

SPATIO-TEMPORAL MODELING OF URBAN EXTREME RAINFALL EVENTS AT HIGH RESOLUTION

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²INRAE, BioSP, Avignon

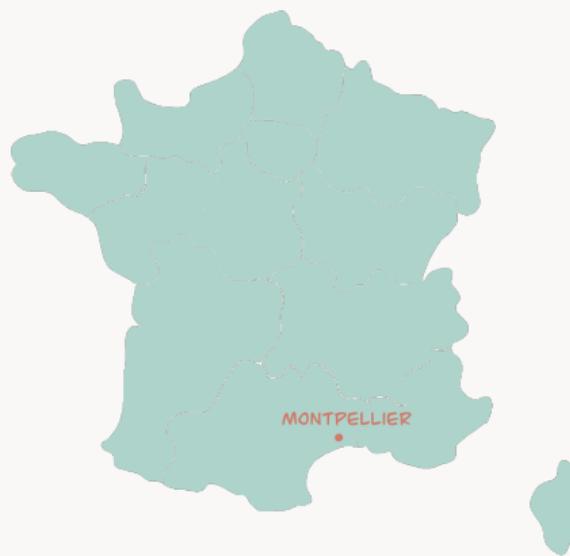


GENERAL CONTEXT - MONTPELLIER, FRANCE



Floods in Montpellier, September 2022 and August 2015 (*Midi Libre*)

- ▶ Mediterranean events, localized rainfall
- ▶ Urban area, flood risks



RAIN GAUGES NETWORK - OMSEV¹



- ▶ **Study area:** Verdanson water catchment
- ▶ **Source:** Urban observatory of HydroScience Montpellier (HSM)²
- ▶ **Time period:** [Sept.2019, Jan.2025[
- ▶ **Temporal resolution:** 5 minutes
- ▶ **Spatial resolution:** 77 m to 2259 m

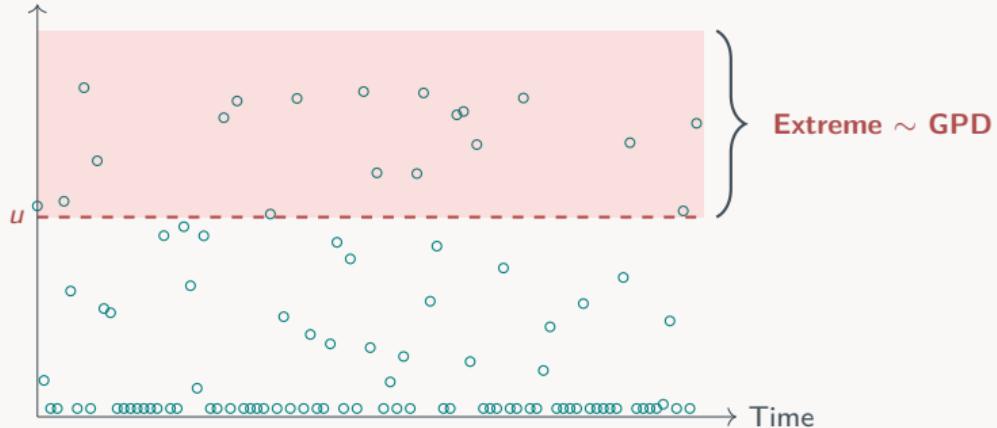
¹Observ. Montpelliérain et au Sud de l'Eau dans la Ville

²FINAUD-GUYOT et al., 2023

$$\mathcal{S} = \{\text{20 rain gauges}\} \subset \mathbb{R}^2 \text{ and } \mathcal{T} \subset \mathbb{R}_+$$

MODELING UNIVARIATE PRECIPITATION

Rainfall $X > 0$

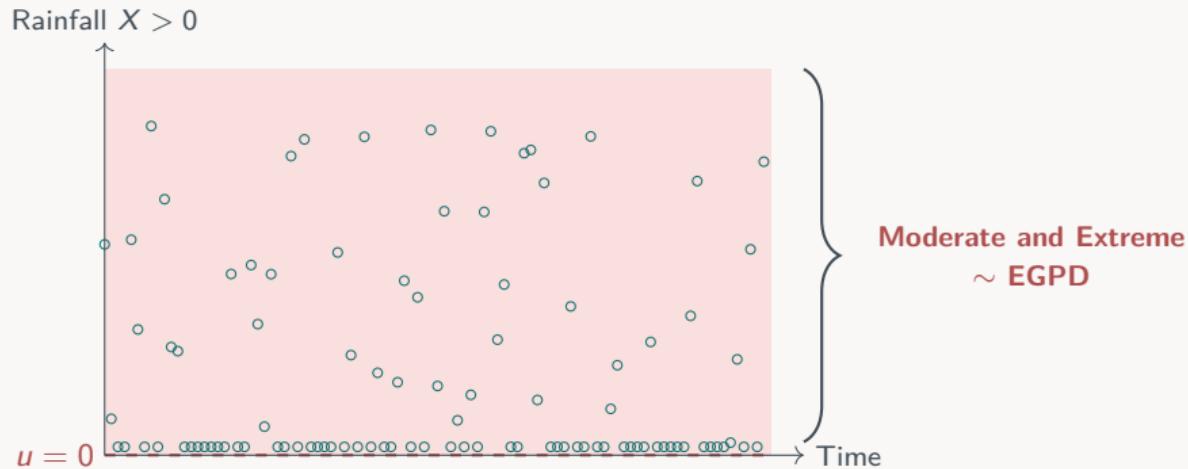


► Generalized Pareto Distribution (GPD)¹

$$X \mid X > u \sim \underbrace{H_\xi}_{\text{GPD}(\xi, \sigma, u)} \text{ with } \xi \in \mathbb{R}, \sigma > 0.$$

¹PICKANDS III, 1975

MODELING UNIVARIATE PRECIPITATION

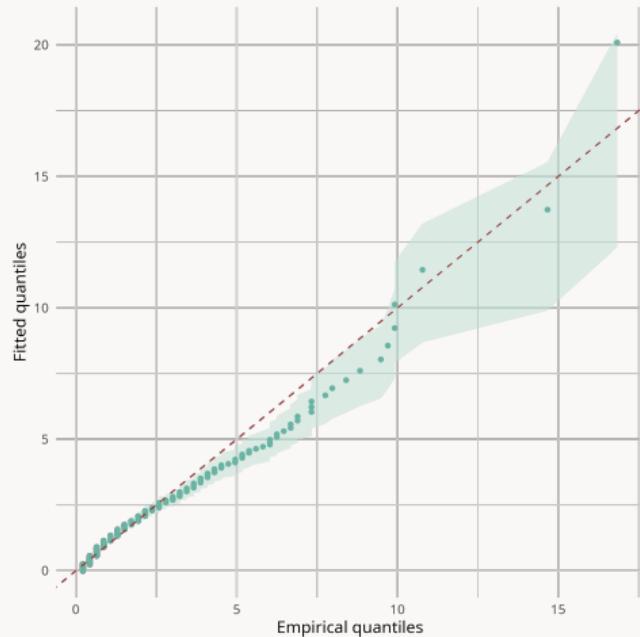
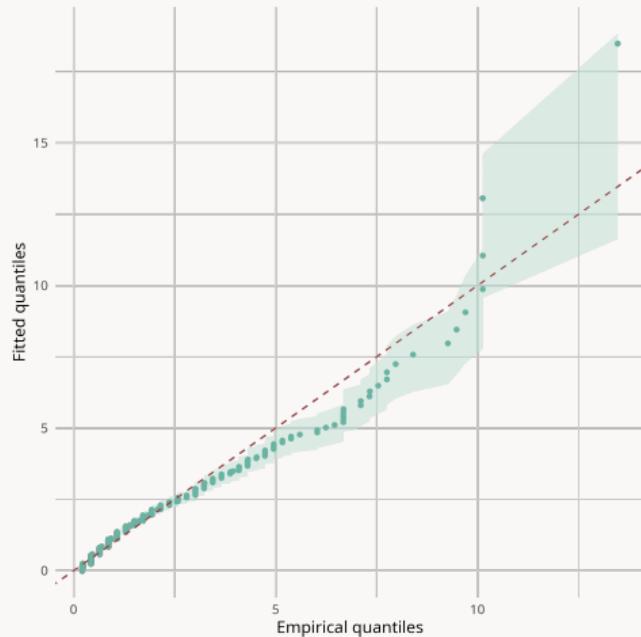


