

3 standard glacier models:
are they well-posed?

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Glaciers Seminar

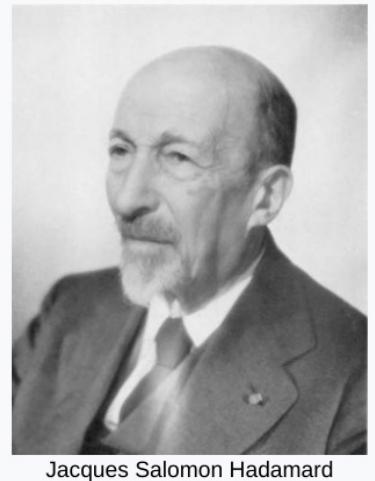
Well-posed problem

Article Talk

From Wikipedia, the free encyclopedia

In [mathematics](#), a **well-posed problem** is one for which the following properties hold:^[a]

1. The problem has a solution (*existence*)
2. The solution is [unique](#)
3. The solution's behavior changes [continuously](#) with the [data of the problem](#)



Jacques Salomon Hadamard

ex 1: well posed!

$$\begin{cases} -u''(x) = f(x) \\ u(0) = 0, \quad u(1) = 0 \end{cases}$$



best-known
proof that it is
well-posed is via
minimizing

$$J(u) = \int_0^1 \frac{1}{2} u''(x)^2 - f(x) u(x) dx$$

over functions satisfying
the b.c.s

solution:

$$u(x) = \int_0^1 g(x,y) f(y) dy$$

$$\text{where } g(x) = \begin{cases} y(1-x), & x \leq y \\ x(1-y), & y > x \end{cases}$$

ex 2: usually well-posed

$$\begin{cases} -u''(x) + au(x) = f(x) \\ u(0) = 0, \quad u(1) = 0 \end{cases}$$

issue: if e.g.
 $a = -\pi^2$ then
 $u(x) + c \sin(\pi x)$
is also a solution
for any c

well-posed if $a \neq -\pi^2, -4\pi^2, -9\pi^2, \dots$

otherwise multiple solutions

Uniqueness can fail

ex 3: rarely well-posed ← or "usually ill-posed"?

$$\begin{aligned} u'(x) &= f(x) \\ u(0) = 0, \quad u(1) &= 0 \end{aligned}$$

← issue:
 $u(x) = \int_0^x f(y) dy,$
but then only rarely
(for special f) does
 $0 = u(1) = \int_0^1 f(y) dy$
hold

existence can fail

but this is not a
math class ...

outline

problem 1: Glen-Stokes model to find \bar{u}, p for a fixed glacier geometry well-posed!

problem 2: shallow ice approximation for steady glacier geometry uniqueness?

problem 3: time-dependent Glen-Stokes for evolving glacier geometry open!

problem 1 is to use the Glen-Stokes equations, within ice of known geometry, and ^{with} dynamical boundary conditions, to determine \vec{u}, p



Symbols (for all problems)

$x = (x_1, x_2)$ horizontal coordinate(s)

z vertical coordinate

s ice surface elevation

b bedrock elevation

\vec{u} ice flow velocity

p ice fluid pressure

a surface mass balance

$D\vec{u}$ strain rate tensor

$|D\vec{u}|$ Frobenius norm of $D\vec{u}$ ($\|D\vec{u}\| = \sqrt{\frac{1}{2} (D\vec{u})^T D\vec{u}}$)

$\nu = \nu(D\vec{u})$ ice viscosity (assumed isothermal)

ρ_i ice density

\vec{g} gravity

ν_o viscosity scale

$n=3$ Glen exponent

equations within ice Λ :

$$\nu(D\vec{u}) = \nu_0 \left(|D\vec{u}|^2 + \varepsilon^2 \right)^{-\frac{2}{3}}$$

$$-\nabla \cdot (2\nu(D\vec{u}) D\vec{u}) + \nabla p = \rho_i \vec{g}$$

$$\nabla \cdot \vec{u} = 0$$

on boundary

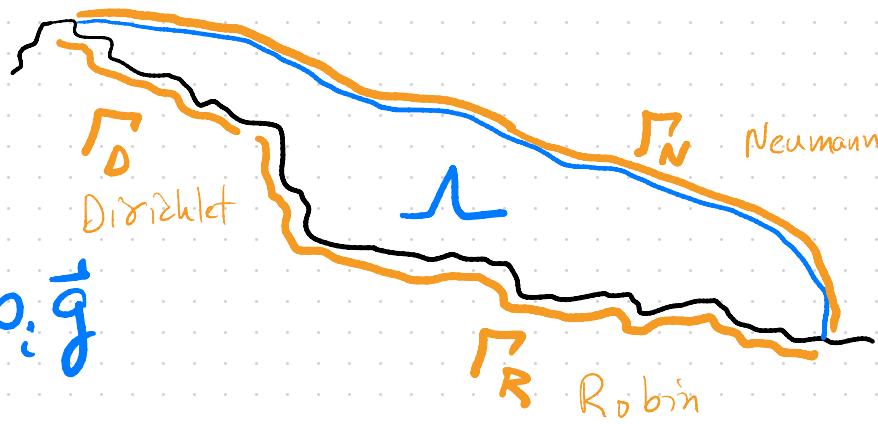
$$\sigma \hat{n} = 0$$

on Γ_N

$$\vec{u} = 0$$

on Γ_D

$$\begin{aligned} & (\sigma \hat{n})_t + \beta \vec{u}_t = 0 \\ & \vec{u} \cdot \hat{n} = 0 \end{aligned} \quad \left. \right\} \text{on } \Gamma_R$$



