

Seasonality in governing dynamics at submesoscales

Cesar et al.¹

Key Points.

- Upper-ocean submesoscale (10-100 km) turbulence and inertia-gravity waves undergo strong seasonal cycles that are out-of-phase.
- Submesoscale turbulence dominates the horizontal velocity and sea-surface height variability in late Winter/early Spring.
- Inertia-gravity waves dominate the the horizontal velocity and sea-surface height variability in late Summer/early Fall.

Two new high-resolution numerical simulations with embedded tides suggest a strong modulation in the governing near-surface dynamics at submesoscales (roughly 10-100 km). Consistent with recent studies, deep late Winter mixed layers are prone to baroclinic instabilities and submesoscale turbulence prevails in late Winter/early Spring. Inertia-gravity waves, on the other hand, are enhanced near the surface when the upper ocean is strongly stratified. As a consequence of reduction of submesoscale turbulence and enhancement of inertia-gravity waves, the latter strongly dominate the submesoscale surface kinetic energy and sea-surface height variance in late Summer/early Fall.

1. Introduction

Recent interest in upper-ocean dynamics has focused on the strong seasonal cycle of shallow baroclinic instabilities and their role in submesoscale roughly 1-100 km) turbulence and mesoscale modulation [Sasaki et al., 2014; Qiu et al., 2014; Callies et al., 2015; Thompson et al., 2016; Buckingham et al., 2016]. Contemporary studies have also suggested that inertia-gravity waves contribute significantly to the near-surface variability at submesoscales [Richman et al., 2012; Bühler et al., 2014; Rocha et al., 2016], but their seasonality has not been investigated.

Using the output of two new $1/24^\circ$ and $1/48^\circ$ global numerical simulations with embedded tides, we show that inertia-gravity waves undergo a strong seasonality near the surface in the Kuroshio Extension region. The seasonal cycle of inertia-gravity waves is out-of-phase with the seasonal cycle of submesoscale turbulence. Consistent with previous studies, deep late Winter mixed layers are prone to shallow baroclinic instabilities that are roughly in geostrophic balance and flux energy upscale [Sasaki et al., 2014; Callies et al., 2016], driving a mild seasonal modulation of the mesoscales [Sasaki et al., 2014; Qiu et al., 2014].

Inertia-gravity waves, on the other hand, peak in late Summer/early Fall, when the upper ocean is strongly stratified. Thus there exists a strong seasonal modulation in the dominant upper-ocean submesoscale dynamics: submesoscale turbulence dominates the upper-ocean dynamics

in late Winter/early Spring, whereas quasi-linear inertia-gravity waves prevail in late Summer/early Fall. Because submesoscale turbulence is severely diminished in late Summer/early Fall, the present results suggest that inertia-gravity waves account for most of the summertime submesoscale SSH variability.

2. The LLC numerical simulations

We use the output of two latitude-longitude polar cap (LLC) MITgcm realistic numerical simulations. The outputs analyzed here, LLC2160 and LLC4320, are the $1/24^\circ$ and $1/48^\circ$ — 90-vertical-levels simulations from a hierarchy of global forward numerical solutions on a LLC grid [Forget et al., 2015] that were spun up from a ECCO2 adjoint-method state estimate. Both simulations were forced by tides and 6-hourly surface fluxes. The LLC2160 output spans two years from March 2011 to April 2013; the LLC4320 was spun up from the LLC2160 simulation, spanning one year from September 2011 to September 2012. The LLC4320 simulation is an extension of the then 5-month long output used by Rocha et al. [2016].

A key aspect of the LLC simulations is that they were forced by the 16 most significant tidal components. Because barotropic tides interact with topography and generate internal tides that project on mesoscales to submesoscales [e.g., Rocha et al., 2016], tidal forcing fundamentally distinguishes our analysis from recent modeling studies that investigated upper-ocean dynamics and its seasonality [Sasaki et al., 2014; Qiu et al., 2014]. Details of the LLC simulations are provided in the supplemental material.

To study seasonal variations in the upper-ocean dynamics, we focus on the northwest Pacific, in the vicinities of the Kuroshio Extension, where previous studies have suggested strong submesoscale seasonality. We analyze a sub-domain of the LLC4320 and LLC2160 simulations of about 2000 km^2 spanning $155\text{--}175^\circ\text{E}$; $25\text{--}40^\circ\text{N}$ (figure 1a). The stratification in this mesoscale-rich subtropical region undergoes a strong seasonal cycle: wintertime enhance small-scale turbulence de-stratify the upper ocean, yielding mixed layers as deep as 300m. In late Spring/early Summer enhanced solar radiation and mixed-layer instabilities re-stratify the upper ocean, yielding mixed layers as shallow as 40m. The stratification and its seasonality is well-captured by both LLC simulations: the skillful comparisons of model stratification to Argo climatology lends confidence to the present analysis (supplemental material).

