

MAMMOTH  
C. Thieulot  
Utrecht Universiteit

# 1 Install & run procedure

To obtain the code, type the following in the terminal:

```
git clone https://github.com/cedrict/mammoth.git
```

This creates a *mammoth* folder. Do

```
cd mammoth
```

To compile the code, simply open a terminal and make sure you have the gfortran compiler installed. Then simply do

```
make
```

Unless there is an error, this should generate the executable file *mammoth*. To run the executable, simply do

```
./mammoth
```

If you wish to redirect the standard output to a file (e.g. *opla*), then:

```
./mammoth > opla
```

If you wish to clean all the data files after a run, use the *cleandata* script as follows:

```
./cleandata
```

To clean the code off data and other compilations files:

```
make clean
```

When the code runs it produces ascii data (*\*.dat* files) and paraview files (*\*.vtu* files) in the *OUT* folder. The former can be visualised with gnuplot, while the latter should be opened with paraview (or VisIt or MayaVi).

## 2 Physical equations

The code solves the heat transport equation in the absence of advective processes:

$$\rho c_P \frac{\partial T}{\partial t} = \nabla \cdot k \nabla T + H$$

where  $\rho$  is the mass density (kg/m<sup>3</sup>),  $c_P$  is the heat capacity (J/kg·K),  $k$  is the heat conductivity (W/m·K), and  $H$  is a heat production term.  $T(x, y, z, t)$  is the temperature field. The domain is designated by  $\Omega$  and its boundary by  $\Gamma$ . This PDE must be supplied with boundary conditions: at a given point of the boundary  $\Gamma$  either a temperature can be prescribed or a heat flux.

## 3 Discretisation

The heat transport equation is solved by means of the Finite Element Method, following the methodology presented in [?]. The FE equation corresponding to this PDE is

$$\mathbf{M} \cdot \frac{\partial \mathbf{T}}{\partial t} + \mathbf{K}_d \cdot \mathbf{T} = \mathbf{F}$$

where

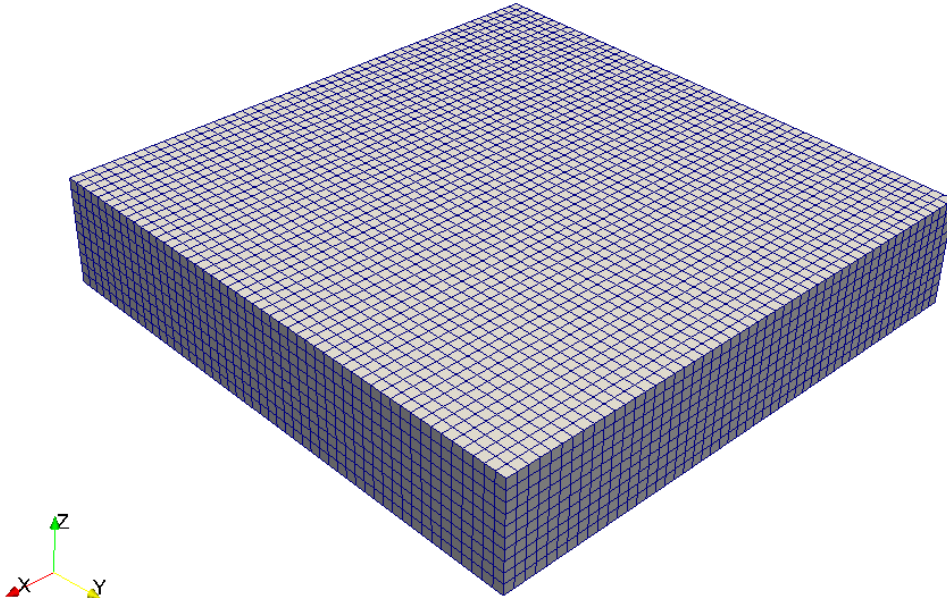
$$\mathbf{M} = \int_{\Omega} \mathbf{N}^T \rho c_P \mathbf{N} dV$$

$$\mathbf{K}_d = \int_{\Omega} \mathbf{B}^T k \mathbf{B} dV$$

$$\mathbf{F} = \int_{\Omega} \mathbf{N}^T H dV$$

where  $\mathbf{T}$  is the vector of nodal temperatures.

