A photograph of a large, light-colored glacial erratic boulder resting on a bedrock outcrop in a deciduous forest. The boulder is surrounded by smaller rocks and fallen tree branches. Bare trees stand in the background under a clear sky.

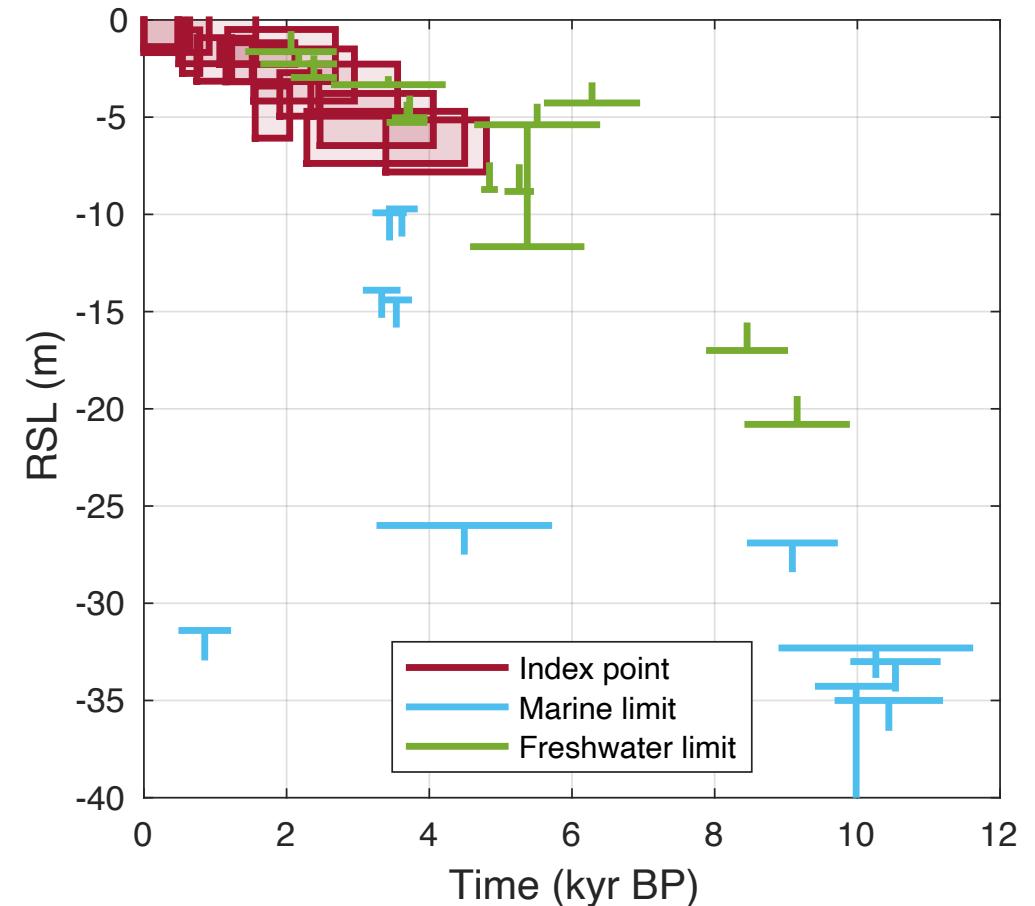
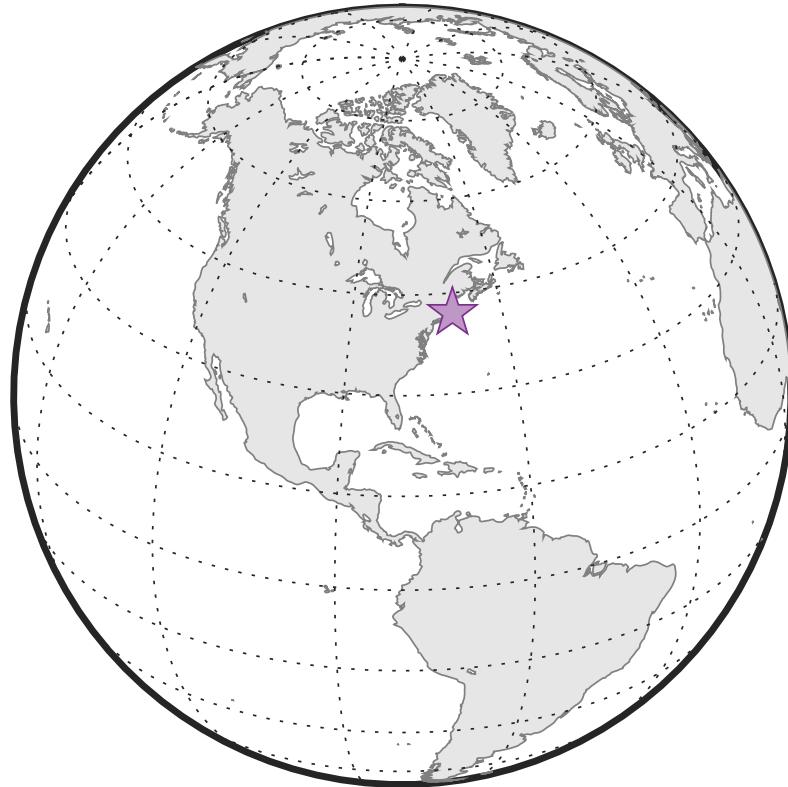
# Doing Bayesian analysis for glacial isostatic adjustment

Chris Piecuch, Woods Hole Oceanographic Institution

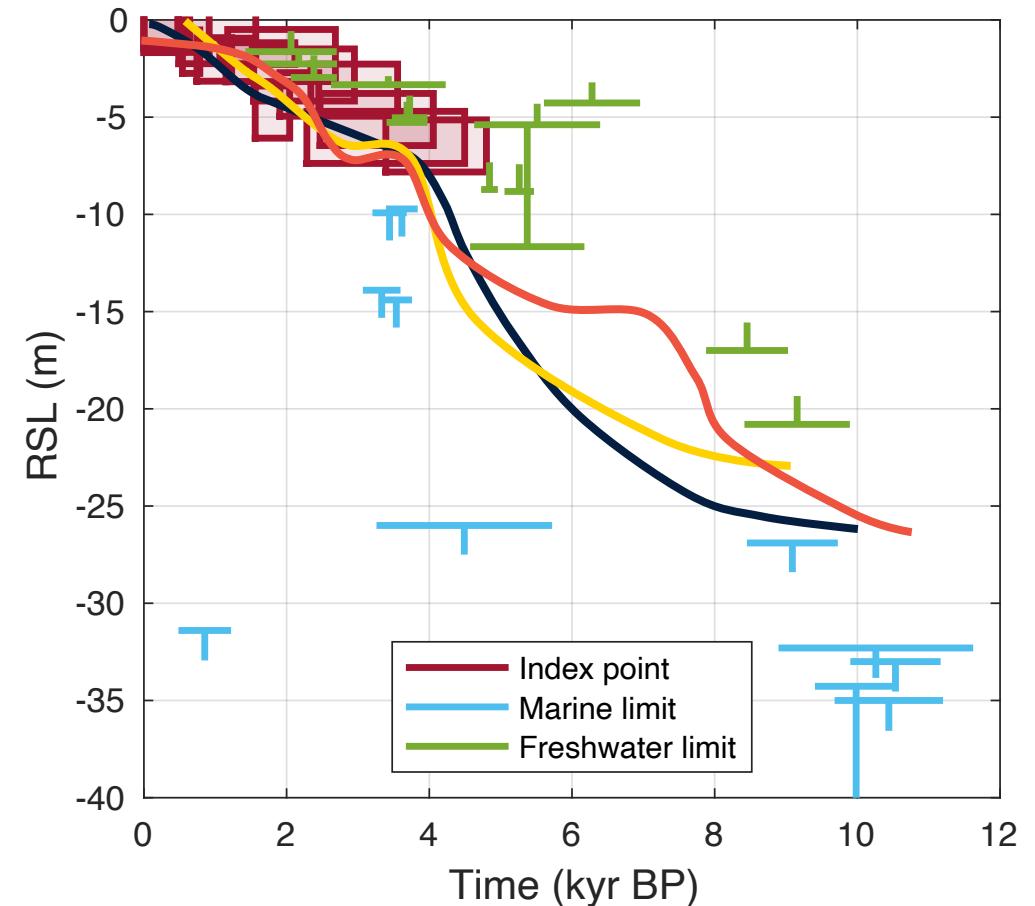
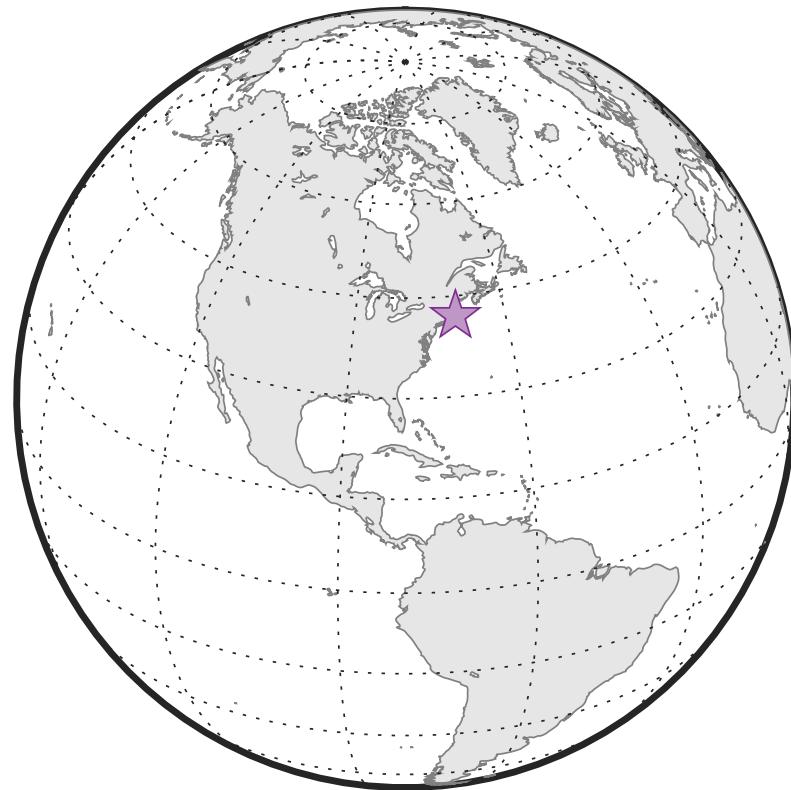
**What do we have?**

**What do we want?**

# We have uncertain proxies; we want past rates of change



# We have uncertain proxies; we want past rates of change

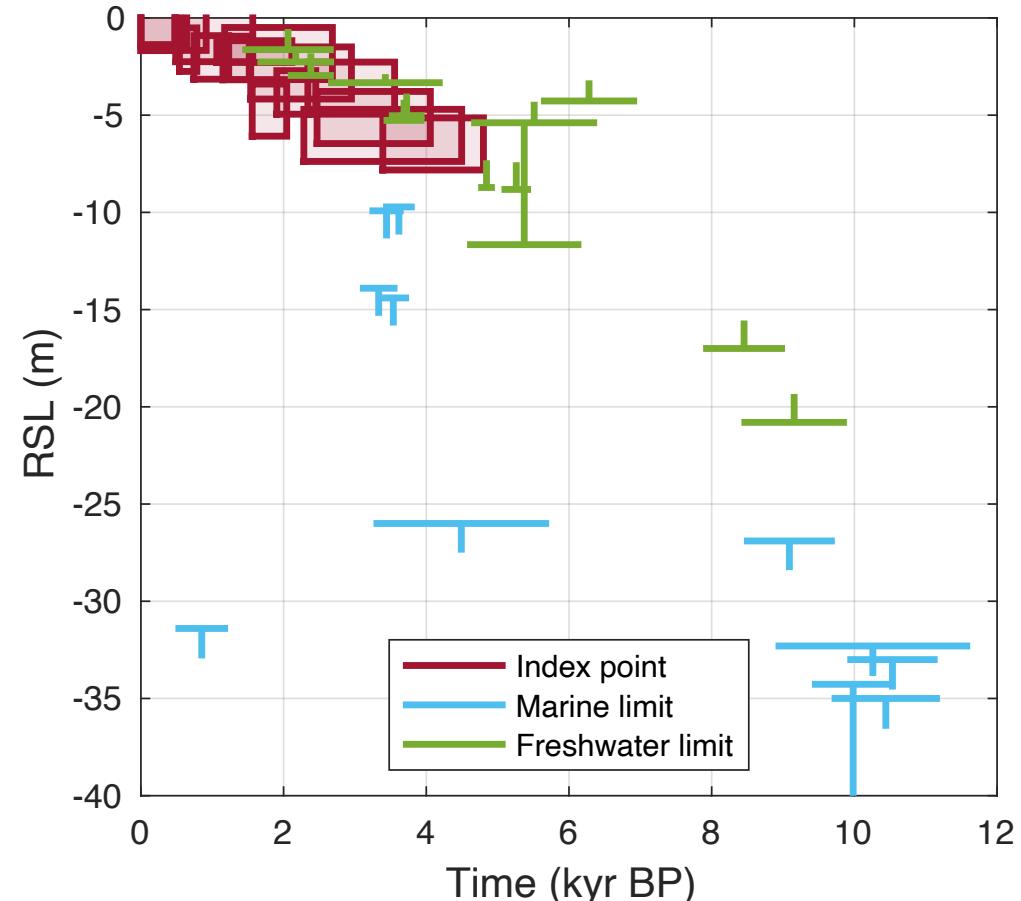


# We have uncertain proxies; we want past rates of change

So how do we do this?

