Slow particle remineralization, rather than suppressed disaggregation, likely drives efficient flux transfer through the Eastern Tropical North Pacific Oxygen Deficient Zone

11 May 2021

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# Key Points

* The upper mesopelagic of the oligotrophic Eastern Tropical North Pacific Oxygen Deficient Zone (ODZ) has low flux attenuation.
* Comparison of these observations to models suggests that the breakdown of particles of all sizes is slow throughout the ODZ.
* Zooplankton appear to transport organic matter into, and disaggregate particles within, the ODZ above 500 m.

# Abstract (limit 250 words)

Models and observations suggest that particle flux attenuation is lower across the mesopelagic zone of anoxic environments compared to oxic ones. Flux attenuation is controlled by microbial metabolism and aggregation and disaggregation by zooplankton, which also shape the relative abundance of different sized particles. Observing and modeling particle spectra can provide information about the contributions of these processes. We measured particle size spectrum profiles at one station in the oligotrophic Eastern Tropical North Pacific Oxygen Deficient Zone (ETNP ODZ) using an underwater vision profiler (UVP), a high-resolution camera that counts and sizes particles. Measurements were taken at different times of day, over the course of a week. Comparing these data to particle flux measurements from sediment traps collected over the same time-period allowed us to constrain the particle size to flux relationship, and to generate highly resolved depth and time estimates of particle flux rates. We found that particle flux attenuated very little throughout the anoxic water column, and at some time-points appeared to increase. Comparing our observations to model predictions suggested that particles of all sizes remineralize more slowly in the ODZ than in oxic waters, and that large particles disaggregate into smaller particles, primarily between the base of the photic zone and 500 m. Acoustic measurements of multiple size classes of organisms suggested that many organisms migrated, during the day, to the region with high particle disaggregation. Our data suggest that diel-migrating organisms both actively transport biomass and disaggregate particles in the ODZ core.

# Plain Language Summary (limit 200 words)

Marine snow are microscopic particles that form in the surface of the ocean and sink into the deep ocean. Most of these particles are the remains of dead algae and feces of tiny animals (zooplankton). The deeper the particles sink into the ocean before microbes or animals eat them, the longer it takes before the carbon in those particles can return to the atmosphere. In parts of the ocean where there is very little oxygen more particles sink to greater depths for reasons that are not well-understood. We used an underwater camera to observe marine snow particles in a part of the ocean (just west of Mexico) where there is very limited oxygen at depth. We compared the observations to predictions from different computer simulations to see which simulations were most accurate. Our measurements suggest that one reason that particles sink to deeper depths here is because microbes consume the particles slowly when there is no oxygen. Meanwhile, tiny animals break large particles into smaller ones and produce fecal pellets in these low oxygen waters.

# Introduction

The biological pump, in which sinking particles transport carbon from the surface into the deep ocean, is a key part of the global carbon cycle (Neuer et al., 2014; Turner, 2015). Organic matter flux into the deep ocean is a function both of export from the photic zone into the mesopelagic (export flux), and the fraction of that flux that crosses through the mesopelagic (transfer efficiency) (Francois et al., 2002; Passow & Carlson, 2012; Siegel et al., 2016). The transfer efficiency of the biological pump may affect global atmospheric carbon levels (Kwon & Primeau, 2008). Thus, understanding the processes that shape organic matter degradation in the mesopelagic is critical.

Zooplankton modulate carbon flux through the mesopelagic (Jackson & Burd, 2001; Steinberg & Landry, 2017; Turner, 2015), and by extension the efficiency of the biological pump (Archibald et al., 2019; Cavan et al., 2017). They affect particle flux through four processes: *repackaging*, *respiration*, *active transport* and *disaggregation*. Zooplankton *repackage* particles into fecal pellets that have different properties from the original particles they ingest (Wilson et al., 2008). The repackaged particles may be egested at greater depth in the water column during the zooplankton's migration, resulting in the *active transport* of organic carbon over depth (Archibald et al., 2019; Bianchi et al., 2013; Hannides et al., 2009; Steinberg et al., 2000; Stukel et al., 2018, 2019). In addition to repackaging, zooplankton may *respire* some proportion of the particles’ biomass consume particles in the mesopelagic and *respire* some of their biomass (Stukel et al., 2019). It was found that this rate of consumption did not substantially attenuate flux in the California Current system (Stukel et al., 2019). Zooplankton break large particles into smaller ones, likely by generating turbulence when they swim (Dilling & Alldredge, 2000; Goldthwait et al., 2005). This *disaggregation* can lead to increased remineralization of particles because those smaller particle pieces sink more slowly and so have longer residence times in the mesopelagic, causing them to be consumed before reaching deep waters (Goldthwait et al., 2005). This fragmentation has been shown in some cases to explain around 50% of flux attenuation (Briggs et al., 2020). In this manuscript we focus on active transport and disaggregation in particular.

