Application of the Dirichlet problem (Upward continuation integral)

$$U(x_0,y_0,z_0) = \frac{1}{2\pi} \int \left(U(x_0,y_0,z_0) \left(-\frac{1}{2} \frac{1}{Y} \right) dx dy \right)$$

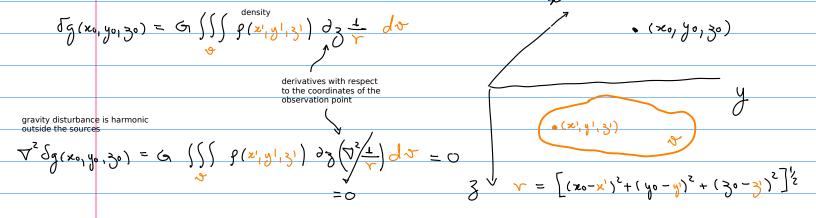
$$\frac{3\sqrt{1}}{1} = \left(-\frac{\lambda_3}{20-3c}\right)(-T) = -99\frac{\lambda_3}{1}$$

vertical component of the gravitational attraction produced by a point of mass

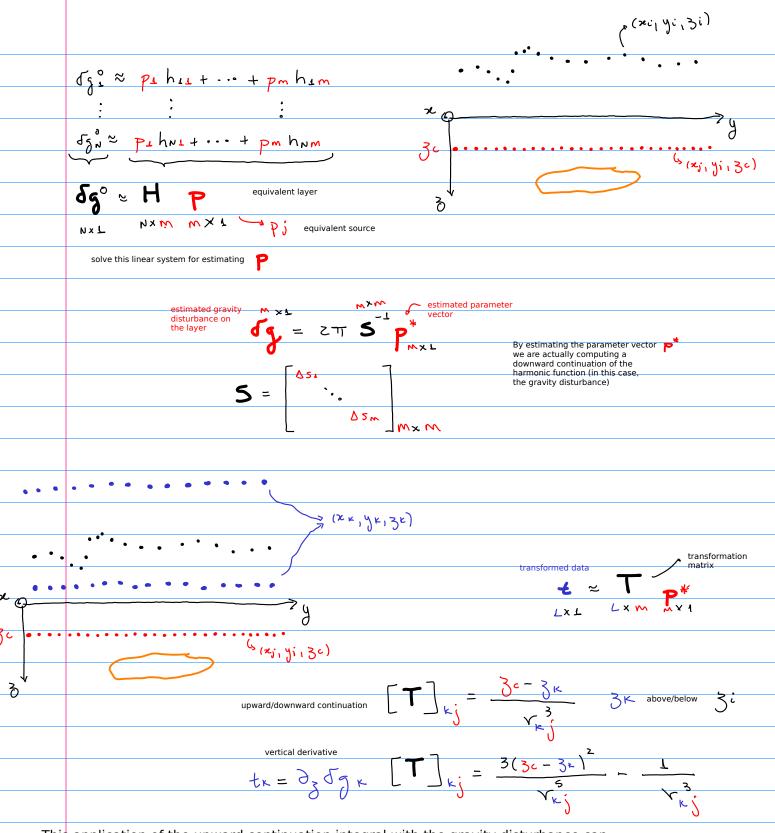
$$U(x_0, y_0, z_0) = \int \int P(x_1y_1z_0) dz \frac{1}{y} dx dy$$

$$P(x_1y_1z_0) = \frac{U(x_1y_1z_0)}{z_T}$$

Consider that U(1001190130) represents the gravity disturbance Jg(100130)



Consider now a set of gravity disturbance values at coordinates (\varkappa_i, y_i, z_i) , $z < z_c$



This application of the upward continuation integral with the gravity disturbance can be made with any harmonic function (i.e., gravitational potential and its derivatives, components of magnetic induction, approximated total-field anomaly, etc.)

References:

^{*} Dampney, C. N. G., 1969, The equivalent source technique: GEOPHYSICS, 34, 39–53. doi: 10.1190/1.1439996
* Leão, J. W. D., and J. B. C. Silva, 1989, Discrete linear transformations of potential field data: GEOPHYSICS, 54, 497–507. doi: 10.1190/1.1442676
* Mendonça, C. A., and J. B. C. Silva, 1994, The equivalent data concept applied to the interpolation of potential field data: GEOPHYSICS,

^{59, 722-732.} doi: 10.1190/1.1443630

* Oliveira Jr., V. C., V. C. F. Barbosa, and L. Uieda, 2013, Polynomial equivalent layer: GEOPHYSICS, 78, G1-G13. doi: 10.1190/geo2012-0196.1

* Siqueira, F., V. C. Oliveira Jr., and V. C. F. Barbosa, 2017, Fast iterative equivalent-layer technique for gravity data processing: A method grounded on excess mass constraint: GEOPHYSICS, 82, G57-G69. doi: 10.1190/GEO2016-0332.1

* Diego Takahashi, Vanderlei C. Oliveira Jr., and Valéria C. F. Barbosa, 2020, Convolutional equivalent layer for gravity data processing, GEOPHYSICS 2020 85:6, G129-G141. doi: 10.1190/geo2019-0826.1