

The role of satellite remote sensing in climate change studies

Jun Yang¹, Peng Gong^{1,2,3*}, Rong Fu⁴, Minghua Zhang⁵, Jingming Chen^{6,7}, Shunlin Liang^{8,9}, Bing Xu^{8,10}, Jiancheng Shi² and Robert Dickinson⁴

Satellite remote sensing has provided major advances in understanding the climate system and its changes, by quantifying processes and spatio-temporal states of the atmosphere, land and oceans. In this Review, we highlight some important discoveries about the climate system that have not been detected by climate models and conventional observations; for example, the spatial pattern of sea-level rise and the cooling effects of increased stratospheric aerosols. New insights are made feasible by the unparalleled global- and fine-scale spatial coverage of satellite observations. Nevertheless, the short duration of observation series and their uncertainties still pose challenges for capturing the robust long-term trends of many climate variables. We point out the need for future work and future systems to make better use of remote sensing in climate change studies.

Observational data and model simulations are the foundations of our understanding of the climate system¹. Satellite remote sensing (SRS) — which acquires information about the Earth's surface, subsurface and atmosphere remotely from sensors on board satellites (including geodetic satellites) — is an important component of climate system observations. Since the first space observation of solar irradiance and cloud reflection was made with radiometers onboard the Vanguard-2 satellite in 1959², SRS has gradually become a leading research method in climate change studies³.

The use of satellites allows the observation of states and processes of the atmosphere, land and ocean at several spatio-temporal scales. For instance, it is one of the most efficient approaches for monitoring land cover and its changes through time over a variety of spatial scales^{4,5}. Satellite data are frequently used with climate models to simulate the dynamics of the climate system and to improve climate projections⁶. Satellite data also contribute significantly to the improvement of meteorological reanalysis products that are widely used for climate change research, for example, the National Center for Environmental Prediction (NCEP) reanalysis⁷. The Global Climate Observing System (GCOS) has listed 26 out of 50 essential climate variables (ECVs) as significantly dependent on satellite observations⁸. Data from SRS is also widely used for developing prevention, mitigation and adaptation measures to cope with the impact of climate change⁹.

Despite the aforementioned contributions of SRS, there are concerns about the suitability of satellite data for monitoring and understanding climate change¹⁰. Climate change studies require observations to be calibrated/validated and consistent, and to provide adequate temporal and spatial sampling over a long period of time¹¹. However, satellite data often contain uncertainties caused by biases in sensors and retrieval algorithms, as well as inconsistencies between continuing satellite missions with the same sensors. The use of satellite observations in climate change studies requires a clear identification of such limitations.

In this Review we discuss the contribution of SRS to our understanding of climate change and the processes involved. We focus on SRS-enabled discoveries that have substantiated or challenged our fundamental knowledge of climate change. Our main goal is to reveal the unique contributions and major limitations of SRS as used in these studies. Technical details on instrumentation and retrieval methods can be found in a recent review¹².

Observations of the climate system

Conventional land-based observations are typically collected at fixed intervals with limited spatial coverage, whereas SRS allows for continual monitoring on the global scale. This has greatly enhanced our understanding of the climate system and its variations (Fig. 1).

Global warming. The warming trend of the Earth's mean surface temperature since the late nineteenth century has provided evidence for anthropogenic influences on global climate¹³. This trend was first identified by analysing anomalies in time series of near-surface air temperature over the land that were recorded by weather stations¹⁴. However, the existence of the trend was consistently challenged due to the biases in weather records¹⁵ caused by such things as poor siting of the instrumentation and the influence of land-use/land-cover changes. Satellite data provides an independent way to investigate global temperature trends, particularly at the ocean surface and in the atmosphere.

The sea surface temperatures (SSTs) of the oceans — which are directly related to heat transfer between the atmosphere and oceans — serve as important indicators of the state of the climate system¹⁶. The Advanced Very High Resolution Radiometer on board the National Oceanic and Atmospheric Administration (NOAA) satellites allows us to monitor the SST worldwide. An increase in SST has been observed in all ocean basins since the 1970s, with an average estimated increase of 0.28 °C from 1984 to 2006¹⁷. This

¹Ministry of Education Key Laboratory for Earth System Modeling, Center for Earth System Science, Tsinghua University, Beijing 100084, China, ²State Key Lab of Remote Sensing Science, jointly sponsored by the Institute of Remote Sensing and Digital Earth, Chinese Academy of Science, and Beijing Normal University, Beijing 100101, China, ³Department of Environmental Science, Policy and Management, University of California, Berkeley, California 94720, USA, ⁴Jackson School of Geosciences, University of Texas, Austin, Texas 78712, USA, ⁵School of Marine and Atmospheric Sciences, State University of New York, Stony Brook, New York 11794, USA, ⁶International Institute for Earth System Science, Nanjing University, Nanjing 210093, China, ⁷Department of Geography and City Planning, University of Toronto, Toronto, Ontario M5S 3G3, Canada, ⁸College of Global Change and Earth System Science, Beijing Normal University, Beijing 100875, China, ⁹Department of Geography, University of Maryland, College Park, Maryland 20742, USA, ¹⁰College of Environmental Science and Engineering, Tsinghua University, Beijing 100084, China. *email: peng Gong@berkeley.edu

