### Lecture 3: Discussion of *Kenyon and Turcotte (1985)*

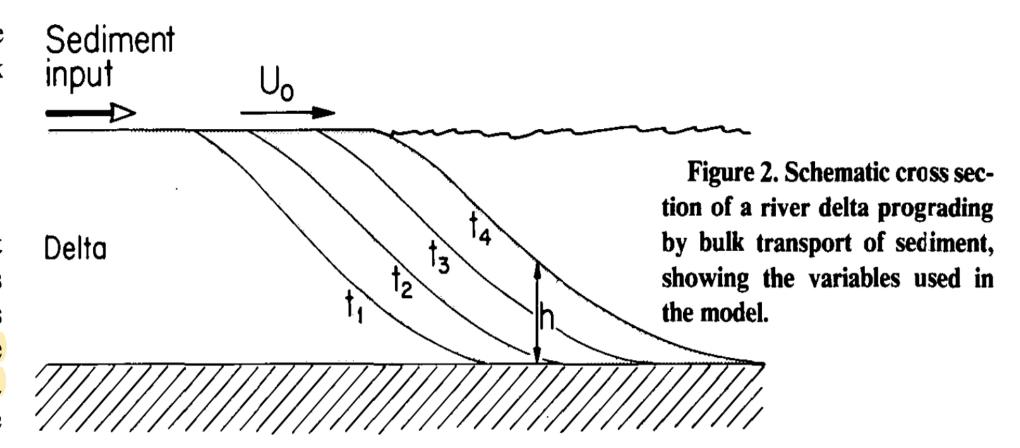
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KENYON AND TURCOTTE

deltas which are made up of the relatively fine grained sediments in which transport by bulk movements might be expected to occur.

#### PROGRADATION MODEL

Consider a river delta prograding into a quiet basin with a horizontal floor (Fig. 2). Sediments are brought to the delta by the river which is forming it, and the greatest share of them are dropped near the landward edge of the delta front (Wright and Coleman, 1974). From there



We acknowledge and respect the  $l \ni k^{\vec{w}} \ni j \ni n$  peoples on whose traditional territory the university stands and the Songhees, Esquimalt and  $V_2 \circ N_2 \circ N_3 \circ N_4 \circ N$ 





### Lecture 3: Discussion of Kenyon and Turcotte (1985)

- working pairs will be selected at random to lead discussion on different parts of the paper
- you can use the board to talk through concepts
- if a slide heading is in blue, then the picked group should take the lead (esp. questions in **bold**)
- if a slide heading is in red, then I will take the lead
- sections to be discussed:
  - Introduction and Motivation (or, why was this paper written)?
  - Setting-up the model: applying the diffusion equation (why is this an appropriate approach?)
  - Making the model: getting to an equation for delta morphology
  - Using the model: comparing to real world



What is being shown here, and how does it relate to the paper topic?