How do we get what we want  
given what we have?

Simplest thing is to fit a line  
through the best-estimate  
ages and relative sea levels  
from the index points using  
ordinary least squares (OLS)



# We have uncertain proxies; we want past rates of change

We specify the model

$$y_n = \alpha x_n + \beta + \varepsilon_n$$

where

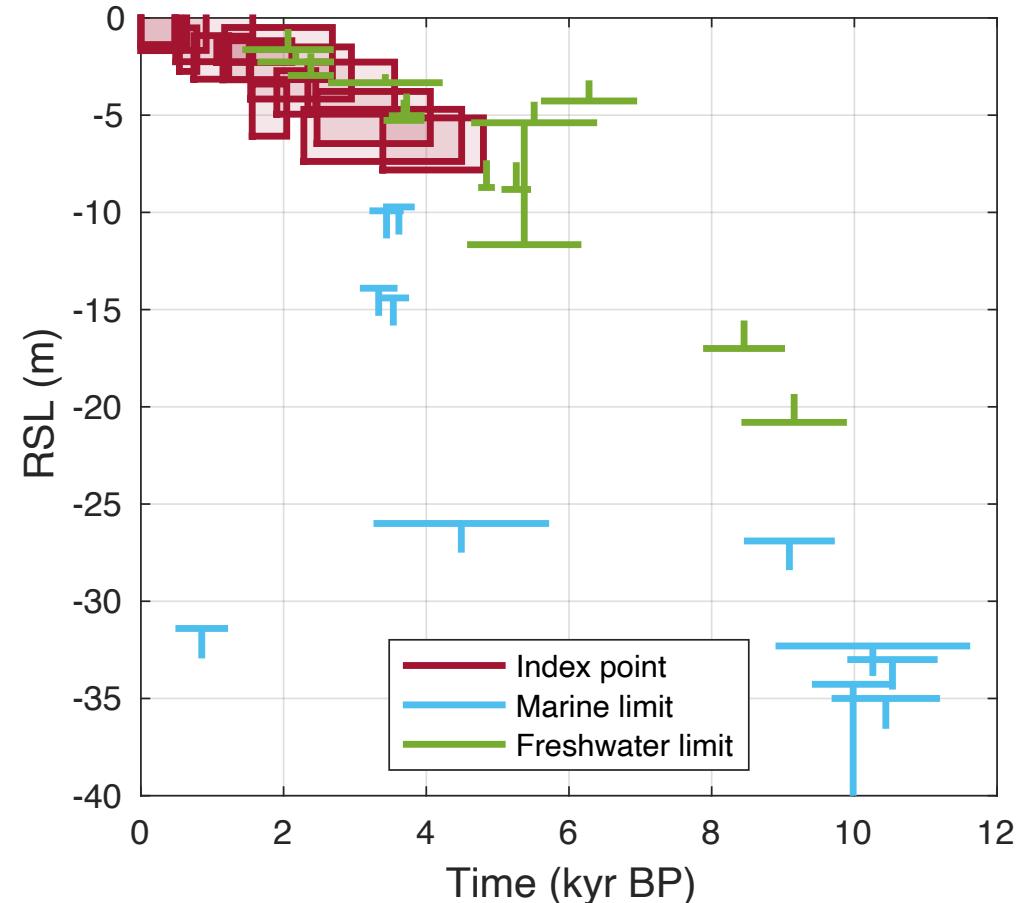
$y_n$  is the nth value of sea level

$x_n$  is the nth value of time

$\alpha$  is the slope

$\beta$  is the intercept

$\varepsilon_n$  is the nth residual (random  
unstructured noise)



# We have uncertain proxies; we want past rates of change

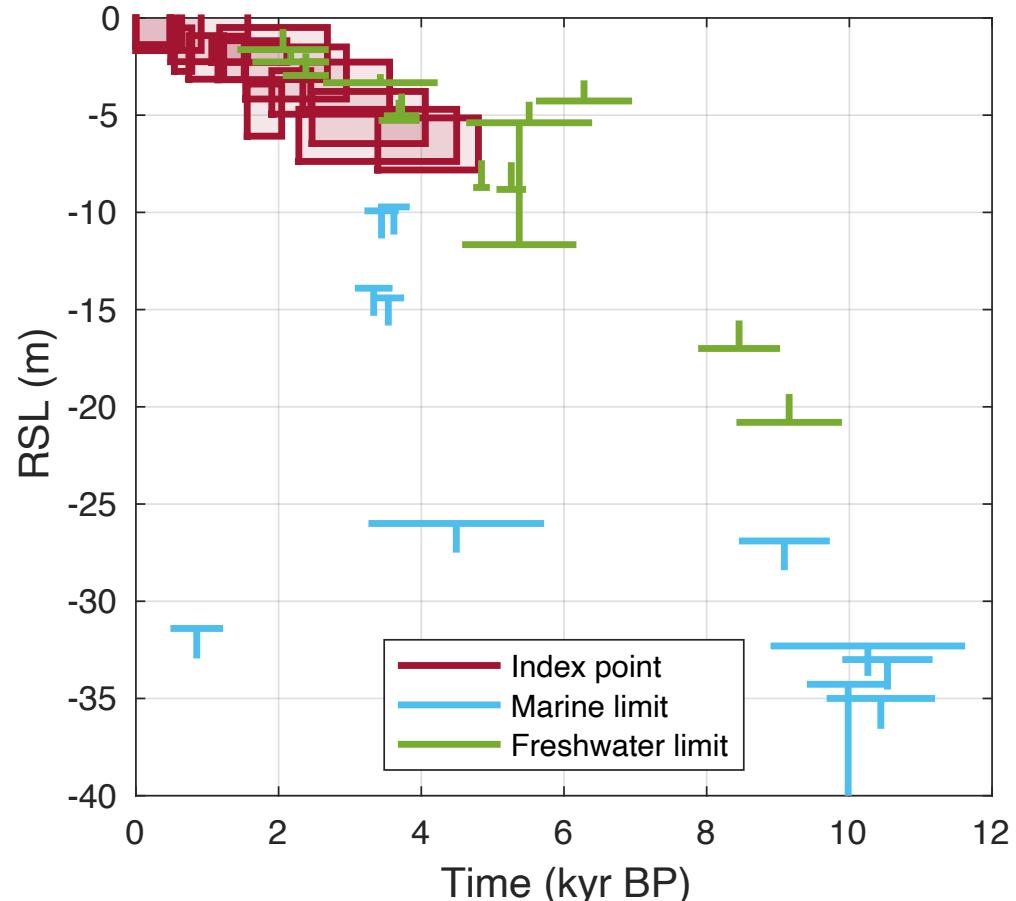
We solve it by minimizing the sum of the squared residuals

$$\sum_n (y_n - \alpha x_n - \beta)^2$$

OLS gives solutions  $\hat{\alpha}$  for  $\hat{\beta}$

$$\hat{\alpha} = \frac{\sum_n (x_n - \bar{x})(y_n - \bar{y})}{\sum_n (x_n - \bar{x})^2}$$

$$\hat{\beta} = \bar{y} - \hat{\alpha} \bar{x}$$



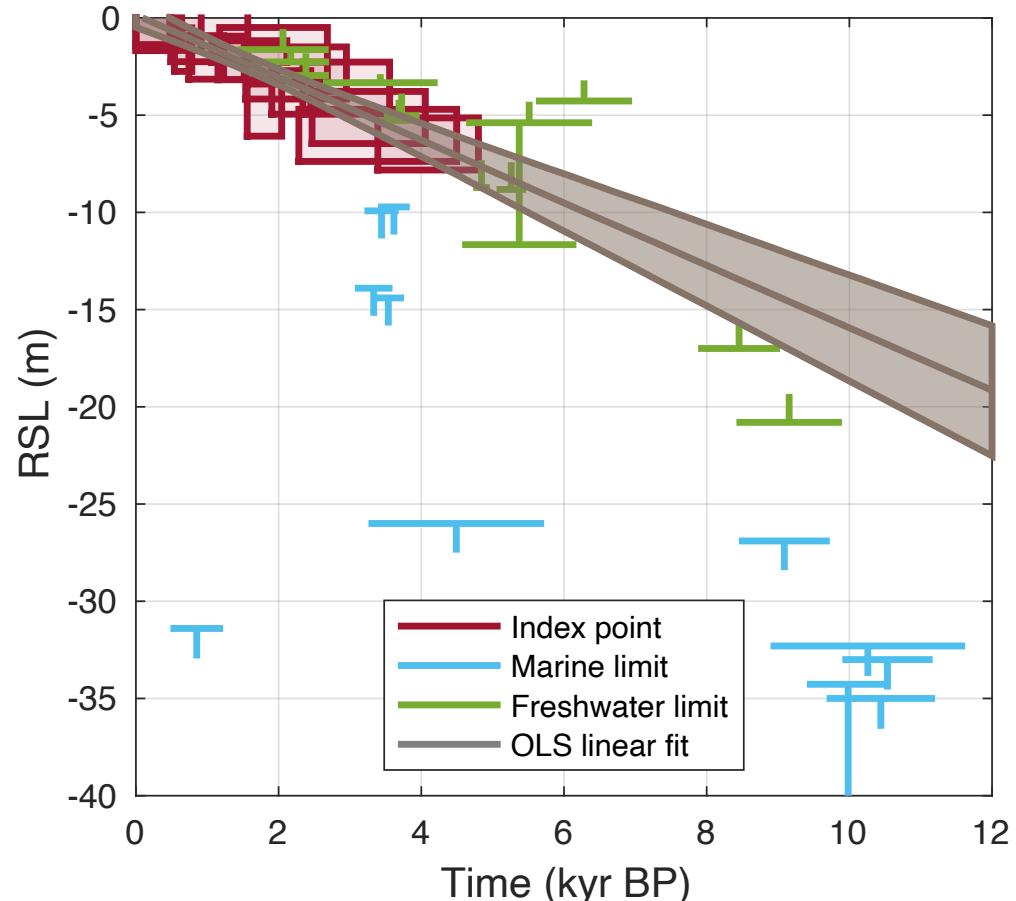
# We have uncertain proxies; we want past rates of change

When we do this, we get the  
value  $\hat{\alpha} = 1.6 \pm 0.3 \text{ mm yr}^{-1}$

But what about the error bars  
on the observations?

Do we have to prescribe a  
particular functional form?

And how do we incorporate  
limiting or bounding data?



**Think probabilistically**

**Think hierarchically**

**Think conditionally**

**Observations** are conditioned on **processes**

**Processes** are conditioned on **parameters**

**Everything is uncertain!**

# Probability postulates

Say that  $S$  is the set (“sample space”) of all possible events and  $A$  is an event (“a subset of  $S$ ”)

## Postulate #1

- The probability of  $A$  is a non-negative real number

$$P(A) \geq 0$$

## Postulate #2

- Something must happen

$$P(S) = 1$$

## Postulate #3

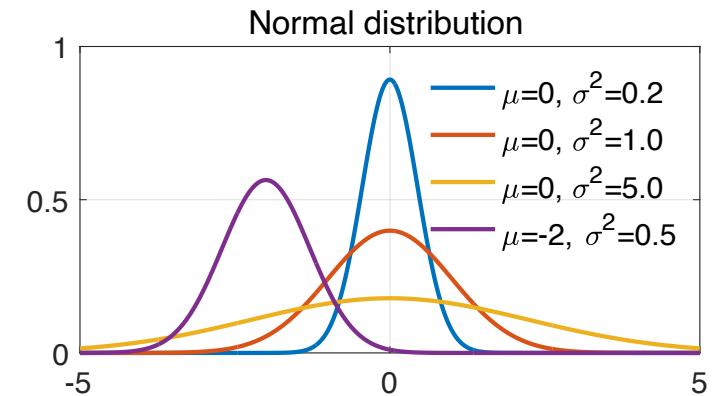
- If  $A, B, C, \dots$  are mutually exclusive events, then

$$P(A \cup B \cup C \cup \dots) = P(A) + P(B) + P(C) + \dots$$

# Some standard distributions

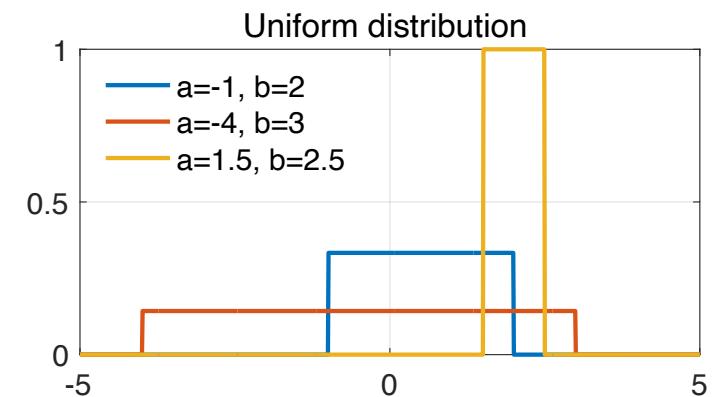
## Normal

$$\bullet P(x) = \frac{1}{\sigma\sqrt{2\pi}} e^{-\frac{1}{2}\left(\frac{x-\mu}{\sigma}\right)^2}$$