► Extended Generalized Pareto Distribution (EGPD)²

$$X \sim \underbrace{G(H_\xi)}_{\text{EGPD}(\xi, \sigma, \kappa)} \text{ with } G(x) = x^\kappa, \kappa > 0$$

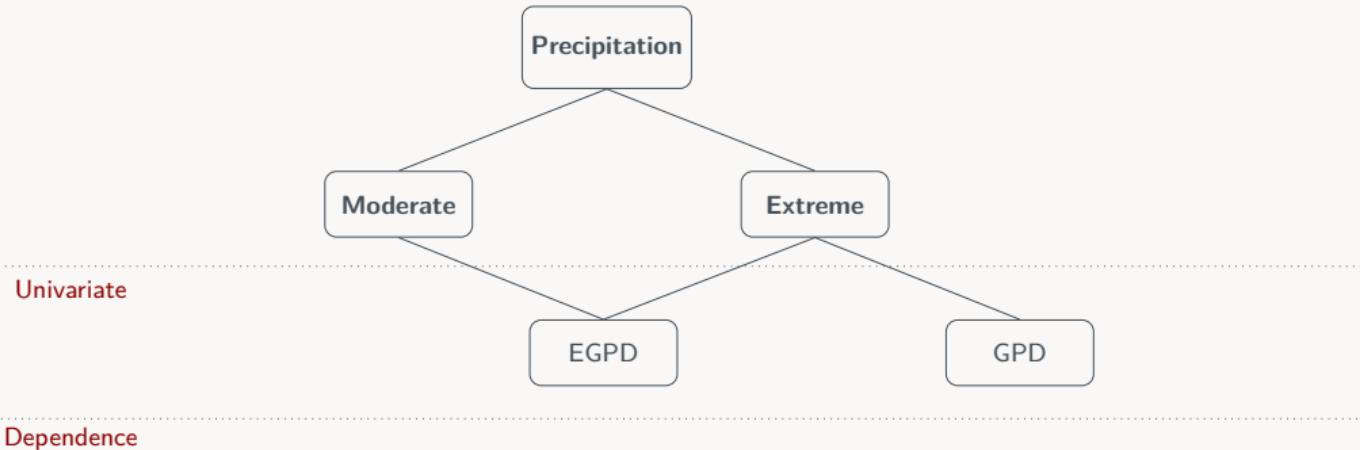
²NAVEAU et al., 2016

EGPD FITTING



EGPD fitting for two rain gauges, CEFE (left) and IEM (right) with left-censoring and 95% CI

SPATIO-TEMPORAL DEPENDENCE MODELING



SPATIO-TEMPORAL DEPENDENCE MEASURES

Rainfall random field: $\mathbf{X} = \{X_{s,t}, (s, t) \in \mathcal{S} \times \mathcal{T}\}$

Let $\Lambda_{\mathcal{S}} \subset \mathbb{R}_{+}^2$ and $\Lambda_{\mathcal{T}} \subset \mathbb{R}_{+}$ be sets of spatial and temporal lags respectively.

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Variogram (MATHERON, 1963)

$$\gamma(\mathbf{h}, \tau) = \frac{1}{2} \text{Var}(X_{s,t} - X_{s+\mathbf{h}, t+\tau})$$

- Quantifies variability
- Higher $\gamma(\mathbf{h}, \tau) \rightarrow$ weaker dependence

$$\mathbf{h} \in \Lambda_{\mathcal{S}}, \tau \in \Lambda_{\mathcal{T}}$$

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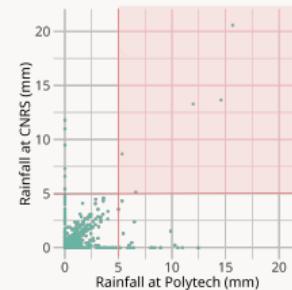
- Quantifies variability
- Higher $\gamma(\mathbf{h}, \tau) \rightarrow$ weaker dependence

$\mathbf{h} \in \Lambda_{\mathcal{S}}, \tau \in \Lambda_{\mathcal{T}}$, $X_{s,t}^{*}$ uniform margins.

Extremogram (DAVIS and MIKOSCH, 2009)

$$\chi(\mathbf{h}, \tau) = \lim_{q \rightarrow 1} \mathbb{P}(X_{s,t}^{*} > q \mid X_{s+\mathbf{h}, t+\tau}^{*} > q)$$

- Measures tail dependence
- Higher $\chi(\mathbf{h}, \tau) \rightarrow$ stronger dependence

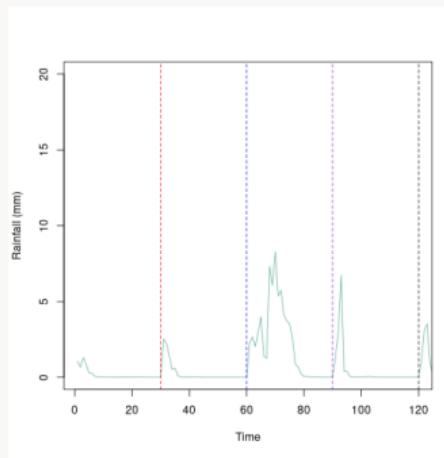


Definition (de FONDEVILLE and DAVISON, 2018)

For all $s \in S$ and $t \in T$, a risk function $r(\mathbf{X}) = X_{s_0, t_0}$,

$$u^{-1} X_{s,t} \mid X_{s_0,t_0} > u \xrightarrow{d} Y_{s,t} \quad \text{with} \quad Y_{s,t} = R e^{W_{s,t} - W_{s_0,t_0} - \gamma(s-s_0, t-t_0)},$$

where (s_0, t_0) is a space-time location, u is a high threshold, $R \sim \text{Pareto}(1)$, $W_{s,t}$ is a Gaussian process.



Random simulation

Spatio-temporal extremogram with a Brown-Resnick dependence

Let $\mathbf{h} \in \Lambda_S$ and $\tau \in \Lambda_T$. We have

$$\chi(\mathbf{h}, \tau) = 2 \left(1 - \phi \left(\sqrt{\frac{1}{2} \gamma(\mathbf{h}, \tau)} \right) \right)$$

with ϕ the std normal c.d.f. and γ the variogram of \mathbf{W} .

Spatio-temporal extremogram with a Brown-Resnick dependence

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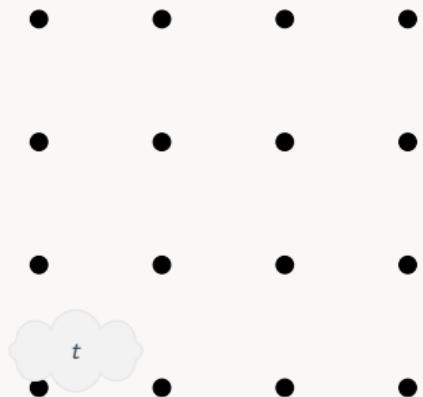
Separable model: Fractional Brownian motion with additive separability.

$$\frac{\gamma(\mathbf{h}, \tau)}{2} = \beta_1 \|\mathbf{h}\|^{\alpha_1} + \beta_2 |\tau|^{\alpha_2}$$

with $0 < \alpha_1, \alpha_2 \leq 2$, $\beta_1, \beta_2 > 0$.

