# A1. CARIACO model equations

The ecosystem model equations are similar to those used in Acevedo-Trejos et al. (2016). Most significant changes are that multiple phytoplankton and zooplankton functional types have been added, and that the grazing formulation was expanded to include preferential feeding on certain functional types.

Nitrogen N (and Silicate Si for Diatoms) is assimilated by the phytoplankton types  $P_i$ , which are grazed by several zooplankton types  $Z_j$ . Mortality of and excretion from plankton, and sloppy feeding by zooplankton contribute to a Detritus D. The phytoplankton types include Nanoflagellates  $P_{Nano}$ , Diatoms  $P_{Dia}$ , Coccolithophorse,  $P_{Cocco}$  and Dinoflagellates  $P_{Dino}$ . There are two Zooplankton types split by size class, names Mikrozooplankton  $Z_{mikro}$  and Mesozooplankton  $Z_{meso}$ .

$$\frac{\partial N}{\partial t} = K \cdot (N_0 - N) + D \cdot \delta_D^N - \sum_{i=1}^{i_{max}} [Gains_i \cdot P_i]$$

$$\frac{\partial Si}{\partial t} = K \cdot (Si_0 - Si) - Gains_{Dia} \cdot P_{Dia}$$

$$\frac{\partial P_{Dia}}{\partial t} = Gains_{Dia} \cdot P_{Dia} - \sum_{j=1}^{j_{max}} [I_j^{tot} \frac{p_{Diaj} P_{Dia}}{R_j} Z_j]$$

$$\frac{\partial P_{Nano}}{\partial t} = 1$$

$$\frac{\partial Z_{mikro}}{\partial t} = 1$$

$$\frac{\partial D}{\partial t} = 1$$

#### where:

 $\begin{array}{l} \mathbf{u} = (u,v,w), \text{ velocity in physical model,} \\ \kappa = \text{Mixing coefficients used in physical model,} \\ \mu_j = \text{Growth rate of phytoplankton } j \text{ (see below),} \\ M_{ij} = \text{Matrix of Redfield ratio of element } i \text{ to P for phytoplankton } j \\ \zeta_{jk} = \text{Grazing efficiency of zooplankton } k \text{ on phytoplankton } j \text{ (represents sloppy feeding),} \\ g_{jk} = \text{Grazing of zooplankton } k \text{ on phytoplankton } j \text{ (see below),} \\ m_j^P = \text{Mortality/Excretion rate for phytoplankton } j, \\ m_k^Z = \text{Mortality/Excretion rate for zooplankton } k, \\ w_j^P = \text{Sinking rate for phytoplankton } j, \\ w_{POi} = \text{Sinking rate for POM } i, \\ r_{DOM_i} = \text{Remineralization rate of DOM for element } i, \\ r_{POM_i} = \text{Remineralization rate of POM for element } i, \\ S_{N_i} = \text{Additional source or sink for nutrient } i \text{ (see below),} \\ S_{DOM_i} = \text{Source of DOM } i, \text{ for element } i \text{ (see below),} \\ \end{array}$ 

 $S_{POM_i} =$ Source of POM i, for element i (see below),

### A1.1. Phytoplankton growth:

$$\mu_j = \mu_{max_j} \gamma_j^T \gamma_j^I \gamma_j^N$$

where

 $\mu_{max_{j}} = \text{maximum growth rate of phytoplankton } j,$ 

 $\gamma_i^T$  =Modification of growth rate by temperature for phytoplankton j,

 $\gamma_j^I$  =Modification of growth rate by light for phytoplankton j,

 $\gamma_i^N = Modification of growth rate by nutrients for phytoplankton j.$ 

Temperature modification (Fig. ??a):

$$\gamma_j^T = \frac{1}{\tau_1} (A^T e^{-B(T - T_o)^c} - \tau_2)$$

where coefficients  $\tau_1$  and  $\tau_2$  normalize the maximum value, and  $A, B, T_o$  and C regulate the form of the temperature modification function. T is the local model ocean temperature.

Light modification (Fig. ??b):

$$\gamma_j^I = \frac{1}{F_o} (1 - e^{k_{par}I}) e^{-k_{inhib}I}$$

where  $F_o$  is a factor controlling the maximum value,  $k_{par}$  is the PAR saturation coefficient and  $k_{inhib}$  is the PAR inhibition factor. I is the local PAR, that has been attenuated through the water column (including the effects of self-shading).

Nutrient limitation is determined by the most limiting nutrient:

$$\gamma_j^N = \min(N_i^{lim})$$

where typically  $N_i^{lim}=\frac{N_i}{N_i+\kappa_{N_{ij}}}$  (Fig. **??**c) and  $\kappa_{N_{ij}}$  is the half saturation constant of nutrient i for phytoplankton j.