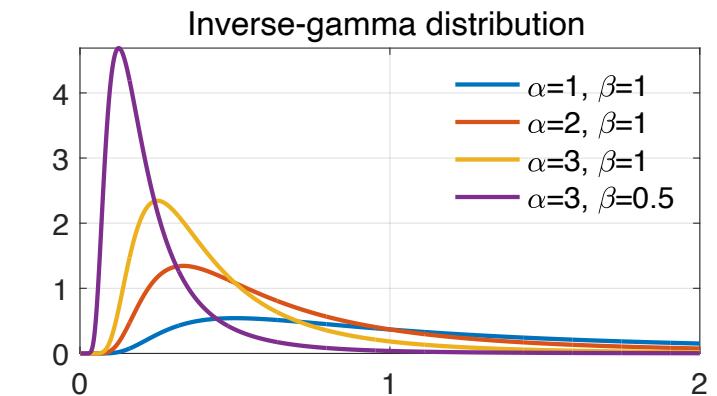
## Uniform

$$\bullet P(x) = \begin{cases} \frac{1}{b-a}, & x \in [a, b] \\ 0, & x \notin [a, b] \end{cases}$$



## Inverse-gamma

$$\bullet P(x) = \frac{\beta^\alpha}{\Gamma(\alpha)} x^{-\alpha-1} e^{-\beta/x}$$



# Probability distributions

Marginal distribution

- $P(A)$  is the probability that  $A$  occurs

Joint distribution

- $P(A, B)$  is the probability that  $A$  and  $B$  both occur

Conditional distribution

- $P(A|B)$  is the probability that  $A$  occurs given that  $B$  is certain to occur (or already occurred)

These three distributions are related according to

$$P(A|B) = P(A, B)/P(B)$$

# Bayes theorem

## Bayes theorem

$$P(A|B) = P(B|A)P(A)/P(B) \propto P(B|A)P(A)$$

↑      ↑      ↗  
Posterior Likelihood Prior Normalizing constant

Example: what is the temperature  $\theta$  outside?

- Prior (climatology)  $p(\theta) = N(\mu, \sigma^2); \mu = 17^\circ\text{C}, \sigma = 2^\circ\text{C}$
- Likelihood (data)  $p(T|\theta) = N(\theta, \delta^2); T = 20^\circ\text{C}, \delta = 1^\circ\text{C}$
- Posterior (temperature estimate  $\theta$  given the data  $T$ )

$$p(\theta|T) = N\left(\frac{\frac{\mu}{\sigma^2} + \frac{T}{\delta^2}}{\frac{1}{\sigma^2} + \frac{1}{\delta^2}}, \frac{1}{\frac{1}{\sigma^2} + \frac{1}{\delta^2}}\right) = N(19.4^\circ\text{C}, 0.64^\circ\text{C}^2)$$

**Back to the data**

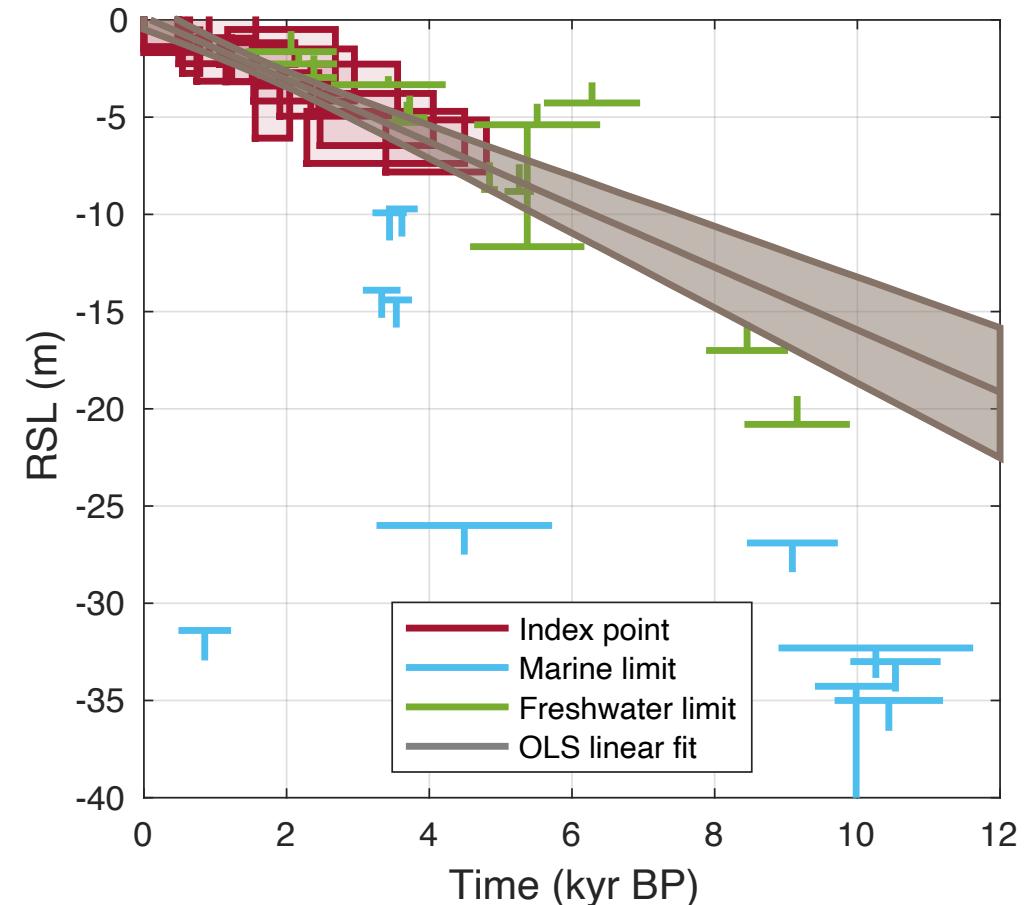
# Let's think about what we did earlier

When we got the OLS slope, we did two things without realizing

1. We made a **modeling** choice
2. We made an **analysis** choice

What if we make different modeling and analysis choices?

Does that allow us to address the questions we had before?



# A series of four Bayesian models

Start—  
Ordinary least squares;  
parametric;  
perfect data;  
without limiting dates



End—  
Bayesian;  
nonparametric;  
uncertain data;  
with limiting dates

**First model: linear/parametric process,  
perfect data, without limiting data**

# Model 1: linear/parametric process, perfect observations, without limiting data

Let's use the same model but  
change the analysis approach

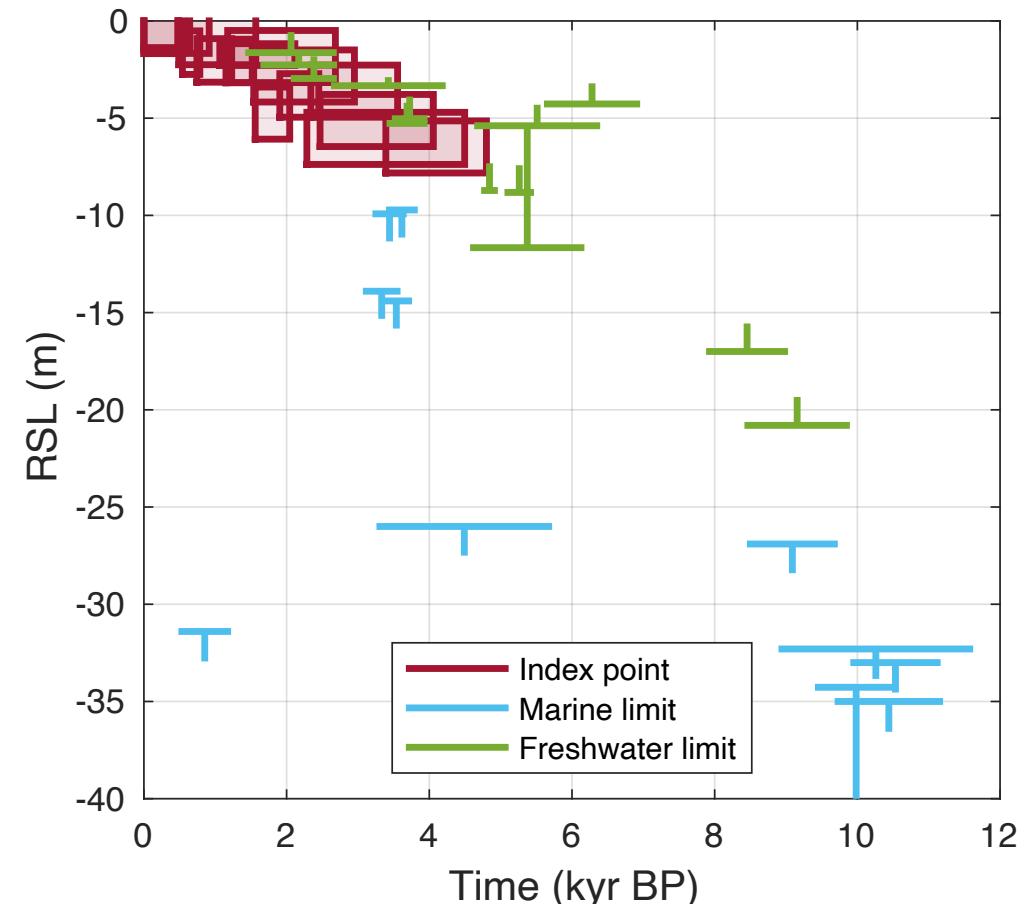
We want the **posterior** distribution

$$p(\alpha, \beta, \gamma^2 | x_1, \dots, y_1, \dots)$$

To get that, Bayes rule says that we  
need **likelihood & prior** distributions

The likelihood is our linear model,  
now written in a probabilistic form

$$p(y_n | x_n, \alpha, \beta, \gamma^2) = N(\alpha x_n + \beta, \gamma^2)$$



# Model 1: linear/parametric process, perfect observations, without limiting data

The prior distributions express our belief about the parameters before we consider the data; let's choose

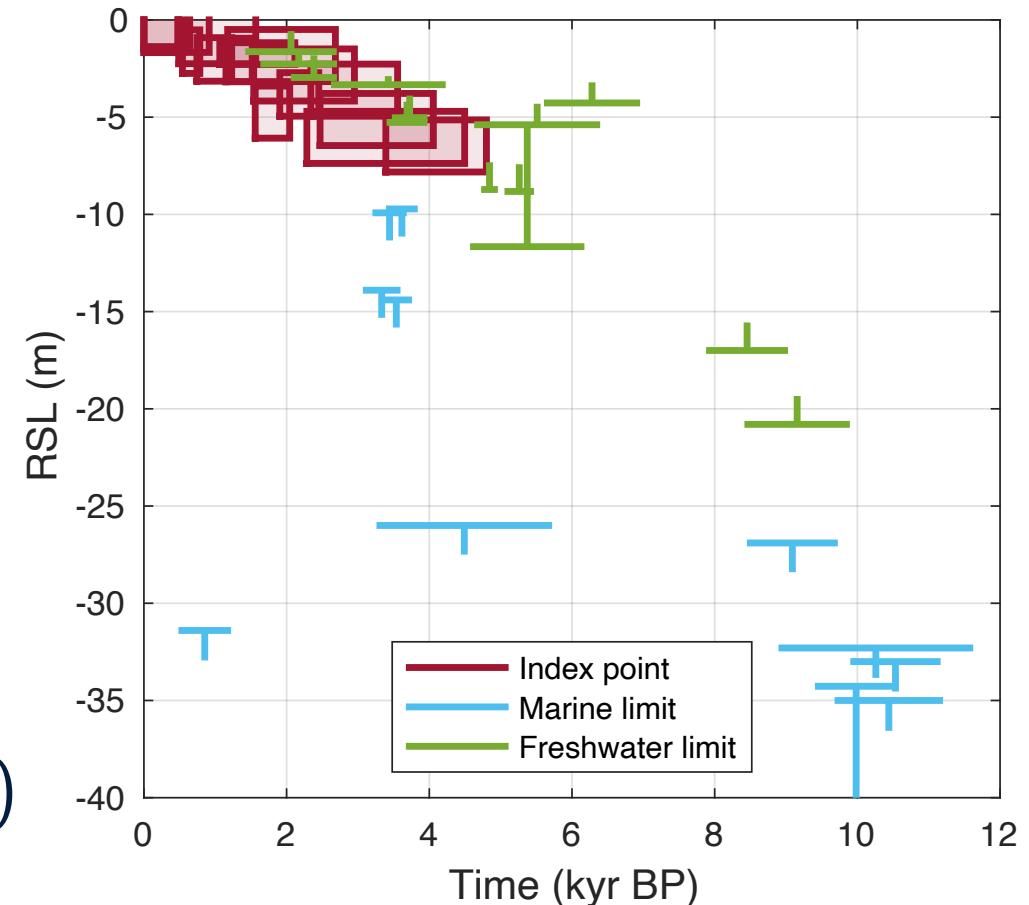
$$p(\alpha) = N(\tilde{\eta}_\alpha, \tilde{\zeta}_\alpha^2)$$

$$p(\beta) = N(\tilde{\eta}_\beta, \tilde{\zeta}_\beta^2)$$

$$p(\gamma^2) = G^{-1}(\tilde{\xi}_\gamma, \tilde{\chi}_\gamma)$$

Bringing the pieces together, we find

$$\begin{aligned} p(\alpha, \beta, \gamma^2 | x_1, \dots, y_1, \dots) \\ \propto p(\alpha)p(\beta)p(\gamma^2) \prod_n p(y_n | x_n, \alpha, \beta, \gamma^2) \end{aligned}$$



