AS1

February 10, 2020

1 Assignment 1

```
Name: Zhi Li
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In [1]: from netCDF4 import Dataset, num2date
        import numpy as np
        import matplotlib.pyplot as plt
        import matplotlib.dates as mdates
        import pandas as pd
        from netCDF4 import Dataset,num2date
        from datetime import datetime
        from dateutil.rrule import rrule, MONTHLY
        from scipy import stats
        from mpl_toolkits.basemap import Basemap, maskoceans, interp, shiftgrid
        import scipy
        from scipy.integrate import quad
        import sys
        sys.path.append('/Users/allen/Documents/Python/PlotGallary')
        from matplotlibconfig import basic
        #configure plot
        basic()
In [102]: # some utility functions
          def visualize(lon, lat, data, stipple=None, **figkwargs):
              HHHH
              Arqs:
              _____
              :figkwarqs - dict; {
                                   'ylabel': '',
                                   'cmap': 'seismic',
                                  'cRange': (-1,1),
                                  'title': '',
                                   'extent': tuple; (llclon, llclat, urclon, urclat)
```

```
'projection': str; default 'npstere'
11 11 11
ylabel= figkwargs.get('ylabel', '')
cmapName= figkwargs.get('cmap', 'seismic')
cRange= figkwargs.get('cRange', (-1,1))
title= figkwargs.get('title', '')
lllon, lllat, urlon, urlat= figkwargs.get('extent', (-180, -90, 180, 90))
proj= figkwargs.get('projection', 'npstere')
if proj!='npstere':
    rnd= False
else: rnd=True
cmin = cRange[0]; cmax = cRange[1];
cint = 0.2; clevs = np.round(np.arange(cmin,cmax,cint),1)
nlevs = len(clevs)-1
plt.gca()
cmap = plt.get_cmap(name= cmapName,lut=nlevs)
m = Basemap(llcrnrlon=lllon, llcrnrlat=lllat, urcrnrlon=urlon, urcrnrlat=urlat,
            projection=proj, lon_0=-100, boundinglat=20, round=rnd)
x,y = m(lon, lat)
m.drawcoastlines(linewidth=3)
m.drawmapboundary(linewidth=2)
  m.drawmeridians(range(-90, 90, 10))
  m.drawparallels(range(-180, 180, 20))
if stipple is not None:
    m.plot(x[stipple][::3],y[stipple][::3],'o',color='Gold',markersize=1.5) #
cs = m.contourf(x, y, data)
cbar = m.colorbar(cs)
cbar.ax.set_ylabel(ylabel)
plt.title(title, name='Arial', weight='bold', size=20)
return m
```