Oxygen concentrations, and in particular the geographic and depth range of anoxic regions of the water column, appear to modulate particle flux through the mesopelagic. Observations of particle flux in the region of the Eastern Tropical North Pacific (Hartnett & Devol, 2003; Van Mooy et al., 2002; Weber & Bianchi, 2020) that is near the Mexican coast, the Eastern Tropical South Pacific (Pavia et al., 2019), and Arabian Sea (Keil et al., 2016) have suggested lower flux attenuation in these ODZ systems. Models have shown that accounting for oxygen limitation in ODZs is necessary to fit global patterns of particle transfer (Cram et al., 2018; Timothy DeVries & Weber, 2017). Analysis of remineralization tracers also shows evidence of slow flux attention in the ODZs (Weber & Bianchi, 2020). The oxygen content of the ocean is decreasing (Breitburg et al., 2018), and the spatial extent and depth range of ODZs, including the ETNP ODZ are likely to change, though there is disagreement over whether they are expanding or undergoing natural fluctuation (Deutsch et al., 2014; Horak et al., 2016; Stramma et al., 2008). Changes to ODZ ranges are likely to affect ocean chemistry, the habitat of marine organisms, and the interactions between organisms and chemistry (Gilly et al., 2013). Models and chemical data suggest that ODZs may enhance carbon transport to the deep ocean, by inhibiting microbial degradation of sinking marine particles (Cram et al., 2018). However, biological organic matter transport is also modulated by zooplankton (Steinberg et al., 2008; Steinberg & Landry, 2017) which feed on, produce and disaggregate particles, and whose interactions on particle flux in pelagic ODZs are only beginning to be explored (Kiko et al., 2020).

Models of particle transfer through the mesopelagic oceans predict that particle size and mineral content, ocean temperature, and oxygen concentrations are the dominant factors modulating particle flux attenuation (Cram et al., 2018; Timothy DeVries & Weber, 2017). These models, however, do not account for active transport or disaggregation by zooplankton. As a result of this assumption, the models predict that small particles preferentially attenuate with depth, which is often not borne out by observations (Durkin et al., 2015). Therefore, these models’ predictions provide a useful null hypothesis of expected particle size distributions in the absence of zooplankton effects, which can be compared to observed distributions of particles to explore the magnitude of zooplankton effects.

Underwater vision profilers - cameras that can count and size many particles over large water volumes (Marc Picheral et al., 2010) - provide valuable information about particle distributions and transport. When deployed in concert with particle traps in some regions, they can be used to predict flux in other regions where traps have not been deployed (Guidi et al., 2008; Kiko et al., 2020). UVPs can furthermore provide resolved information about particle flux variability across space and time (Guidi et al., 2008; Kiko et al., 2017). Connecting UVP and trap data can furthermore inform about relationships between particle size, biomass, composition, and sinking speed, as well as the contributions of the different particle sizes to flux (Guidi et al., 2008).

A recent study combined new particle size tracking, mockness tows, and acoustic data, all collected at one site, with previously collected trap measurements from nearby locations to explore zooplankton transport in the Eastern Tropical North Atlantic, a hypoxic, but not fully anoxic, Oxygen Minimum Zone (Kiko et al., 2020)(Kiko et al. 2020). The authors found a particle concentration maximum in the mesopelagic and contended that this feature suggests transport by zooplankton, and/or mortality of migrating zooplankton. The authors suggest that in more anoxic and larger ODZs, such as the modern day ETNP, and in particular as hypoxic water shifts to anoxia there might be reduced active transport into the mesopelagic, since many migratory organisms would presumably not migrate into the anoxic water and would be less active. In this manuscript we provide data from such a fully anoxic region.

A recent modeling study posed three hypotheses to explain why particle flux attenuates slowly in ODZs (Weber & Bianchi, 2020)(Weber and Bianchi 2020), which are susceptible to testing with UVP data. These are: **HA1:** *All* particles in ODZs remineralize more slowly than in oxic water, regardless of their size, due to slower carbon oxidation during denitrification than aerobic respiration. **HA2:** Disaggregation by zooplankton is slower in ODZs than elsewhere. **HA3:** Large particles remineralize more slowly in ODZs, but smaller ones do not, because carbon oxidation in large particles can become limited by the diffusive supply of oxygen and nitrate. In this case, respiration can only proceed by thermodynamically inefficient sulfate reduction (Bianchi et al., 2018; P. Lam & Kuypers, 2011). Sulfide and organic matter sulfurization have been found on particles at this site (Raven et al., 2021). Microbial analysis of particles found sulfate reducers and S-oxidizing denitrifiers at low abundance (Fuchsman et al., 2017; Saunders et al., 2019). Each of the hypotheses outlined above were predicted to leave distinct signatures in particle size distributions in the core of ODZ regions (Weber & Bianchi, 2020). The model with slow attenuation of all particles, predicts an increase in the abundance of small particles in the ODZ core, while the other two models, predict a decrease in small particle abundance, because small particles are either not replaced by disaggregation of large particles (Model 2) or because those particles are remineralized more quickly than larger particles (Model 3). However, the necessary particle size data from an ODZ was not available to support any hypothesis at the exclusion of the others. In this manuscript we present a new dataset that is sufficient to test these three hypotheses (hereafter called Weber-Bianchi models).

While UVP and traps have been sampled together (Guidi et al., 2008), combined trap and UVP measurements have not been taken together previously in an ODZ. Most of the volume of the ETNP ODZ is below regions of very low surface productivity (Fuchsman et al., 2019; Pennington et al., 2006). Meanwhile most flux data has been measured in higher productivity regions of the ETNP (Hartnett & Devol, 2003; Van Mooy et al., 2002). Furthermore, the degree to which zooplankton swimming or other processes lead to particle disaggregation, both in ODZs and elsewhere in the ocean, is unknown.

To provide the data to test hypotheses and illuminate zooplankton particle interactions in oligotrophic ODZs, we collected particle size data at high temporal resolution over the course of a week in an anoxic site typical of the oligotrophic ETNP ODZ, well away from the high productivity zone in the coast. We integrated this size data with observed flux measurements, and acoustic data. We quantified, throughout the water column, how changes in size distribution deviate from changes that would be predicted by remineralization and sinking only models.

