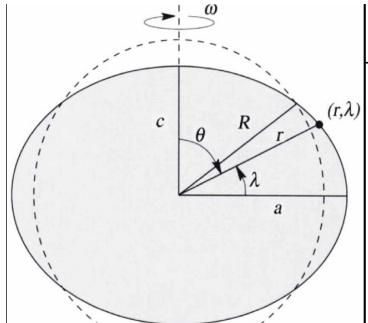


ELLIPSOID



Surface of constant gravity on a homogenous Earth

ellipsoid or spheroid

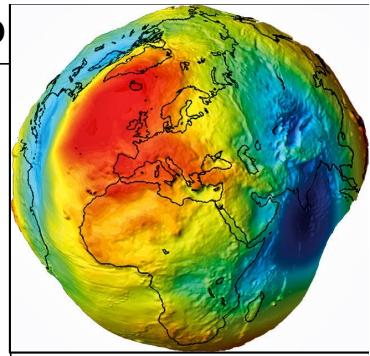
geoid

R/V Sally Ride

g_{obs}
ocean

[1] Definitions

GEOID



Surface of constant gravity on a non-homogenous Earth

deflection of the vertical

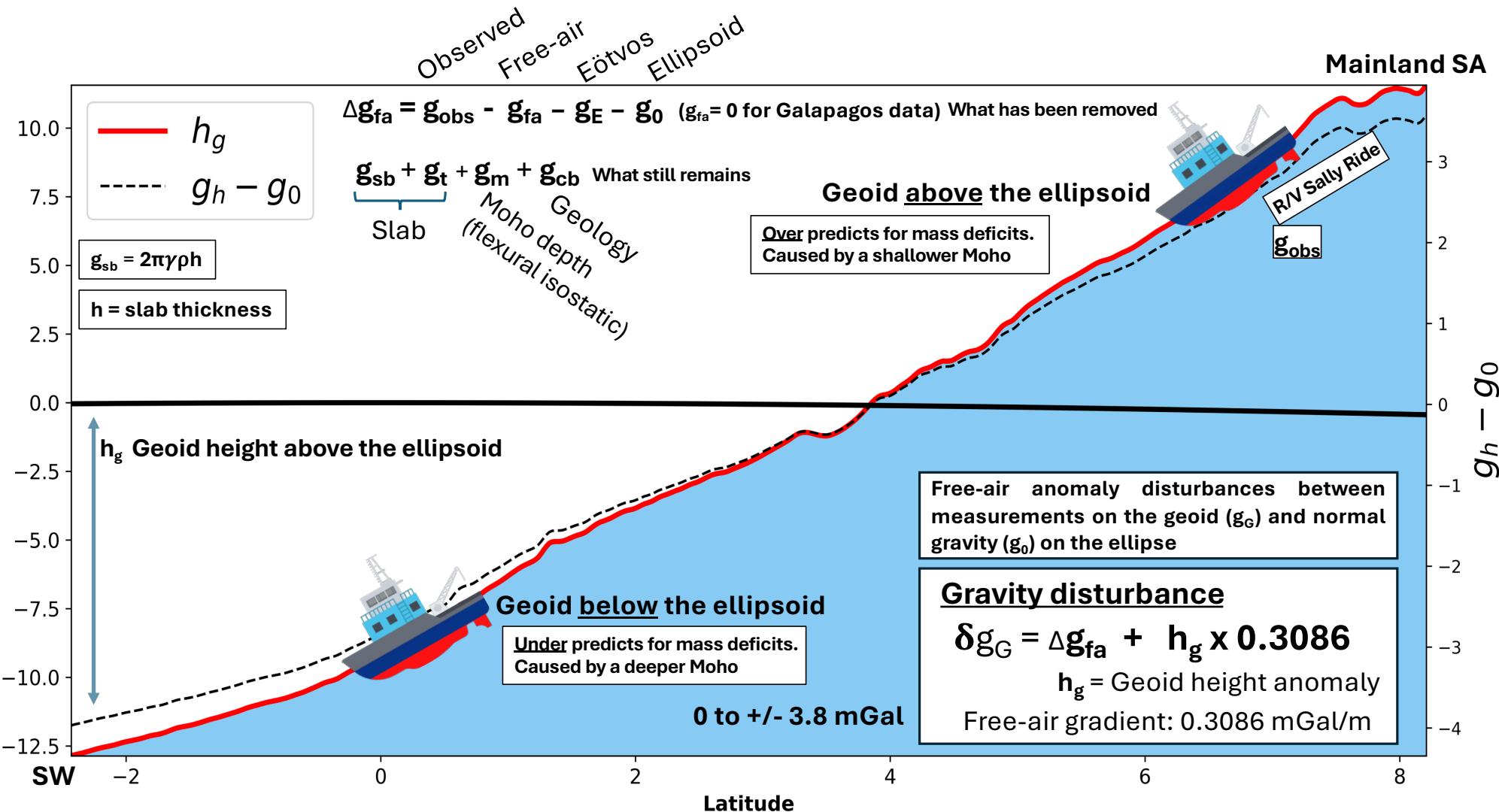
rock surface

geoid height

Eq. 7.12 (Blakely)