1.1 Question 1

This problem examines the statistical relationship between 500 hPa heights and the El Niño-Southern Oscillation (ENSO) during boreal winter (December, January and February). For this question, you will use two files: (1) ENSO.txt, which contains monthly-mean standardized values of an index measuring ENSO strength and (2) Z500.nc, which contains gridded monthly-mean 500 hPa geopotenal heights globally from 1950 - 2019. The steps for the analysis and final product will be as follows:

```
In [113]: # read in geopotential heights data
                        with Dataset('GPH500.nc', 'r') as nc:
                                  lonHeights= np.array(nc.variables['lon'][:])
                                  latHeights= np.array(nc.variables['lat'][:])
                                  time= nc.variables['time'][:]
                                  timeUnits= nc.variables['time'].units
                                  heightDate= num2date(time, timeUnits, calendar='standard')
                                  heightMonth= np.array([date.month for date in heightDate])
                                  heightYear= np.array([date.year for date in heightDate])
                                  heights= np.array(nc.variables['hgt'][:])
In [114]: heights[heights<0] = np.nan</pre>
In [115]: # read in ENSO data
                        ENSOfile= np.loadtxt('ENSO.txt', skiprows=2)
                        ENSOyear= ENSOfile[:,0]
                        ENSOmonth= ENSOfile[:,1]
                        ENSO= ENSOfile[:,2]
                       # select monthly height from 1981 to 2010
indices= np.where((heightYear>=1981) & (heightYear<=2010))
" commute the climatology and corresponding date

con(heights[indices], axis=0)

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      1. Calculate 500 hPa height anomalies, use the 1981 – 2010 base period as your climatology /
           mean.
In [116]: # select monthly height from 1981 to 2010
     2. Calculate the Dec - Feb mean 500 hPa geopotential height anomaly field and ENSO index
           from each complete winter season in the dataset
In [117]: # subselect DJF anomalies and corresponding date
                        anoIndex= np.where((ENSOmonth==12) | (ENSOmonth==1) | (ENSOmonth==2))
anoHeights= anoHeights[anoIndex]

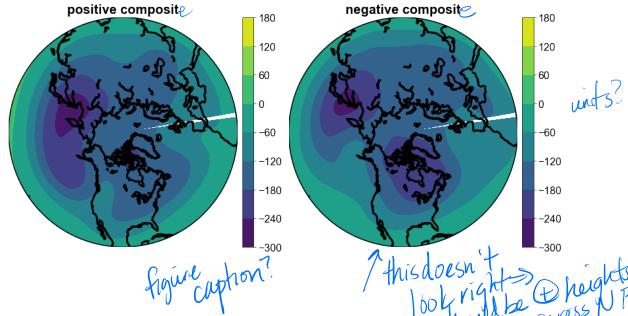
ENSO= ENSO[anoIndex]

In [118]: posENSO= ENSO[ENSO>= np.std(ENSO)]

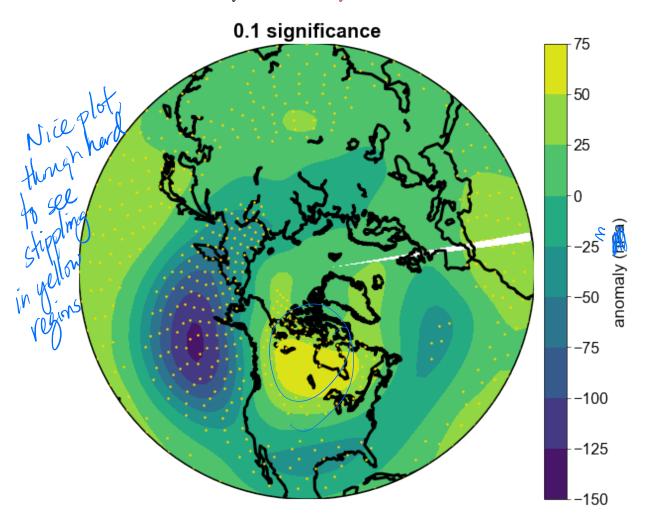
negENSO= ENSO[ENSO<= -np.std(ENSO)]

Standard FMSO>= np.std(ENSO))
                        negIndex= np.where(ENSO<= -np.std(ENSO))</pre>
     3. Generate composite DJF-mean 500 hPa height anomaly plots for times when ENSO index
           exceeds 1\sigma and exceeds -1\sigma. Use shaded contours for plotting.
In [119]: posComp= anoHeights[posIndex]
                        negComp= anoHeights[negIndex]
In [120]: posCompMean= np.array(posComp).mean(axis=0)
```

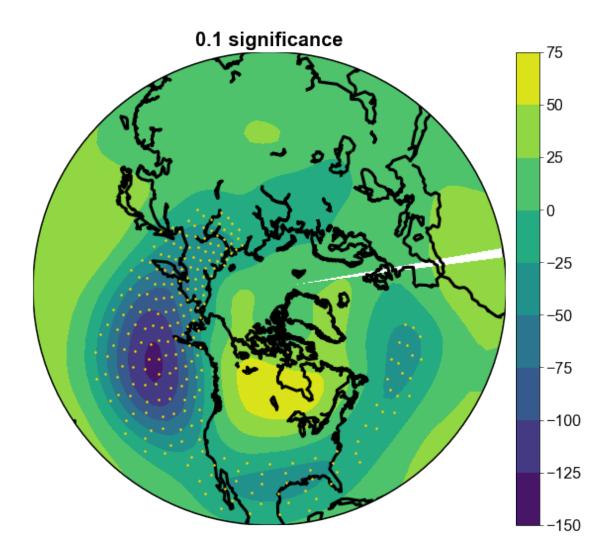
negCompMean= np.array(negComp).mean(axis=0)



