

HWRS 561b: Physical Hydrogeology II

Two-phase flow and Richards equation (Part 1)

Today:

1. Mathematical Representation of Two-Phase Flow
2. Unsaturated Flow and the Richards' Equation

Reading: Chapter 11 (Pinder & Celia, 2006).

- ❖ Physical meaning of the SWC models (Brooks-Corey vs. van Genuchten)
- ❖ Specific yield and drainable porosity in the vadose zone
- ❖ Leverett J function/scaling and Miller-Miller scaling
- ❖ Unsaturated permeability (physical meaning) and mathematical description
- ❖ Extended Darcy's Law for unsaturated flow (Buckingham, 1907)

$$\mathbf{q}_w = -\frac{k_{r,w}(S_w)\mathbf{k}}{\mu_w} \nabla(p_w + \rho_w g z)$$

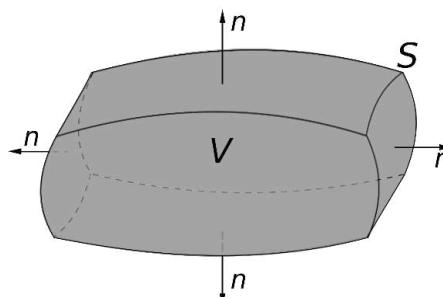
- ❖ How to derive unsaturated permeability from SWC?

Steady-state saturated flow

Derive the flow equation:

Recall: Divergence theorem

$$\int_{\partial\Omega} \mathbf{q} \cdot \mathbf{n} \, ds = \int_{\Omega} \nabla \cdot \mathbf{q} \, dV$$

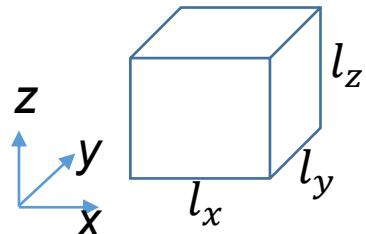


=> Divergence theorem converts a surface integral to a volume integral.

Mass conservation: Change of mass storage = mass in – mass out.

Rate of mass change:

$$\frac{d}{dt} \int_{\Omega} \rho \phi \, dV = \frac{d}{dt} \int_{\Omega} \rho \phi \, dV$$



Net fluxes:

$$\begin{aligned}
 & \int_0^{l_y} \int_0^{l_z} q_x|_{x=0} \, dy \, dz - \int_0^{l_y} \int_0^{l_z} q_x|_{x=l_x} \, dy \, dz \\
 & + \int_0^{l_x} \int_0^{l_z} q_y|_{y=0} \, dx \, dz - \int_0^{l_x} \int_0^{l_z} q_y|_{y=l_y} \, dx \, dz \\
 & + \int_0^{l_x} \int_0^{l_y} q_z|_{z=0} \, dx \, dy - \int_0^{l_x} \int_0^{l_y} q_z|_{z=l_z} \, dx \, dy
 \end{aligned}$$

Steady-state saturated flow

$$\Rightarrow - \int_{\Omega} \tilde{g} \cdot \tilde{n} dS = - \int_{\Omega} \nabla \cdot \tilde{g} dv \quad \text{divergence theorem}$$

Thus, $\frac{d}{dt} \int_{\Omega} \rho \phi dv = - \int_{\Omega} \nabla \cdot \tilde{g} dv \quad (\Omega \neq \Omega(t))$

$$\Rightarrow \int_{\Omega} \frac{d}{dt} (\rho \phi) dv = - \int_{\Omega} \nabla \cdot \tilde{g} dv$$

$$\Rightarrow \int_{\Omega} \left[\frac{d}{dt} (\rho \phi) + \nabla \cdot \tilde{g} \right] dv = 0$$

$$\Rightarrow \int_{\Omega} \left[\frac{\partial}{\partial t} (\rho \phi) + \nabla \cdot \tilde{g} \right] dv = 0$$

$$\Rightarrow \frac{\partial}{\partial t} (\rho \phi) + \nabla \cdot \tilde{g} = 0$$

Integral = 0, Ω is arbitrary ⇒ Integrand = 0

$$\text{Barai's law: } \tilde{g} = - \rho K \nabla H \quad \left. \Rightarrow \frac{\partial}{\partial t} (\rho \phi) + \nabla \cdot (-\rho K \nabla H) = 0 \right\}$$

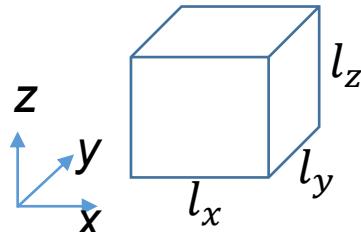
General 3D flow equation for a water-saturated porous medium.