# Model 1: linear/parametric process, perfect observations, without limiting data

We **cannot** sample this joint posterior,  
but we **can** sample the full conditional  
posteriors of each of the parameters

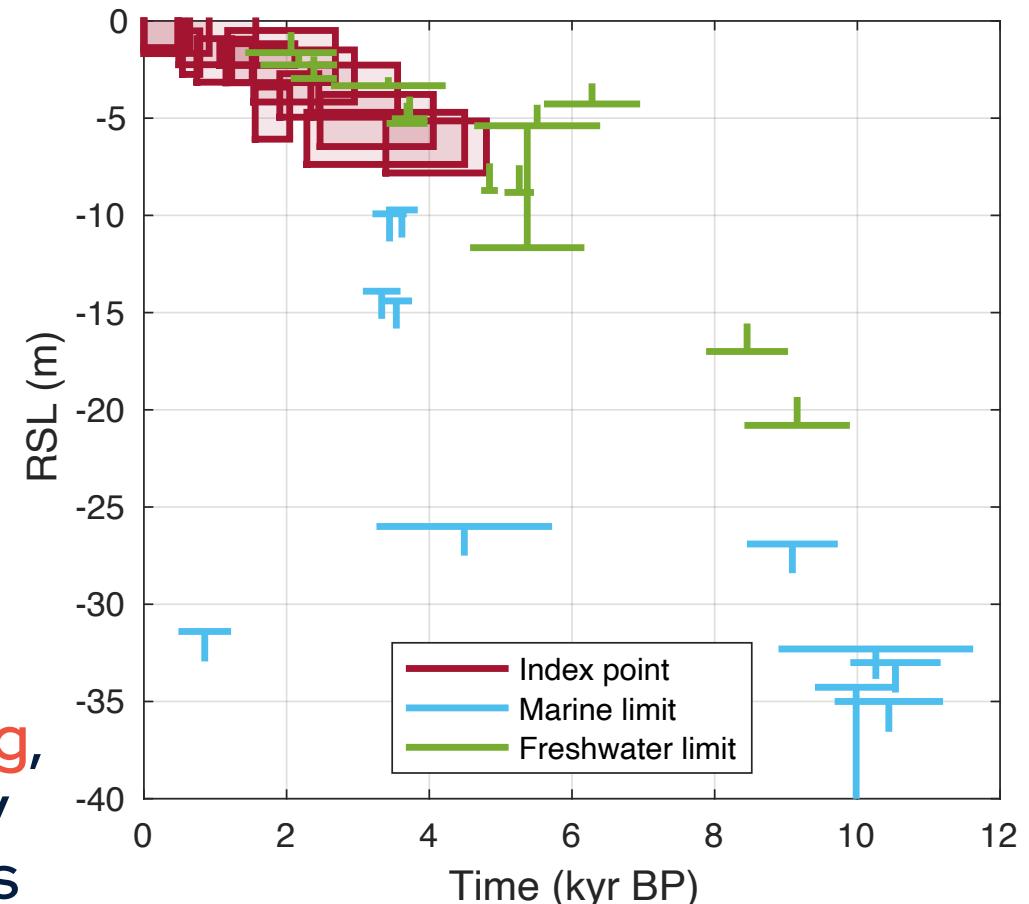
For example, for  $\alpha$ , we have

$$p(\alpha | \cdot) = N(\psi_\alpha V_\alpha, \psi_\alpha)$$

$$V_\alpha = \tilde{\zeta}_\alpha^{-2} \tilde{\eta}_\alpha + \gamma^{-2} \sum_{k=1}^n x_k (y_k - \beta)$$

$$\psi_\alpha = 1 / (\tilde{\zeta}_\alpha^{-2} + \gamma^{-2} \sum_{k=1}^n x_k^2)$$

This forms the basis of **Gibbs sampling**,  
which is a powerful way to numerically  
sample from known distribution forms



# Model 1: linear/parametric process, perfect observations, without limiting data

## Gibbs sampler

First, make initial guesses for the parameters you seek  $\alpha_{(0)}, \beta_{(0)}, \gamma^2_{(0)}$

Second, iteratively sample from the full conditional posterior distributions

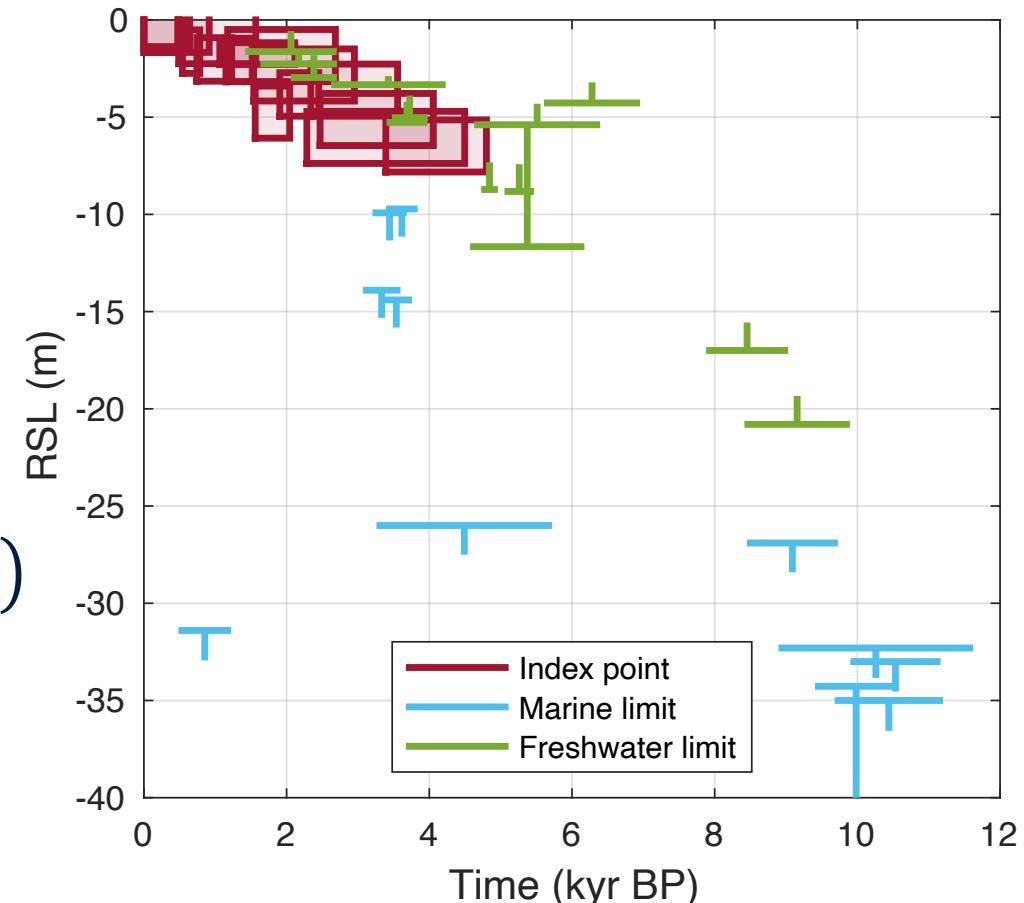
for  $m = 1, \dots, M$

draw  $\alpha_{(m)}$  from  $p(\alpha | \beta_{(m-1)}, \gamma^2_{(m-1)}, \dots)$

draw  $\beta_{(m)}$  from  $p(\beta | \alpha_{(m)}, \gamma^2_{(m-1)}, \dots)$

draw  $\gamma^2_{(m)}$  from  $p(\gamma^2 | \alpha_{(m)}, \beta_{(m)}, \dots)$

end



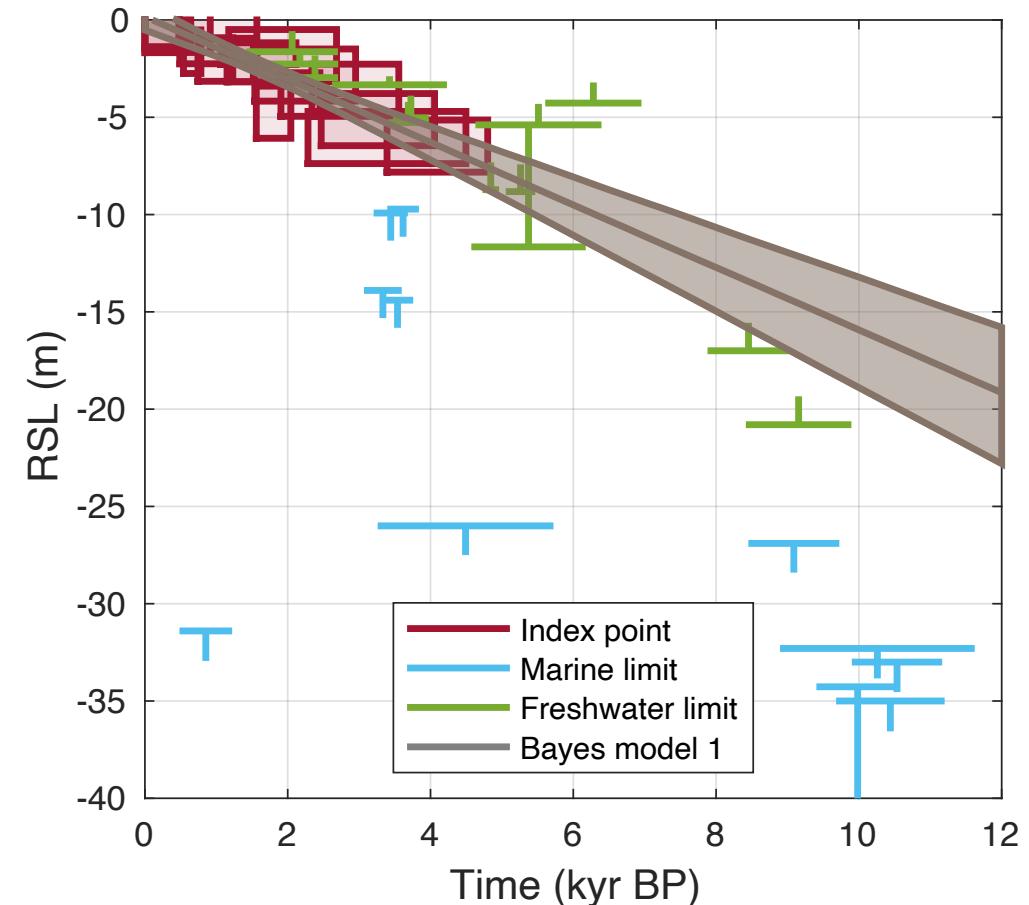
# Model 1: linear/parametric process, perfect observations, without limiting data

We choose wide uninformative priors

When we run the Gibbs sampler, we  
get  $\alpha = 1.6 \pm 0.3 \text{ mm yr}^{-1}$  (95% CI)

Did we just do a lot of work to get the  
same thing we had earlier with OLS?

Yes & no. You'll see the advantages  
of the Bayesian approach shortly.



**Second model: linear/parametric process,  
imperfect data, without limiting data**

# Model 2: linear/parametric process, imperfect observations, without limiting data

How do we factor in the errors  
on the proxy reconstructions?