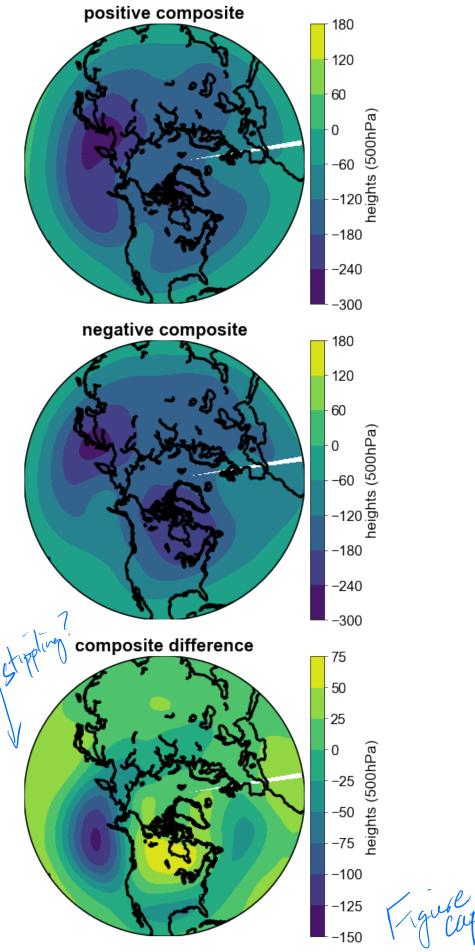
4. Compute the composite difference. Identify regions where the composite difference is significant at the p<0.1 level using either a t-test or a Monte Carlo test. Consider each event as a degree of freedom for the calculation. Outline or stipple these significant regions on the composite difference plot



```
for n in range(niter):
            c1 = np.random.choice(djf, len(posComp), replace=False)
            c2 = np.random.choice(djf, len(negComp), replace=False)
            mcPos = np.nanmean(y[c1,:],0)
            mcNeg = np.nanmean(y[c2,:],0)
            mcCompDiff[n,:] = mcPos-mcNeg
         # Calculate the pvalue for each point
        c12 = compDiff.reshape(J*I,order='F')
        pval = np.array([stats.percentileofscore(mcCompDiff[:,i],c12[i])/100.
                        for i in range(c12.size)])
In [129]: alpha= 0.1
        stipple= pval<=alpha
        fig= plt.figure(figsize=(8,8))
        ax= fig.add_subplot(111)
        map= visualize(lons, lats, compDiff, stipple.reshape(lons.shape, order='F'),
                      title='0.1 significance')
```



5. Assemble the plots for presentation as a 3-panel plot, with the positive ENSO composite on top, the negative ENSO composite in the middle, and the composite difference on the bottom. Make the plot only for the Northern Hemisphere (i.e., 20-90 N).



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As from the significance test, both the t-test and Monte-Carlo test show that the differences of geopotential heights for positive ENSO and negative ENSO are significant in North Pacific Ocean, southern U.S., and part of North Atlantic Ocean. In these areas, the difference is significantly skewed towards negative, which indicates the negative ENSO indicator strongly affect this extratropical places, especially decreasing the 500 hPa geopotential height. As a consequence, this impact will change the global weather pattern (i.e., influence the precipitation distribution).

There are some differences between t-test and Monte-Carlo simulation as well. we can see the t-test result (stipples) cover more areas in the globe, suggesting the ENSO factors (positive and Ls yes, but just a couple of missleps negative) have more places under significant differences.