DEPENDENCE MODELING

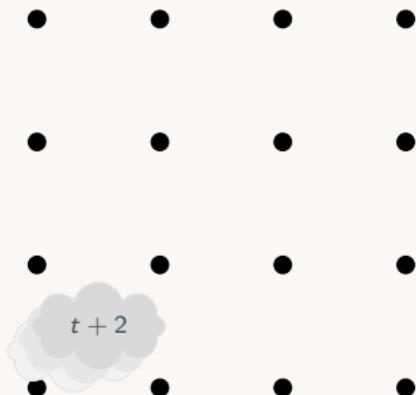
Separable model



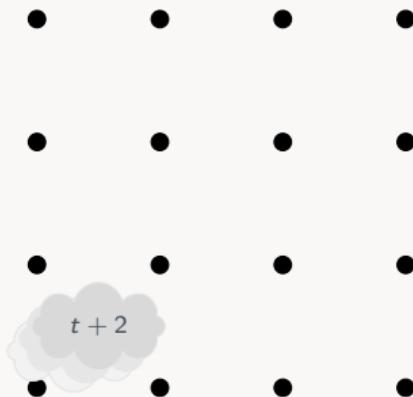
Separable model



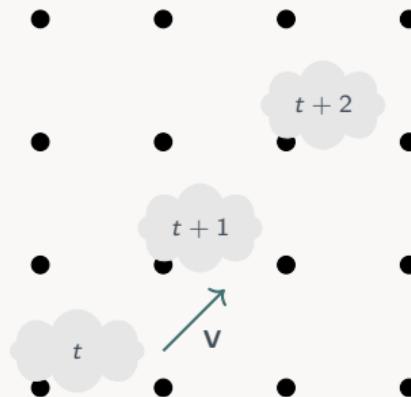
Separable model



Separable model



Non-separable model



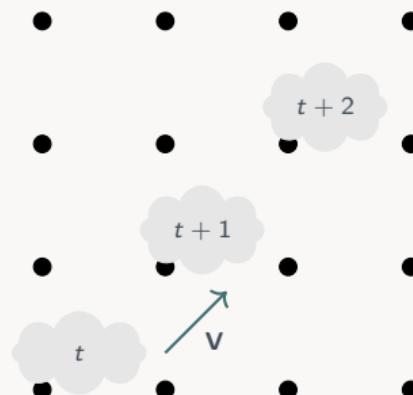
Towards more realistic modeling: introduce advection \mathbf{V} to relax separability

$$\gamma_L(\mathbf{h}, \tau) = \gamma(\mathbf{h} - \tau \mathbf{V}, \tau)$$

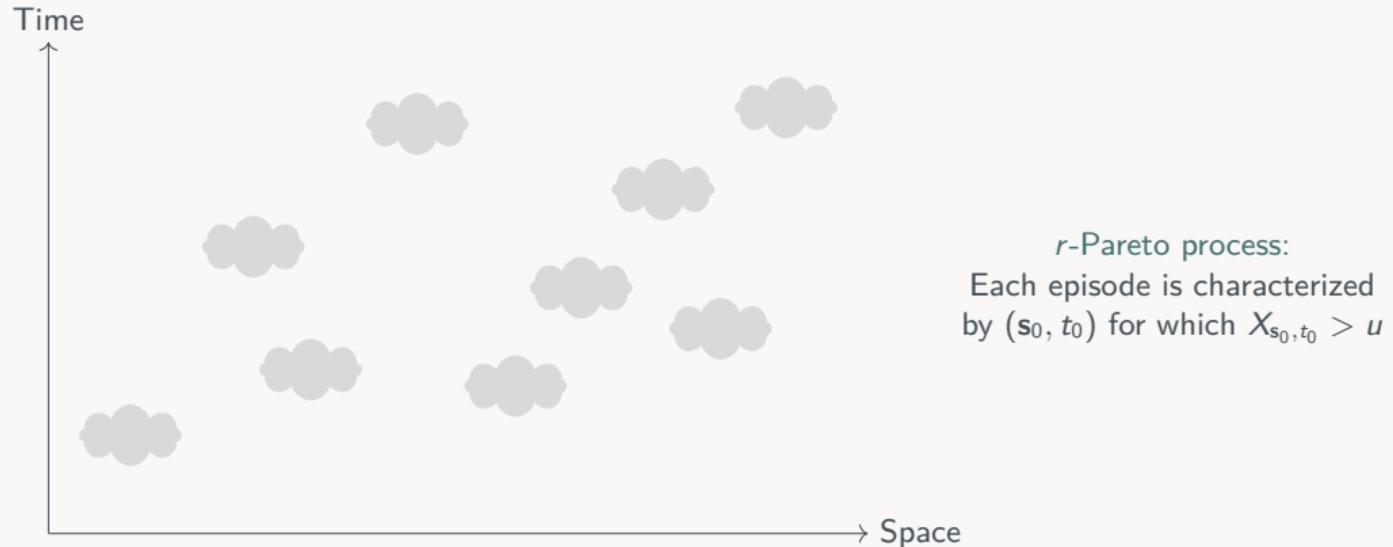
$$\Rightarrow \frac{1}{2}\gamma_L(\mathbf{h}, \tau) = \beta_1 \|\mathbf{h} - \tau \mathbf{V}\|^{\alpha_1} + \beta_2 |\tau|^{\alpha_2}$$

- ▶ Parameters: $\Theta = (\beta_1, \beta_2, \alpha_1, \alpha_2, \mathbf{V})$

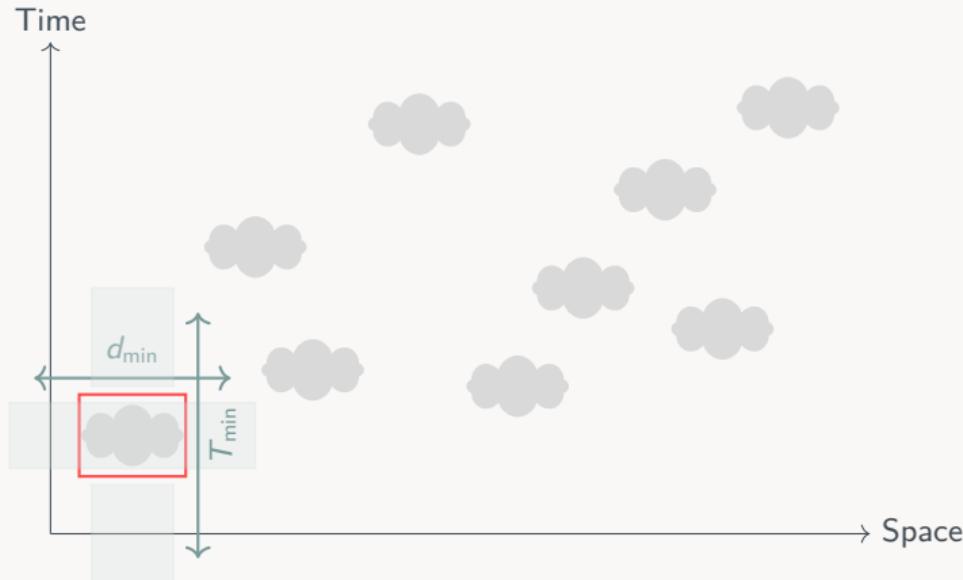
Non-separable model



ALL RAINFALL EVENTS



SELECT EPISODES



Episode selection:

Only episodes separated by

- spatial distance $\geq d_{\min}$
- temporal gap $\geq T_{\min}$

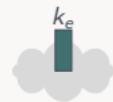
\Rightarrow reduces dependence
between **selected episodes** $\in \mathcal{E}$.

SELECT EPISODES



COUNT JOINT EXCEEDANCES PER EPISODE

Time



Space

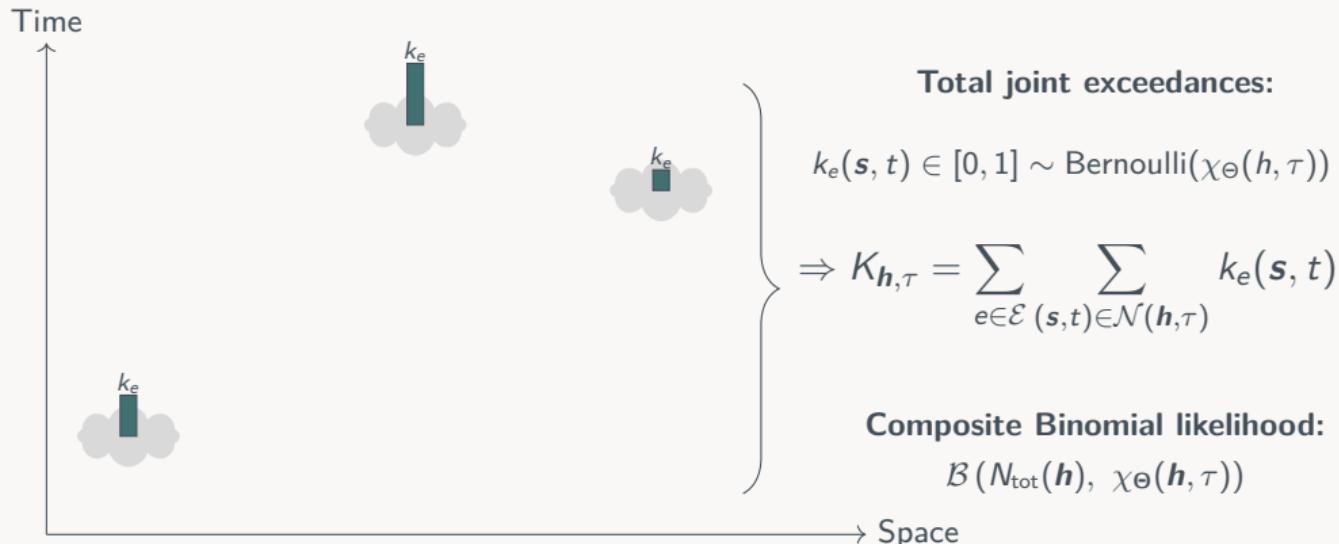
Joint exceedances

$$k_e(s, t) = \mathbb{1}_{\{X_{s_0, t_0} > u, X_{s, t} > u\}}, \\ \text{with } (s, t) \in \mathcal{N}(h, \tau)$$

