

Needs the climada module `etopo`¹

This module implements a raw version of a tropical cyclone storm surge model². It's based on climada's core tropical cyclone (TC) module. Hence in order to run `tc_surge`, make sure you've made yourself familiar (to some extent at least) with the core climada tropical cyclone hazard module. Please recall that climada stores the hazard information for each event_i at each centroid_j as `hazard.arr(event_i,centroid_j)`. Hence any climada hazard event set (stored in a structure `hazard`) contains all hazard event footprints for a given region (i.e. all tropical cyclones in the North Atlantic basin). Further to that, probabilistic hazard events can be generated and hence a full probabilistic hazard event set contains not only the historic events, but also a full set of probabilistic (storm) events for a given hazard in a given region. See the core climada manual for details.

In essence, `tc_surge` infers a proxy storm surge height (measured in m above terrain, not MSL³) from the maximum windspeed (peak gust in m/s) for each TC event at each centroid, i.e. it takes `hazard.arr(event_i,centroid_j)` and converts the maximum windspeed into a proxy surge height. It then uses a global bathymetry (and topography) dataset⁴ to infer land surface height at each centroid and hence calculates the remaining surge given the height of any given centroid.

The relation between maximum windspeed and surge height is implemented in the code `tc_surge_hazard_create`, where

$$\text{surge height [m]} = \max(f(\text{wind speed [m/s]}) - \text{elevation [m]}, 0)$$

where you find `f(wind speed)` documented further below and the elevation is the elevation of the centroid, inferred from the bathymetry data.

We do present a crude approach here, but given the many other uncertainties in the full economics of climate adaptation calculations, the method provides good enough a basis for the intended purpose.

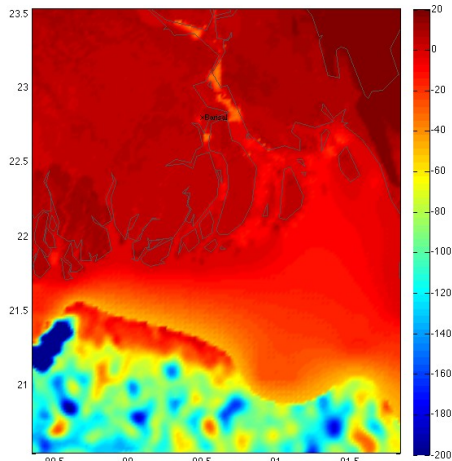
Please use `tc_surge_TEST` to test this module. The code `tc_surge_hazard_create` creates a storm surge (TS) hazard event,

¹ https://github.com/davidnbresch/climada_module_etopo

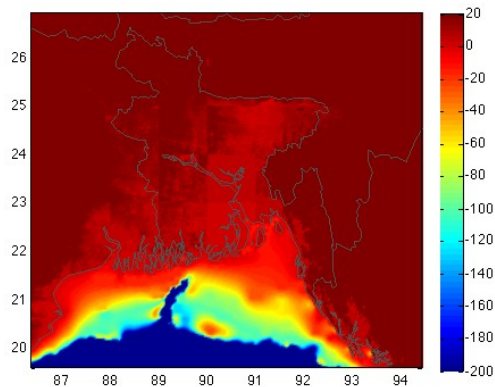
² Please note that a similar module exists for torrential rain, named `tc_rain`

³ This convention is most useful, as the damage-relevant surge height is the height of the surge at each asset's location, i.e. the centroid. This way, the elevation of the centroid does not need to be known at the time of damage calculation (in code climada), but only at the time of the generation of the storm surge hazard event set.

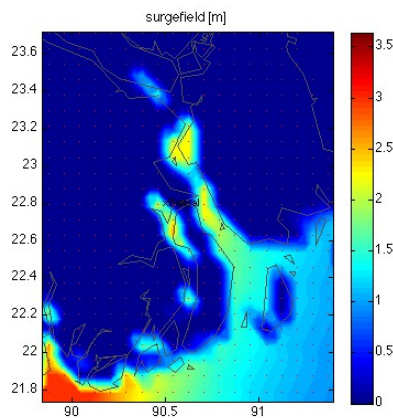
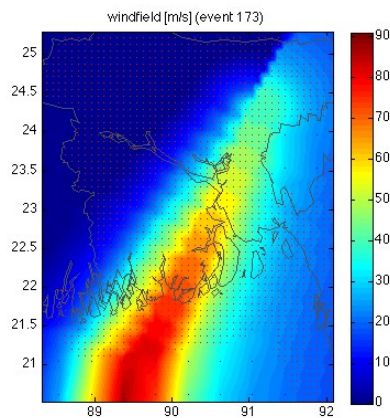
⁴ It uses ETOPO dataset, see <http://www.ngdc.noaa.gov/mgg/global/global.html> and the readme files in `.../etopo/data/` (get https://github.com/davidnbresch/climada_module_etopo)