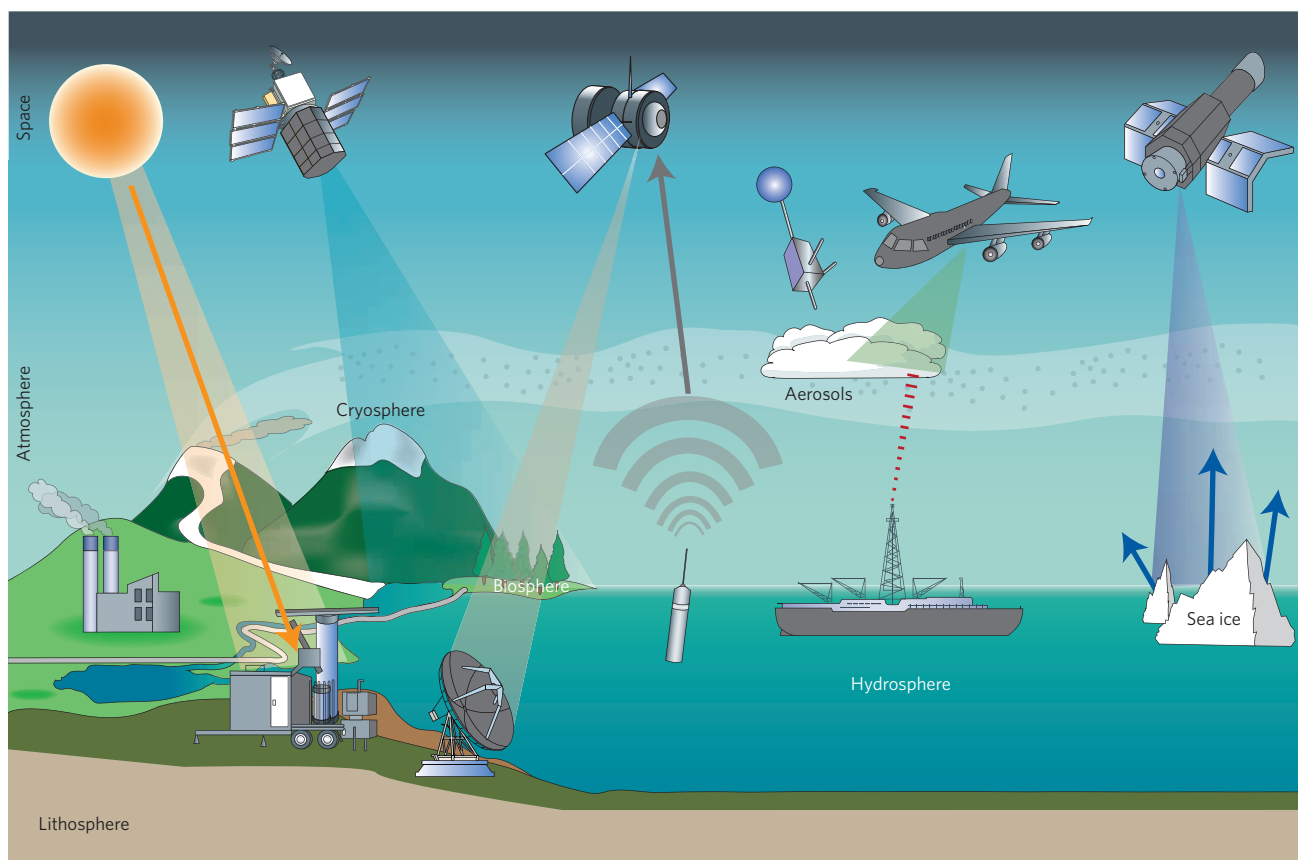


Figure 1 | Remote sensing of the climate system. Remote sensing is carried out by sensors aboard different platforms, including plane, boat and Argo floats. Ground-based instruments are also used, for example, sun spectral radiometers measure solar radiation. However, satellite remote sensing is capable of providing more frequent and repetitive coverage over a large area than other observation means. Figure courtesy of R. He, Hainan University.

global trend is similar to that of the near-surface air temperature over land, thus providing a decisive corroboration of the warming identified from weather records, although the magnitude still has considerable uncertainty¹⁸. Satellite observations also reveal an uneven warming pattern. The ocean warming trend is highest in the middle latitudes of both hemispheres¹⁹, stronger east–west zonal SST gradients have been detected across the equatorial Pacific Ocean²⁰. The regional variability of SST is closely linked with that of other climate parameters, such as precipitation²¹. All of these discoveries have led to a better understanding of the role of the oceans in influencing global climate.

Surface and tropospheric warming has been predicted as a major response of the climate system to escalating concentrations of greenhouse gases²², but initial analyses found no obvious trend in the tropospheric temperature records obtained using microwave sounding units (MSUs) on polar-orbiting satellites²³ — a finding that has challenged both the reliability of surface temperature records and our understanding of the reaction of the climate system to increased greenhouse gas concentrations²⁴. However, after removing known problems with the sensors, accounting for the influence of stratospheric cooling and reducing biases in retrieval methods²⁵, new versions of satellite time series all show a warming trend in the troposphere, except in the Antarctic region²⁶ (Fig. 2). This progress, together with improved climate models, has led to the conclusion that there is no fundamental disagreement between observed and modelled tropospheric temperature trends of warming (provided that uncertainties in both are accounted for)²⁴. Nevertheless, some questions remain. The observed scaling ratio (that is, the ratio of atmospheric trend to surface trend) is still significantly lower than model projections²⁷. The differences between

observed temperatures in the tropical upper- and lower-middle troposphere are also significantly smaller than those simulated by climate models²⁸. These remaining differences indicate that either model deficiencies or observational errors that were identified in previous studies have not been fully resolved.

Snow and ice. The retreat of snow- and ice-cover is an important indicator of global warming. Melting of seasonal snow- and ice-cover can cause a positive feedback by lowering the albedo of the Earth's surface, and the latter contributes to sea-level rise (SLR)¹³. Data from SRS has played a crucial role in monitoring the dynamics of snow extent and ice covers.

