

An aerial photograph of a glacier system with several icebergs floating in the water. The glacier is a mix of white and blue ice, with dark rock visible at the edges. The water is a pale blue-grey color.

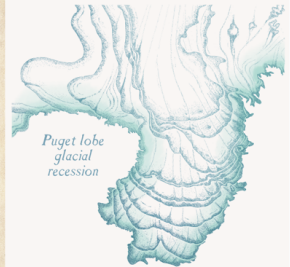
# Finite element errors in Stokes models of glacier evolution

---

Ed Bueler, University of Alaska Fairbanks

JMM 2025, Seattle

# Puget Sound glaciation



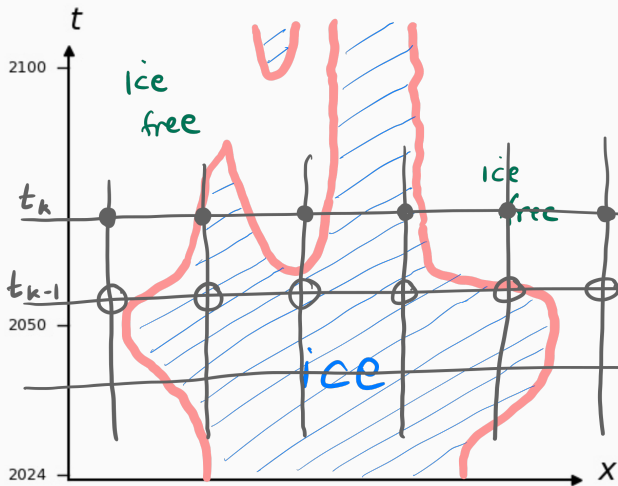
# glacier evolution in space-time

- a) what is true within the ice?
- b) what is true on bare land?
- c) what is true on free boundary?



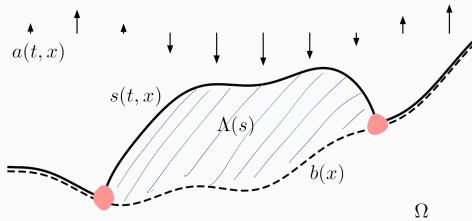
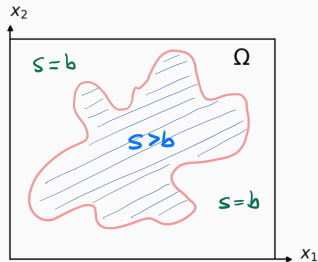
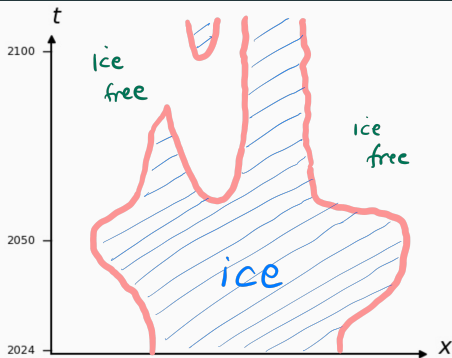
# glacier evolution in space-time ...implicitly?

- a) what is true within the ice?
- b) what is true on bare land?
- c) what is true on free boundary?



1. introduction to the standard glacier model
2. well-posed implicit steps?
3. finite element approximation

# the 3 views



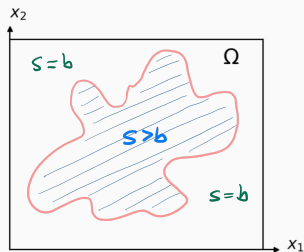
## what is true everywhere in $[0, T] \times \Omega$ ?