Question 2

```
In [231]: import pandas as pd
In [247]: precip= pd.read_csv('BostonMaxDailyPrecip.txt', sep='\t+', skiprows=2, header=None)
/usr/local/lib/python3.7/site-packages/ipykernel_launcher.py:1: ParserWarning: Falling back to
```

"""Entry point for launching an IPython kernel.

```
In [248]: precip.columns= ['precipInch', 'date', 'precipMM']
In [252]: precip.date= pd.to_datetime(precip.date)
```

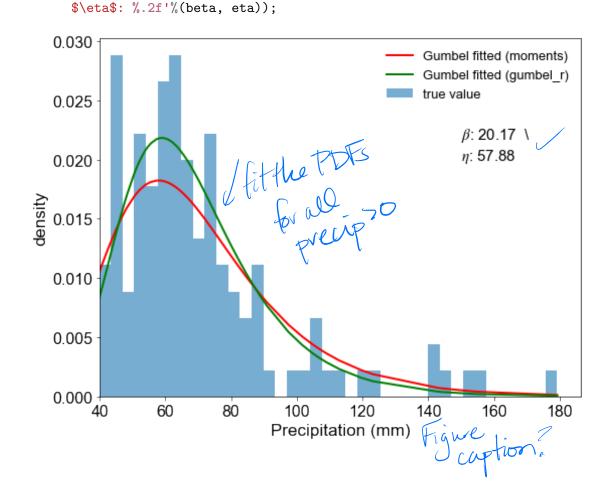
1. Using the data provided, fit a Gumbel distribution to the Boston precipitation data in mm. Use the methods of moments to estimate the parameters for the distribution.

$$Gumbel = \frac{1}{\beta}exp(-exp(-\frac{x-\eta}{\beta}) - \frac{x-\eta}{\beta})$$

```
In [269]: def gumbel(x, eta, beta):
              pdf= 1/beta*np.exp(-np.exp(-(x-eta)/beta)-(x-eta)/beta)
              return pdf
          def est_beta(x):
              return np.std(x)*6**.5/np.pi
          def est_eta(x, beta):
              return x.mean()-0.5772*beta
In [270]: beta= est_beta(precip.precipMM)
          eta= est_eta(precip.precipMM, beta)
In [272]: print('estimated beta and eta are: %.2f, %.2f'%(beta, eta))
estimated beta and eta are: 20.17, 57.88
```

```
In [273]: gumbel_pdf= gumbel(precip.precipMM, eta, beta)
```

2. Make a plot with a histogram of the annual daily-maximum precipitation (in mm) and the fitted Gumbel distribution overlaid on the histogram. On your plot, indicate the values of β and η used for the fit.



3. Using your fitted distribution, determine what the probability is that Boston will experience a year with a daily maximum precipitation amount (i) greater than 130 mm (about 5 in) and (ii) less than 50 mm (about 2 in).

$$P(x >= 130) = \int_{130}^{+\inf} p df(x) dx$$

$$P(x \le 50) = \int_0^{50} p df(x) dx$$

The probability exceeds 130 mm from Gumbel function is 0.03, with numerical error of 1e-9 The probability blow 50 mm from Gumbel function is 0.23, with numerical error of 5e-9

In [431]: # cross-check the answer with python inner module stats.gumbel_r

from scipy.stats import gumbel_r
loc, scale= gumbel_r.fit(precip.precipMM)
print(r'fitted beta: %.2f, eta: %.2f, the difference is (%.2f, %.2f) for (beta, eta)
 relative error (%.2f%, %.2f%)'%(scale, loc, beta-scale, eta-loc, (beta-scale)

fitted beta: 16.85, eta: 58.90, the difference is (3.32, -1.02) for (beta, eta)\
relative error (16.48%, -1.76%)

you checked the fitted to the second to th

In [432]: print('The true probability beyond 130 mm is %.2f, below 50 mm is %.2f'%(
quad(lambda x: gumbel(x, loc, scale), 130, np.inf)[0], quad(lambda x: gumbel(x, loc, scale),

The true probability beyond 130 mm is 0.01, below 50 mm is 0.18

- 4. Calculate the daily maximum precipitation amounts corresponding to a 1-in-50 year, 1-in-100 year, and 1-in-500 year event in Boston.