When we include the nitrogen as a potential limiting nutrient (EXP2) we modify  $N_i^{lim}$  to take into account the uptake inhibition caused by ammonium:

$$N_N^{lim} = \frac{NO_2}{NO_2 + \kappa_{IN}} e^{-\psi NH_4} + \frac{NH_4}{NH_4 + \kappa_{NH4}} \tag{nsource=1}$$

$$N_N^{lim} = \frac{NH_4}{NH_4 + \kappa_{NH_4}}$$
 (nsource=2)

$$N_{N}^{lim} = \frac{NH_{4}}{NH_{4} + \kappa_{NH4}}$$
 (nsource=2) 
$$N_{N}^{lim} = \frac{NO_{3} + NO_{2}}{NO_{3} + NO_{2} + \kappa_{IN}} e^{-\psi NH_{4}} + \frac{NH_{4}}{NH_{4} + \kappa_{NH4}}$$
 (nsource=3)

where  $\psi$  reflects the inhibition and  $\kappa_{IN}$  and  $\kappa_{NH4}$  are the half saturation constant of  $IN = NO_3 +$  $NO_2$  and  $NH_4$  respectively.

### A1.2. Zooplankton grazing:

$$g_{jk} = g_{max_{jk}} \frac{\eta_{jk} P_j}{A_k} \frac{A_k}{A_k + \kappa_k^P}$$

where

 $g_{max_{jk}} = \text{Maximum grazing rate of zooplankton } k \text{ on phytoplankton } j,$ 

 $\eta_{ik}$  = Palatibility of plankton j to zooplankton k,

 $A_k = \text{Palatibility (for zooplankton } k)$  weighted total phytoplankton concentration,

 $=\sum_{j}[\eta_{jk}P_{j}]$ 

 $\kappa_k^P$  =Half-saturation constant for grazing of zooplankton k,

## A1.3. Inorganic nutrient Source/Sink terms:

 $S_{N_i}$  depends on the specific nutrient, and includes the remineralization of organic matter, external sources and other non-biological transformations:

$$S_{PO4} = r_{DOP}DOP + r_{POP}POP$$

$$S_{Si} = r_{POSi}POSi$$

$$S_{FeT} = r_{DOFe}DOFe + r_{POFe}POFe - c_{scav}Fe' + \alpha F_{atmos}$$

$$S_{NO3} = \zeta_{NO3}NO_2$$

$$S_{NO2} = \zeta_{NO2}NH4 - \zeta_{NO3}NO_2$$

 $S_{NH4} = r_{DON}DON + r_{PON}PON - \zeta_{NO2}NH_4$ 

where:

 $r_{DOM_i}$  =Remineralization rate of DOM for element i, here P, Fe, N,

 $r_{POM_i}$  =Remineralization rate of POM for element i, here P, Si, Fe, N,

 $c_{scav}$  =scavenging rate for free iron,

Fe' =free iron, modelled as in Parekh et al (2004),

alpha =solubility of iron dust in ocean water,

 $F_{atmos}$  =atmospheric deposition of iron dust on surface of model ocean,

 $\zeta_{NO3} = \zeta_{NO3}^0 (1 - I/I_0)_+ =$ oxidation rate of NO<sub>2</sub> to NO<sub>3</sub>,

 $\zeta_{NO2} = \zeta_{NO2}^0 (1 - I/I_0)_+ =$ oxidation rate of NH<sub>4</sub> to NO<sub>2</sub> (is photoinhibited),

 $I_0$  =critical light level below which oxidation occurs,

The remineralization timescale  $r_{DOi}$  and  $r_{POi}$  parameterizes the break down of organic matter to an inorganic form through the microbial loop.

#### A1.3.1 Fe chemistry:

$$Fe' = FeT - FeL$$

$$FeL = L_{tot} - \frac{L_{stab}(L_{tot} - FeT) - 1 + \sqrt{(1 - L_{stab}(L_{tot} - FeT))^2 + 4L_{stab}L_{tot}}}{2L_{stab}}$$

(Fe' may be constrained to be less than  $Fe'_{max}$  while preserving FeT).

