



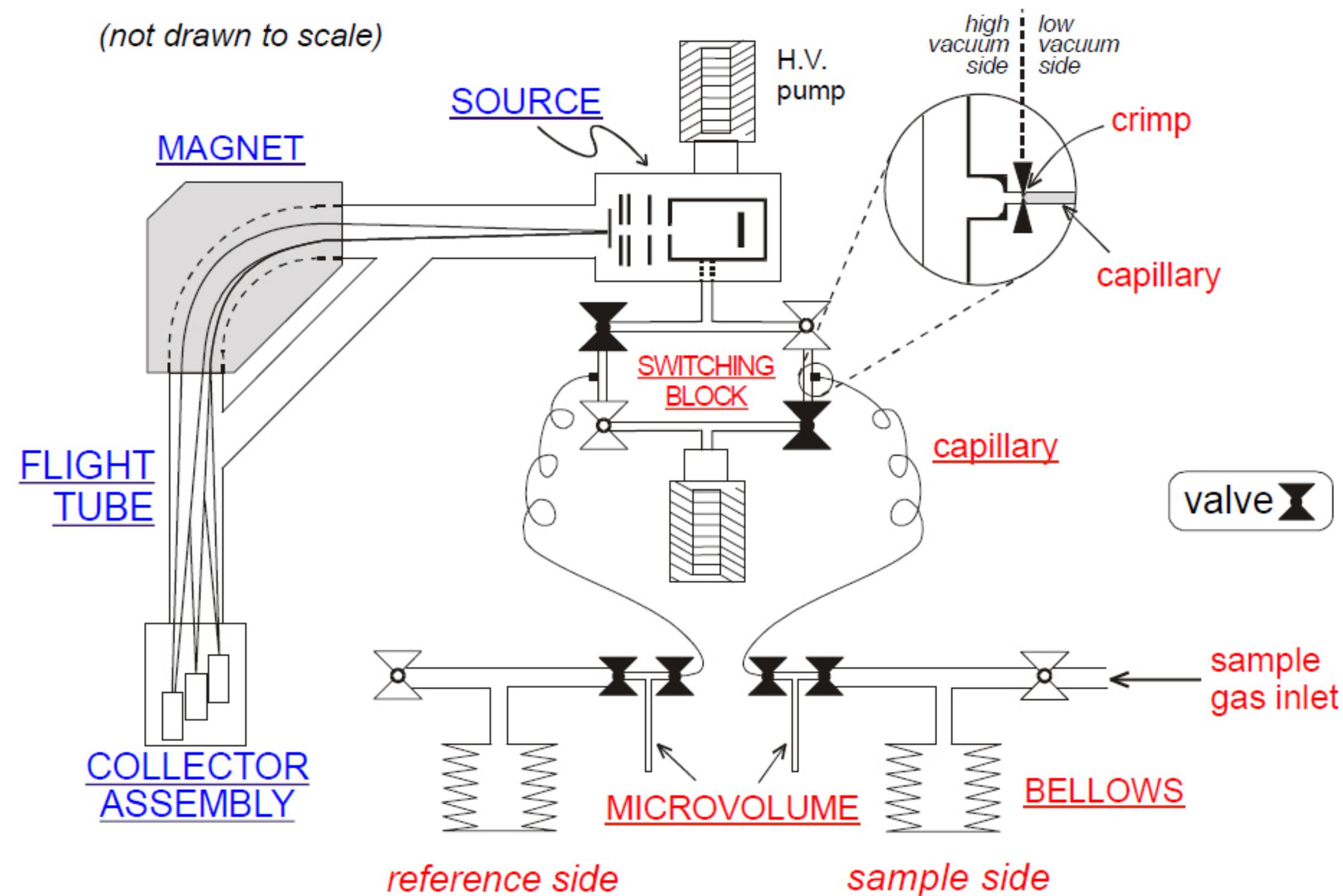
# Lecture 16: Stable isotopes

1. Equilibrium fractionation of stable isotopes
2. Pleistocene climate

*We acknowledge and respect the lək'ənən peoples on whose traditional territory the university stands and the Songhees, Esquimalt and WSÁNEĆ peoples whose historical relationships with the land continue to this day.*



*(not drawn to scale)*



*"Before the discovery of isotopes, it was generally assumed that all atoms of an element were identical in all respects. With their discovery it was evident that such atoms may differ in atomic weights, but it was believed that their chemical and physical properties were identical except for those properties directly related to mass such as densities of gases and condensed phases, rates of diffusion and evaporation, and others of this kind. As a result of the theoretical and experimental studies reviewed in this paper, we now know that isotopes and isotopic compounds differ in their thermodynamic properties. These differences are small except in the case of the hydrogens and they generally decrease with increasing atomic weight. These small differences... may have important applications as a means of determining the temperatures at which geological formations were laid down."*