notation

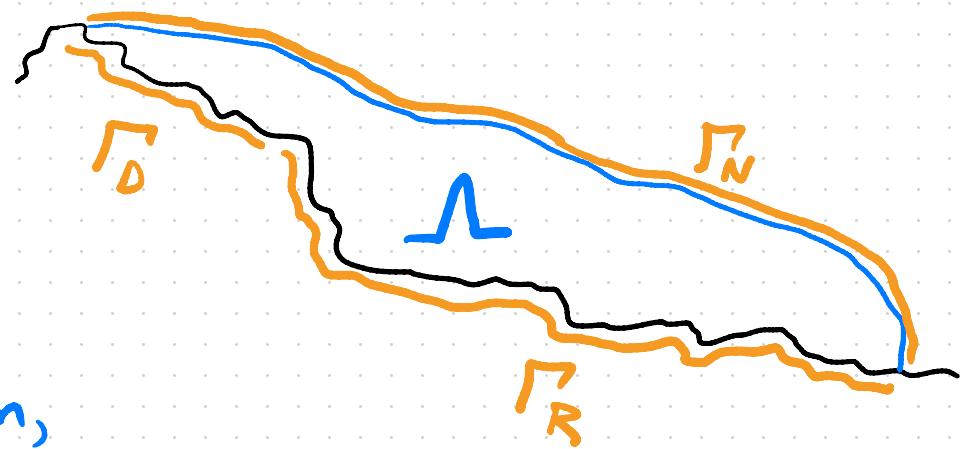
$$\sigma = 2\nu(D\vec{u})D\vec{u} - pI$$

\hat{n} = (unit outward normal)

$$\vec{w}_t = (I - \hat{n}\hat{n}^T)\vec{w}$$

= (tangential part
of \vec{w})

- key idea in problem 1 is that the domain Λ itself is the "input data" to the model



- the surface elevation, defining Γ_N , is data in problem 1

Theorem (Jouvet & Rappaz 2011, Isaac et al 2015)

if $s(x)$ and $b(x)$ are continuous and Lipschitz,
and if $\varepsilon > 0$, and if either Γ_0 is positive measure,
or if the sliding law on Γ_R gives sufficient
resistance, then there are unique fields
 \vec{u}, p satisfying the weak form of the
Glen-Stokes model

$$\text{Lip: } |s(x) - s(y)| \leq L \|x - y\|$$

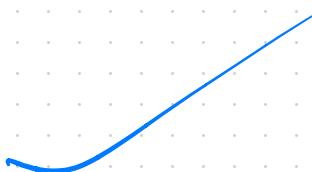
What, physically-speaking, are these caveats about?

" $\varepsilon > 0$ ": avoid ∞ viscosity or zero strain.
(and uniqueness)

"the sliding law on Γ_B gives sufficient resistance":

so the glacier doesn't slide off

" $s(x)$ and $b(x)$ are continuous and Lipschitz":