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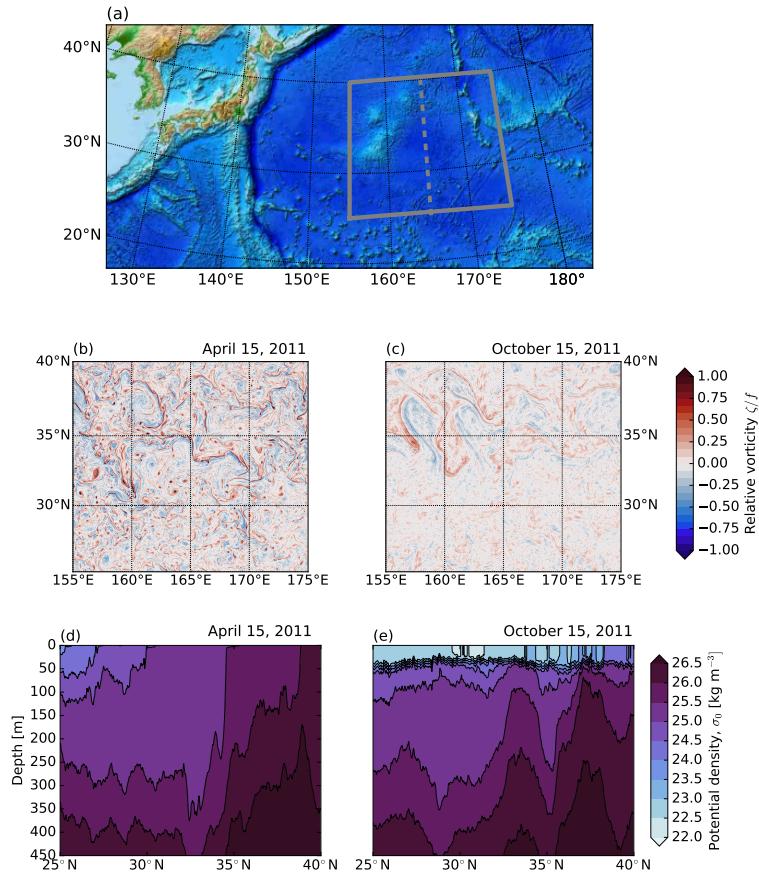


Figure 1. (a) The study region with the subregion where the LLC outputs are analyzed. LLC 4320 ($1/48^\circ$) snapshots of surface vorticity (b and c) and a transect of potential density at 165°E (d and e). The snapshots were taken at 00:00 UTC.

3. Statistics of surface lateral velocity tensor

To study the seasonality in the surface velocity, we calculate the lateral velocity tensor

$$\nabla_h \mathbf{u}_h = \begin{bmatrix} u_x & u_y \\ v_x & v_y \end{bmatrix}. \quad (1)$$

The components of the lateral velocity tensor are calculated using a centered second-order finite differences scheme. We then calculate the vertical vorticity $\zeta \equiv v_x - u_y$, lateral rate of strain $\alpha \equiv [(u_x - v_y)^2 + (u_y + v_x)^2]^{1/2}$, and horizontal divergence $\delta \equiv u_x + v_y$. These diagnostics highlight the submesoscale structures in the flow and their analysis hint on the governing dynamics in the upper ocean [e.g., Capet et al., 2008; Shcherbina et al., 2013].

Figures 1b-c show snapshots of vertical vorticity ζ in early Spring (April 15) and Fall (October 15) in the LLC4320 ($1/48^\circ$) simulation. The model solutions depict seasonality in vorticity: large values of fine-grained vertical vorticity are observed in early Spring with maximum values as large as $4f$, where f is the local planetary vorticity, and root-mean-square (RMS) of about 0.4. In early Fall, the situation is the opposite: the vertical vorticity is relatively coarse-grained, and its local maximum and RMS are no larger than $0.5f$ and $0.2f$, respectively. Indeed, the seasonal cycle of vorticity and rate of strain is very strong (figure 2): in both simulations, the RMS vorticity and strain are about twice as large in late Winter/early Spring than in late Summer/early Fall. Because the wintertime vorticity and strain rate are dominated by the smallest scales in the flow, increasing the resolution from $1/24^\circ$ to $1/48^\circ$ significantly increases the wintertime RMS vorticity and strain by about 40%.

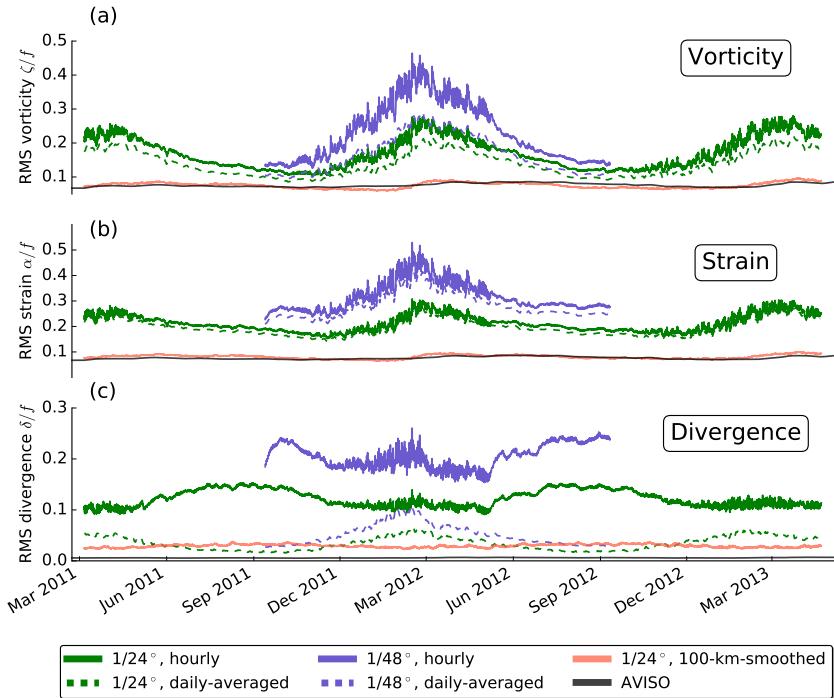


Figure 2. Time series of the root-mean-square (RMS) of vorticity (a) and lateral rate of strain (b), and horizontal divergence (c) in the LLC outputs and gridded AVISO data.