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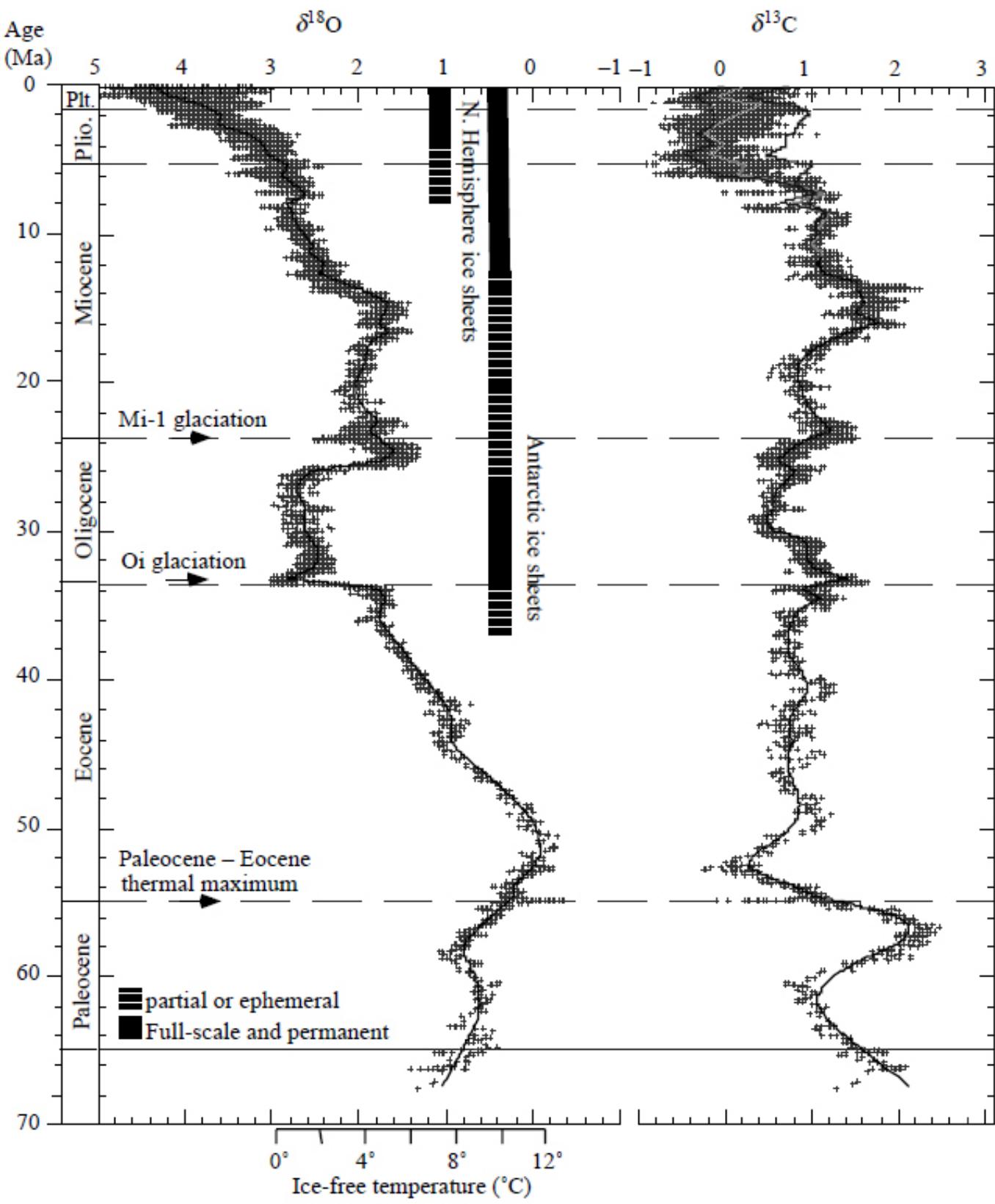
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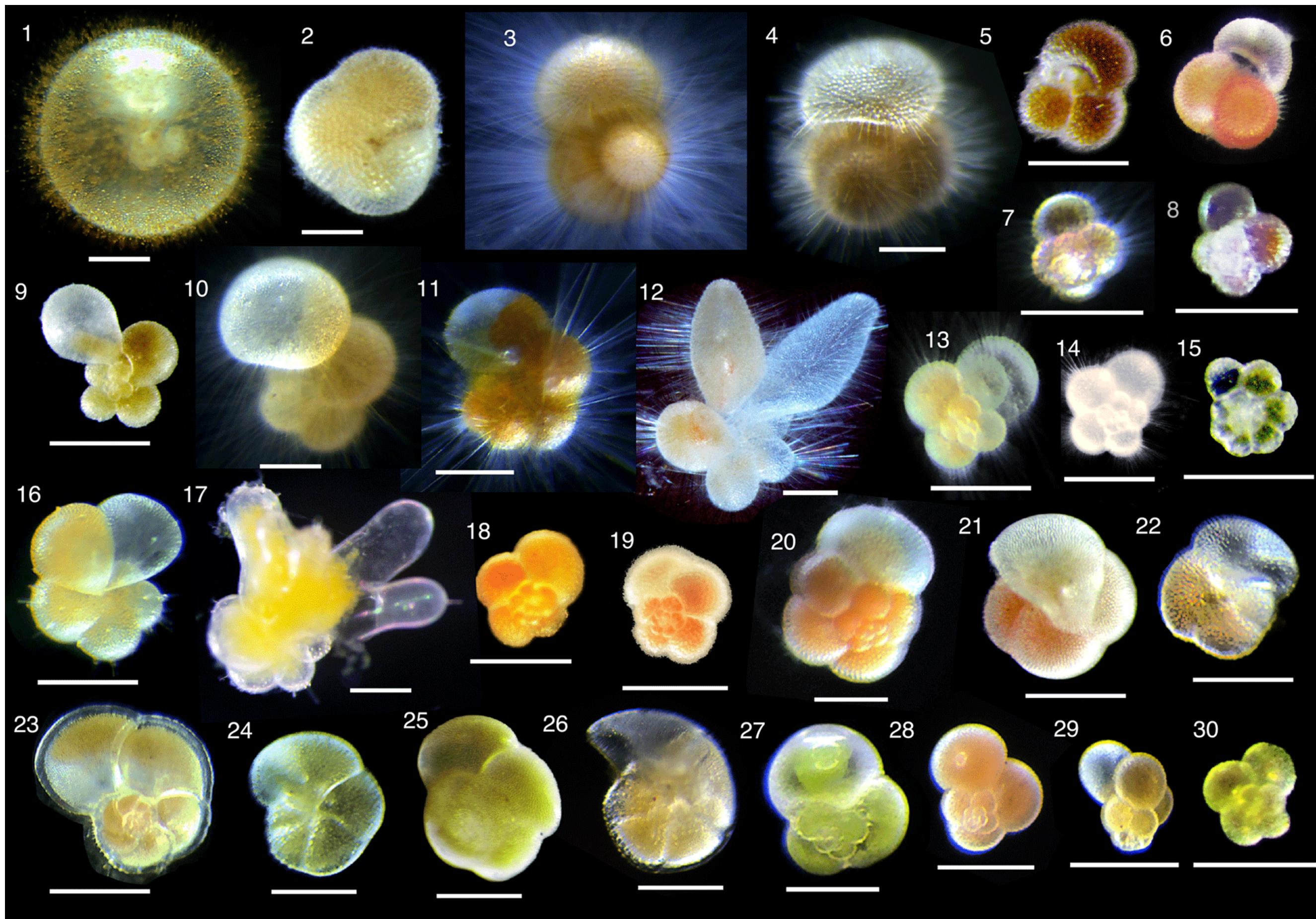