We ask the following three questions:

**Question A:** Do our data support any of the three Weber and Bianchi (2020) models (**HA1-HA3**)?

**Question B:** How do the particle size distribution at one location in the oligotrophic Eastern Tropical North Pacific vary with respect to depth and time?

Within the scope of this question, we have three hypotheses: **HB1:** Temporal day to day variability in particle number, particle size distribution slope and flux would be evident. **HB2:** This variability would relate to the location of migratory zooplankton, with a combination of increased particle flux and disaggregation present where zooplankton occur. **HB3:** Disaggregation and particle production by zooplankton might lead to particle size patterns that cannot be explained by remineralization and sinking alone.

**Question C:** Do our data suggest regions of the oxygen deficient zone with disaggregation like processes, and if so, do these co-occur with migratory zooplankton?

# Methods

Unless specified otherwise, measurements were taken on board the R/V Sikuliaq from 07 January 2017 through 13 January 2017 at 16.5°N 106.9°W, located in an oligotrophic region of the Eastern Tropical North Pacific Oxygen Deficient Zone (ETNP Station P2; Figure 1A). Data are compared against measurements taken at 16.5°N 152.0°W on 08 May 2015, collected on the GO-SHIP CLIVAR/CARBON P16N Leg 1 Cruise (CCHDO Hydrographic Cruise 33RO20150410). This station was at the same latitude as ETNP Station P2, west of the ODZ, which was not anoxic (P16 Transect Station 100; Figure S1).

## Water property measurements

We measured water properties of temperature, salinity, fluorescence, oxygen concentration and turbidity using the shipboard SeaBird 911 CTD. Auxiliary sensors included a WetLabs C-Star (beam attenuation and transmission) and a Seapoint fluorometer. Data were processed with Seabird Software, (programs–data conversion, align, thermal mass, derive, bin average and bottle summary) using factory supplied calibrations. Data was analyzed and visualized in *R* (Team 2011).

## Water mass analysis

Evans et al. (2020) previously employed optimum multiparameter analysis to map the percent identity of the water observed at each depth to three water masses: the 13 Degree Celsius Water 13CW, North Equatorial Pacific Intermediate Water (NEPIW), and Antarctic Intermediate Water (AAIW). We subset and examine only the portion of these data that correspond to our site.

## Particle size measurements

Particle size data were collected by Underwater Vision Profiler 5 (UVP) that was mounted below the CTD-rosette and deployed for all CTD casts shallower than 2500 m. A UVP is a combination camera and light source that quantifies the abundance and size of particles from to several centimeters in size (Marc Picheral et al., 2010). Visual inspection of images larger than 1 mm suggests that particles are primarily “marine snow” but about 5% are zooplankton. UVP data were processed using custom MATLAB scripts, uploaded to EcoTaxa (Marc Picheral et al., 2017), and analyzed in R. The UVP provided estimates of particle abundances of particles in different size-bins, as well as information about the volumes over which those particle numbers had been collected.

## Flux measurements

Free floating, surface tethered particle traps were used to quantify carbon fluxes from sinking particles.. Arrays, consisting of a surface float, and two traps were deployed, and allowed to float freely for at least 21-96 hours, during which time they collected and incubated particles. Trap deployments began on 07 January, concurrently with the beginning of the UVP sampling, and continued through January 12th. Trap collection began on the 08 January and continued through 13 January. Trap depths spanned the photic zone and mesopelagic, with the shallowest 60 m and the deepest at 965 m. Traps deployments lasted for between 21 hours, and 93 hours, with deeper traps left out for longer, to collect more biomass. Two types of traps were deployed. One set of traps, generally deployed in shallower water, had a solid cone opening with area 0.46 m2. The second set had larger conical net with opening of 1.24 m2 area made of 53 μm nylon mesh similar to the description in Peterson et al. (2005). In all cases, particles collected in the net or cone fell into one of two chambers. The bottom chamber collected particles from the net and incubated them -- we did not use these samples in this study. The top chamber collected particles for 22-27 hours, and then returned immediately to the surface. For some traps, there was a malfunction such that the door between the bottom and top chamber did not close, in which case we sampled from the combined volume. This meant that these traps had a longer effective collection time of 33-35 hours (69m, 365m 452 m). The trap at 965 m was deployed for an extended period, and also suffered from the trap door malfunction, such that collection time was effectively 91 hours. No poisons were used, and living zooplankton, or ‘swimmers’, were manually removed from collection bottles, while zooplankton carcasses were retained.

Sediment trap material was filtered immediately upon trap recovery onto pre-combusted GF-75 45mm filters and preserved until further analysis at -80 °C. These filters were split into several fractions for other analyses not discussed here. Carbon content of particles in each trap was measured by isotope ratio mass spectrometry. Elemental analyses for particulate carbon and nitrogen quantities as well as 13C and 15N isotopic compositions were conducted at the U.C. Davis Stable Isotope Facility (http://stableisotopefacility.ucdavis.edu) on acidified freeze-dried trap samples to capture organic elemental contributions. Mass spectrometry failed to detect carbon above detection limits in four surface traps, which were excluded from the analysis. Traps at similar depths did detect carbon, lending confidence to the idea that these non-detections were technical in nature, rather than reflecting environmental conditions.