- $\Omega \subset \mathbb{R}^2$  fixed simulation domain
- $x = (x_1, x_2) \in \Omega$
- $a(t, x)$  surface mass balance (SMB)
- $b(x)$  bed elevation
- $s(t, x)$  *solution* surface elevation
- $\mathbf{n}_s = (-\nabla s, 1)$  surface-normal vector
- $\mathbf{u}|_s(t, x)$  surface value of ice velocity, extended by zero to bare land
- nonlinear complementarity problem (NCP) in  $[0, T] \times \Omega$ :

$$s - b \geq 0$$

$$\frac{\partial s}{\partial t} - \mathbf{u}|_s \cdot \mathbf{n}_s - a \geq 0$$

$$(s - b) \left( \frac{\partial s}{\partial t} - \mathbf{u}|_s \cdot \mathbf{n}_s - a \right) = 0$$



# the free-boundary problem for glacier surface elevation

- surface kinematical equation (SKE) in glaciology:

$$\frac{\partial s}{\partial t} - \mathbf{u}|_s \cdot \mathbf{n}_s - a = 0$$

- our time-dependent, nonlinear complementarity problem (NCP) is the underlying *free-boundary* meaning of the SKE:

$$s - b \geq 0$$

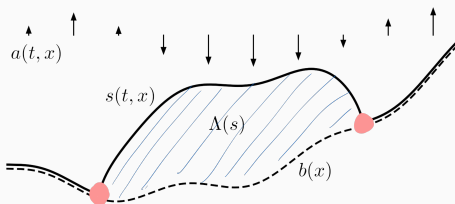
$$\frac{\partial s}{\partial t} - \mathbf{u}|_s \cdot \mathbf{n}_s - a \geq 0$$

$$(s - b) \left( \frac{\partial s}{\partial t} - \mathbf{u}|_s \cdot \mathbf{n}_s - a \right) = 0$$

- applies regardless of dynamical model within the ice
  - this NCP appears in (Calvo et al 2003), but only for shallow ice
- state-of-the-art numerically: first-order explicit time-stepping



# what is true within the ice?



- fix  $t$  and define:

$$\Lambda(s) = \{(x, z) : b(x) < z < s(t, x)\} \subset \mathbb{R}^3$$

- Glen-Stokes equations in  $\Lambda(s)$  with  $p = (n + 1)/n \approx 4/3$ :

$$-\nabla \cdot (2\nu(D\mathbf{u}) D\mathbf{u}) + \nabla p = \rho_i \mathbf{g}$$

$$\nabla \cdot \mathbf{u} = 0$$

$$\nu(D\mathbf{u}) = \nu_0 |D\mathbf{u}|^{p-2} \quad \leftarrow \text{needs regularization}$$

- boundary conditions:

$$(2\nu(D\mathbf{u}) D\mathbf{u} - pI) \mathbf{n}_s = \mathbf{0} \quad \text{on } \Gamma_s \subset \partial\Lambda(s)$$

$$\mathbf{u} = \mathbf{0} \text{ or } f(\mathbf{u}, D\mathbf{u}) = 0 \quad \text{on } \Gamma_b \subset \partial\Lambda(s)$$

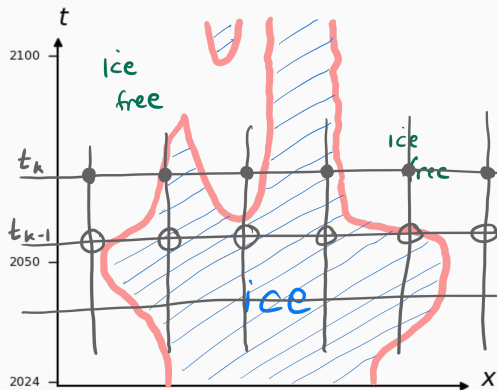
## the standard model for glacier evolution

$$\begin{aligned}s - b &\geq 0 && \text{in } \Omega \\ \frac{\partial s}{\partial t} - \mathbf{u}|_s \cdot \mathbf{n}_s - a &\geq 0 \\ (s - b) \left( \frac{\partial s}{\partial t} - \mathbf{u}|_s \cdot \mathbf{n}_s - a \right) &= 0 \\ -\nabla \cdot (2\nu_0 |D\mathbf{u}|^{p-2} D\mathbf{u}) + \nabla p &= \rho_i \mathbf{g} && \text{in } \Lambda(s) \\ \nabla \cdot \mathbf{u} &= 0 \\ (2\nu(D\mathbf{u})D\mathbf{u} - pI) \mathbf{n}_s &= \mathbf{0} && \text{on } \Gamma_s \subset \partial\Lambda(s) \\ \mathbf{u} = \mathbf{0} \quad \text{or} \quad f(\mathbf{u}, D\mathbf{u}) &= 0 && \text{on } \Gamma_b \subset \partial\Lambda(s)\end{aligned}$$