Snow-cover extent (SCE) over the Northern Hemisphere has been routinely monitored since 1967 using visible-band sensors and passive microwave-band sensors carried by satellites²⁹. Reconstructed time series based on the NOAA SCE data set and *in situ* measurements have shown that, overall, the SCE over the Northern Hemisphere has been reduced by 0.8 million km² per decade in March and April from 1970–2010. A comparison of recent SCE with pre-1970 values indicates a 7% and 11% decrease in March and April, respectively³⁰. This trend is consistent with observed global surface warming. The satellite observations also display strong regional patterns of SCE that are affected by regional climate variability. No significant decrease was observed over North America in March from 1970–2010³⁰, and the decrease of SCE inside Russia ceased after 1990³¹.

The extent of sea ice is primarily monitored by passive microwave sensors such as the special sensor microwave/imagers (SSM/I). The latest pattern identified from the satellite records shows that the Antarctic sea-ice extent has increased by $1.5 \pm 0.4\%$ per decade from

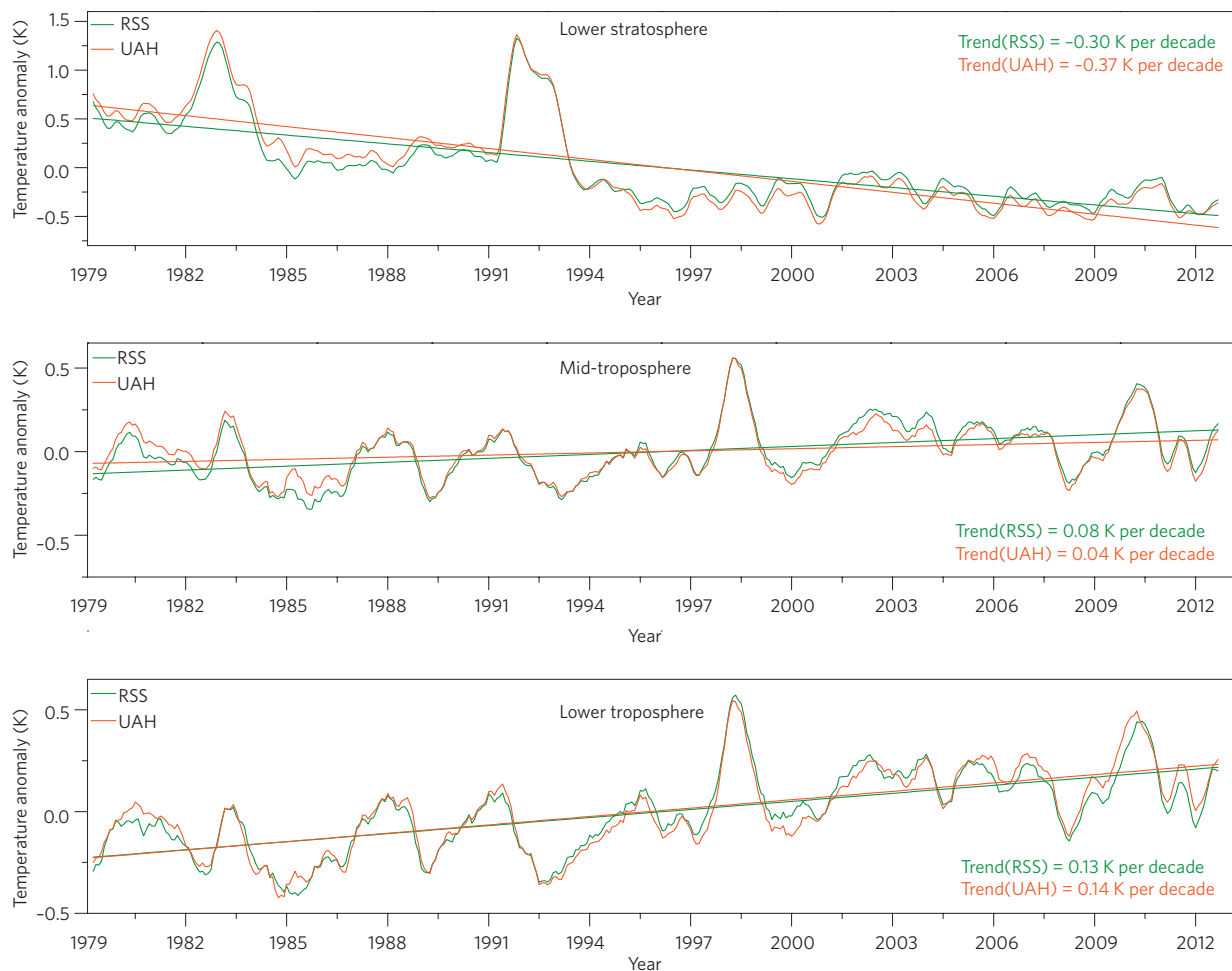


Figure 2 | Upper atmosphere temperature trends between 1979 and 2012 based on MSU data sets. Data from the University of Alabama at Huntsville (UAH) and the Remote Sensing Systems (RSS) are shown here. A 6-month moving window smoothing has been applied to the data sets. Both data sets now show a warming trend, whereas UAH formerly reported a cooling trend in the lower troposphere²³. Data from the National Climate Data Center; <http://vlb.ncdc.noaa.gov/temp-and-precip/msu>. Figure courtesy of C. Huang, Beijing Forestry University.

