reconstructing values calculated from measurement data from other meteorological stations in Wrocław. Only the data for 1945 were estimated from the values of Polish (Cracow, Gniezno, Puławy) and foreign (Berlin, Prague) stations. The problem of restoring the homogeneity of the available initial data was taken into consideration because there were changes in location and observation times. The data were verified by comparing the measured values, among others of temperature, with the data from the nearest as well as other meteorological stations in Warsaw, Berlin, Prague, Vienna, de Bilt. In relation to this, the problem of special and temporal anisotropy of relations between the values of analysed stations was discussed.

Key words: air temperature series, anisotropy, homogeneity, metadata, precipitation series, reconstruction

Introduction

The city of Wrocław (Breslau) boasts one of the earliest histories of permanent measurements of air temperature and other meteorological elements such as precipitation. The first hundred years (1791–1890) of the "Wrocław air temperature series" has been presented in a separate study against an historical background of meteorological measurements in other Silesian stations (Bryś and Bryś 2010). This paper not only continues this reconstruction for the next 117 years up to 2007, but it also presents the reconstruction of the 217-year period (1791–2007) of the "Wrocław precipitation series" conducted on the basis of estimated and measured values of monthly precipitation totals.

Unlike continuous air temperature measurements in Wrocław, which had begun in February 1791, measurements of precipitation at the Breslau-Sternwarte Observatory did not begin until January 1799 (Galle 1857). Earlier, from March 1791, observations of the frequency and types of precipitation were made. Precipitation measurements were conducted on the terrace of the Mathematical Tower of the University of Wrocław and continued until the 1860's. In the second half of the century the measurements were transferred outside the Observatory building. Other important measurement changes frequently occurred in later years. As was the case with measurements of air temperature Tp, changes in the location of measurements also had an influence on the nonhomogeneity of values for precipitation P in Wrocław. Thus, the problem of restoring the homogeneity of the Tp and P series is crucial and is closely linked with the methods applied to their credible reconstruction

Materials and metadata: Historical observations

The analyzed *Tp* and *P* data from the years 1791–2007 were collected from different sources. The data from diurnal measurements of these elements in the Breslau-Sternwarte Observatory (1791–1930) are of major significance for our work. The extant data originate from other meteorological (or precipitation) stations located in different districts of Wrocław (Gądów, Strachowice, Biskupin, Swojec, Psie Pole) and were collected in different years. Only the analyzed values for 1945 were estimated from the values of the other stations in Poland (Cracow, Gniezno, Puławy) and abroad (Berlin, Prague). The more than 200 years of changes in location and other criteria

in *Tp* measurements in Wrocław have been presented in separate studies (Pyka 2003; Bryś and Bryś 2010). A similar history of the primary changes in location in Wrocław measurements of *P* requires closer examination in this study as it is linked with a separate history of changes in methods of measurement.

In particular, historical events in the twentieth century were very important for the reconstruction of the *Tp* and *P* (and also the earlier history for *P*) analyzed in this paper. The framework and scale of the available and analyzed materials from different Wrocław stations were connected with the time limits of these changes. The Breslau-Sternwarte Observatory functioned as Wrocław's main station from 1848 to 1920, first within the Prussian and then the German meteorological network. Its operations were taken over from January 1921 to July 1936 by a new station, Breslau-Krietern (in Polish: Wrocław-Krzyki), located in the suburban part of the town, at a distance of about 5.5 km SSW from the previous station. This station, specializing in geophysical measurements, continued meteorological observations until April 1945, at which time it was burned and partly destroyed during World War II.

From August 1936 until 1964, with only 18 months interruption, the basic measurements for Wrocław for the national meteorological network (Polish from 1946) were recorded at the Wrocław-Gądów (in German: Breslau-Gandau)¹ airport station. (Between April 1944 and December 1945 data came from the Breslau-Krietern station.) The Wrocław-Gądów station was located about 6 km WNW from the old university observatory.

Meteorological observations in Breslau-Sternwarte finished at the end of March 1921, but air temperature measurements were probably taken at least until 1930. A new University meteorological station (*Ergebnisse* 1921–1924) began work on 1 April 1921, at the eastern end of Szczytnicki Park (Scheitniger Park), in a grass garden by the new Astronomical Observatory (about 3.5 km east of the current day building) of the University of Wrocław. The station was also given the name Breslau-Sternwarte (sometimes Neue Sternwarte), and the data collected in the form of monthly averages (or totals) from the years 1934–1936 were published in the annals of the German meteorological service (*Deutsches Meteorologisches Jahrbuch* 1934–1936).

¹ The German aerodrome station in 1935–1944 included two autonomous stations: one synoptic and one aerologic (see: next section)

The monthly average values of air temperature from the years 1851–1930 and monthly precipitation totals from 1881–1930 for Breslau-Sternwarte were published in *Klimakunde*...(1939)². The remaining German data of the *Tp* and *P* from the years 1891–1944 used in this paper were published in *Ergebnisse* (1891–1933) and *Deutsches Meteorologisches Jahrbuch* (1934–1944) or in *Atlas Klimatyczny* (1931–1944). Original measurement materials from the German stations mentioned which had been left in various places, probably did not survive either military operations during World War II or subsequent fires and looting in the first post-war months (Pyka 2003; Bryś and Bryś 2010).

From 1887 until 1944, observations were made at: 7.00, 14.00, and 21.00 local time. After World War II, from 1946 until 1965, observation times at Wrocław-Gądów were changed from 14.00 to 13.00 (12.00 UTC). From 1966 meteorological observations, which had continued since 1965 at the Wrocław-Strachowcie airport (about 10 km W of the first Breslau -Sternwarte station), were made at 8 fixed synoptic times: 3.00, 6.00, 9.00, 12.00, 15.00, 18.00, 21.00, 24.00 UTC.³ The remaining Wrocław meteorological stations (Biskupin, Swojec) used the climatologic times 6.00, 12.00, and 20.00 UTC until 1978 (Swojec) and 1979 (Biskupin) and next: 0.00, 6.00, 12.00, and 18.00 UTC (at 0^h – reading of thermohigrograms; at the remaining times – values of standard measurements).

The Polish *Tp* and *P* data from Gądów (1946–1964) and Strachowice (1960–2007 for *P* and 1963, 1965–2007 for *Tp*), which was used in this paper, originated in several sources (*Atlas Klimatyczny Polski; Biuletyn Agrometeorologiczny; Rocznik Meteorologiczny*). The temperature and precipitation data from about a dozen other Polish stations which used the Wrocław data for estimation or verification, also originated from the abovementioned sources. The longest Polish homogeneity *Tp* data from Warsaw (1779–1998) and Cracow (1792–1995), used here for comparative purposes,

² However, here we are not certain whether the results of the activities of the old Observatory from the years 1921–1930, which were published in *Klimakunde* (1939) in the form of mean monthly air temperature values and monthly precipitation totals, are solely measurement data or are partly estimated data.

³ A military airport Strachowice was opened at the beginning of 1938. In the period from June 1945 to December 1946 and from 1958 up to the present there has also been a civil airport at Wrocław-Strachowice. Probably the first meteorological measurements were made there in the 1930s, but the earliest temperature data (NOAA data resources) only date from 1963.

originated from works by Lorenc (2000) and Trepińska (1997). These values were supplemented up to the year 2007 on the basis of temperature data from the Polish meteorological sources mentioned earlier. Other historical data (de Bilt, Berlin, Vienna, Prague), used for the same purpose in this paper, were published by Wetterzentrale (http://www.wetterzentrale.de/klima/). They were completed by the authors of this paper up to the year 2007 on the basis of Tp data from Internet sources of other national meteorological services (KNMI, DWD, CHMU, ZAMG). The data for de Bilt was corrected according to Engelen and Nellestijn, and Errenwijlens (http://home.casema. nl/errenwijlens/co2/homogen.htm) and the data for Praha-Ruzyne was enriched and verified by the Tp data from Praha-Klementinum (for which we thank Rudolf Brázdil). Additionally, temperature data from the NOAA Internet base of meteorological data (ftp://ftp.ncdc.noaa.gov/pub/data/gsod/) was used for verification of the Wrocław series.

Precipitation measurements conducted first at a height of 33.5 m above ground level on the Observatory terrace of the University of Wrocław were transferred to the University courtyard in 1854. Measurements of *P* (at a height of 2.0 m above ground level) were conducted at the new location for only 4 years. In the summer of 1858, measurements were transferred to the nearby Wrocław University Botanical Garden, about 800 m east of the Mathematical Tower (measurements at a height of 1.5 m AGL). Simultaneously, for comparative purposes, measurements were continued until 1868 on the terrace of the Mathematical Tower. The measurements showed a difference of about 30% between the annual totals of *P* at the terrace and in the Botanical Garden. Galle (1879), who was interested in the homogeneity of these data, also precisely described the rain-gauge recorders in use at that time and the history of their changes in the Observatory.

The precipitation measurements taken in the Botanical Garden were not interrupted when meteorological observations were transferred from Breslau-Sternwarte to other stations in Wrocław (Fig. 1). As a precipitation station (or post), precipitation measurements at the Botanical Garden survived until the beginning of the 1960s (with several years of interruption during and just after the war). However, only part of the data was published in the precipitation annals of German or Polish meteorological service. It is worth noting that there were other precipitation stations (or posts) operating in Wrocław apart from this station, but most of them (both German and Polish) functioned for only a few years. Only two – Stabłowice (Breslau-

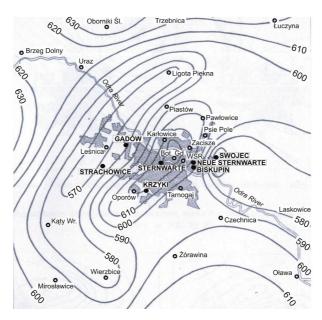


Fig. 1. The location of meteorological and precipitation stations and spatial distribution of average annual totals of precipitation (for the years1957–1966) in Wrocław and its neighbourhood. The isopleths show the effects of urban precipitation shade in relation to the prevailing wind direction from WNW-NW (after Schmuck 1967)

Stablowitz) in the western part of the city (near Leśnica) and Psie Pole (Breslau-Hunzfeld) on its north-eastern outskirts – worked for a longer time (probably starting earlier than 1920, but published data are only available from 1920 to 1940).

The Meteorological Observatory of the University of Wrocław in Wrocław-Biskupin has been operating since February 1946, and is located in Szczytnicki Park about 400 m SE from the previous station (Neue Sternwarte). Basic observation data obtained there from the years 1946–1965 and 1970–1974 were published and elaborated in annual reports (*Biuletyn Meteorologiczny; Prace* ...). Together with other data from the years 1881–2000, they were useful for the reconstructions of long-term runs of monthly air temperature averages and precipitation totals (Pyka 1991, 1998a, b), and also for the studies of the climate of Wrocław (Dubicka 1994; Dubicka, Pyka 2001). Thanks to Krzysztof Migała from the University of Wrocław, the authors for the purpose of this paper were able to supplement analyzed

data of monthly values of Tp and P for the years 1946–2007 with unpublished data (1996–2007) from the University Observatory.

The history of agrometeorological measurements began in Wrocław as early as the mid-1930s, including, among others, the measurements of air temperature and precipitation. Measurements were made in the garden of the experimental station of the Agricultural Institute of the University of Wrocław, located in a suburban, agricultural district – Wrocław-Swojec – east of the city. Pre-war results were never published. After an interval during World War II, measurements were not taken until May 1947 at the Agricultural Institute, which was the first part of the University of Wrocław and then at the Higher School of Agriculture (HAS) in Wrocław. A specialized Observatory of Agro- and Hydrometeorological was established in 1962 at the HAS (later known as the Agricultural University of Wrocław and now the Wrocław University of Environmental and Life Sciences). The scale of observations and measurements there has been systematically extended since 1960. However, the data for Tp and P for the period June 1947 – December 1959 from the station were not complete and revealed numerous missing fragments (about 55% for Tp and 15% for P). The data for the years 1960–2007, on the other hand, were complete.

Reconstruction difficulties and methodological dilemmas

Previous reconstructions of long-term temperature and precipitation courses in Wrocław (Pyka 1991, 1998a, b; Pyka and Dubicka 2001) did not exceed, or exceeded by just a few years, a 100-year period. The greatest difficulty in the reconstructions was to join together different series of measurements due to location changes of the meteorological stations mentioned above and the difficulty of filling the measurement gap for the year 1945.