### standard model

*an NCP in  $[0, T] \times \Omega$ , for fixed  $\Omega \subset \mathbb{R}^2$ , coupled to a Glen-Stokes problem within the ice, in  $\Lambda(s) \subset \mathbb{R}^3$*

# the standard model wants “fully-implicit” time-stepping



standard model as a dynamical system

an inequality-constrained **differential algebraic system** in  $\infty$  dimensions

introduction to the standard glacier model

well-posed implicit steps?

finite element approximation

# implicit time-step problem is a variational inequality

- weak form of time-discretized NCP; details found in preprint\*
- $\mathcal{X}$  denotes Banach space of elevation functions on  $\Omega$ ;  $b \in \mathcal{X}$
- for  $\Delta t > 0$  define:

$$\mathcal{K} = \{r \in \mathcal{X} : r \geq b \text{ a.e. } \Omega\}$$

$$\ell^k(x) = s(t_{k-1}, x) + \int_{t_{k-1}}^{t_k} a(t, x) dt$$

$$F(s)[q] = \int_{\Omega} (s - \Delta t \mathbf{u}|_s \cdot \mathbf{n}_s) q$$

## Definition

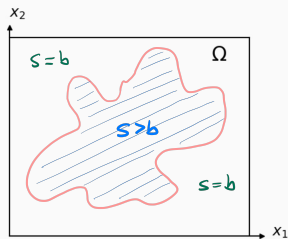
the *backward Euler time-step problem*, for  $t_k \in [0, T]$ , is to find the surface elevation  $s \approx s(t_k, x) \in \mathcal{K}$  solving the variational inequality (VI)

$$F(s)[r - s] \geq \ell^k[r - s] \quad \text{for all } r \in \mathcal{K}$$

---

\*Bueler (2024). *Surface elevation errors in finite element Stokes models for glacier evolution*, [arxiv 2408.06470](https://arxiv.org/abs/2408.06470)

# well-posedness is only conjectural



## Conjecture (B '24)

for some  $q > 2$ , with  $\mathcal{X} = W^{1,q}(\Omega)$ , and defining the admissible surface elevation subset

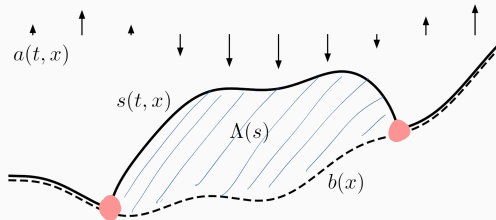
$$\mathcal{K} = \{r \geq b\} \subset \mathcal{X},$$

the backward Euler time-step VI problem,

$$F(s)[r - s] \geq \ell^k[r - s] \quad \text{for all } r \in \mathcal{K},$$

is well-posed for  $s \in \mathcal{K}$

# what's behind the conjecture? 1



## Theorem (Jouvet & Rappaz, 2011)

over a fixed domain  $\Lambda$ , the Glen-Stokes problem

$$-\nabla \cdot (2\nu_0 |D\mathbf{u}|^{p-2} D\mathbf{u}) + \nabla p = \rho_i \mathbf{g} \quad \text{in } \Lambda$$