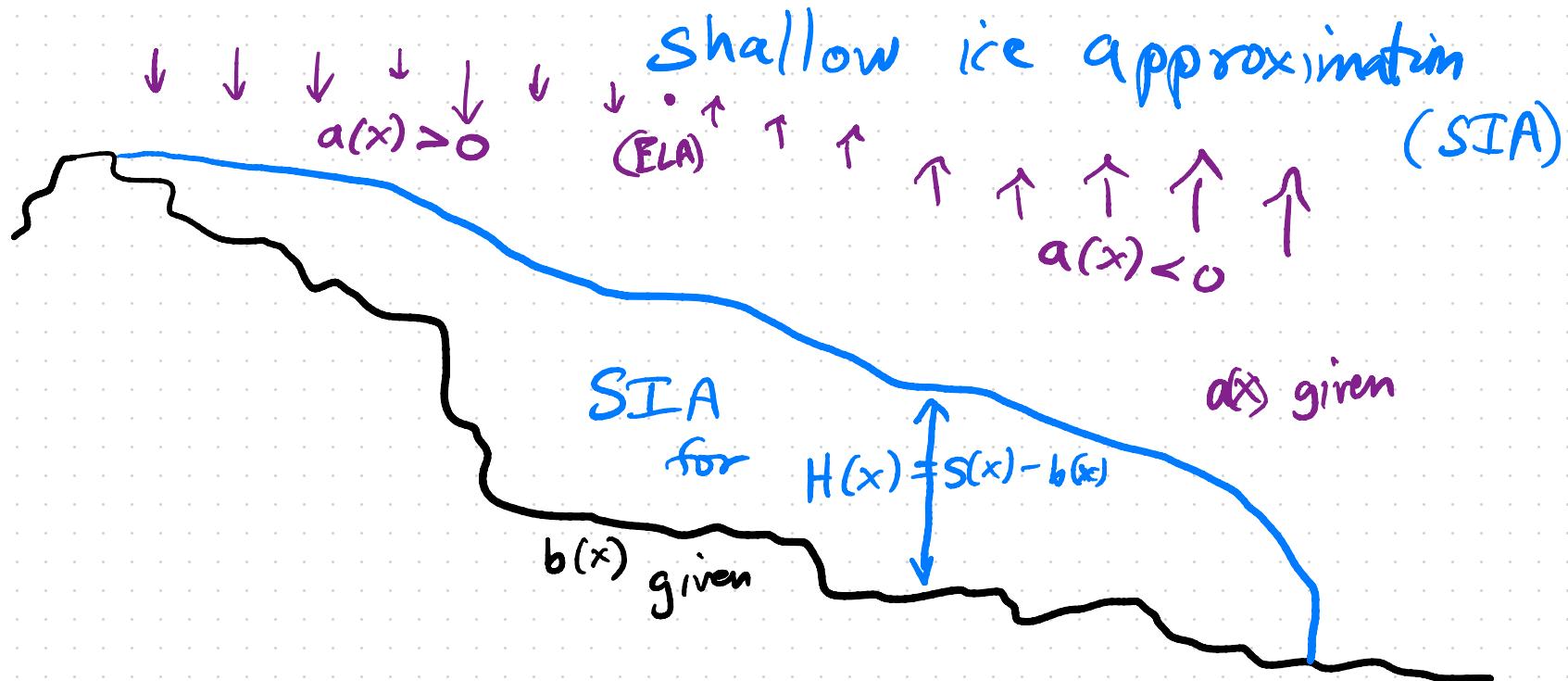
problem 1 summary

if you know the geometry of a glacier
and choose a reasonable sliding law, then
you can find its velocity and pressure,
thus its stress field

Comments:

- problem 1 is instantaneous, but it is not
Steady state
- if you solve problem 1 numerically,
your implementation may make errors,
(shouldn't?)
but you won't [^] be confused about
their sources

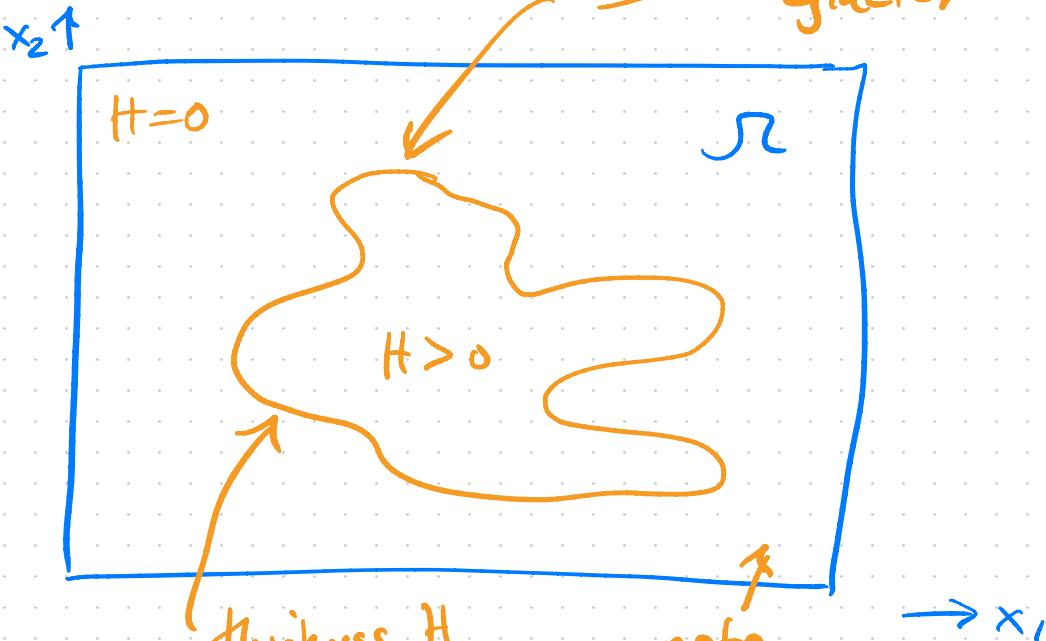
problem 2 is to find the steady state thickness, in a steady climate (SMB $a(x)$), of a glacier modeled using the isothermal