The bulk of vertical vorticity and lateral strain rate is associated with subinertial flows ($T_f = 2\pi/f_0 \approx 23.5$ h, where f_0 is the inertial frequency at the mean latitude): daily-averaging the velocity fields reduces the RMS vorticity by less than 40% and the RMS strain by less than 10%; the seasonal cycle remains strong (see dashed lines in figures 2a-b). Most of this seasonal cycle is associated with submesoscale flows: smoothing the velocity fields with a Hanning filter with cut-off scale of 100 km dramatically reduces the RMS vorticity and strain. The reduction in variance is about 80% in winter, yielding RMS vorticity and strain roughly consistent the ones diagnosed from AVISO gridded geostrophic velocities (compare blue lines with black lines in figures 2a-b). The overall picture that emerges is consistent with recent studies: shallow baroclinic instabilities energize the submesoscales in late Winter, drawing from the available potential energy stored in deep mixed layers [Sasaki *et al.*, 2014; Callies *et al.*, 2015, 2016].

The seasonal cycle of the horizontal divergence, however, showcases the complexity of the upper-ocean annual variability. If submesoscale eddies and fronts domi-

nated the near-surface variability all year, then the seasonal cycle of horizontal divergence, vertical vorticity, and lateral strain rate would be in phase [e.g., Sasaki *et al.*, 2014]. While there is a clear wintertime peak in divergence of daily-averaged velocity (see dashed lines in figure 2a; RMS divergence ~ 0.22 f in the $1/48^\circ$ simulation), the hourly fields show a stronger enhancement of lateral divergence in late Summer/early Fall (RMS divergence ~ 0.22 f the $1/48^\circ$ simulation). Because the $1/48^\circ$ simulation better resolve smaller-scale submesoscale flows, the winter peak in divergence is nearly as strong the summertime's. Of course, part of these submesoscale fronts and eddies evolve relatively fast and there is no clear temporal and spatial scale separation between those motions and inertia-gravity waves: daily-averaging the velocity fields efficiently suppresses the summertime horizontally divergent flows, but also reduces the wintertime lateral divergence by about 50%. As for the vorticity and strain rate, most of the lateral divergence is associated with submesoscale flows: smoothing the velocity fields with a 100km-cut-off suppresses more than 80% of the RMS divergence.

4. Seasonality in governing submesoscale dynamics

The overall results of figure 2c strongly suggest that the summertime submesoscale surface variability stems from different dynamics than then wintertime variability. To better understand these differences, we calculate vorticity-divergence and vorticity-strain joint probability distributions (jPDF).

The April vorticity-strain jPDF have a shape characteristic of submesoscale turbulence (see figures 3f-g). The alignment of vorticity and strain $\alpha \sim \zeta$ with strong positive skewness are fingerprints of submesoscale fronts [Shcherbina et al., 2013; McWilliams, 2016]. The shape of the vorticity-strain jPDF is similar for hourly, daily-averaged, and residual velocities, although the vorticity skewness reduces 1.4 to 1.13 from hourly to daily-averaged. The April vorticity-divergence jPDFs, however, have a egg-like shape that which again highlight the strong skewness in vorticity. Thus wintertime submesoscale surface velocity is strongly dominated by submesoscale turbulence — this is true even for superinertial submesoscale currents.

The October vorticity-strain jPDF show much weaker skewness (the vorticity skewness is 0.68 and 0.67 for hourly and daily-averaged velocities). The shape of the vorticity-strain jPDF appears to be a combination of two half-ellipses centered about $\zeta = 0$, one with a 45° slope (characteristic of submesoscale fronts that persist in Summer) and one with a very steep slope. The vorticity-divergence jPDF is significantly different than the April one. In particular, the jPDF for daily-averaged currents is still egg-shaped but hourly and residual are ellipses about $\zeta = 0$, likely associated with inertia-gravity waves.

Time series of vorticity and divergence PDFs (supplemental material) suggest a strong oscillation between these two regimes. In late Winter/early Spring the vorticity is strongly positively skewed, whereas the divergence is moderately negatively skewed as predicted by frontogenesis [Capet et al., 2008; McWilliams, 2016]. In late Summer/early Fall, the divergence is more stronger, but PDFs are much less skewed, consistent with linear inertia-gravity waves.

5. Projection onto horizontal scales

To better understand the projection of these flows onto different horizontal scales, we calculate wavenumber spectra of kinetic energy and sea-surface height (SSH) variance. Before calculating the spectra, time-mean and spatial linear trend were removed, and the ensuing fields were multiplied by a Hanning spectral “window”. We discuss only spectra for the $1/48^\circ$,

which simply extends the $1/24^\circ$ towards smaller scales (supplemental material).