1979 to 2010³². Researchers hypothesize that this slight increase was caused by reduced upward heat-transport in the oceans and increased snowfall³³. The situation is reversed in the Arctic, where satellite observations show an overall negative trend of $-4.1 \pm 0.3\%$ per decade in that period, with the largest negative trend occurring in summer³⁴. This concurs with climate model projections, but its magnitude is larger³⁵. Whether the Arctic as a whole has reached a 'tipping point' (a point in time when the decrease of ice in the Arctic cannot be reversed) is still a matter of debate, but some researchers believe that satellite observations show that a 'regional tipping point' has been reached, given that the average age of multi-year ice is decreasing, and the percentage of oldest ice declined from 50% of the multi-year ice pack to 10% from 1980 to 2011³⁶.

The mass losses of the Antarctic and Greenland ice-sheets have been estimated by measuring surface elevation changes from satellite altimetry data (collected by satellites such as ERS-1/-2, Envisat, ICESat and CryoSat-2) or by measuring ice-mass changes using the Gravity Recovery and Climate Experiment (GRACE) satellite data. Both estimates show that the Antarctic and Greenland ice-sheets are losing mass^{37,38}. The latest studies find that these polar ice-sheets contributed an average of $0.59 \pm 0.2 \text{ mm yr}^{-1}$ to the rate of global SLR since 1992³⁸. However, estimates of the mass-loss rates are highly divergent (Supplementary Table 1). Both GRACE-based and altimeter-based methods have limitations, and the efforts to combine the strengths of the two have achieved only partial success³⁷. Along with

the mean trend, the acceleration of ice discharge in the Antarctic and Greenland was revealed through the satellite-borne interferometric synthetic-aperture radar records³⁹. This finding and subsequent studies indicate that ice-ocean interaction drives much of the recent increase in mass loss from the Antarctic and Greenland ice-sheets⁴⁰.

The retreat of global mountain glaciers and ice caps (GICs) has been cited as a clear sign of global warming. However, satellite observations show that the extent and magnitude of melt for some glacier systems have been less than predicted. Based on Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) and Satellite Pour l'Observation de la Terre (SPOT) data, the dynamics of glaciers in the Himalayan region were highly variable from 2000–2008. Although 65% of the monsoon-influenced glaciers were retreating, 50% of the glaciers in the Karakoram region of the northwestern Himalayas were advancing or stable⁴¹. The mass loss of global GICs that was assessed using GRACE measurements during 2003–2010 (excluding the Greenland and Antarctic peripheral GICs) was 30% less than the *in situ* mass-balance result. In the high mountains of Asia, the glacier loss-rate in terms of sea-level contribution was estimated to be only $0.01 \pm 0.06 \text{ mm yr}^{-1}$, significantly lower than the $0.13\text{--}0.15 \text{ mm yr}^{-1}$ calculated by other studies⁴². In the Karakoram region, the mass balance of glaciers was found to be positive, at $0.11 \pm 0.22 \text{ m yr}^{-1}$ of water equivalent between 1999 and 2008⁴³. Although some of these evaluations still contain high

uncertainties, they show non-uniform glacier retreat and regional variations that have been verified by *in situ* observations⁴⁴. They also indicate that the estimated contribution of water from melting glaciers to SLR may be less than predicted in certain regions. Evidently, causal factors other than surface temperature must be considered in projecting the future evolution of glaciers.

Sea-level change. Sea level is driven by climate conditions which are influenced by climate change and variability⁴⁵. On the basis of monitoring sites with good-quality tide-gauge records, the global averaged SLR was estimated as $1.9 \pm 0.4 \text{ mm yr}^{-1}$ since 1961⁴⁶. Satellite altimetry observations using the TOPEX/Poseidon satellite launched by NASA and Centre National d'Études Spatiales (CNES), which mapped ocean surface topography from 1992 to 2006, as well as its follow-on and other missions, observed a global mean SLR of around $3.2 \pm 0.8 \text{ mm yr}^{-1}$ between 1992 and 2010⁴⁷. Satellite altimetry observations have also been combined with *in situ* or tide-gauge measurements to reconstruct long-term sea-level time series with global coverage⁴⁶. Strong regional SLR variations have