domain and data for problem 2

$a(x)$ and $b(x)$
are given on
all of the
"map" $\Omega \subset \mathbb{R}^2$

solution is the
glacier



thickness H
and flux \vec{q}
go to zero
along
free boundary

note
 $\vec{q} = 0$
where ice
free

equations and inequalities on "map" \mathcal{S}

$$H \geq 0$$

$$\vec{g} = -\Gamma H^{n+2} (\nabla(H+b))^{n-1} \nabla(H+b) \quad \left. \begin{array}{l} \text{vertically-} \\ \text{integrated} \\ \text{ice flux} \end{array} \right\}$$

$$\nabla \cdot \vec{g} - a \geq 0$$

$$H(\nabla \cdot \vec{g} - a) = 0$$

the surface
elevation

where

$\Gamma > 0$ constant, a rescaled ice softness

note: this is the non-sliding version, but certain sliding laws are possible ...

Theorem (Jouret & Bueler, 2012)

if $b(x)$ has an integrable gradient over \mathbb{R} ,
and if $a(x)$ is (merely) integrable, then there
exists a steady thickness $H(x)$, for which
 $H(x)^{8/3} \leftarrow (n=3 \text{ case})$ is a well-behaved function with
an integrable gradient, satisfying the
weak form of problem 2

What, physically-speaking, are these caveats about?

" $H(x)^{8/3}$ is a well-behaved function...": } next slide

" $b(x)$ has an integrable gradient over S^2 ":

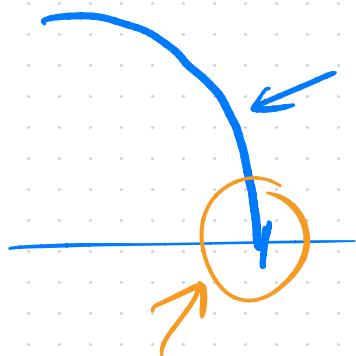
2 slides a head

↑
Ravant (1970)

"there exists ...":

uniqueness not known!

re garding $\frac{8}{3}$ power:



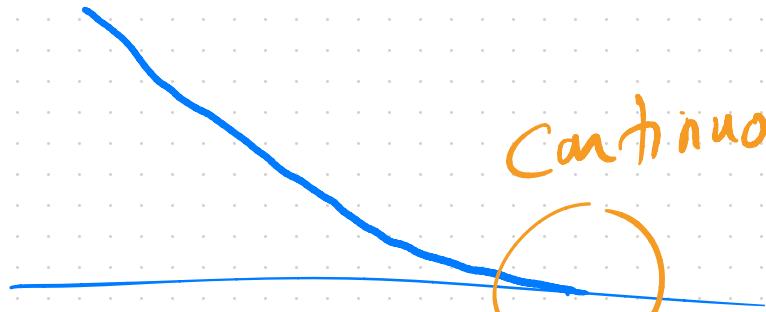
f. shape

$\frac{6}{3}$ power



infinite
gradient

$\frac{8}{3}$ power



continuous
gradient

regarding irregular $b(x)$:

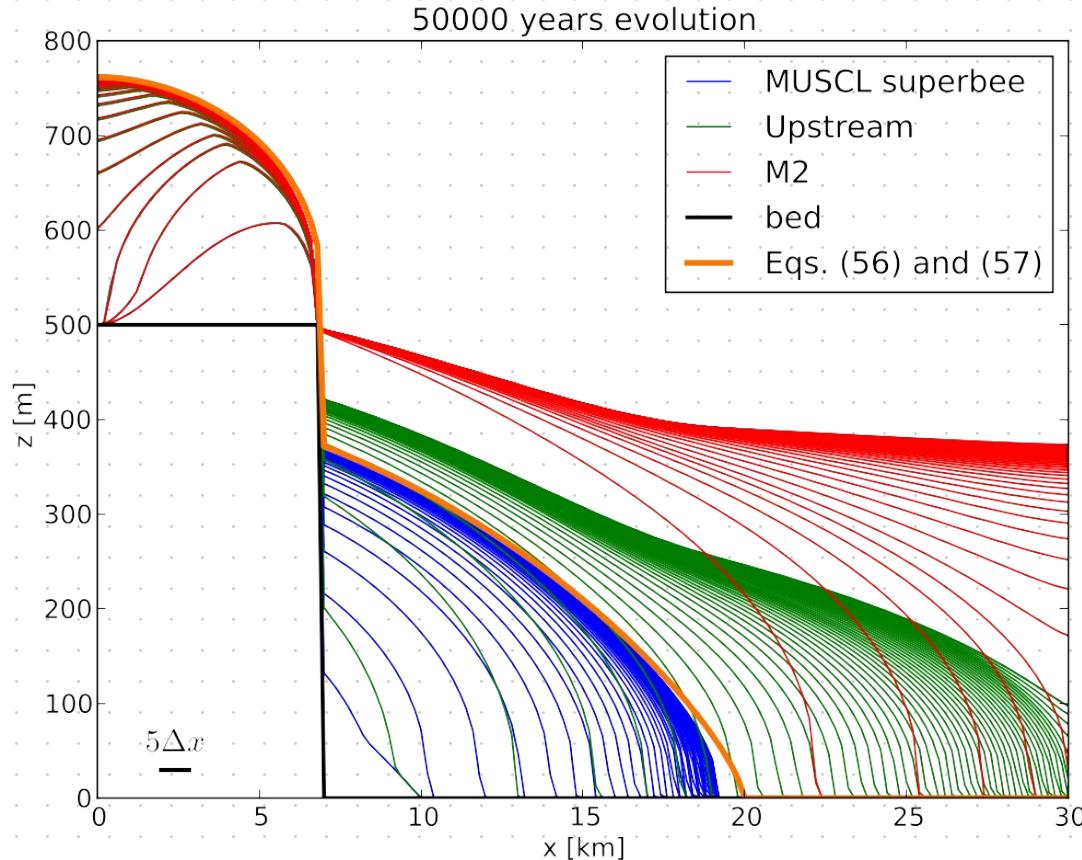


figure from Jarosch et al (2013)

numerical confusion?

- problem 2 remains hard to solve numerically on sufficiently - rough bed topography
- the methods which work best must aggressively enforce edge continuity of the flux \vec{q} (Brinkhoff...)
- I think the underlying numerical issue is the "same" as the uniqueness issue

- how about the time dependent case?

Theorem (Calvo et al 2003 in 1D, Piersanti & Temam 2023 in 2D)

if the bed is flat, and if $a(t, x)$ is integrable, then there is a unique $H(t, x)$, with $H(t, x)^{8/3}$ well-behaved, satisfying the time dependent version of problem 2

- this is flat-bed and only SIA

- note

$$a = a(x)$$

in problem 2, is not

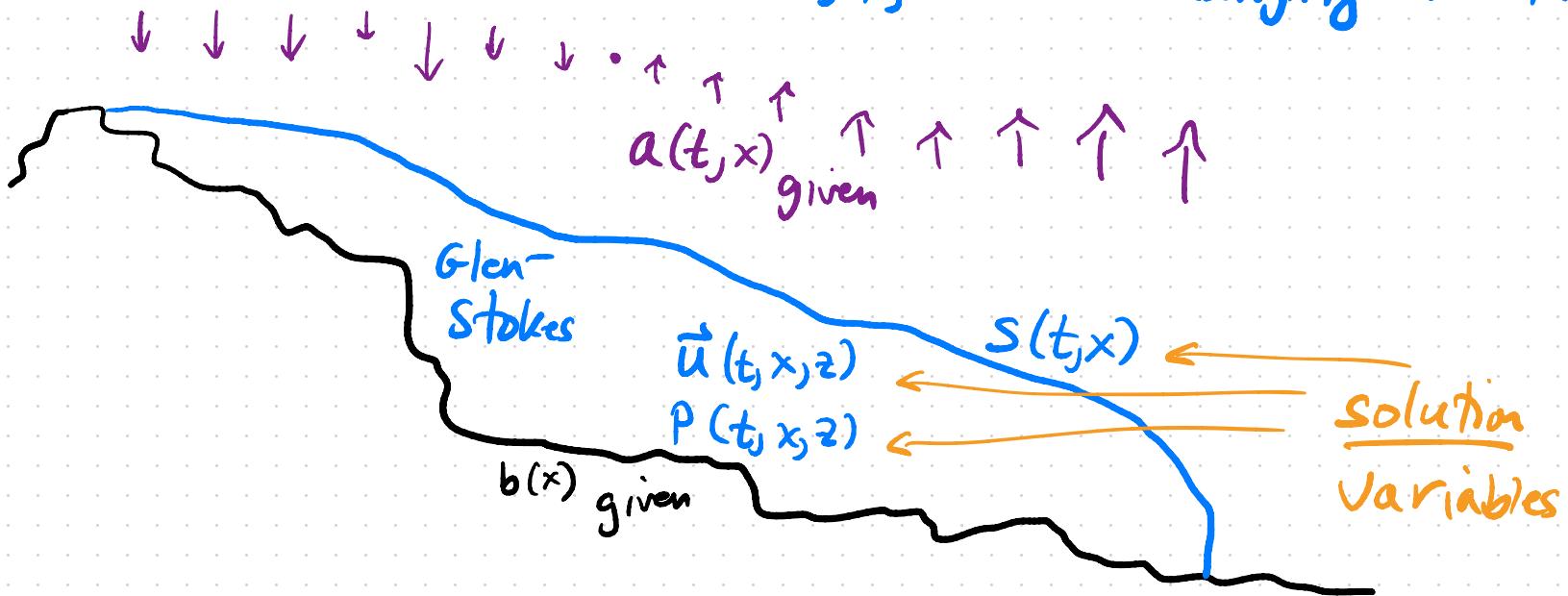
$$a = a(x, z) = a(x, s)$$

- the fact that uniqueness is open in problem 2

is not about elevation - accumulation

feed back (for which actual non-uniqueness
can be shown)

problem 3 is to use the Glen-Stokes equations,
and the surface kinematical equation, to
find the evolution of the surface elevation s ,
with co-evolution of \vec{u}, p , in a changing climate