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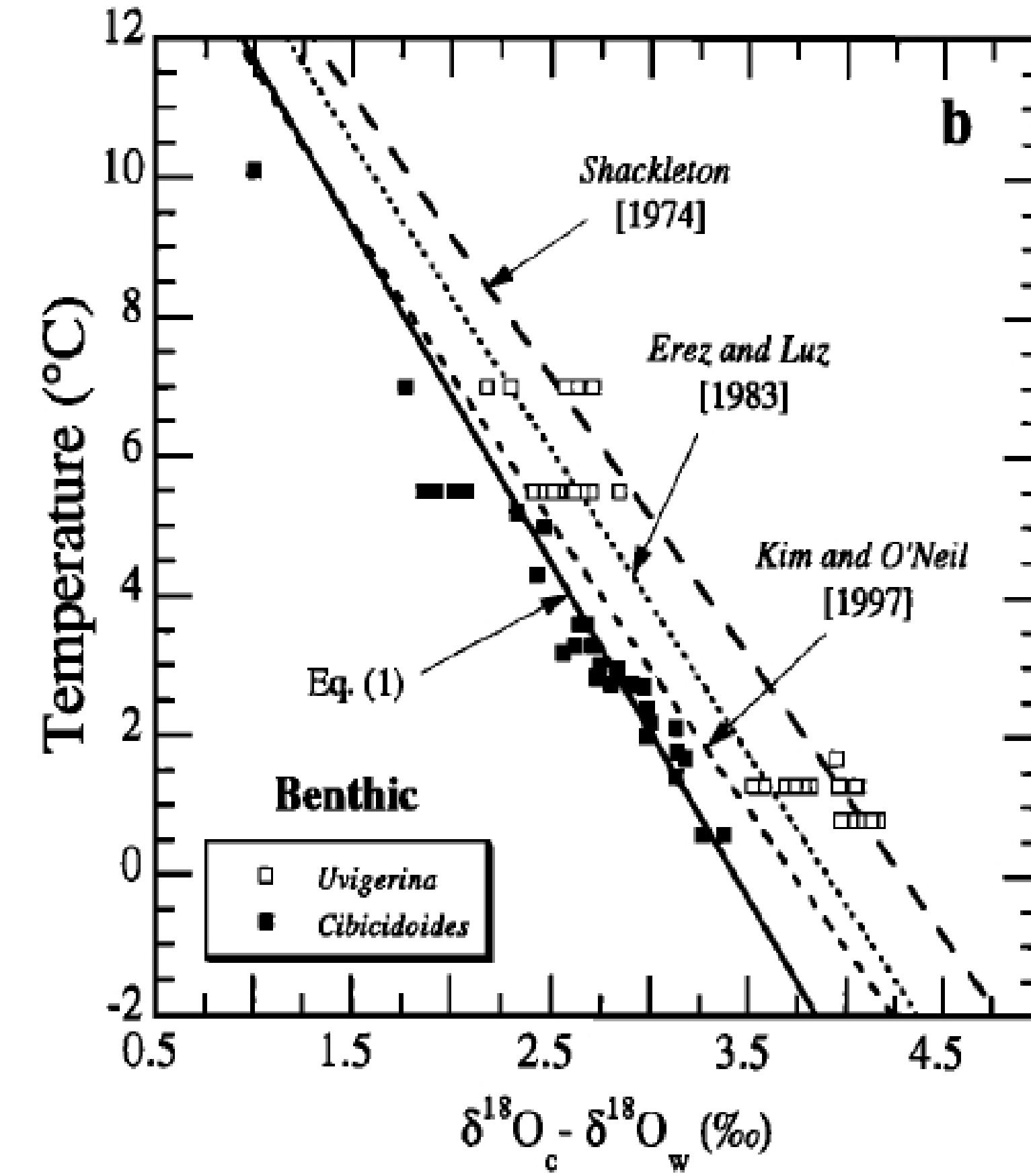
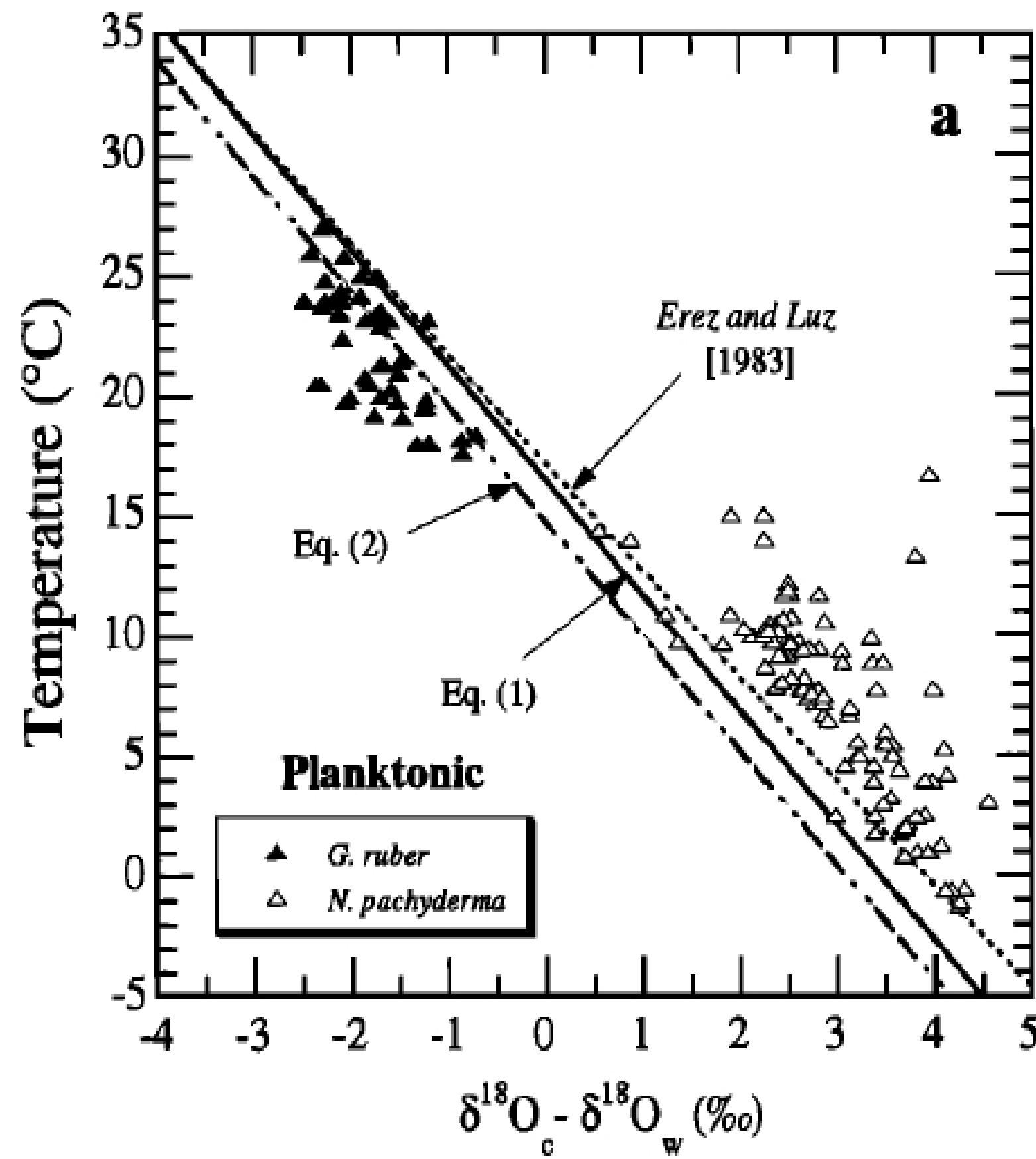