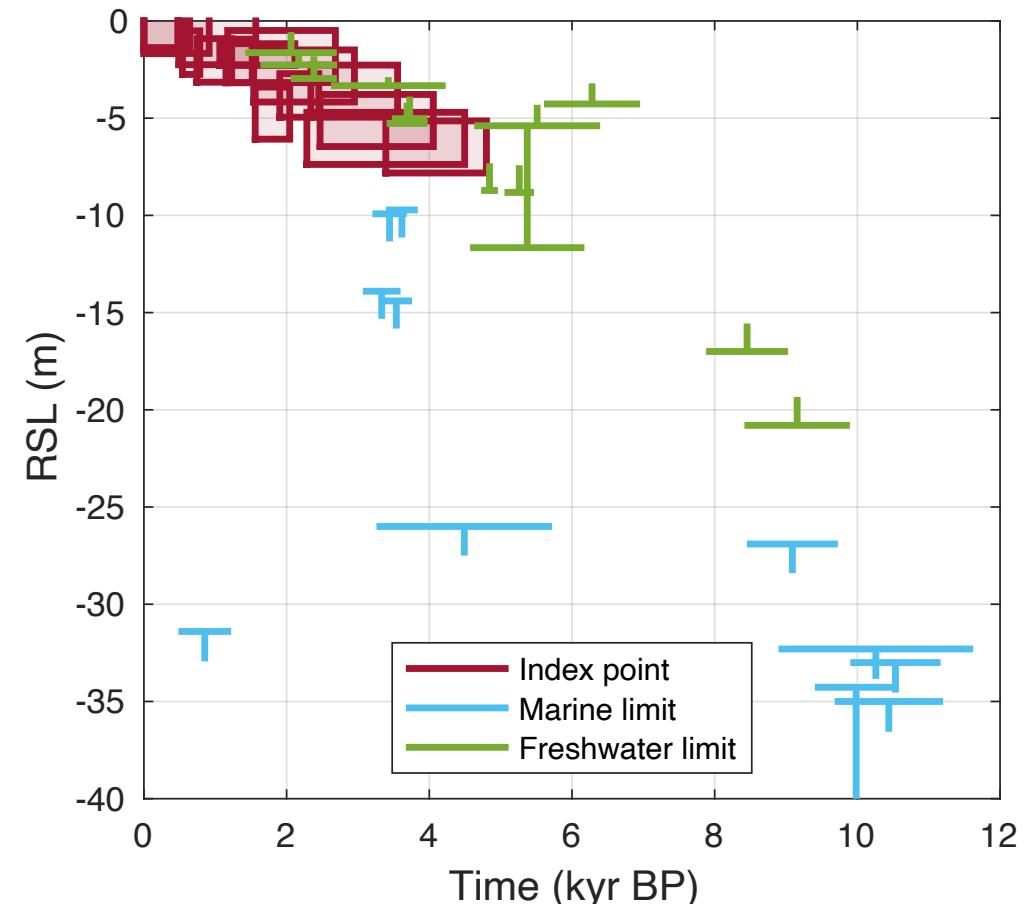
Earlier, we had a **process level**  
and a **prior level**; now let's add  
more equations in a **data level**

Say that index points are noisy  
versions of the true processes

$$p(z_n) = N(y_n, \delta_n^2)$$

$$p(w_n) = N(x_n, \epsilon_n^2)$$

where  $z_n, w_n$  are sea level, age  
data, and  $\delta_n^2, \epsilon_n^2$  are the errors



## Model 2: linear/parametric process, imperfect observations, without limiting data

All other equations are the same

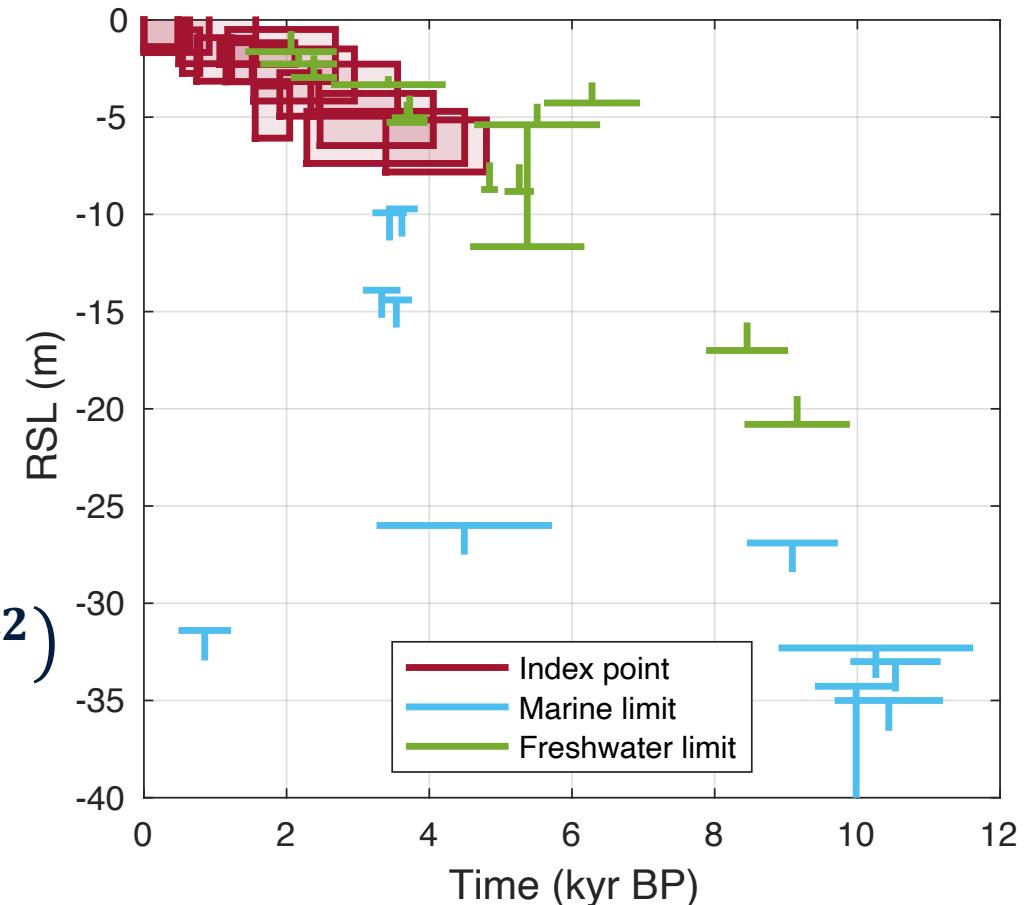
Now  $x_n$  and  $y_n$  are uncertain(!),  
so our posterior distribution is

$$p(\alpha, \beta, \gamma^2, x_1, \dots, y_1, \dots | w_1, \dots, z_1, \dots)$$

Using Bayes, & our prior, data, &  
process equations, this becomes

$$\propto p(\alpha)p(\beta)p(\gamma^2) \prod_n [p(y_n|x_n, \alpha, \beta, \gamma^2) \\ \times p(z_n|y_n) p(w_n|x_n)]$$

Draw solutions for  $\alpha, \beta, \gamma^2, x, y$   
again based on Gibbs sampling

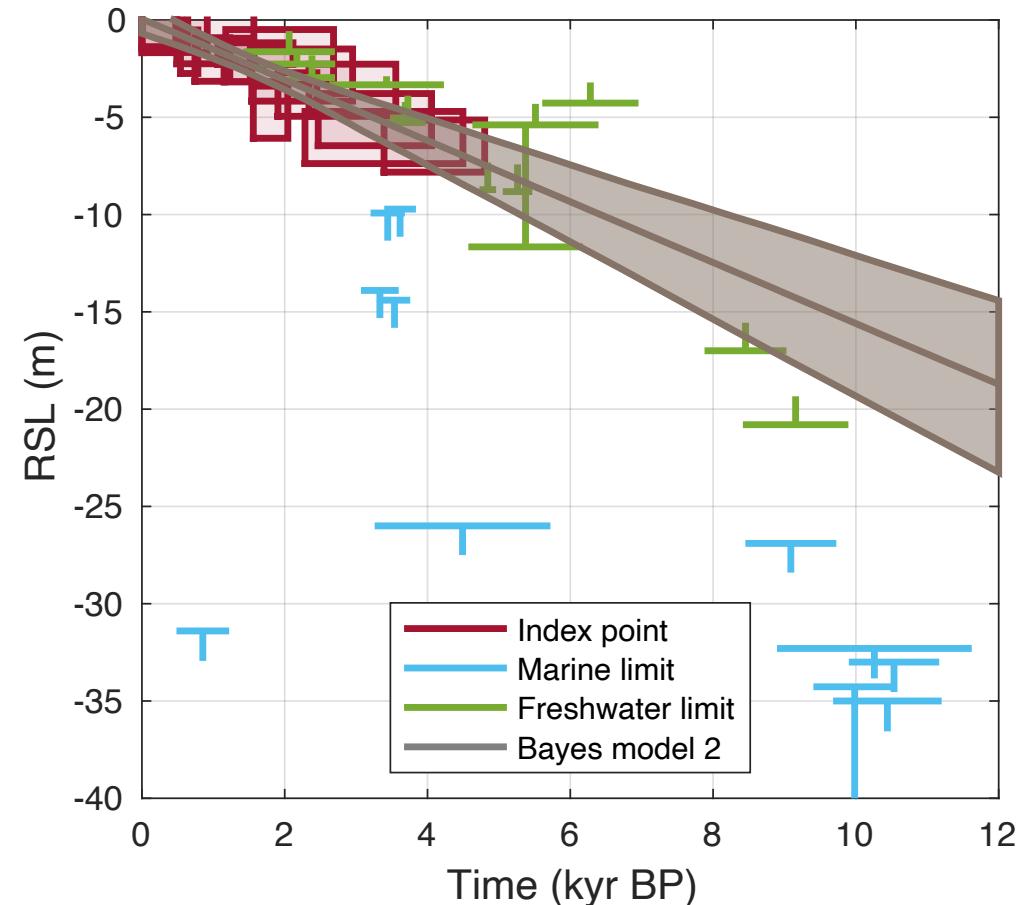


# Model 2: linear/parametric process, imperfect observations, without limiting data

Now we determine a sea level rate of  
 $\alpha = 1.6 \pm 0.4 \text{ mm yr}^{-1}$  (95% CI)

The error bars are now wider because  
we include observational uncertainty  
in the design of our Bayesian model

We arguably obtain a more realistic  
linear trend in sea level, but is a linear  
trend the most appropriate model?



**Third model: nonlinear/nonparametric process,  
imperfect data, without limiting data**

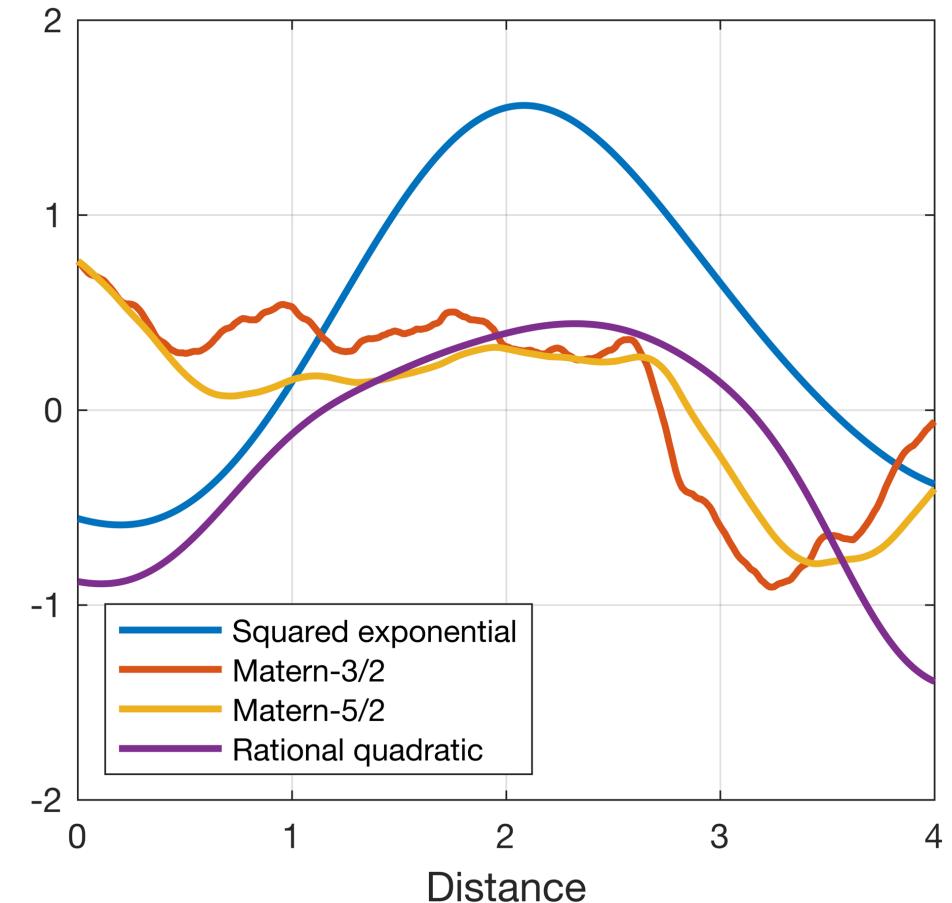
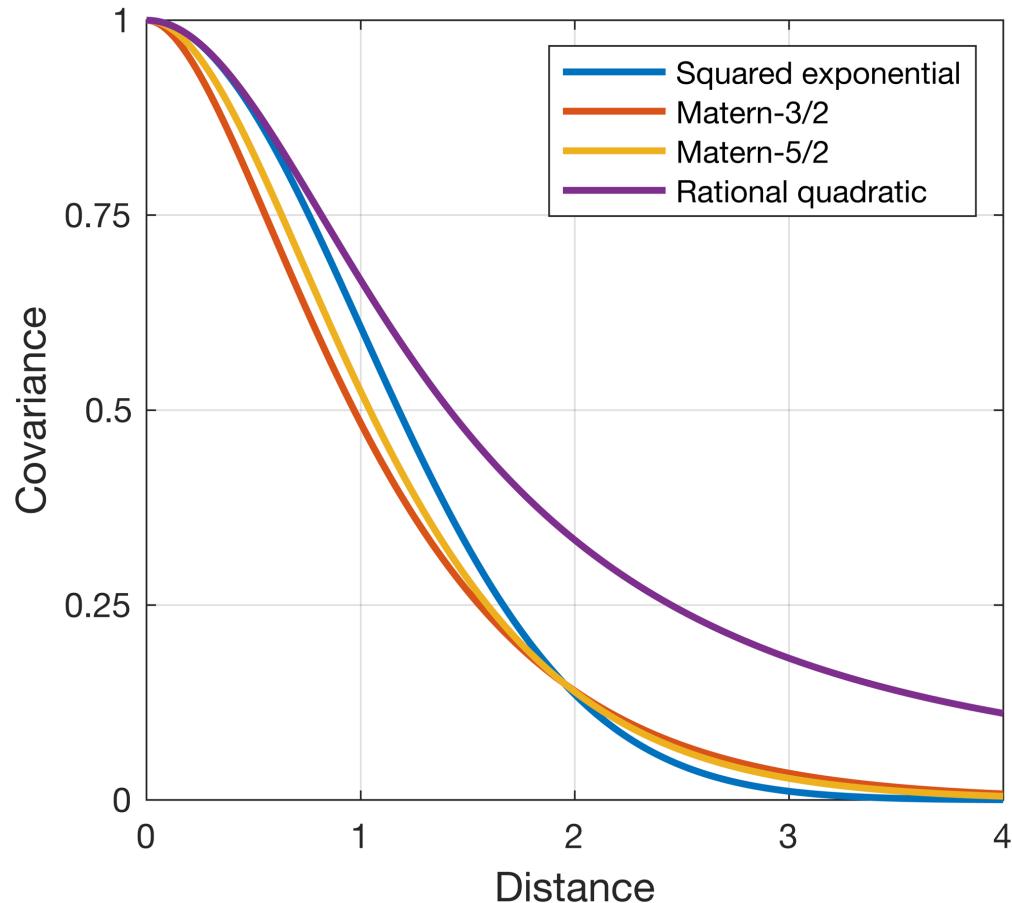
We must change how we view the process

We will still use a Gaussian distribution, but ...