- In [451]: print('The daily maximum precipitation for 1-in-50 year, 1-in-100-year and 1-in-500 year, fifty, hundred, fivehundred))

contemporary (

The daily maximum precipitation for 1-in-50 year, 1-in-100-year and 1-in-500 year are 136.60 m

```
In [452]: print('The true daily maximum precipitation for 1-in-50 year, 1-in-100-year and 1-in-gumbel_r.ppf(1-0.02, loc, scale), gumbel_r.ppf(1-0.01, loc, scale), gumbel_r.ppf(1-0.02, loc, scale)
```

The true daily maximum precipitation for 1-in-50 year, 1-in-100-year and 1-in-500 year are 124

1.3 Question 3

A popular topic in wintertime subseasonal and seasonal forecasting is using the state of the stratospheric polar vortex to make skillful predictions of tropospheric weather regimes. One mode of climate variability that forecasters seek to predict is the Northern Annular Mode (NAM) in both the stratosphere and the troposphere. For stratospheric levels, strongly positive (negative) values of the NAM index (e.g., NAM at 50 hPa; NAM50) typically mean that the stratospheric polar vortex is stronger (weaker) than normal i.e., climatological westerlies in the stratosphere are stronger (weaker) than average. In the troposphere, strongly positive (negative) values of the near-surface NAM index (e.g., NAM at 1000 hPa; NAM1000) indicate that the mean position of the polar jet stream is more poleward (equatorward) than normal, suggesting warmer (colder) than normal temperatures across the Northern Hemisphere mid-latitudes. In this exercise, you will investigate how the probability density function (PDF) of NAM1000 may change given the background state of the stratospheric circulation.

1. Calculate the clamatological PDF of NAM_{1000} (i.e., all NDJFMA days). Then, on the same plot, overlay two other conditional PDFs:

```
The PDF of NAM_{1000} for days when NAM_{50} > +1\sigma
The PDF of NAM_{1000} for days when NAM_{50} < -2\sigma
```

```
In [503]: #composite
    ndjfma= np.where((NAM1000.month>=11) | (NAM1000.month<=4))[0]

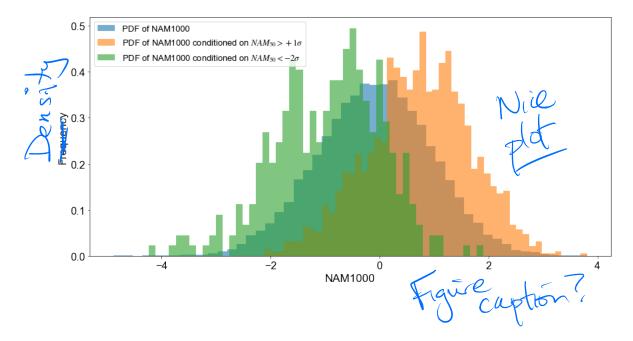
climNAM1000= NAM1000.iloc[ndjfma, :]
    climNAM50= NAM50.iloc[ndjfma, :]

posInds= np.where((climNAM50.iloc[:,-1]> np.std(climNAM50.iloc[:,-1])))[0]
    negInds= np.where((climNAM50.iloc[:,-1]< -2* np.std(climNAM50.iloc[:,-1])))[0]

posClimNAM1000= climNAM1000.iloc[posInds, -1]
    negClimNAM1000= climNAM1000.iloc[negInds, -1]

#make plot
fig, ax= plt.subplots(1,1, figsize=(12,6))
    climNAM1000.iloc[:,-1].plot.hist(alpha=.6, bins=50, density=True,label='PDF of NAM10</pre>
```