$$k_e(s, t) \in [0, 1] \sim \text{Bernoulli}(\chi_\Theta(h, \tau))$$

$$\text{with } \mathcal{N}(h, \tau) = \{(s, t) \in \mathcal{S} \times \mathcal{T} \mid s - s_0 = h, t - t_0 = \tau\}$$

JOINT EXCEEDANCES



with $\mathcal{N}(h, \tau) = \{(s, t) \in \mathcal{S} \times \mathcal{T} \mid s - s_0 = h, t - t_0 = \tau\}$
and $N_{\text{tot}}(h) = |\mathcal{E}| \times \#\{s \in \mathcal{S} \mid s - s_0 = h\}$

Bernoulli contributions (large u):

$$k_e(\mathbf{s}, t) = \mathbf{1}_{\{X_{s_0, t_0} > u, X_{s, t} > u\}}, \quad (\mathbf{s}, t) \in \mathcal{N}(\mathbf{h}, \tau),$$

each treated as

$$k_e(\mathbf{s}, t) \sim \text{Bernoulli}(\chi_\Theta(\mathbf{h}, \tau)).$$

Composite Binomial likelihood:

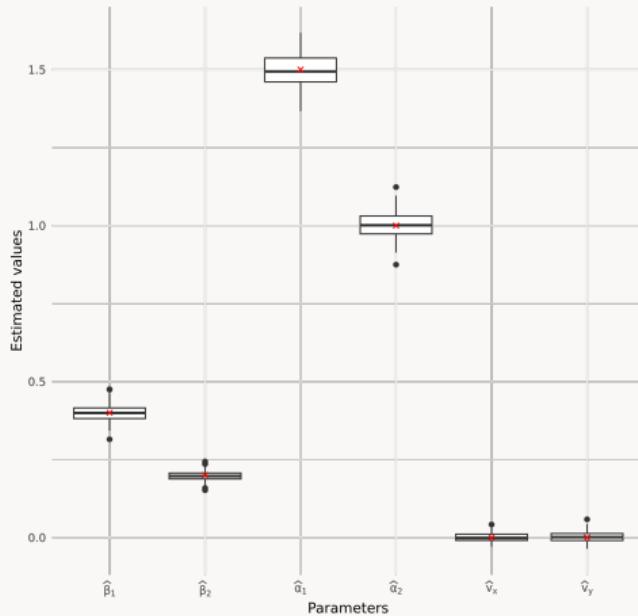
$$\mathcal{B}(N_{\text{tot}}(\mathbf{h}), \chi_\Theta(\mathbf{h}, \tau)), \quad N_{\text{tot}}(\mathbf{h}) = |\mathcal{E}| \#\{\mathbf{s} \in \mathcal{S} : \mathbf{s} - \mathbf{s}_0 = \mathbf{h}\}.$$

Composite log-likelihood

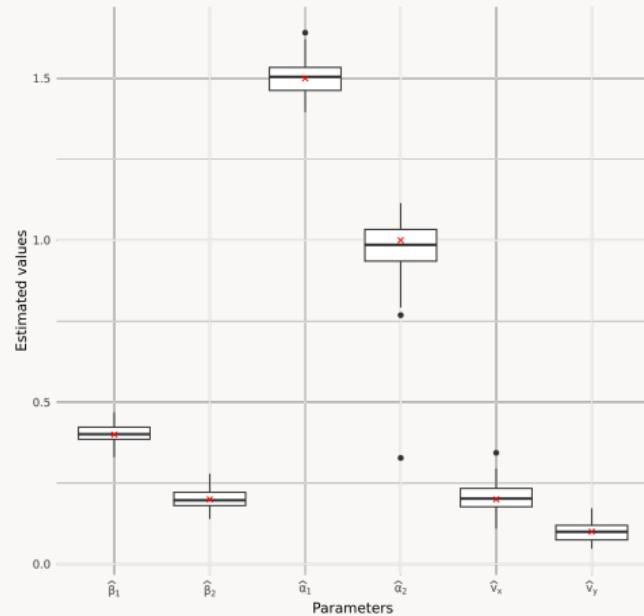
$$\ell_C(\Theta) \propto \sum_{e \in \mathcal{E}} \sum_{(\mathbf{h}, \tau) \in \Lambda_S \times \Lambda_T} \sum_{(\mathbf{s}, t) \in \mathcal{N}(\mathbf{h}, \tau)} k_e(\mathbf{s}, t) \log \chi_\Theta(\mathbf{h}, \tau) + (1 - k_e(\mathbf{s}, t)) \log(1 - \chi_\Theta(\mathbf{h}, \tau)).$$

- Optimization via maximization of $\ell_C(\Theta)$.

VALIDATION: CONSTANT ADVECTION



Without advection



With advection

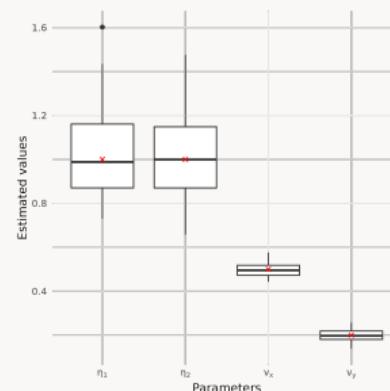
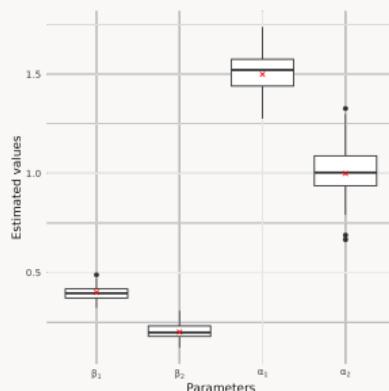
50 simulations, 500 replicates, 25 sites, 30 time steps

EPISODE-WISE ADVECTION TO GLOBAL MODEL

$$\mathbf{V}^{\text{emp}} = \begin{bmatrix} \mathbf{V}^{(1)} \\ \mathbf{V}^{(2)} \\ \vdots \\ \mathbf{V}^{(N)} \end{bmatrix} \in \mathbb{R}^{N \times 2} \xrightarrow{\text{global model}} \mathbf{V}^{\text{final}} = \eta_1 \text{sign}(\mathbf{V}^{\text{emp}}) \odot |\mathbf{V}^{\text{emp}}|^{\odot \eta_2} \in \mathbb{R}^{N \times 2}$$

unified transformation $\Rightarrow \eta$ to estimate

one component = one episode



50 simulations, 500 replicates, 25 sites, 30 time steps

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$\underbrace{\qquad\qquad\qquad}_{\text{unified transformation} \Rightarrow \eta \text{ to estimate}}$

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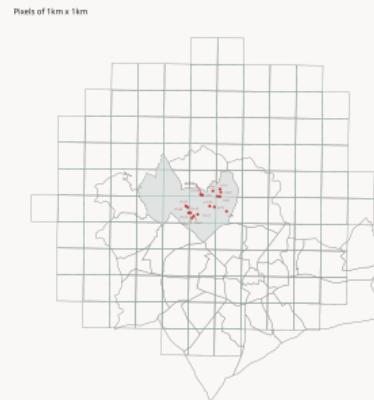
Challenge:

OMSEV data \Rightarrow limited information for advection

Approach:

Fusion COMEPHORE–OMSEV \rightarrow more reliable advection

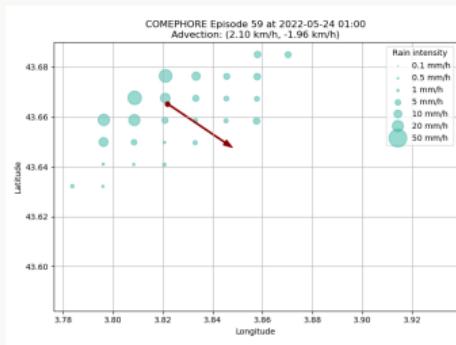
Advection classes \rightarrow class-dependent $\eta = (\eta_1, \eta_2)$



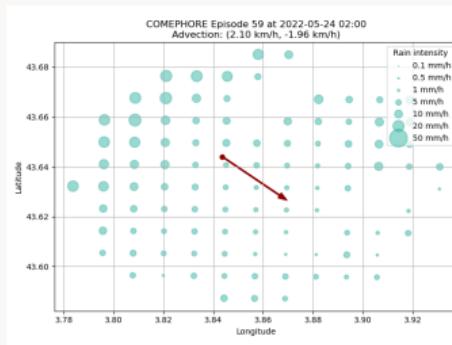
COMEPHORE pixels
Météo France

ADVECTION ESTIMATION ON REAL DATA

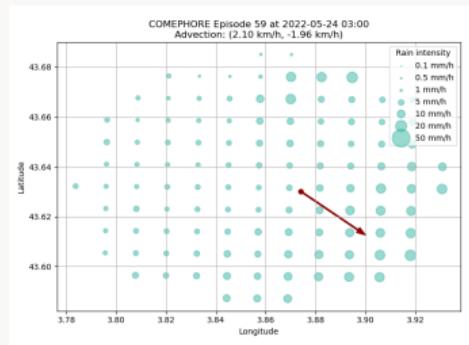
For one episode (COMEPhORE):



$t_0 - 1$



t_0

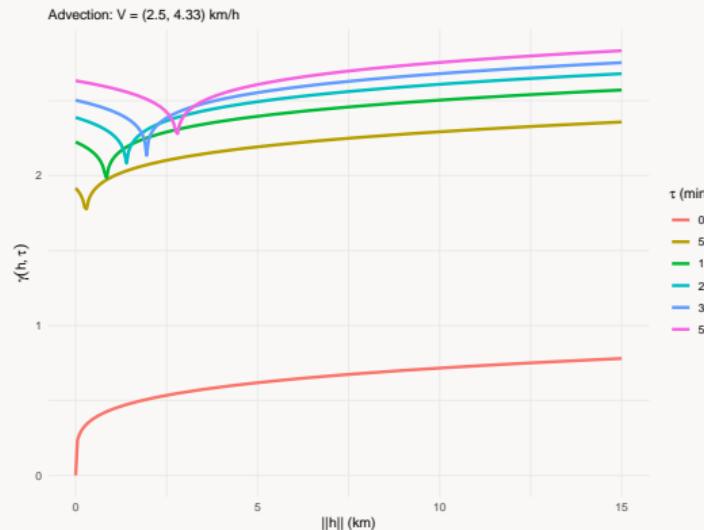


$t_0 + 1$

\mathbf{V}^{emp} is estimated from **rain storm barycenter displacement** within a time window.

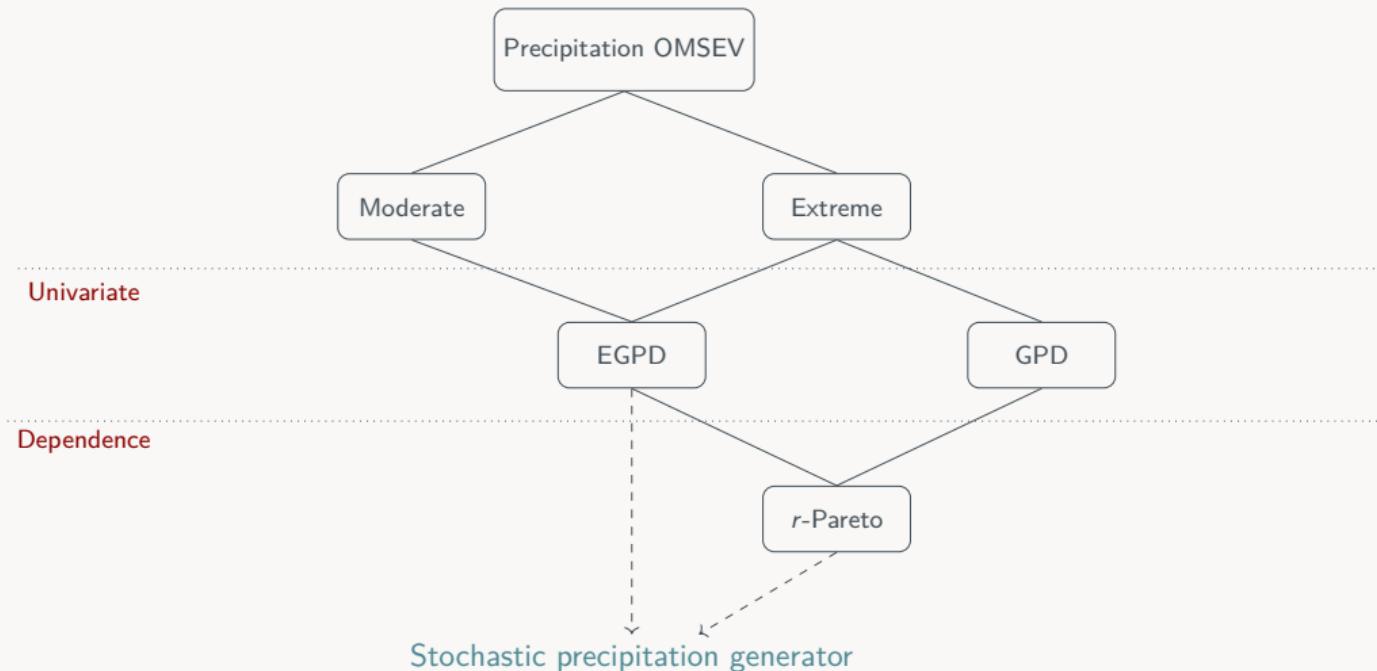
OMSEV RESULTS

- Advection group: relatively strong ($\|\mathbf{V}\| > 2$ km/h), north direction \Rightarrow 46 episodes
- Fixed η from COMEPHORE data: $\hat{\eta}_1 = 0.667$, $\hat{\eta}_2 = 1.757$ (km/h).
- OMSEV estimates (m/5 min): $\hat{\beta}_1 = 0.066$, $\hat{\beta}_2 = 0.747$, $\hat{\alpha}_1 = 0.480$, $\hat{\alpha}_2 = 0.691$

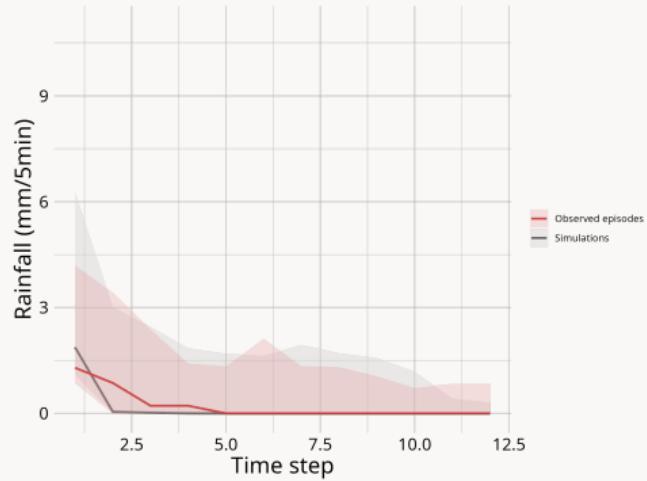
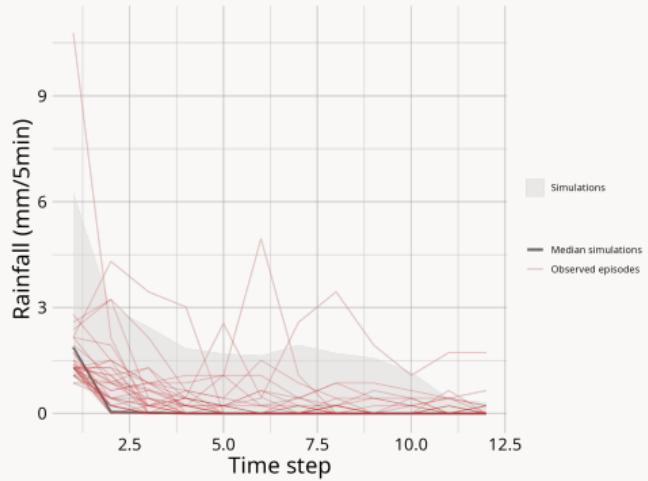