Figure 3a depicts the horizontal wavenumber spectra of surface KE in April and October. At scales larger than 20km, the spectra based on hourly velocity snapshots are nearly indistinguishable within 95% confidence level (compare solid lines in figure 3a). Consistent with the results of Rocha et al. [2016] who analyzed a 5-month output of the LLC4320 simulation in Drake Passage, there is significant high-frequency variability at submesoscales. Daily-averaging the velocity field suppresses spatial variability at scales smaller than about 250 km both in April and October (compare solid lines against dashed lines in figure 3a). But this suppression is more dramatic in October, when the horizontally divergent flows peak. At scales smaller than 100 km, 39% of the surface KE in April is accounted for by super-inertial flows as opposed to 79% in October. The seasonality of subinertial submesoscale flows is much more dramatic than the seasonal cycle of the total flow. This is consistent with the results of Sasaki et al. [2014], which are based on daily-averaged velocity fields of a different model without tidal forcing (P. Klein, personal communication).

These high-frequency submesoscale flows significantly project on the sea-surface. In October, when the inertia-gravity waves are stronger near the surface and submesoscale turbulence is weak, there is a dramatic difference between the spectra based on hourly and daily-averaged SSH: at scales smaller than 100 km, the spectra of hourly SSH roughly follows a -2 power-law, whereas the spectra of daily-averaged SSH roughly follows a -5 power-law. At scales smaller than 100 km, 33% of the SSH variance in April is accounted for by suprainertial flows as opposed to 83% in October! The out-of-phase seasonal cycle of submesoscale turbulence and near-surface inertia-gravity waves conspire to yield very weak seasonality in the spectra of KE and SSH variance based on hourly field.

6. Summary and Conclusion

Our main finding is that near-surface upper-ocean inertia-gravity waves, undergo a strong seasonal cycle that is out-of-phase with the seasonal cycle of the submesoscale turbulence. D'Asaro [1978] showed that linear internal waves velocity in the mixed-layer strongly depends on the density jump at the mixed-layer base. In Summer, the shallow mixed-layer overlays a strong seasonal pycnocline, and thus the internal waves projection onto the mixed-layer may be stronger according to D'Asaro [1978]'s arguments. An alternative explanation is that the shape baroclinic modes experience significant seasonal modulation, with significant summertime peak in their surface amplitude. If the internal wave source has weak seasonal dependence, then a

strong seasonal cycle in the near-surface expression of internal waves is expected. Internal tide generation is unlikely to vary significantly seasonally [e.g., *Alford*, 2003], and therefore it is plausible to expect a strong seasonality in the the near-surface expression of internal tides and other small-scale internal waves generated through internal tide interactions. Near-inertial wave generation peaks in Winter, but those waves project on large horizontal scales [e.g., *Qi et al.*, 1995].

Our work adds to recent studies that presented modeling [*Sasaki et al.* [2014] and observational [*Callies et al.*, 2015; *Buckingham et al.*, 2016] evidence of vigorous seasonality in submesoscale turbulence. We suggest

that the dominance at inertia-gravity waves in Summer is a consequence of both suppression of submesoscale turbulence and enhancement of inertia-gravity waves due to restratification of the upper ocean. Of course, our study focuses on a single patch of ocean in the vicinities of the Kuroshio Extension, which may be typical of mesoscale-rich subtropical regions, but not representative of other regions such as eastern boundary currents and the middle of the subtropical gyre: care must be taken in overgeneralizing these results. We plan to report on the geographic variability of submesoscale seasonality in a future study.