We **were** putting all structure in the **mean**;  
now we'll put the structure in the **variance**

We **were** prescribing a **functional form**;  
now we'll specify a **relation between points**

# Covariance functions (kernels)



# Model 3: nonlinear/nonparametric process, imperfect observations, without limiting data

We'll use the same data model, but write a new process model

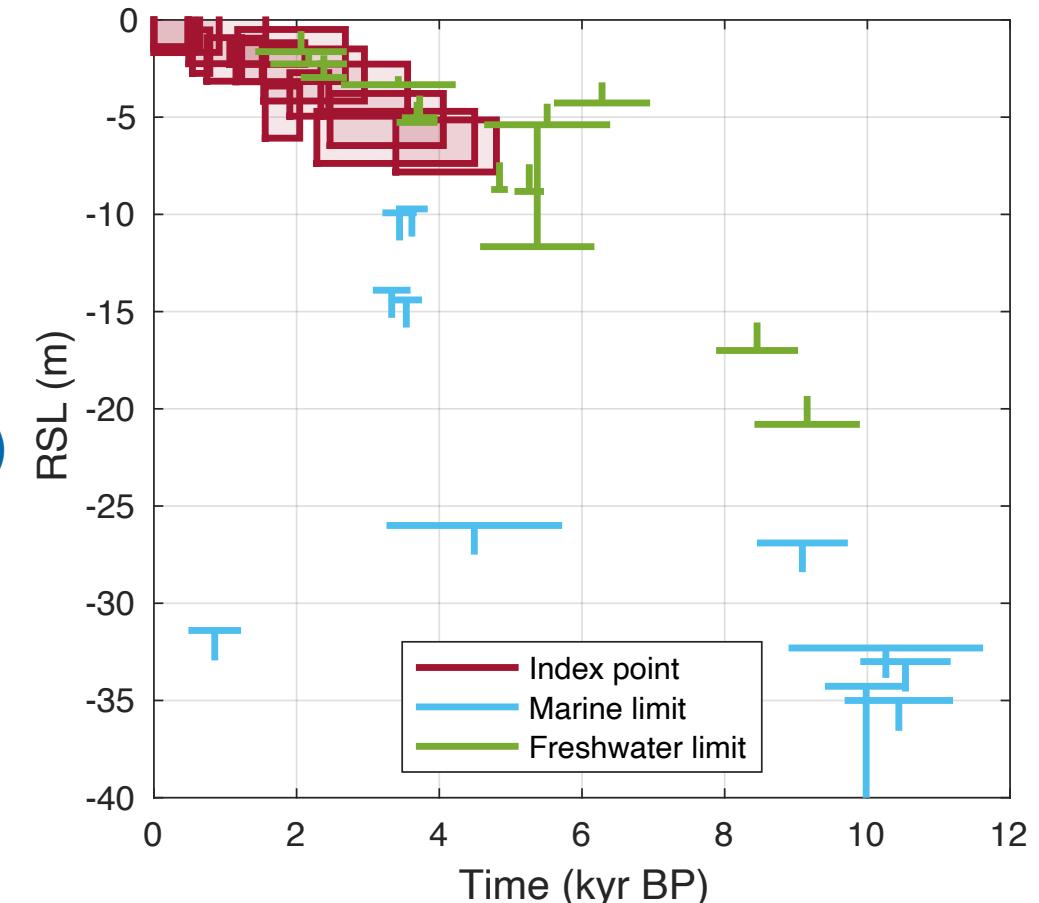
Let's represent sea level  $y$  as a random process with structure

$$y \sim N(0, L + G) \quad \text{local (uncorrelated)}$$

$$L_{jk} = \lambda^2 \delta(x_j, x_k) \quad \text{regional (correlated)}$$

$$G_{jk} = \rho^2 C(x_j, x_k, \varphi)$$

where  $\delta$  is the Kronecker delta and  $C$  the covariance function (compare to Khan et al. 2015)



# Model 3: nonlinear/nonparametric process, imperfect observations, without limiting data

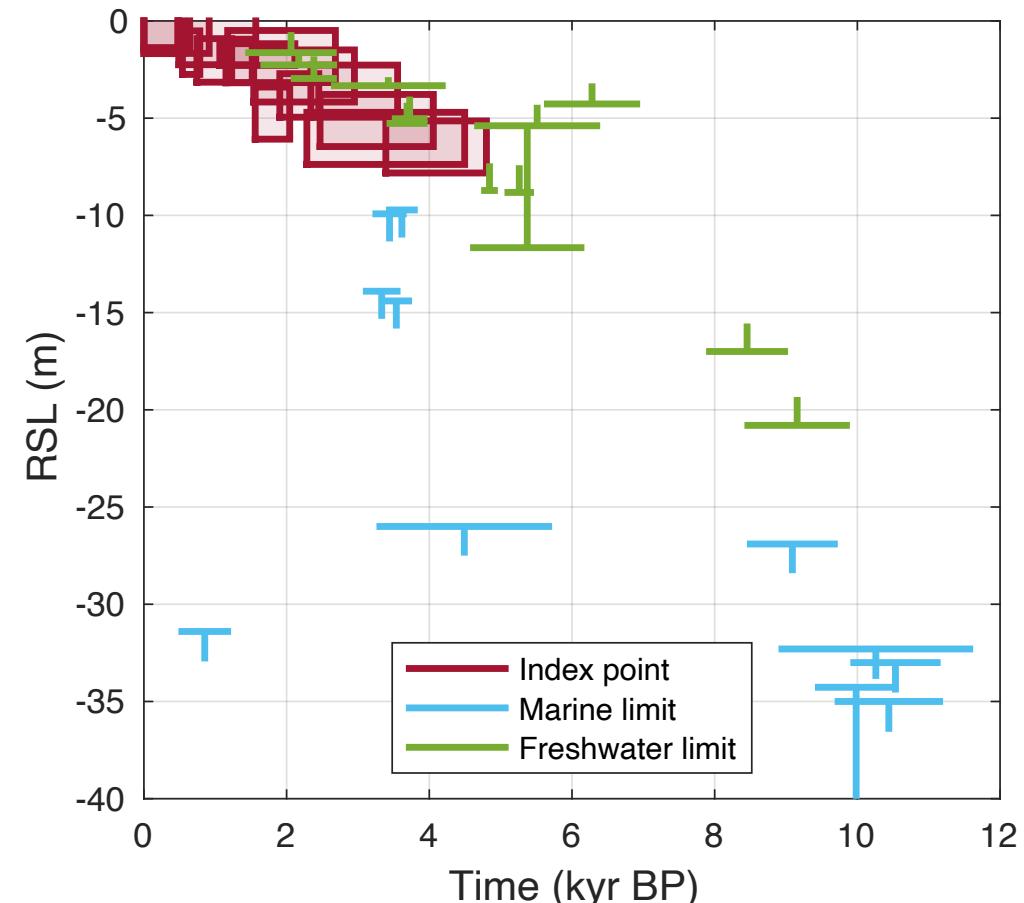
Place priors on  $\lambda^2, \rho^2, \varphi$

The new posterior distribution is

$$\begin{aligned} p(\lambda^2, \rho^2, \varphi, x_1, \dots, y_1, \dots | w_1, \dots, z_1, \dots) \\ \propto p(\lambda^2)p(\rho^2)p(\varphi) \prod_n [p(z_n|y_n) \\ \times p(w_n|x_n)p(y_n|x, y_{m \neq n}, \lambda^2, \rho^2, \varphi)] \end{aligned}$$

Posterior is complicated, and full conditional posteriors don't all have standard distributions

We can't use Gibbs sampler, but we can use Metropolis sampler



# Model 3: nonlinear/nonparametric process, imperfect observations, without limiting data

## Metropolis sampler for parameter $\vartheta$

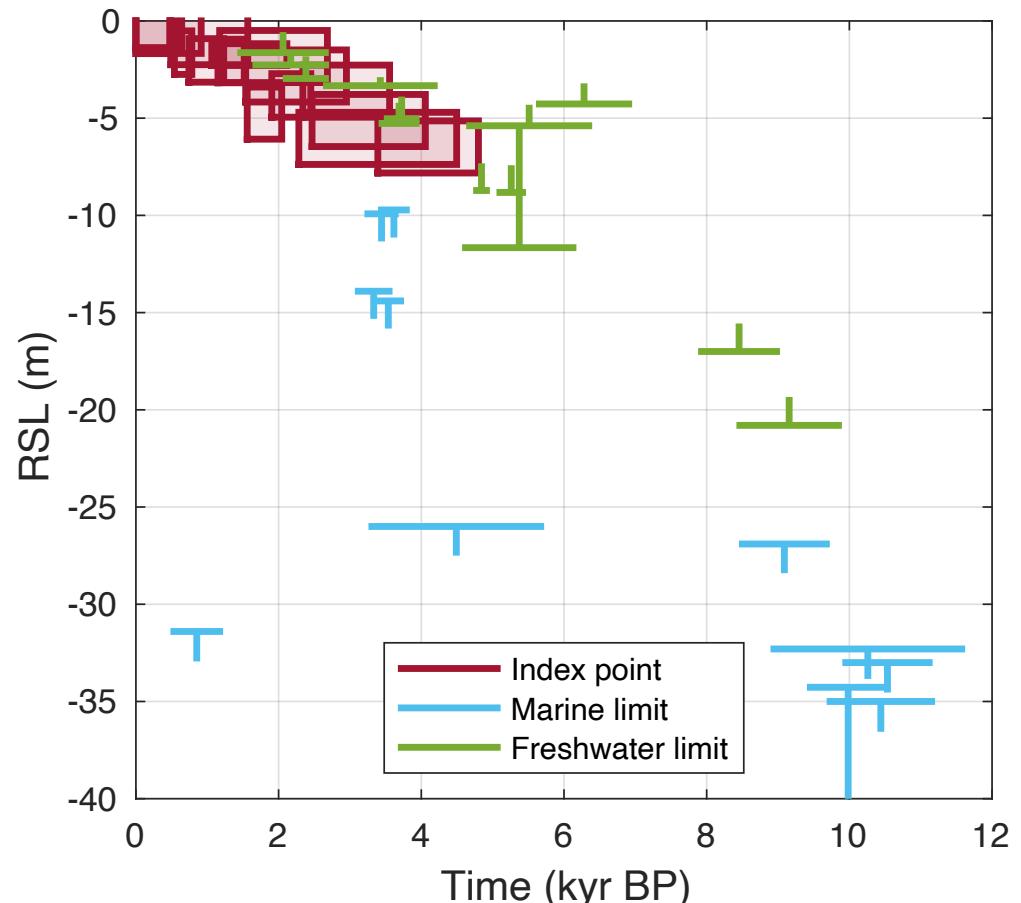
Goal is to sample a new estimate  $\vartheta^{new}$  based on the current estimate  $\vartheta^{now}$