Estimated variogram with $\mathbf{V} = (2.5, 4.3)$ km/h

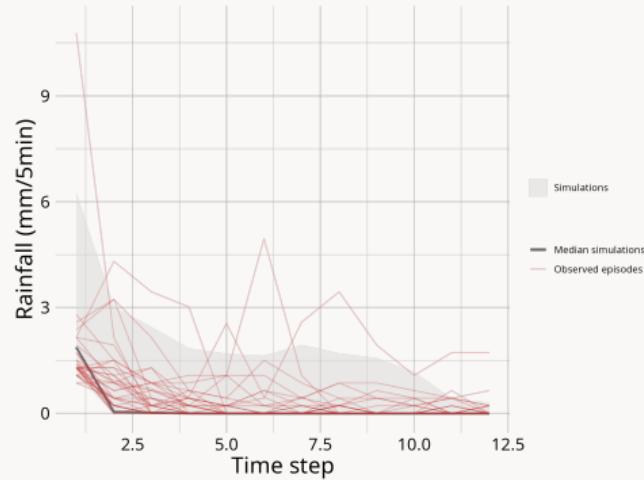
STOCHASTIC SIMULATION OF PRECIPITATION EVENTS



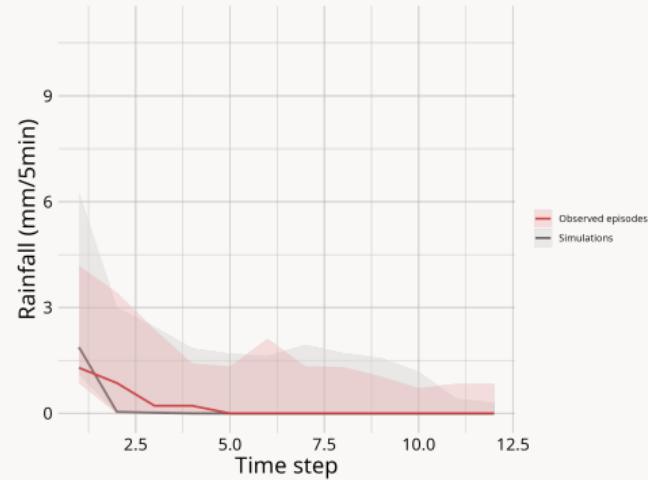
STOCHASTIC SIMULATION OF PRECIPITATION EVENTS



STOCHASTIC SIMULATION OF PRECIPITATION EVENTS



Moderate advection intensity
1000 simulations vs **34 observations**

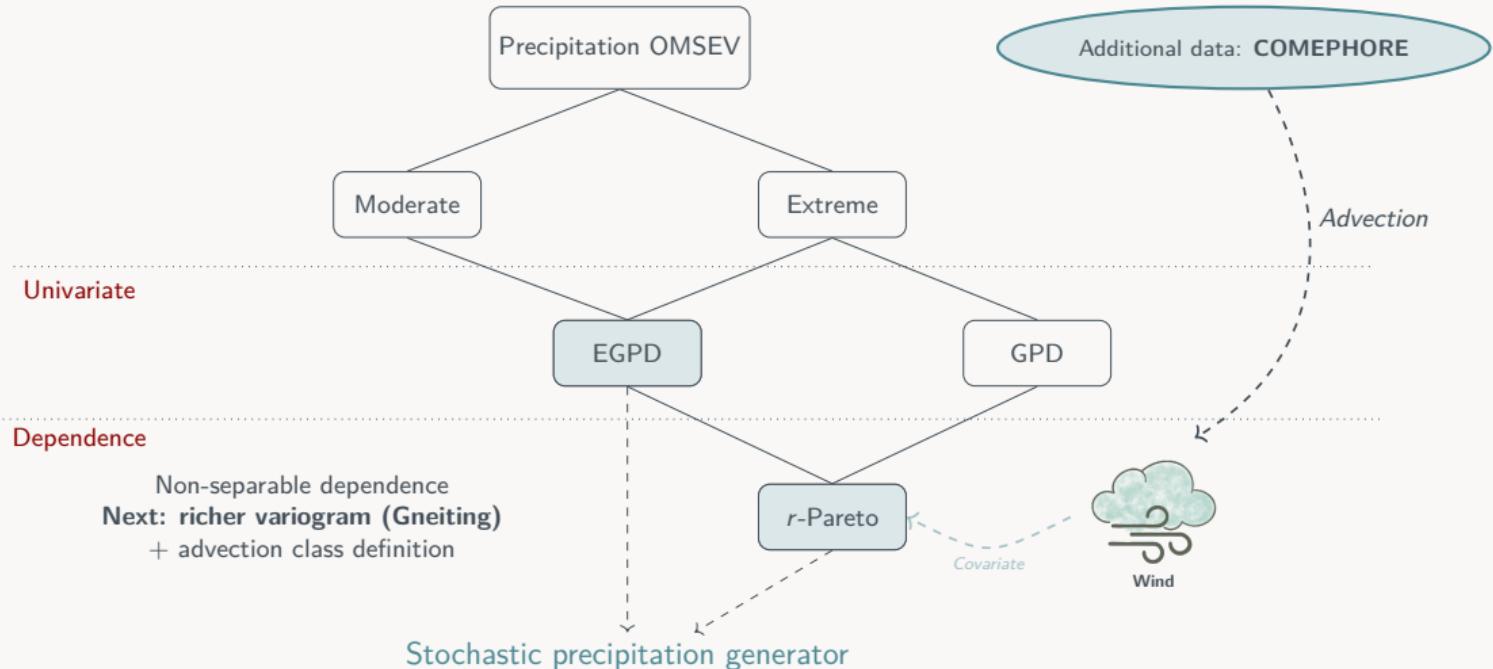


Moderate advection intensity
95% CI

Challenge:

- ▶ Limited information to estimate advection: incorporate wind data
- ▶ Advection class definition still needs refinement

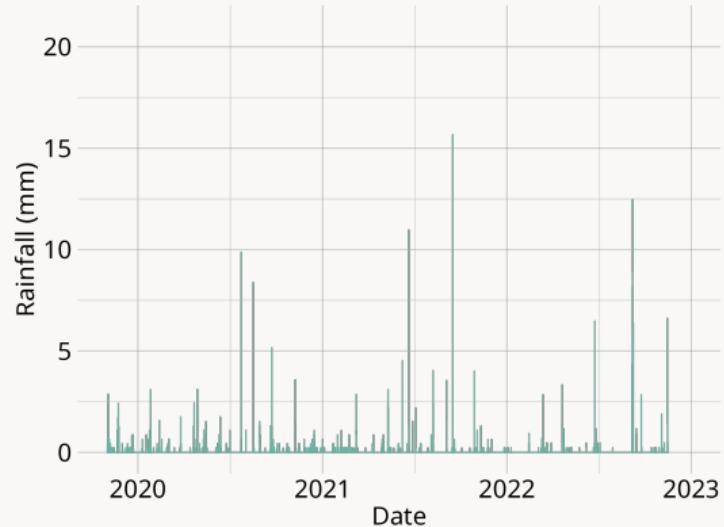
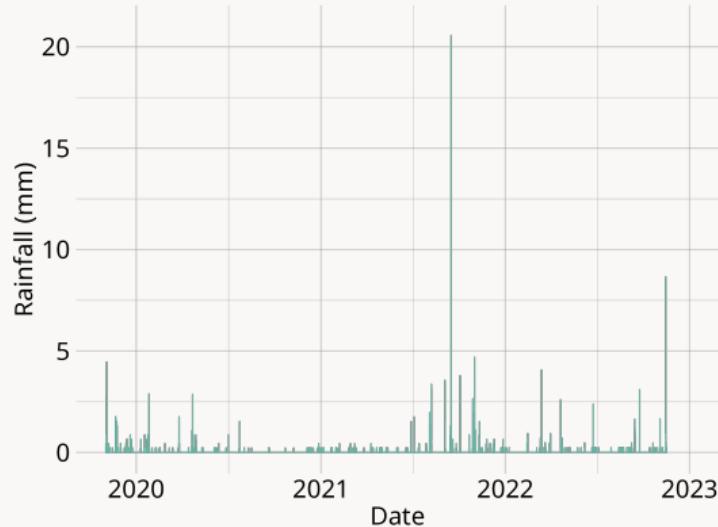
CONCLUSION & OUTLOOK