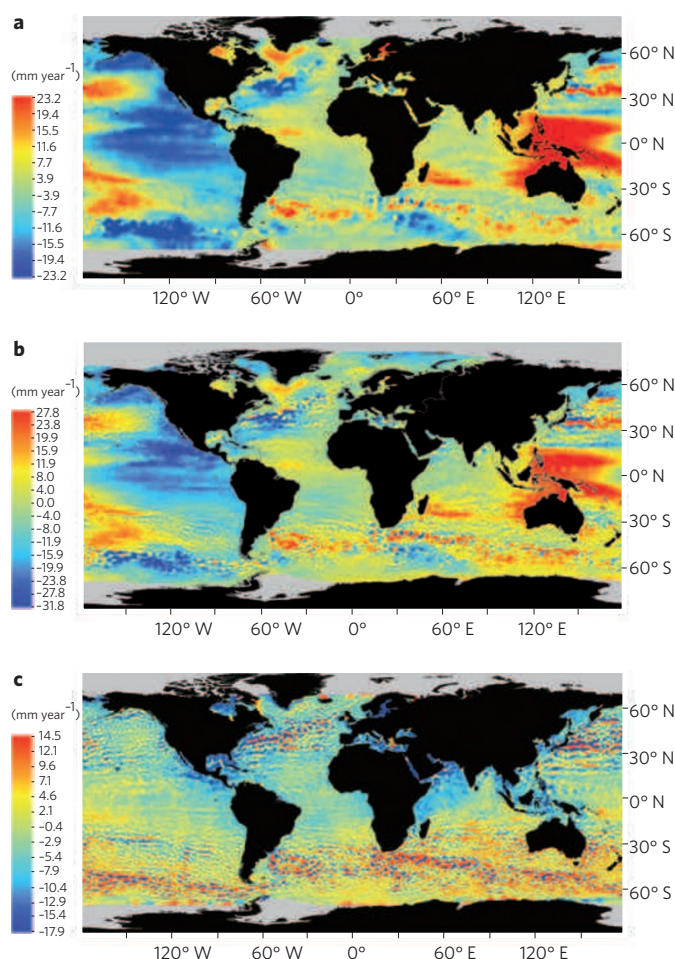


Figure 3 | Map of sea-level trends between 1993 and 2012. **a**, Combined data sets (66°N to 66°S) of TOPEX/Poseidon, Jason-1 and Jason-2 produced by the Laboratory for Satellite Altimetry, NOAA (<http://ibis.grdl.noaa.gov/SAT/SeaLevelRise>). **b**, Sea level trends for 85°N to 85°S from the same three satellites produced by the archiving, validation and interpretation of satellite oceanographic data (AVISO). Data from <http://www.aviso.oceanobs.com/en/news/ocean-indicators/mean-sea-level>. **c**, The difference between the LSA/NOAA and AVISO data. AVISO and LSA/NOAA have close estimates of global mean SLR (without GIA corrections), 2.72 ± 3.01 and $2.71 \pm 2.30 \text{ mm yr}^{-1}$, respectively. However, the estimated magnitudes of regional SLR were different in the two data sets.

been uncovered by satellite observations (Fig. 3). However, these spatial patterns were probably transient features influenced by El Niño/La Niña or longer-period signals⁴⁵. The prevalence of mesoscale eddies (radius around 100 km or smaller) in oceans has also been detected by satellite observations, profoundly changing our understanding of the relationship between sea level and ocean circulations⁴⁸.

The global mean SLR since the last decade of the twentieth century estimated from satellite data has been much higher than that calculated from tide-gauge data during the twentieth century. However, owing to the short data span (~2 decades) of satellite altimetry, the higher estimated rate of SLR could be influenced by interannual or longer oceanic variations, and not necessarily accelerated SLR. Knowing why the rate of SLR has changed is important for its robust future projections^{46,47}. Verification of altimetry observations against tide-gauge data has ruled out the possibility that this change in the trend is spurious, or resulting from biases in satellite data. Tide-gauge calibrations did not show any trend drift in the combined 18-year satellite altimetry data⁴⁷. The global mean SLR from tide-gauge records alone was $2.8 \pm 0.8 \text{ mm yr}^{-1}$ for the period of 1993–2009, which is not significantly different from the values estimated from satellite altimetry⁴⁶. Values of SLR contain steric change (from alterations in total heat content and salinity) and mass change components. Therefore, in studies of sea-level budget, the total SLR measured from satellite altimeters should be equal to the sum of the SLR equivalent of mass changes observed by GRACE and the steric sea-level change measured by the *in situ* network of hydrography data collected, for example, by the Argo system (a global array of 3,000 free-drifting profiling floats that measure the temperature and salinity of the upper 2,000 m of the ocean). Initial results show a general agreement in regions where all three observation systems are in operation. However, the sea-level budget cannot be fully closed with existing data sets. Systematic errors in each observation system, and the error in the glacial isostatic adjustment (GIA) modelling (accounting for post-glacial rebound of the Earth's surface), must be addressed, and, longer observation periods are needed^{47,49}.

Accurately estimating trends in global mean SLR is still a challenging task. A longer period of observation is needed to distinguish the interannual and decadal variability from the long-term trend in sea-level records. Satellite altimetry measurements of mean sea level for regions outside the 66°N–66°S zone are not readily available⁴⁷. A recent effort to use multisatellite altimetry data covering the ice-free ocean to generate weekly gridded sea-level data for regions between 66 and 82°N has offered a partial solution⁵⁰.

Solar radiation. It is important to track changes in the sun's luminosity (measured in total solar irradiance; TSI), to determine whether the natural variation in solar radiation has contributed significantly to recent climate change. Satellite observation of TSI started in 1978 using active-cavity electrical-substitution radiometers. Climate simulations indicate that the variation in TSI has had little influence on global warming since 1880⁵¹. One study determined that the sun might have contributed 25–35% of the 1980–2000 global warming on the basis of two alternative satellite TSI composites⁵². A nearly negligible influence was found in a later study using the same data, and limitations of the retrieval methodology that was previously used were suggested to be the source of the overestimation⁵³.

Recent studies have identified possible ways that variability in solar spectral irradiance influences climate^{54,55}. Measurements from the spectral irradiance monitor (SIM) on board the Solar Radiation and Climate Experiment (SORCE) satellite showed that ultraviolet radiation has decreased over the solar magnetic energy cycle four to six times more than had been expected from model calculations. This reduction is partially compensated by an increase in radiation at visible wavelengths⁵⁴. These spectral changes were linked to a significant decline in stratospheric ozone below an altitude of

45 km from 2004 to 2007, which caused a ripple effect throughout the atmosphere and contributed to global surface warming⁵⁵. The variation in solar ultraviolet irradiance has also been linked to cold winters in northern Europe and Canada by using the SIM measurements as inputs to run climate models⁵⁶. However, others believe that the observed variation is a consequence of undercorrection of instrument response to changes during early on-orbit measurements or undetected drifts in the sensitivity of instruments⁵⁷. The true change in the solar spectrum and its impact can only be confirmed when longer-period SIM measurements become available and more factors are considered in climate model simulations⁵⁸.