based on an existing tropical cyclone (TC) hazard event set (see e.g. `climada_tc_hazard_set`)
 The high-resolution bathymetry dataset (ETOPO1) for a region in Bangladesh (the city of Barisal is denoted). Note the very high resolution of the elevation information.

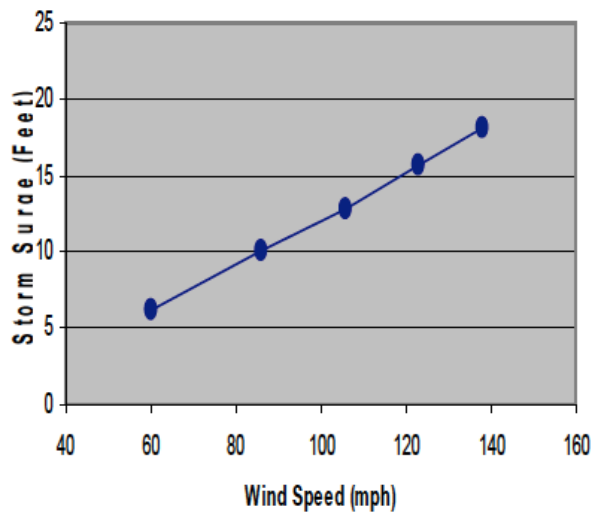


For comparison, the mid-resolution bathymetry dataset (ETOPO2) for the same region (a bit larger area shown). Note the kind of 'tiles' in the land area, leading to artificial 'steps' in elevation of a few decimeters... Hence the use of ETOPO1 is highly recommended.

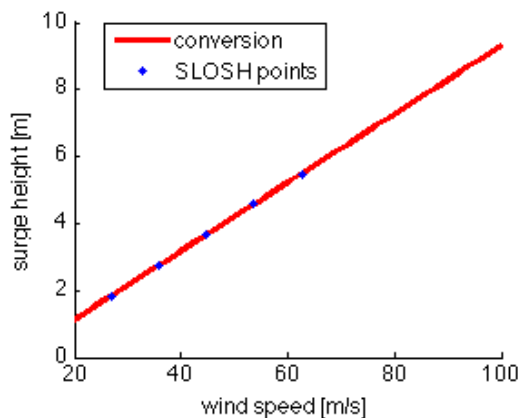


Comparison of maximum gust windfield (left) and resulting surge height field (right, zoomed in, city of Barisal marked) for the strongest recent event in Bangladesh (Nov 2007).

Wind speed surge height relationship



The simple linear relationship between wind speed and surge height⁵. Simple conversion in SI units⁶ reveals a function like $\text{surge height [m]} = a * (v - v_0) + b$ with $a=0.1023 \text{ [m/ms}^{-1}\text{]}$, $b=1.8288 \text{ [m]}$, $v_0=26.8224 \text{ [m/s]}$, and v the 1-min surface sustained wind in m/s. Not applicable to wind speeds much lower than 20 m/s. The wind speed is the 1-min surface sustained wind (usually expressed in knots [kt], hence the mph looks a bit strange in the graph).