The reconstruction of the *Tp* made by Pyka (1991, 1998a), covering the years 1881–1995, made dubious connections between post-war data from the Wrocław-Biskupin Observatory and earlier German measurements. This procedure involved assuming the similarity of meteorological data from the nearby but much older Neue Sternwarte station with university data obtained after World War II from the Biskupin station. While this might seem to be feasible, a more precise examination points to the contrary. This analysis became the basis for the revision of Pyka's thesis, presented later in this study.

Serious doubts about the homogeneity of Pyka's reconstruction of air temperature series arose from local on-site observations made by the authors, who analyzed differences in the location of these two stations. Differentiations in various topoclimatic conditions in both locations are clearly visible. The Neue Sternwarte area represents specific, local climatic conditions, formed by the eastern border, primarily consisting of the orchards of the Szczytnicki Park. The Park covered with tall, old trees that provide not only a very specific precipitation shade but also form a distinct and milder temperature structure within the Park and in the garden of the Astronomical Observatory as well. Topoclimatic conditions differ therefore from those observed in the nearby area of low buildings, where the Wrocław-Biskupin station is located. This difference has a significant influence on the measurements of temperature and precipitation coming from the two stations, though they are not extremely different. One cannot therefore assume, without making the appropriate adaptative corrections, that the results from these stations, as conjoined by Pyka (as a simple compilation), constitute a representative homogeneous series.

The comparison of *Tp* from the Wrocław-Gądów station with temperature values from the Wrocław-Biskupin station also gave rise to doubts undermining Pyka's assumptions. The average temperature in the years 1946–1964 reached 8.4°C for Gądów and 8.1°C for Biskupin. In preparing his series, Pyka (1998a) lowered (on an average of about 0.4 °C) the pre-war data from Gądów and the previous data from Krzyki and the Astronomical Tower as well. Yet results from a comparison of *Tp* from Gądów (and also Strachowice) with Biskupin indicate a different set of dependencies. Therefore, Pyka's homogeneity trial is based on two false assumptions.

The first assumes that air temperatures at the Neue-Sternwarte station are identical with the *Tp* for the Biskupin station. The second assumes that the location of the post-war airport station in Wrocław-Gądów was exactly the same as that of the pre-war climatology station. The situation in the second case is more complex. All the most important Polish meteorological and climatology sources (for example: *Atlas...*) also identify the post-war station with the pre-war one. However, a precise verification of available data from meteorological annals (*Ergebnisse*, *Rocznik Meteorologiczny*) and the cartographical analysis (Fig. 2) of different archival sources indicate that there were two meteorological stations at the Gądów airfield before the war: 1) the synoptic station (16° 59'E and H=116 m alt.), 2) the aeorological-



Fig. 2. The most likely location of the two meteorological stations on the Gądów (Gandau) airfield (Flughafen) before the second World War: C – climatology station, S – synoptic station

climatology station (16°58'E and H=118 m alt.). The first was probably on the future site of the Polish station (1946–1964). The average daily temperatures (recorded at 8 synoptic times) for the Gądów synoptic station in the years 1935–1941, used in this paper, are available from the easily accessible NOAA data resources (*internet sources*). The German meteorological annals (*Ergebnisse* 1936–1938, 1959) published measurement data from Gądów for an 8-year period (August 1936 – March 1944). Most of these data (aside from wind and precipitation) were probably connected with the aeorologic-climatology station.⁴ The temperature data from these two sources differ about 0.3°C on average, because topoclimatic conditions in various places near the northern border of the airfield were different. The monthly *Tp* values in the years 1935–1936 recorded at the synoptic station were higher on average by about 0.1–0.2°C than the temperatures for Krzyki.

⁴ There is a note in *Ergebnisse* (1936) that climatological measurements were made at the northern border of the airfield 100m W from the big plane hangar, but wind and precipitation measurements were made in other places near the new airport building. The cartographical analysis of different German and Polish archival sources points to the fact that the first location was probably near the NW border of the airfield, and the second over 500 m E of the first.

On the basis of this, the authors concluded that by applying various linear and multiple regression equations between the analyzed stations, it is possible to link the pre-war Tp data with post-war Polish measurements from the Gądów and Strachowice stations in a logical and statistically correct manner.

The average monthly Tp values from Gadów and Strachowice are also not identical. The average differences noted in the period 1960-1965 amounting to about 0.2°C (values are higher for Gądów), are representative for a longer period of time. This can be proven by comparing Tp values for Gadów and Strachowice with the synchronous values for the nearby stations (Swojec, Biskupin) and those further away at a distance of 30-60 km (Legnica, Oborniki, Szprotawa, Wińsko). The resulting estimated temperature differences from these two stations oscillated around 0.2 °C (in relation to Legnica, Oborniki, Swojec) and 0.4°C (in relation to Biskupin, Wińsko) and are of a different character. A comparison of monthly values of Tp Biskupin and Tp Strachowice in the years 1962–2007 reveals a systematic decrease in temperature difference between these stations. The Tp differences between Swojec and Strachowice in this same period were stable with insignificant oscillations. This indicates that Tp in Biskupin was under the strong influence of UHI, values of which increased in comparison with the stations on the outskirts of the city or located in agricultural areas. Thus, slight irregular differences were often connected with different methods applied to calculate Tp averages. For example, a comparison of the average from 8 synoptic fixed times with the average from 3 standard climatological once or the average calculated with the formula $(Tp \ 6 \ UTC + Tp \ max + Tp \ min + Tp)$ 18 UTC)/4, which has been in use since 1998 in the the Polish meteorological IMGW net, revealed a difference of about 0.1°C (occasionally 0.2°C).

A very characteristic feature of Silesian and Polish meteorological stations (except for Cracow and several mountain stations) is the general lack of homogeneity for longer time series as well as the presence of long interruptions in measurements during the second World War and for some years immediately after the war. For example, most Silesian stations (as was also the case in Wrocław) had two or three location changes in the period between the first and second World Wars. Therefore, it is very difficult to find a homogeneous 20-year temperature series taken in the climatic background surrounding Wrocław (within a maximum radius of 100–150 km) in order to make representative comparisons. Before 1945, only the following

mountain stations in Lower Silesia functioned in the same location for an extended period of time: Śnieżka (Schnnegebirge), Miłków (Arnsdorf) and several others in the Sudetes. However, the importance of these stations for the reconstruction of Tp in the Wrocław series is limited due to varying and specific mountain topoclimatic conditions and their greater distance from Wrocław than from lowland stations.

In this situation, the fact that the homogeneity of many analyzed series was interrupted greatly complicates the reconstruction of the Wrocław Tp series. The most representative series from the other Wrocław station are only partly compatible (sometimes only by a 2-year period), as far as the synchronous period of measurement is concerned, with the analyzed data from Wrocław base station operating at that period. This resulted in the fact that only one linear regression equation could be deduced, as a common statistical tendency for all the estimated monthly values (for example: the correlation of Krzyki with the Gadów synoptic station was possible only for 1935-1936 with R²=0.9996). When possible, we took into consideration multiple regression equations deduced from several analyzed stations instead of a simple correlation with one station.⁵ Such reconstructed values were additionally verified by means of a comparison with several other Silesian and Polish stations (including Cracow) against a more distant climatic background. Finally, the reconstructed runs of temperature data (1791–2007) for Wrocław were verified through comparison with the same long-term distance runs of *Tp* from Warsaw, Berlin, Prague, Vienna, De Bilt. The problem of spatial and time anisotrophy in relations between the values from the stations analyzed has been discussed in this context.

Having applied Alexandersson's method (Alexandersson 1986) to reconstruct homogenous series of Tp and P, the authors nonetheless took a critical approach to the nearly mythical status of the reconstruction capabilities of this method. Complications arise when the data that is compared comes from distant locations in differing climatic backgrounds (see sections below). Reconstruction of precipitation patterns proved to be difficult

⁵ The climatic data were often derived from such station measurements where there had been changes of location as well as of observation times. Therefore, they required careful homogenizing which assumes that nonhomogeneity of the series is not possible if it is not obvious, and if the apparent interruption of continuity of measurement values is constituted by physically explicable changes in the station's natural surroundings. As a basic principle, the authors referenced the checked series to the values occurring in the immediate vicinity of the station examined.

because it requires taking into consideration the important influence of the UHI (Urban Heat Island) effect (Landsberg 1981) and of urban precipitation shade (Fig. 1) on spatial variability, even on a fragmentary and small areas of Wrocław. These influences were the subject of many detailed analyses and studies (among others: Zipser-Urbańska 1968; Schmuck 1967; Dubicka 1994; Szymanowski 2004; Bryś 2007). The differentiation of the precipitation pool in a big city has a dynamic character and is connected with the dominating direction of rain-cloud mass advection. In the case of Wrocław, the largest frequency of rain-cloud mass advection from the western sector (W-NW) is counterbalanced by the largest precipitation efficiency (P intensity) from the SE sector (despite the low frequency of rain mass from this direction), connected with the influence of summer Genoese lows. A study by Dubicka (1994) closely analyzed the connection between precipitation in Wrocław and air circulation.

The fact that there were at least several precipitation stations (or posts) in the city or its environs at different times over the twentieth century allowed the authors to establish the fundamental and strongest spatial connections for P. Linear and multi-elemental regression equations were deduced to arrive at monthly precipitation values observed in the city center which could credibly be merged with those registered at its borders. Gaps in the observation material, such as the year 1945, were filled in by interpolations based on the quotient method and precipitation totals from neighboring stations further away. The results of estimates made by Wiszniewski's group (Atlas Klimatyczny Polski 1979) were incorporated. Similarly, in the years when the references made to other Silesian stations were credible (Bryś and Bryś 2010; Galle 1857, 1879), the authors carried out a general verification of the earlier measuring materials for Wrocław. On the basis of this, they first reconstructed the precipitation series for the years 1799-2007. Only the P values for topoclimatic conditions of the central-urban reference station (Botanical Garden) were reconstructed in this paper. Not only had the location of precipitation measurement changes to be taken into account, but so too did the influence of various reception surfaces and the construction of measurement instruments. These latter changes occurred in various years throughout the nineteenth century and were noted by Galle (1879). He also analyzed their influence on the quantity of precipitation measured in Wrocław. This study makes use of the results of Galle's analysis, which in turn were based on corrected precipitation values deduced by him for the

years 1851–1876. In making his corrections, Galle used only one quotient coefficient for the entire year, but the authors proceeded differently. Based on his research and relations obtained from previously mentioned, parallel precipitation measurements taken on the terrace of the Mathematical Tower and in the Botanical Garden, the authors deduced statistically substantial relations in the form of linear regression equations. The separate equations were used for each month to correct precipitation totals for the years 1799–1850. The authors converted the erstwhile precipitation notation from Paris lines to millimeters, using the formula: 1 Paris line = 2.256 mm.

This basic measurement series was completed with estimated P monthly totals for the years 1791–1798. In this case, the authors used statistically substantial relations of monthly precipitation frequency and types with their monthly efficiency (multiple regression equations were deduced separately for each month) noted for the nearest 30 years. The 217 year Wrocław precipitation series for the years 1791–2007 which they obtained was divided into 4 sub-series: 1791–1842, 1843–1894, 1895–1945, 1946–2007 and verified by the Bartlett test (Twardosz 1996), which did not produce any differences that would have disqualified a homogeneity series deduced by this method. The worst results were obtained for July and May, but the value of X^2 calculated did not exceed the threshold of significant value to the level of p=0.05.

Comparison with long-term precipitation courses noticed for other central European stations (Warsaw, Cracow, Prague, Berlin) showed slight deviations for certain years, most probably connected with local determinants. Their closer analysis requires a longer, separate study. In this article the authors have limited themselves to charts, as in the case of temperature, showing only the pattern of monthly, seasonal, half-year and annual precipitation totals in Wrocław in the years 1791–2007 and to a short discussion of these patterns.

It is worth bearing in mind the unpredictable nature of heavy storm-produced precipitation, the influence of which, in certain summer situations could create a considerabe distance between precipitation results (even those deduced from the most reliable statistical methods) and their real value. Uncertainty about results reconstructed in this manner, considering the fluctuation of the deviation mark (+ or -) from real values and their mutual compensation, which negates their influence over a longer period, should not have a greater significance for the reality of deduced long-term precipi-

tation trends. Being aware of this "uncertainty principle" in the pursuit of the greatest precision in the reconstruction of long-term series of *P* grants us a perspective on the re-occurring practice of replacing precipitation measurement values (even with some errors) with estimated values.