g_{obs} Observed gravity =

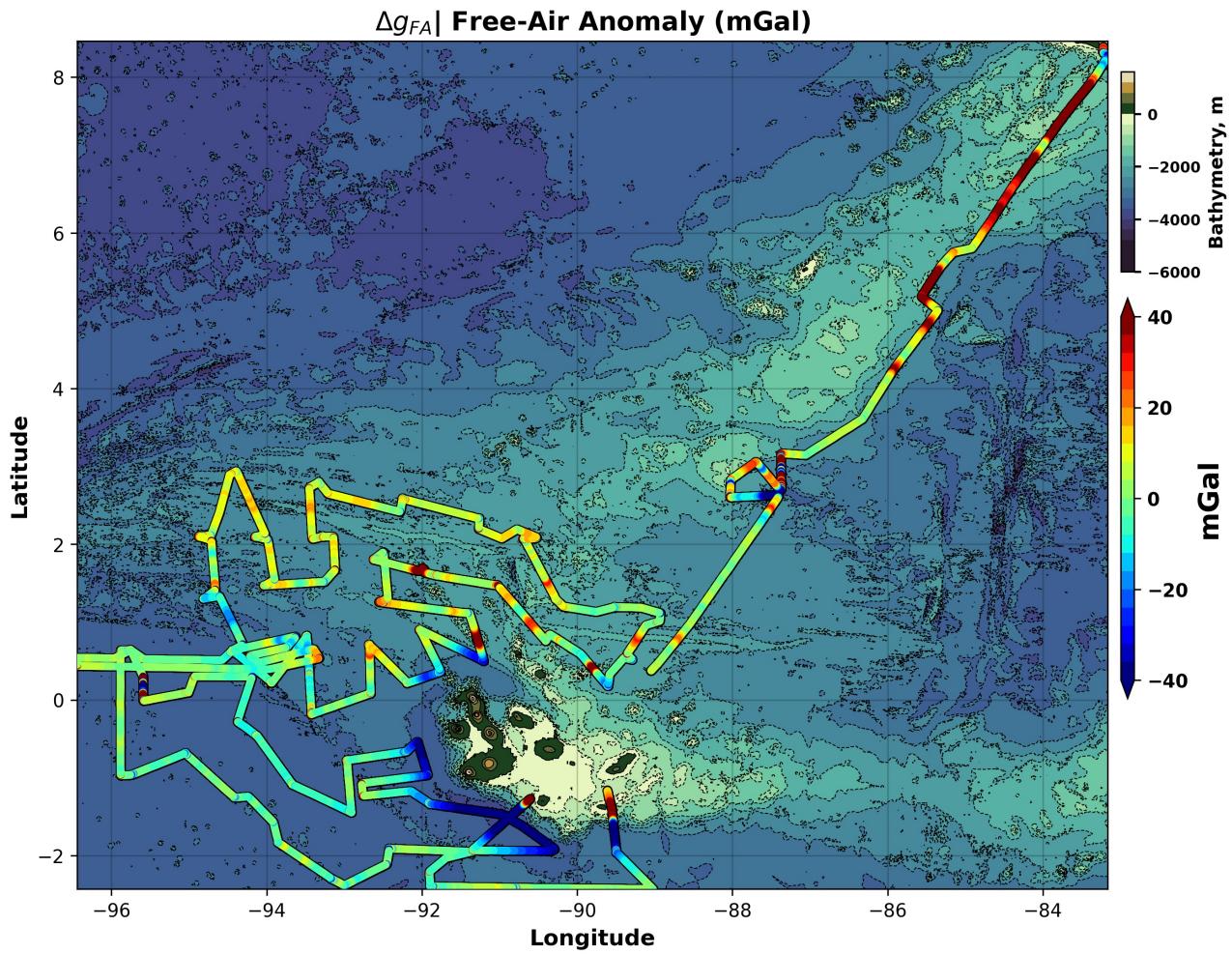
- g_0 Attraction of reference ellipsoid
- + g_{fa} Elevation above sea level (**free-air** = 0 for the Sally Ride)
- + g_E Eötvos (**centrifugal force** of rotating non-uniform sphere)
- + $g_{sb} + g_t$ Mass above sea level (g_t accounting for terrain)
- + g_m Isostatic residual compensation (Airy Moho depth)
- + g_d Gravity tides, $< 0.3 \text{mGal} \approx 0$
- + g_{cb} Crust and Mantle Lateral density variations (**geology**)



[2] Crossing the Geoid Anomaly & Gravity Disturbances

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

($g_{fa} = 0$ for Galapagos survey)

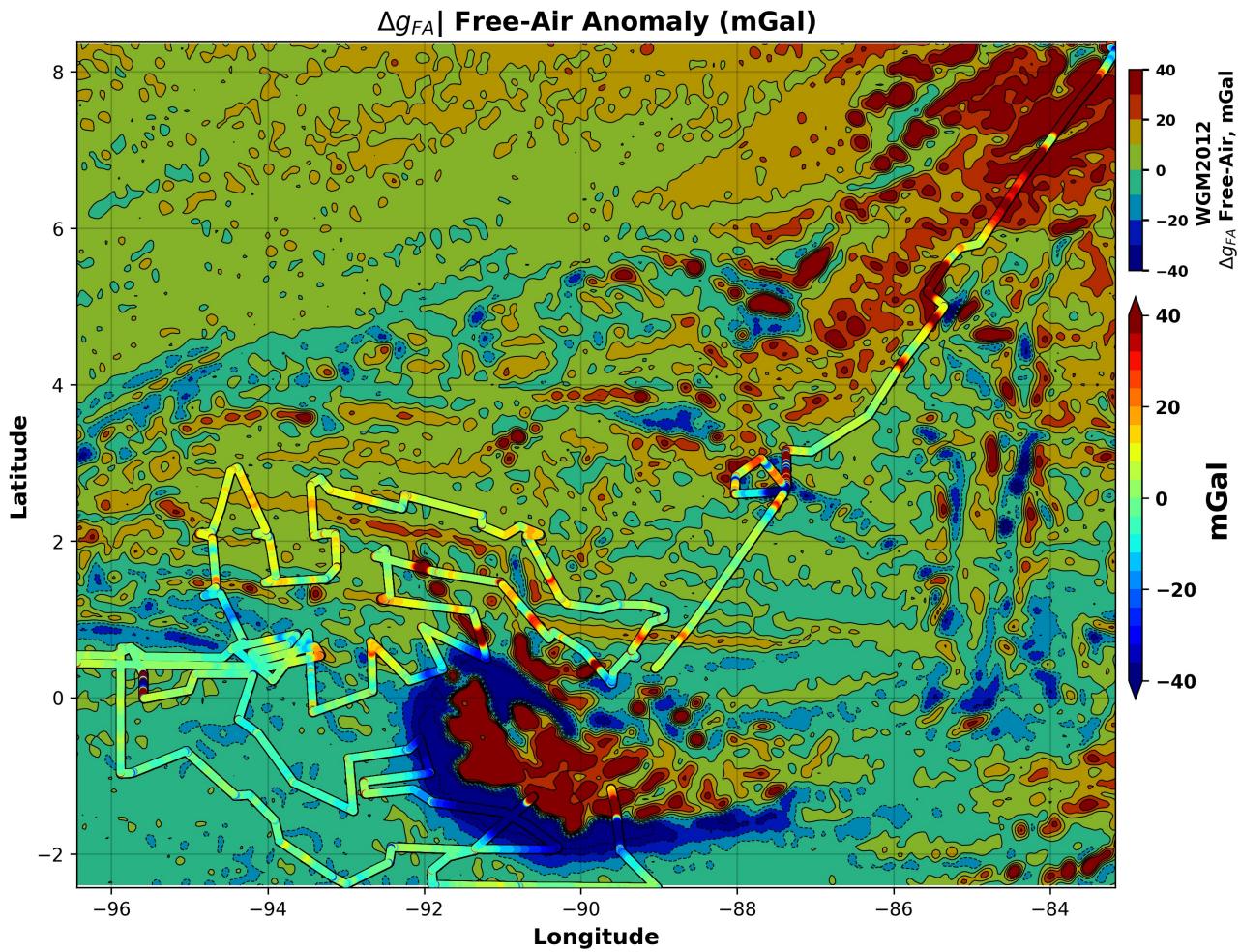


[3] Iguana Survey Free Air Anomaly

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

($g_{fa} = 0$ for Galapagos survey)

WGM2012 (Bonvalot, 2012) Free-Air Anomalies



Bonvalot, S., Balmino, G., Briais, A., M. Kuhn, Peyrefitte, A., Vales N., Biancale, R., Gabalda, G., Reinquin, F., Sarrailh, M., 2012. World Gravity Map. Commission for the Geological Map of the World. Eds., Paris.