Aerosols. Particles in the atmosphere known as aerosols can generate a cooling effect on the climate system, counteracting the warming effects of anthropogenic greenhouse gases by affecting both atmospheric radiation and cloud–precipitation processes^{59,60}. Recent changes in atmospheric aerosol concentration have been identified through aerosol optical depth (AOD), which is derived from observations recorded by visible and infrared optical sensors on board various satellites. Since 1982 these data show a negative trend in the troposphere over North America and most of Europe, and a positive trend over South and East Asia. The overall combined effect of these regional changes probably amounts to a small negative trend^{61,62}, which is attributed to efforts in North America and Europe to control air pollution and rapid industrialization in Asian countries⁶². The findings are important, as aerosols in the troposphere produce direct and indirect radiative forcing that can significantly alter the regional patterns of climate change. The concentrations of aerosols in the stratosphere have increased by as much as 10% since 2000, in contrast to assumptions used in climate models that the background stratospheric aerosol layer is constant. This increase might have caused a negative radiative forcing of about -0.1 W m^{-2} , implying a global cooling of about -0.07°C and consequently about 10% less SLR since 2000⁶³. The increase was attributed to small-scale volcano emissions⁶⁴.

Satellite observations of the direct and indirect climate forcing by aerosols provide independent comparisons for climate models. The direct radiative forcing by anthropogenic aerosols is calculated by combining satellite observations of the radiation budget and AOD. The estimated value is around $-1.0 \pm 0.34 \text{ W m}^{-2}$, greater than the $-0.5 \pm 0.4 \text{ W m}^{-2}$ ‘consensus’ value from the Fourth Assessment Report of the Intergovernmental Panel on Climate Change and most model results. Explanations for this disagreement include biases in satellite data and the deficiency of models in simulating the physical, chemical and optical properties of aerosols⁶⁰. The indirect effects of aerosols that are assessed using satellite data range from -0.2 to -0.5 W m^{-2} , which is three to six times smaller than model estimates⁶⁵. One reason for this difference is that satellite-based methods use the present-day relationship between observed cloud water content (droplet number) and AOD to determine the preindustrial values for droplet concentration but the primary reason is probably overestimation by models, because they still cannot accurately simulate cloud processes⁶⁶.

The interactions between cloud, aerosols and precipitation are also better understood with SRS observations. Cloud–aerosol lidar data show that dust particles lifted to the cold cloud layer promote the formation of ice crystals in supercooled clouds and decrease cloud albedo — which, in turn, leads to decreased reflection of solar radiation and decreased cooling effects⁶⁷. The simultaneous multisensor observation capability provided by the A-train mission (a formation of six Earth-observing satellites) has allowed the detection of the suppression effect of aerosols on precipitation of ice clouds⁶⁸.

Clouds. Climate forcing and feedback of clouds adjust the energy flow throughout the Earth’s system. Net cloud forcing (NCF) is

estimated to be -21 W m^{-2} by combining model simulations with the Earth Radiation Budget Experiment (ERBE) and Clouds and the Earth’s Radiant Energy System (CERES) observations⁶⁹. The cloud feedback to short-term climate variations is calculated to be a positive value of $0.54 \pm 0.74 \text{ W m}^{-2} \text{ K}^{-1}$ on the basis of satellite observations of the top-of-atmosphere radiation budget⁷⁰. These estimates provide good constraints for climate models. The model outputs of short-term (decadal scale) cloud feedback are in the same range as the satellite observations, indicating that climate models can simulate the response of clouds to short-term climate variations⁷⁰. However, cloud feedback is considered to be the most complex and least understood climate phenomenon. For example, measurements made with the Multiangle Imaging SpectroRadiometer (MISR) on board the Terra satellite showed a decreasing global effective cloud height between 2000 and 2010. Such a cloud top lowering could lead to a negative cloud feedback not previously considered⁷¹. However, satellite cloud products still lack the accuracy and duration required to detect trends in cloud properties that are critical for understanding the long-term (centennial-scale) feedback to climate changes⁷². To better understand cloud feedback, long-term stable satellite observations of cloud types and their physical properties, as well as radiative effects, are needed.

Water vapour and precipitation. Water vapour is an important greenhouse gas as it contributes around 50% of the present-day global greenhouse effect⁷³. Models predict that climate warming will increase atmospheric specific humidity (resulting in a positive feedback) and, in turn, strongly amplify the warming⁷⁴. From 1988 to 2003, an increase of $0.4 \pm 0.09 \text{ mm}$ per decade of precipitable water in the troposphere over the ocean has been observed using the SSM/I data⁷⁵. A strong interannual correlation between tropospheric water-vapour content and surface temperature over the ocean–land combination was also detected during the period of 1988–2008 using a combination of SSM/I and radiosonde (instrumentation hosted on weather balloons) data⁷⁶. High-Resolution Infrared Radiation Sounder (HIRS) records from 1979 to 2009 showed an average increase in water-vapour content in the upper troposphere over the equatorial tropics⁷⁷. Although regional, seasonal, interannual and even longer-period variations that respond to decadal climate events such as the El Niño Southern Oscillation (ENSO) have been observed^{77,78}, satellite observations consistently support a positive tropospheric water-vapour feedback on a long timescale⁷⁴. A different trend was detected in observations of stratospheric water-vapour content. Satellite data sets from the HALogen Occultation Experiment (HALOE), Stratospheric Aerosol and Gas Experiment II (SAGEII) and the Microwave Limb Sounder (MLS) indicated that the stratospheric water-vapour content increased persistently from 1980 to 2000, and then decreased by 10% between 2000 and 2009, contributing to the flattening of the global warming trend⁷⁹. Satellite observations also support the theory that the stratospheric water-vapour content is tightly regulated by the tropical tropopause temperature⁸⁰, even though decoupling may occur on a short timescale⁸¹. These results provided powerful validation of the abilities of climate models to simulate the feedback of water vapour.