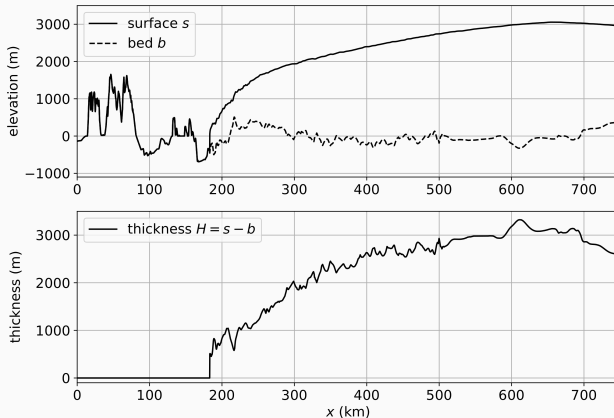
$$\nabla \cdot \mathbf{u} = 0$$

$$(2\nu(D\mathbf{u})D\mathbf{u} - pI) \mathbf{n}_s = \mathbf{0} \quad \text{on } \Gamma_s \subset \partial\Lambda$$

$$\mathbf{u} = \mathbf{0} \quad \text{or} \quad f(\mathbf{u}, D\mathbf{u}) = 0 \quad \text{on } \Gamma_b \subset \partial\Lambda$$

is well-posed for  $(\mathbf{u}, p) \in W_0^{1,p}(\Omega; \mathbb{R}^3) \times L^{p'}(\Omega)$

## what's behind the conjecture? 2



- question: use ice surface elevation  $s$  or thickness  $H = s - b$  for geometry in the standard model?
- answer: prefer  $s$  for smoothness reasons



## what's behind the conjecture? 3



- the conjecture is that  $s$  exists in  $\mathcal{X} = W^{1,q}(\Omega)$  for some  $q > 2$
- however, margin shape of glaciers is a hard modeling problem

# what's behind the conjecture? 4

## Conjecture (actual; B '24)

For  $s \in \mathcal{K} = \{r \geq b\} \subset \mathcal{X} = W^{1,q}(\Omega)$ , for some  $q > 2$ , let  $\Lambda(s) = \{b < z < s\} \subset \mathbb{R}^3$ , and suppose  $\mathbf{u}$  solves Glen-Stokes over  $\Lambda(s)$ . Define

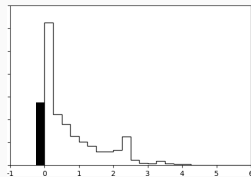
$$\Phi(s) = -\mathbf{u}|_s \cdot \mathbf{n}_s,$$

a map  $\Phi : \mathcal{K} \rightarrow \mathcal{X}'$ . Then  $\Phi$  is  $q$ -coercive: there is  $\alpha > 0$  so that

$$(\Phi(s) - \Phi(r)) [r - s] \geq \alpha \|r - s\|_{\mathcal{X}}^q \quad \text{for all } r, s \in \mathcal{K}$$

## numerical evidence?

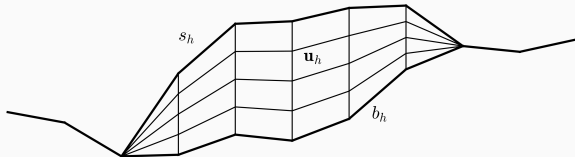
- from surface elevation samples  $r, s$ , compute numerical ratios  $\frac{(\Phi(r) - \Phi(s))[r - s]}{\|r - s\|_{\mathcal{X}}^q}$



introduction to the standard glacier model

well-posed implicit steps?

finite element approximation



## continuum problem

find  $s \in \mathcal{K} = \{r \geq b\}$  which solves the backward-Euler step VI:

$$F(s)[r - s] \geq \ell^k[r - s] \quad \text{for all } r \in \mathcal{K}$$

- proposed FE spaces:

- $b_h, s_h \in P_1$
- $\mathbf{u}_h, p_h \in P_2 \times P_1$ , on an extruded mesh

← one possibility

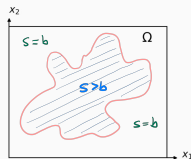
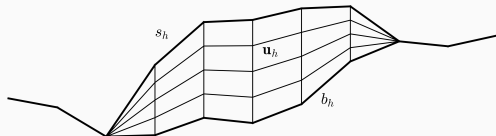
- FE method for  $s_h \in \mathcal{K}_h = \{r_h \geq b_h\}$ :

$$F_h(s_h)[r_h - s_h] \geq \ell^k[r_h - s_h] \quad \text{for all } r_h \in \mathcal{K}_h$$