## Analysis

All analyses focused on the mesopelagic, defined here as the region between the base of the secondary chlorophyll maximum layer (175 m) (hereafter photic zone), which is below the upper oxycline, and 1000 m. For many analyses, including visualizations that include the photic region, particles were binned by depth with 20 m resolution between the surface and 100 m, 25 m resolution between 100m and 200 m depths and 50m resolution below 200m. This increasing coarseness of the depth bins helped account for more scarce particles deeper in the water column, while maintaining higher depth resolution near the surface. To perform this binning, particle numbers, and volumes of water sampled of all observations within each depth bin were summed prior to other analyses.

Two normalized values of particle numbers were calculated. In the first, particle numbers were divided by volume sampled, to generate values in *particles/m3*. In the second, particles were divided by both volume sampled and the width of the particle size-bins to generate values in *particles/m3/mm*.

### Particle size distribution

We determined the slope and intercept of the particle size distribution spectrum by fitting a power law to the data, which is a common function for fitting particle size distributions (Buonassissi & Dierssen, 2010). Because large particles were infrequently detected, we used a general linear model that assumed residuals of the data followed a negative-binomial (rather than normal) distribution. We fit the equation

to solve for the Intercept () and particle size distribution slope (). The term on the left describes the expected number of particles seen in a given depth and particle size bin `E(Total Particles)` assuming a negative binomial distribution of residuals. (Date, 2020; Ooi, 2013). *Volume* indicates the volume of water sampled by the UVP, or in the case of depth-binned data, the sum of the volumes of all UVP images in that depth interval. *Binsize* indicates the width of the particle-size bin captured by the UVP. Thus, if particles between .1 and 0.12 mm are in a bin, the *Binsize* is 0.2. On the right hand size, *Size* corresponds to the lower bound of the particle size bin. We use the lower bound of a particle size-in, rather than its midpoint, because, due to the power-law particle size distribution slopes, the average size of particles in each size-bin is closer to the size-bin’s lower bound rather than its midpoint.

### Estimating particle flux

We estimated particle flux throughout the water column, by fitting particle data to trap measurements. We assumed that particle flux in each size bin (j) followed the equation

(Eqn. 1)

Such that flux at a given depth is the sum of all bin specific values.

We used the optimize() function in *R*’s stats package to find the values of and that yielded closest fits of the UVP estimated flux to each particle trap.

We also estimated the exponent of the particle size to biomass exponent and size to sinking speed exponent per the equations and . This is done by assuming a spherical drag profile, in which case and (Guidi et al., 2008).

### Size specific information

We separately analyzed total particle numbers, particle size distribution, and particle flux for particles larger than or equal to 500 μm, and those smaller than 500 μm, to determine the relative contributions of these two particle classes to particle properties. 500 μm was chosen as it has been previously defined as the cutoff point between microscopic “microaggregates” and macroscopic “marine snow” (Simon et al., 2002).

## Variability

To explore the timescales of temporal variability in the POC flux, we determined how well the flux at each depth horizon can be described by the sum of daily and hourly temporal modes. This was achieved by fitting the general additive model of form

(Eqn. 2)

This model explored whether estimated flux levels appeared to vary by day and hour, holding the effects of depth constant, in the 250 m to 500 m region. The smooth terms *s* for “Depth” and “Day” were thin plate splines, while the *s* term for “Hour” was a cyclic spline of 24 hour period.

## Smoothing for Comparison to Model Results

Normalized particle abundance data, from the one cast that traversed the top 2000m of the water column, taken on January 13 at 06:13, was smoothed with respect to depth, time, and particle size using a general additive model of form

(Eqn. 3)

In this case, there is a single, two dimensional, smooth term, rather than additive one dimensional terms as in Eqn. 2 so that the smooth term can consider interactions between the two parameters, rather than assuming that the terms are additive. The predicted particle numbers at each particle size and depth, as well as particle size distribution spectra, and estimated particle masses of all particles smaller than 500 um and all particles larger than or equal to 500 um were then compared to each of Weber and Bianchi’s (2020) models, corresponding to our **HA1-HA3**.