Data of problem 3:

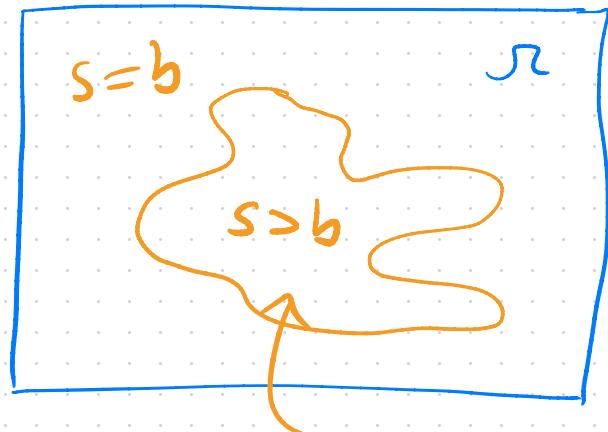
Same as time-dependent form of
problem 2,
at least in the non-sliding
case, namely

1. $a(t, x)$ SMB

2. $b(x)$ topography,

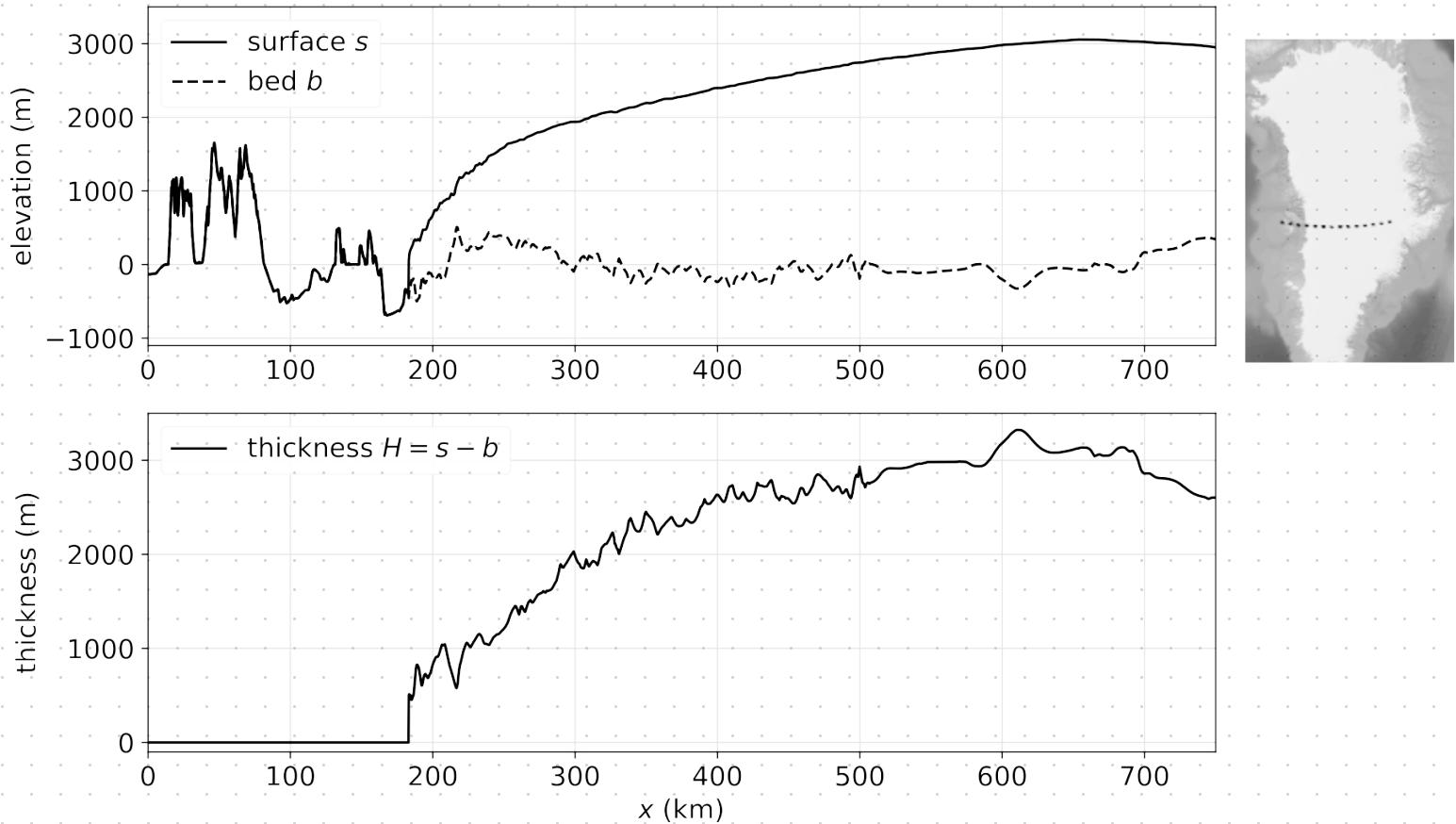
with initial surface elevation

3. $s_0(x)$ satisfying $s_0 \geq b$



Solution of
problem 3 is the
evolving glaciated
area, and the evolving
 $s(t, x)$, and the
co-evolving \vec{u}, p

why not phrase in terms of thickness $H = s - b$?