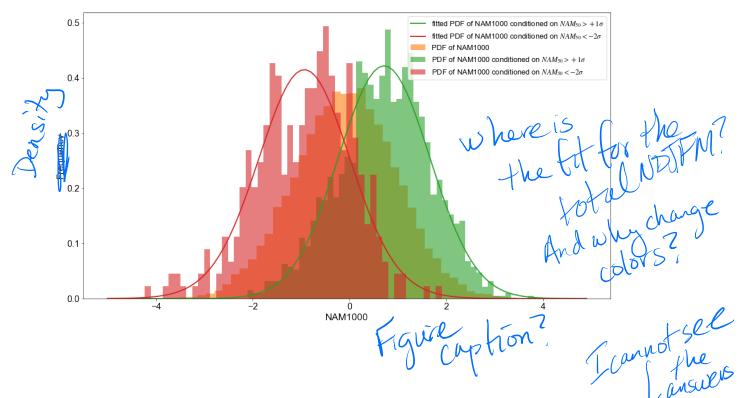
```
posClimNAM1000.plot.hist(alpha=.6, bins=50, density=True,label='PDF of NAM1000 condit
negClimNAM1000.plot.hist(alpha=.6,bins=50, density=True,label='PDF of NAM1000 condit
ax.legend()
ax.set_xlabel('NAM1000');
```



By firstly looking at the distribution for three data, they behave more or less like normal. Thus, we decided to fit into a t-distribution. Then the next thing we need to do is to calibrate those parameters.

Assume the degree of freedom for each case is the sample size.

```
In [521]: paramsPos= stats.t.fit(posClimNAM1000.values, len(posClimNAM1000)-1)
          paramsNeg= stats.t.fit(negClimNAM1000.values, len(negClimNAM1000)-1)
                                           Sample Sizes PDF
In [522]: posDist= stats.t(*paramsPos)
          negDist= stats.t(*paramsNeg)
In [532]: fig, ax= plt.subplots(1,1, figsize=(14,8))
          ax.plot(np.arange(-5,5,0.1), posDist.pdf(np.arange(-5,5,0.1)), color= 'C2',
                  label='fitted PDF of NAM1000 conditioned on $NAM_{50}>+1\sigma$')
          ax.plot(np.arange(-5,5,0.1), negDist.pdf(np.arange(-5,5,0.1)),color='C3',
                  label='fitted PDF of NAM1000 conditioned on $NAM_{50}<-2\sigma$')
          climNAM1000.iloc[:,-1].plot.hist(alpha=.6, bins=50, density=True,
                                           label='PDF of NAM1000', color= 'C1')
          posClimNAM1000.plot.hist(alpha=.6, bins=50, density=True, color= 'C2',
                                   label='PDF of NAM1000 conditioned on $NAM_{50}>+1\sigma$')
          negClimNAM1000.plot.hist(alpha=.6,bins=50, density=True,color= 'C3',
                                   label='PDF of NAM1000 conditioned on $NAM_{50}<-2\sigma$')
          ax.set_xlabel('NAM1000')
          plt.legend();
```



In [546]: print('probability that \$NAM_{1000}\$ smaller than \$-1\sigma\$ conditioned on \$NAM_{500} '%(posDist.cdf(-np.std(posClimNAM1000.values))))

probability that NAM_{1000} smaller than $-1 \simeq \infty$ conditioned on $NAM_{50} > +1 \simeq 0$.

In [547]: print('probability that \$NAM_{1000}\$ smaller than \$-1\sigma\$ conditioned on \$NAM_{50}' (negDist.cdf(-np.std(negClimNAM1000.values))))

probability that NAM_{1000} smaller than $-1 \simeq conditioned on <math>NAM_{50}<-2 \simeq conditioned$

K-S test

In [559]: stats.ks_2samp(posDist.rvs(size=100), negDist.rvs(size=100))

Out[559]: Ks_2sampResult(statistic=0.63, pvalue=6.314161651442317e-19)

From the K-S test results, we can reject the null hypothesis that two samples are identical with almost 100% confidence.

The NAM_{50} (stratospheric polar vortex) is a good indicator to describe northern hemisphere surface weather since is has distinct behavior (from the PDFs) among positive NAM_{50} and negative NAM_{50} . Basically, positive NAM_{50} pushes NAM_{1000} more towards positive and negative NAM_{50} pushes NAM_{1000} towards negative. Hence, it could be an indicator to predict northern hemisphere surface weather.

Nie interpretation
of the stats and tes

Q1: 29/35 Q2: 30/30 Q3: 29/35 No figure coptions: -6 Total: 8276