

## 2. Methods

### 2.1. Model design

We modeled the vertical accretion of a tidal platform ( $d\eta/dt$ ) using a zero-dimensional mass balance approach initially described by Krone (1987) and validated by subsequent studies (Allen, 1990; French, 1993; Temmerman et al., 2003, 2004). The rate of vertical accretion is described as

$$\frac{d\eta(t)}{dt} = \frac{dS_m(t)}{dt} + \frac{dS_o(t)}{dt} + \frac{dP(t)}{dt} + \frac{dM(t)}{dt}, \quad (1)$$

where  $dS_m(t)/dt$  is the rate of mineral sedimentation,  $dS_o(t)/dt$  is the rate of organic matter sedimentation,  $dP(t)/dt$  is the rate of compaction of the deposited sediment, and  $dM(t)/dt$  is the rate of tectonic subsidence. Studies show only ~3 % of organic matter is preserved in the region so we chose to neglect organic sedimentation. We set the rate of tectonic subsidence to 6 mm which is consistent with measured rates for the region Higgins et al. (2014). We also effectively internalize the compaction term by using dry bulk density within the mineral sedimentation term. Thus, we set terms  $dS_o(t)/dt$ ,  $dP(t)/dt$ , and  $dM(t)/dt$  to zero.

In order to solve  $dS_m(t)/dt$ , we began by conceptualizing a tidal platform periodically inundated by cyclical tides. We first defined depth to be

$$h(t) = \zeta(t) - \eta(t), \quad (2)$$

where  $\zeta(t)$  is the water-surface elevation and  $\eta(t)$  is the sediment-surface elevation which also implies that

$$\frac{dh(t)}{dt} = \frac{d\zeta(t)}{dt} - \frac{d\eta(t)}{dt}. \quad (3)$$

Independently, we assume when  $h(t) > 0$ , the rate of mineral sedimentation is

$$\frac{dS_m(t)}{dt} = \frac{w_s C(t)}{\rho_b}, \quad (4)$$

where  $w_s$  is the nominal settling velocity of sediment grain,  $C(t)$  is the depth-averaged suspended sediment concentration (SSC) in the water column, and  $\rho_b$  is the bulk density of the sediment. We assumed no resuspension of mineral sediment which is practical and consistent with previous studies (Krone, 1987; Allen, 1990; French, 1993; Temmerman et al., 2003, 2004). We used Stoke's law to determine  $w_s$ . Stoke's law assumes unhindered settling which likely overestimates actual settling rates and, therefore, mineral sedimentation rates. However, we only considers a median grain size which likely underestimates mineral sedimentation rates from coarser material. Model calibration further corrects for these errors. Thus, the  $w_s$  given by Stoke's law should be considered an imprecise, but reasonable approximation.

In order to solve for  $C(t)$  in eq. (4), we first defined a mass balance of sediment within the water column as

$$\frac{d}{dt}[h(t)C(t)] = -w_s C(t) + C_b \frac{dh(t)}{dt}, \quad (5)$$

which can be expanded and rearranged as

$$\frac{dC(t)}{dt} = -\frac{w_s C(t)}{h(t)} - \frac{1}{h(t)}[C(t) - C_b] \frac{dh(t)}{dt}. \quad (6)$$

The mass flux from the boundary term was constrained to only occur during flood tide ( $d\zeta/dt > 0$ ) which is consistent with previous studies (Krone, 1987; Allen, 1990; French, 1993; Temmerman et al., 2003, 2004). We formalized this mathematically using a Heaviside function which serves as a binary switch controlling the flux of sediment from the boundary. The Heaviside function is given as

$$S = \frac{d\zeta}{dt}, H(S) = \begin{cases} 0 & \text{if } S < 0 \\ 1 & \text{if } S \geq 0. \end{cases} \quad (7)$$

Need ref

Need to investigate further. Look into Steckler papers from 2010 and 2013.

eq. (6) then becomes

$$\frac{dC(t)}{dt} = -\frac{w_s C(t)}{h(t)} - \frac{H(S)}{h(t)} [C(t) - C_b] \frac{dh(t)}{dt}. \quad (8)$$

Finally, we solved eqs. (1), (4) and (8) in that order. Our approach differs from previous studies in that we resolve each eqs. (1), (4) and (8) at every timestep. The system of equations was solved using an explicit Runge-Kutta method of order 5(4) (Dormand and Prunce, 1980) and implemented in Python using SciPy (Virtanen et al., 2020).

## Software and/or data availability

## Acknowledgements

## References

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