1. Propose a new value  $\vartheta^* = \vartheta^{now} + \xi$  with  $\xi$  from a symmetric distribution

2. Then compute the ratio

$$R_M = \frac{P(\vartheta^* | \cdot)}{P(\vartheta^{now} | \cdot)}$$

3. Accept the proposition and define  $\vartheta^{new} = \vartheta^*$  with probability  $\min(1, R_M)$



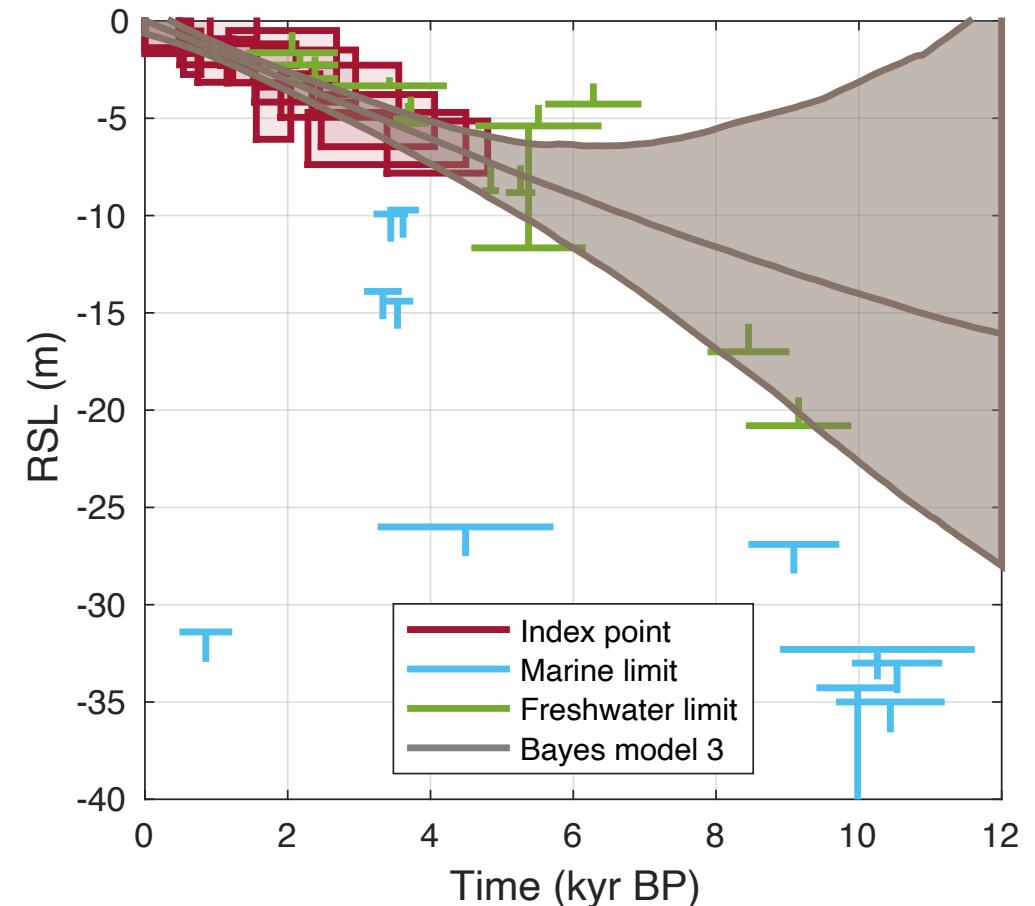
## Model 3: nonlinear/nonparametric process, imperfect observations, without limiting data

Now we determine a variable rate of sea-level change over the Holocene

Rates are  $1.5 \pm 0.6$ ,  $1.5 \pm 0.7$ , and  $1.2 \pm 1.9 \text{ mm yr}^{-1}$  at 0, 4, and 8 kyr

Model more completely represents the structure of the index points, and the uncertainties characterizing them

But we're still omitting limiting dates!



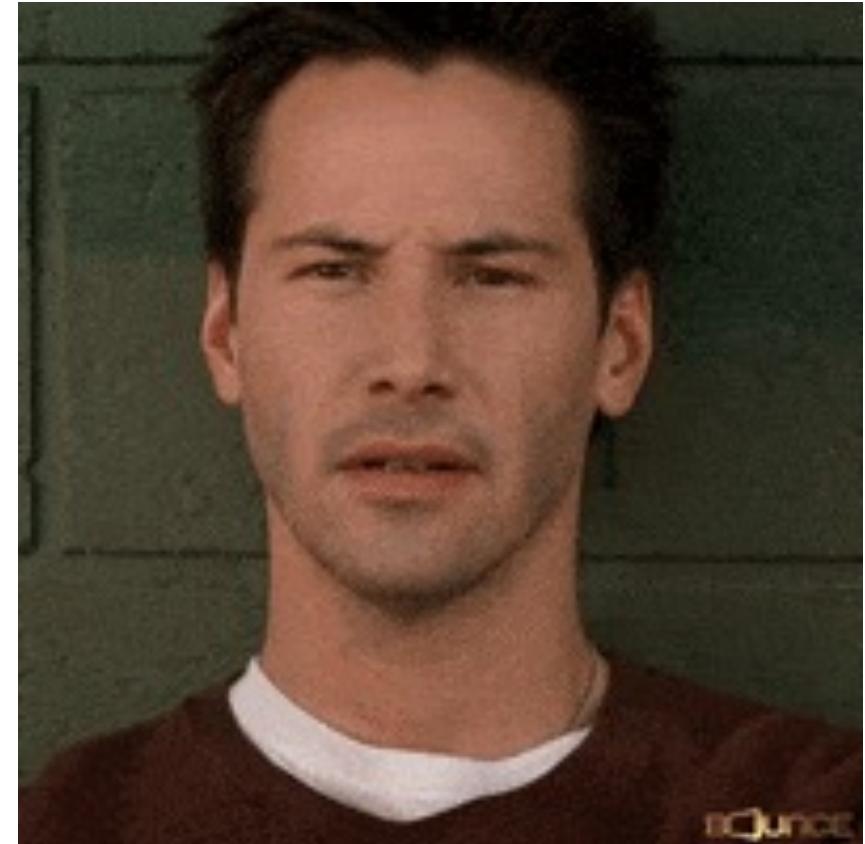
## Model 3: nonlinear/nonparametric process, imperfect observations, without limiting data

Hold up. We didn't specify a function for sea level, so how did I just estimate sea-level values at regular time points?

To do this, I used posterior prediction: knowing sea levels  $y$  and times  $x$ , I can estimate sea levels  $y^*$  at other times  $x^*$

$$p(y^*|x, y, x^*) = N \left( C(x^*, x) C(x, x)^{-1} y, \rho^2 C(x^*, x^*) - \rho^4 C(x^*, x) [C(x, x) + \lambda^2 I]^{-1} C(x, x^*) \right)$$

This follows from Gaussian identities and is at the core of Gaussian Process regression



**Fourth model: nonlinear/nonparametric  
process, imperfect data, with limiting data**

# Model 4: nonlinear/nonparametric process, imperfect observations, with limiting data

Last model! Almost there!

Adjust data equation and add process equations for marine limit  $m$  and freshwater limit  $f$

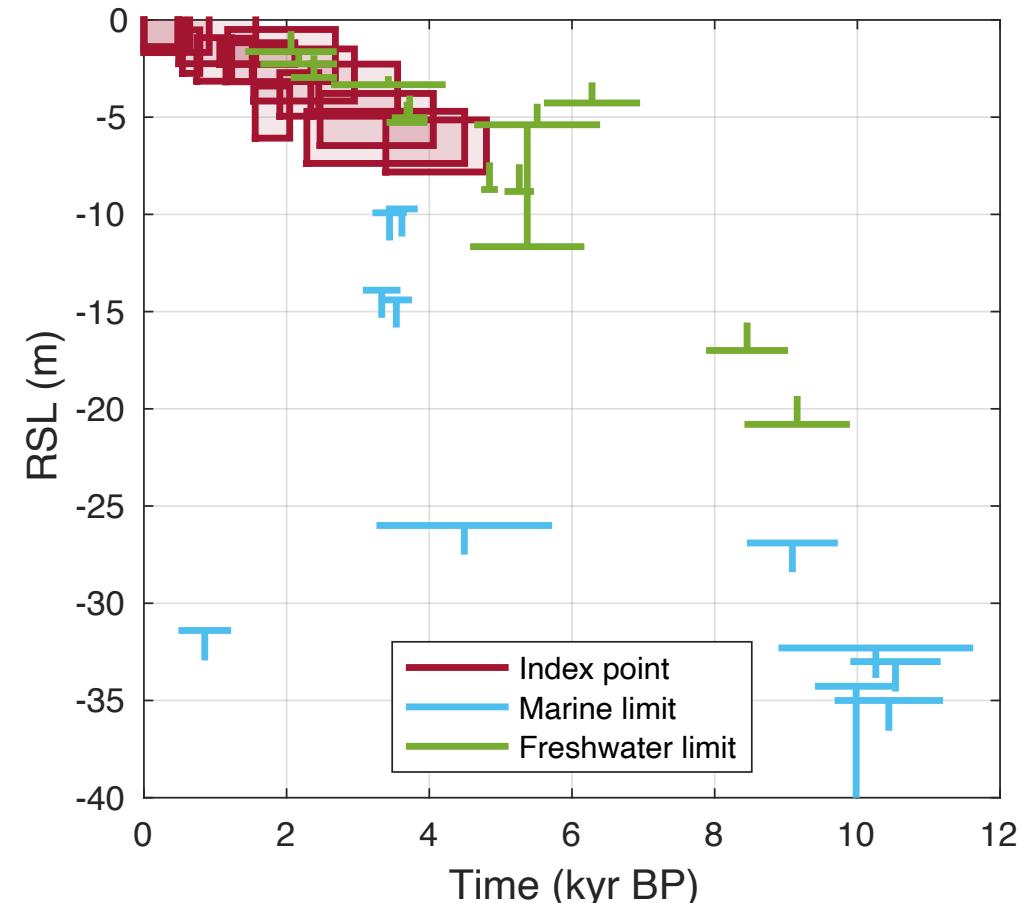
Data—

$$p(z_n) = \begin{cases} N(m_n, \delta_n^2) \\ N(y_n, \delta_n^2) \\ N(f_n, \delta_n^2) \end{cases}$$

Process—

$$p(m_n) = U(-\infty, y_n)$$

$$p(f_n) = U(y_n, +\infty)$$



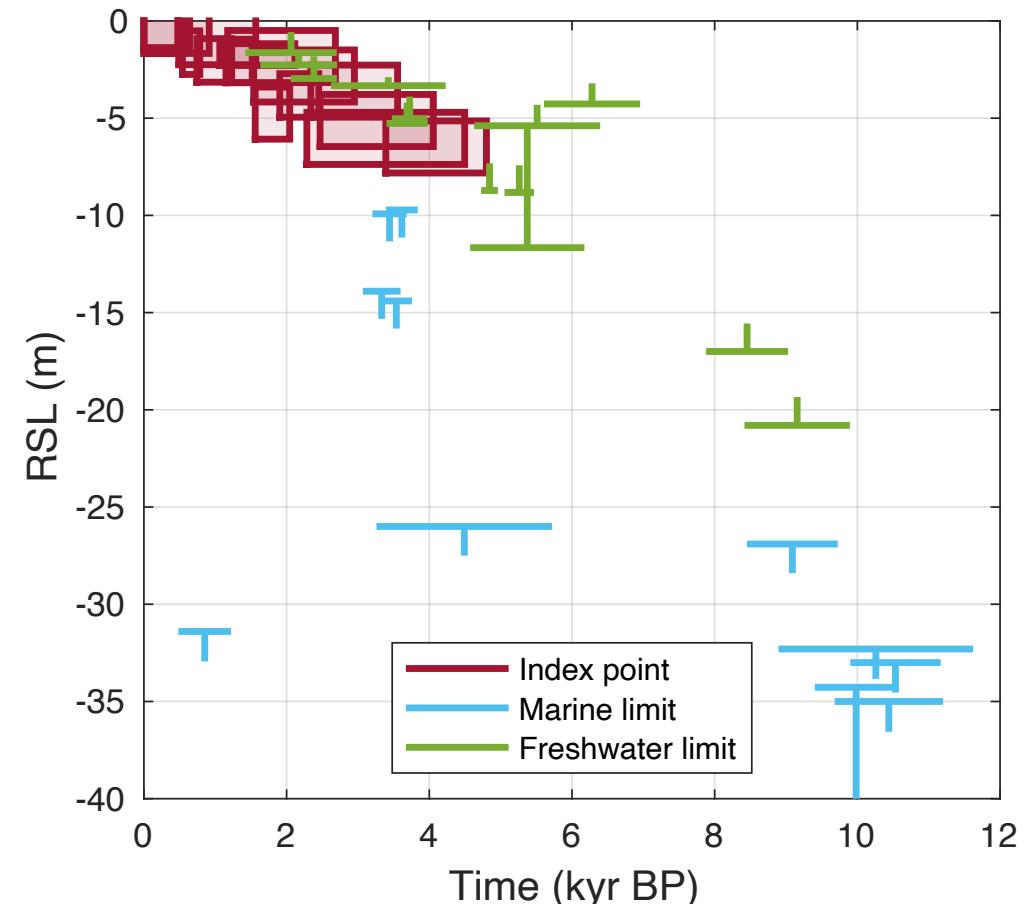
## Model 4: nonlinear/nonparametric process, imperfect observations, with limiting data