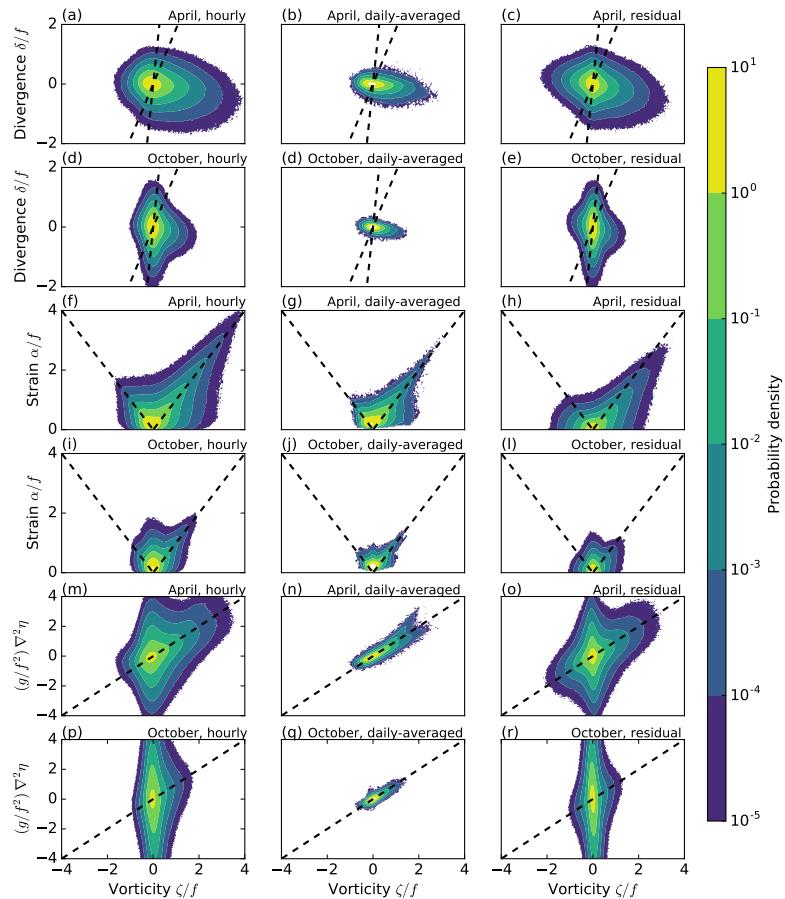


Figure 3. Seasonal variation of joint probability distributions: vorticity vs. divergence (a through e), vorticity vs. strain rate (f through l), and vorticity vs. laplacian of sea-surface height (m through r). Dashed lines in (a) through (e) represent $\delta = \frac{\omega}{f_0} \zeta$ where $\omega = 2\pi/(12 \text{ h}, 3 \text{ h})$ and f_0 is the inertial frequency in the middle of the domain (32.5° N). Dashed lines in (f) through (l) represent pure strain flow $\alpha = \zeta$. Dashed lines in (m) through (o) represent geostrophic flow $\zeta = \frac{g}{f} \nabla_h^2 \eta$.

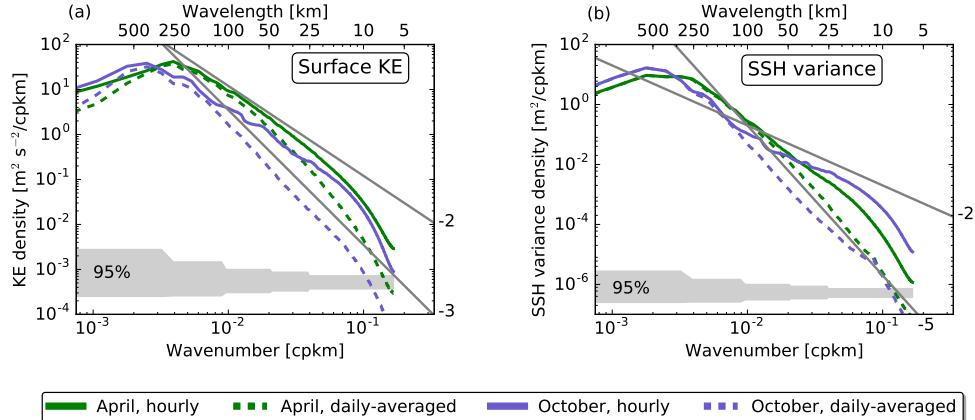


Figure 4. Surface (horizontal) KE (a) and SSH variance wavenumber spectra (b) in the 1/48° simulation. Solid lines are spectra based on hourly snapshots, dashed lines are spectra based on daily-averaged fields.

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Appendix A: Supplemental material 1: Spectral errors

Appendix B: Supplemental material 2: On the LLC simulations

The time-step was 45s for the LLC2160 simulation and 25s for the LLC4320 simulation. Details of the LLC simulations are provided in the supplemental material.

LLC hierarchy setup and control files are available online (at http://mitgcm.org/viewvc/MITgcm/MITgcm_contrib/llc_hires). For the LLC hierarchy, the sea-surface height is calculated using a linearized equation. While the nominal horizontal resolution of the LLC2160 simulation is $1/24^\circ$, horizontal wavenumber spectra suggests an effective resolution of about 20 km. (see figure 3).