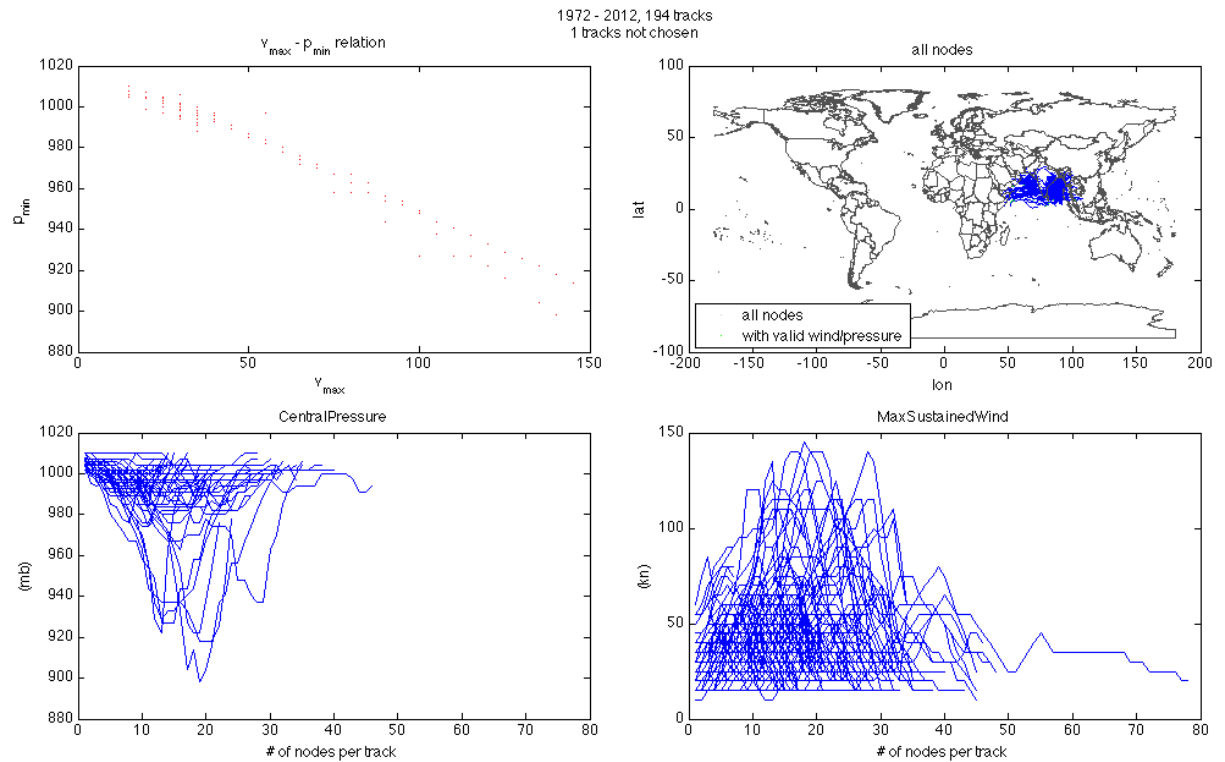
We kept the parameters as cited above, well aware of a certain level of 'mock precision'.

⁵ Liming Xu, 2010: A SIMPLE COASTLINE STORM SURGE MODEL BASED ON PRE-RUN SLOSH OUTPUTS. 29th Conference on Hurricanes and Tropical Meteorology, 10–14 May 2010, Tucson, Arizona.

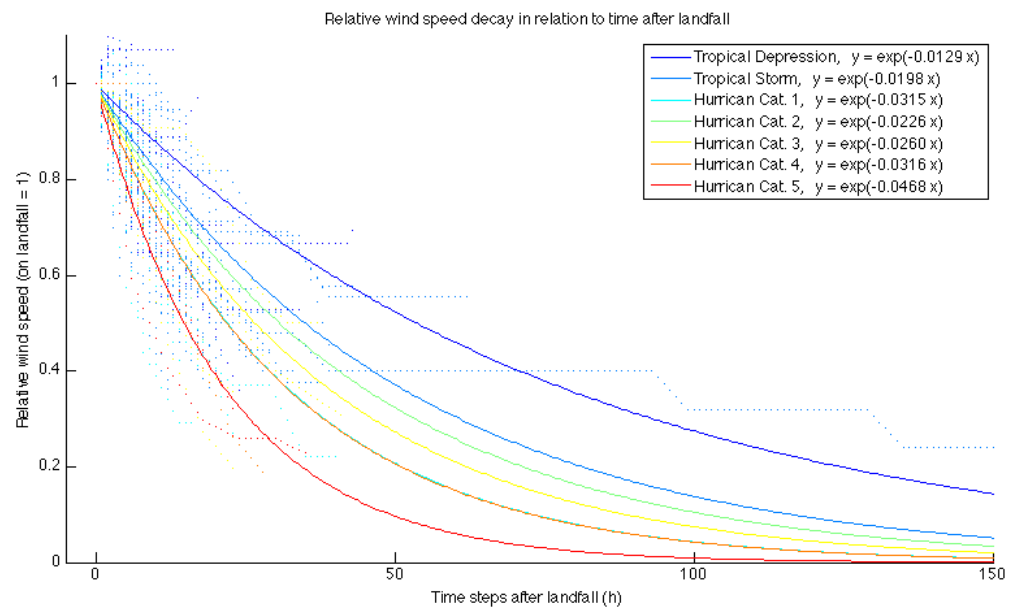
⁶ 1 mph = 0.44704 m/s, 1 foot = 0.3048 m