It is worth noting that in the case of small measurement errors and in many controversial situations it is better to leave the data unchanged, neither erasing nor disturbing their whole information value, which is very significant for climatic research. Apart from quantitative information, the qualitative information reflected by this data is also important (for example: fluctuation patterns, relation with neighboring data etc.), which is also a reason for it being preserved in its entirety, in spite of its imperfections. Thus one avoids creating "smoothed over" and artificial meteorological data series that are statistically correct, but discordant with the climatic reality being reconstructed. Research into long-term climatic trends cannot impoverish the available sources on short-term climate dynamics.

On certain aspects of the anisotropy of the temperature and precipitation fields in central Europe

Owing to its location in the range of moderate geographical latitudes, the climate of central Europe is subject to the teleconnectional impact of long-lasting changes of barometric tendencies in the source Atlantic region reflected by NOA and AO indicators. Seasonally variable fluctuations in intensity dependent on complex long-term dynamics and the spatial range of these influences modify, among others, the temperature field in central Europe and determine its variability in different spatial and temporal scales. A specific feature of this variability is the phenomenon of thermal anisotropy and its temporal and territorial differentiation.

Thermal anisotropy is the representation of various thermal features in different directions reflected in geographical space. This study focuses solely on the analysis of air temperature anisotropy of the ground layer (i.e. temperature taken in a meteorological cage at a standard height of 2.0 m). In the case of Poland, anisotropy thus understood and climate continentalism increases as one moves from over Denmark towards the east and advancing inland. Therefore, this evinces a close correspondence (inversely proportional) with the gradual decrease in the impact of thermal oceanism. Additional modification is introduced by local physical and geographical features as well as the nature of the active area surface. Among these factors, vegeta-

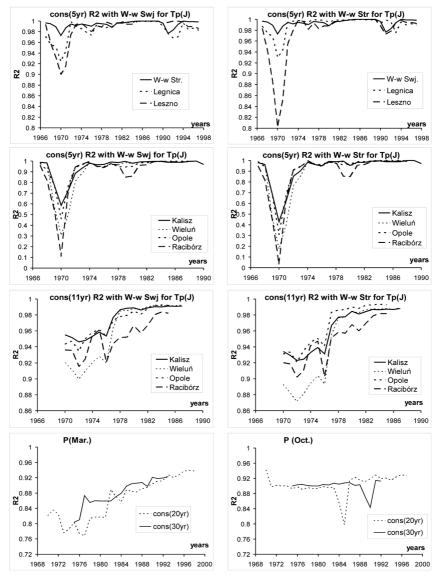


Fig. 3. Temporal anisotropy for monthly values of air temperature (Tp) as exemplified by 5- and 11-year consecutive (cons) correlations (R2 – coefficient of determination) between Wrocław-Swojec (W-w Swj) or Wrocław-Strachowice (W-w Str) and other meteorological stations in January (J). Similar phenomenon for monthly totals of precipitation (P) using the example of 20- and 30-year consecutive correlations between Wrocław-Swojec and Wrocław-Strachowice in March (Mar.) and October (Oct.). Explanation: *Tp* data for the years 1965–1999, *P* for the period January 1961- March 2008

tion cover and areas that are urbanized or have otherwise undergone strong anthropopressure, play a significant and dynamic role. Vegetation cover, differently influencing the effectiveness of area evaporation in each stage of its development (phenophase or agriphenophase), also functions as a thermal buffer. It restricts solar radiation energy inflow into soil and likewise limits explicit heat stream outflow from the ground into the air. Consequently, it affects the redistribution of heat accumulated in soil and is characterized by considerable seasonal and long-term variability. It exerts a disruptive effect on thermal records of the meteorological station as a result of its territorial expansion as well as the growth of perennial plants near the station. This may also lead to an increase in the roughness of the atmospheric substratum. Furthermore, urbanized (or industrialized) areas influence local thermal conditions through the UHI effect, through a rise in radiation extinction as well as alterations in the wind field and the precipitation field (caused, among other things, by accumulated effects and precipitation shadows).

According to Flohn (1981, 1993), horizontal climate variations are both less substantial and significant than temporal ones in medium size areas (such as Poland or its regions). This phenomenon is clearly visible when one looks at modifications of the temperature field. Similar phenomena appear to an even greater degree in variations of the precipitation field (Fig. 3). Its effects are reflected both in seasonal changes of thermal (or precipitation) anisotropy as well as in different aspects of its long-term variability. The primary role of the macrocirculation factor in modifying thermal (or precipitation) features of regions exceptionally "sensitive" to climate variability (i.e., regions of so-called transitory climate) is particularly evident during periods of circulation changes. Records of long-term correlation changes between different stations attest to this fact, and we make use of this in our homogeneity analysis presented below.

The spatial correlation of climate changes, or "teleconnection", provides evidence of a certain spatial order in the variability of different climate elements, including air temperature and precipitation, among others. The analysis carried out in this work on the teleconnectional relations of each of these elements (in a selected part of central Europe) has been rendered by use of the R² determination coefficient in its values for different (5- to 35-year) dynamic correlations. Long-term records (1946–2007 or at least a 35 year period within this span) of correlation changes of monthly mean air temperature (or monthly precipitation totals) between stations in Wrocław and

its environs were compiled separately for each month. Relations between Gądów, Strachowice, Swojec, Biskupin or Poznań, Opole, Zielona Góra and these background stations have been examined in a manner similar to those between stations in Wrocław (the graphs in Fig. 3, for greater clarity include only 2 such stations: Strachowice and Swojec and 6 stations of the climatic background: Legnica, Leszno, Kalisz, Wieluń, Opole, Racibórz).

Contesting the myth concerning the reliability of certain homogeneity methods

The application of the R² coefficient (Fig. 3) reveals not only the statistical significance of these correlations, but also the character of their temporal and spatial variability. It also enables the identification of periods of relative stability and greater fluctuation (which are different for various months) as well as the characteristic, "funnel-shaped" refraction forms of their values. The latter phenomenon can be associated both with a continuity disruption in the measurement methodology (nonhomogeneity) in one of the compared stations, as well as with the occurrence, as time elapses, of different statistical aspects of climate changes (e.g., climate discontinuity, periodicity or climate rhythm) or macrocirculation changes. The occurrence of temporal consistency of these forms in different examined (base) stations in juxtaposition to the same compared background (referential) stations excludes the disruption of measurement homogeneity as a cause of their topographical features. Therefore, the significant role of climate variability in the formation of R² discontinuities in Tp (or P) teleconnections in different spatial mesoscales can be inferred. This, in turn, influences (in some years quite considerably) the accuracy of the reconstructed data. Simultaneously, it undermines the primary assumption of those homogeneity tests that involve weighed mean values, being linear correlation coefficients between temperatures (or P) in the base stations and their adjacent stations (or exponential function of their distance), as a measure of a constant long-term influence of the reference station on the examined station (and vice versa).

Alexanderson's test (Alexandersson 1986; Brázdil and Stepanek 1998; Miętus 1998; Lorenc 2000), which applies the so-called referential series (a temporal sequence of mean values weighted from values of an examined meteorological element recorded in adjacent stations) has gained popularity for the restoration of data homogeneity. It is based on the assumption

that there exist, for the purpose of homogeneity, reliable statistical relations between the examined station and its adjacent stations. These, it argues, can be evaluated by the use of linear regression and weighted mean value, which results in a reference sequence (for comparisons with the sequence of values from the examined station). It is only necessary to be certain that in these adjacent stations, the continuity and homogeneity of the observation series are maintained in the period examined. The homogeneity criterion is the most difficult to fulfill due to the fact that, for objective reasons, there is no absolute possibility of retaining the same observation rules in the course of long-term (e.g., decadal) measurements conducted by different observers, using various equipment and not always in the same place and time. Location changes were often considerable (vertical and horizontal relocations of the measurement station), while methods of data compilation underwent substantial alterations (e.g. an alteration in the method of calculating a mean, day value connected with changing the quantity or observation times). Detailed and thorough information (metadata) concerning these alterations is not always available. Consequently, even the stations regarded as homogeneous do not possess fully homogeneous observation sequences. The concept of the referential series diminishes these objective methodological inconveniences as it focuses on the assumption that if one takes into consideration a greater number of reference stations for the creation of the referential series one can thus reduce measurement homogeneity disruptions in any of these stations. The simultaneous occurrence of significant nonhomogeneity in the majority of stations is highly improbable (apart from rare instances when system changes were introduced simultaneously in all of the stations) and therefore ought not to be taken into consideration.

As indicated, however, by the results cited earlier, of the correlation changes in the years 1946–2007 between several examined stations and their reference stations (the same for all), the use of the referential series does not guarantee an increase in the estimation accuracy of the missing data. This series (as well as its relations to the value of the examined station) is subject to previously mentioned climate changes that destabilize weight coefficients, particularly in periods of macrocirculation changes which are not taken into account in the calculation of weights. As is assumed in the criticized method of restoring homogeneity, Alexandersson's average stability of weights or weights in the form of average stable correlation coefficients is at variance with natural reality. The temporal variability of spatial air temperature rela-

tions between Wrocław (or other examined stations) and the stations of closer or more outlying surroundings, measured by means of R² value, indicates the necessity of considering separately possible periods of greater estimation unreliability and inaccuracy in the homogeneity procedures. This is indispensable and should be applied as a constant regular practice determining the possible reliability limits of the reconstructed series (or its respective parts), and not those stated on the basis of some theoretical, oversimplifying assumptions. The formal and statistical methodological criteria need to be in accordance with the over-arching concrete criteria closely related to the complexity of natural reality and the physical processes that define it.

Ignoring these methodological precepts can result in the creation of fictional segments of the examined temporal sequences that are divorced from natural reality, which only appear to be sequences of numerical values constituting a reliable reconstruction of measurement continuity of the meteorological element examined. They imply an exaggerated and physically inexplicable accuracy of value estimation in the whole study period. This implication can be achieved through high R² values acquired by means of standard (and therefore non-shifting, long-term periods) correlation comparisons of longer sequences of the examined data, as well as through the application of adequate linear regression equations. One often must confront the "fetishization" of the applied method (eg., Mietus 1998), an odd result of which is the homogeneity of monthly mean temperature with an accuracy of up to 0.01°C. Equally irrational is the "mythical" conviction that this method is efficient in the detection of nonhomogeneity periods and in the precise estimation of correction quantity. This method is thus unreliable in the process of restoration of an inhomogeneous value to its methodological consistency with former homogeneous values. The occurance of temporal anisotrophy is important phenomenon of climate variability. The research results about anisotrophy presented in this paper by the auhors show unequivocally that application of Γ parameter for examing the inhomogeneity of the temperature series leads to an ambiguous interpretation of the genesis of Γ max rather than to the detection of measurement continuity disruptions by of Alexandersson's test. The correlation connections presented herein, analyzed also in different one hundred year periods for Wrocław and outlying stations, are of a non-stationary nature, depending to a large extent on macrocirculation changes. The role of these changes is generally demonstrated in the regular or quasi-regular effect of temperature variations in

the analyzed long-term series, including, among others, the detection of approximately 8 and 12-14-year oscillations in the Czech mean monthly Tp series (Pisoft et al. 2004). This paper also took into consideration other aspects of this dependency.

It is probable that the spatial distribution of this nonstationarity is connected with the phenomenon of its migration over time, which in turn is in central Europe a repercussion, with ever-increasing delay (as one moves further away from the source area of teleconnection), of the climate changes underway in the North Atlantic and subtropical zones. In the mesoscale area examined, this is reflected in the spread of certain features of the activization of zonal or meridional circulation along with spatially and temporally diverse effects of their conflicting influences. The previously mentioned phenomena of local character connected in turn with differentiation of the active area as well as its radiation, temperature and evaporation interactions with different air masses are of secondary importance. Among them vegetation cover and anthropopressure play an important yet secondary role compared with the impact of circulation.