## Modeling remineralization and sinking

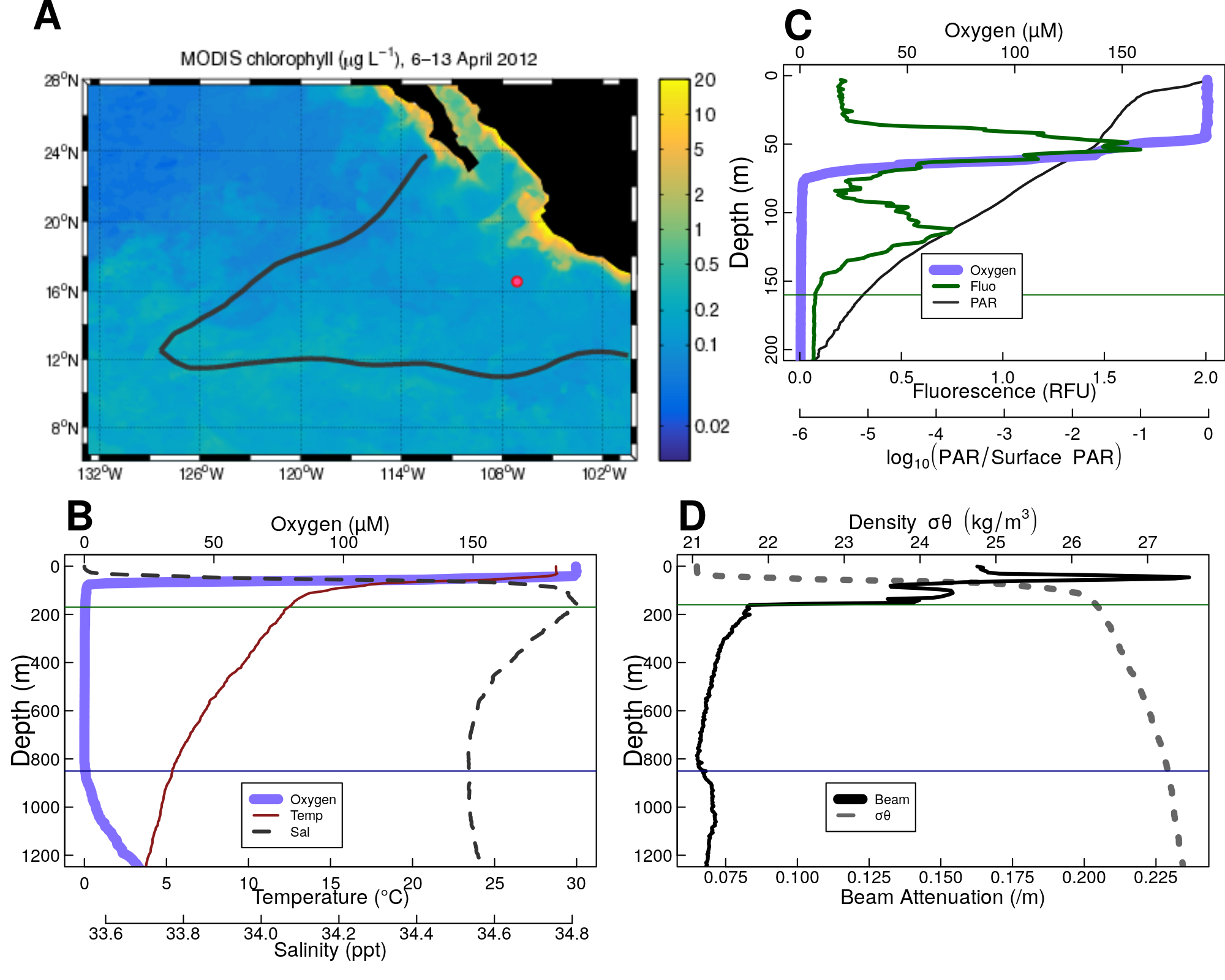
To quantify disaggregation, our goal was to compare the particle size-abundance spectrum at each depth to a prediction of the null hypothesis, that it is simply governed by the effects of sinking and remineralization reshaping the spectrum observed shallower in the water column. This prediction is generated using the particle remineralization and sinking model (PRiSM), modified from DeVries et al. (T. DeVries et al., 2014), applied to the shallower spectrum as an initial condition. The difference between the null hypotheses prediction and observation indicates of the role of processes not accounted for in PRiSM, such as disaggregation, aggregation, and active or advective transport of particles with a different size spectrum than the ones seen at the deeper depth.

In practice we expanded the previous numerical implementation of PRiSM to allow for particle size distribution spectra with particle-size bins that match those obtained by the UVP, and to return estimates of the number of particles in those same size bins (Text S1). The model was given at inputs with particle size distributions at each depth, and calculated flux losses between that depth and the one deeper in the water column. The model uses optimization to determine the particle remineralization rate that would allow the PRiSM model to generate a new size spectrum that has the

# Results

## Physical and Chemical Data

The anoxic zone, characterized by undetectable oxygen levels less than 1 μM, as measured by the CTD, extends from 90 m to 900 m depth, with a sharp upper oxycline and a gradual lower oxycline (Figure 1B-C). The upper oxycline tracks a sharp pycnocline (Figure 1C-1D), set by the high salinity of the 13CW water mass (Figure S2), characterized by an abrupt drop in temperature below the mixed layer, and an increase in salinity (Figure 1B). The site is characterized by two fluorescence maxima (Figure 1C). The larger, shallower fluorescence peak is positioned just above the oxycline, with fluorescence from this peak and oxygen attenuating together. The smaller, lower peak is positioned entirely inside of the anoxic zone. This fluorescence peak does not attenuate until photosynthetically active radiation (PAR) is < 10-5 of surface PAR levels. We note that this photic zone depth is deeper than conventional definitions, in which the base of the photic zone corresponds with 10-2 (90 m) or 10-3 (120 m) of surface PAR. Turbidity tracks the two chlorophyll peaks in the surface and has a tertiary maximum at the lower oxycline (Figure 1D). Water mass analysis indicated that water in the top part of the ODZ is dominated by the 13CW water mass, while water below 400m is primarily from the NEPIW.