[4] Iguana Survey Free Air Anomaly compared with Global Gravity Models (WGM2012)

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

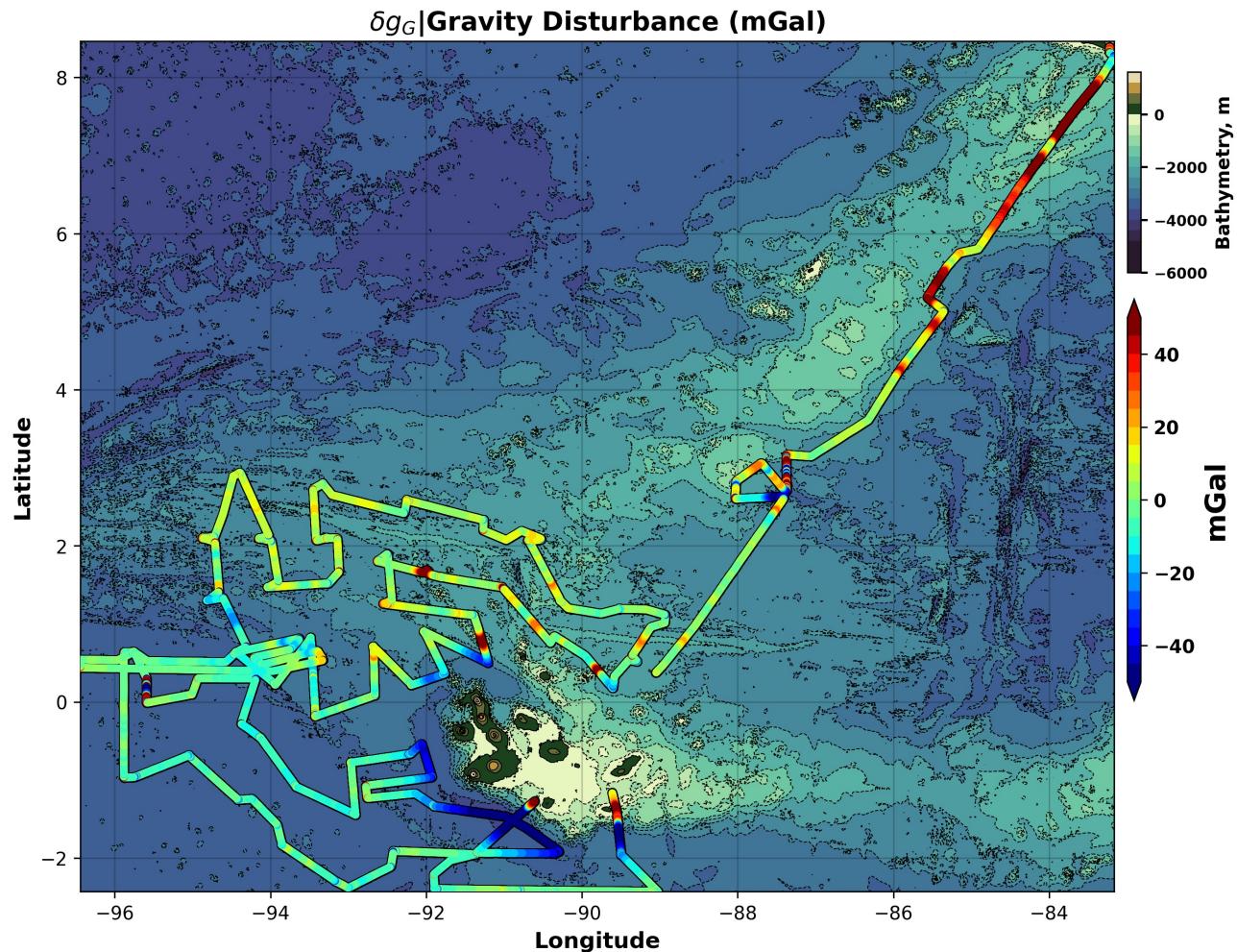
($g_{fa} = 0$ for Galapagos survey)

WGM2012 (Bonvalot, 2012) Free-Air Anomalies

$$\delta g_G = \Delta g_{fa} + h_g \times 0.3086$$

Gravity Anomalies: The (geodetic) gravity anomaly is defined as the difference between gravity on the geoid and normal gravity on the reference ellipsoid (Heiskanen & Moritz, 1967)

Gravity Disturbances: The gravity disturbance is defined as the difference of two fields at the same point on the reference ellipsoid (Hackney & Featherstone, 2003)



[5] Iguana Survey Geoid Gravity Disturbance, δg_G

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

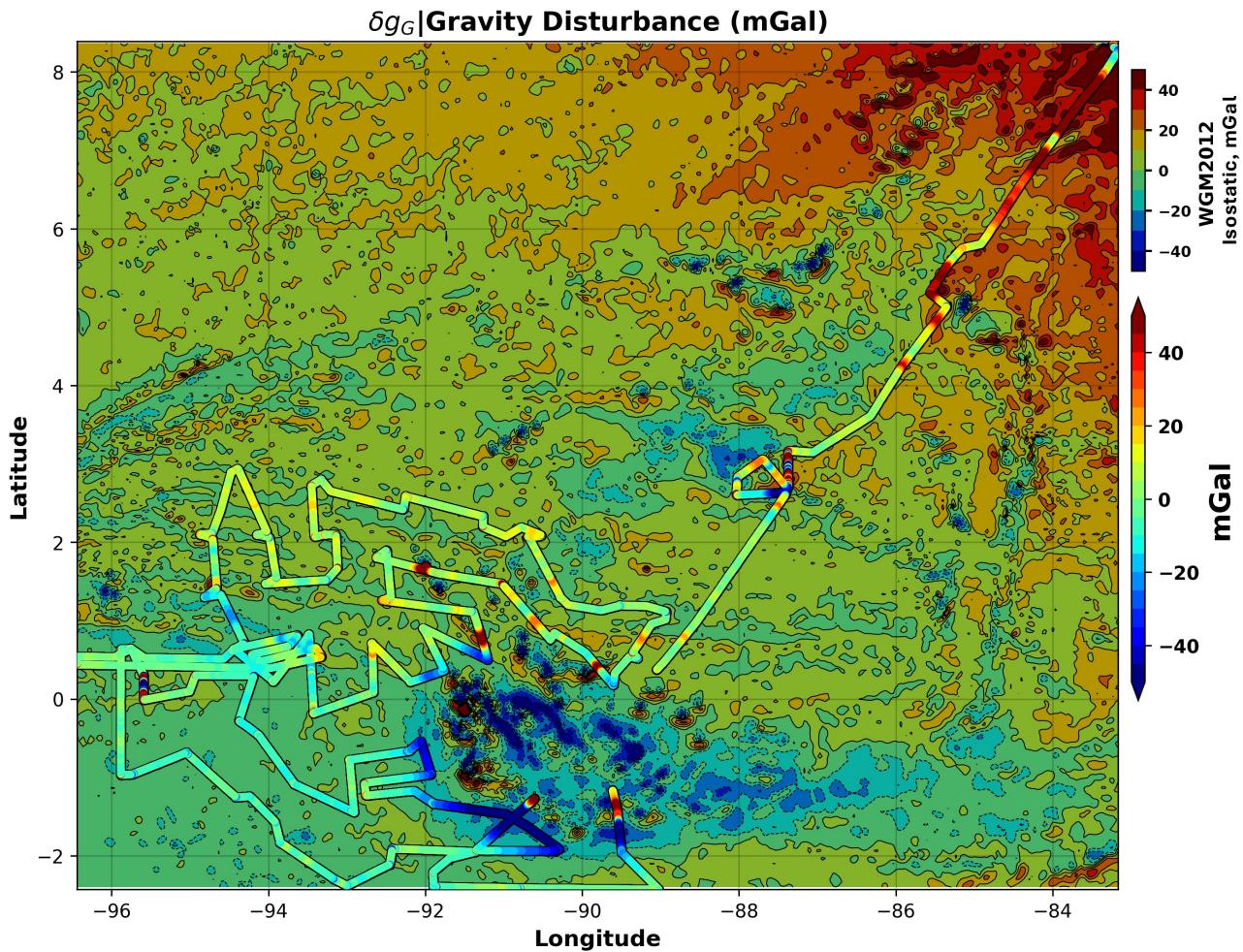
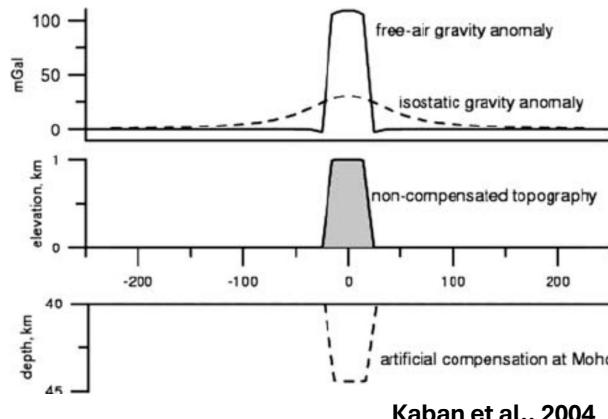
$(g_{fa} = 0 \text{ for Galapagos survey})$

