

Iterative approach to atmospheric temperature profiles using retrievals from the LAFE field campaign

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Abstract

Temperature is an important parameter when considering the state of the atmosphere and its processes. A method of atmospheric temperature retrieval discussed in Bösenberg [1] involves using differential absorption lidar (DIAL) as a means of finding the temperature sensitive absorption coefficient of a molecule with a known mixing ratio, such as diatomic oxygen. However the absorption coefficient has a non-linear dependence on temperature and therefore cannot be solved analytically; an iterative approach must then be used to complete the temperature retrieval. The accuracy of this retrieval method has been demonstrated using a modeled atmosphere in Bunn et al. [2]. This paper will use this same retrieval method, but applying it to atmospheric data collected from the LAFE (Land Atmosphere Feedback Experiment) field campaign, which took place in August of 2017 at the Southern Great Plains atmospheric observatory in Oklahoma. The data from this experiment includes aerosol and molecular backscatter measurements from an HSRL as well as water vapor profiles from a water vapor DIAL, which are ancillary information necessary for finding the absorption coefficient and temperature profile. We intend to investigate the convergence rate of this method using LAFE data against that of a reasonable modeled atmosphere, such as the one presented in Bunn et al. [2].

I. INTRODUCTION

The DIAL method of temperature profiling relies on the measurement of the temperature dependent absorption coefficient of a molecule with a known mixing ratio (number of molecules of gas divided by total number of molecules of air). The absorption lines of oxygen in the near-infrared around 770 nm are suitably temperature sensitive [1] and the mixing ratio of oxygen is reliably stable, known to be 20.95% [3]. The absorption coefficient is found by measuring the backscatter signal from two closely spaced wavelengths; one associated with the absorption feature (known as the online wavelength) and the other is unaffected by it (known as the offline wavelength). Comparing these two signals and accounting for the Doppler-broadened Rayleigh scattering, Mie scattering off of aerosols, and spectroscopic details of the absorption feature will allow for the retrieval of the absorption coefficient. The returns from an O₂ DIAL that will be used in this paper are the same model used in a manuscript currently being prepared by Bunn et al. [2]

The DIAL equation used to calculate the absorption coefficient is given by [1]

$$\frac{\ln \left(\frac{N_1(r+\Delta R)N_2(r)}{N_1(r)N_2(r+\Delta R)} \right)}{\Delta R} = -\alpha_{u,eff,1}(r) + \alpha_{u,eff,2}(r) - \alpha_{d,eff,1}(r) + \alpha_{d,eff,2}(r) + G_1(r) - G_2(r) \quad (1)$$

where $N(r)$ is the return signal from range r and the subscripts 1 and 2 indicate online and offline wavelengths respectively. The differential range $\Delta R = \frac{c\tau}{2}$ is the bin size where c is the speed of

light and τ is the pulse duration. The effective absorption coefficients for outgoing and return signals (indicated by subscripts u and d respectively) for the online wavelength are given by

$$\alpha_{u,eff,1}(r) = \frac{\int h_1(\nu) \alpha_{m,1}(\nu, r) T_{m,1}(\nu, r) d\nu}{\int h_1(\nu) T_{m,1}(\nu, r) d\nu} \quad (2)$$

$$\alpha_{d,eff,1}(r) = \frac{\int g_1(\nu, r) E(\nu) \alpha_{m,1}(\nu, r) T_{m,1}(\nu, r) d\nu}{\int g_1(\nu, r) E(\nu) T_{m,1}(\nu, r) d\nu} \quad (3)$$

where $h_1(\nu)$ is the online laser lineshape, $\alpha_{m,1}(\nu, r)$ is the molecular absorption coefficient, $T_{m,1}(\nu, r) = \exp[-\int_0^r \alpha_{m,1}(\nu, r') dr']$ is the one-way transmission due to molecular absorption, and $E(\nu)$ is the DIAL receiver spectral filter transmission function. In Equation 3, the term

$$g_1(\nu, r) = \frac{\beta_a(r)}{\beta(r)} h_1(\nu) + \frac{\beta_m(r)}{\beta(r)} (h_1(\nu) * l(\nu, r)) \quad (4)$$

is the lineshape of the backscattered light where $\beta_a(r)$, $\beta_m(r)$, and $\beta(r)$ are the aerosol, molecular, and total backscatter profiles respectively, and $l(\nu, r)$ is the Doppler-broadened lineshape. The effective absorption coefficients for the offline wavelength, $\alpha_{u,eff,2}(r)$ and $\alpha_{d,eff,2}(r)$, have the same form as Equations 2 and 3.

The terms $G_1(r)$ and $G_2(r)$ are correction terms that help account for changes in the spectral distribution due to Doppler broadening of the scattered return signal, given by

$$G_1(r) = \frac{\int \frac{dg_1(\nu, r)}{dr} E(\nu) T_{m,1}(\nu, r) d\nu}{\int g_1(\nu, r) E(\nu) T_{m,1}(\nu, r) d\nu}. \quad (5)$$

The correction term $G_2(r)$ is written in the same form for the offline wavelength.

This paper will rely on a perturbative solution to the DIAL equation (Equation 1) developed in a manuscript that is currently being prepared by Bunn et al. [2], which calculates the absorption coefficient while providing a means of accounting for the Doppler broadening of the elastically scattered light. This solution also relies on the aerosol and molecular backscatter profiles, which will be provided by HSRL measurements [4] taken at the LAFE field campaign.