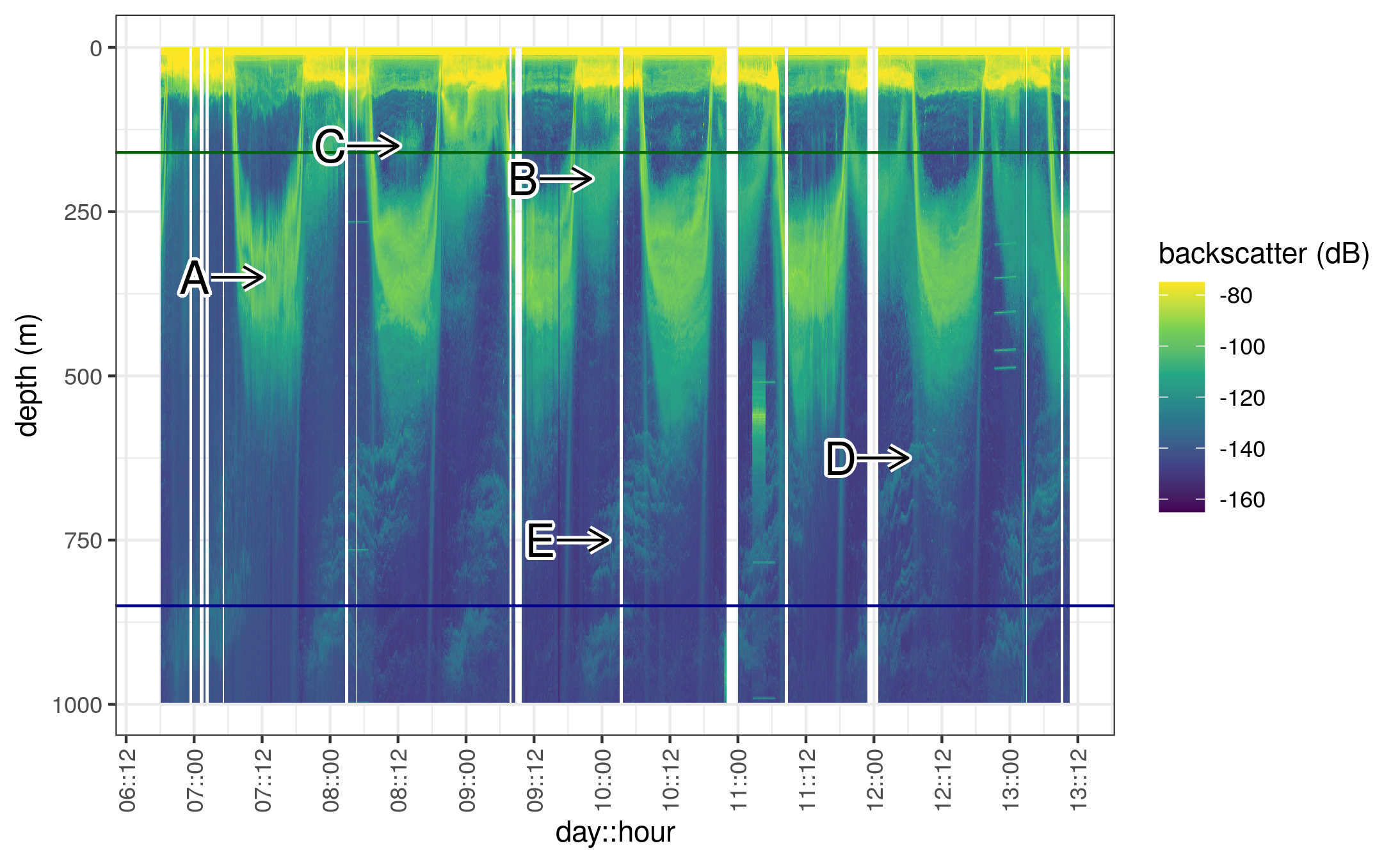


**Figure 1.** Overview of the geography, physics and chemistry of ETNP station P2 (A) Map of the ETNP Oxygen Minimum Zone and the location of station P2. Colors indicate chlorophyll concentrations at the surface as determined by MODIS satellite in 2012, while the black outline signifies the region containing <10 μM oxygen at 300m. The red circle indicates the location of Station P2 (modified from Fuchsman et al 2019, credit Hilary Palevsky). (B-D) Oceanographic parameters collected from a cast at 2017-01-13 12:15 CST (local time). The thin horizontal green line shows the location of the base of the photic zone (160m), defined by the complete attenuation of the in-situ fluorescence, while the horizontal blue line shows the base of the ODZ. (B) highlights temperature and salinity. (C) fluorescence and photosynthetically available radiation (PAR), focusing on the top 200m of the water column and photosynthetically active radiation, and (D) beam attenuation and density.

## Acoustic data reveal diel migration patterns

Acoustic data, produced by the shipboard EK60 (Andersen, 2001), suggest the presence of multiple cohorts of migratory organisms. We focus initially on backscattering measurements from the EK60’s lowest frequency 18000 Hz signal, corresponding to organisms the size of small fish, because it travels furthest into the water column and has the best resolution of the channels. Most migratory organisms appeared to leave the surface at dawn and return at dusk, spending the day between 250 and 500m (Figure 2A). There appeared to be two local maxima in backscattering intensity at mid-day, one at ~300m and one at ~375 m (Figure 2A). There also appeared to be organisms that migrated downward at dusk and upward at dawn, spending the night at ~300m (Figure 2B). There was a peak of organisms that appeared, at mid-day, on some but not all days, without any visible dawn or dusk migration, just above the base of the photic zone. (Figure 2C). Some diel migrators appeared to cross the ODZ and spend the day below the detection range of the EK60 (Figure 2D), as well as organisms that appeared between 500m and 1000m but did not appear to migrate to or from that depth at our site, but rather traveled through the EK60’s field of view (Figure 2D).