Compilation from Zachos *et al.*, 2001









$$Q = \frac{[\text{CaC}^{16}\text{O}_3][\text{H}_2^{18}\text{O}]}{[\text{CaC}^{18}\text{O}_3][\text{H}_2^{16}\text{O}]} \approx \frac{99.76}{0.24} \cdot \frac{0.24}{99.76} \approx 1$$

$^{16}\text{O} = 99.76\%$   
 $^{18}\text{O} \approx 0.24\%$

What if  $^{18}\text{O}$  more likely to be in  $\text{CaCO}_3$ ?

$$\ln 1 = 0 \quad Q \neq 1$$

at equilibrium:

$$\Delta G^\circ = -RT \ln Q \quad \text{if } Q \neq 0 \text{ then } \Delta G^\circ \neq 0$$

recall from L3 that  $\Delta G^\circ$  is the sum of the chemical potential of the products minus the sum of the chemical potential of the reactants

$$\underline{\mu_{\text{CaCO}_3}^{16}} + \underline{\mu_{\text{H}_2\text{O}}^{18}} - \underline{\mu_{\text{CaCO}_3}^{18}} - \underline{\mu_{\text{H}_2\text{O}}^{16}} = \Delta G^\circ \neq 0$$

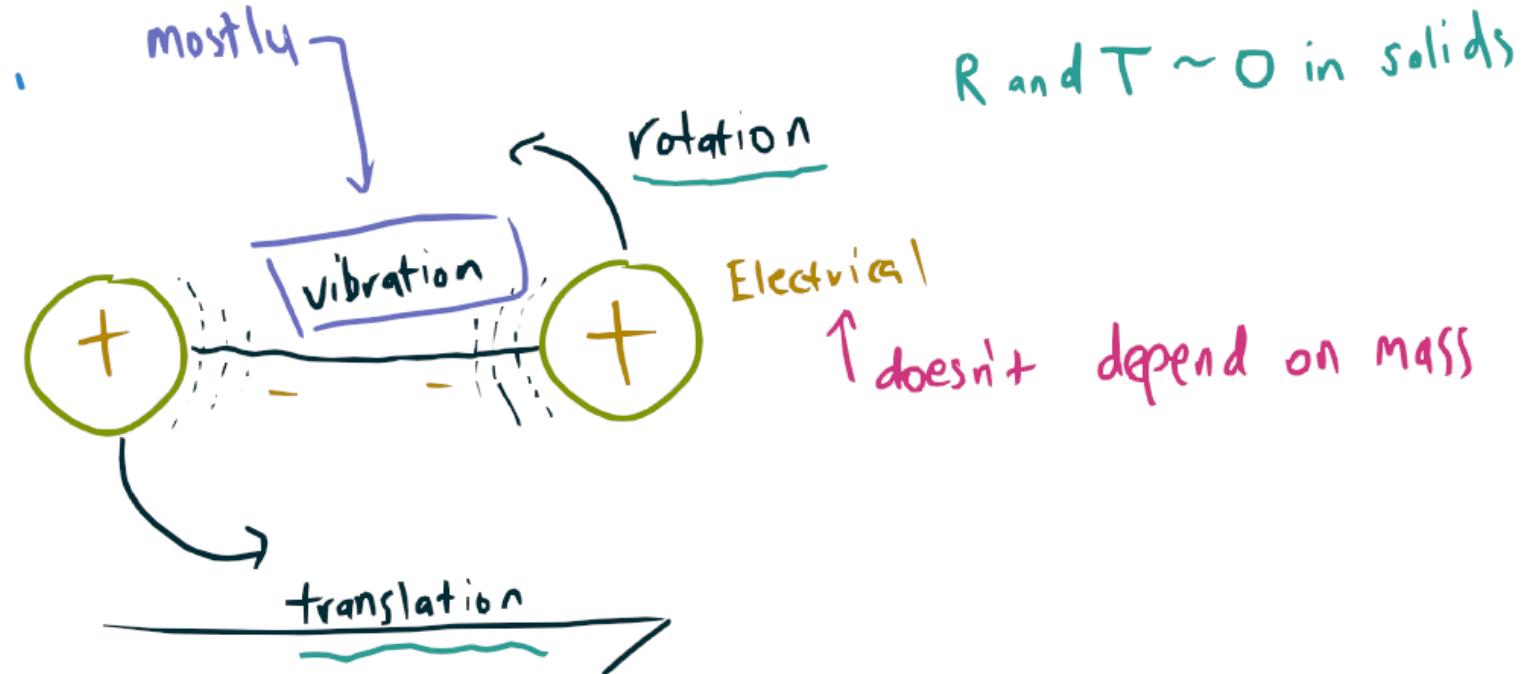
$\underline{\mu}$  must depend on mass



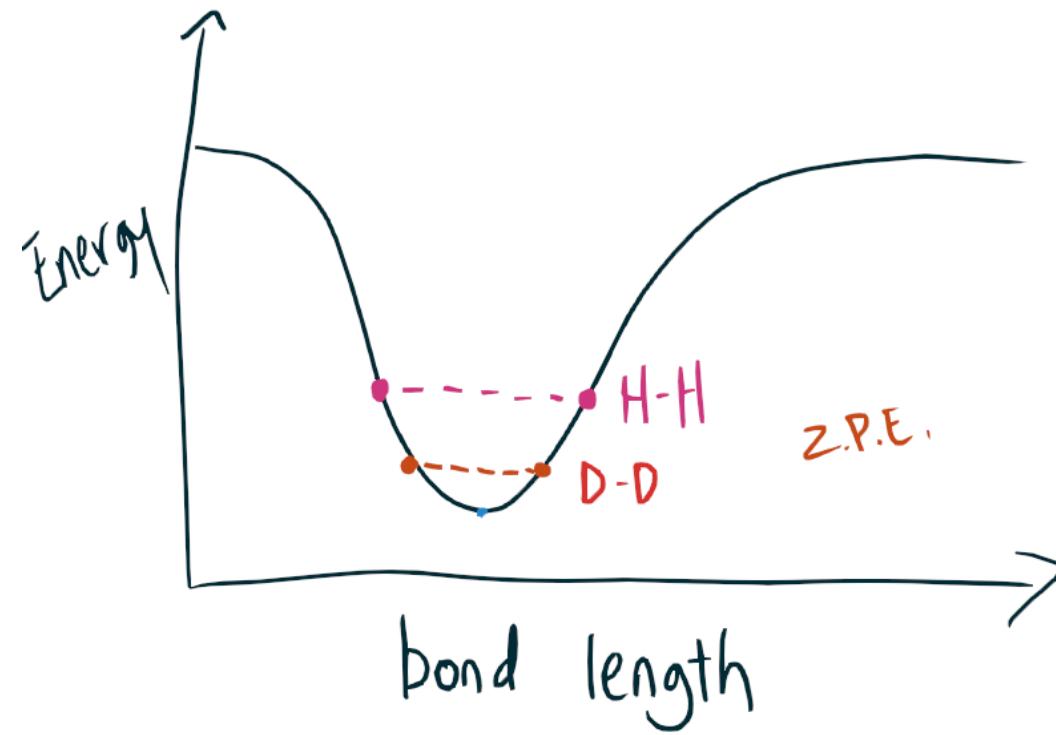
$\Delta \mu$  must depend on mass



what is this difference in energy?



Consider  $H_2$



Vibrations only exist in specific levels  
(energy is quantized)

lowest level is the zero point energy

lower  $\nu$



Spring constant

$$F = -K X$$

2 masses on a  
spring

difference from Bond length  
at the true minimum E

freq of vibration  
is directly  
proportional  
to energy

high  $\nu$  = high E

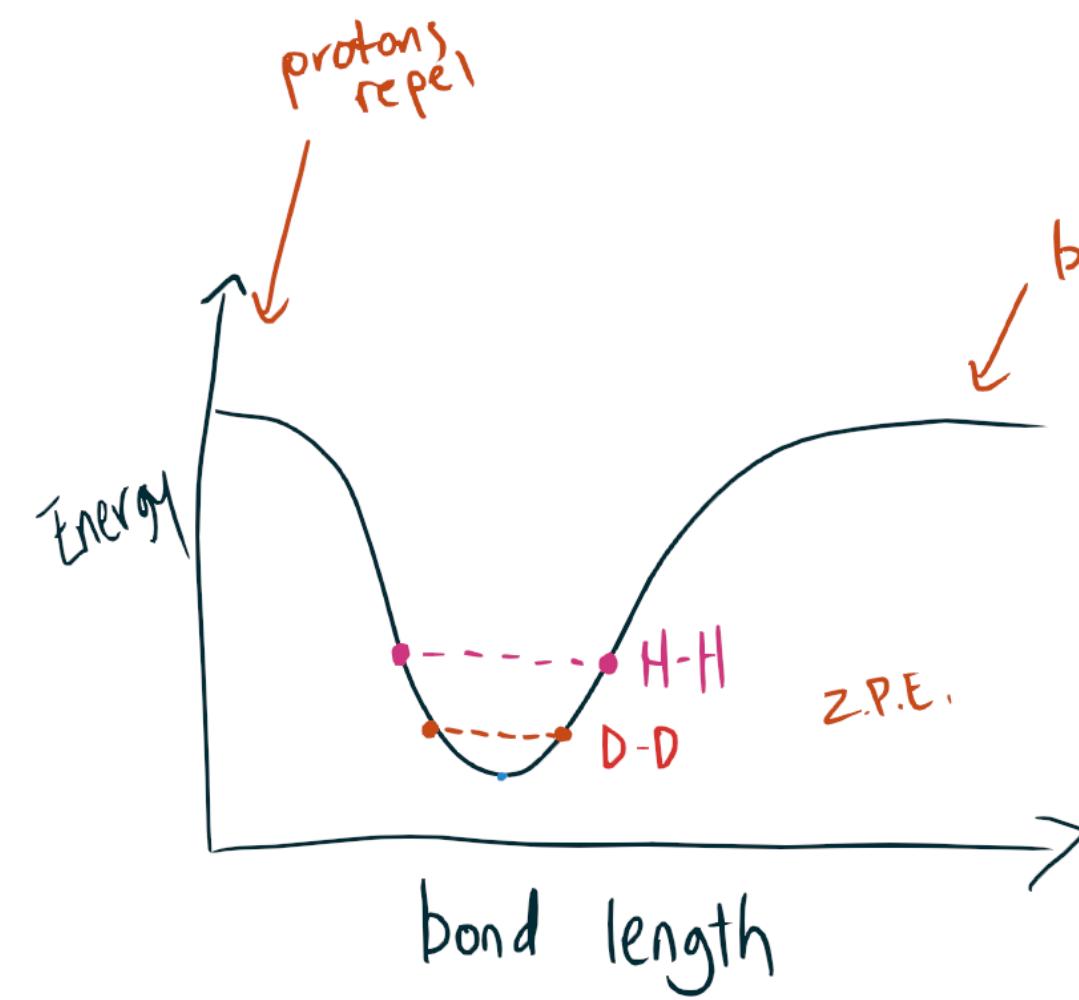
$$\nu = \frac{1}{2\pi} \sqrt{\frac{k}{\mu}}$$