#### A1.4 DOM and POM Source terms:

 $S_{DOM_i}$  and  $S_{POM_i}$  are the sources of dissolved and particulate organic detritus arising from mortality, excretion and sloppy feeding of the plantkon. We simply define that a fixed fraction  $\lambda_m$  of the the mortality/excretion term and the non-consumed grazed phytoplankton  $(\lambda_g)$  go into the dissolved pool and the remainder into the particulate pool.

$$\begin{split} S_{DOM_i} &= \sum_{j} [\lambda_{mp_{ij}} m_{j}^{p} P_{j} M_{ij}] + \sum_{k} [\lambda_{mz_{ik}} m_{k}^{z} Z_{ik}] + \sum_{k} \sum_{j} [\lambda_{g_{ijk}} (1 - \zeta_{jk}) g_{ij} M_{ij} Z_{k}] \\ S_{POM_i} &= \sum_{j} [(1 - \lambda_{m_{ij}}) m_{j}^{p} P_{j} M_{ij}] + \sum_{k} [(1 - \lambda_{mz_{ik}}) m_{k}^{z} Z_{ik}] + \sum_{k} \sum_{j} [(1 - \lambda_{g_{ijk}}) (1 - \zeta_{jk}) g_{ij} M_{ij} Z_{k}] \end{split}$$

### A1.4 Geider light limitation model:

The phytoplankton growth rate is given by the carbon-specific photosynthesis rate (rate of carbon synthesized per carbon present),

$$\mu_j = P_i^C$$

The carbon-specific photosynthesis rate

$$P_j^C = P_{m,j}^C \begin{cases} 1 - e^{-\alpha_j^{\rm Chl} I \theta_j / P_{m,j}^C} & \text{if } I > 0.1 \\ 0 & \text{otherwise} \end{cases}$$

depends on the carbon-specific, light-saturated photosynthesis rate

$$P_{m,j}^C = P_{\text{MAX}j}^C \gamma_j^N \gamma_j^T$$

and the Chl a to carbon ratio

$$\theta_j = \left[\frac{\theta_j^{\max}}{1 + \theta_j^{\max}\alpha_j^{\text{Chl}}I/(2P_{m,j}^C)}\right]_{\theta_i^{\min}}^{\theta_j^{\max}}$$

The chlorophyll concentration is

$$Chl_j = P_j R_j^{PC} \theta_j$$

The light limitation factor can be diagnosed

$$\gamma_i^I = P_i^C / P_{m,i}^C$$

$$\alpha_j^{\rm Chl} = mQ_j^{\rm yield}A_{\rm Chl,ave}^{\rm phy}$$

### Parameters:

ameters:  $P_{\mathrm{MAX}j}^{C} = \text{Maximum C-spec. photosynthesis rate at reference temperature of phytoplankton } j$   $\theta_{\mathrm{MAX}j}^{\mathrm{max}} = \text{Maximum Chl a to C ratio if phytoplankton } j$   $R_{j}^{PC} = \text{Carbon to phosphorus (!) ratio of phytoplankton } j$   $\alpha_{j}^{\mathrm{Chl}} = \text{Chl a-specific initial slope of the photosynthesis-light curve}$   $mQ_{j}^{\mathrm{yield}} = \text{slope of the photosynthesis-light curve per absorption}$   $A_{\mathrm{Chl,ave}}^{\mathrm{phy}} = \text{absorption } (m^{-1}) \text{ per mg Chl a}$ 

# A2 Diagnostics:

Total phytoplankton biomass:

$$P_{\text{tot}} = \sum_{j} P_{j}$$

name	definition		units
PhyTot	$P_{ m tot}$		$\mu MP$
PhyGrp1	Total biomass of small phytoplankton with ${ t nsrc}=1$		$\mu \mathrm{M}\mathrm{P}$
PhyGrp2	Total biomass of small phytoplankton with $nsrc = 2$		$\mu \mathrm{M}\mathrm{P}$
PhyGrp3	Total biomass of small phytoplankton with $nsrc = 3$		$\mu \mathrm{M}\mathrm{P}$
PhyGrp4	Total biomass of large non-diatoms		$\mu \mathrm{M}\mathrm{P}$
PhyGrp5	Total biomass of diatoms		$\mu \mathrm{M}\mathrm{P}$
PP	Primary production		$\mu\mathrm{M}\mathrm{P}\mathrm{s}^{-1}$
Nfix	Nitrogen fixation		$\mu \mathrm{M}\mathrm{N}\mathrm{s}^{-1}$
PAR	Photosynthetically active radiation		$\mu\mathrm{Ein}\mathrm{m}^{-2}\mathrm{s}^{-1}$
Rstar01	$R^*_{\mathrm{PO4}}$ of Phytoplankton species #1,		$\mu \mathrm{M}\mathrm{P}$
Diver1	Number of species with $P_j > 10^{-8}  \mu \mathrm{MP}$	where $P_{\rm tot} > 10^{-12}$	
Diver2	Number of species with $P_i > 0.1\% P_{\text{tot}}$	where $P_{\rm tot} > 10^{-12}$	
Diver3	Number of species that constitute 99.9% of $P_{ m tot}$	where $P_{\rm tot} > 10^{-12}$	
Diver4	Number of species with $P_j > 10^{-5} \cdot \max_j P_j$	where $P_{\rm tot} > 10^{-12}$	