- mostly conforming, especially if one pretends  $b_h = b$

- extruded-mesh Stokes solver infrastructure:  
[github.com/bueler/stokes-extrude](https://github.com/bueler/stokes-extrude)
- solver in Firedrake/PETSc:
  - VI-adapted, reduced-space Newton method
  - for admissible surface elevation iterates  $s_h^{(j)} \geq b_h$ , each residual evaluation  $F(s_h^{(j)})$  requires solving the Glen-Stokes model on  $\Lambda(s_h^{(j)})$  to get  $\mathbf{u}_h^{(j)}$
- accelerate using new FASCD nonlinear multigrid scheme (Bueler & Farrell, 2024), using above solver as the smoother

# error theorem for a backward Euler step



## Theorem (B '24)

Make assumptions sufficient for well-posedness. Let  $\Omega_A(s)$  be the exact active set, and  $\Pi_h$  the interpolation into  $\mathcal{K}_h = \{r_h \geq b_h\}$ . Then the FE error in the new  $(t_k)$  surface elevation is bounded by 3 terms:

$$\begin{aligned} \|s_h - s\|_{W^{1,q}}^q &\leq \frac{c_1}{\Delta t} \int_{\Omega_A(s)} (b - \ell^k)(b_h - b) \\ &\quad + C(s_h) \|u_h - u\|_{W^{1,p}} \\ &\quad + c_0 \|\Pi_h(s) - s\|_{W^{1,q}}^{q'} \end{aligned}$$

*Proof.* Make the Falk (1974) technique for VIs fully nonlinear.

## summary

- implicit time-stepping makes sense for geometry-evolving, Stokes models for glaciers
  - both a differential-algebraic system and a free-boundary problem
- **conjecture.** for some  $q > 2$ , the surface motion  $-\mathbf{u}|_s \cdot \mathbf{n}_s$  from the Glen-Stokes problem is  $q$ -coercive over admissible  $s \in W^{1,q}(\Omega)$
- **theorem.** supposing  $q$ -coercivity, each continuous-space, backward Euler time step problem is well-posed
- **theorem.** supposing well-posedness, the surface elevation FE error is bounded by a sum of 3 terms:
  1. discretizing the bed elevation ( $b_h$  versus  $b$ )
  2. numerically solving the Stokes equations ( $\mathbf{u}_h$  versus  $\mathbf{u}$ )
  3. Cea's lemma for the surface elevation ( $s_h$  versus  $\Pi_h(s)$ )



---

E. Bueler (2024). *Surface elevation errors in finite element Stokes models for glacier evolution*, [arxiv:2408.06470](https://arxiv.org/abs/2408.06470)



E. Bueler (2024). *Surface elevation errors in finite element Stokes models for glacier evolution*, [arxiv:2408.06470](https://arxiv.org/abs/2408.06470)

- 
- E. Bueler & P. Farrell (2024). *A full approximation scheme multilevel method for nonlinear variational inequalities*, SIAM J. Sci. Comput. 46 (4), [doi:10.1137/23M1594200](https://doi.org/10.1137/23M1594200)
  - N. Calvo and others (2003). *On a doubly nonlinear parabolic obstacle problem modelling ice sheet dynamics*, SIAM J. Appl. Math. 63 (2), 683–707 [doi:10.1137/S0036139901385345](https://doi.org/10.1137/S0036139901385345)
  - R. Falk (1974). *Error estimates for the approximation of a class of variational inequalities*, Mathematics of Computation 28 (128), 963–971
  - G. Jovet & J. Rappaz (2011). *Analysis and finite element approximation of a nonlinear stationary Stokes problem arising in glaciology*, Adv. Numer. Analysis 2011 (164581) [doi:10.1155/2011/164581](https://doi.org/10.1155/2011/164581)