Two-Phase Flow

Flux Law for two-phase flow (Muskat and Meres, 1936)

$$\left. \begin{aligned} q_w &= -\frac{k_{r,w}(S_w)\mathbf{k}}{\mu_w} \nabla(p_w + \rho_w g z) \\ q_{nw} &= -\frac{k_{r,nw}(S_w)\mathbf{k}}{\mu_{nw}} \nabla(p_{nw} + \rho_{nw} g z) \end{aligned} \right\} \quad \begin{aligned} q_\alpha &= -\frac{k_{r,\alpha}(S_w)\mathbf{k}}{\mu_\alpha} \nabla(p_\alpha + \rho_\alpha g z), \quad \alpha = nw \text{ or } w \\ &= -\frac{k_{r,\alpha}(S_w)\mathbf{k}}{\mu_\alpha} (\nabla p_\alpha - \rho_\alpha \mathbf{g}) \quad \nabla(\rho_\alpha g z) = -\rho_\alpha \mathbf{g} \end{aligned}$$

The derivation of the governing equations for two-phase flow shares the same procedure as that used for single-phase flow. The only difference is that a governing equation is needed for each fluid phase.



Rate of mass change:

$$\frac{\partial}{\partial t} \int_0^{l_x} \int_0^{l_y} \int_0^{l_z} \rho_\alpha(x, y, z, t) \phi(x, y, z, t) \Sigma_\alpha(x, y, z, t) dx dy dz = \frac{\partial}{\partial t} \int_\Omega \rho_\alpha \phi \Sigma_\alpha dV$$

net fluxes:

$$\int_0^{l_y} \int_0^{l_z} \rho_\alpha \bar{q}_x \Big|_{x=0} l_y l_z - \int_0^{l_y} \int_0^{l_z} \rho_\alpha \bar{q}_x \Big|_{x=l_x} l_y l_z$$

$$\int_0^{l_x} \int_0^{l_z} \rho_\alpha \bar{q}_y \Big|_{y=0} l_x l_z - \int_0^{l_x} \int_0^{l_z} \rho_\alpha \bar{q}_y \Big|_{y=l_y} l_x l_z$$

$$\int_0^{l_x} \int_0^{l_y} \rho_\alpha \bar{q}_z \Big|_{z=0} l_x l_y - \int_0^{l_x} \int_0^{l_y} \rho_\alpha \bar{q}_z \Big|_{z=l_z} l_x l_y$$

$$\begin{aligned} &\Rightarrow - \int_\Omega \rho_\alpha \bar{q}_\alpha \cdot \bar{n} dS \\ &= - \int_\Omega \nabla \cdot (\rho_\alpha \bar{q}_\alpha) dV \quad \text{Divergence theorem} \end{aligned}$$

Two-Phase Flow

$$\Rightarrow \frac{\partial}{\partial t} \int_{\Omega} \rho_2 \phi S_2 \, dv = - \int_{\Omega} \nabla \cdot (\rho_2 \underline{g}_2) \, dv$$

$\downarrow \Omega \neq \Omega (*)$

$$\Rightarrow \int_{\Omega} \frac{\partial}{\partial t} (\rho_2 \phi S_2) \, dv = - \int_{\Omega} \nabla \cdot (\rho_2 \underline{g}_2) \, dv$$

$$\Rightarrow \int_{\Omega} \left[\frac{\partial}{\partial t} (\rho_2 \phi S_2) + \nabla \cdot (\rho_2 \underline{g}_2) \right] \, dv = 0$$

$\downarrow \Omega \text{ is arbitrary}$

$$\Rightarrow \frac{\partial}{\partial t} (\rho_2 \phi S_2) + \nabla \cdot (\rho_2 \underline{g}_2) = 0 \quad \left. \begin{array}{l} \\ \end{array} \right\} \Rightarrow$$

Flux law: $\underline{g}_2 = - \frac{k_{r,2} \underline{k}}{\mu_2} \nabla (P_2 + \rho_2 g z)$

$$\boxed{\frac{\partial}{\partial t} (\rho_2 \phi S_2) - \nabla \cdot \left[\rho_2 \frac{k_{r,2} \underline{k}}{\mu_2} \nabla (P_2 + \rho_2 g z) \right] = 0}$$

General governing equation for two-phase flow

Two-Phase Flow

The governing equations for the two fluid phases:

Wetting phase $\frac{\partial}{\partial t}(\rho_w \phi S_w) - \nabla \cdot \left[\rho_w \frac{k_{r,w} \mathbf{k}}{\mu_w} \nabla(p_w + \rho_w g z) \right] = 0 \quad (1)$

Nonwetting phase $\frac{\partial}{\partial t}(\rho_{nw} \phi S_{nw}) - \nabla \cdot \left[\rho_{nw} \frac{k_{r,nw} \mathbf{k}}{\mu_{nw}} \nabla(p_{nw} + \rho_{nw} g z) \right] = 0 \quad (2)$

Unknowns: $p_w, p_{nw}, S_w, S_{nw}, k_{r,w}, k_{r,nw}$ (Assuming $\rho_\alpha, \mu_\alpha, \phi, \mathbf{k}$ are known.)

=> 6 unknowns, 2 equations. We need 4 more equations to have a mathematically closed system.

Constitutive equations (e.g., VG or B-C models) $S_w + S_{nw} = 1 \quad (3)$

$$\begin{cases} p_{nw} = p_w + p_c(S_w) & (4) \\ k_{r,w} = k_{r,w}(S_w) & (5) \\ k_{r,nw} = k_{r,nw}(S_w) & (6) \end{cases}$$

- 6 equations, 6 unknowns. The system of equations is mathematically closed.
- Substituting Eqs. (3-6) to Eqs. (1-2) => 2 equations and 2 unknowns.
- Several options for the primary variables.
Two common options:
 - ✓ One phase pressure + one phase saturation (e.g., p_w, S_{nw})
 - ✓ Two phase pressures (e.g., p_w, p_{nw})

Unsaturated Flow: Richards' Equation

Let's now think about the air-water system in the vadose zone.

At a common condition: 1 atm, 20 °C

$$\left. \begin{array}{l} \rho_w \approx 1000 \text{ kg/m}^3 \\ \rho_a \approx 1 \text{ kg/m}^3 \end{array} \right\} \rho_a \ll \rho_w \quad \left. \begin{array}{l} \mu_w \approx 10^{-3} \text{ Pa} \cdot \text{s} \\ \mu_a \approx 1.8 \times 10^{-5} \text{ Pa} \cdot \text{s} \end{array} \right\} \mu_a \ll \mu_w$$

Air flux: $q_a = -\frac{k_{r,a}\mathbf{k}}{\mu_a} \nabla(p_a + \rho_a g z)$

Richards' assumptions $\rightarrow \mu_a \ll \mu_w$. Air is essentially inviscid. In that sense, almost no gradient of energy potential is required to drive air flow.

$$\Rightarrow \nabla(p_a + \rho_a g z) \approx 0 \Rightarrow p_a + \rho_a g z \approx \text{const}$$

At the ground surface ($z=0$), $p_a=0$

$$p_a(z) = 0 - \rho_a g z \quad \left. \begin{array}{l} p_a \approx 1 \text{ kg/m}^2 \\ \rho_a g z \ll \text{atm} \end{array} \right\} \Rightarrow p_a \approx 0 \text{ in the unsaturated zone.}$$

$\frac{1 \times 10 \times 100}{\text{kg/m}^3 \text{ m/s}^2 \text{ m}}$

Richards' Equation

Governing equations for air-water flow in porous media:

$$\frac{\partial}{\partial t}(\rho_a \phi s_a) - \nabla \cdot \left[\frac{\rho_a k_{r,a} \mathbf{k}}{\mu_a} \nabla(p_a + \rho_a g z) \right] = 0 \quad (1)$$

$$\frac{\partial}{\partial t}(\rho_w \phi s_w) - \nabla \cdot \left[\frac{\rho_w k_{r,w} \mathbf{k}}{\mu_w} \nabla(p_w + \rho_w g z) \right] = 0 \quad (2)$$

$$s_a + s_w = 1 \quad (3)$$

$$k_{r,a} = k_{r,a}(s_w) \quad (4)$$

$$k_{r,w} = k_{r,w}(s_w) \quad (5)$$

$$p_a - p_w = p_c(s_w) \quad (6)$$

p_a = 0. Thus, no need to solve equation (1)
now, this is the only eqn we need to solve. It's essentially the Richards' eqn, but in a more general form.

If we simplify by assuming constant density and porosity:

$$\phi \frac{\partial s_w}{\partial t} - \nabla \cdot \left[\frac{k_{rw} k}{\mu_w} \nabla(p_w + \rho_w g z) \right] = 0$$

Richards' Eqn.