reduced mass

$$\mu = \frac{M_1 \cdot M_2}{M_1 + M_2}$$

$$\rightarrow \mu_{H-D} = \frac{1 \cdot 2}{3} = 0.6 \quad \mu_{H-H} = \frac{1 \cdot 1}{2} = 0.5$$





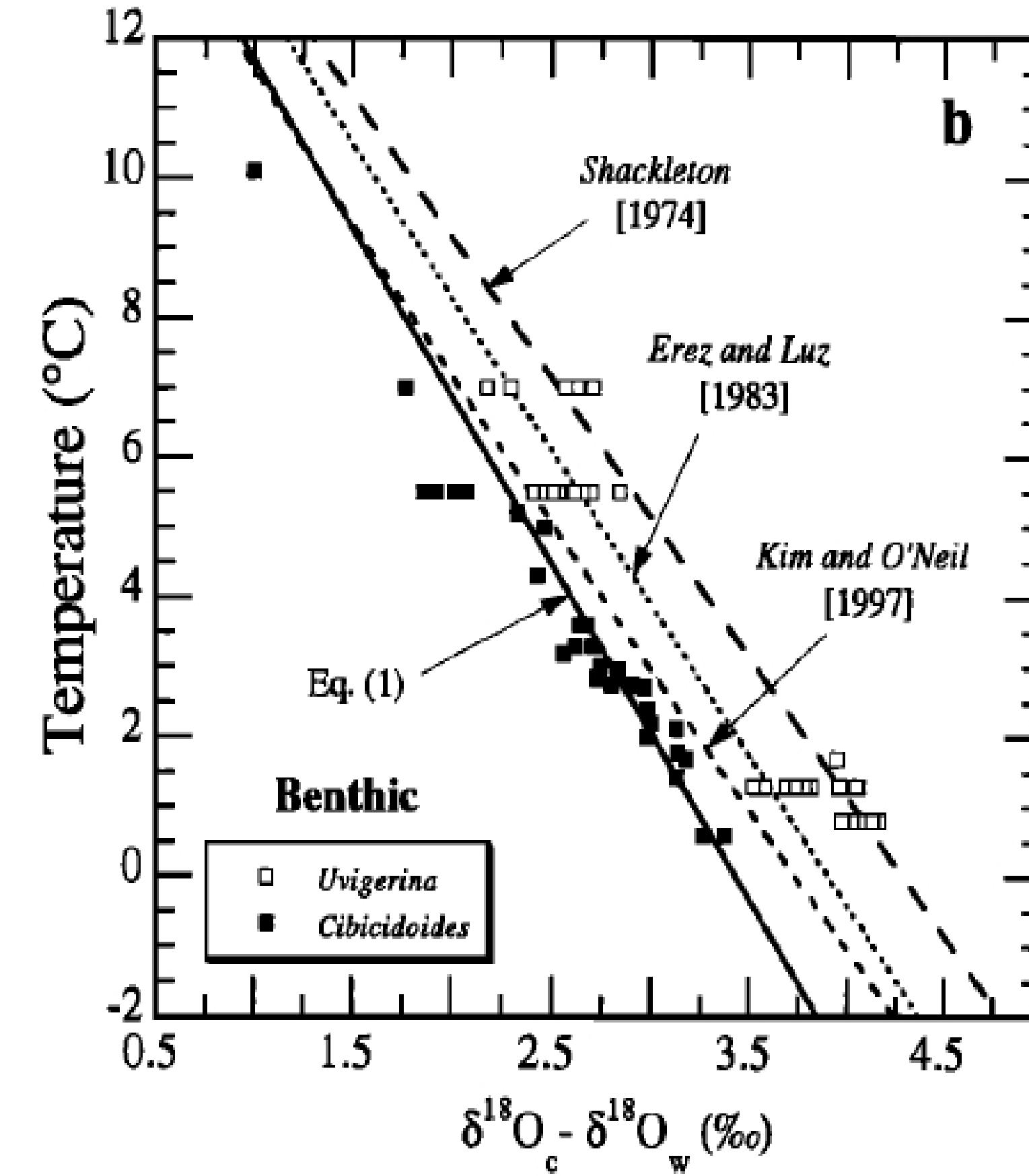
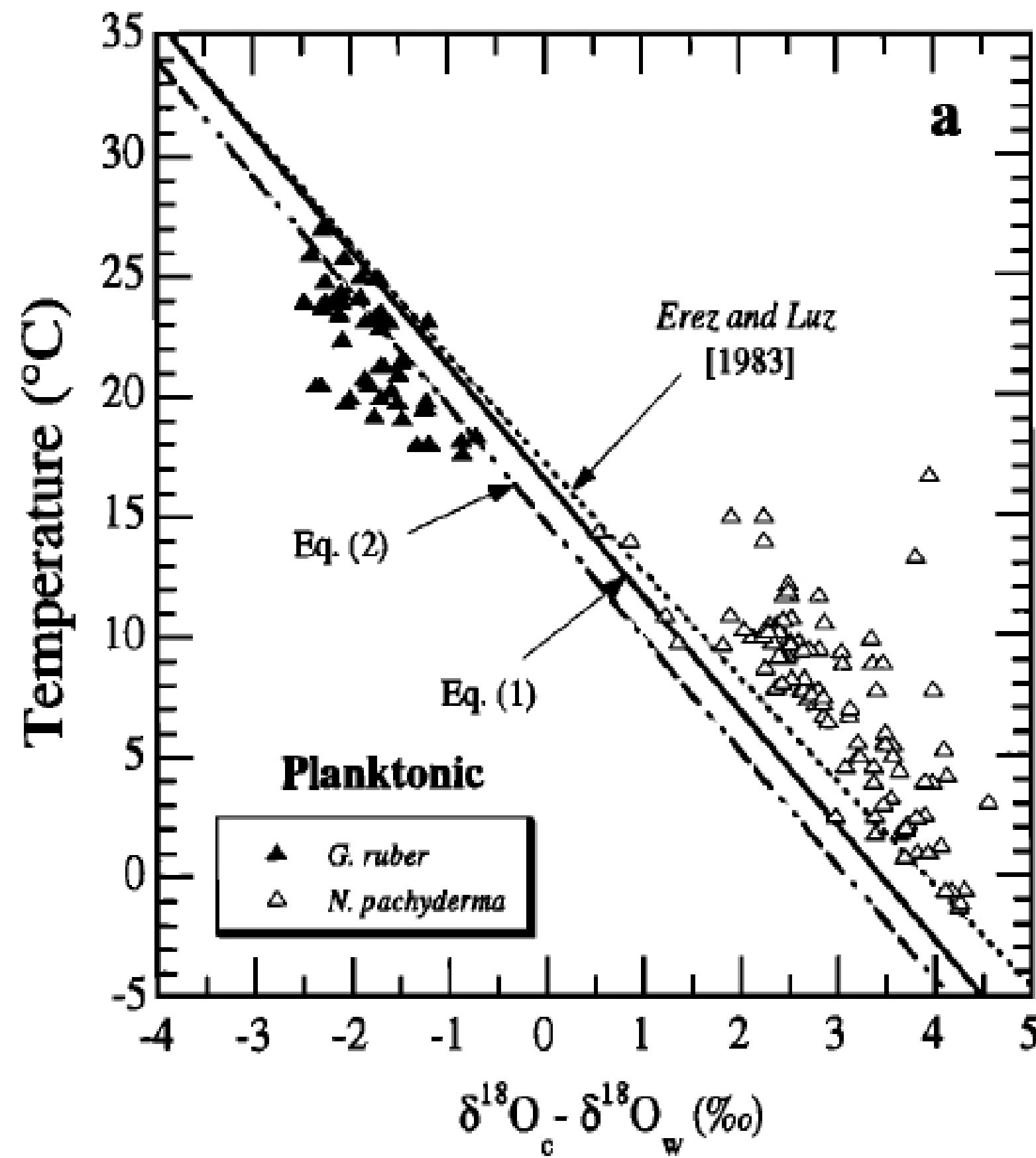
bond breaks