REFERENCES

- DAVIS, R. A., & MIKOSCH, T. (2009). The extremogram: A correlogram for extreme events.
- de FONDEVILLE, R., & DAVISON, A. C. (2018). High-dimensional peaks-over-threshold inference.
Biometrika, 105(3), 575–592.
- FINAUD-GUYOT, P., GUINOT, V., MARCHAND, P., NEPPEL, L., SALLES, C., & TOULEMONDE, G. (2023). Rainfall data collected by the HSM urban observatory (OMSEV).
- MATHERON, G. (1963). Principles of geostatistics. *Economic geology*, 58(8), 1246–1266.
- NAVEAU, P., HUSER, R., RIBEREAU, P., & HANNART, A. (2016). Modeling jointly low, moderate, and heavy rainfall intensities without a threshold selection. *Water Resources Research*.
- PICKARDS III, J. (1975). Statistical inference using extreme order statistics. *the Annals of Statistics*, 119–131.

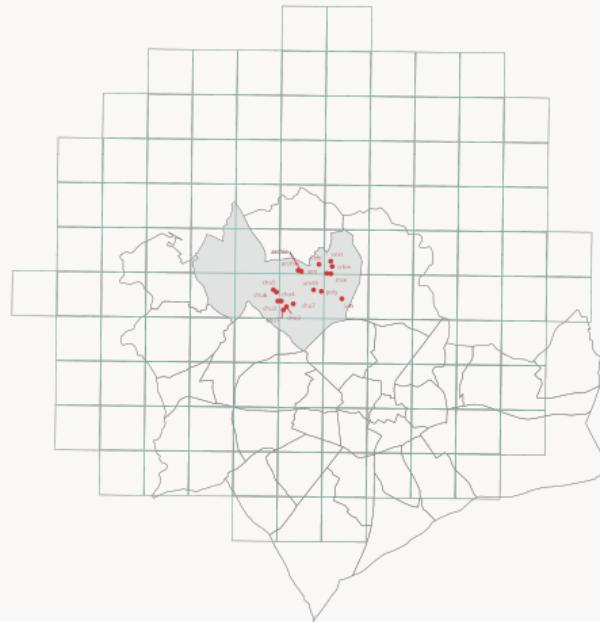
RAINFALL DATA - OMSEV



Rainfall amounts on CNRS and Polytech rain gauges

ADDITIONAL DATA

Pixels of 1km x 1km



- ▶ **Source:** COMEPHORE, Météo France
- ▶ **Time period:** [1997, 2023[
- ▶ **Temporal resolution:** Every hour
- ▶ **Spatial resolution:** 1 km²

More consistent data: Both datasets + Neural Network Downscaling.

Generalized Pareto Distribution



Extended GPD¹

$$\bar{H}_\xi \left(\frac{x-u}{\sigma} \right) = \begin{cases} \left(1 + \xi \frac{x-u}{\sigma} \right)_+^{-1/\xi} & \text{if } \xi \neq 0, \\ e^{-\frac{x-u}{\sigma}} & \text{if } \xi = 0, \end{cases}$$

where $a_+ = \max(a, 0)$, $\sigma > 0$, $x - u > 0$

$$F(x) = G \left(H_\xi \left(\frac{x}{\sigma} \right) \right),$$

where $G(x) = x^\kappa$, $\kappa > 0$

- ▶ Models extreme precipitation
- ▶ Depends on a threshold choice

- ▶ Models **moderate and extreme** precipitation
- ▶ Avoids a threshold choice

¹NAVEAU et al., 2016

ESTIMATION OF THE SEPARABLE VARIOGRAM PARAMETERS

Case of additive separability: $\frac{\gamma(\mathbf{h}, \tau)}{2} = \beta_1 \|\mathbf{h}\|^{\alpha_1} + \beta_2 |\tau|^{\alpha_2}, \quad 0 < \alpha_1, \alpha_2 \leq 2, \quad \beta_1, \beta_2 > 0$

Spatio-temporal

$$\chi(\mathbf{h}, \tau) = 2 \left(1 - \phi \left(\sqrt{\frac{1}{2} \gamma(\mathbf{h}, \tau)} \right) \right)$$

Transformation:

$$\zeta(\chi) = 2 \log \left(\phi^{-1} \left(1 - \frac{1}{2} \chi \right) \right)$$

Spatial

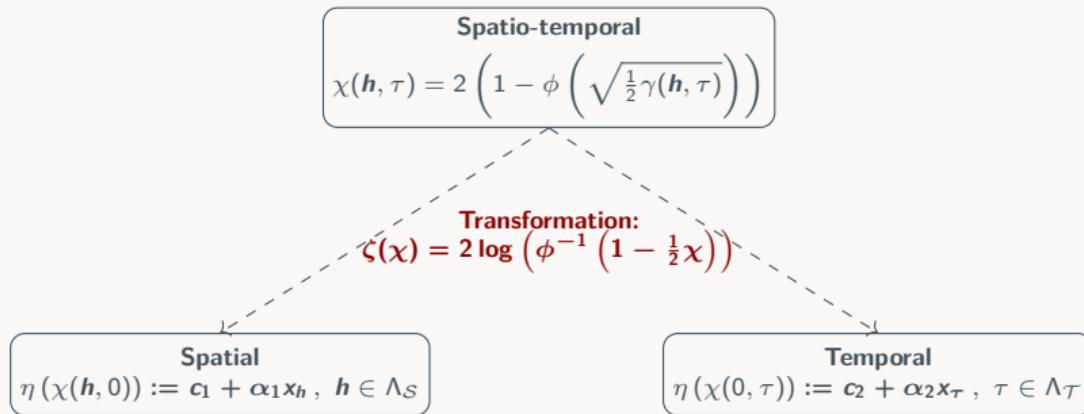
$$\zeta(\chi(\mathbf{h}, 0)) = \log \beta_1 + \alpha_1 \log \|\mathbf{h}\|, \quad \mathbf{h} \in \Lambda_S$$

Temporal

$$\zeta(\chi(0, \tau)) = \log \beta_2 + \alpha_2 \log \tau, \quad \tau \in \Lambda_T$$

ESTIMATION OF THE SEPARABLE VARIOGRAM PARAMETERS

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Weighted Least Squares Estimation (WLSE)

$$\begin{pmatrix} \hat{c}_i \\ \hat{\alpha}_i \end{pmatrix} = \underset{c_i, \alpha_i}{\operatorname{argmin}} \sum_x w_x \left(\zeta \left(\hat{\chi} \right) - (c_i + \alpha_i x) \right)^2$$