B1. Seasonality in the upper-ocean stratification

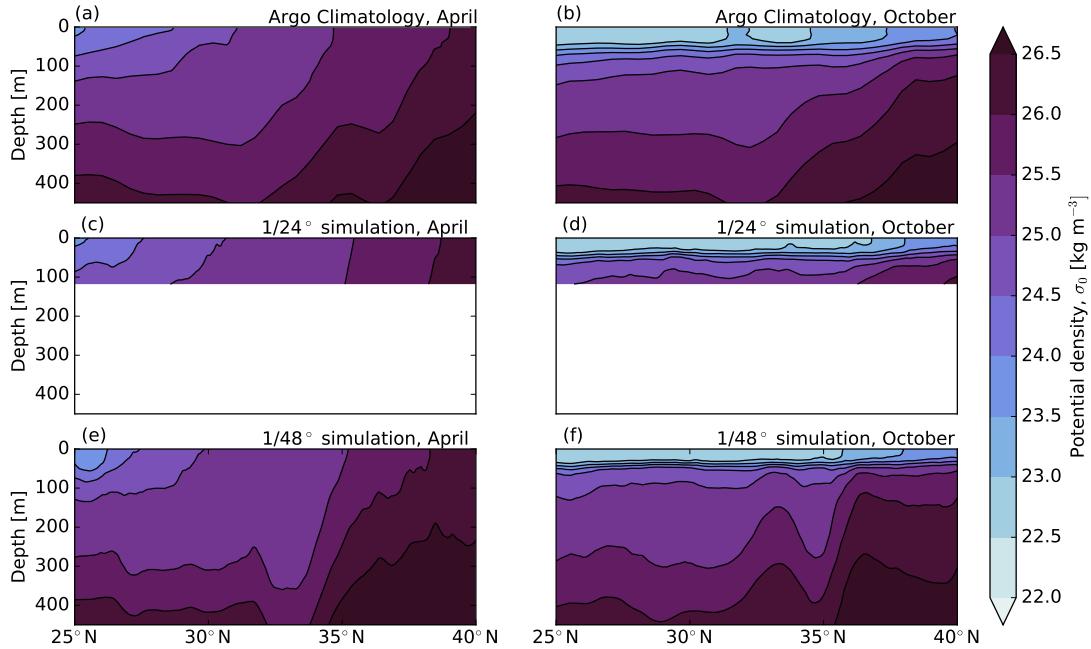


Figure 5. Potential density: a comparison between Roemmich-Gilson Argo climatology (a, b) and the LLC simulations $1/24^\circ$ (c, d) and $1/48^\circ$ (e, f). Need to re-plot the $1/24^\circ$ figures once extraction (currently down) becomes available.

2.2. Spectra

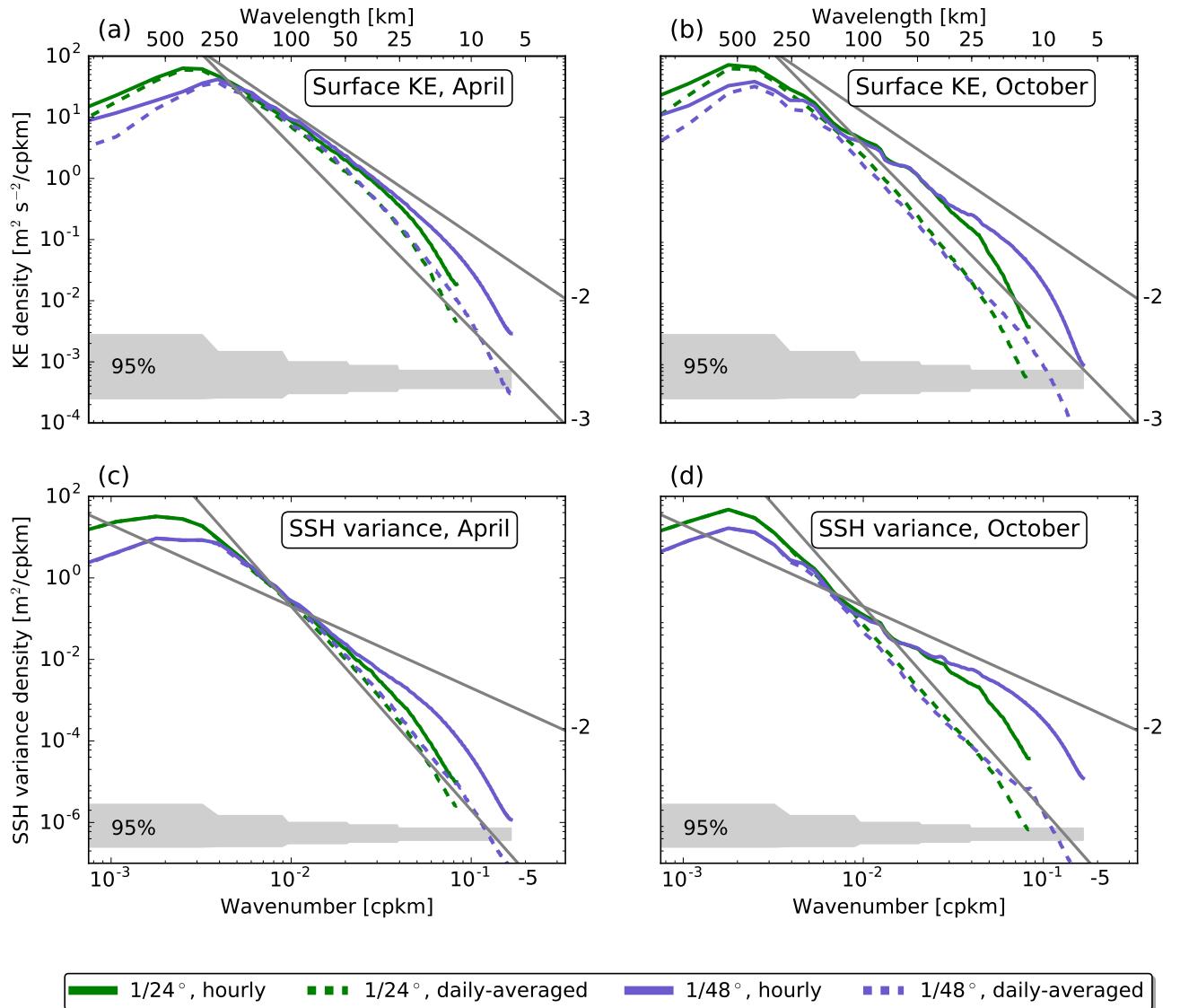


Figure 6. A comparison between spectra of the two LLC simulations $1/24^\circ$ (green) and $1/48^\circ$ (purple). The statistical errorbars represent conservative significance levels considering that there are only four independent relations for the monthly spectra.