Precipitation plays a primary role in the global water and energy cycle, and its variations are closely linked to climate change. The spatial and temporal variability of precipitation on the global scale can be retrieved from observations made by infrared sensors on board geostationary satellites, passive microwave sensors carried by the polar-orbiting satellites and recently active radars on board the TRMM satellite and its successors. Satellite records between 1987 and 2006 indicated that the response of precipitation to global warming was $7\% \text{ K}^{-1}$ of surface warming, much higher than the $1\text{--}3\% \text{ K}^{-1}$ predicted by climate models⁸². The observation has generated an intense debate on whether a large discrepancy exists between observed and modelled trends of precipitation. Although studies

Table 1 | Time span of climate ECVs retrieved from satellite observations.

Time span (yr)	Atmospheric ECV	Oceanic ECV	Terrestrial ECV
0–9		Ocean salinity	Biomass, glacier and ice caps
10–19	Wind speed and direction (upper air), CO ₂ and ozone	Ocean color, sea state (including wave height, direction, wavelength and time period)	Land cover, fire disturbance
20–29	Radiation budget, wind speed and direction (surface), water vapour, cloud properties and aerosol properties	Sea level	Albedo, lakes (water levels and areas)
30–39	Precipitation, upper air temperature	SST, sea-ice extent	Fraction of Absorbed Photosynthetically Active Radiation (fAPAR) LAI, soil moisture
40–49			Snow cover

*Data sets, providers and access links for these ECVs are listed in the Supplementary Table 2.

using longer satellite time series did produce smaller rates of increase, these results are still regarded as inconclusive due to the brevity of the series^{83,84}. The correlation between satellite observed precipitation and global surface temperature anomalies at inter-annual scale was also found to be weak during 1988–2008 if large-scale forcing including ENSO and volcanic eruptions were removed⁷⁶. A survey of available satellite-based long-term precipitation products showed mostly little to no trend in global precipitation⁸⁵. These divergent findings illustrate the problems of detecting a robust global mean trend of precipitation, a consequence of high variability of precipitation, systematic biases associated with instruments, and inadequate interpretation of the surface and atmospheric properties in the retrieval algorithms^{84,86}. Although there is still uncertainty regarding a general trend, satellite observations have greatly enhanced our understanding of the climate processes that control the variability of precipitation. For example, the effect of SST on the regional precipitation pattern at multi-annual and multidecadal timescales²¹, and the prevalence of ‘wet-gets-wetter’ and ‘dry-gets-drier’ trends in tropical regions that are caused by the intensification of summer monsoons, were detected using satellite precipitation products⁸⁷.

Limitations and solutions

Three recurring limitations can be identified from these studies: short data spans of satellite records, biases associated with instruments and uncertainties in retrieval algorithms.

Short spans of satellite data sets. Researchers who use short time series of satellite data might have trouble reliably separating the long-term trends from inter-annual and decadal variability. Intuitively, to most accurately detect trends in climate change, satellite observations should have long-term continuity⁸⁸, consistency, and homogeneity (here this means that the non-climatic factors such as instruments and observing methods have homogeneous influences on all observed data). However, there is no uniform requirement for observation length. Some studies have specified minimum requirements for climate variables, ranging from 40 years for tracking the change of ocean colour⁸⁹ to 60 years for determining SLR⁹⁰. The GCOS suggested 30 years for satellite observation of the climate system⁸. More studies are needed to determine the lengths of satellite time series required for detecting reliable trends in other climate variables. An examination of the lengths of ECVs constructed from satellite observations shows that some time series are already longer than 30 years (Table 1). The availability of more time series with adequate length will depend on our ability to maintain the continuity of existing satellite missions.

Biases associated with instruments. Observations of many climate variables are made by satellite sensors that were originally designed for meteorological observations. The coarse-resolution

sensors carried by some satellites cannot capture climate processes occurring at finer spatial scales. For example, satellite sensors are unable to provide the observations at high enough resolutions to characterize the small-scale ‘turbulence’ that is associated with the variability of atmospheric temperature and water vapour⁹¹. In addition, satellite sensors do not have the required accuracy for detecting cloud-property trends⁷². This requires new sensors with sufficiently high spatial resolution and accuracy for observing the climate phenomena of interest. Existing systems provide inadequate temporal coverage with which to study quickly evolving climate processes. This shortfall must be addressed. A possible solution is a constellation of small satellites that observe the same location over a given time interval^{9,72}.

The findings of satellite records that are used for trend detection and retrieval of the absolute levels of climate variables are greatly affected by how well the uncertainties associated with the sensors are resolved. This important step is underscored in the debate on the trend in solar radiation. Undetected drifts in sensor sensitivity have been cited as the main reason for the apparent spectrum of change⁵⁷. Satellite sensors gradually lose radiometric sensitivity and stability during their operation, so good calibration is essential. Some satellite sensors cannot be recalibrated after launch due to the lack of accurate on-board or on-orbit calibrations. Procedures have been developed to calibrate these types of sensors but may still contain uncertainties⁹². Biases caused by instrument drifts are also common in satellite data. Satellites go through a slow change of crossing time at the local equator and a decay of orbital height, adding a spurious effect to detected trends⁹³. Such biases must be addressed by applying a diurnal correction procedure to the data⁹³ or by determining the precise orbit position of the satellites⁹⁴.