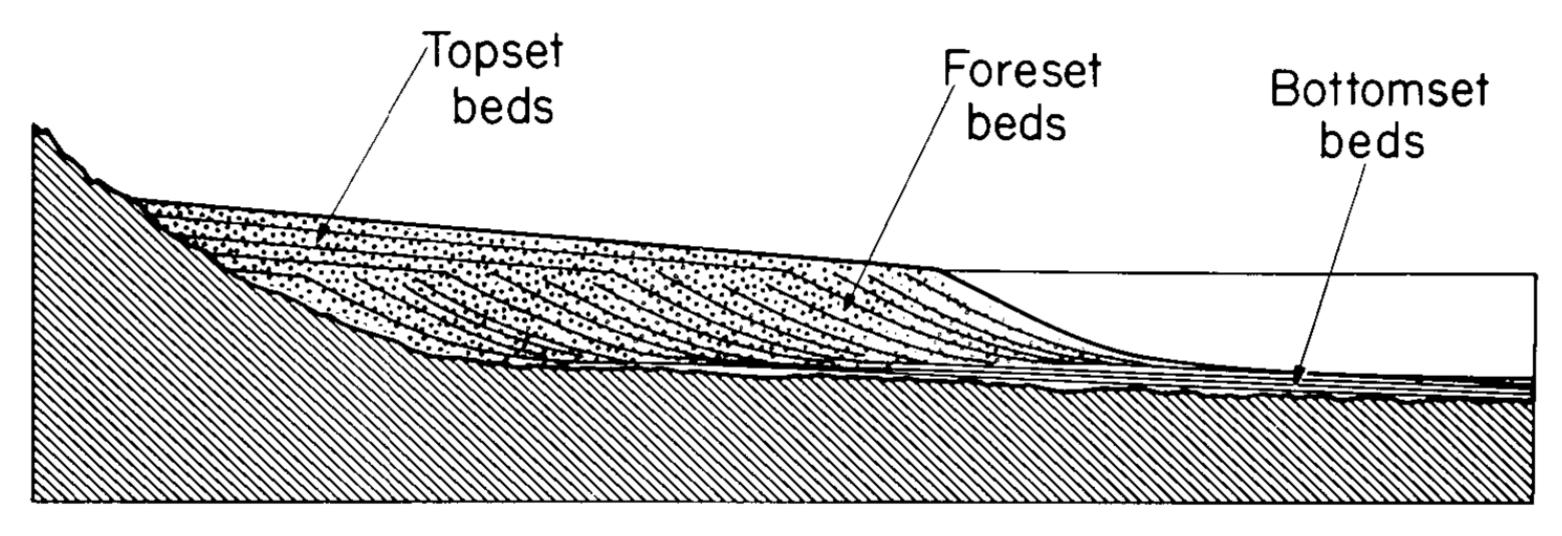
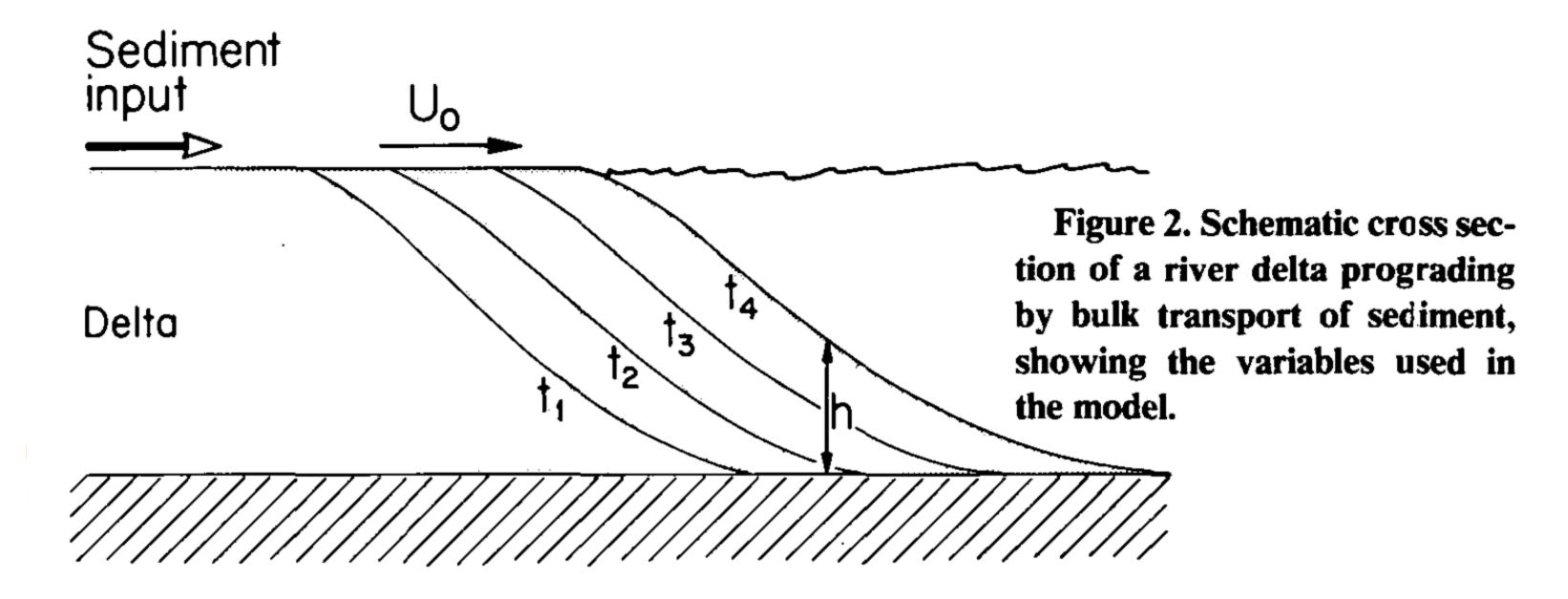


Figure 1. Cross section of a Gilbert-type delta (after Gilbert, 1890).

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This cartoon shows the model set-up. What are the assumptions, and how does it relate to a Gilbert-type delta?