With the absorption coefficient found from the perturbative method, we now use the following equation to retrieve the temperature profile:

$$\alpha(r) = q_{O_2}(1 - q_{H_2O}) \frac{P}{k_B T} S(T, \epsilon) \Lambda_V(\nu - \nu_0, P, T) \quad (6)$$

where q_{O_2} and q_{H_2O} are the mixing ratios of O_2 and H_2O (the latter of which is highly variable and will be provided by water vapor DIAL measurements [5] taken at the LAFE field campaign), P is the pressure, k_B is the Boltzmann constant, T is the temperature, $S(T, \epsilon)$ is the temperature dependent line strength of the absorption feature, ϵ is the ground state energy, and $\Lambda_V(\nu - \nu_0, P, T)$ is the normalized Voigt lineshape function.

The temperature dependent line strength is given by [6]

$$S(T, \epsilon) = S_0 \frac{T_0}{T} \exp \left[-\frac{\epsilon}{k_B} \left(\frac{1}{T} - \frac{1}{T_0} \right) \right] \quad (7)$$

where S_0 is the line strength at temperature T_0 .

The perturbative solution relies on an initial guess at the temperature profile in order to complete the inversion. If this guess differs from the actual profile, errors in the retrieval of the absorption coefficient will occur, so an iterative approach must be used. The initial guess at the temperature profile will give an initial absorption coefficient, which is then used to retrieve the temperature profile. This new temperature profile is then used to complete the inversion again, providing an updated absorption coefficient and temperature profile. This process is repeated until the input temperature profile matches the retrieved profile. This iterative solution to Equation 6 is given here and has been shown to converge rapidly [6]:

$$T_{i+1} = \frac{\epsilon/k_B}{\ln[Cq_{O_2}(1 - q_{H_2O})\Lambda_V(T_i)] - \ln(\alpha T_i^2/P)} \quad (8)$$

where $C = \frac{T_0}{k_B} S_0 \exp\left(\frac{\epsilon}{k_B T_0}\right)$.

II. LAFE DATA PROCESSING

It is necessary to have an initial guess of the temperature profile. The LAFE data fulfills this requirement because it includes temperature and pressure profiles based off of a surface measurement at the instrument's location and assuming a standard atmosphere with a lapse rate of -6.5 K/km for the entire troposphere. The LAFE data also includes HSRL measurements such as a combined (aerosol plus molecular) backscatter channel signal and a molecular backscatter channel signal, the latter of which was separated out through the use of a rubidium cell as explained in Hayman and Spuler [4]. These signals are then used to find the aerosol and molecular backscatter coefficients necessary for calculating the O_2 absorption coefficient $\alpha(r)$. A water vapor DIAL was also present at the LAFE field campaign, so we have an absolute humidity profile with which we may calculate the water vapor mixing ratio $q_{H_2O}(r)$.

Once the aerosol and molecular backscatter coefficients are found, they are used in the perturbative method of calculating the O_2 absorption coefficient. This, along with the water vapor mixing ratio, gives us everything we need to calculate the updated temperature profile from Equation 8.

I. Aerosol and Molecular Backscatter Coefficients

The aerosol backscatter coefficient is found using [4]

$$\beta_a(r) = \left[\frac{\hat{S}_c(r)}{\hat{S}_m(r)} - 1 \right] \beta_m(r) \quad (9)$$

where $\hat{S}_c(r)$ and $\hat{S}_m(r)$ are the modified combined and molecular channel observations [4], and $\beta_m(r)$ is the estimated molecular backscatter coefficient, calculated from [7]

$$\beta_m(r) = 5.45 \times 10^{-32} \frac{P(r)}{k_B T(r)} \left(\frac{550 \text{ nm}}{\lambda} \right)^4 \quad (10)$$

where $P(r)$ and $T(r)$ are the pressure (Pa) and temperature (K) profiles that were estimated from surface measurements, k_B is the Boltzmann constant, and λ is the HSRL laser wavelength (780 nm).

II. Water Vapor Mixing Ratio

The water vapor DIAL LAPE data is given as absolute humidity $A(r)$ (g/m^3) and we need the mixing ratio $q_{\text{H}_2\text{O}} = m_{\text{wv}}/m_d$ where m_{wv} is the mass of water vapor and m_d is the mass of dry air. The density of dry air may be calculated from the ideal gas law as a function of temperature and pressure

$$\rho_d(r) = \frac{P(r)}{R_d T(r)} \quad (11)$$

where $R_d = 287.058 \frac{\text{J}}{\text{kg K}}$ is the specific gas constant for dry air. The water vapor mixing ratio is then found to be

$$q_{\text{H}_2\text{O}}(r) = \frac{A(r)}{\rho_d(r)}. \quad (12)$$

III. ABSORPTION COEFFICIENT

With the initial temperature and pressure profiles and aerosol and molecular backscatter coefficients found, we model the O_2 DIAL and apply the perturbative method in order to retrieve an initial estimate of the absorption coefficient.

IV. ITERATIVE TEMPERATURE RETRIEVAL

V. COMPARISON TO MODEL ATMOSPHERE

VI. CONCLUSIONS

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