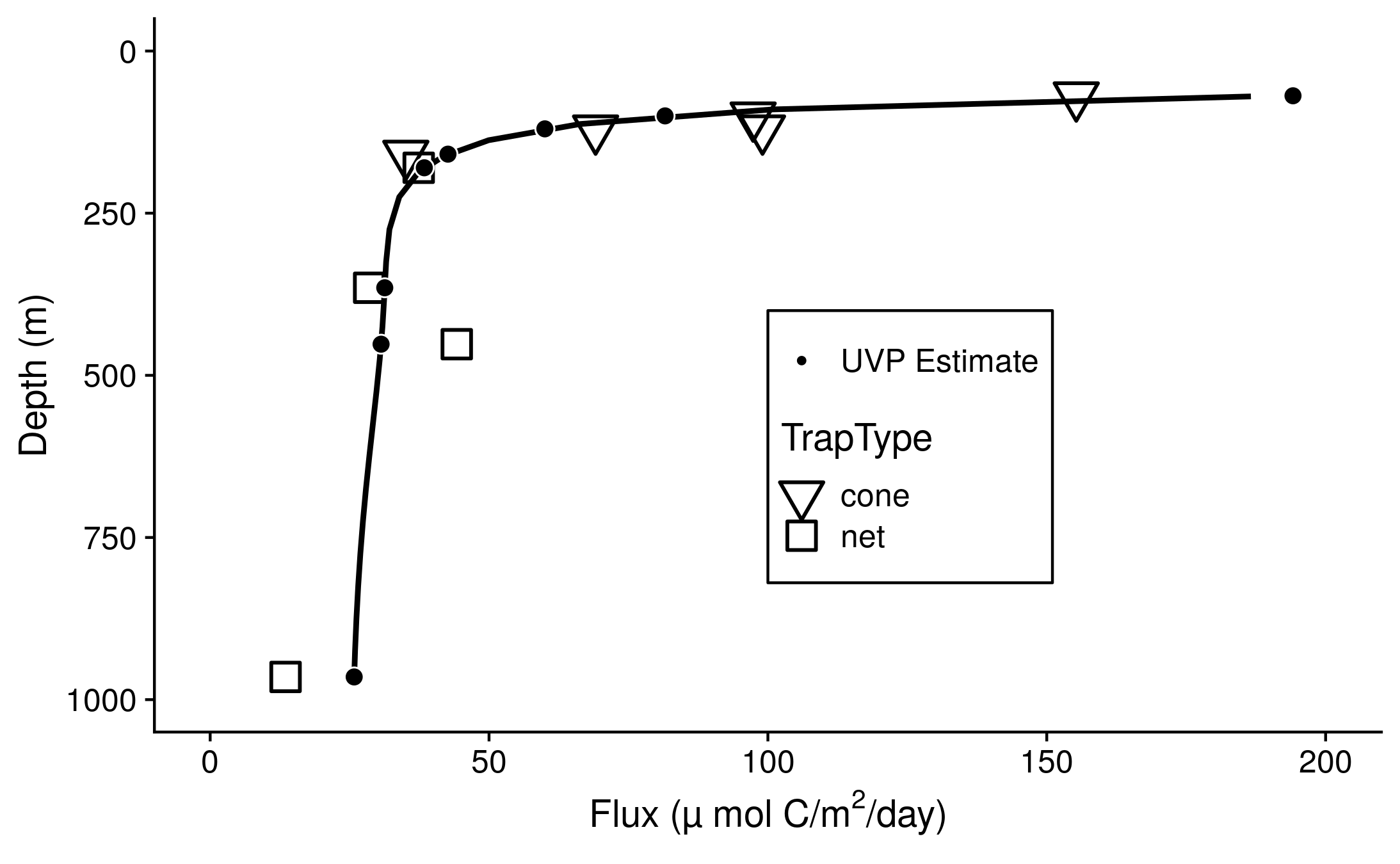
Similar patterns were evident in each of the other measured frequencies, with better resolution by the lower frequencies (Figure S3).



**Figure 2.** Acoustic data, measured by EK60over the course of the experiment. Shown are data from the 18000 Hz frequency band, which have highest depth penetration, but which appear to co-occur with data from other frequency bands (see Figure S3). Values are in return signal intensity and have not been normalized to observed biomass. Several interesting patterns can be seen. **A.** Two bands of organisms can be seen leaving the surface at dawn, spending the day between 250 and 500m and returning to the surface at dusk. **B.** Another group of nocturnally migrating organisms can be seen leaving the surface at dusk, spending the night near 250m and returning at dawn. **C.** Some organisms appear at the base of the photic zone, during some, but not all mid days, and then disappear in the evening. **D.** A group of very deep migrating organisms appears to leave the surface with the diel migrators and pass all the way through the ODZ and out of the EK60’s field of view. It returns at dusk. **E.** Swarms of organisms appear between 500 and 1000m disappearing later in the day. Swarms appear in the deepest layers at night and appear progressively shallower as the day progresses.

## Flux data from traps

Flux measurements at station P2 were consistent between the different particle trap types, showing a profile that broadly followed a power law with respect to depth, with the exception that flux appeared to increase in one trap at 500 m (Figure 3).