- Heavy isotopes have lower ZPE
- \* can we use this fact to predict which phases heavy isotopes prefer?

reactions minimize energy of the system so heavy isotopes prefer the phase with the greatest difference in ZPE<sub>LIGHT</sub> and ZPE<sub>HEAVY</sub>

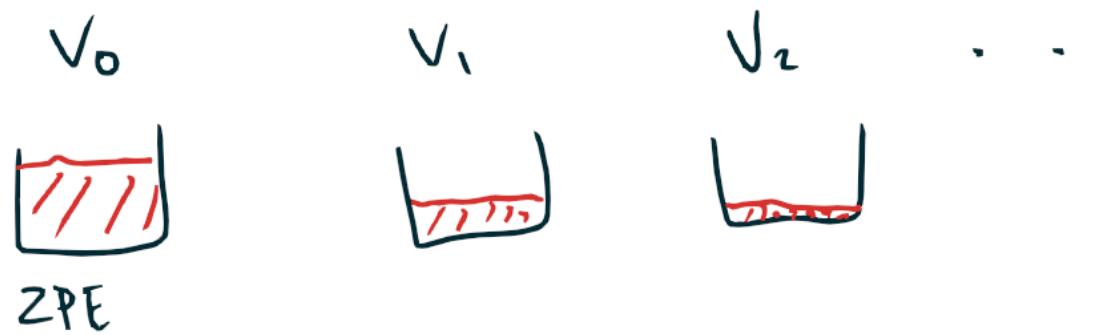
- often the strongest bonds ie mineral over H<sub>2</sub>O





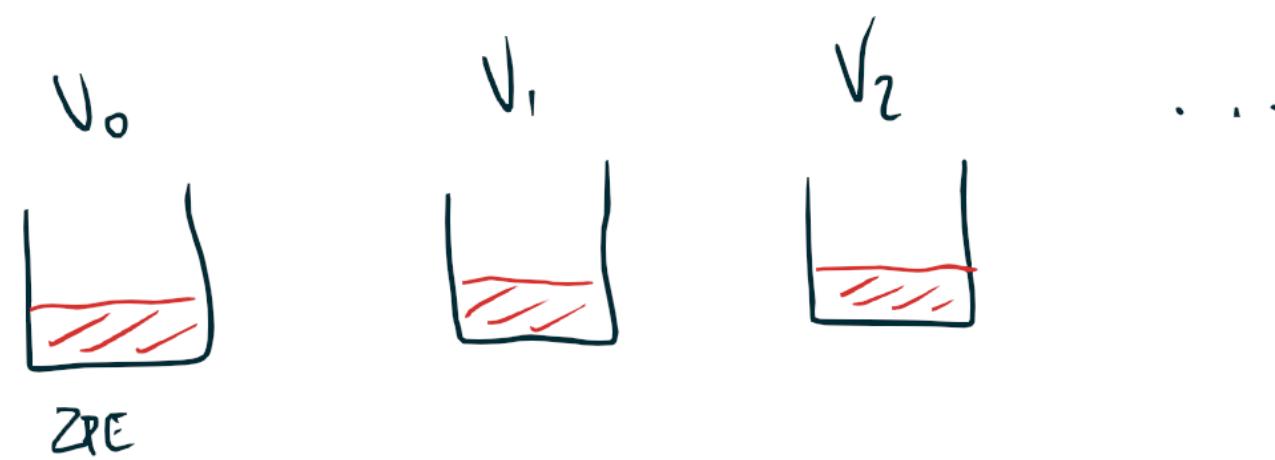
What about the sensitivity to Temp?