2.3. Seasonality in statistics: probability density functions

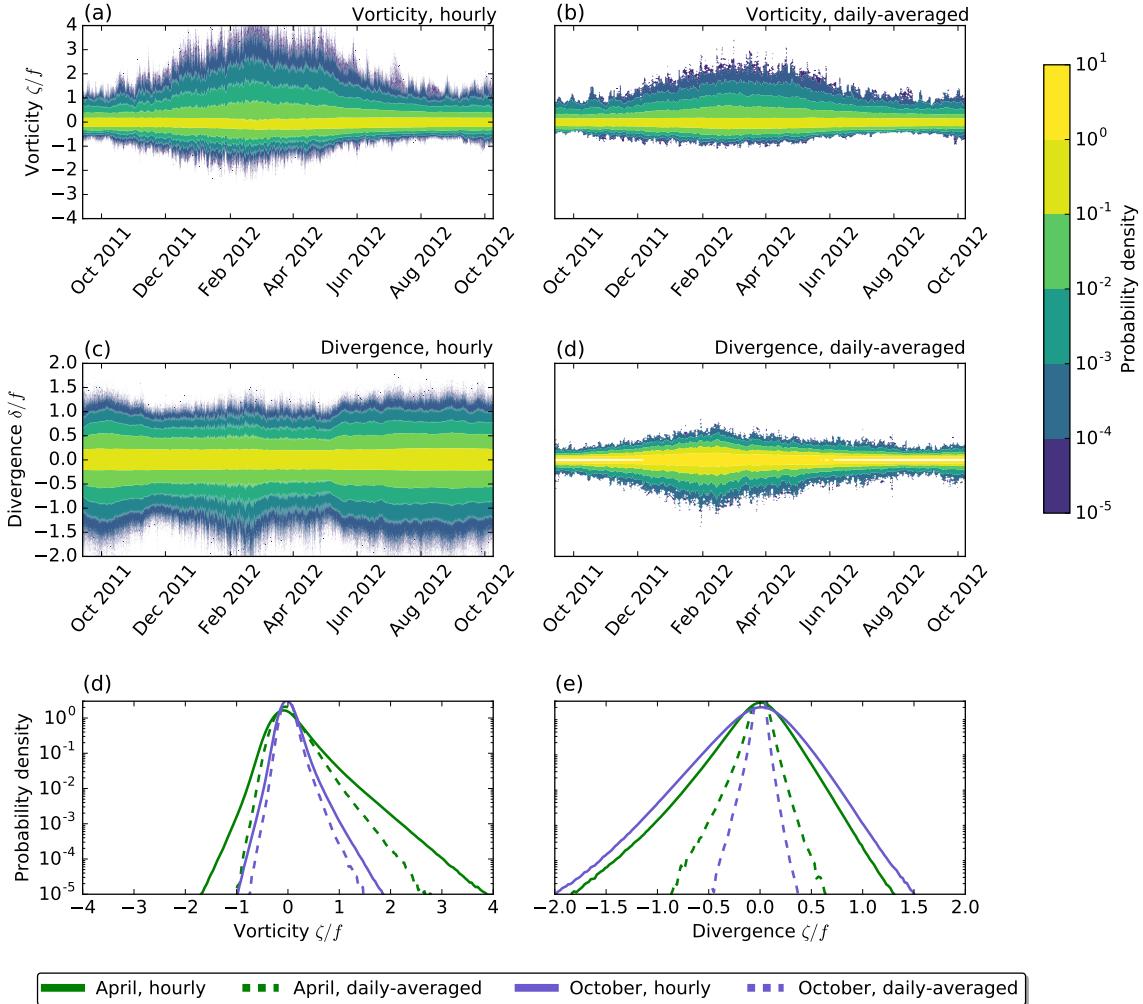


Figure 7. Time dependence of probability density functions of vertical vorticity (a, b) and horizontal divergence (c, d). Monthly averages in April and October are also shown (e, f).

2.4. How well does the model captures high-frequency modes?

To assess how well the LLC2160 represents the high-frequency variability, we compare the model near-surface velocity field against two available mooring data: KEO (the data is available online at <http://www.pmel.noaa.gov/ocs/KEO>) and

KESS mooring #7 (the data is available online at <http://uskesse.whois.edu>). For proper comparison, we use model data closest to these two moorings (west of the region considered in the present study). Focusing on high-frequencies, we calculate frequency spectral estimates every month. There are about 48 independent spectral realizations.

Figure 8 shows KE frequency spectra for the model and mooring data. At the uppermost common depth (20 m), very low frequencies (periods > 10 days) and

the inertial, diurnal, and semi-diurnal frequencies agree well — the KESS data are more energetic at intermediate frequencies. At 40 m, the model and observations have consistent spectra at periods larger than about 6 h. The data from moorings are more energetic at higher frequencies and tidal harmonics are not observed, except for the first two harmonics that are apparent in the

KESS data at 40 m. Three reasons for the lack of tidal harmonics in the mooring data are: 1) The high-frequency variability is dominated by high-frequency noise due to mooring vertical excursions; 2) The model does not have enough resolution to resolve the full spectrum of ultraviolet catastrophic wave-wave interactions; 3) The model is not forced at frequencies higher than 6 h.