WGM2012 (Bonvalot, 2012) Free-Air Anomalies

$$\delta g_G = \Delta g_{fa} + h_g \times 0.3086$$

$$\text{WGM2012, } \delta g_I = \Delta g_{fa} - A^T + A_C$$

A_C : Isostatic compensation ($\rho_c = 2670 \text{ kg/m}^3$)
 A^T : Topographic gravity



Bonvalot, S., Balmino, G., Briais, A., M. Kuhn, Peyrefitte, A., Vales N., Biancale, R., Gabalda, G., Reinquin, F., Sarraih, M., 2012. World Gravity Map. Commission for the Geological Map of the World. Eds., Paris.

[6] Comparing the δg_P with δg_I , the WGM Isostatic Gravity Disturbance

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

($g_{fa} = 0$ for Galapagos survey)

WGM2012 (Bonvalot, 2012) Free-Air Anomalies

$$\delta g_G = \Delta g_{fa} + h_g \times 0.3086$$

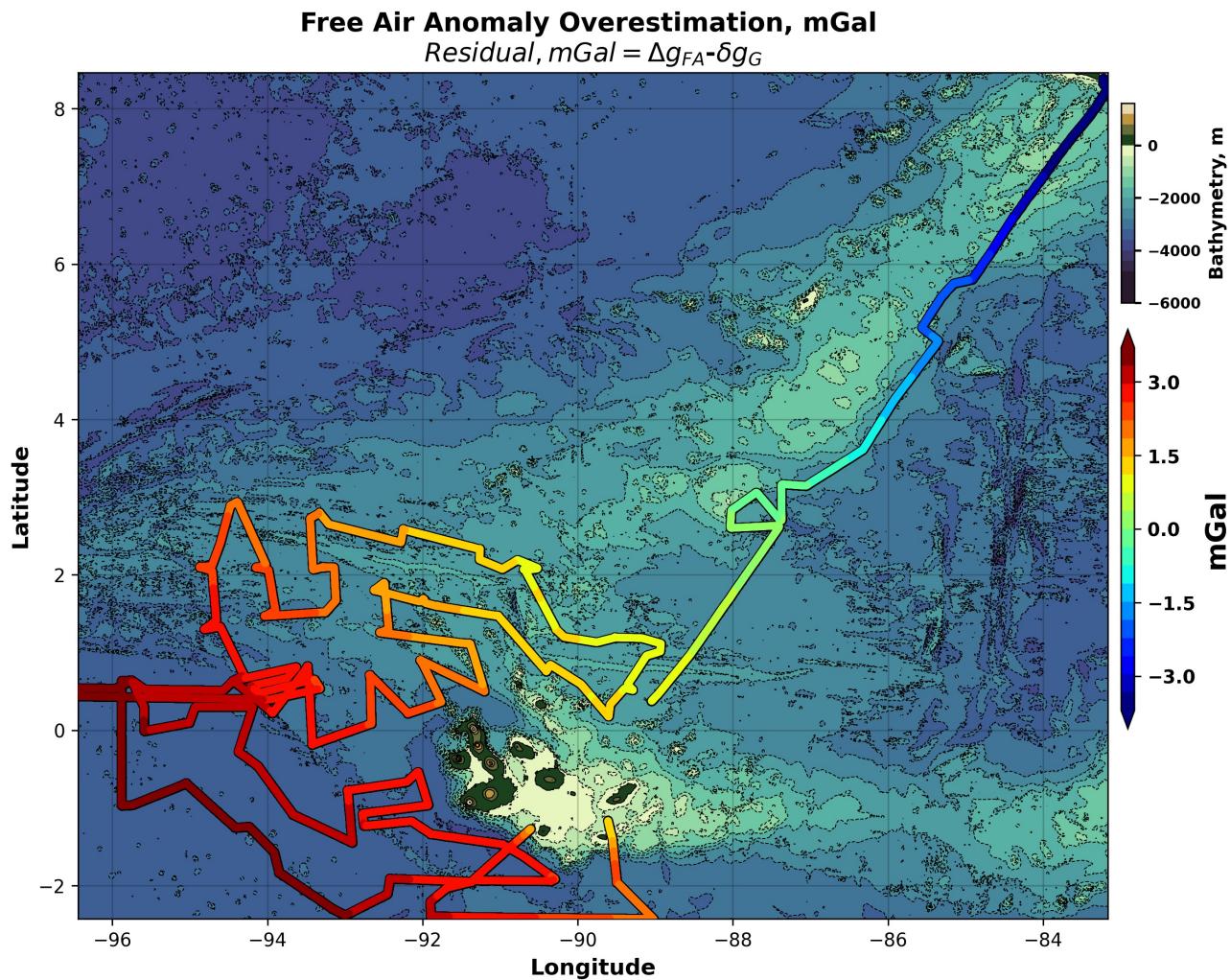
$$WGM2012, \delta g_I = \Delta g_{fa} - A^T + A_C$$

A_C : Isostatic compensation ($\rho_c = 2670 \text{ kg/m}^3$)

A^T : Topographic gravity

$$\Delta g_{fa} - \delta g_P = \Delta g_{fa} - \Delta g_{fa} + h_g \times 0.3086$$

$$= h_g \times 0.3086$$



[7] $\Delta g_{fa} - \delta g_P$ (The Geoid Gravity anomaly)