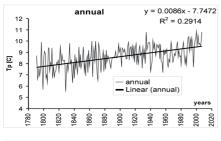
Reconstruction results and a short discussion

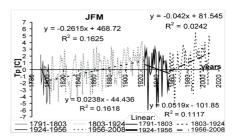
The reconstruction of the Tp values made by the authors, in contrast to that of Pyka (1998a), shows a distinct increase of temperature in Wrocław over the twentieth century as well as over the entire analysed 217-year period. In the years 1791-2007, the warming tendency, for topoclimatic conditions of the former Breslau-Sternwarte Observatory, amounted to 0.85°C/100 years (Fig. 4). The strongest secular trend appeared in the years 1829-1939 and amounted to 1.43°C/100 years. A relatively small increase in temperature occurred during the twentieth century, because its trend only amounted to 0.39°C/100 years. This was the result of a relatively strong decreasing tendency in the years 1934-1956, which for 22 years interrupted the process of climatic warming. A very strong increasing tendency began again after 1956. The increase in warming for the years 1956–2007 amounts to 1.65°C, which is equal to a tendency of 3.2°C/100 years. Of decisive importance here were changes in the winter time. In the period from January to March (JFM) a warming tendency in the years 1791–2007 amounted to 1.24°C/100 years, while in the periods from December to February (DJF) 1.22°C/100 years, and November to January (NDJ) 1,23 °C/100 years. These were the strongest

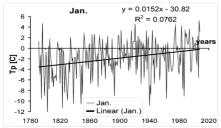
trends from all analysed 3-month periods of the year. The weakest tendency $0.45^{\circ}\text{C}/100$ years appeared in the period of August to October (ASO), while in the period of July to September (JAS) it amounted to $0.56^{\circ}\text{C}/100$ years, and from September to November (SON) $0.6^{\circ}\text{C}/100$ years. The weak tendency ($0.66^{\circ}\text{C}/100$ years) was visible also in the summer (JJA), and relatively strong ($0.99^{\circ}\text{C}/100$ years) in the spring (MAM). The result of these tendencies was that the trend in the warm half-year (April-September) amounted to $0.64^{\circ}\text{C}/100$ years. The largest increase occured in January ($1.50^{\circ}\text{C}/100$ years), March ($1.43^{\circ}\text{C}/100$ years) and December ($1.23^{\circ}\text{C}/100$ years). The smallest one occurred in August ($0.42^{\circ}\text{C}/100$ years), September ($0.43^{\circ}\text{C}/100$ years) and October ($0.45^{\circ}\text{C}/100$ years). All these trends are statistically significant to a level of α =0.05.

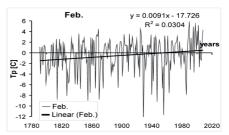
The warming trend in winter (DJF, and also JFM) started in 1803; a cooling tendency was apparent earlier. However, the period 1803–1838 had a particular significance for temperature changes in Wrocław. The first three decades of the twentieth century were the final culmination of the "little ice age" in Wrocław and a period of transition to a new era of progressive warming (Bryś and Bryś 2010). Of crucial importance for this new tendency was the strong secular trend of increase in winter temperature in the years 1803–1924 (2.35°C/100 years), and especially strong in 1838–1921 (2.86°C increase, i.e., 3.42°C/100 years). After 1924, we see the beginning of a dynamic period of greater oscillations of the winter *Tp* from year to year (compared with 1791–1924). This period consists of two different stages: 1) 1924–1956 with a strong decreasing tendency (1.35°C of the drop, i.e. 4.16°C/100 years); and 2) 1956–2007 with a strong increasing trend (2.71°C of the growth, i.e. 5.27°C/100 years).

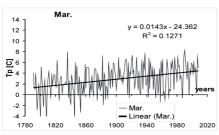
The reconstructed *Tp* and *P* series for Wrocław represent long-term modifications of natural climate oscillations and tendencies by specific urban conditions, especially UHI and many phenomena connected with it. In addition, these modifications are to a large extent formed by local, anthropogenic specificity (urban and industrial physiography), which overlaps with local natural conditions (location of a city and its natural physiography, such as relief, types and distribution of plant cover, water reservoirs and tides etc.). This results in the strong impact of differentiated topoclimates in large urban areas on the principle climatic features observed in meteorological stations located in various places around the city analyzed. Relations between the base station and measuring stations do not have a stable character but are

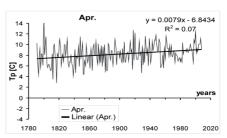


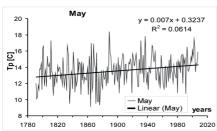


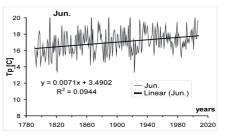


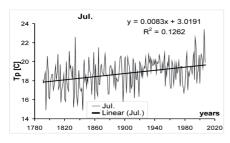


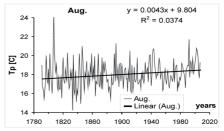












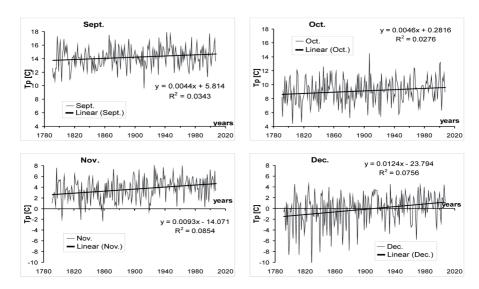


Fig. 4. The course of the mean annual, selected 3-month period (JFM) and monthly (Jan., Feb., Mar., ...Dec.) temperature values (Tp) and their linear tendencies for the years 1791–2007 in Wrocław

rather variable due to the dynamics of conditions (sometimes in a specific range of values) from year to year (time anisotrophy). Their directional, progressive variability in time most often reflects the dynamics of anthropopressure on the urban landscape. Greater instability of a macro-circulatory provenance may appear alongside locally conditioned instability. All these factors have an influence on the precision of the reconstructed series Tp and P in the base station of reference (Breslau-Sternwarte), the estimated values of which originate from the interpolation of data from other measuring stations.

However, the tendencies and the structure of climatic oscillations presented here were shaped primarily by macroscale changes in circulation. In the analyzed 217-year history of thermal changes in Wrocław there are two long-term climatic variations that are especially important. They appear, similar to those presented in winter period (Fig. 4), also in the long-term course of annual Tp values. The most important change began during the first three decades of the twentieth century and was connected with the previously mentioned transition to a new era of progressive warming. The second change, which has taken place over the last five decades, following

a 22-year (1934–1956) period of cooling, is connected with a new phase of warming and a sub-phase of significant increase of the Tp in the years 1987–2007. In the authors' opinion, there is an analogy to be made with past temperature changes in Wrocław. In 1797–1829, there was a period of cooling. The period of warming in the years 1829–1956 had two major phases with two sub-phases. In the first phase one can distinguish two larger sub-phases: 1) a longer one of increased Tp (1829–1868); 2) a shorter one of decreased Tp (1868–1888). The second phase was similar: 1) a longer sub-phase with Tp increase (1888–1934); and 2) a shorter sub-phase with Tp decrease (1934–1956). Profound and rapid changes appeared in the years immediately before the culmination of warming began, as was the case in the sub-phases of cooling.

The latest period (phase) of warming has so far been characterized by a longer sub-phase of temperature increase. However, it seems that this sub-phase culminated between July 2006 – June 2007 and we should now probably expect, as was the case previously, a sub-phase of decrease in the average annual temperature. The earlier changes had a characteristic form – their proximal stage was longer than the distal stage. The first stage was a period of about 40–50 years and the distal stage (the years of post-culmination decrease) appeared as a relatively steep and short period (about 22–23 years, sometimes about 30–35 years). Is the present a new transition period and a turning point in climatic change? Possibly, but it may also for some time be a continuation of the increase sub-phase. On the basis of analyses of other long-term *Tp* series (Berlin, de Bilt, Cracow, Warsaw, Prague, Vienna), the authors would argue that the turning point has already occurred or will occur during the next few years.

The increasing tendencies are also evident in the reconstructed totals of P (Fig. 5). For annual precipitation totals the linear trend amounted to 68.9 mm/100 years, i.e., relatively 12.8% of the growth. For hydrological year (Nov. – Oct.) the trend was a little stronger, because it amounted to 69.5 mm/100 years (12.9% of the growth). In the warmer half-year (Apr.-Sept.), this tendency amounted to 39.8 mm/100 years (11.5%), but for April-June only 13.2 mm/100 years (8.3%) and for July-Sept. 29.4 mm/100 years (12.3%). In the remaining 3-month periods, nearing

⁶ The highest average annual Tp (10.9 °C) in Wrocław appeared in 2000. The value is over 5 °C higher than the smallest average annual Tp (5.5 °C) for 1829 (see: Table 1).

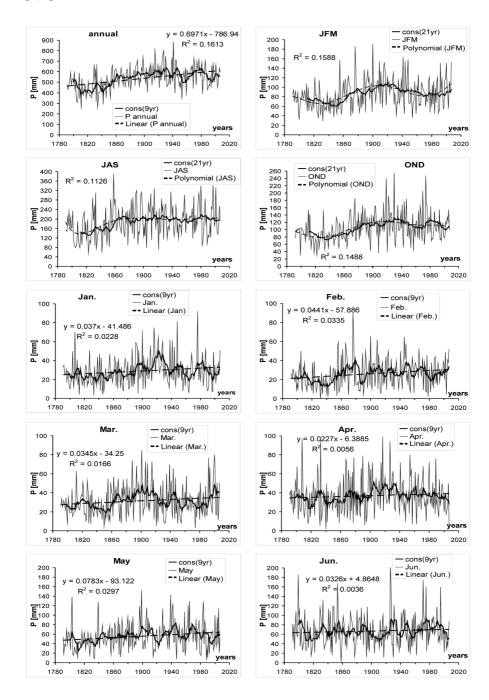
calendar seasons, it amounted to 11.4 mm/100 years (13.3%) for Jan.-Mar. and 17.7 mm/100 years (16.8%) for Oct.-Dec. For comparison, in the climatological seasons it amounted to 13.3 mm/100 years (15.2%) in the winter (DJF), 13.4 mm/100 years (10.8%) in the spring (MAM), 27.8 mm/100 years (13.3%) in the summer (JJA) and 14.9 mm/100 years (12.6%) in the autumn (SON). From the midst of remaining 3-month periods of the year, the weak trend was visible also for Feb.-April (10.0 mm/100 years and 10.7%), and relatively strong for Nov.-Jan. (15.3 mm/100 years and 15.8%). For May-July the values amounted to 25.4 mm/100 years and 12.8%, while for Aug.-Oct. only 18.46 mm/100 years and 12.5%.

The largest monthly increases appeared in July (14.4 mm/100 years and 19.0%) and August (10.1 mm/100 years and 15.3%). The relatively strong ones are also in November (6.5 mm/100 years and 18.5%), October (6.3 mm/100 years and 17.0%) and February (4.4 mm/100 years and 17.0%). The lowest increase appeared in September (2.1 mm/100 years and 4.6%), April (2.2 mm/100 years and 6.0%) and June (3.3 mm/100 years and 4.9%). Except for September, all these trends are statistically significant to a level of $\alpha = 0.05$.

The largest annual totals amounted to more than (or about) 800 mm and appeared in a phase of pluvial maximum between 1880 and 1945 (881 mm in 1941, 822 mm in 1915 and 799 mm in 1927). In addition, a shorter period of high values, sometimes above 700 mm, appeared between 1970 and 1986. The smallest annual totals were under 300 mm and appeared in a phase of pluvial minimum between 1804 and 1852 (247 mm in 1811, 261 mm in 1834 and 289 mm in 1842). The first three decades of the phase saw the transition period mentioned above, leading to the most important climatic reversal in the twentieth century, whose thermal and precipitation increase effects continue to the present.

A period of stagnation or slight decrease in the monthly and seasonal totals of *P* during the warm half-year has been observed over the past two decades. An opposite tendency appears in the period January-March, as a strong increase trend has been noted for March beginning in the 1990's.

The reconstructed values of *Tp* and *P* presented in detail in Table 1 and Table 2 also show their long-term quasi-cycle oscillations. These are related not only to circulation changes but also to solar variability effect. The authors discussed these issues at length in another article (Bryś and Bryś 2007).