Appendix: Output of the TEST run for Bangladesh

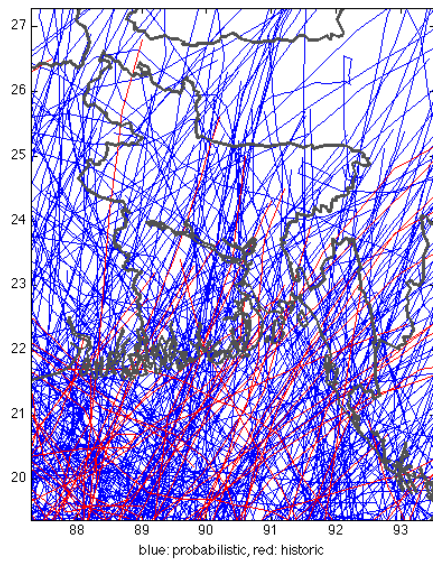
First, we obtain the full historic tropical cyclone dataset:



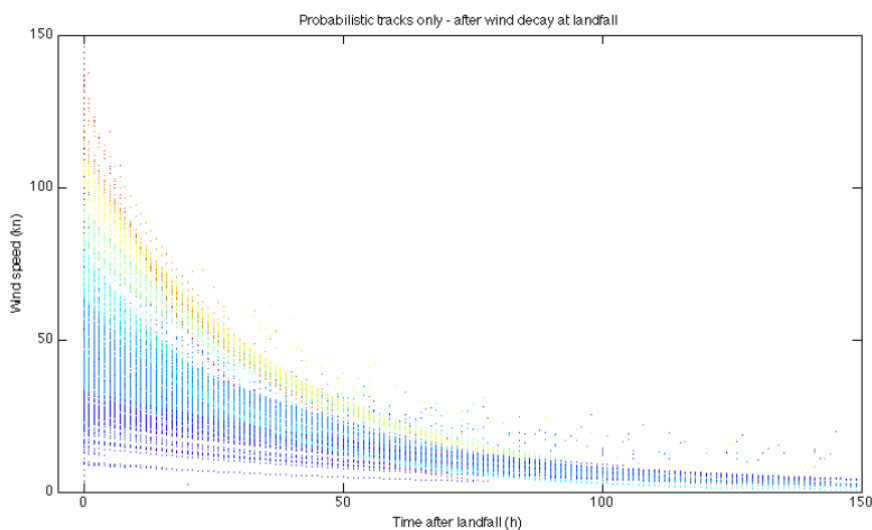
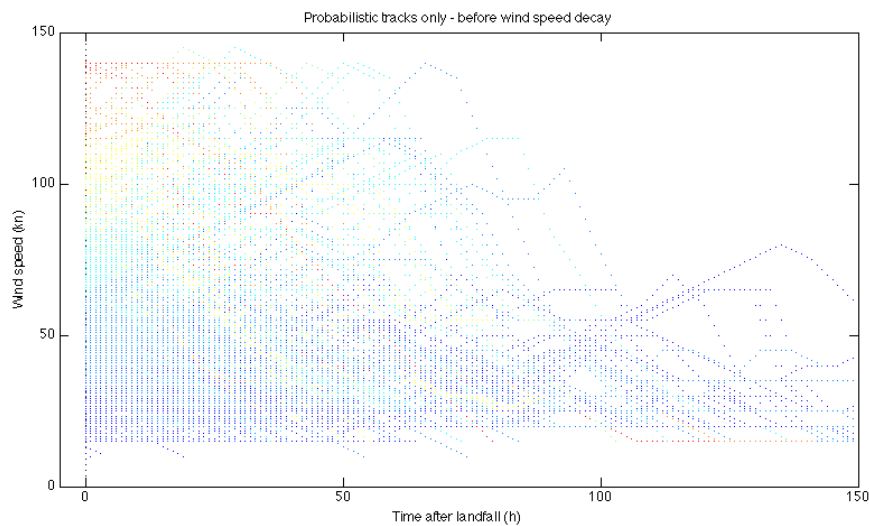
Second, we establish the wind speed decay relationship based on all historical events:



Third, we generate the probabilistic tracks:

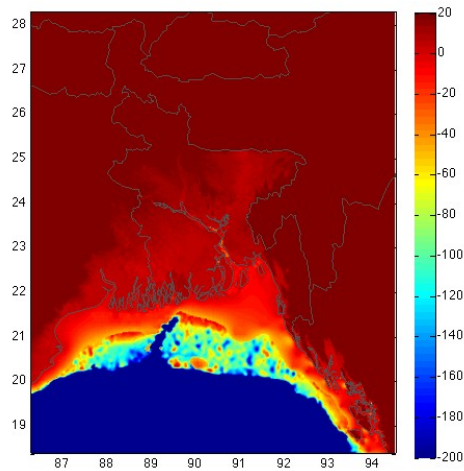


Fourth, we apply the decay to correct for probabilistic storms to show the same decay relationship as the historic events (correct for probabilistic storms be 'moved' inland with unrealistic 'open water' windspeeds):

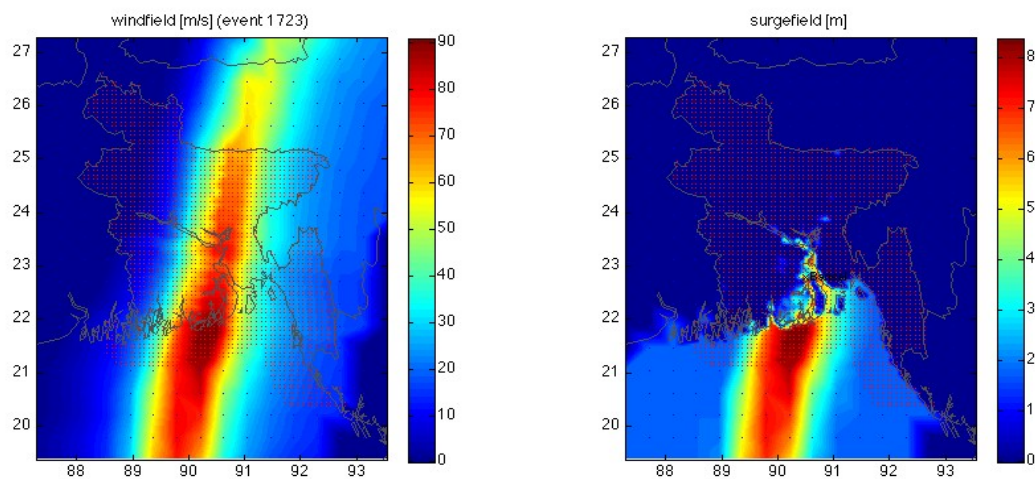


We then create the windfields for all events (both historic and probabilistic).
Not shown.

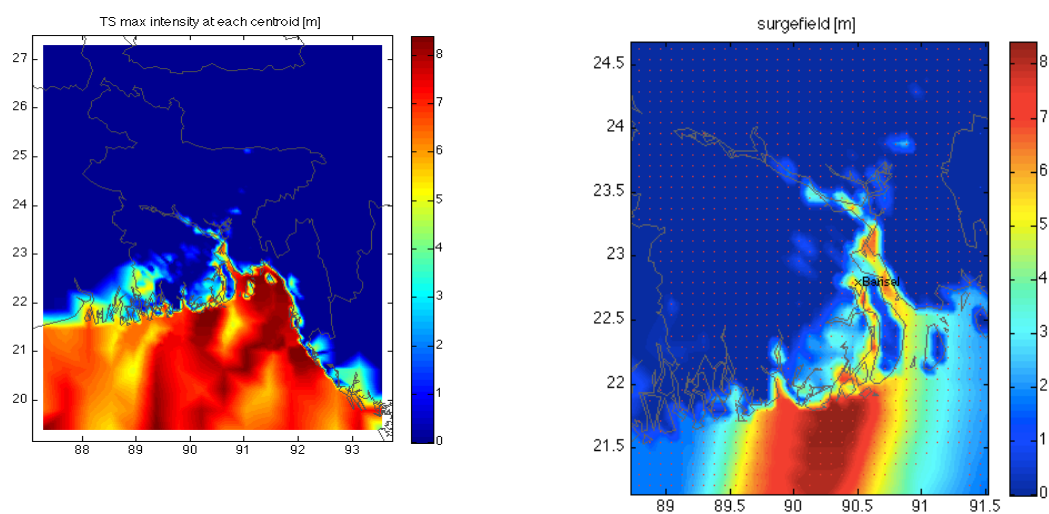
Fifth, we obtain the bathymetry:



Sixth, convert event-by-event into storm surge footprint:

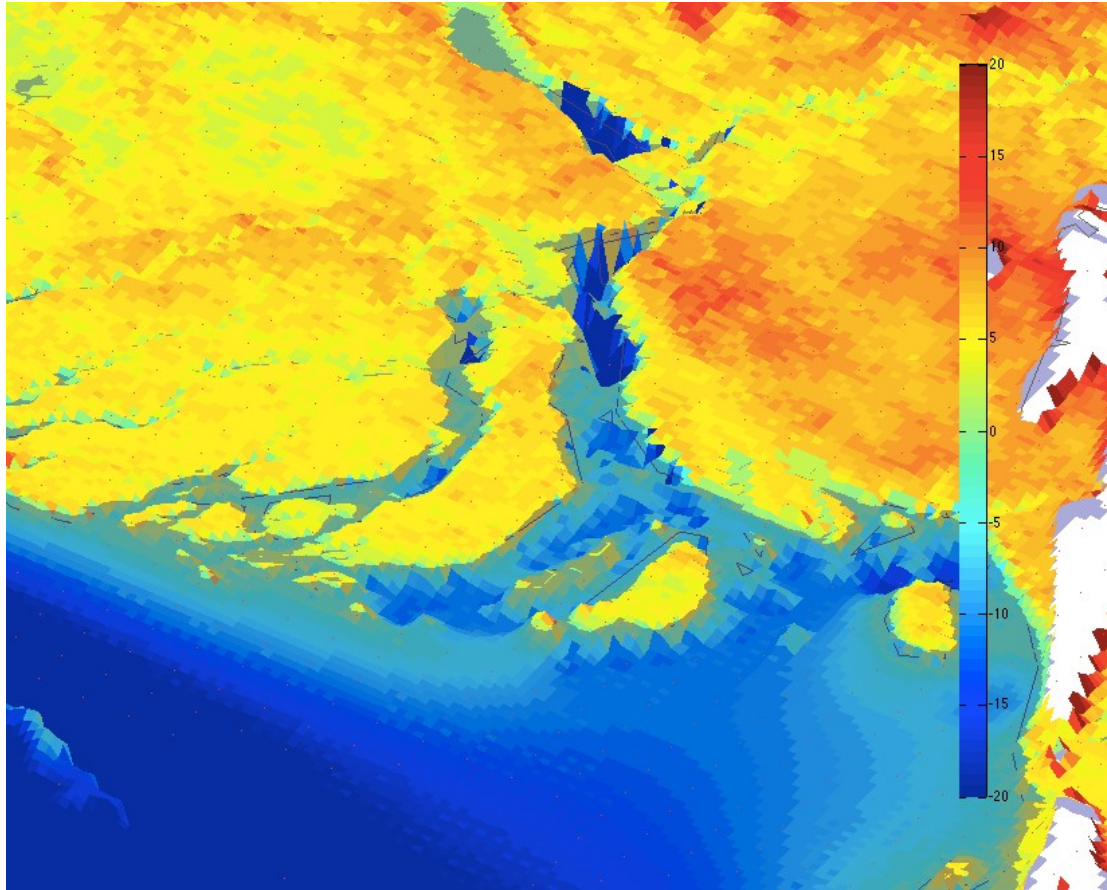


Seventh (and last), the maximum storm surge height obtained in the whole basin (and a zoom into the surge height above right):



Appendix – 3D plot

See `tc_surge_plot_3d` for 3D plots of surge height, with the surge level 'overshadowing' the elevation:



The bathymetry (light green and blue), topography (yellow to red) and the surge as semi-translucent surface (in a kind of light blue) to indicate the extent of surge (the maximum surge height of all events, not a particular one). 1 voxel 1' x 1' or approx. 1.7km x 1.7km, area approx. 3km². Tiny red dots indicate centroids, black line the seashore.