low T



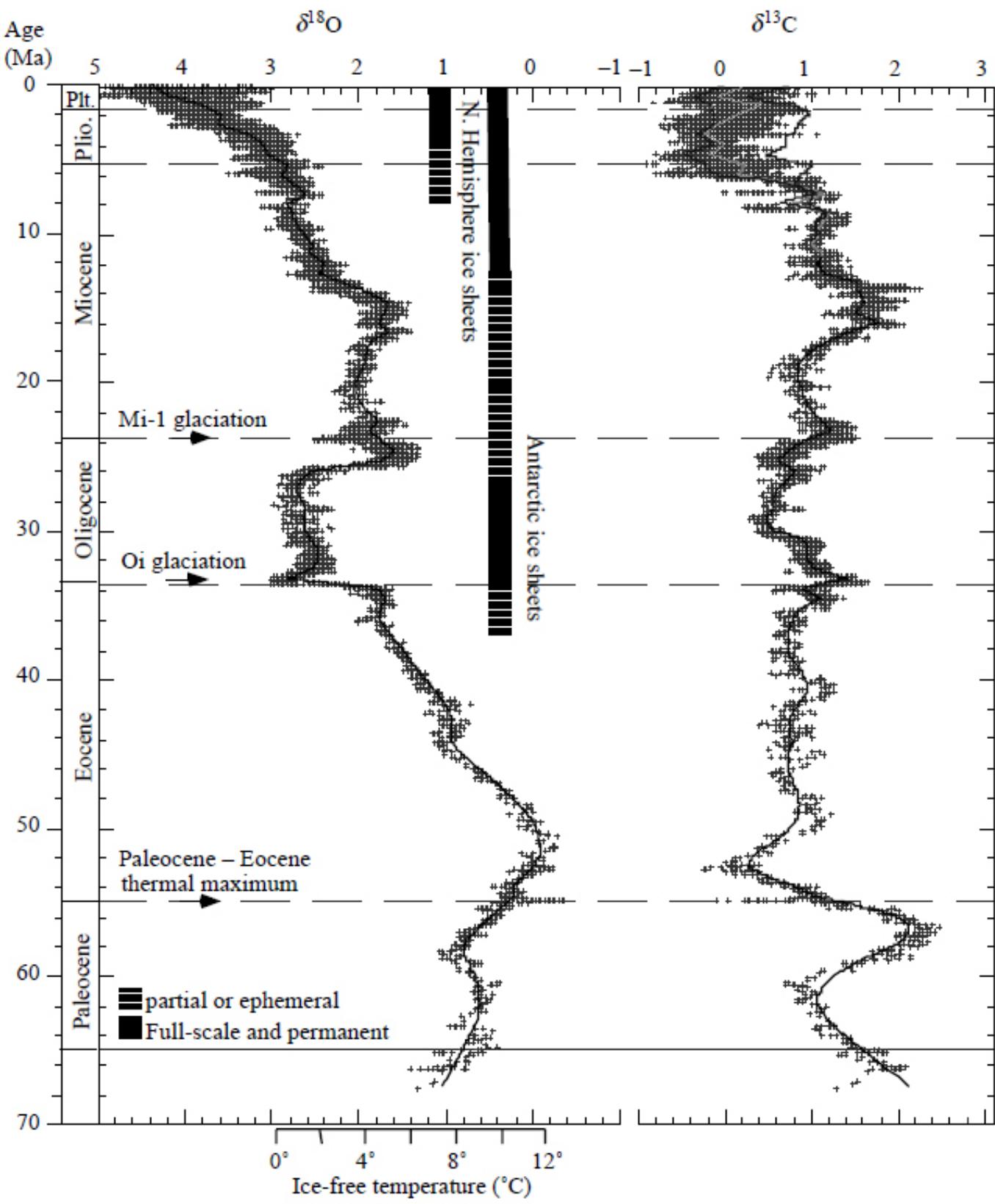
high T

even distribution  
at  $T = \infty$



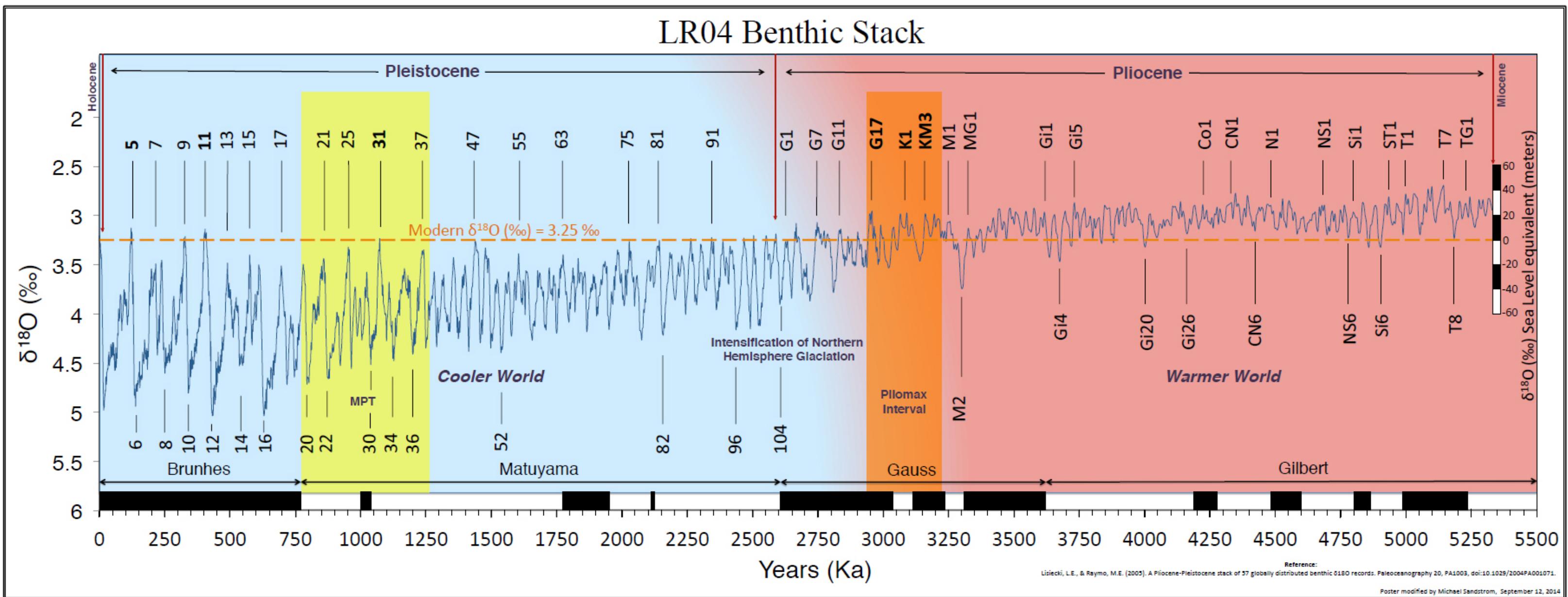
- MORE bonds vibrating at  $V_0$ ,  
MORE energy difference between  
 $H-H$  and  $H-D$