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

$(g_{fa} = 0 \text{ for Galapagos survey})$

WGM2012 (Bonvalot, 2012) Free-Air Anomalies

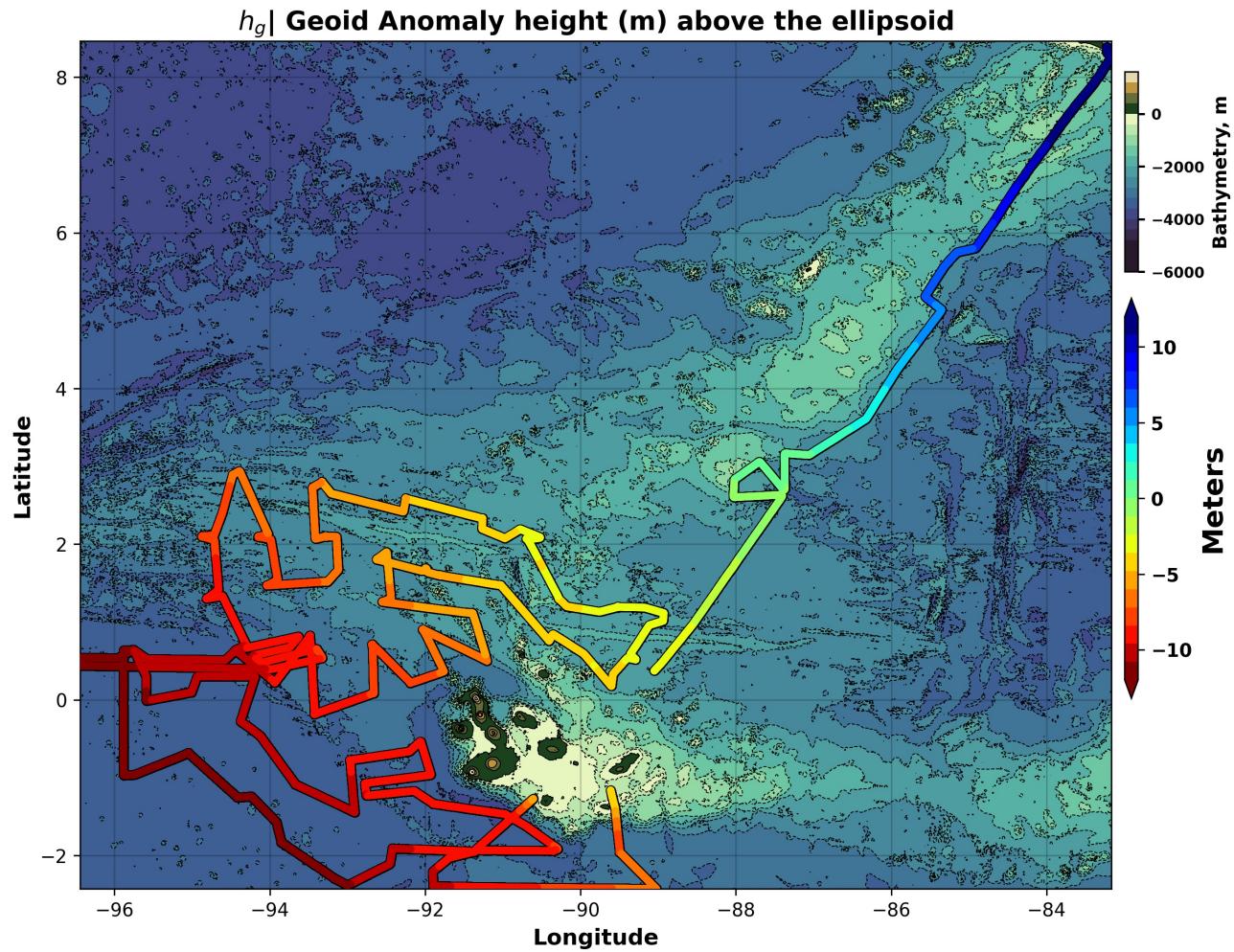
$$\delta g_G = \Delta g_{fa} + h_g \times 0.3086$$

$$\text{WGM2012, } \delta g_I = \Delta g_{fa} - A^T + A_C$$

A_C : Isostatic compensation ($\rho_c = 2670 \text{ kg/m}^3$)
 A^T : Topographic gravity

$$\Delta g_{fa} - \delta g_P = \Delta g_{fa} - \Delta g_{fa} + h_g \times 0.3086$$

$$= h_g \times 0.3086$$



[8] The Geoid Height anomaly

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

$(g_{fa} = 0 \text{ for Galapagos survey})$

WGM2012 (Bonvalot, 2012) Free-Air Anomalies

$$\delta g_G = \Delta g_{fa} + h_g \times 0.3086$$

$$\text{WGM2012, } \delta g_I = \Delta g_{fa} - A^T + A_C$$

A_C : Isostatic compensation ($\rho_c = 2670 \text{ kg/m}^3$)
 A^T : Topographic gravity

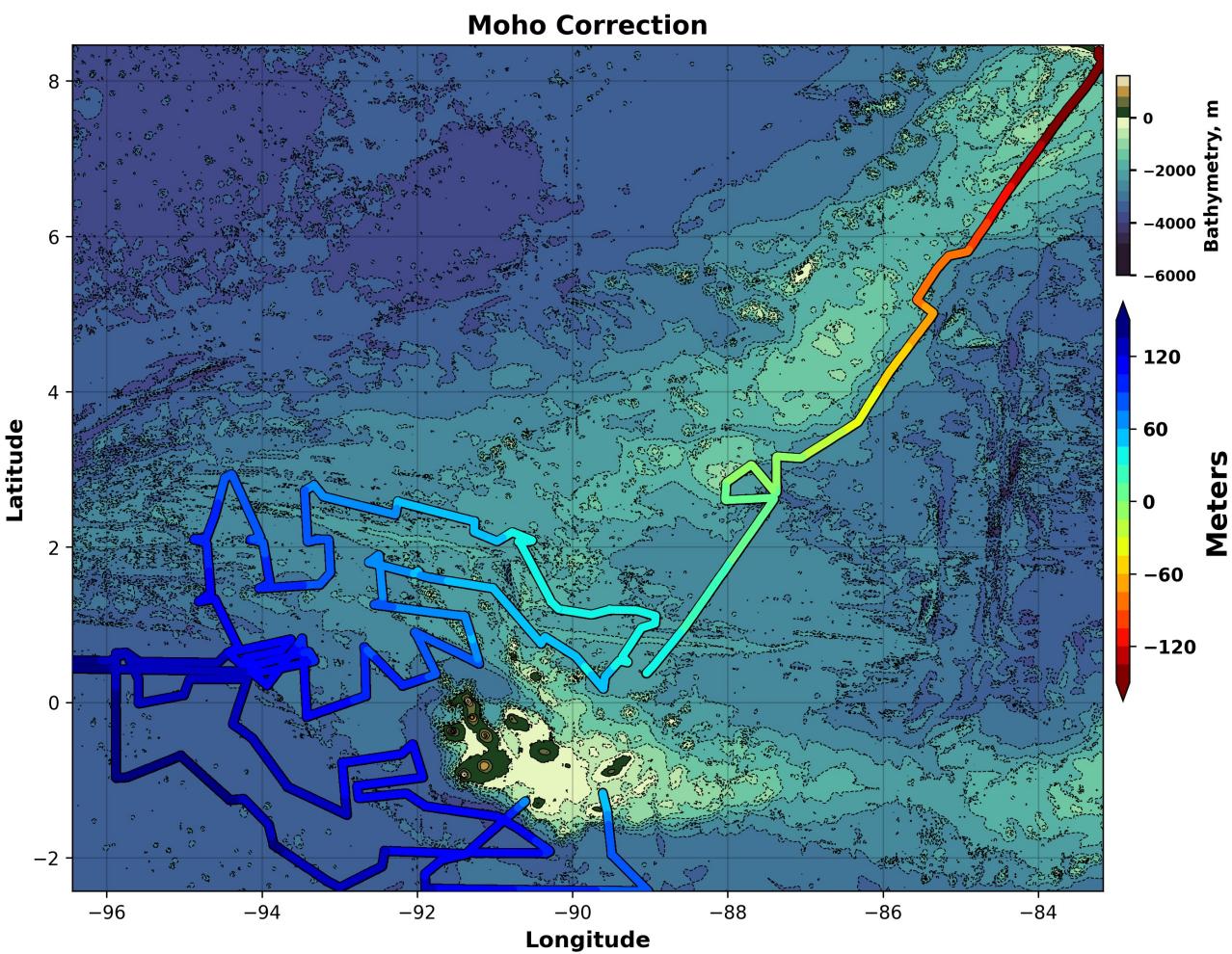
$$\Delta g_{fa} - \delta g_P = \Delta g_{fa} - \Delta g_{fa} + h_g \times 0.3086$$

$$= h_g \times 0.3086$$

$$\Delta g_{fa} - \delta g_P = h_g \times 0.3086$$

$$\Delta \rho = 600 \text{ kg/m}^3$$

$$\delta h_m = (\Delta g_{fa} - \delta g_P) / 2\pi\gamma\Delta\rho$$



[9] Moho depth corrections caused by the Geoid

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

($g_{fa} = 0$ for Galapagos survey)

WGM2012 (Bonvalot, 2012) Free-Air Anomalies

$$\delta g_G = \Delta g_{fa} + h_g \times 0.3086$$

$$WGM2012, \delta g_I = \Delta g_{fa} - A^T + A_C$$

A_C : Isostatic compensation ($\rho_c = 2670 \text{ kg/m}^3$)
 A^T : Topographic gravity

$$\Delta g_{fa} - \delta g_P = \Delta g_{fa} - \Delta g_{fa} + h_g \times 0.3086$$

$$= h_g \times 0.3086$$

$$\Delta g_{fa} - \delta g_P = h_g \times 0.3086$$

$$\Delta \rho = 600 \text{ kg/m}^3$$

$$\delta h_m = (\Delta g_{fa} - \delta g_P) / 2\pi\gamma\Delta\rho$$

$$\Delta g_{SB} = \Delta g_{fa} - 2\pi\gamma\rho_c h$$

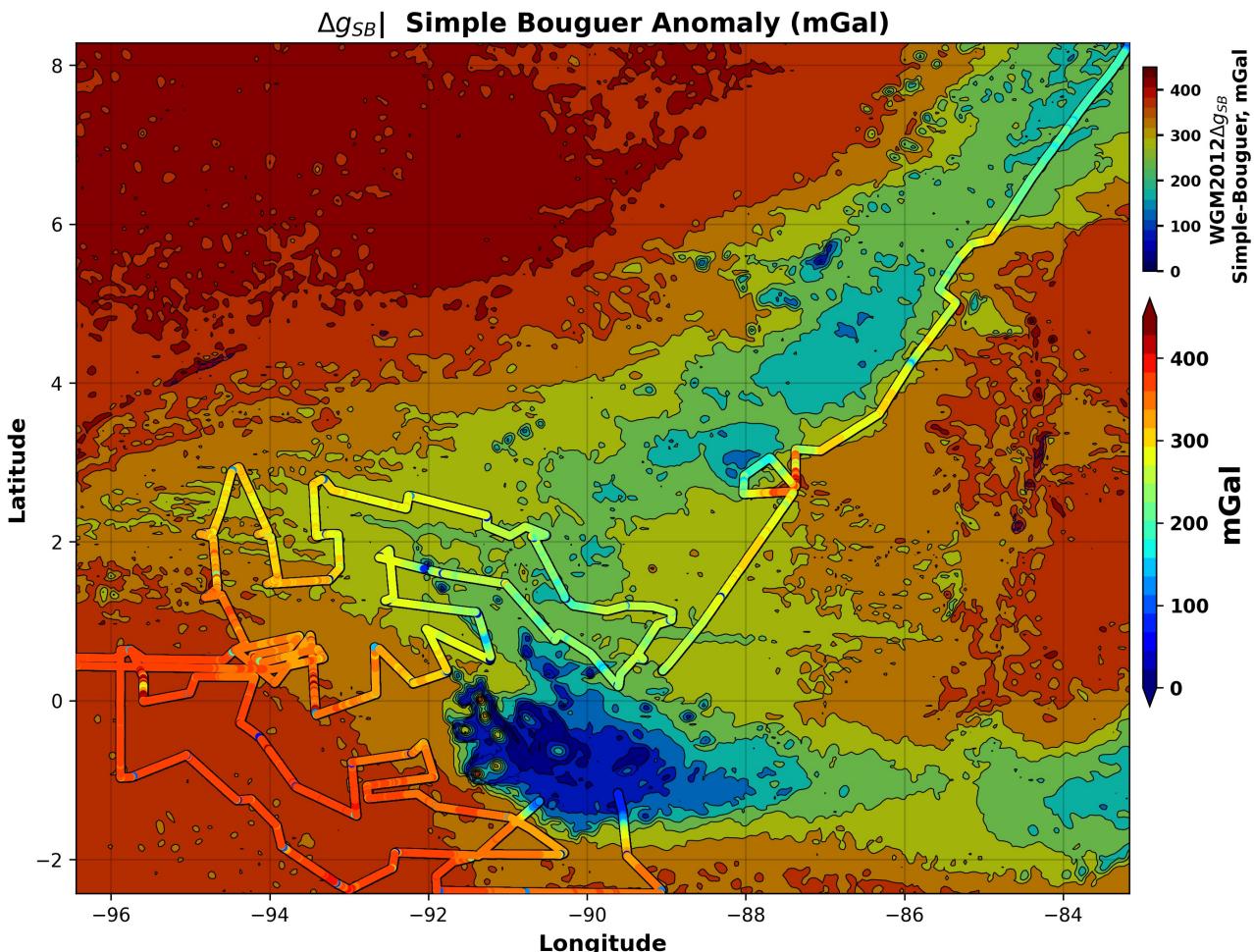
h = Slab Thickness
 $(\rho_c = 2670 \text{ kg/m}^3)$

$$\Delta g_{CB} = g_{obs} - g_{fa} - g_{sb} - g_t - g_0 = \Delta g_{SB} - g_t$$

z = Depth

$$(\Delta \rho = \rho_c - \rho_w = 1670 \text{ kg/m}^3)$$

$$g_t = \Delta \rho 2\pi\gamma [R - \sqrt{(R^2 + z^2) + z}] \quad (\text{LaFehr, 1991})$$



R = Average spacing (1,000m)

LaFehr, T. R. "Standardization in gravity reduction." Geophysics 56.8 (1991): 1170-1178.

[11] Simple Bouguer with Topo

$$\Delta g_{fa} = g_{obs} - g_{fa} - g_E - g_0$$

($g_{fa} = 0$ for Galapagos survey)

WGM2012 (Bonvalot, 2012) Free-Air Anomalies

$$\delta g_G = \Delta g_{fa} + h_g \times 0.3086$$

$$WGM2012, \delta g_I = \Delta g_{fa} - A^T + A_C$$

A_C : Isostatic compensation ($\rho_c = 2670 \text{ kg/m}^3$)

A^T : Topographic gravity

$$\Delta g_{fa} - \delta g_P = \Delta g_{fa} - \Delta g_{fa} + h_g \times 0.3086$$

$$= h_g \times 0.3086$$

$$\Delta g_{fa} - \delta g_P = h_g \times 0.3086$$

$$\Delta \rho = 600 \text{ kg/m}^3$$

$$\delta h_m = (\Delta g_{fa} - \delta g_P) / 2\pi\gamma\Delta\rho$$

$$\Delta g_{SB} = \Delta g_{fa} - 2\pi\gamma\rho_c h$$

h = Slab Thickness

$$(\rho_c = 2670 \text{ kg/m}^3)$$

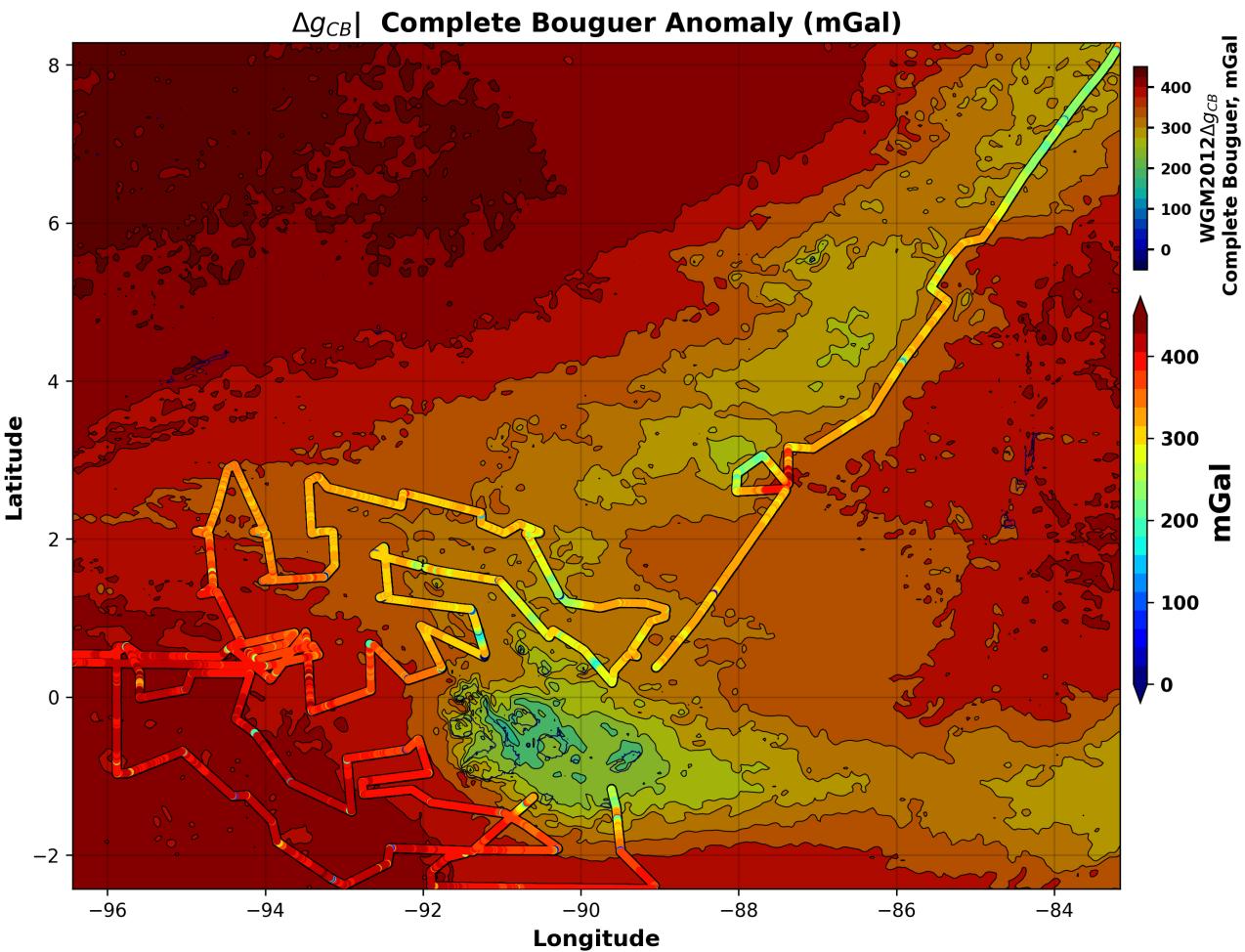
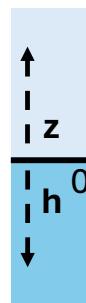
$$\Delta g_{CB} = g_{obs} - g_{fa} - g_{sb} - g_t - g_0 = \Delta g_{SB} - g_t$$

z = Depth

$$(\Delta \rho = \rho_c - \rho_w = 1670 \text{ kg/m}^3)$$

$$g_t = \Delta \rho 2\pi\gamma [R - \sqrt{(R^2 + z^2) + z}]$$

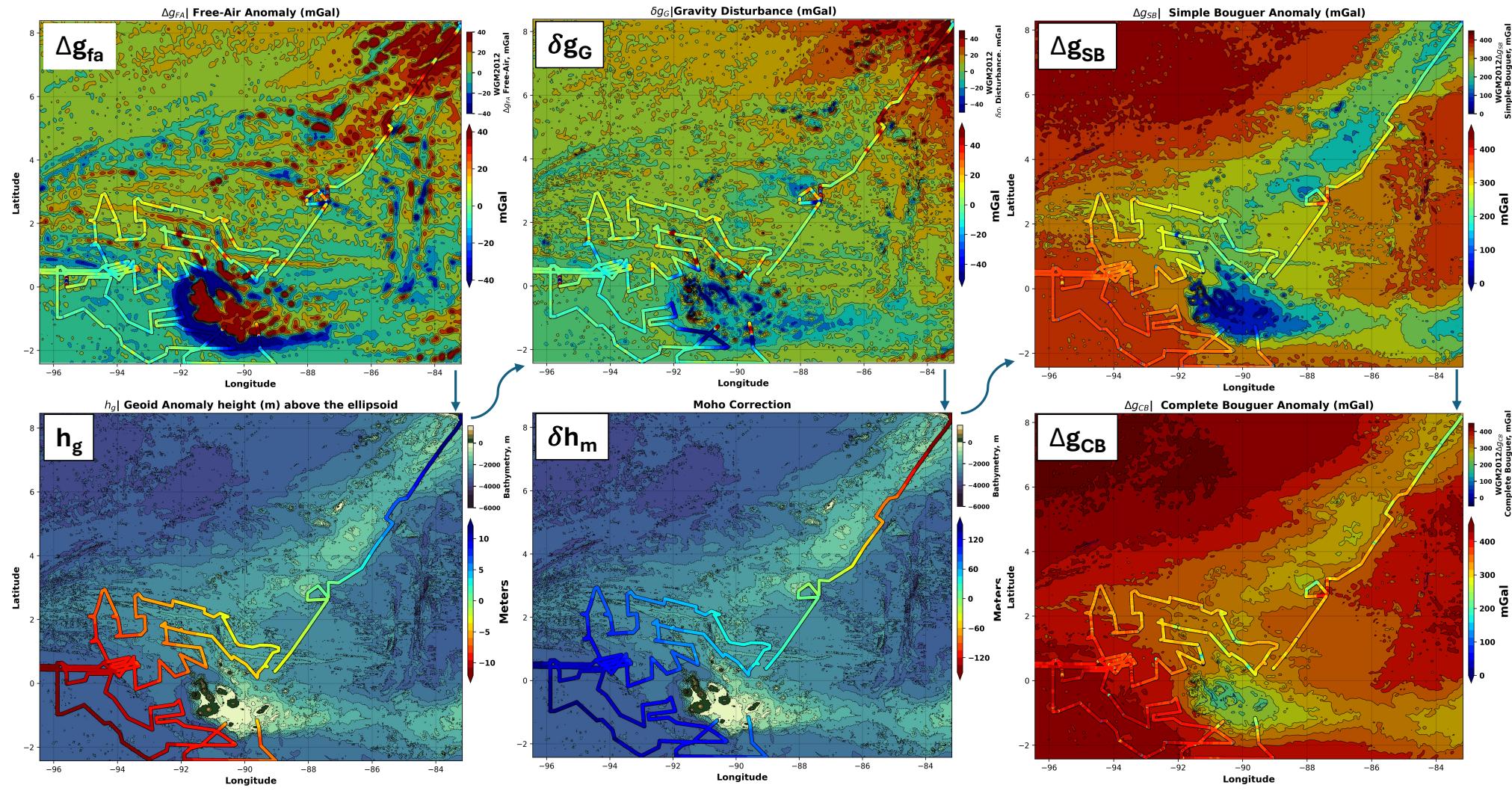
(LaFehr, 1991)



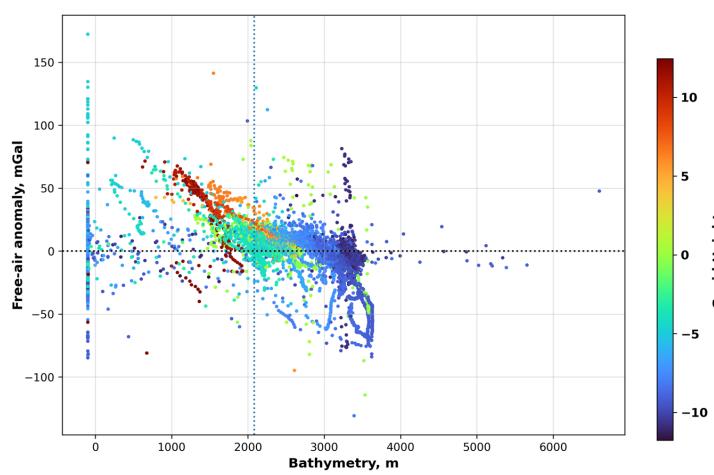
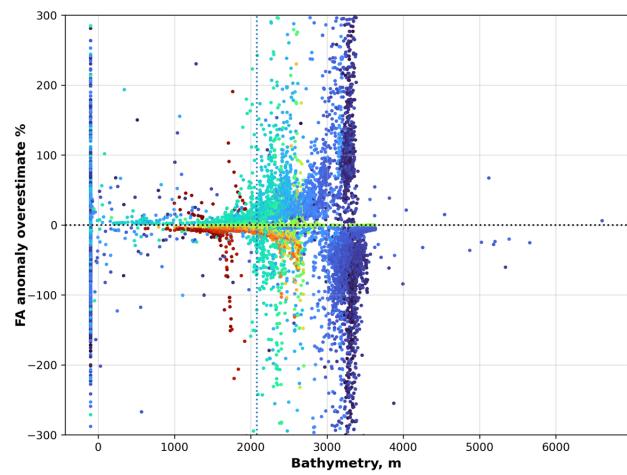
R = Average spacing (1,000m)

LaFehr, T. R. "Standardization in gravity reduction." Geophysics 56.8 (1991): 1170-1178.

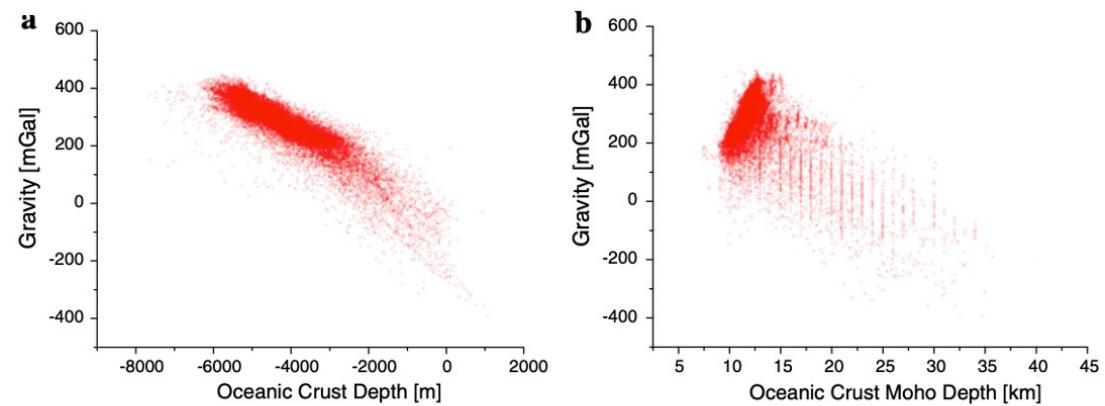
[12] Complete Bouguer



[13] End



Gravity disturbances	Correlation with	
	Solid topography	Moho geometry
δg	-0.01	-0.09
δg^T	-0.55	-0.68
δg^{TB}	-0.89	-0.95
δg^{TBI}	-0.92	-0.97
δg^{TBIS}	-0.94	-0.97
δg^{cs}	-0.95	-0.98



[Supplement] Moho depth controls