Final posterior distribution is

$$p(\lambda^2, \rho^2, \varphi, x_1, \dots, y_1, \dots, m_1, \dots, f_1, \dots | w_1, \dots, z_1, \dots)$$

$$\propto p(\lambda^2)p(\rho^2)p(\varphi) \prod_n [p(w_n|x_n) \\ \times p(z_n|y_n, m_n, f_n) p(m_n|y_n) \\ \times p(f_n|y_n) p(y_n|x, y_{m \neq n}, \lambda^2, \rho^2, \varphi)]$$

We use Metropolis algorithm and posterior prediction to generate solutions at regular points in time

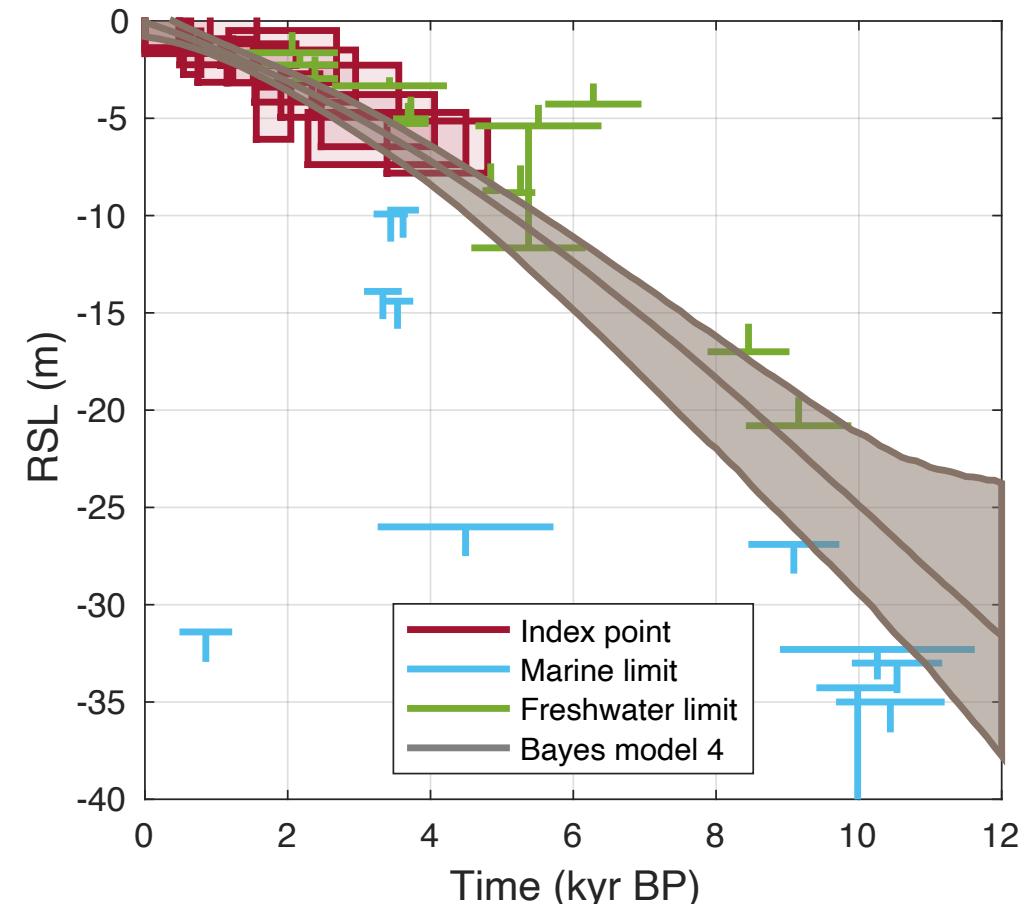


## Model 4: nonlinear/nonparametric process, imperfect observations, with limiting data

Now we determine variable rates of sea-level change over the Holocene that are consistent with all data types

Model shows higher rates earlier in time, which is more consistent with our physical intuition

Rates are  $1.3 \pm 0.8$ ,  $2.4 \pm 0.6$ , and  $3.1 \pm 0.9 \text{ mm yr}^{-1}$  at 0, 4, and 8 kyr

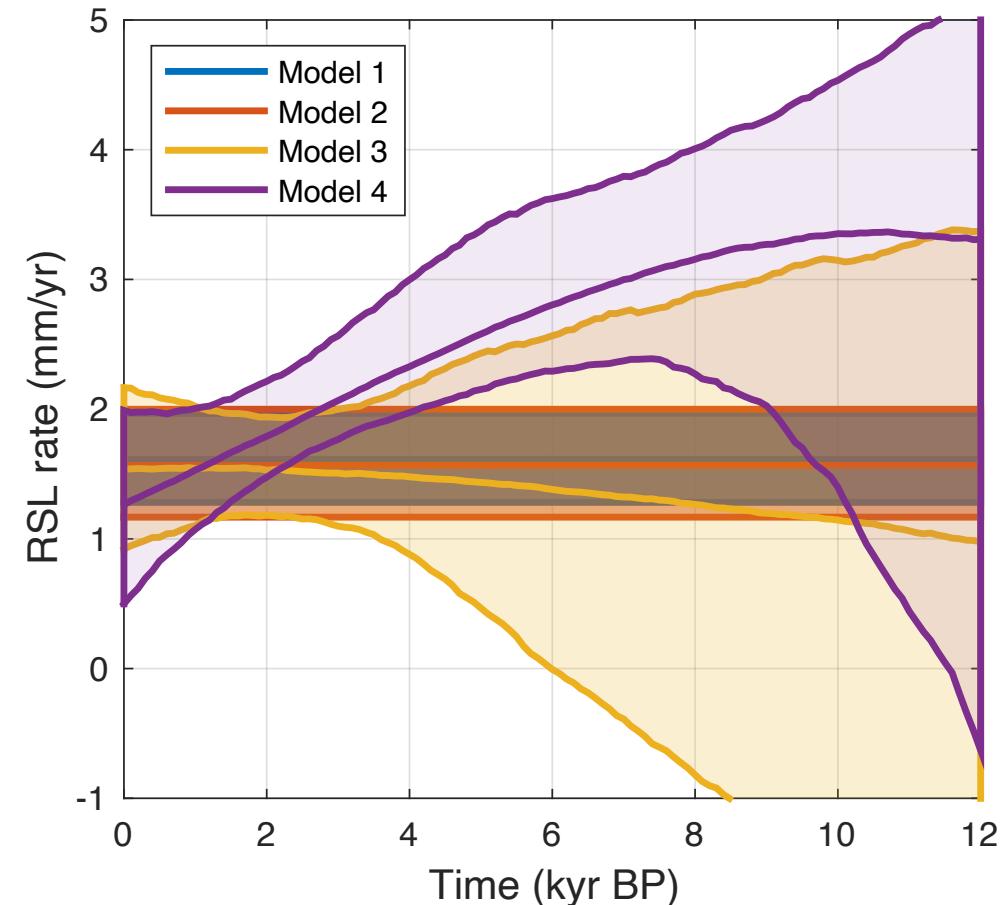


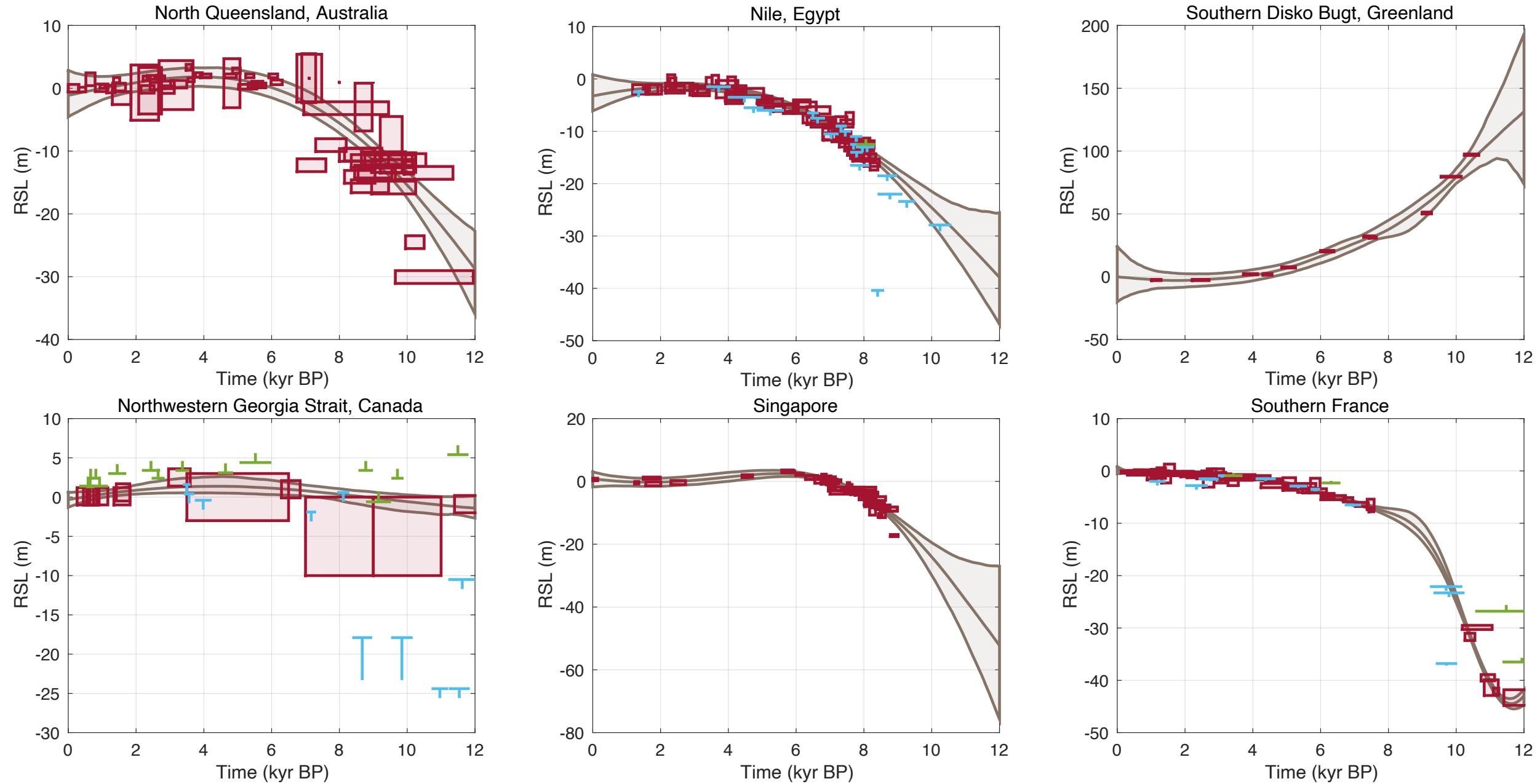
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# Summary

Bayesian modeling involves thinking about problems probabilistically, hierarchically, and conditionally

This approach offers a flexible framework for analyzing sea-level reconstructions and quantifying uncertainty

You don't need to be a statistician to understand these methods and use them informedly in your own research

# Summary

codes/slides available at

[github.com/christopherpiecuch/bayesGavle](https://github.com/christopherpiecuch/bayesGavle)

get/stay in touch

[cpiecuch@whoi.edu](mailto:cpiecuch@whoi.edu)

[twitter.com/chrisspiecuch](https://twitter.com/chrisspiecuch)

# Literature (freely available books)

**Cressie and Wikle (2019)**

<https://spacetimewithr.org/>

**Gelman et al. (2013)**

<http://www.stat.columbia.edu/~gelman/book/BDA3.pdf>

**Rasmussen and Williams (2006)**

<https://gaussianprocess.org/gpml/chapters/RW.pdf>

# Literature (a very incomplete list)

- Ashe et al. (2019) <https://www.sciencedirect.com/science/article/pii/S0277379118302130>
- Ashe et al. (2022) <https://ascmo.copernicus.org/articles/8/1/2022>
- Cahill et al. (2015) <https://www.jstor.org/stable/24522592>
- Cahill et al. (2016) <https://cp.copernicus.org/articles/12/525/2016>
- Creel et al. (2022) <https://doi.org/10.1016/j.quascirev.2022.107422>
- Hay et al. (2015) <https://www.nature.com/articles/nature14093>
- Hay et al. (2013) <https://www.pnas.org/doi/10.1073/pnas.1117683109>
- Khan et al. (2015) <https://doi.org/10.1007/s40641-015-0029-z>
- Khan et al. (2019) <https://doi.org/10.1016/j.quascirev.2019.07.016>
- Kopp et al. (2016) <https://www.pnas.org/doi/10.1073/pnas.1517056113>
- Kopp (2013) <https://agupubs.onlinelibrary.wiley.com/doi/full/10.1002/grl.50781>
- Kopp et al. (2009) <https://www.nature.com/articles/nature08686>
- Piecuch et al. (2018) <https://www.nature.com/articles/s41586-018-0787-6>

# Open-source code from others

**Erica Ashe** <https://github.com/ericaashe>

**Niamh Cahill** <https://github.com/ncahill89>

**Bob Kopp** <https://github.com/bobkopp>

**Maeve Upton** <https://maeveupton.github.io/reslr/>

A photograph of a large, light-colored boulder resting on a rocky ground. The background consists of many bare, leafless trees, suggesting a late autumn or winter setting. The lighting is natural, with sunlight filtering through the branches.

# Questions?