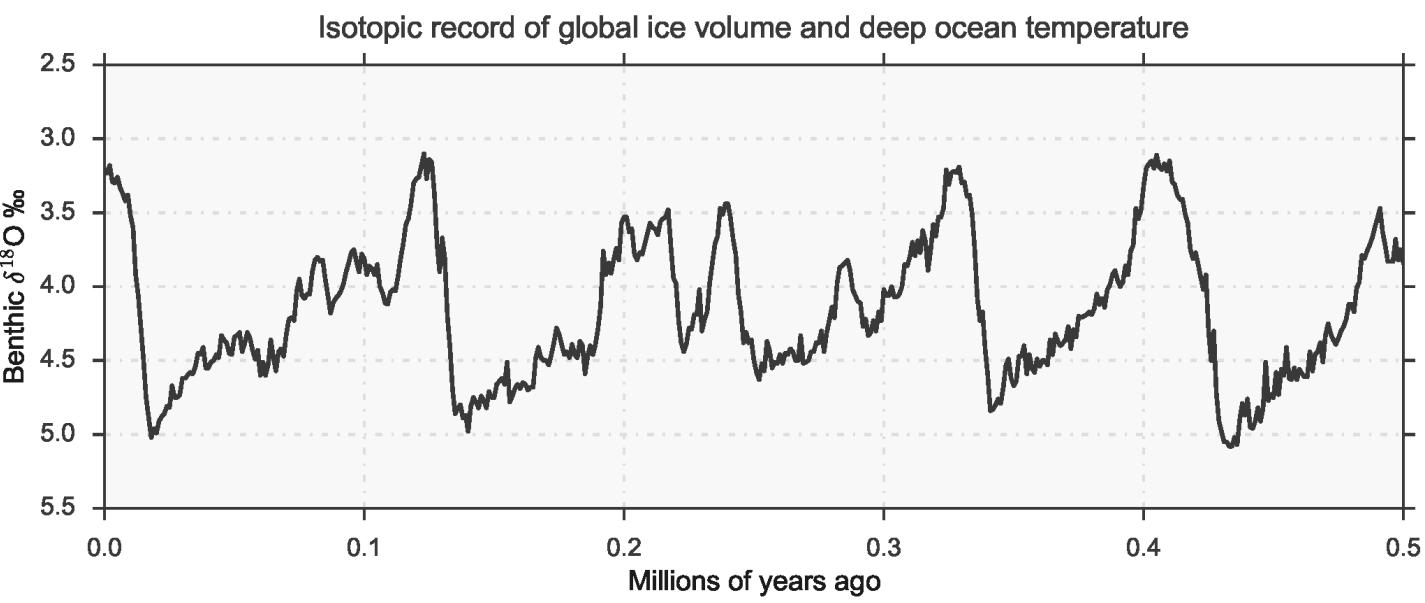
Compilation from Zachos *et al.*, 2001





from Lisicki and Raymo, 2004

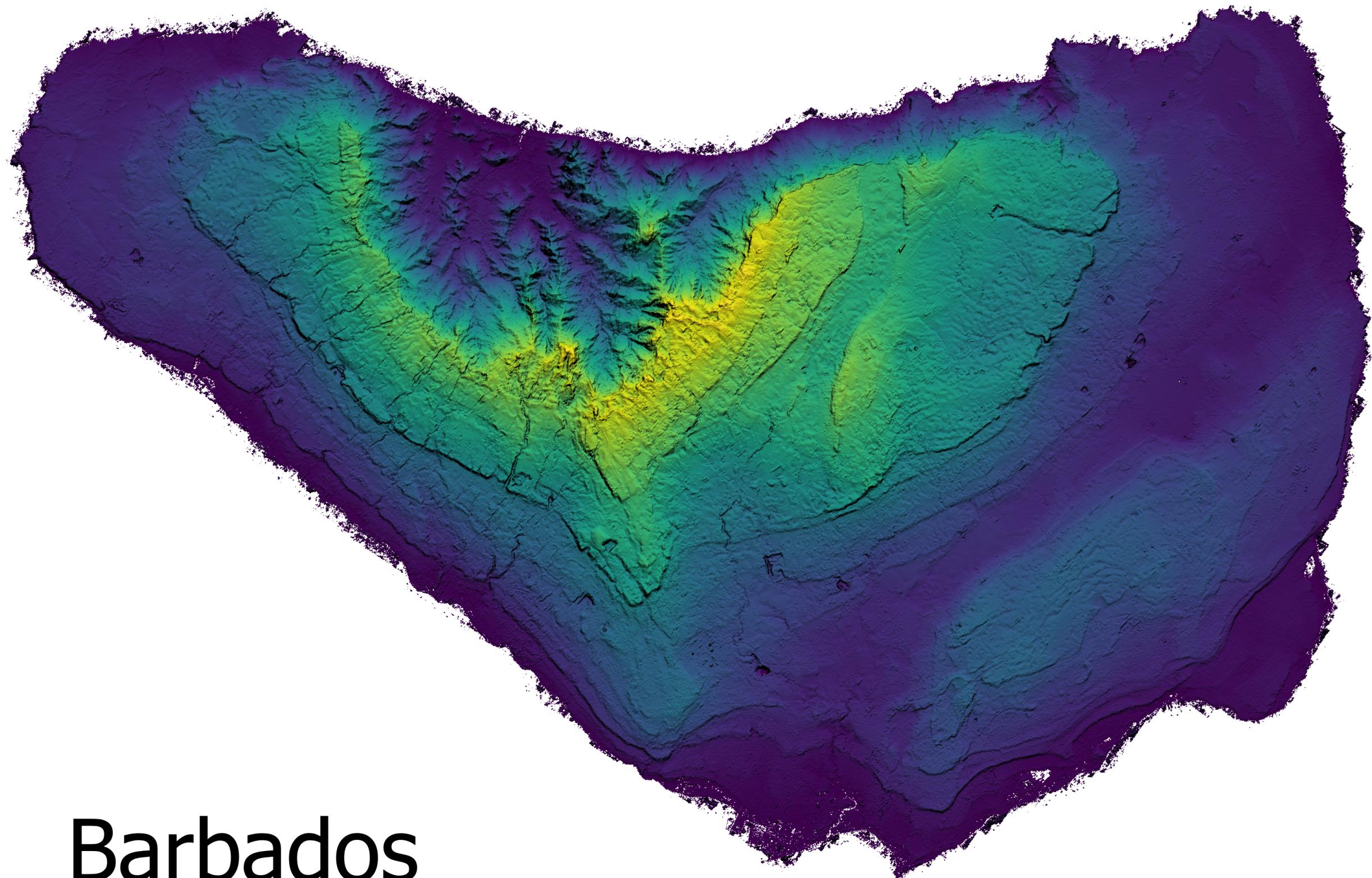






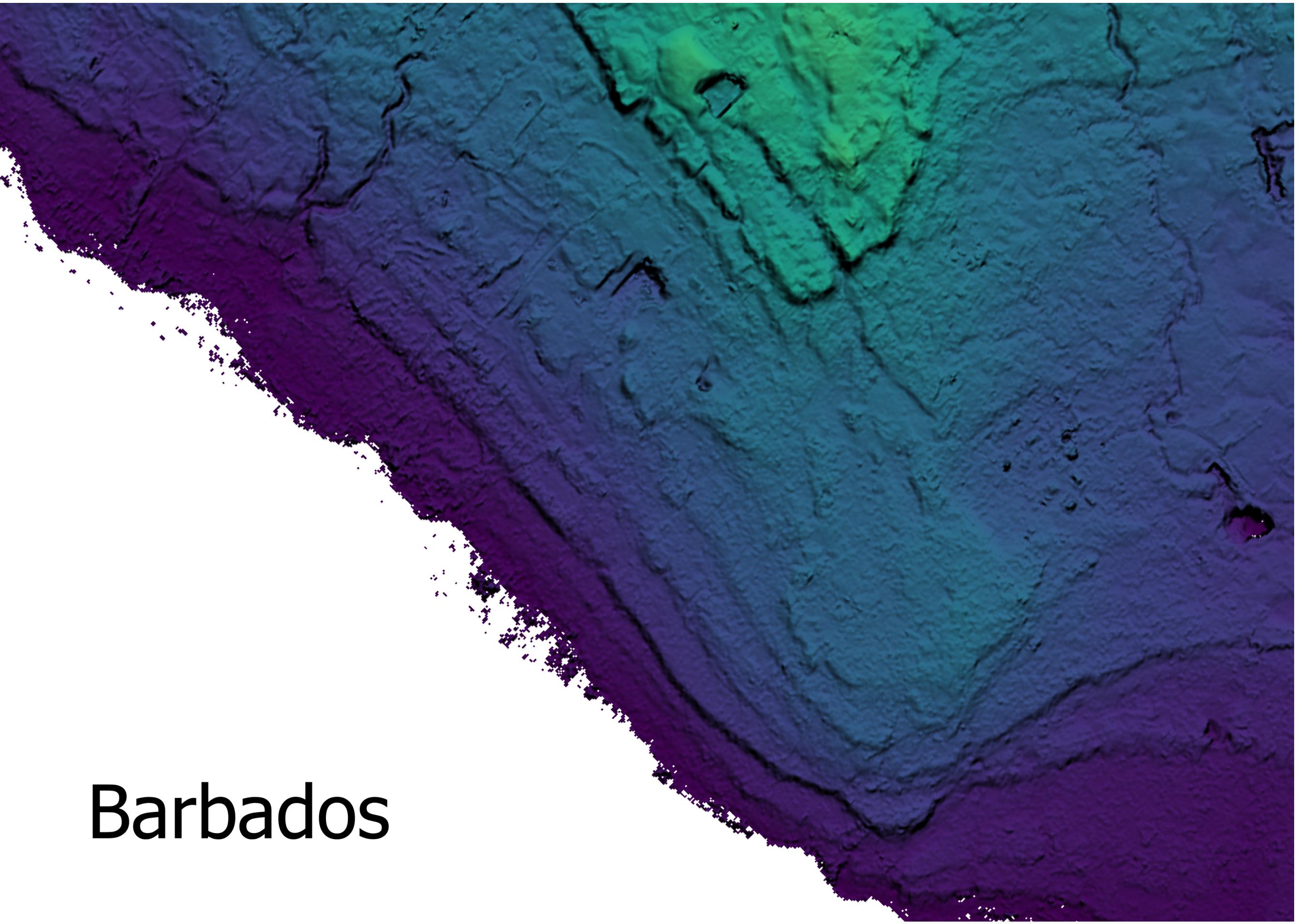
# Barbados





Barbados





Barbados