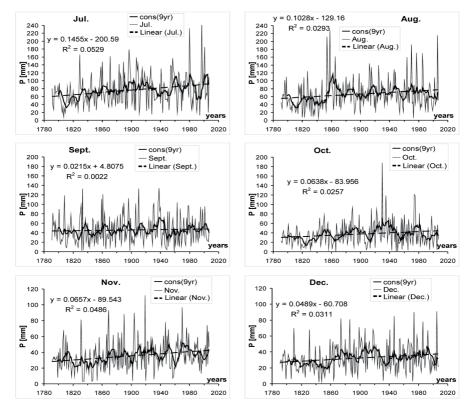


Fig. 5. The course of the annual and selected 3-month periods: January-March (JFM), July-September (JAS), October – December (OND), and also monthly (Jan., Feb., Mar., ...Dec.) precipitation totals (P) for the years 1791–2007 in Wrocław and their polynomial of the 6th degree or linear trends. The values also are smoothed over by 9- or 21-year moving (cons) means

Table 1. Monthly (Jan., Feb., Mar.,...Dec.), annual (January-December; ann.), warm half-year (April-September; A-S) and seasonal (DJF, MAM, JJA, SON) mean values [°C] of air temperature in Wrocław in the years 1791–2007. Explanation: bold italics for estimated or interpolated values, bold and colouring for extreme values (max - gold, min - blue)

A-S DJF MAM JJA SON
0.0
3.1 -2.6 6.1 16.7
-2.6 6.1
6.6 13.1 7 12.5 7.6 13.4
0.1
1.6
7
ŀ
16.7 11.3
1.3

10.2 10.2 10.1 8.6 8.2 9.7 δ. ∞ 6.2 8.6 8.2 7.7 7.7 9.1 8.1 10 ∞ ∞ 16.1 18.4 16.4 16.4 16.6 16.8 18.6 16.9 17.3 16.3 16.4 15.5 18.2 20.1 17.7 17.7 16.7 16. 18. 19. 6.8 8.2 8.6 9.9 9.9 8.3 8.2 6.3 7.8 8.3 7.6 6.8 9.5 8.2 8.5 6.7 10 0 ∞ -2.8 -3.9 -4.5 -2.4 6.0--0.3 -3.8 -2.5 -3.5 -4.3 8.8 -1.6 0.8 -1.6 -2.5 1.5 -5.1 0.8 1.1 0 15.9 14.8 14.6 13.4 14.1 13.9 13.5 13.5 13.9 14.3 14.4 15.2 14.4 14.3 15.2 14.8 14.2 13.2 14.2 15.2 16.7 4 15 8.9 7.8 7.4 8.1 8.3 8.5 9.5 8.8 8.3 8.2 8.3 7.9 5.5 7.3 8.4 ∞ ∞ ∞ -8.6 -4.8 -3.2 -2.5 4.0 -1.4 -2.2 2.5 2.9 4.5 8.0 0.2 -1.4 -1.5 1.3 0.7 3.7 7: 1.4 ņ 4 3.2 6.4 4.5 4.4 5.2 -0.2 3.6 -2.1 1.6 2.4 4.2 2.1 1.6 3.1 4 5 11.5 11.5 7.2 6.7 8.6 4.6 7.9 8.5 9.1 11.2 9.6 9.6 9.6 7.4 6.2 7.4 11.3 8.6 7.8 12.1 7.7 7.2 ∞ 10.9 14.4 13.6 14.2 13.5 11.3 12.9 14.5 14.3 14.2 14.6 13.5 12.2 14.1 12.9 14.1 13.1 13.1 14.1 13.1 13.1 15 16 20.3 16.8 16.3 16.5 16.4 18.2 19.8 18.6 16.5 18.4 18.4 14.2 19.1 16.7 17.7 17.7 15.5 17.4 17.1 17.7 17.3 17.5 17.5 16.9 17.5 19.7 16.9 17.5 17.9 18.7 15.8 17.9 17.3 16.9 19.3 19.6 18.7 18.3 15.2 15.7 17.1 21.1 16.1 20 20. 19. 16.9 18.3 20.5 16.5 14.6 14.4 17.9 15.9 17.9 13.8 13.4 15.8 16.6 15.3 18.4 15.8 15.5 15.7 17.4 16.7 15.4 19.1 17 10.5 13.5 11.6 12.8 17.3 13.7 13.2 11.6 13.3 12.5 14.2 12.6 13.2 12.9 16.7 13.1 15.2 13.1 12.1 11.8 11.1 4 4 9.6 8.3 10.1 2.8 8.3 9.6 8.2 6.8 6.6 8.9 5.8 7.7 9 6 7 9 $\stackrel{\leftarrow}{\vdash}$ -0.5 0.2 4.3 2.5 1.2 3.7 2.7 3.7 3.7 1.4 6.4 2.8 1.2 -1:1 3.4 3.4 1.9 1.8 1.9 2.3 1.7 \sim 8.8 -0.8 0.8 -0.2 -1.8 -2.2 2.6 -2.7 2.6 3.3 -1.6 -3.4 -5.3 1.5 2.1 1.3 -8.1 -5.7 2.7 0.1 7 -4.6 -5.6 -0.4 -0.4 -0.2 -11.3 -0.4 -8.4 -5.3 8.6--6.7 -5.1 1. -3.1 9 7: 7 0 1823 1824 1825 1826 1828 1829 1830 1813 1814 1815 1816 1817 1818 1819 1820 1822 1833 1811 1812 1821 1827 1831

Fable 1. cor

Fable 1. cont

10.2 10.2 8.6 8.2 8.0 9.3 9.2 9.2 9.6 9.1 16.4 17.8 16.9 18.2 16.9 19.7 17.3 18.3 16.7 18.2 19.5 15.3 17.2 16.1 17.7 17 17 16. 10.1 6.4 7.6 8.5 9.3 7.8 6.3 5.7 5.5 5.5 9.5 6.7 5.1 8.1 -0.9 -0.5 -2.8 -0.2 -3.6 -2.6 -2.3 -5.5 -3.1 -0.1 0.7 -1.7 -1.5 1.8 0.9 7 9 4 7 14.4 14.5 13.3 13.2 13.4 14.3 13.6 13.4 14.3 15.9 14.2 15.5 14.2 14.3 14.5 13.7 15.1 14.1 15.4 14.1 15 8.8 8.8 9.5 6.9 8.3 6.3 6.7 7.6 8.4 7.6 9.5 7.6 8.5 8.1 8.1 7:1 8.1 8.1 ∞ 6.0--5.2 -0.9 -2.5 8.0 -3.2 -1.5 ÷. 3.2 1.6 0.9 0.2 -7.1 9.0 2.4 0.7 4 5.7 1.6 2.7 7 5 **0.**2 0.7 3.2 5.2 4.4 0.5 4.6 5.9 2.1 3.5 4.5 5.2 6.1 1.7 10.5 11.9 **9.1** 12.6 10.2 11.8 8.6 8.8 9.4 6.5 8.3 9.6 7.4 9.8 9.3 7.2 6.1 \vdash 12 6 14.9 13.8 12.6 13.6 13.2 15.5 13.4 15.9 14.4 14.2 11.8 12.5 14.4 12.2 12.2 14.7 12.9 12.8 14.1 12.1 13.1 15 12 19.4 16.6 20.1 15.9 16.4 17.6 16.2 17.6 18.2 16.4 18.4 15.9 20.2 14.9 **16.4** 21 19.3 17.2 18.2 17.7 15.1 18 17 16.9 20.5 16.6 18.6 15.5 16.6 19.1 17.2 17.2 14.9 20.2 17.9 18.6 17.3 18.3 17.8 18.9 18.4 19.1 17.1 19 16. 18.3 16.8 16.6 15.6 17.9 15.3 16.3 15.8 15.6 17.9 14.8 18.6 15.4 18.1 17.3 19.1 16.7 18.2 16.1 17.2 15.7 16 17.7 12.8 13.4 14.6 10.6 14.7 13.5 11.3 15.9 13.8 10.9 14.4 12.4 14.5 13.2 13.1 11.7 12.3 15.1 9.5 12.1 13 5. 10.3 6.8 9.7 6.6 6.7 4.7 8.2 4.1 9.1 7.4 ∞ 6.1 9 \sim 9.0--0.5 -3.3 0.3 6.0 -1.7 2.6 3.6 **5** 1.2 0.5 1.9 5.4 6.0 3.4 2.3 ∞ 0 9 -6.8 -0.3 -2.3 6.0 8.0 0.1 -1.3 -1.5 -1.3 0.5 2.3 2.8 9 9 7 -8.6 -0.3 -3.2 -1.8 -0.7 -2.3 -2.3 2.6 1840 1848 1856 1836 1844 1845 1849 1855 1834 1839 1841 1843 1852 1853 1837 1851

10.5 10.4 8.5 9.2 9.2 9.5 9.3 8.4 9.5 8.7 8.5 11.5 9.7 6.9 8.1 8.1 19.5 19.1 18.6 17.4 16.6 18.3 17.5 19.4 18.9 17.6 17.6 17.2 17.2 19.1 18.4 17.7 17.1 18 18 17 17 19 8 8.8 8.6 6.9 10.3 8.6 9.9 9.9 6.3 9.5 8.9 7.7 6.7 6.9 6.3 7.7 6 9 -5.8 -2.2 -0.7 -3.2 -1.5 -4.1 -1.8 -2.9 8.0 -2.1 1.5 1.8 -3.1 0.2 0.1 ψ 1: 7 _ 7 $\overline{}$ 14.6 16.6 15.4 14.6 14.5 15.3 15.6 15.6 14.5 15.4 14.2 13.5 15.4 14.4 14.9 15.4 15.1 15.1 15.1 14.7 14.1 5 15 8.5 7.8 6.5 8.7 9.2 8.5 6.8 8.5 9.4 8.2 8.9 6.9 9.8 9.2 8.5 8.2 8.5 7.3 9.1 10 10 -2.4 -0.5 0.8 -6.7 -4.4 -4.2 -0.5 -7.8 3.3 9.0 0.9 -1.5 0.2 -1.3 -1:1 1.9 3.5 <u>ڄ</u> 1.7 ψ 7 ÷ 7 8.0 4.8 9.0 6.3 4.2 9.0 0.7 4.7 5.4 4:1 2.1 2.2 3.2 5:1 1.5 \sim 7 7 10.8 11.9 10.3 11.9 10.7 10.7 9.9 9.4 7.3 8.9 6.3 9.9 6.4 9.1 11:1 7:7 7.4 6 9 ∞ ∞ 14.9 16.8 15.8 16.4 15.2 14.2 14.7 13.9 17.5 14.2 15.5 12.4 15.3 13.4 13.6 14.9 13.7 14.1 13.7 13.2 16.1 $\frac{1}{2}$ 16.5 17.5 19.9 18.5 20.2 16.9 19.8 16.4 19.2 18.3 20.2 18.7 17.5 15.7 17.3 16.7 18.1 19.3 ∞ 18.1 19 18 8 17.6 20.8 18.5 16.6 18.6 18.6 19.7 18.3 16.9 16.4 21.5 17.6 19.5 19.7 18.9 18.7 18.6 18.6 18.4 16.6 21.1 16 20 16.9 17.5 18.8 14.2 20.2 16.5 18.6 15.9 14.9 15.9 16.8 16.8 19.6 18.3 19.2 16.4 14.7 17.4 19.1 17.7 17.1 17.1 8 14.8 10.6 16.9 13.4 13.2 13.9 11.9 13.9 15.8 13.6 12.4 15.4 9.6 12.7 11.1 9.5 9.3 10.7 9.7 4 9.1 17 12 10.6 10.2 8.3 7.5 9.3 7.9 10.1 8.2 6.3 8.8 9.9 9.8 9.6 9.2 7.3 6.7 5.9 8.1 5.4 5.1 ∞ -0.5 -0.1 0.8 4.8 -1.2 0.3 5.7 1.2 5.1 5.1 1.9 0.7 3.6 1.4 4.2 4.9 4.8 2.2 2.6 2.1 4 8.8 -3.5 -7.5 -0.4 -0.5 -2.2 -2.4 -6.7 4.5 -1.2 -6.7 -1.2 2.9 2.7 2.5 2.8 3.4 3.1 0.3 1.8 0 9.0--0.5 -3.8 9.0 6.0 -6.3 -7.3 -2.5 -5.2 -3.3 -2.3 9.0 8.0 -6.7 2.6 -0.7 -2.2 0.2 1.8 4 1860 1868 1869 1859 1863 1864 1866 1874 1878 1879 1857 1861 1862 1865 1867 1870 1871 1872 1873 1875 1876 1877

Fable 1.