Uncertainties can also increase when combining observations from different satellite systems to form long-term records. If the procedures for merging data from different systems are not well developed and calibrated, the uncertainties can potentially be high in combined data sets. The report of an abrupt increase in Antarctic sea-ice area was found to be a false detection caused by the shift from the SSM/I system to the Advanced Microwave Scanning Radiometer for the Earth Observing System (AMSRE)⁹⁵. This type of problem can be reduced by allowing an overlap of operating time between instruments so inter-instrument calibrations can be conducted to find out the relative bias⁹⁶. For instance, launch times of the Topex/Poseidon satellite and its successors allow for tandem measurements and intersatellite calibrations. Many new satellite missions are adopting this approach¹².

Uncertainties in retrieval algorithms. Retrieval algorithms are key for converting the electromagnetic signals from satellite sensors to measurements of climate variables. Uncertainties can therefore

affect the magnitude of detected trends and even change their direction — for example, the positive and negative trends of tropospheric temperature derived from the same satellite data by using different retrieval algorithms^{23,93}. Practices such as producing several satellite data sets of climate variables independently and conducting intercomparison studies should continue and be expanded to help identify potential problems.

Recent studies show that the common inputs used in retrieval algorithms are an important source of uncertainty. One main factor in the divergent estimates of the mass-loss rate of ice sheets discussed earlier is the different GIA values adopted by different research groups³⁸. In studies of the concentrations of aerosols in the atmosphere, surface emissivity and albedo, cloud information and aerosol properties are frequently cited inputs that significantly affect the accuracy of retrieved AOD⁹⁷. Researchers are faced with many choices for these common inputs. For example, seven types of global land-cover data derived from satellite data are available for land surface modelling⁵. Studies are needed to evaluate the quality of the common inputs used in retrieval algorithms and to quantify their effects on the uncertainties in the final products.

Besides making improvements in instrumentation and retrieval algorithms, high-quality validation data will help researchers to calibrate instruments, tune algorithms and gauge the level of uncertainties in satellite data sets^{21,94}. There is an urgent need for more global reference networks for calibrating satellite data and validating data products¹¹. However, the spatial and temporal mismatch between satellite observations and validation data sets must be noted and accounted for in this process⁹⁸. Guided by better knowledge of errors in instruments and algorithms, rigorous reanalysis should be conducted regularly to remove errors in long-term remotely sensed data. This approach provides the best hope for producing high-quality climate records from data collected by existing satellites and their predecessors. For example, a temporally consistent global product of the Leaf Area Index has been developed using a well calibrated algorithm⁹⁹.

Conclusions

In this Review we have demonstrated that SRS has made crucial contributions to our understanding of the climate system and its variations. It provides an independent source of observations to validate climate models and climate theories. Although the shortness of satellite time series and associated uncertainties have so far limited the detection of robust long-term trends in some climate variables, the progress made in both instrumentation and retrieval algorithms, accompanied by the accumulation of satellite records, can remedy this. By combining the strength of passive and active remote sensing — for example, the formation of A-train — better insights into the complicated climate system have been obtained. Innovative use of existing satellite data is also a path to production of long-term climate records. A good example is the NOAA's use of archived geostationary satellite data (which are not used as frequently as polar-orbiting satellite data in producing climate records¹⁰⁰) to produce 30-year brightness temperature data — that is, the effective temperature of a black body radiating the same amount of energy per unit area at the same wavelengths as the observed body.

Along with advancing the science and technology of SRS, strong international collaboration and governmental support are needed to enhance its role in climate change studies. Several international initiatives — notably the Global Observing Systems Information Center (GOSIC), the Global Geodetic Observing System (GGOS) and the Global Earth Observation System of Systems (GEOSS) — have been implemented to coordinate efforts to produce and disseminate high-quality satellite climate records. The USA, European Union and several other countries have planned new satellite missions for climate observation: in total, 17 satellite missions that can

provide improved climatological measurements are scheduled for launch by 2020¹². However, the recent cancellation of the Climate Absolute Radiance and Refractivity Observatory (CLARREO) and the Deformation, Ecosystem Structure and Dynamics of Ice (DESDynI) missions show that these initiatives depend on a high level of government commitment. Clearly, our 'eyes in space' would contribute more towards capturing real trends of climate change if the aforementioned activities are fully implemented.

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Author contributions

J.Y. and P.G. designed the framework of the Review. All authors contributed to writing.

Additional information

Supplementary information is available in the online version of the paper. Reprints and permissions information is available online at www.nature.com/reprints. Correspondence should be addressed to P.G.

Competing financial interests

The authors declare no competing financial interests.

CORRIGENDUM

The role of satellite remote sensing in climate change studies

Jun Yang, Peng Gong, Rong Fu, Minghua Zhang, Jingming Chen, Shunlin Liang, Bing Xu, Jiancheng Shi and Robert Dickinson

Nature Clim. Change **3**, 875–883 (2013); published online 15 September 2013; corrected after print 27 September 2013.

In the version of this Review Article originally published, on page 877, the mass balance of glaciers in the Karakoram region should have been $0.11 \pm 0.2 \text{ m yr}^{-1}$. On page 879, in the second paragraph of the 'Aerosols' section, the estimated value for the direct radiative forcing should have been $-1.0 \pm 0.34 \text{ W m}^{-2}$. These errors have now been corrected in the HTML and PDF versions of the Review Article.

The role of satellite remote sensing in climate change studies

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In the version of this Review Article originally published, the temperature anomaly trends for RSS and UAH in Fig. 2b,c should have been positive values. These errors have now been corrected in the online versions of the Review Article. In the previous corrigendum, the mass balance of glaciers in the Karakoram region should have read $0.11 \pm 0.22 \text{ m yr}^{-1}$.