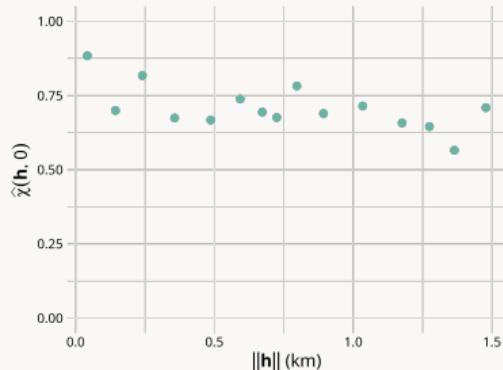
Empirical spatial extremogram

For a fixed $t \in \mathcal{T}$ and q a high quantile,

$$\hat{\chi}_q^{(t)}(\mathbf{h}, 0) = \frac{\frac{1}{|N_{\mathbf{h}}|} \sum_{i,j \mid (\mathbf{s}_i, \mathbf{s}_j) \in N_{\mathbf{h}}} \mathbb{1}_{\{X_{\mathbf{s}_i, t}^* > q, X_{\mathbf{s}_j, t}^* > q\}}}{\frac{1}{|\mathcal{S}|} \sum_{i=1}^{|\mathcal{S}|} \mathbb{1}_{\{X_{\mathbf{s}_i, t}^* > q\}}},$$

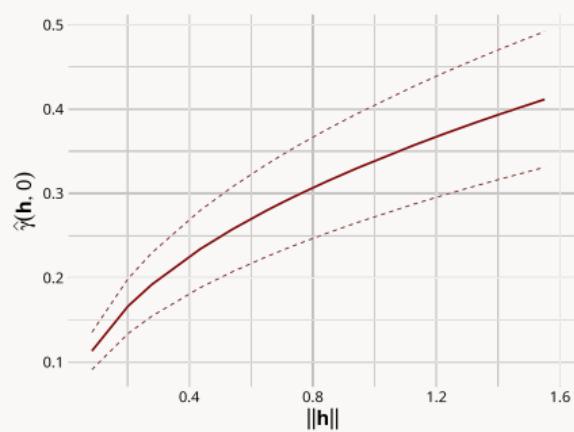
where $C_{\mathbf{h}}$ are equifrequent distance classes and

$$N_{\mathbf{h}} = \left\{ (\mathbf{s}_i, \mathbf{s}_j) \in \mathcal{S}^2 \mid \|\mathbf{s}_i - \mathbf{s}_j\| \in C_{\mathbf{h}} \right\}.$$



Transformation and WLSE

\implies



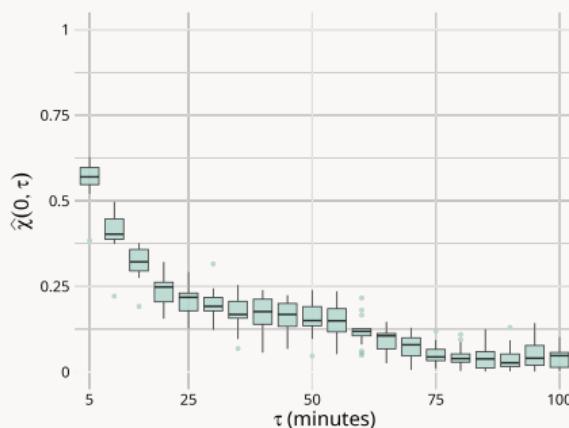
$$\text{Spatial variogram } \hat{\gamma}(\mathbf{h}, 0) = 2\hat{\beta}_1 \|\mathbf{h}\|^{\hat{\alpha}_1}$$

TEMPORAL DEPENDENCE ESTIMATION

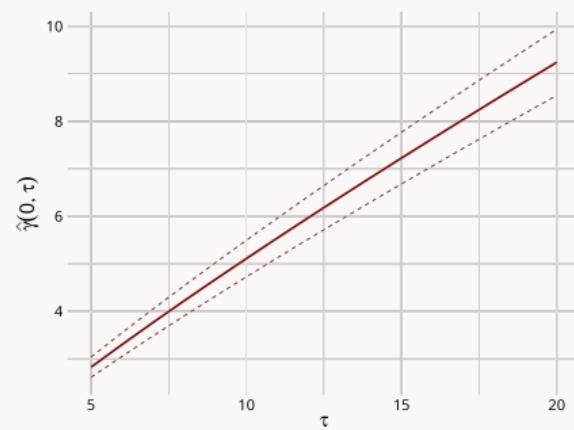
Empirical temporal extremogram

For a location $s \in \mathcal{S}$, a high quantile q and $t_k \in \{t_1, \dots, t_T\}$,

$$\hat{\chi}_q^{(s)}(\mathbf{0}, \tau) = \frac{\frac{1}{T-\tau} \sum_{k=1}^{T-\tau} \mathbb{1}_{\{X_{s,t_k}^* > q, X_{s,t_k+\tau}^* > q\}}}{\frac{1}{T} \sum_{k=1}^T \mathbb{1}_{\{X_{s,t_k}^* > q\}}}$$

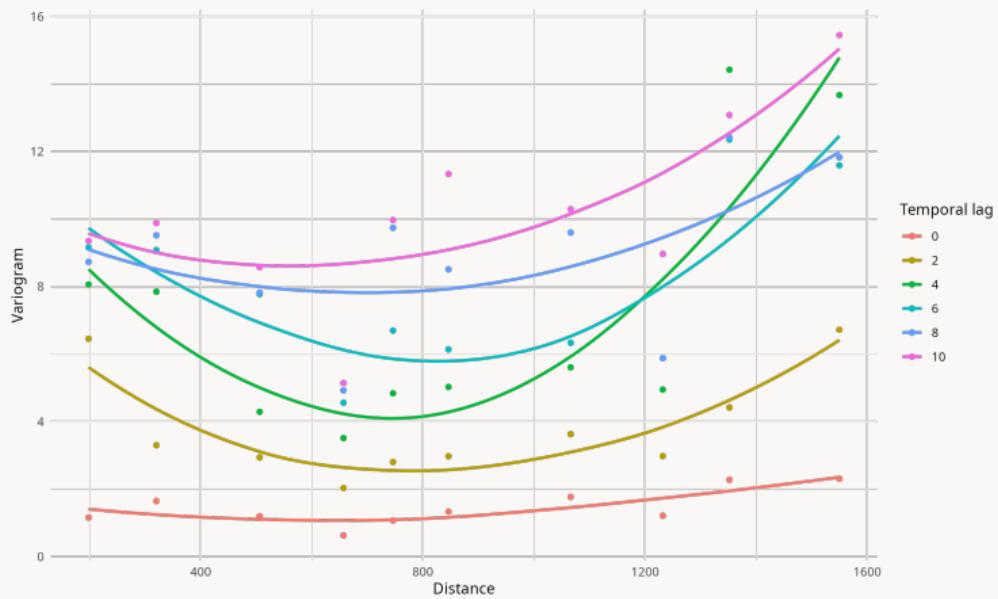


Transformation and WLSE



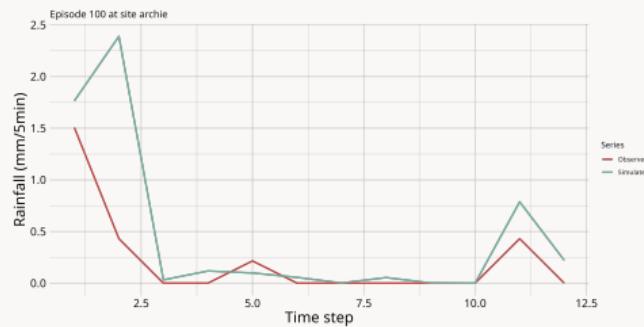
Temporal variogram $\hat{\gamma}(\mathbf{0}, \tau) = 2\hat{\beta}_2|\tau|^{\alpha_2}$

NON-SEPARABILITY

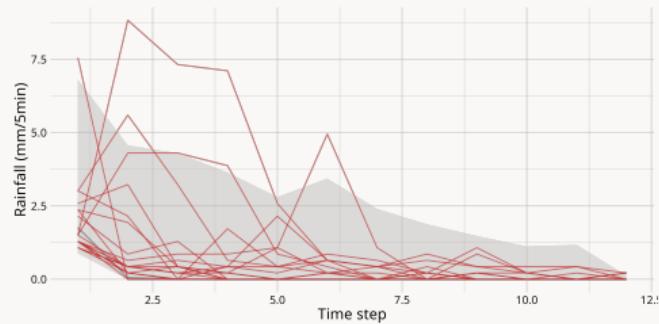


Empirical spatio-temporal variogram $\gamma(\mathbf{h}, \tau)$ on OMSEV data

STOCHASTIC SIMULATION OF PRECIPITATION EVENTS



Simulation vs Observation - Episode 100



High advection intensity - South direction
95% CI over 1000 simulations vs **20 observations**