10.2 9.6 9.4 8.6 9.6 8.7 8.1 8.3 9.8 8.6 8.2 8.1 9:1 16.8 19.1 18.6 18.9 16.6 17.5 17.5 17.8 18.3 19.2 18.2 17.3 16.7 18.7 17.7 17 17 8.6 9.9 9.4 10.1 8.4 80.00 8.4 8.9 8.5 8.5 9.3 8.2 5.3 7.7 7.3 9.2 6.9 ∞ 6.0--2.4 -0.2 -4.2 -3.2 -1.9 -3.5 0.3 -1.9 0.1 -3.1 -1.3 1.5 2.6 1.6 7 0 14.8 14.6 14.6 14.5 15.2 14.5 14.3 15.8 15.3 15.9 16.2 14.5 15.3 15.4 15.8 13.5 14.1 14.7 15.1 4 4 15 4 8.5 8.8 8.8 δ. ∞ 9.2 8.7 8.5 8.3 7.9 7.6 8.3 8.5 8.5 8.5 8.5 9.7 9.4 7.7 8.1 9.1 6 9.0 -0.2 -0.7 -6.7 -0.5 -3.6 0.2 8.0 -1.9 0.5 0.3 2.4 1.6 3.3 2.4 0.1 7 $\overline{}$ $\overline{}$ 7 5.2 4.8 2.9 4.8 4.3 6.2 9.0 4.7 3.7 7: 2.7 2.5 3.1 1.7 2.2 2.3 7 9 11.6 8.8 9.6 8.6 6.7 9.8 8.6 9.2 8.3 8.2 9.6 11.1 5.2 9.1 11.1 8.7 9.1 ∞ ∞ ∞ 16.9 13.8 13.8 14.6 15.2 15.9 14.6 11.4 13.8 13.9 11.8 14.3 15.2 13.2 12.4 15.7 14.3 15.1 16.1 14 4 4 16.2 16.5 17.2 20.2 18.8 18.2 16.5 17.2 17.5 16.7 15.4 17.9 16.9 17.3 21.3 17.9 17.5 16.7 19.2 17.3 19.1 17 15.8 20.4 20.6 18.8 20.5 19.6 19.5 18.6 19.2 18.4 17.5 20.1 18.3 18.1 17.9 19.5 19.4 18.2 20.3 16.7 17 14.6 14.5 18.6 20.5 15.9 17.6 17.6 15.3 18.1 18.5 18.5 16.9 17.8 16.6 15.7 15.2 17.1 16.1 15 18 17 17 10.8 12.5 15.6 13.8 14.4 14.3 13.5 12.7 13.4 14.1 18.4 15.4 13.5 12.8 12.3 15.2 11.7 13.7 13.1 13.1 11.4 12 10.2 8.9 4.8 8.4 9.8 8.2 7.1 8.3 9.4 6.2 8.8 6.1 8.7 8.7 5.2 5.1 \vdash 6.0 -0.4 6.0 2.3 3.9 3.3 1.5 5.8 3.9 1.1 4.7 1.4 5.9 5.9 4.9 3.2 2.5 3.3 9.9--0.4 -4.4 -0.4 -2.2 -2.5 0.7 -4.1 -3.4 -0.2 0.2 1.8 2.6 0.7 1.3 1.6 2.1 2.1 1: -5 7 -5.8 -4.2 -4.5 -5.6 -3.2 -2.2 -3.3 -3.3 1.8 -2.3 7 1894 1886 1889 1902 1880 1882 1883 1885 1887 1888 1890 1891 1892 1893 1895 1901 1881

Fable 1. cont

10.3 10.1 10.3 8.6 8.2 10.1 6.2 8.5 9.6 8.4 6.7 9.2 8.2 9.7 9 10 ∞ 10 17.9 16.3 16.8 19.5 16.6 17.6 17.4 18.6 17.3 19.2 18.7 17.8 16.6 18.4 16.6 17.4 19.2 17.7 17.4 18.1 17.7 17.7 10.4 10.4 8.3 9.6 8.6 9.8 9.9 8.6 8.8 8.9 9.3 8.5 8.3 8.5 8.9 9.3 9.8 7.8 7.3 9.7 -0.2 -2.6 -2.3 -3.5 0.3 1.2 -2.1 0.3 2.3 -0.1 0.7 1.8 2.5 -1.2 0.5 1.3 1.4 1.6 2.4 7: 7: 0 $\overline{}$ 14.9 15.3 15.4 15.4 14.8 14.5 14.9 13.8 14.2 15.6 14.5 15.8 15.4 15.9 15.6 14.5 14.4 15.1 14.9 બ 14.7 16.1 14.1 4 8.7 8.8 8.4 9.6 9.8 9.4 9.2 9.3 8.4 9.5 9.6 8.3 9.3 9.5 9.6 8.5 8.1 9.3 7.9 8.0 9.3 6 ∞ 9.0 -2.8 ·0.3 -0.4 0.5 2.6 3.5 3.4 3.4 -1:1 0.3 0.4 0.5 2.2 1.5 2.3 3.5 2.5 3.1 7 7 7 -0.5 9.0 5.6 9.9 3.6 5.3 9.0 2.6 2.9 2.4 2.2 5.7 3.5 \sim \sim 7 10.5 10.5 14.5 11.6 11.6 10.1 8.9 5.6 9.4 8.8 8.4 8.9 9.8 7.6 5.4 8.7 6.7 7.3 5.7 9.7 10 6 9 14.6 15.6 13.9 12.8 15.1 12.9 15.8 13.4 12.5 15.9 15.4 16.7 14.2 14.9 14.3 13.4 9.6 13.7 12.7 12.1 14 13 4 20.5 17.5 16.2 18.5 17.2 16.4 16.2 18.8 17.7 18.5 18.7 17.6 16.1 17.3 17.2 17.6 19.8 17.2 17.5 17.1 17.1 19 17 18.3 20.5 19.9 19.5 16.7 17.2 20.3 16.2 19.9 18.2 17.8 16.3 19.8 18.5 19.5 18.4 19.1 18.1 19 19 20 16.3 16.9 16.2 17.5 18.6 16.5 18.8 16.8 17.8 16.5 50.6 14.8 16.4 17.5 15.5 19.1 17.1 15.2 15.4 15.7 17.3 19.1 14.6 10.8 13.5 13.5 14.4 14.4 15.3 14.2 14.3 12.9 15.1 11.8 14.3 14.4 14.5 15.7 15.7 15.4 12.3 15.7 15.1 15 10.6 10.5 9.9 12.2 9.3 8.5 8.6 6.9 9.2 8.5 12.7 8.8 6.7 7.9 9.1 5.3 ∞. 9.0-4.8 9.9 6.9 4.8 6.5 9.9 2.8 3.1 2.5 3.4 2.1 3.9 4.2 3.7 4.6 1.2 2.1 6 0 5 4 -3.6 -4.5 2.6 -1.8 3.6 -4.7 0.2 0.5 -2.5 4.5 3.5 1.6 1:3 1.4 1.6 1:3 0.7 1.3 7 0 \sim **-**0.4 -0.2 -4.3 -3.6 -2.4 -2.5 0.7 -1.7 1.4 9.0 3.3 4.6 -4.1 2.3 1.3 7 ٠ 1909 1903 1905 1907 1910 1912 1913 1914 1915 1916 1917 1918 1919 1920 1922 1924 1925 1921 1911

Fable 1.

10.9 10.8 10.1 10.4 10.6 10.3 10.3 10.3 8.6 8.0 10.1 8.0 8.3 9.3 9.8 9.6 8.1 18.5 18.6 19.1 18.5 19.4 19.6 18.5 18.6 18.8 18.2 19.2 18.5 18.9 19.1 19.1 19.1 18.1 18 19 8 10.6 10.6 10.3 10.7 10.2 10.1 8.2 8.9 6.9 8.2 9.7 8.9 8.3 9.1 -0.8 -6.8 9.0--0.5 -0.3 -6.8 -4.2 -0.4 -6.3 -3.5 0.8 -1.8 0.8 0.3 0.8 1.3 -1.1 1.3 1.3 0.3 1.4 16.8 15.6 16.8 14.8 15.5 14.7 15.5 15.3 16.2 15.4 14.4 15.9 16.3 16.6 15.1 14.7 14.7 16.1 15.7 16 15 16.1 17 10.8 9.5 9.8 10.1 8.6 8.7 7.6 9.5 8.5 9.2 9.2 9.6 9.7 7.3 7.9 8.3 9.3 9.2 9.7 6 ∞ -4.3 -0.8 -0.2 9.0--0.3 -4.7 9.0 0.2 4.5 1.8 -1.5 -1.5 -3.2 -1.4 0.8 1.4 2.3 0.2 1.7 \sim Ϋ́ 5.7 6.3 3.9 5.6 7.7 3.8 4.5 3.4 5.5 2.7 2.4 6.1 1.2 4 4 10.9 10.3 10.1 10.7 10.7 10.7 9.9 8.9 9.2 9.6 9.2 9.8 9.6 11.1 7.5 8.6 8.3 9.5 6.2 8.9 5.7 0 16.3 15.8 16.3 17.6 14.6 15.1 14.4 11.6 15.4 14.2 17.9 14.5 15.5 15.7 15.3 17.1 14.2 15.1 14.1 5 4 20.7 18.9 18.3 17.5 19.3 18.2 18.3 17.2 20 20 20.1 16.8 17.8 20.2 20.7 17.6 18.1 18.1 17.5 17.1 18.7 16. 19.8 20.8 19.5 20.5 20.4 20.2 20.2 20.2 19.3 19.4 18.5 19.4 19.9 19.6 19.6 19.7 20.2 18.8 19.9 18.8 19 20 15.8 16.5 19.8 17.8 16.3 19.4 18.1 19.3 17.9 16.5 16.7 16.8 17.5 17.9 19.4 17.4 15.7 15.7 16.1 19 11.6 14.9 17.8 13.3 14.8 16.4 13.2 17.3 13.2 15.5 13.5 16.7 13.4 11.2 12.1 15 12 14 15 15 10.6 10.9 10.7 8.8 8.6 9.8 6.2 9.9 8.9 11.2 6.5 8.9 9.2 ∞ 9 ∞ -0.5 6.8 0.3 4.8 7.8 4.9 7: 0.1 6.2 1.8 1.9 3.5 1.1 4.5 3.3 -9 9 4 \sim 7 9.0-6.8 -0.3 -9.2 0.1 -0.5 -13 0.1 1.1 -0.1 1.3 1.5 1.5 6.0 3.4 2.1 3.1 'n m -10.9 -8.8 -0.3 -5.6 -6.7 1.6 0.7 -2.2 3.3 -4.1 9.0 2.2 3.4 2.1 <u>'</u>-÷ 0 1940 1944 1926 1936 1939 1948 1928 1929 1932 1933 1934 1935 1937 1941 1945 1927 1931