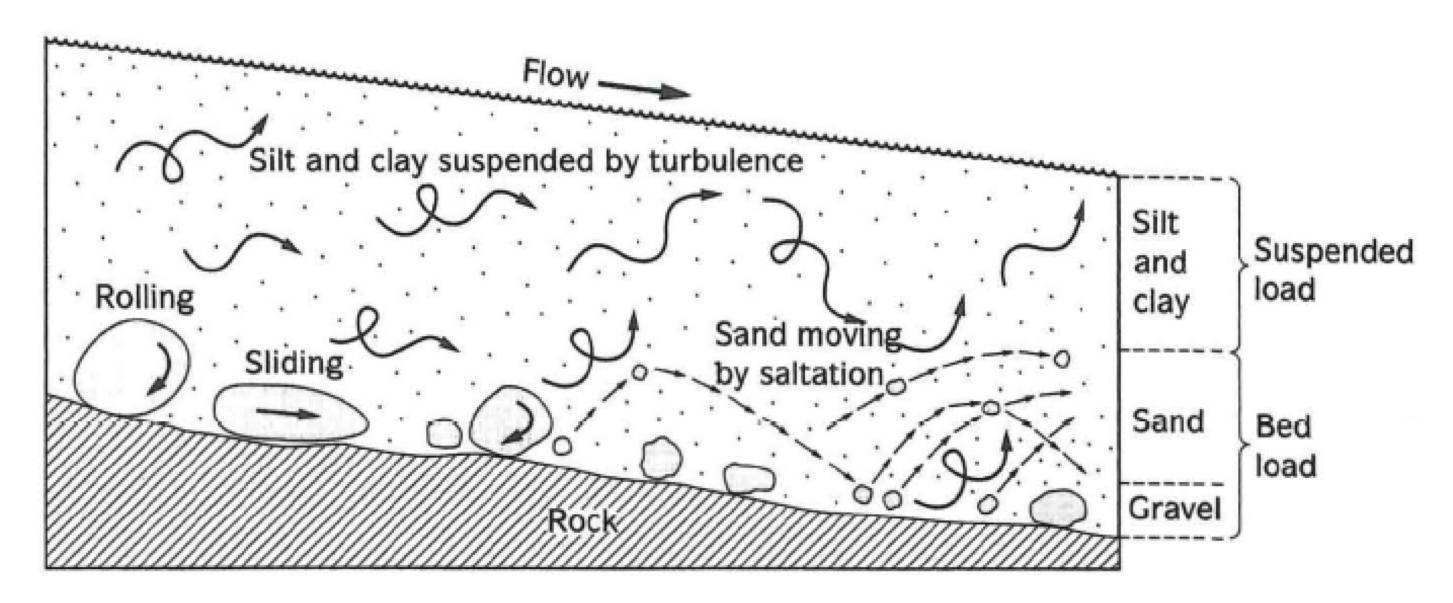


#### some key passages from the paper:

- "Recently, a number of papers have pointed out the prevalence of bulk-sediment movements, such as **creep** and **landslides**, on the subaqueous portions of large deltas (and in other subaqueous environments), as opposed to the movements of individual particles stressed by Gilbert."
- "In subaerial environments, the geomorphic forms characteristic of sediment transport by bulk motion, and in particular by creep, differ substantially from those developed by the sliding of individual particles down a slope lying at the angle of repose (Carson and Kirkby, 1972). This difference occurs because, in landforms governed by bulk sediment movement, slope is variable and adjusts to produce the rate of transport necessary to accommodate a given supply of sediment. In contrast, in landforms governed by angle of repose, slope is fixed, and the transport rate adjusts to maintain the slope. As yet, the consequences of this difference have not been considered for subaqueous features such as deltas."







- 1. traction: sediment dragged/rolled along bed surface
- 2. saltation: sediments "jump" along bed surface
- 3. suspension: sediments float along with flow



] suspended | load





Now, let's start building the model, starting with a way to model creep. The authors introduce this equation:

$$\frac{\partial h}{\partial t} = -\frac{\partial S}{\partial x}$$
 (equation 2)

In words, what is this equation saying?

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Equation 2 is the **continuity equation**, and "expresses the law of conservation of mass and, neglecting compaction, should be valid for any of the bulk-transport processes we are considering."





Next, the authors introduced this equation:

$$S = -D\frac{\partial h}{\partial x}$$
 (equation 3)

In words, what is this equation saying?

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"In theoretical geomorphology, equation 3 is derived by considering the downslope bias introduced by gravity into the random motions of soil particles induced"

subaerial: subaqeuous: freeze-thaw, bioturbation bioturbation, movement by waves





Just like we saw in lecture yesterday, if we take the derivative of equation 3 with respect to x:

$$S = -D\frac{\partial h}{\partial x}$$
 (equation 3)  

$$\frac{\partial S}{\partial x} = \frac{\partial}{\partial x} \left( -D\frac{\partial h}{\partial x} \right)$$
  

$$\frac{\partial S}{\partial x} = -D\frac{\partial^2 h}{\partial x^2}$$

We can now combine with equation 2:

$$\frac{\partial h}{\partial t} = -\frac{\partial S}{\partial x}$$
 (equation 2)  
$$\frac{\partial h}{\partial t} = D\frac{\partial^2 h}{\partial x^2}$$
 (the diffusion equation)





### Setting-up the model: landslides

Okay, so then what about landslides? What is this equation saying, and how do we relate it to diffusion?