The domain  $\Omega$  of size  $L_x \times L_y \times L_z$  is divided into  $nel = nelx \times nely \times nelz$  elements as shown hereunder. The computational grid therefore counts  $np = (nelx + 1)(nely + 1)(nelz + 1)$  nodes and as many degrees of freedom. A Cartesian coordinate system is used with its origin at the bottom corner of the domain.



## 4 Setting up an experiment

### 4.1 Initial temperature field

### 4.2 Magma chambers

The code can model as many magma chambers as needed. This number is controlled by the `nchamber` parameter. Each chamber is parametrised by:

- its **shape**: 1 is a spheroid, 2 is a cuboid and 3 is a cylinder;
- its center coordinates **xc**, **yc**, **zc**;
- its spatial extent in the  $x$ ,  $y$ ,  $z$  directions: **a**, **b** and **c**;
- its initial temperature **T**.

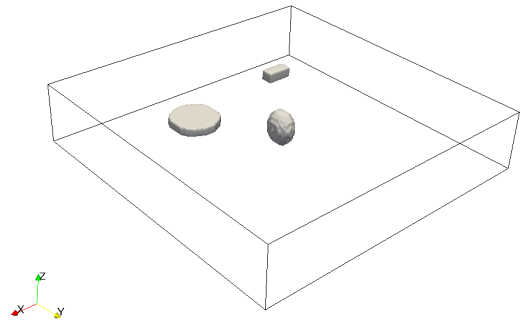
**Example** The following figure is obtained by setting three magma chambers as follows:

```
mc(1)%shape=1
mc(1)%xc=0.5*Lx
mc(1)%yc=0.5*Ly
mc(1)%zc=0.5*Lz
mc(1)%a=4d3
mc(1)%b=5d3
mc(1)%c=6d3
mc(1)%T=500
```

```
mc(2)%shape=2
mc(2)%xc=0.25*Lx
mc(2)%yc=0.2*Ly
mc(2)%zc=0.4*Lz
mc(2)%a=10d3
mc(2)%b=5d3
mc(2)%c=4d3
mc(2)%T=500
```

```
mc(3)%shape=3
mc(3)%xc=0.7*Lx
mc(3)%yc=0.33*Ly
```

```
mc(3)%zc=0.6*Lz
mc(3)%a=8d3
mc(3)%b=8d3
mc(3)%c=3d3
mc(3)%T=500
```



### 4.3 Boundary conditions

## A Steady state solutions

### A.1 One layer

At steady state:  $\partial T/\partial t = 0$  so that one has to solve

$$\Delta T = -\frac{H}{k}$$

Assuming the domain to be infinite in the  $x$  and  $y$  direction this equation becomes:

$$\frac{d^2 T}{dz^2} = -\frac{H}{k}$$

The generic solution to this ODE is given by:

$$T(z) = -\frac{H}{2k}z^2 + Az + B \quad (1)$$

where  $A$  and  $B$  are integration constants to be determined by means of the specified boundary conditions.

At the surface a constant temperature is maintained, i.e.  $T(z = L) = T_s$ , so that

$$T(L) = -\frac{H}{2k}L^2 + AL + B = T_s$$

which allows us to express  $B$  as a function of the rest and insert back into Eq. (1):

$$T(z) = -\frac{H}{2k}(z^2 - L^2) + A(z - L) + T_s$$

**Temperature prescribed at the bottom** In this case  $T(z = 0) = T_b$  so that

$$T(z) = -\frac{H}{2k}L^2 + AL + T_s = T_b$$

which yields

$$A = \frac{T_b - T_s}{L} + \frac{HL}{2k}$$

**Heat flux prescribed at the bottom** In this case  $k dT/dz(z = 0) = \phi_b$  which yields  $kA = \phi_b$  so that

$$T(z) = -\frac{H}{2k}(z^2 - L^2) + \frac{\phi_b}{k}(z - L) + T_s$$

### A.2 Two layers

In each layer  $i$  the solution is given by:

$$T(z) = -\frac{H_i}{2k_i}z^2 + A_i z + B_i \quad (2)$$

Aside from the two previous boundary conditions we need to enforce that the temperature and the heat flux are continuous fields, i.e.  $T(z = l^-) = T(z = l^+)$  and  $\phi(z = l^-) = \phi(z = l^+)$ .

**B** todo