Fable 1. co

10.3 10.2 10.3 10.4 10.7 9.2 9.5 9.8 8.7 8.0 9.7 10.1 9.1 의 ∞ 7 18.5 18.8 19.3 18.1 17.8 18.3 18.1 18.2 17.4 19.2 17.4 17.2 17.1 17.1 19 18 19 18 ∞. 18. 8 8.6 7.6 9.2 8.2 8.4 9.9 7.8 9.9 9.9 8.3 9.2 9.5 8.9 10 7.9 7.5 9 \sim -6.9 9.0--0.7 -0.9 9.0-**4.**0--2.5 -0.7 -2.4 -4.1 0.7 0.7 0.4 1.3 -0.1 0.1 7: 4-Ÿ 7: 7 0 0 14.8 14.4 14.6 14.4 15.5 16.2 15.2 16.1 14.9 14.6 15.4 15.3 15.8 15.4 14.3 15.2 15.5 15.3 14.7 15.3 15.7 15.1 4 8.9 9.6 9.4 9.7 8.4 9.7 7.2 9.2 8.8 9.2 8.9 9.3 8.4 9.3 9.6 8.9 8.3 8.1 8.1 8.1 7.7 ∞ -0.8 -0.4 -2.8 -3.5 -2.8 -5.8 8.0 3.5 -3.1 0.5 6.4 3.4 -1.1 2.5 1.4 2.5 2.2 2.7 1.1 1.7 1.1 7 4 4.6 6.4 5.6 4.8 4.6 4.8 5.8 3.9 3.1 4.1 3.7 4.1 3.7 6.1 3.1 9 2 ∞ 7 12.3 10.5 11.3 12.4 11.3 7.4 7.5 9.8 8.6 8.8 10.1 10.1 8.3 8.9 7.6 9.6 9.2 9.2 9.4 ∞ ∞ 6 ∞ 15.8 14.9 12.8 16.2 14.6 14.6 11.9 15.7 14.7 14.3 13.4 15.9 13.2 14.9 13.3 12.2 14.5 16.1 14.2 12.3 15.2 13.7 14 18.6 19.3 17.6 18.3 16.3 16.6 17.8 18.4 16.8 16.6 17.9 19.7 18.2 17.7 17.4 17.4 17.9 17.7 17.3 17.7 19.7 19 17 19.1 17.8 18.8 19.9 19.7 20.9 19.3 18.5 19.9 19.5 20.7 17.2 17.2 20.1 19.5 17.6 18.4 19.9 18.1 19.8 18.8 19 18.8 16.6 18.7 16.4 16.4 18.9 16.3 16.3 15.4 18.1 19.4 17.7 16.2 19.2 17.3 17.7 18.1 18 18.1 17.2 18 18 17 13.8 15.6 12.6 12.4 13.9 13.9 11.5 14.9 15.9 13.6 13.9 13.2 15.5 11.9 12.2 12.6 14.7 11.4 11.3 14.1 13.7 12 14 10.8 10.4 11.6 10.1 8.8 6.3 9.4 7.3 9.3 9.8 8.2 5.6 5.7 5.9 7.7 7 ∞ 9.0-9.9 -0.4 9.0 4.6 4.9 1.9 3.8 0.2 1.6 4.9 6.3 3.7 6.4 0.7 0 7 2 4 -11.7 9.0--2.6 -2.3 -7.6 -0.2 -1.8 -0.2 -6.7 -3.3 1.5 3.2 1.6 3.2 -1.4 2.3 1.7 1.7 2.1 1.1 ņ -0.9 -0.3 -9.5 -3.3 -2.4 -0.7 0.2 0.7 -2.7 0.1 0.3 -1.8 0.7 -3.1 -4.5 -4.1 1949 1955 1956 1958 1959 1960 1962 1963 1970 1952 1953 1954 1957 1961 1967 1951 1971

Table 1.

10.9 10.2 8.4 9.3 9.6 9.6 9.8 9.5 9.1 ∞ 16.8 18.2 18.3 17.3 18.4 17.6 18.2 17.4 18.3 17.6 17.7 17.1 19.1 18 19 ∞. 18 18. 9 10.6 10.2 8.6 8.4 8.8 9.9 9.8 8.6 9.6 9.2 9.2 6.7 9.6 9.7 8.3 10.1 8.1 9 6 -4.3 -1.6 -2.8 -0.3 3.6 9.4 -2.4 0.5 0.5 0.2 0.1 1.5 0.2 2.6 3.1 0.5 2.1 0 15.6 16.8 16.3 14.7 14.9 14.2 13.9 15.2 16.2 14.4 15.6 15.2 15.1 15.1 13.7 15.7 14.7 15.7 5 5 10.3 10.4 10.2 10.3 8.6 8.6 9.5 9.9 8.7 9.5 9.6 9.2 8.7 7.9 9.2 8.6 7.9 9.7 9.1 9.3 ∞ 6 -0.1 -0.5 9.0-0.0 0.3 6.0 -1.7 -0.7 -0.2 2.9 2.3 4.2 1.2 3.4 0.2 3.2 2.1 7 4 7 4.5 4.8 4.8 5.6 3.8 5.6 2.6 5.3 0.3 5.7 4:1 3.2 1.6 9 10.5 10.2 11.6 9.9 11.2 9.9 8.5 9.6 9.6 9.3 8.9 9.6 9.6 9.6 8.4 9.7 7.7 10 6 16.6 14.6 14.8 15.8 15.2 13.3 12.3 13.2 14.5 13.5 13.6 13.9 14.7 15.3 17.2 12.2 15.2 15.7 12.7 14 15 14 18.8 18.9 16.2 19.2 19.3 16.6 18.9 19.4 17.2 18.4 17.2 16.7 17.7 17.4 17.9 18.1 17.7 18.4 17.7 19 19 18.1 20.1 18.8 19.5 19.5 17.3 16.7 17.4 18.5 20.1 20.9 18.9 18.5 19.5 19.2 18.7 20.4 22.5 17 19 16.9 15.5 16.8 17.9 16.3 19.5 18.1 17.6 17.5 15.8 15.2 16.9 16.4 16.5 16.1 19.5 17.4 17.1 17 17 10.8 14.8 13.9 14.6 15.6 16.5 13.5 13.5 14.3 14.7 13.3 15.3 14.4 15.1 10.7 14.9 13.4 12.1 13.4 12.7 15.3 13 12 8.8 8.8 10.1 8.4 6.9 7.5 6.7 9.3 6.7 ∞ 6 6 9.0-3.6 6.9 8.3 5.9 4.9 6.4 0.9 5.2 4.3 2.3 5.2 5.4 2.7 4.9 6.2 ^ $\boldsymbol{\omega}$ 3 / 9.0--0.4 -2.6 -0.7 -0.3 -0.7 0.8 0.2 3.1 2.1 0 0 -5.1 4.2 5.9 -3.2 -0.4 -3.2 2.2 4.2 -4.1 -7.7 -9.1 2.2 1.3 0.1 2.7 1980 1986 1978 1979 1982 1988 1990 1972 1973 1974 1975 1976 1977 1983 1984 1985 1987 1981 1991

Fable 1. cont

10.4 10.3 10.1 10.1 8.6 80.00 8.4 9.1 9.1 18.8 18.5 20.2 18.9 18.5 18.6 20.3 18.9 18.6 19.6 9 10.3 11.3 10.9 8.8 8.5 9.2 9.7 9.5 8.9 8.5 10 -1.4 -2.2 8.0 3.1 2.1 1.8 1.7 0.5 4.2 -7 $\overline{}$ 16.9 16.8 16.6 14.6 16.9 15.7 15.4 16.1 16.5 16.3 15.4 15.9 16.1 10.8 10.3 10.9 10.5 6.6 9.6 9.8 9.6 9.6 9.5 9.2 9.7 -2.6 -5.6 -3.7 0.2 2.9 -1.1 4.4 1.9 1.8 2.1 1.1 2 6.4 3.8 1.2 7.3 5.4 3,3 4 9 10.5 11.6 10.4 13.2 9.4 9.8 8.6 8.9 13 6.3 # 14.8 14.9 16.2 16.9 13.6 13.9 14.5 17.7 13.9 14.2 13 15 20.8 20.5 18.6 18.1 19.8 17.7 19.3 18.3 19.7 19.3 19.7 18.1 19.1 20.8 20.3 20.7 20.3 23.4 21.5 18.9 17.7 20.1 19.3 19.9 17.7 19 18.9 16.9 17.6 18.5 17.2 18.6 18.9 17.4 19.7 17.5 16.1 18.1 20 14.5 15.3 17.5 14.5 15.9 13.5 13.9 14.3 14.7 16.1 15.3 16.3 13.4 10.9 10.3 10.1 12.4 10.1 8.5 9.9 8.4 9.3 10.1 11.2 9.1 3.9 3.9 -0.1 4.7 5.9 5.5 3.9 4.2 7.1 9 7 7 -3.7 0.4 -2.7 4.9 3.5 5.2 4.1 1.5 5.2 2.3 ω -2.8 -5.6 -4.5 -3.9 0.9 2.3 -0.1 1.7 2.1 0 2006 1999 2000 2003 2004 1996 1998 2002 2007 1995 2001

Table 1. cont

values are for the normal year (Jan. – Dec.; ann.) and for the hydrological one (Nov. – Oct.; hydro); warm half-year values (Apr.-Sep.; A-S) and seasonal values (DJF, MAM, JJA, SON) are also presented. Explanation: bold italics for Table 2. Monthly (Jan., Feb., Mar.,...Dec.) values [mm] of precipitation totals for Wrocław for the years 1791–2007. Annual estimated or interpolated values, bold and colouring for extreme values (max - gold, min - blue)

2	10 H	7	7	1,017	2		۷۷	000	+	NON	200	2	hydro	V <		77 7 7 7	٧١١	NOU
	+	Mal.	Id	May	Juli.	Jul.	Aug.	oeb.	000	NOV.	חבר:	٠ ا	ııyaı o	7-Y	707	MAM		SON
- 1	56	56	38	54	81	80	45	24	23	29	25	484		322		118	206	9/
35	56	34	42	18	48	40	30	82	19	30	30	434	428	260	86	94	118	131
25	21	25	28	39	81	56	99	56	38	34	28	467	465	296	92	92	173	128
97	31	27	42	34	27	80	99	63	45	31	18	490	503	312	85	103	173	139
31	29	32	33	49	75	87	59	31	31	34	37	528	506	334	78	114	221	96
18	23	27	18	70	108	67	102	63	34	30	35	595	601	428	78	115	277	127
29	18	22	42	39	61	73	52	37	31	32	42	478	469	304	82	103	186	100
29	27	35	23	65	55	73	110	24	16	35	33	272	531	958	86	123	238	75
15	6	18	54	62	186	86	35	40	16	8	9	535	589	463	25	134	307	64
10	16	8	40	53	66	60	45	59	43	40	23	496	447	356	32	101	204	142
14	10	35	68	27	112	94	51	90	57	22	24	604	621	442	47	130	257	169
25	27	56	27	67	34	56	55	20	18	45	37	437	401	259	92	120	145	83
19	37	19	12	139	74	50	30	16	40	22	32	490	518	321	66	170	154	78
50	24	19	44	20	41	29	27	26	38	18	30	366	372	187	106	83	97	82
12	50	19	22	15	27	19	36	22	38	22	27	309	308	141	95	26	82	82
20	13	44	20	22	35	35	22	29	7	56	9	309	296	163	9	86	92	95
70	8	15	23	33	46	8	46	49	22	89	31	419	382	205	84	71	100	139
35	33	18	97	23	113	30	41	119	29	22	30	590	637	423	66	138	184	170
19	30	19	34	18	27	19	35	57	56	33	29	346	336	190	79	71	81	116
4	38	29	20	41	56	44	31	12	22	62	70	436	366	174	108	96	101	96

56																						
	130	134	65	6	123	131	62	22	69	85	50	47	84	64	25	64	56	140	135	183	89	148
96	205	120	128	205	236	166	116	186	144	176	187	113	128	156	138	207	218	338	210	249	144	311
45	129	115	117	76	71	99	114	100	127	143	94	115	66	90	110	116	66	138	124	104	62	152
98	59	46	69	63	54	42	55	55	46	105	92	81	63	69	31	87	71	09	9	48	90	45
127	293	250	251	297	340	321	243	289	264	319	270	226	235	259	249	279	304	536	393	413	201	553
338	489	435	393	427	490	440	363	407	381	509	423	344	358	383	354	467	481	672	534	526	399	669
247	535	397	379	443	480	446	346	421	397	510	384	373	379	352	353	471	488	673	529	628	311	671
22	34	16	16	18	14	22	21	24	35	36	13	30	35	8	25	22	27	24	24	89	15	30
19	53	33	19	33	27	25	6	20	25	25	6	21	37	33	15	22	24	28	23	81	46	3
29	55	48	15	23	41	28	9	16	15	6	6	11	21	8	28	33	43	44	27	4	20	12
8	22	53	31	41	55	78	47	41	29	51	32	15	56	23	14	6	28	89	85	98	2	133
20	31	44	16	90	67	85	50	93	7	40	109	7	59	49	34	124	75	55	28	80	10	70
49	85	45	44	63	93	67	38	54	74	44	57	33	47	38	48	24	95	166	80	79	75	108
27	89	31	89	52	9/	14	28	39	63	95	21	73	52	69	99	59	51	117	102	90	59	133
6	44	40	46	35	41	99	40	48	64	51	28	47	38	64	54	29	16	9	42	41	50	13
14	22	37	46	16	8	11	40	14	27	41	23	51	43	16	43	34	42	70	99	25	5	96
22	63	38	25	25	22	22	34	38	36	51	43	17	18	10	13	53	41	8	56	38	7	43
41	14	∞	27	14	16	14	13	13	6	38	5	39	20	16	7	11	22	17	22	6	0	27
41	23	4	56	33	20	14	20	21	13	32	35	29	13	18	16	51	27	16	14	15	22	3
1811	1812	1813	1814	1815	1816	1817	1818	1819	1820	1821	1822	1823	1824	1825	1826	1827	1828	1829	1830	1831	1832	1833

Table 2. cont.