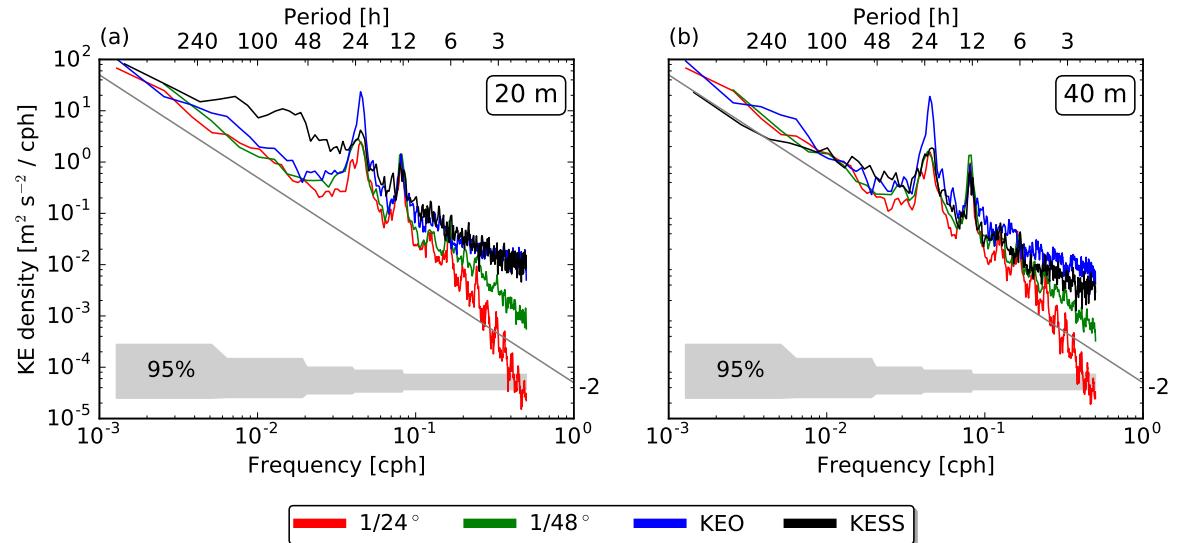


Figure 8. A comparison between frequency spectra of LLC simulations and two available moored current meter records.

Appendix C: Supplemental material 3: Seasonality in pressure modes near-surface amplitude

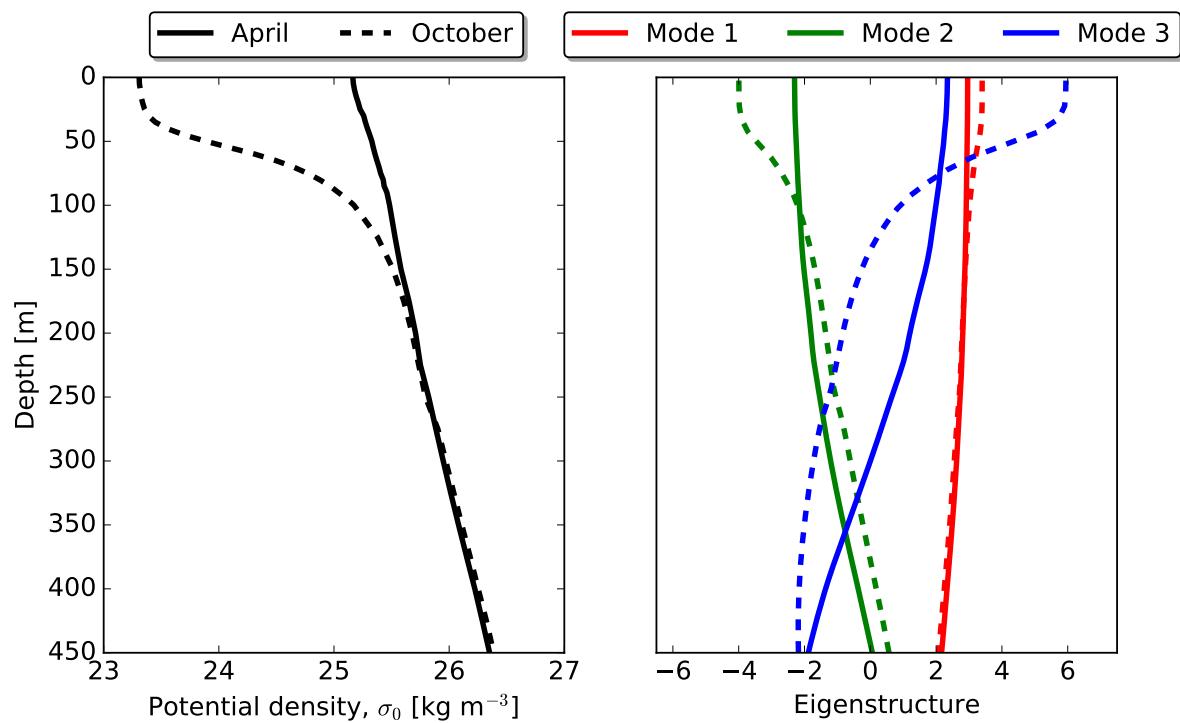


Figure 9. The seasonal variability of the WOA 2013 stratification averaged over the domain and the associated gravest three pressure modes. Only the upper 450 m is shown.