$$\tau_d = \rho gz \sin(\beta)$$
 (equation 4)

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### Setting-up the model: landslides

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at low slope angles (like we encounter in deltas,  $\sim$ 4°), we can say  $sin(\beta) \approx tan(\beta)$ 





#### Setting-up the model: creep and landslides

now we have all the parts needed:

$$S_{L} = -C\frac{\partial h}{\partial x} \qquad \text{(equation 7)}$$

$$\frac{\partial h}{\partial t} = C\frac{\partial^{2} h}{\partial x^{2}} \qquad \text{(equation 8)}$$

$$\frac{\partial h}{\partial t} = (C + D)\frac{\partial^{2} h}{\partial x^{2}} \qquad \text{(equation 9)}$$

$$\frac{\partial h}{\partial t} = K\frac{\partial^{2} h}{\partial x^{2}} \qquad \text{(equation 10)}$$



To get to an analytical (h(x,t) = ...) solution for equation 10, we need to introduce some new variables. What are they?

$$\xi = x - u_0 t$$
 (equation 11)  
 $t' = t$  (equation 12)

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 (equation 11)  
 $t' = t$  (equation 12)

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"An analytic solution of equation 10 for steady progradation can be obtained in a coordinate system moving with the landward edge of the delta front"





How are these new variables used in these new equalities?

$$\frac{\partial h}{\partial x} = \frac{\partial h}{\partial \xi} \qquad \text{(equation 13)}$$

$$\frac{\partial h}{\partial t} = \frac{\partial h}{\partial t'} - u_0 \frac{\partial h}{\partial \xi} \qquad \text{(equation 14)}$$

$$\frac{\partial h}{\partial t'} = 0 \qquad \text{(equation 16)}$$



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$$\frac{\partial h}{\partial t'} = 0 \qquad \text{(equation 16)}$$

$$\frac{\partial h}{\partial t'} - u_0 \frac{\partial h}{\partial \xi} = K \frac{\partial^2 h}{\partial \xi^2} \qquad \text{(equation 15)}$$

$$0 = \frac{\partial^2 h}{\partial \xi^2} + \frac{u_0}{K} \frac{\partial h}{\partial \xi} \qquad \text{(equation 17)}$$



because we are ONLY differentiating with respect to  $\xi$ , we can express equation 17 as an ordinary differential equation ( $\partial \rightarrow d$ ):

$$0 = \frac{d^2h}{d\xi^2} + \frac{u_0}{K} \frac{dh}{d\xi}$$
 (equation 17)

this equation has the general form:

$$0 = y'' + Cy'$$



#### Using the model: comparing to real world

combined with the restraint that  $h=h_0$  at the top of the delta, we now have an equation that predicts how the height of the delta changes with increasing distance into the basin:

$$h(x,t) = h_0 e^{-\frac{u_0}{K}(x-u_0t)}$$

but to better facilitate data-model comparisons, it is helpful to consider the delta front slope:

$$\frac{\partial h}{\partial x} = \frac{-u_0 h}{K} \longrightarrow K = \frac{-u_0 h}{(\partial h/\partial x)}$$

Why is the above equation useful, and what can be done with it?

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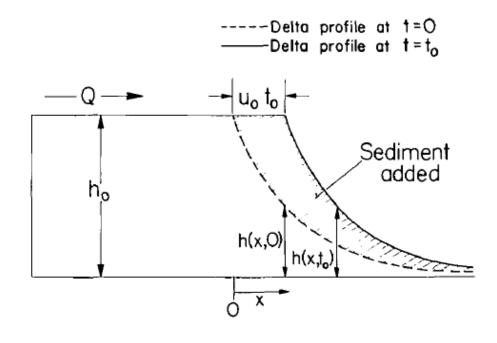


#### Using the model: comparing to real world

However, constraining  $u_0$  for real deltas is rather tricky (why?). By constrast, constraining the time-averaged volumetric sediment (Q) to a delta is more tractable. If we consider a span of time  $t_0$ , the volume of sediment added is given by the shaded region in the figure below. This volume can be calculated as:

$$Qt_0 = h_0 u_0 t_0 \longrightarrow u_0 = \frac{Q}{h_0}$$

Why are the above equalities correct? (hint: no complicated integral calculus is required)



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## Using the model: comparing to real world

So how did this model compare to real data? (I can pull up specific figures on the PDF for discussion)

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## Assignment 1.1: exploring diffusive deltas

This weeks assignment will give you some experience with coding and plotting and further explore the concepts and ideas laid out in Kenyon and Turcotte (1985):

- You are not excluded from working with other groups.
- Each person will submit their own copy of the assignment
- Assignment due date: January 27, 2023 by 1:30PM, via uploaded PDF to Brightspace

In [ ]:



