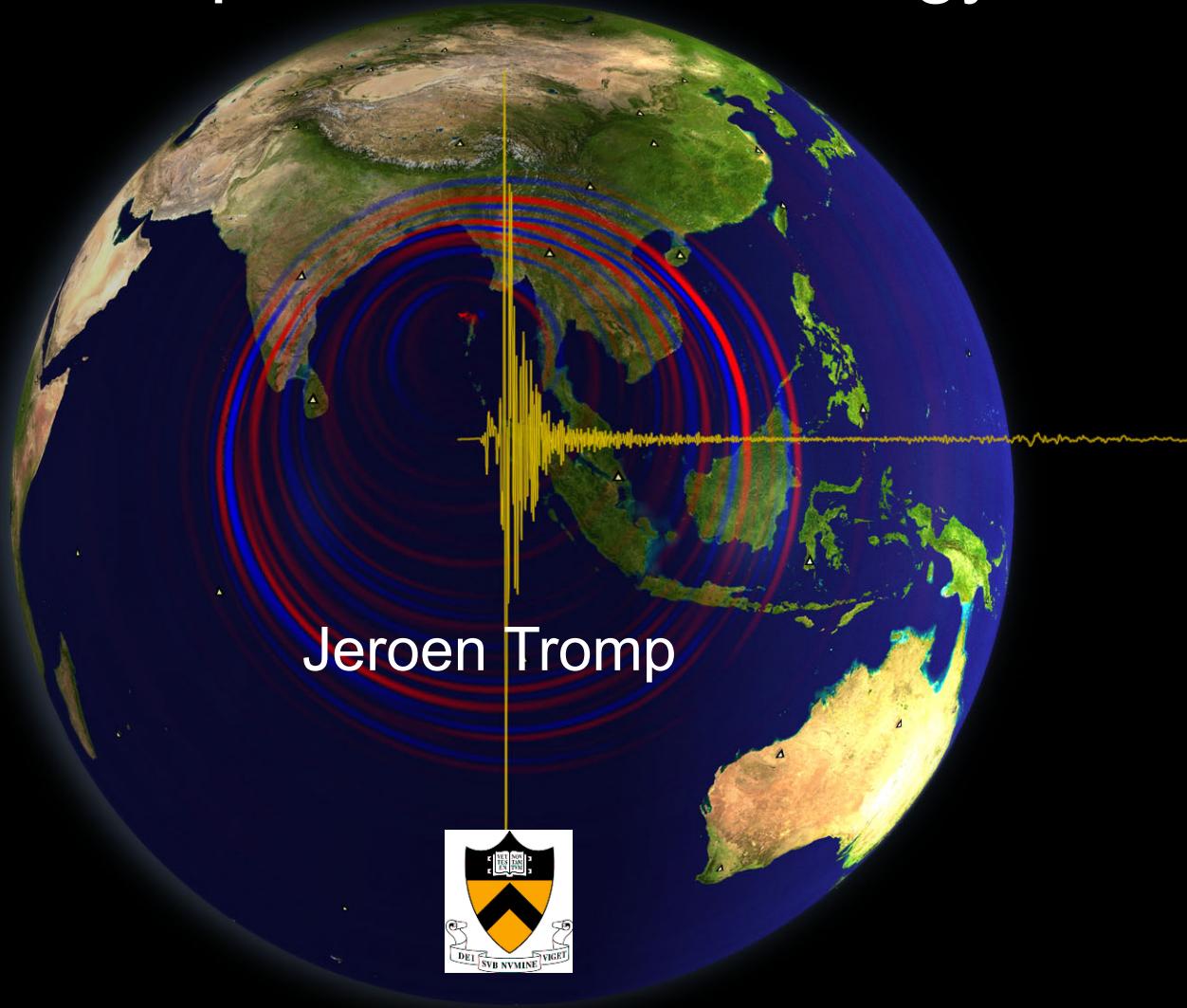


# Computational Seismology





# Governing Equations

Equation of motion:

$$\rho \partial_t^2 \mathbf{s} - \nabla \cdot \mathbf{T} = \mathbf{f}$$

Constitutive relationship (Hooke's law):

$$\mathbf{T} = \mathbf{c} : \boldsymbol{\epsilon}$$

Boundary condition:

$$\hat{\mathbf{n}} \cdot \mathbf{T} = 0$$

Initial conditions:

$$\mathbf{s}(\mathbf{x}, 0) = \mathbf{0}, \quad \partial_t \mathbf{s}(\mathbf{x}, 0) = \mathbf{0}$$

Earthquake source:

$$\mathbf{f} = -\mathbf{M} \cdot \nabla \delta(\mathbf{x} - \mathbf{x}_s) S(t)$$



# Classic Tools

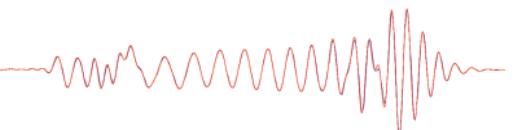
- Semi-analytical methods
  - Layer cake models: reflectivity/discrete wave number
  - Spherically symmetric models: normal modes
- Asymptotic methods
  - Ray theory
  - WKBJ theory
  - Coupled modes

COMPUTATIONAL INFRASTRUCTURE FOR GEODYNAMICS (CIG)



Mineos

User Manual  
Version 1.0



Guy Masters  
Misha Barmine  
Susan Kientz

[www.geodynamics.org](http://www.geodynamics.org)



# Strong Form

- Velocity-stress formulation (first-order PDEs,  $\mathbf{v} = \partial_t \mathbf{s}$ ):

$$\partial_t \mathbf{v} = \rho^{-1} (\nabla \cdot \mathbf{T} + \mathbf{f}), \quad \partial_t \mathbf{T} = \mathbf{c} : \nabla \mathbf{v}$$

- Second-order finite-difference method:

$$\frac{d}{dx} f(n\Delta x) = (f_{n+1} - f_{n-1}) / (2\Delta x), \\ n = 1, \dots, N-2$$

- Fourth-order finite-difference method:

$$\frac{d}{dx} f(n\Delta x) = (-f_{n+2} + 8f_{n+1} - 8f_{n-1} + f_{n-2}) / (12\Delta x), \\ n = 2, \dots, N-3$$

- Challenges:
  - Boundary conditions
  - Numerical dispersion & anisotropy



# Strong Form

- Pseudospectral methods:

$$F(l\Delta k) = \Delta x \sum_{n=0}^{N-1} f(n\Delta x) \exp(-2\pi i nl/N),$$
$$l = 0, \dots, N-1$$

$$\frac{d}{dx} f(n\Delta x) = \frac{1}{N\Delta x} \sum_{l=0}^{N-1} i(l\Delta k) F(l\Delta k) \exp(2\pi i nl/N),$$
$$n = 0, \dots, N-1$$

- Challenges:
  - Boundary conditions
  - Parallel implementation



# Weak Form

$$\int_{\Omega} \rho \mathbf{w} \cdot \partial_t^2 \mathbf{s} d^3 \mathbf{x} = - \int_{\Omega} \nabla \mathbf{w} : \mathbf{T} d^3 \mathbf{x} + \mathbf{M} : \nabla \mathbf{w}(\mathbf{x}_s) S(t)$$

- Weak form valid for any test vector
- Boundary conditions automatically included
- Source term explicitly integrated

Finite-fault (kinematic) rupture:

$$\mathbf{M} : \nabla \mathbf{w}(\mathbf{x}_s) S(t) \rightarrow \int_{S_s} \mathbf{m}(\mathbf{x}_s, t) : \nabla \mathbf{w}(\mathbf{x}_s) d^2 \mathbf{x}_s$$

# Finite Elements

Mapping from reference cube to hexahedral elements:

$$\mathbf{x}(\xi) = \sum_{a=1}^M \mathbf{x}_a N_a(\xi)$$

Volume relationship:

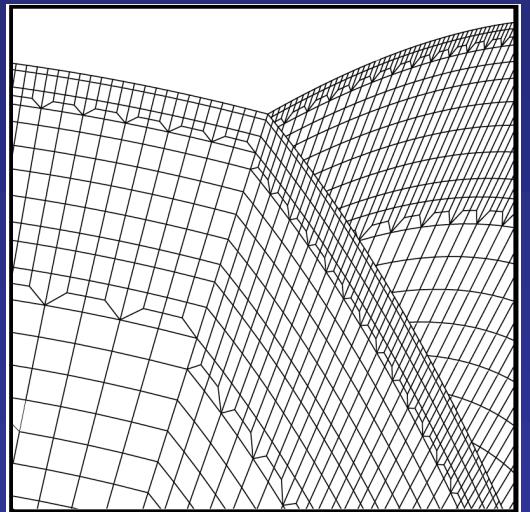
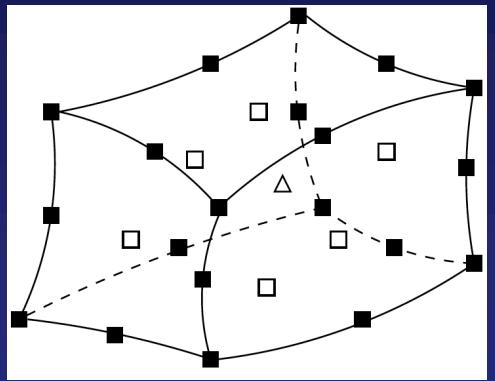
$$d^3\mathbf{x} = dx dy dz = J d\xi d\eta d\zeta = J d^3\xi$$

Jacobian of the mapping:

$$J = \left| \frac{\partial(x, y, z)}{\partial(\xi, \eta, \zeta)} \right|$$

Jacobian matrix:

$$\frac{\partial \mathbf{x}}{\partial \xi} = \sum_{a=1}^M \mathbf{x}_a \frac{\partial N_a}{\partial \xi}$$

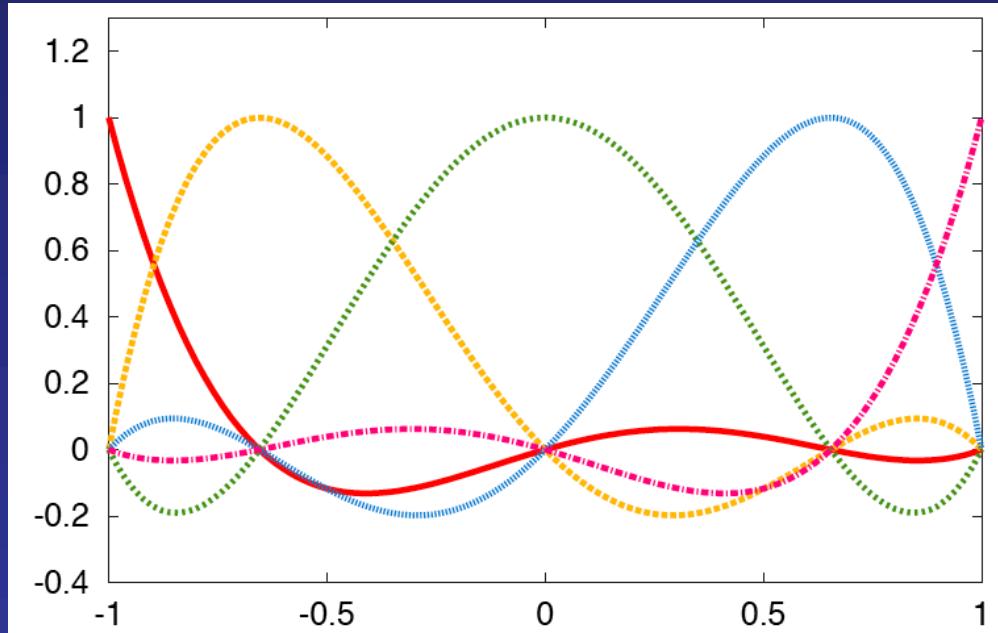


- Challenges:
  - Numerical anisotropy & dispersion
  - Mass lumping/implicit time stepping

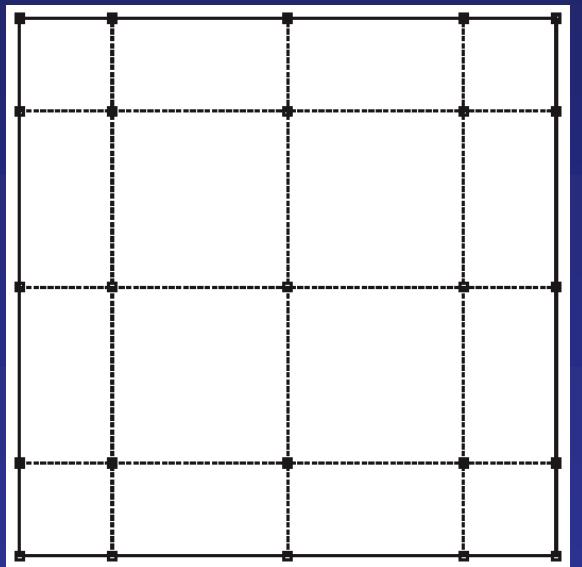


# Spectral Elements

The 5 degree 4 Lagrange polynomials:



Degree 4 GLL points:



GLL points are  $n+1$  roots of:

$$(1 - \xi^2) P'_n(\xi) = 0$$

General definition:

$$h_\alpha(\xi) = \frac{(\xi - \xi_0) \cdots (\xi - \xi_{\alpha-1})(\xi - \xi_{\alpha+1}) \cdots (\xi - \xi_n)}{(\xi_\alpha - \xi_0) \cdots (\xi_\alpha - \xi_{\alpha-1})(\xi_\alpha - \xi_{\alpha+1}) \cdots (\xi_\alpha - \xi_n)}$$

Note that at a GLL point:  $h_\alpha(\xi_\beta) = \delta_{\alpha\beta}$



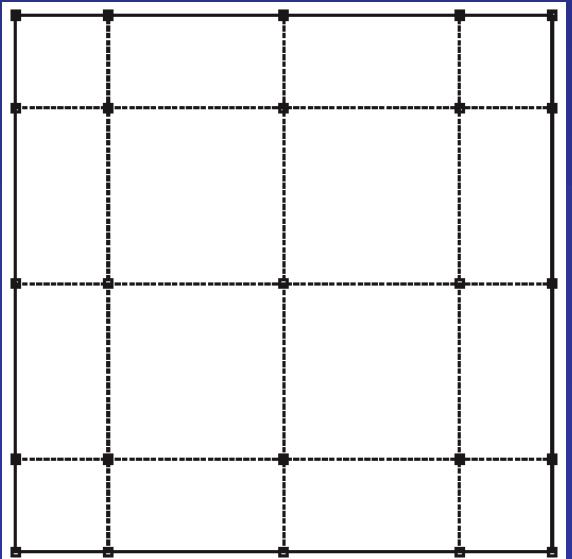
# Interpolation

Representation of functions on an element in terms of Lagrange polynomials:

$$f(\mathbf{x}(\xi, \eta, \zeta)) = \sum_{\alpha=0}^n \sum_{\beta=0}^n \sum_{\gamma=0}^n f^{\alpha\beta\gamma} h_\alpha(\xi) h_\beta(\eta) h_\gamma(\zeta)$$

Gradient on an element:

$$\nabla f(\mathbf{x}(\xi, \eta, \zeta)) = \sum_{i=1}^3 \hat{\mathbf{x}}_i \sum_{\alpha=0}^n \sum_{\beta=0}^n \sum_{\gamma=0}^n f^{\alpha\beta\gamma} [h'_\alpha(\xi) h_\beta(\eta) h_\gamma(\zeta) \partial_i \xi + h_\alpha(\xi) h'_\beta(\eta) h_\gamma(\zeta) \partial_i \eta + h_\alpha(\xi) h_\beta(\eta) h'_\gamma(\zeta) \partial_i \zeta]$$





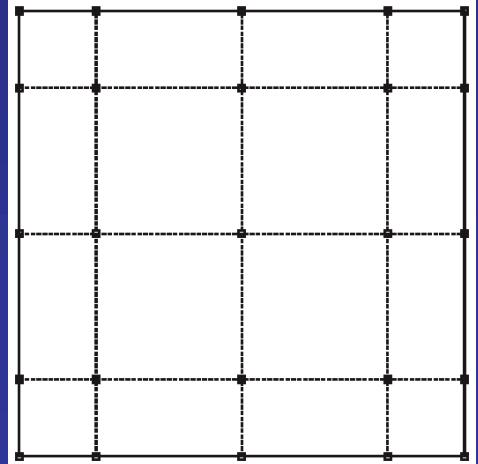
# Integration

Integration of functions over an element based upon GLL quadrature:

$$\int_{\Omega_e} f(\mathbf{x}) d^3\mathbf{x} = \int_{-1}^1 \int_{-1}^1 \int_{-1}^1 f(\mathbf{x}(\xi, \eta, \zeta)) J(\xi, \eta, \zeta) d\xi d\eta d\zeta = \sum_{\alpha=0}^n \sum_{\beta=0}^n \sum_{\gamma=0}^n \omega_\alpha \omega_\beta \omega_\gamma f^{\alpha\beta\gamma} J^{\alpha\beta\gamma}$$

- Integrations are pulled back to the reference cube
- In the SEM one uses:
  - interpolation on GLL points
  - GLL quadrature

Degree 4 GLL points:





# The Diagonal Mass Matrix

Representation of the displacement:

$$\mathbf{s}(\mathbf{x}(\xi, \eta, \zeta), t) = \sum_{i=1}^3 \hat{\mathbf{x}}_i \sum_{\sigma=0}^n \sum_{\tau=0}^n \sum_{\nu=0}^n s_i^{\sigma\tau\nu}(t) h_\sigma(\xi) h_\tau(\eta) h_\nu(\zeta)$$

Representation of the test vector:

$$\mathbf{w}(\mathbf{x}(\xi, \eta, \zeta)) = \sum_{i=j}^3 \hat{\mathbf{x}}_j \sum_{\sigma=0}^n \sum_{\tau=0}^n \sum_{\nu=0}^n w_i^{\alpha\beta\gamma} h_\alpha(\xi) h_\beta(\eta) h_\gamma(\zeta)$$

Weak form:

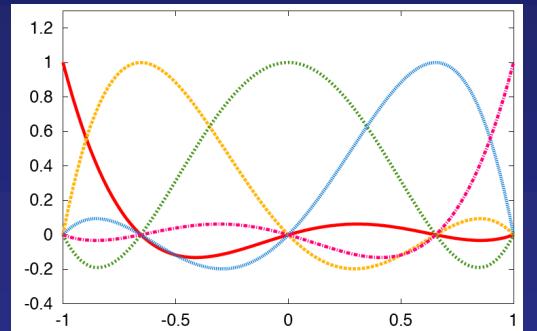
$$\int_{\Omega} \rho \mathbf{w} \cdot \partial_t^2 \mathbf{s} d^3x = - \int_{\Omega} \nabla \mathbf{w} : \mathbf{T} d^3x + \mathbf{M} : \nabla \mathbf{w}(\mathbf{x}_s) S(t)$$

Diagonal mass matrix:

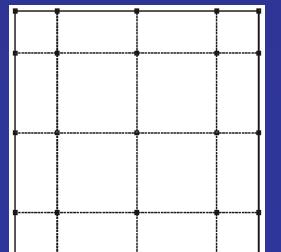
$$\int_{\Omega_e} \rho \mathbf{w} \cdot \partial_t^2 \mathbf{s} d^3x = \int_{-1}^1 \int_{-1}^1 \int_{-1}^1 \rho(\mathbf{x}(\boldsymbol{\xi})) \mathbf{w}(\mathbf{x}(\boldsymbol{\xi})) \cdot \partial_t^2 \mathbf{s}(\mathbf{x}(\boldsymbol{\xi}), t) J(\boldsymbol{\xi}) d^3\xi = \sum_{\alpha=0}^n \sum_{\beta=0}^n \sum_{\gamma=0}^n \omega_\alpha \omega_\beta \omega_\gamma J^{\alpha\beta\gamma} \rho^{\alpha\beta\gamma} \sum_{i=1}^3 w_i^{\alpha\beta\gamma} \ddot{s}_i^{\alpha\beta\gamma}$$

- Integrations are pulled back to the reference cube
- In the SEM one uses:
  - interpolation on GLL points
  - GLL quadrature

Degree 4 Lagrange polynomials:



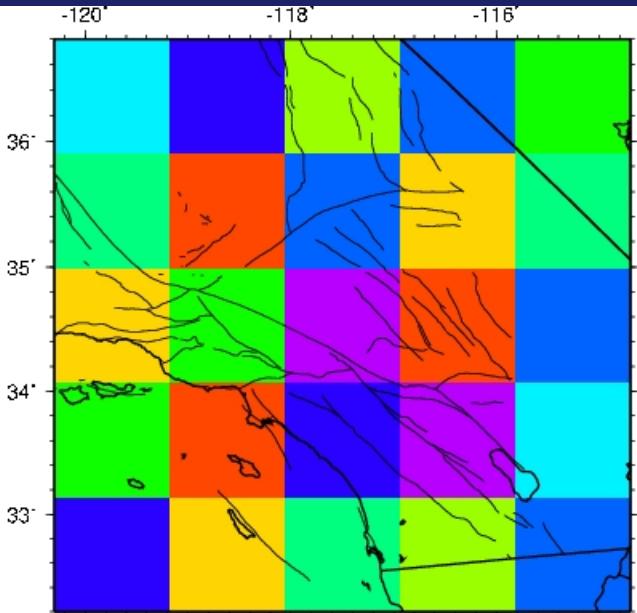
Degree 4 GLL points:



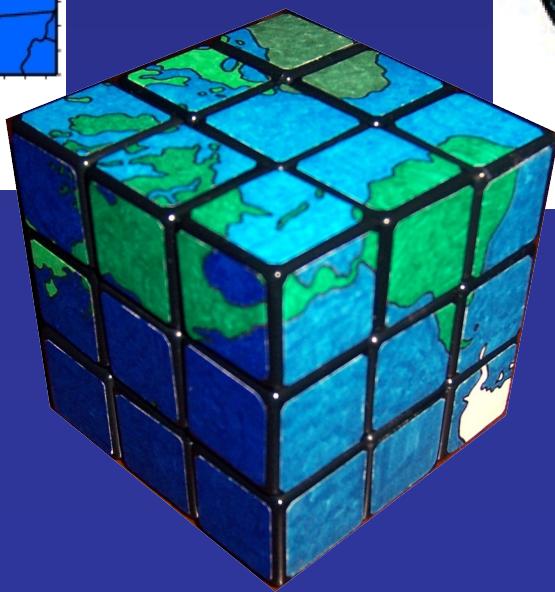


# Parallel Implementation

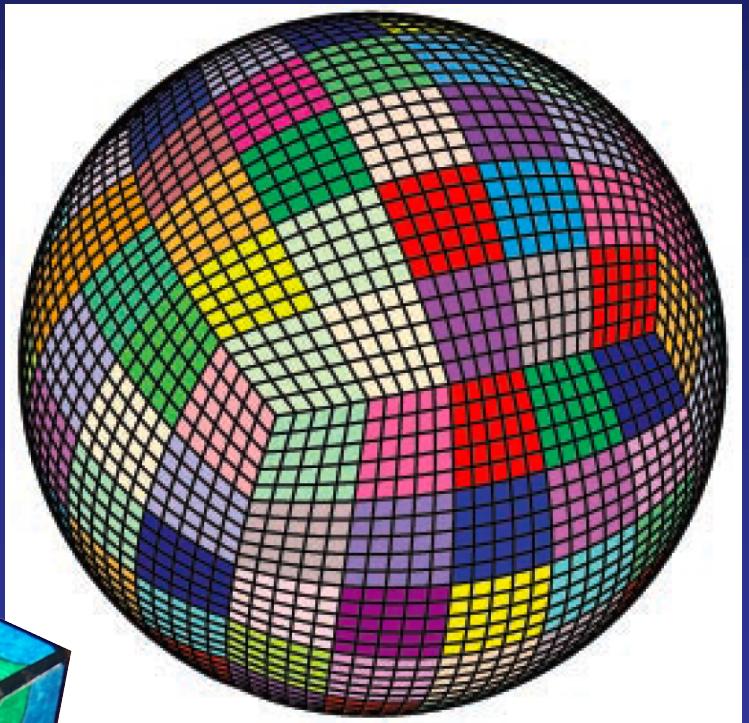
Regional mesh partitioning:



$n \times m$  mesh slices



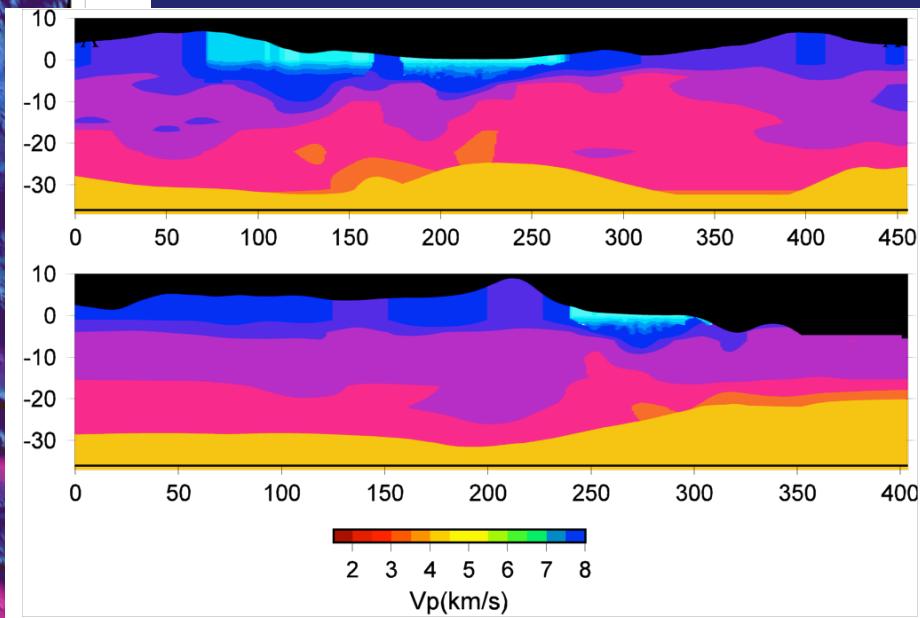
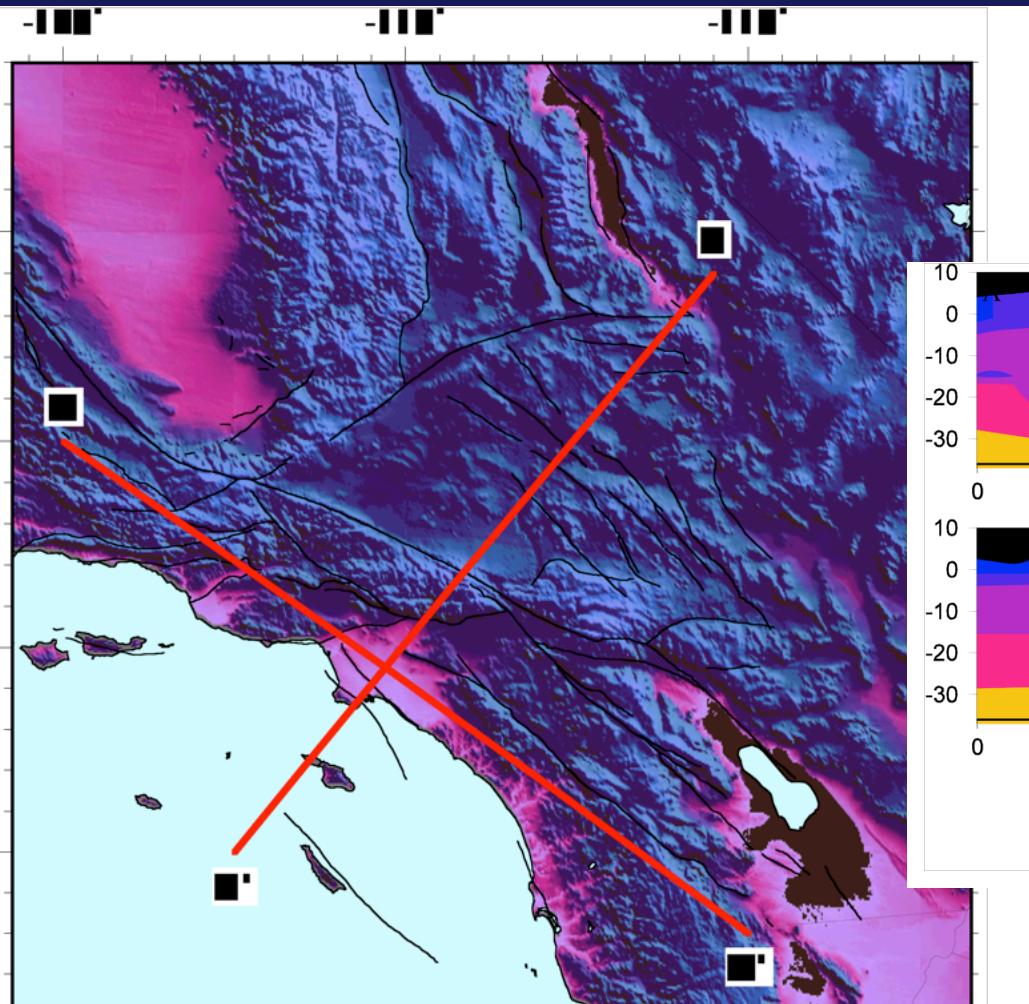
Global mesh partitioning:



Cubed Sphere:  $6 n^2$  mesh slices

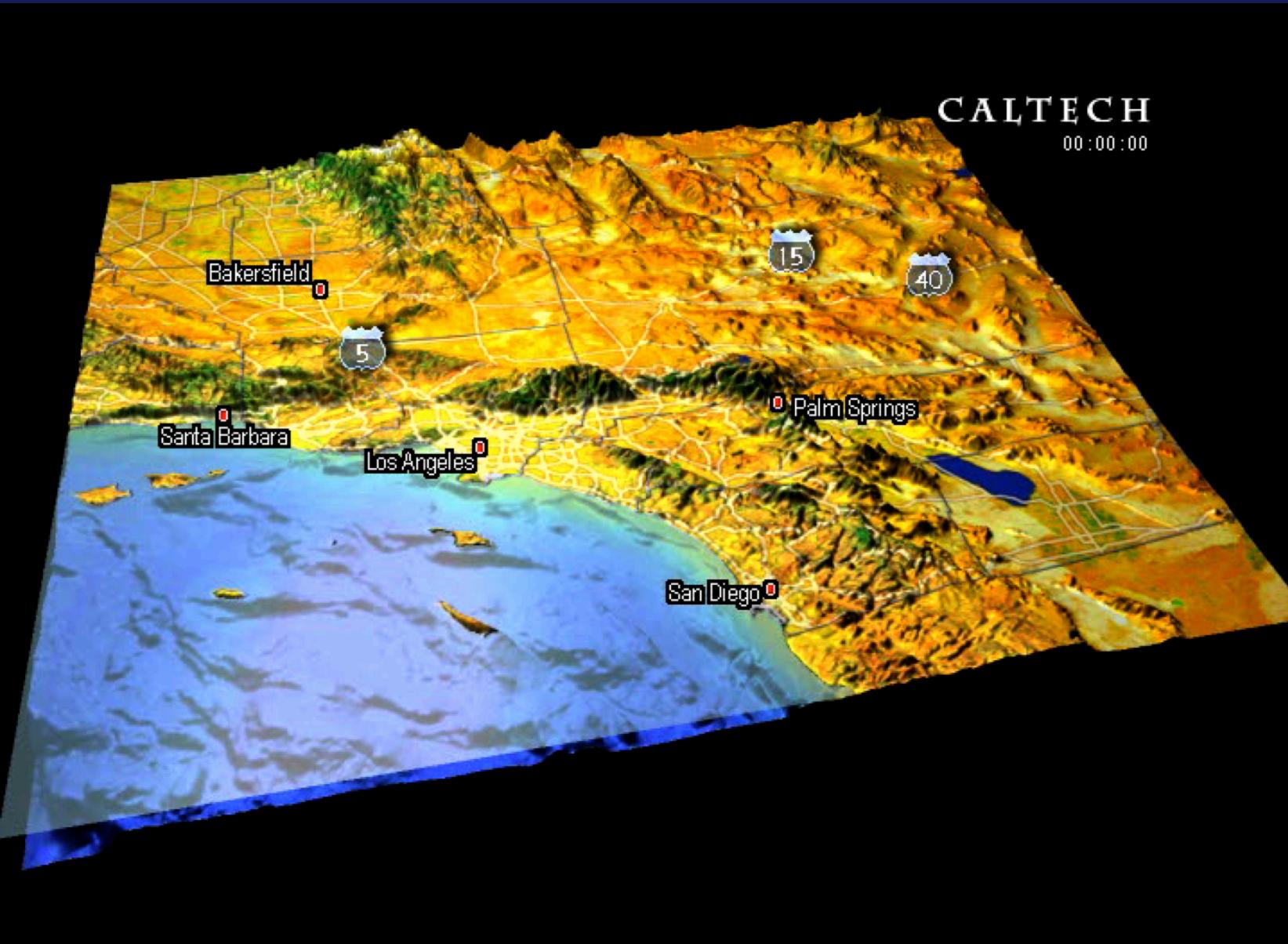


# Southern California Simulations



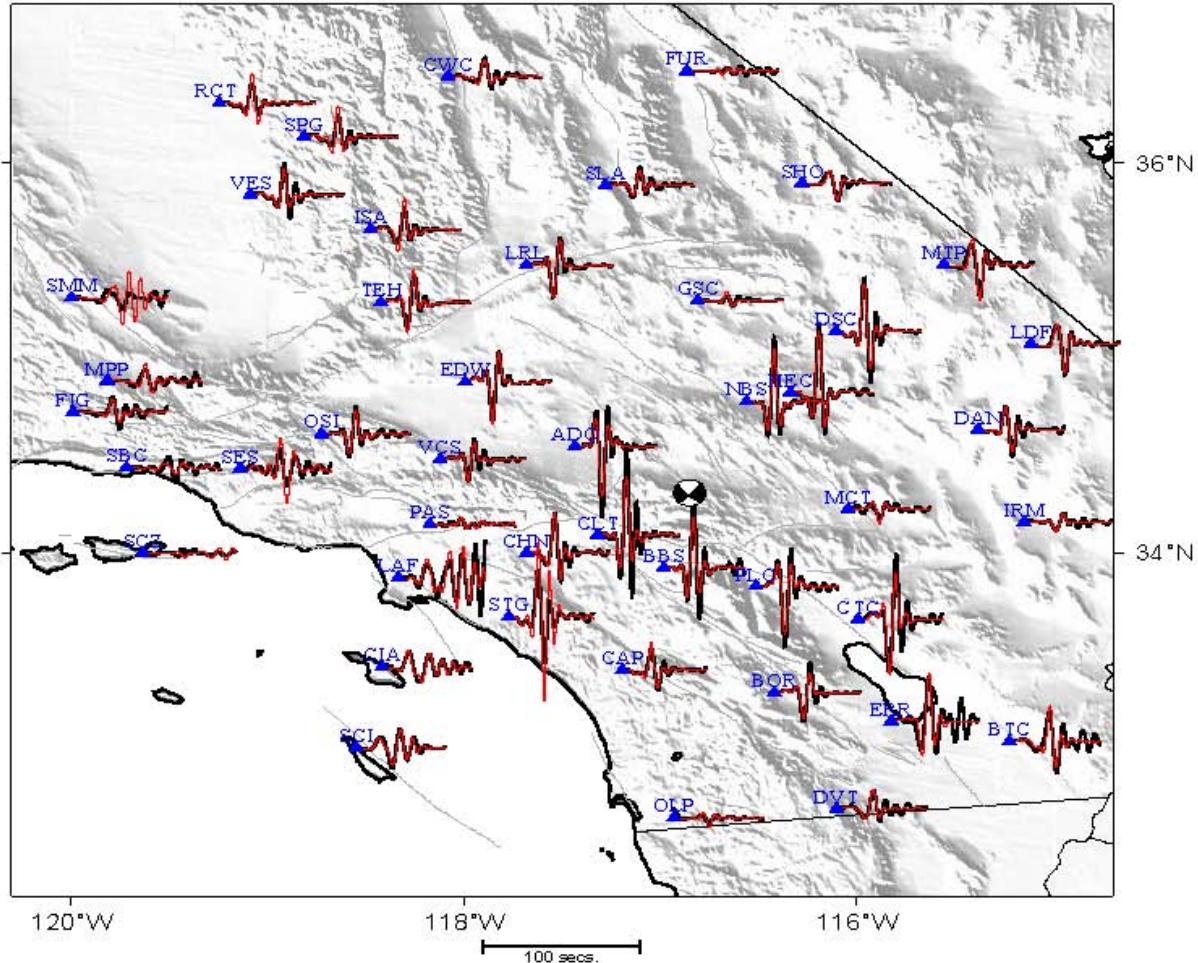


# June 12, 2005, M=5.1 Big Bear





# 3D Regional Forward Simulations



Periods > 6 s

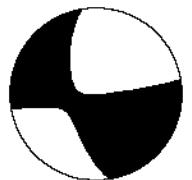
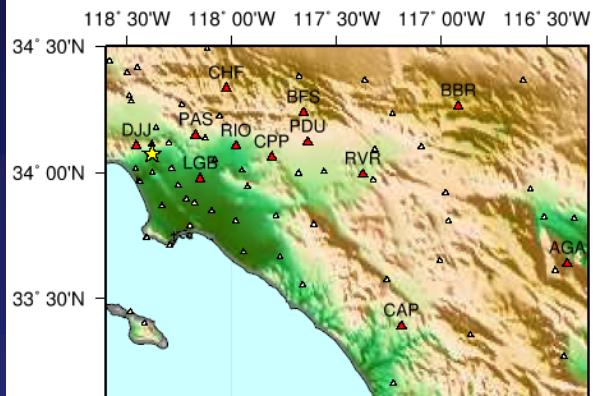
June 12, 2005, M=5.1 Big Bear

Qinya Liu

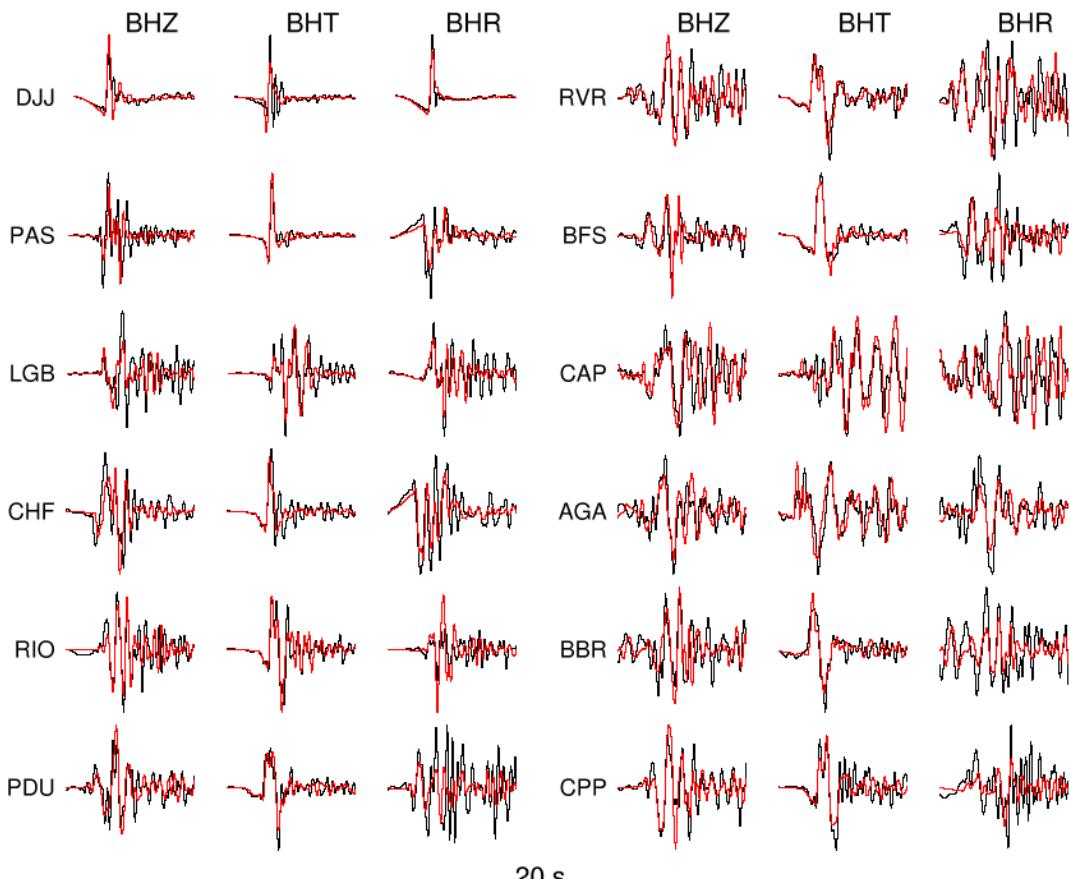


# HOLLYWOOD

September 9, 2001  $M_w = 4.2$

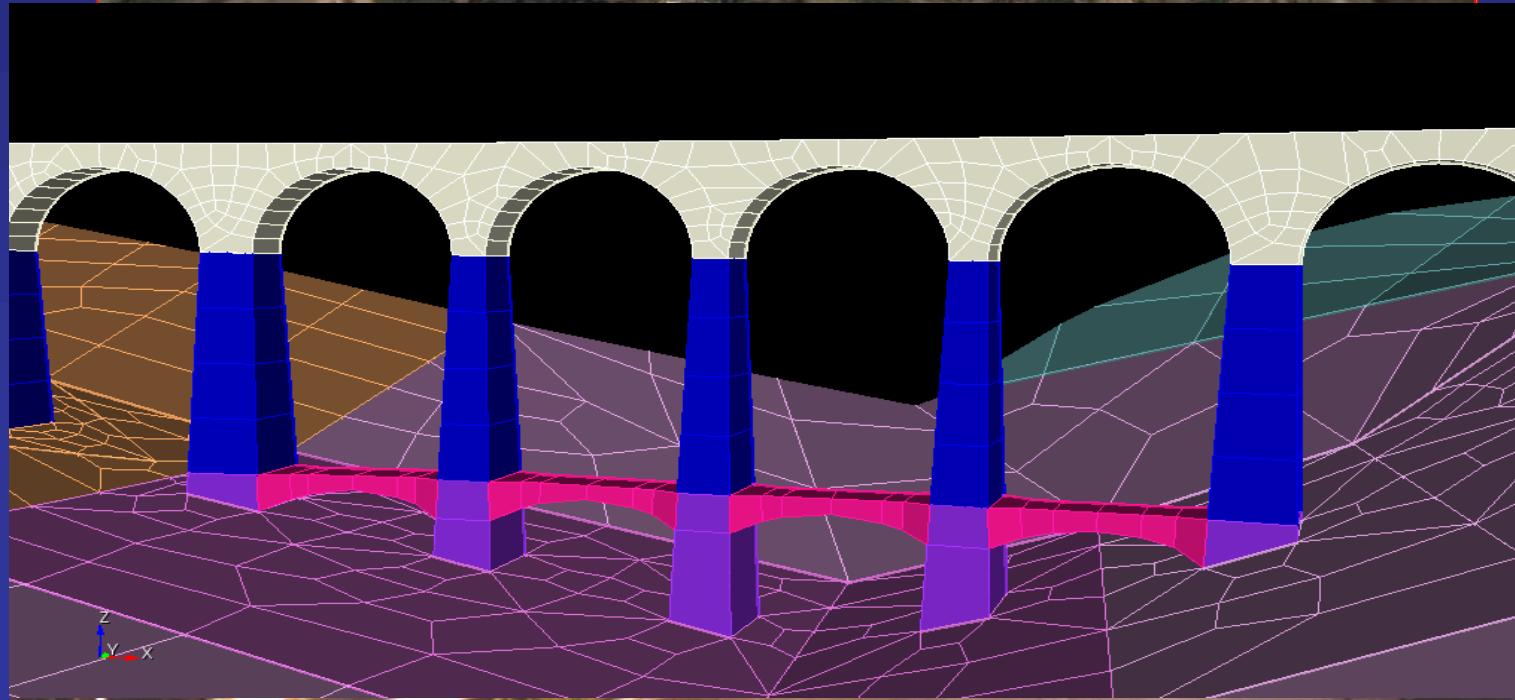


— data  
— 3D synthetics





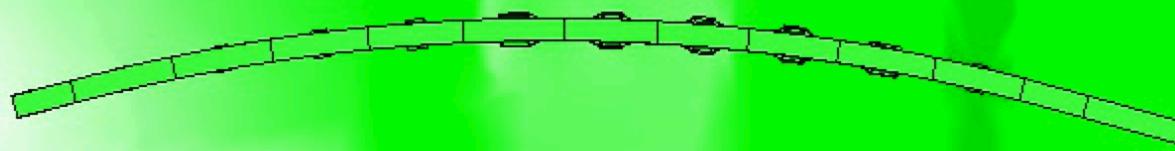
# Soil-Structure Interaction



Stupazzini 2007



# Soil-Structure Interaction (0.33 Hz)

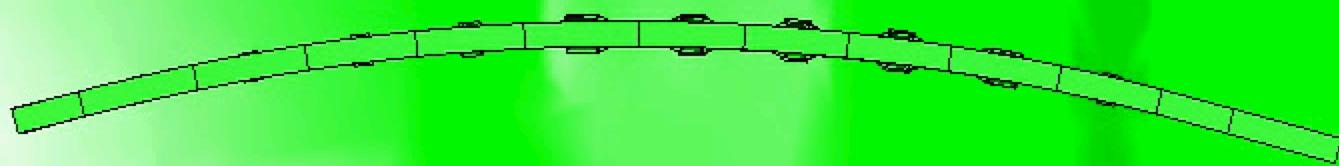


y  
z x

step 0.01  
Display Vectors of displ, |displ| factor 7.5.  
Deformation (x12): displ of TIME ANALYSIS, step 0.01.



# Soil-Structure Interaction (1 Hz)



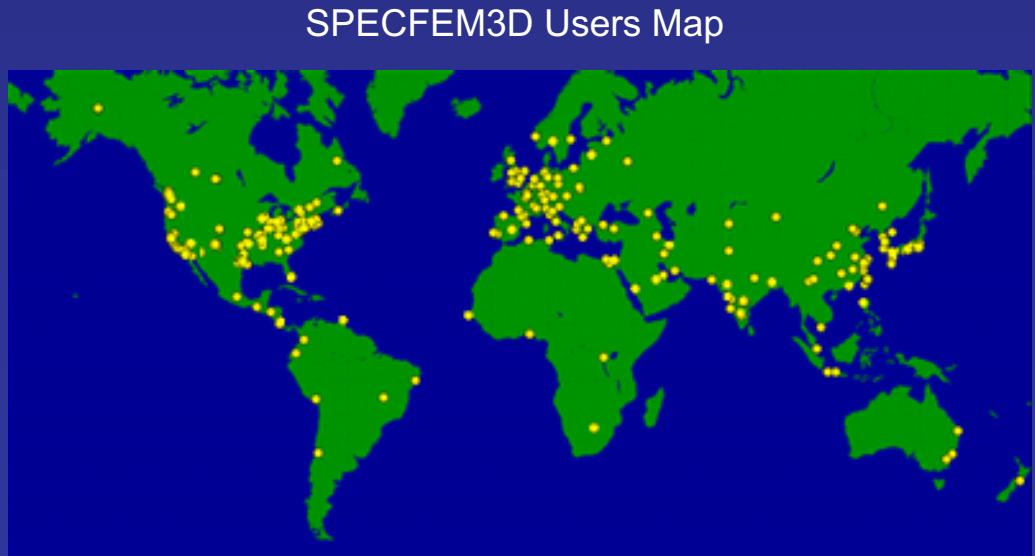
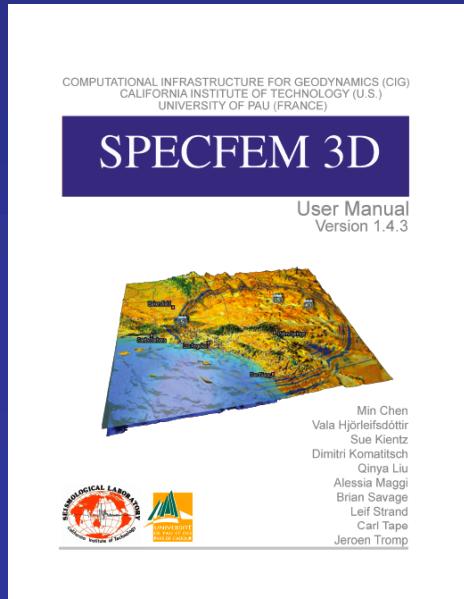
y  
z x

step 0.01  
Display Vectors of displ, |displ| factor 5.  
Deformation (x5): displ of TIME ANALYSIS, step 0.01.



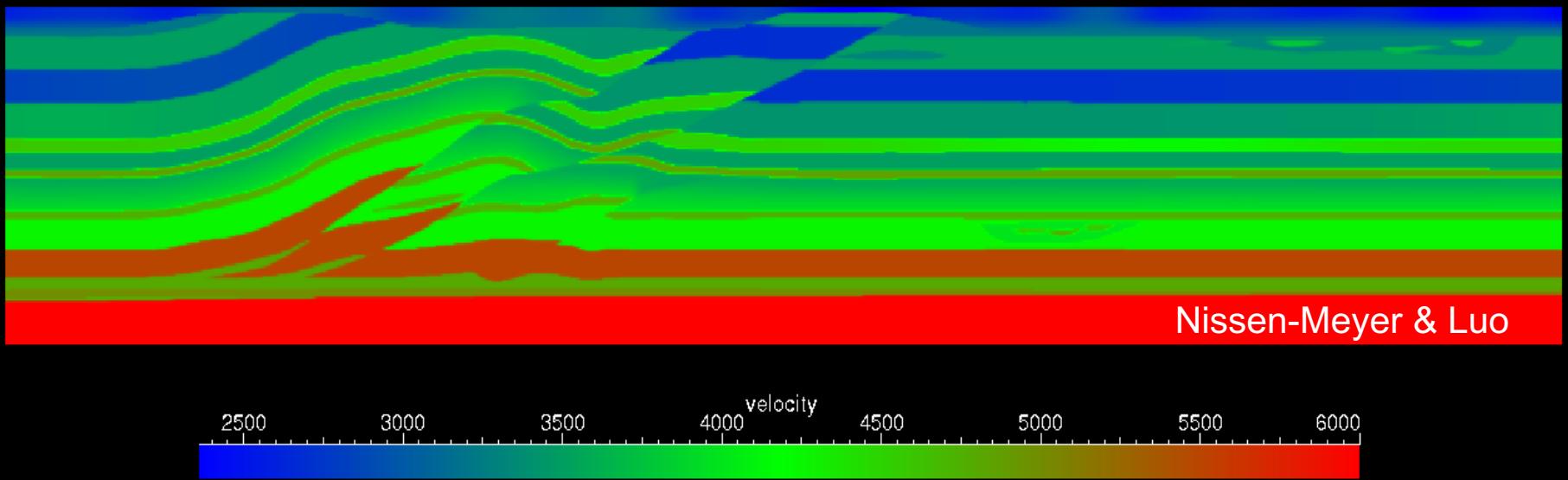
# Recent & Current Developments

- Switch to a (parallel) “GEO”CUBIT hexahedral finite-element mesher
  - Topography & bathymetry
  - Major geological interfaces
  - Basins
  - Fault surfaces
- Use ParMETIS or SCOTCH for mesh partitioning & load-balancing
- 2D mesher & solver are ready (SPECFEM2D)
- Currently developing SPECFEM3D solver (elastic & poroelastic)
- Add dynamic rupture capabilities



# Model Construction & Meshing

2D SEG model



# Model Construction & Meshing

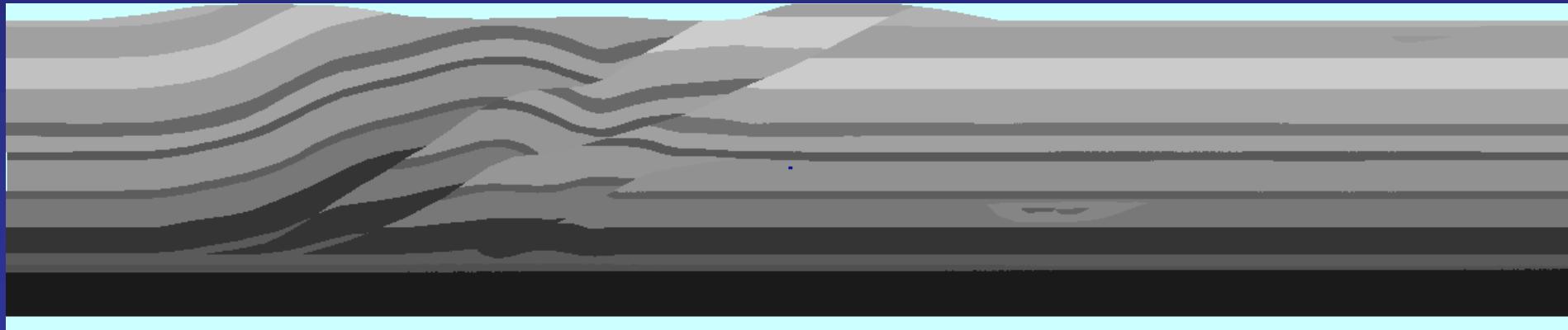
2D SEG mesh





# SEM Simulation

2D SEG model (deep explosive source)

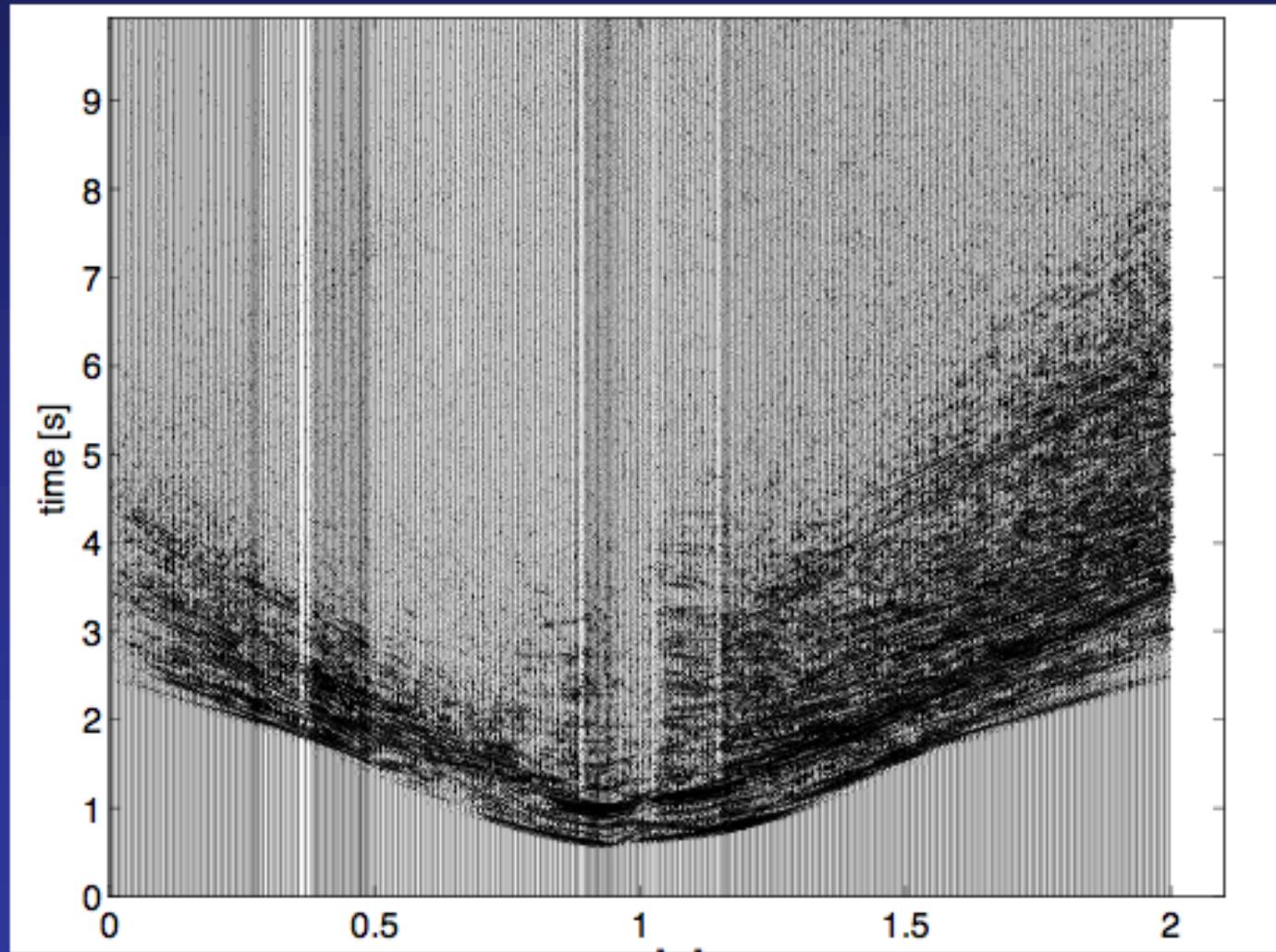


Komatitsch & Le Goff

Nissen-Meyer & Luo



# SEM Seismograms



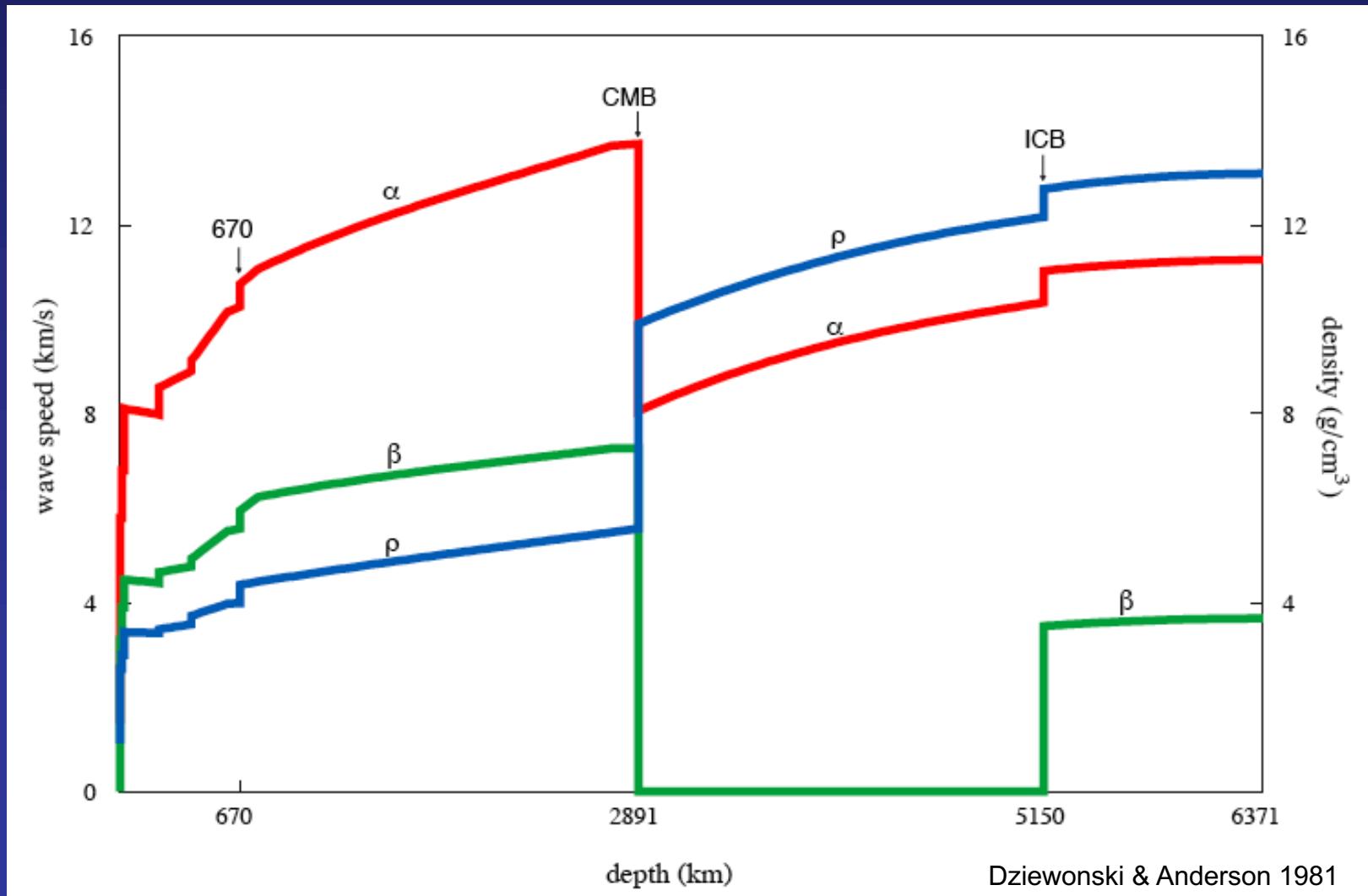
Komatitsch & Le Goff

Nissen-Meyer & Luo



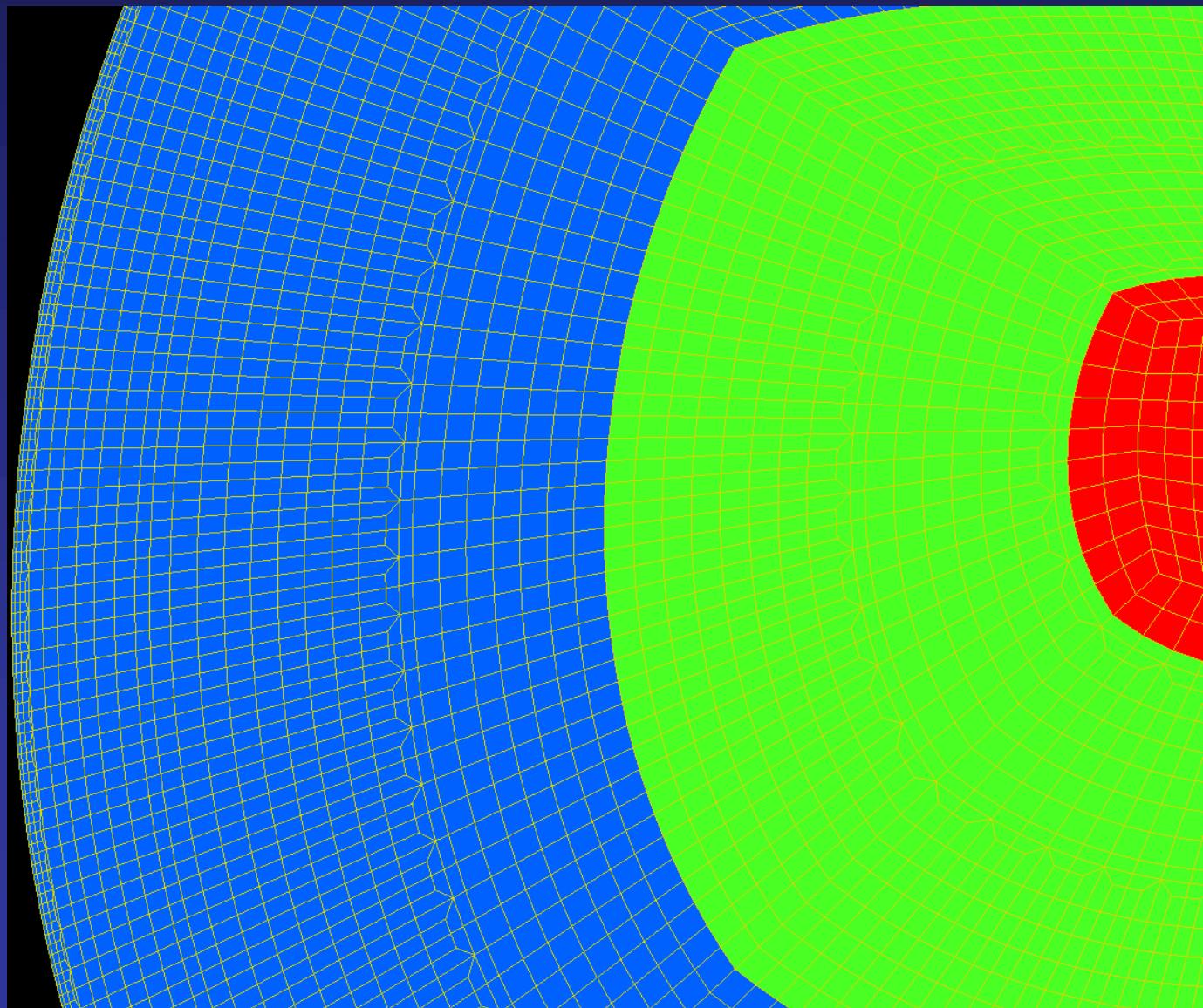
# Global Simulations

## PREM benchmarks





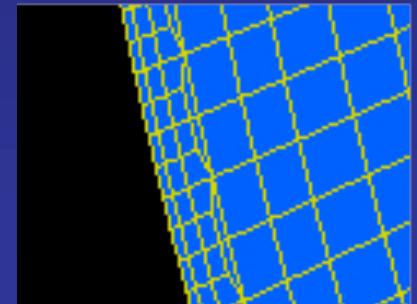
# New V4.0 Mesh



Four doublings:

- below the crust
- in the mid mantle
- two in the outer core

Note: two-layer crust



Michea & Komatitsch



# SEM Implementation of Attenuation

Anelastic, anisotropic constitutive relationship:

$$\mathbf{T}(t) = \int_{-\infty}^{\infty} \partial_t \mathbf{c}(t-t') : \nabla \mathbf{s}(t') dt'$$

Equivalent Standard Linear Solid (SLS) formulation:

$$\mathbf{T} = \mathbf{c}^U : \nabla \mathbf{s} - \sum_{\ell=1}^L \mathbf{R}^\ell$$

Memory variable equation:

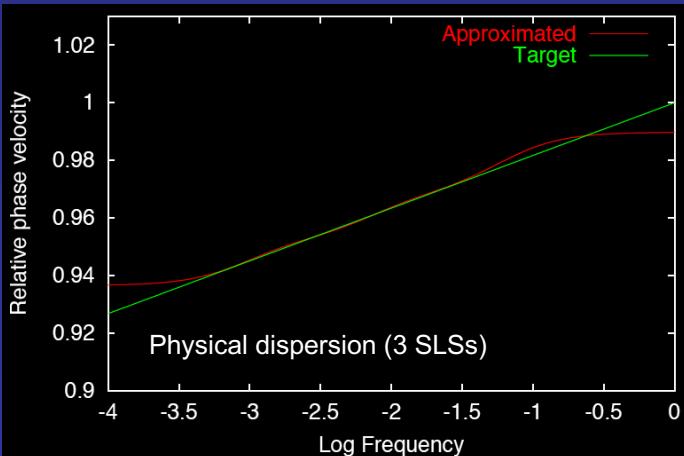
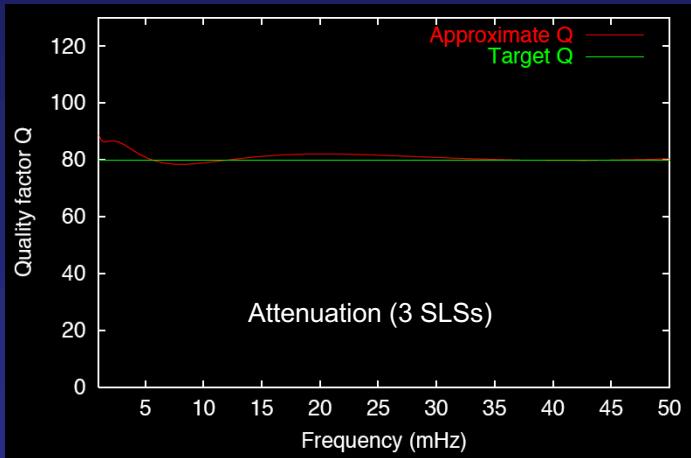
$$\partial_t \mathbf{R}^\ell = -\mathbf{R}^\ell / \tau^{\sigma\ell} + \delta \mathbf{c}^\ell : \nabla \mathbf{s} / \tau^{\sigma\ell}$$

Unrelaxed modulus:

$$c_{ijkl}^U = c_{ijkl}^R \left[ 1 - \sum_{\ell=1}^L (1 - \tau_{ijkl}^{\epsilon\ell} / \tau^{\sigma\ell}) \right]$$

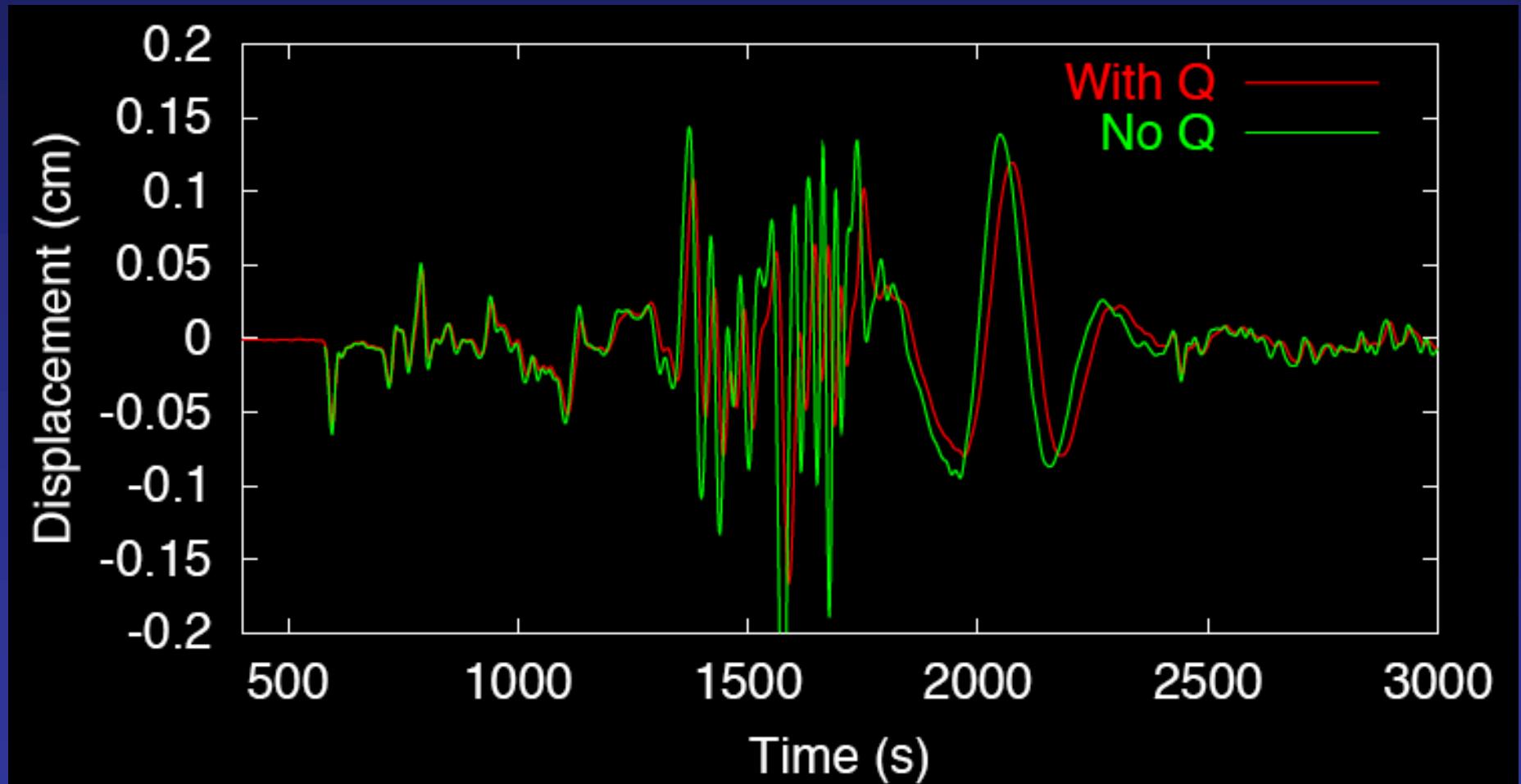
Modulus defect:

$$\delta c_{ijkl}^\ell = -c_{ijkl}^R (1 - \tau_{ijkl}^{\epsilon\ell} / \tau^{\sigma\ell})$$



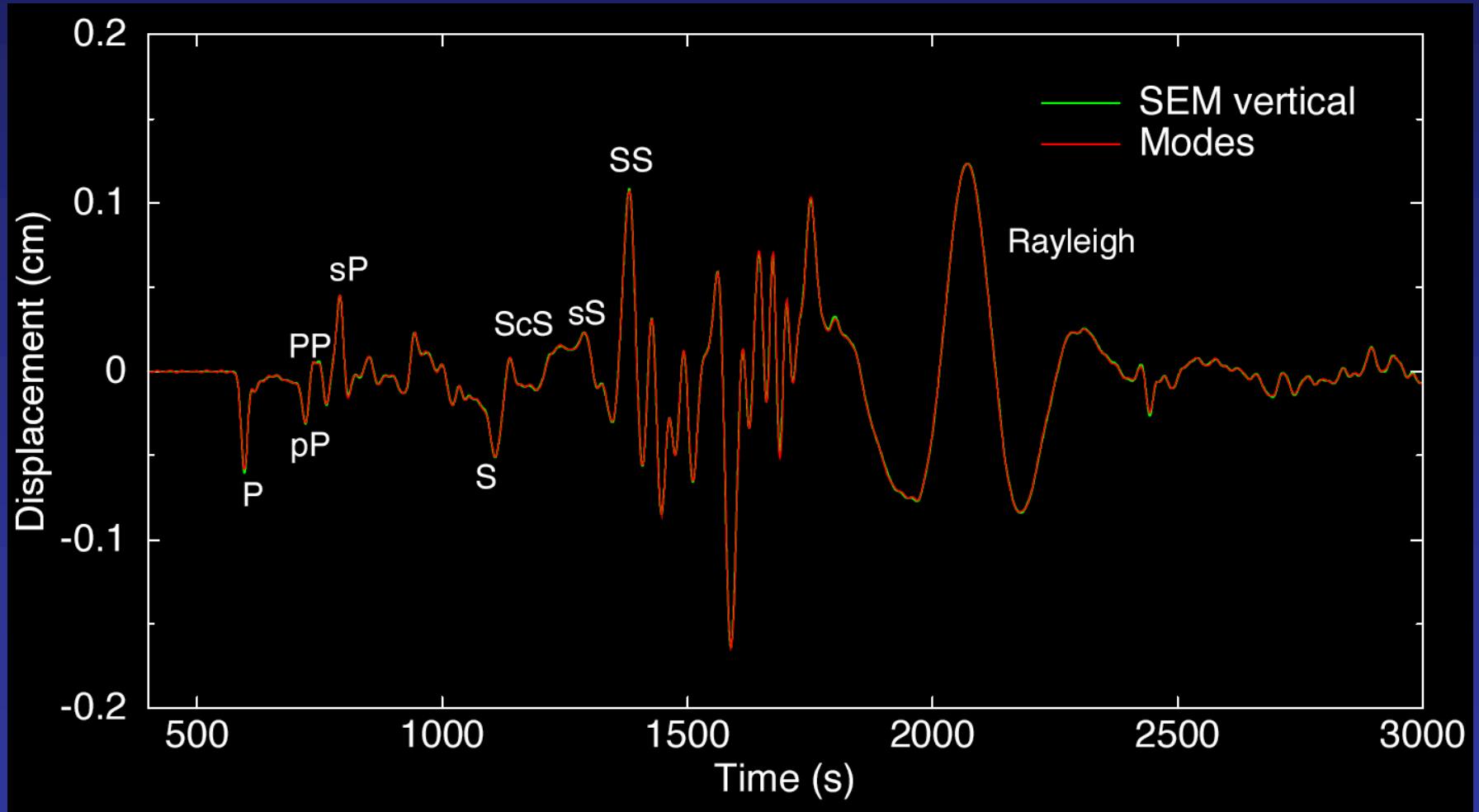


# Effect of Attenuation



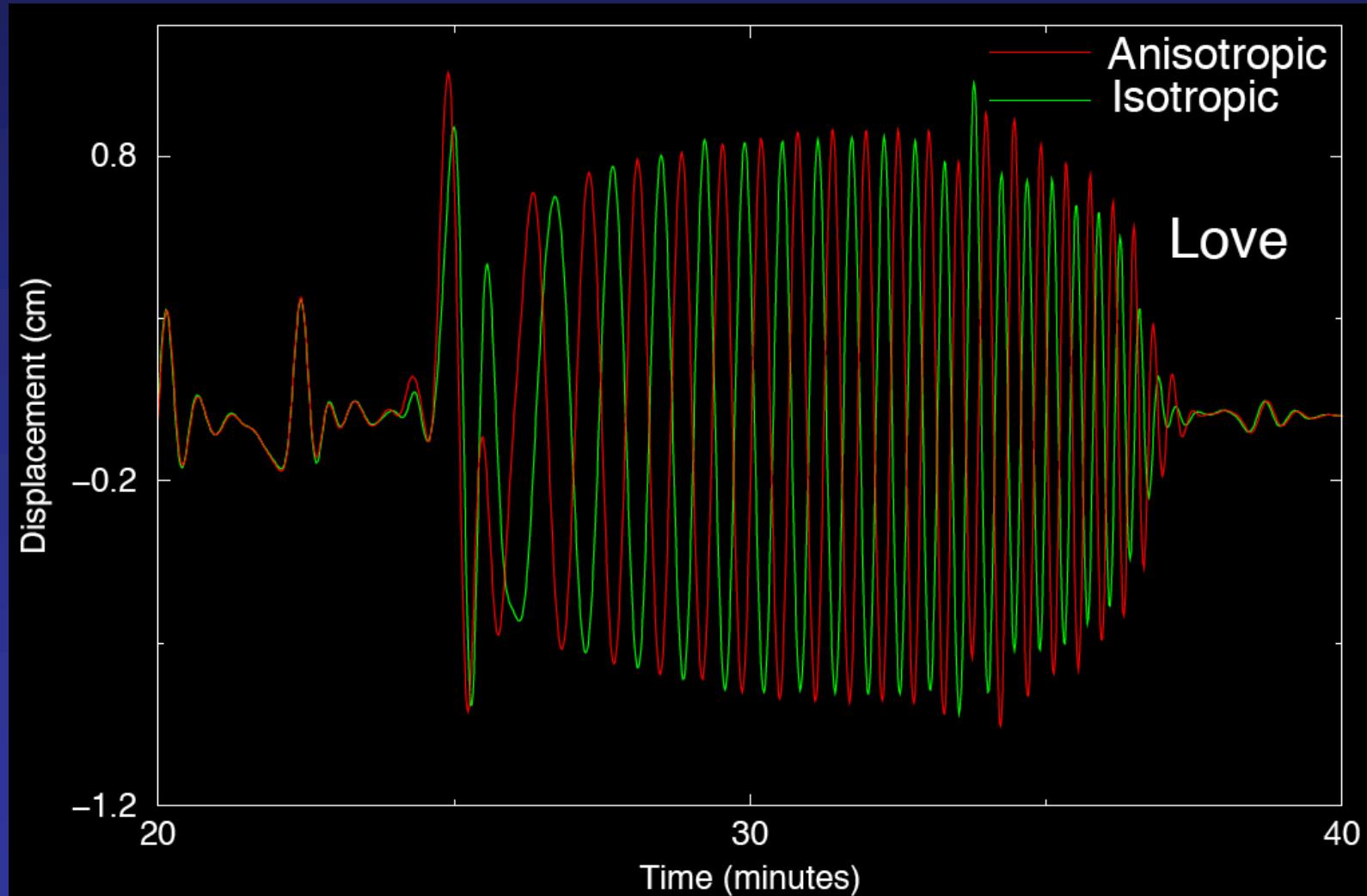


# Attenuation





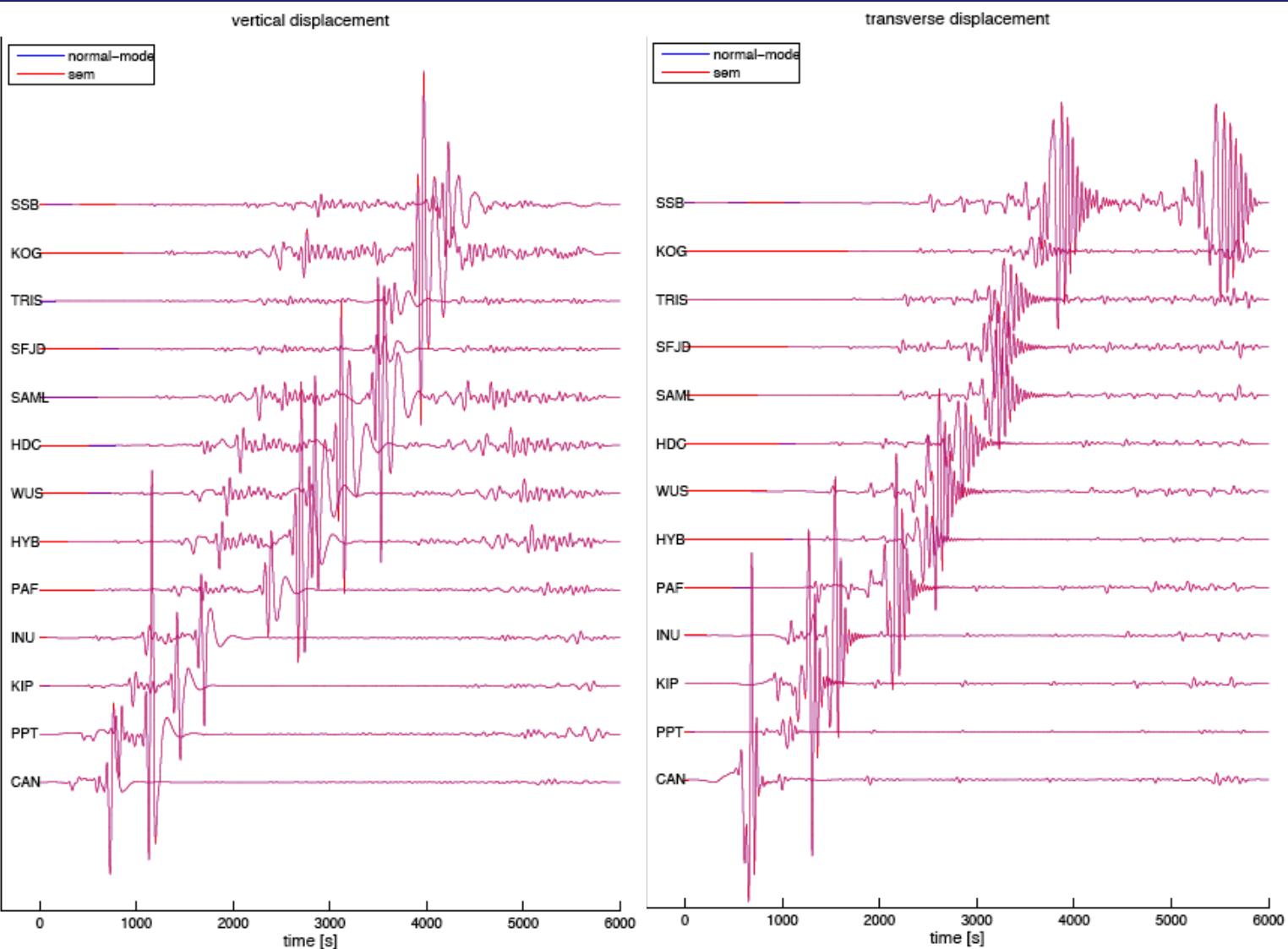
# Effect of Anisotropy



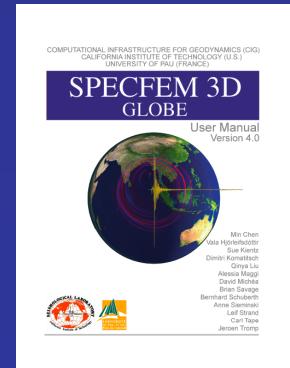


# 11/26/1999 Vanuatu shallow event

## 50 s - 500 s benchmark



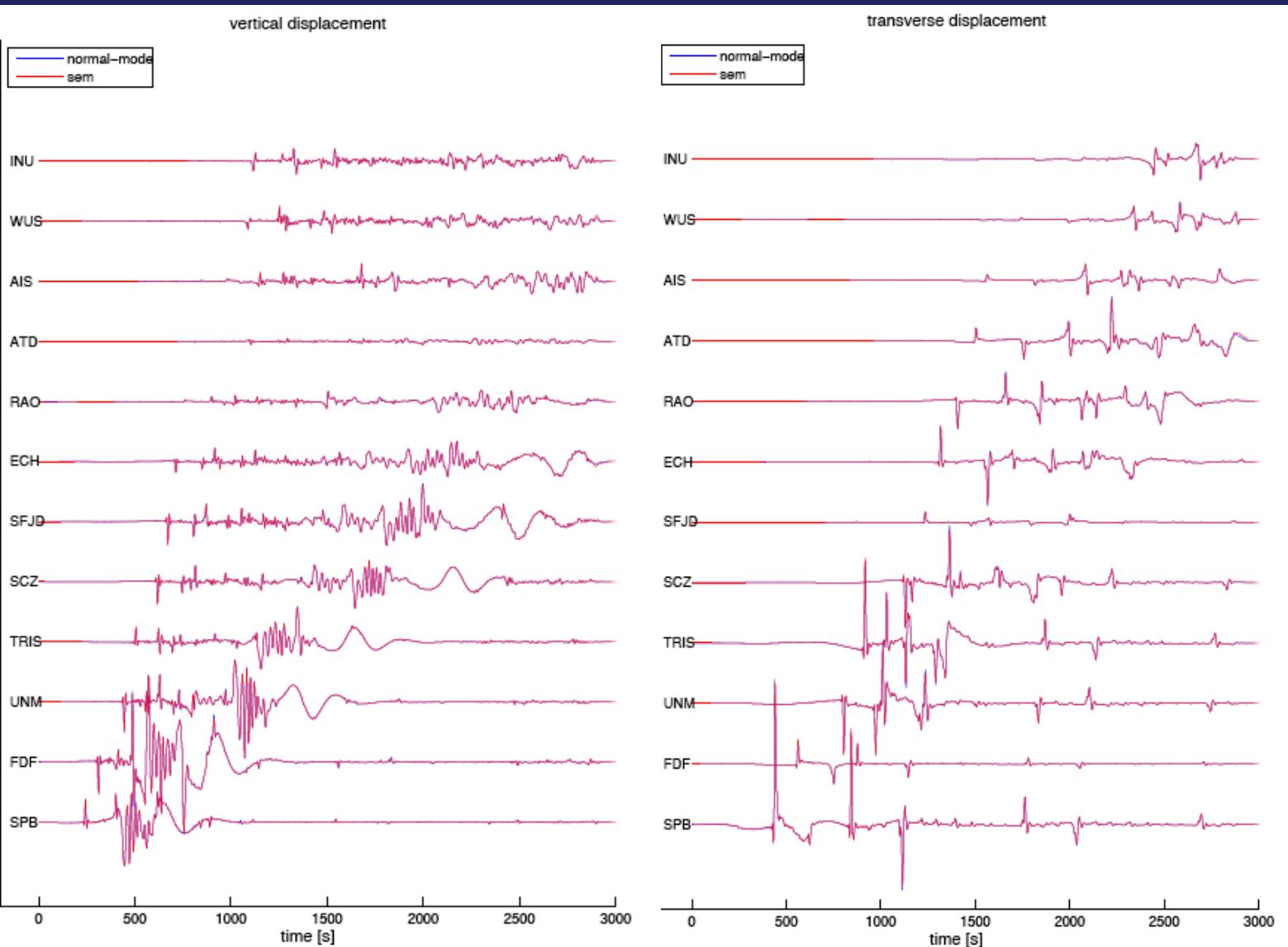
- PREM benchmarks include:
- Attenuation
  - Transverse isotropy
  - Self-gravitation (Cowling)
  - Two-layer crust





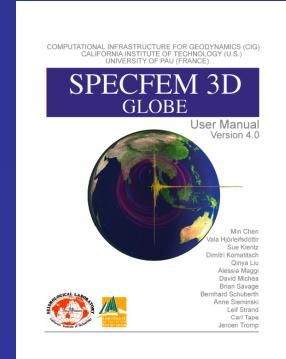
# 6/4/1994 Bolivia Deep Event

## 10 s - 500 s benchmark



PREM benchmarks include:

- Attenuation
- Transverse isotropy
- Self-gravitation (Cowling)
- Two-layer crust





# PKP Phases

15 s - 500 s