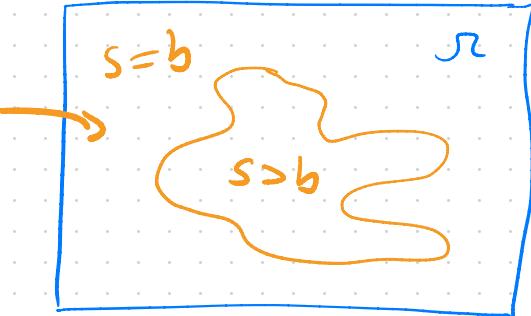
equations in Ω :

$$s \geq b$$

$$\frac{\partial s}{\partial t} - \vec{u}|_s \cdot \vec{n}_s - a \geq 0$$

$$(s-b) \left(\frac{\partial s}{\partial t} - \vec{u}|_s \cdot \vec{n}_s - a \right) = 0$$

set
 $\vec{u}|_{s=0}$
 where
 ice
 free

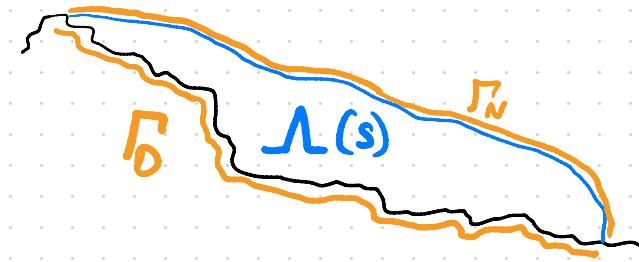


equations in $\Lambda(s) = \{x, z : b < z < s\}$

$$\nu(D\vec{u}) = \nu_0 \left(|D\vec{u}|^2 + \varepsilon^2 \right)^{-\frac{2}{3}}$$

$$-\nabla \cdot (2\nu(D\vec{u}) D\vec{u}) + \nabla p = \rho_i \vec{g}$$

$$\nabla \cdot \vec{u} = 0$$



with nonsliding for simplicity:

$$\sigma \hat{n} = 0$$

$$\vec{u} = 0$$

on Γ_N

on Γ_D

surface kinematical equation (SKE)

$$\frac{\partial s}{\partial t} = \underbrace{\vec{u}|_s \cdot \vec{n}_s}_{①} + \underbrace{a}_{②}$$

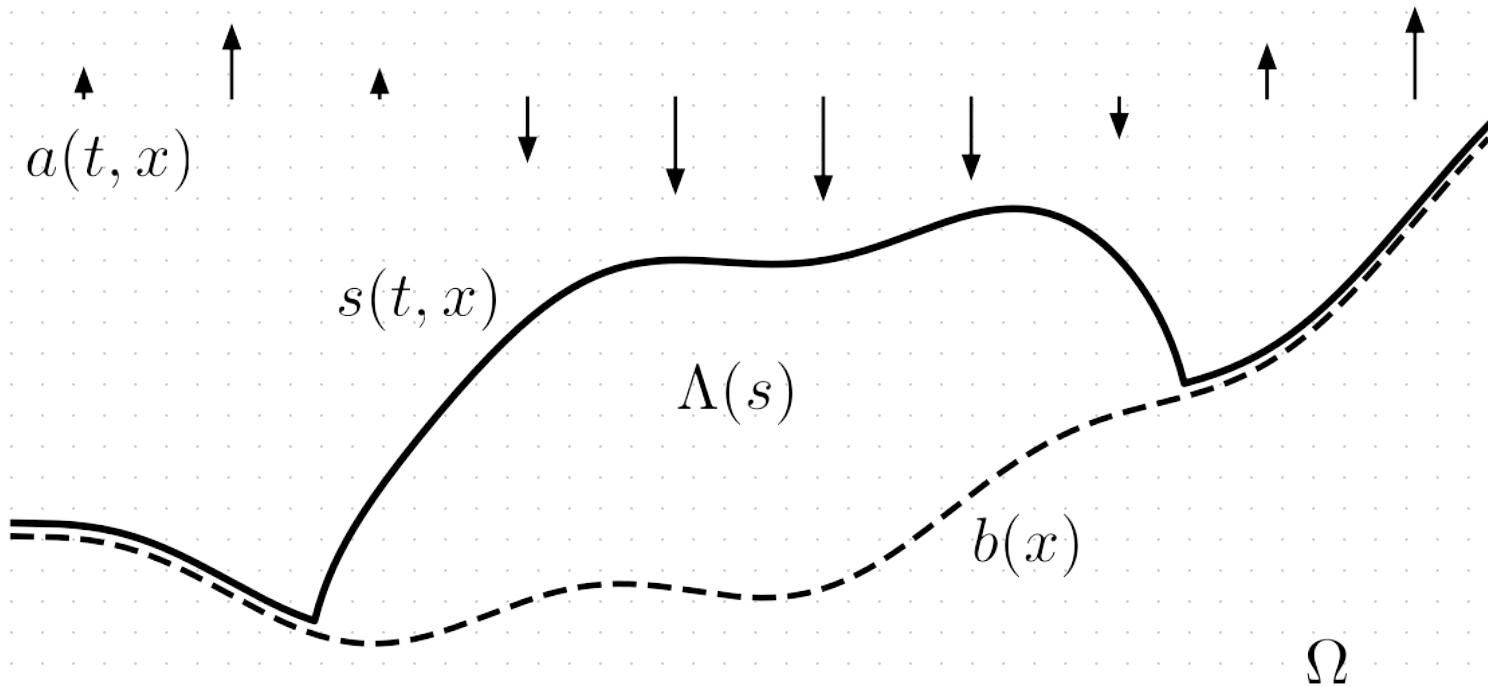
def:

$$\vec{n}_s = \left(-\frac{\partial s}{\partial x}, -\frac{\partial s}{\partial y}, 1 \right)$$

is a surface normal

- the SKE says the surface moves up or down, in an Eulerian view, according to whether the (sloped) surface ice is moving and/or accumulating/ablatting
- it is not anyone's "boundary condition"

domain of ice $\Lambda(s)$ is part of the solution



what "well-posed" would mean for problem 3

given data $s_0(x)$, $a(t,x)$, $b(x)$, then
at least for a short time (?), there
are unique solution fields

$$s(t,x)$$

$$\vec{u}(t,x,z)$$

$$p(t,x,z)$$

with $s(t,x) \geq b(x)$, and \vec{u}, p only defined within
the ice ($x, z \in \Lambda(s(t,x))$), which simultaneously satisfy the
weak form of all inequalities, equations, and b.c.s
of problem 3

- well-posedness, and existence or uniqueness as separate questions, are totally open questions for problem 3
-

do you believe it is well-posed?

- if not, you should either add/subtract equations to make it so, or you should give up entirely on numerical (and ML) glacier simulations

my current strategy for problem 3

- ① only consider a single backward Euler time step (Bueler, 2024 arxiv):

$$s^k \geq b$$
$$\frac{s^k - s^{k-1}}{\Delta t} - \vec{u}|_{s^k} \cdot \vec{n}_{s^k} - a \geq 0$$
$$(s^k - b) \left(\frac{s^k - s^{k-1}}{\Delta t} - \vec{u}|_{s^k} \cdot \vec{n}_{s^k} - a \right) = 0$$

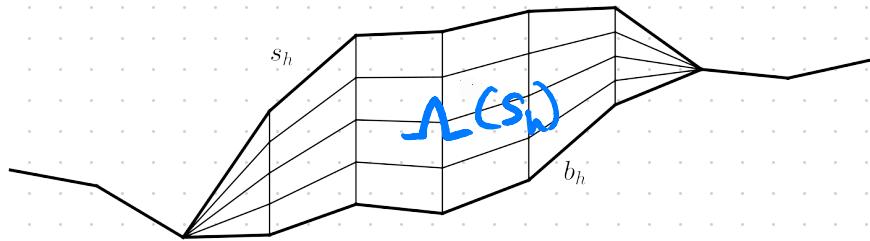
\otimes

s^{k+}
 $\approx s^{(k+1)}$

- ② accept that existence and/or uniqueness may be restricted to small values of $\Delta t > 0$

- ③ work numerically for now, trying to prove that the finite element (FE) approximation is well-behaved (Bueler, 2024)
- ④ use a mixed FE formulation of the SKE complementarity problem \otimes , in which the flux \vec{q} , or the surface gradient ∇S , comes from an edge-continuous "H(div)" FE space
- (see Brinkhoff, 2024 re a choice like this)

- ⑤ use Firedrake
- ⑥ use an extruded mesh for $\Lambda(s)$:



- ⑦ accept that a regularized SKE

$$\frac{\partial s}{\partial t} = \vec{u}|_s \cdot \vec{n}_s + a + \underbrace{s \nabla^2 s}_{\text{regularization}}$$

may be necessary to put s in a reasonable function space (avoids overhangs)

- ⑧ try to sell such a general "fast fluid layer evolution" FE model, with optimal numerical scaling, as a useful scientific computation
- ⑨ ... with first demo being high-resolution 3D Stokes simulation of all glaciers in an Alaskan mountain range

thanks for your attention!

I am happy to try to
answer questions

references

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