Fable 2.

51 59 54 558 ∞ 4 | 33 | 25 | 5 4 \sim 9 35 24 70 50 40 68 57 85 57 66 80 39 68 ∞ 9 42 12 14 = ω ∞

9/ 33 33 5 5 37 50 4 5 25 19 7 17 =9/ 85 83 76 ∞ ∞ 31 ∞ =6 4 \mathcal{L} \sim ∞

Table 2. cor

9/1 9/ 50 6 8 58 32 23 37 12 85 24 10 56 72 8 85 74 12 57 33 33 45 62 62 58 33 29 29 9 8 53 **85** 4 5 8 ∞ 0 8 =

Fable 2. co

4 5 35 17 18 ∞ ∞ \sim 69 36 =9/ 55 45 9/ 95 55 83 55 46 31 27 36 37 13 25 9 4 35
 45

 5

 5

 5

 5

 6

 7

 7

 8

 9

 10

 11

 12

 12

 12

 13

 14

 15

 16

 17

 17

 18

 18

 19

 10

 10

 10

 11

 12

 12

 13

 14

 15

 16

 17

 17

 18

 19

 10

 10

 10

 10

 10

 11

 12

 12

 13

 14

 15

 16

 17

 17

 18

 18

 19

 10

 10

 11

 12

 12

 12

 12

 12

 12

 12

 12

 13

 14

 15

 16

 17

 18

 18

 18

 18

 19

d

Table

9/1 27 35 40 23 17 16 37 38 28 188 73 35 35 3 81 75 21 5 6 6 90 47 62 59 93 82 52 36 55 25 61 33 28 28 7 25 70 13 56 ∞ ∞ 9 25

Fable 2. con

35 30 31 39 53 46 37 97 m 23 35 33 50 49 39 36 75 139 33 33 101 9/ 87 47 63 4 8 _ _ ∞ ∞ 31 6 / ∞

Table 2

86 9/ 520 5 5 31 31 32 25 25 28 33 33 33 26 39 25 17 15 37 19 39 81 ¹⁴ 97 39 62 32 76 19 47 Ħ 34 170 48 35 55 97 12 67 67 93 84 88 56 72 19 54 2 2 9/ 25 51 7 4 ∞ / 33 33 ∞

Fable 2. co

182 35 73 58

Table 2. cont

Summary - More on the Influence of UHI

The biggest difficulty for the reconstruction of an homogenous 217-year Wrocław series, taking into consideration changes of location and missing data for 1945, was to reach a credible reconstruction of the Tp and P values for the years 1931–2007. The reconstructed values after 1930 mainly on the basis of other Wrocław stations, were referenced to and incorporated into the earlier Breslau-Sternwarte series as their most probable continuation. However, these series represent the specific and variable environment in the center of the city. Consequently, the influence of UHI on the values of basic meteorological elements in Wrocław must be taken into consideration and requires a more detailed discussion.

UHI most likely began to exert a greater impact in the 1980s, at which time Wrocław began to undergo dynamic industrial and urban development. Earlier, temperature differences between the city and its peripheries were not very great. Galle (1879) noted similarities in temperature data from the western border of Wrocław and data coming from the Breslau-Sternwarte station. He had access to results from 5-year simultaneous measurements in Goldschmieden (now Złotniki in the western district of Wrocław) in the years 1868–1873. Comparing the results of monthly average Tp, calculated on the basis of extreme temperatures, he was able to deduce their convergence due to the fact that the difference did not exceed 0.2° C (in spite of the stations being 10 km apart).

In the early 1920s, Wrocław reached a maximum population density, particularly in the Old City, in the area adjacent to the Breslau-Sternwarte Observatory. After 1925, a long-term process of population and industrial deconcentration began as people moved from the central part of the growing city towards its suburbs. The influence of UHI on the meteorological values from Sternwarte probably stabilized at the beginning of this process. In this situation, a credible estimation of the *Tp* and *P* values before 1925 for existing meteorological stations (Strachowice, Swojec) located at the peripheries of Wrocław is very complicated and is still open to discussion. However, such a reconstruction would be invaluable in order to separate the influence of UHI on tendencies of values measured in the city.

Acknowledgments

This work has been in partly supported by the grant N 30507332/2594 for 2007–2009 financed by the Ministry of Science and Higher Education – MNiSW (Poland).

The Authors would like also to thank: Rudolf Brázdil (Institute of Geography, Masaryk University, Brno) for temperature data (1771–2005) from Prague (Praha-Klementinum) and discussions about homogeneity of Wrocławian temperature data; Krzysztof Migała (Department of Meteorology and Climatology, Wrocław University) for the use in this study of the *Tp* and *P* data from measurements made in the years 1996–2007 at the University Meteorological Observatory Wrocław-Biskupin.

References

- ALEXANDERSSON H., 1986, A homogeneity test applied to precipitation data, J. Climatol., 6, 661–675.
- ATLAS KLIMATYCZNY POLSKI. Temperatura Powietrza. Część tabelaryczna 2a, 1979, IMGW, WKiŁ, Warszawa, 78 pp.
- BIULETYN AGROMETEOROLOGICZNY, 1965...(1966, 1967...2007). IMGW, Warszawa, 1965–2007.
- BIULETYN METEOROLOGICZNY, 1971...(1972, 1973, 1974), Pyka J.L. (ed.), Acta Univ. Wratislav., No. 287, 305, 569, 768, Wrocław 1984–1988.
- Bràzdil R., Stepanek P., 1998, Kolisani teploty v Brne v obdobi 1891–1995, Geografie Sbornik Ceske Geograficke Spolecnosti, 103, 13–30.
- Bryś K., Bryś T., 2007, Zmienność warunków solarnych klimatu Wrocławia w latach 1875–2004, Pamiętnik Puławski, 144, 13–33.
- Bryś K., Bryś T., 2010, The first one hundred years (1791–1890) of the "Wrocław temperature series", [in:] Przybylak R., Majorowicz J., Brázdil R., Kejna M. (eds.). The Polish Climate in the European context: An Historical Overview. Springer, Dordrecht, 485–524.
- Bryś T., 2007, Badanie reprezentatywności Obserwatorium Wrocław-Swojec jako stacji agrometeorologicznej i klimatologicznej, [in:] Kostrzewski A. & Andrzejewska A. (eds.) Zintegrowany monitoring środowiska przyrodniczego, KPN, Warszawa, 255–268.
- Deutsches Meteorologisches Jahrbuch, 1934 ... (1935, 1936 ...1944), Deutsches Reich, Reichsamt für Wetterdienst, Julius Springer, Berlin 1936...(1937, 1938, 1939): Wetterdienst, Offenbach/Main 1950.

- Dubicka M., 1994, Wpływ cyrkulacji atmosfery na kształtowanie warunków klimatu (na przykładzie Wrocławia), Studia Geograficzne, LX, Acta Univ. Wratisl., 1881, 296 pp.
- Dubicka M., Pyka J.L., 2001, Klimat Wrocławia w XX wieku, Prace i Studia Geogr., 29, 101–112.
- Ergebnisse Der Beobachtungen And Den Stationen II. und III, Ordnung in den Jahren 1914–1918...(1919–1923): im Jahre 1924...(1925, 1926...1933), Veröffentlichungen des Preussischen Meteorologischen Instituts, Julius Springer (Behrend & Co.), Berlin 1925...(1926, 1927...1934).
- Ergebnisse Der Meteorologischen Beobachtungen im Jahre 1891... (1892, 1893...1913), Königlich Preussischen Meteorologischen Instituts, Behrend & Co., Berlin 1892 ... (1893, 1894...1914).
- FLOHN H., 1981, Life on a Warmer Earth Possible Climatic Consequences of Man-Made Global Warming. IIASA Executive Report ER-81–003, Laxenburg, Austria, 75 pp.
- FLOHN H., 1993, Climatic evolution during the last millennium; what can we learn from it? [in:] Eddy J. A., Oeschger H. (eds.), Global Changes in the Perspective of the Past, John Wiley & Sons Ltd., 295–316.
- Galle J.G., 1857, Grundzüge der Schlesischen Klimatologie, Josef Max & Komp., Breslau, XXIII+128 pp.
- Galle J.G., 1879: Mittheilungen der Königlichen Universitäts-Sternwarte zu Breslau, Maruschke und Berendt (print), Breslau, 128 pp.
- KLIMAKUNDE DES DEUTSCHEN REICHES, Bd. II, Tabellen, 1939, Verlag von Dietrich Reimert /Andrews & Steiner/, Berlin, 560 pp.
- Landsberg H. E., 1981, The Urban Climate, Academic Press, New York, x + 275 pp.
- LORENC H., 2000, Studia nad 220-letnią (1779–1998) serią temperatury powietrza w Warszawie oraz ocena jej wiekowych tendencji, Mat. Badaw. IMGW, Meteorologia, 31, Warszawa, 104 pp.
- Miętus M., 1998, O rekonstrukcji i homogenizacji wieloletnich serii średniej miesięcznej temperatury ze stacji w Gdańsku-Wrzeszczu, 1851–1995, Wiad. IMGW, XXI (XLII), 2, 41–63.
- PISOFT P., KALVOWA J., BRAZDIL R., 2004, Cycles and trends in the Czech Temperature Series using wavelet transforms, Int. J. Climat., 24: 1661–1670.
- Prace Obserwatorium Meteorologii I Klimatologii Uniwersytetu Wrocławskiego nr 1 (1946) nr 14 (1958), WTN, nr 15 (1959) nr 20

- (1965), Kosiba A. (ed.), Acta Univ. Wratisl., No. 105, 115, 130, 159, 187, 206, 255, Wrocław 1947–1976.
- Pyka J.L., 1991, Temperatura i opady atmosferyczne we Wrocławiu w latach 1881–1980, Acta Univ. Wratisl., No. 1237, Prace Inst. Geogr., Wrocław, Seria A, VI, 19–54.
- Pyka J.L., 1998a, Temperatura powietrza we Wrocławiu w latach 1881–1995, Acta Univ. Wratisl., 2022, Prace Inst. Geogr., Wrocław, Seria C, Met. i Klimat., V, 25–40.
- Pyka J.L., 1998b, Opady atmosferyczne we Wrocławiu w latach 1881–1995, Acta Univ. Wratisl., 2022, Prace Inst. Geogr., Wrocław, Seria C, Met. i Klimat., V, 41–54.
- Pyka J.L., 2003, Meteorological observations and measurements in Wrocław, [in:] Pyka J. L., Dubicka M., Szczepankiewicz-Szmyrka A., Sobik M., Błaś M. (eds), Man and climate in the 20th century, St. Geogr., 75, 11–22.
- SCHMUCK A., 1967, Wpływ miasta na opady atmosferyczne (na przykładzie Wrocławia). Przegl. Geof., 12 (20), 3–4, 293–310.
- SZYMANOWSKI M., 2005: Miejska wyspa ciepła we Wrocławiu, Acta Univ. Wratisl., 2690, Studia Geogr., 77, 228 pp.
- Trepińska J., Kowanetz L., 1997, Wieloletni przebieg średnich miesięcznych wartości temperatury powietrza w Krakowie (1792–1995), [in:] Trepińska J. (ed.), Wahania klimatu w Krakowie (1792–1995), UJ, Kraków, 99–130.
- Twardosz R., 1997, Homogenizacja serii pomiarów opadów atmosferycznych na Stacji Meteorologicznej w Krakowie, [in:] Janina Trepińska (ed.), Wahania klimatu w Krakowie (1792–1995), UJ, Kraków, 89–95.
- Rocznik Meteorologiczny 1954...(1955, 1956...1965), IMGW, WKiŁ, Warszawa 1959...(1960–1966, 1968, 1970, 1974).
- ZIPSER-URBAŃSKA A., 1968, Wpływ kierunku wiatru na rozkład opadów atmosferycznych w mieście na przykładzie Wrocławia, Czasopismo Geogr., 39, 429–437

Internet Sources:

ftp://ftp.ncdc.noaa.gov/pub/data/gsod/

http://home.casema.nl/errenwijlens/co2/homogen.htm

http://old.chmi.cz/meteo/

http://www.dwd.de/

http://www.kni.nl/klimatologie

http://www.wetterzentrale.de/klima/

http://www.zmag.ac.at/klima/