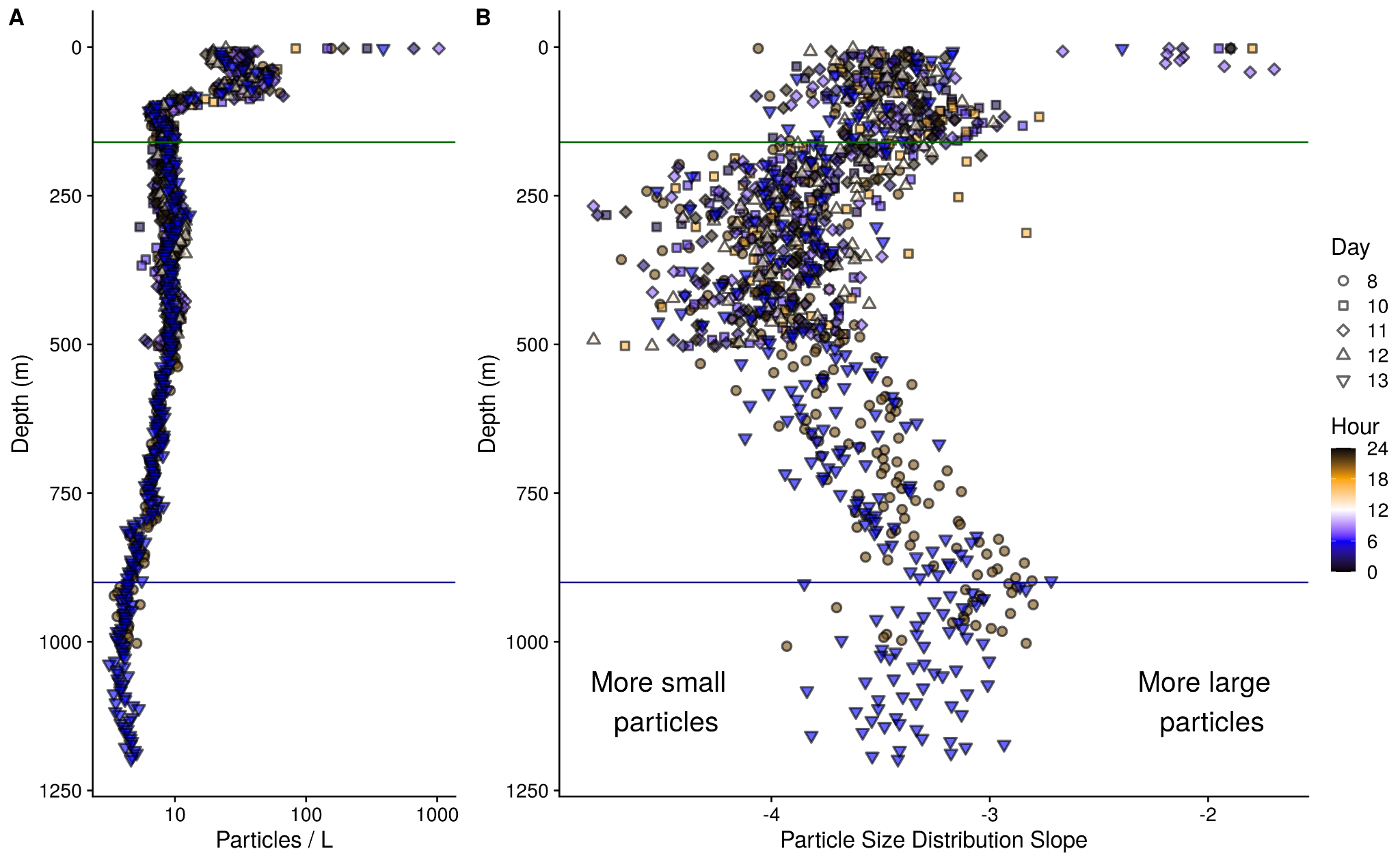
**Figure 3.** Sinking particle flux, measured from surface-tethered sediment traps (large symbols). Trap types are shown by the shape of the large points. Superimposed are binned estimates of particle flux from the UVP generated by fitting the sum of particle numbers all four profiles, binned as in Figure 5, to the trap observed flux. The black line indicates flux predictions made by fitting UVP observations to the trap data. Black circles indicate regions on the black line corresponding to the trap observation depths.

## Particle abundance measurements vary with size and depth

In all profiles, particle abundances were highest at the surface, and highest among the smallest particles (Figure S4). Visual examination of the relationship between particle number and size suggested a power law relationship where the log of volume and bin-sized normalized particle abundance was proportional to the log of the particles’ size (Figure S5). The exception to this pattern were particles larger than 10 mm (Figure S4), which are rare enough that they are usually not detected by the UVP. Generalized linear models that assume a negative-binomial distribution of the data accounted for this under-sampling of large particles to estimate power law slopes, while considering rare occurrences of the large particles at each depth (Figure S5).

Total particle numbers were generally similar between different casts, regardless of which day or hour they were collected (Figure 4A). Particle numbers were highest in the surface and decreased rapidly down to 250m, then remained relatively constant between 250 m to 500 m, before decreasing again until the lower oxycline (Figure 4A).

The particle size distribution slope generally steepened (became more negative) between the surface and 500m, flattened (became less negative) between 500 m and 1000 m, and then steepened again below 1000 m (Figure 4B). Steeper, more negative, slopes indicate a higher proportion of small particles relative to large particles, while flatter, less negative, slopes indicate a higher proportion of large particles.



**Figure 4.** **A.** Observed, volume normalized total particle numbers from 9 casts taken at different times of the day at ETNP station P2. **B.** Calculated particle size distribution slopes of those particles. These data have not been binned by depth. Green line corresponds to the base of the photos zone, while the blue line indicates the base of the lower oxycline.

## Estimated particle flux sometimes increases with depth in the ODZ core

Our optimization approach found best agreement between particle flux measured by traps, and UVP estimated particle flux when per particle flux is given by: Flux = (133 μ mol C / m^2/day) = **133** \* Size (mm) ^ **2.00**. Applying this fit to the UVP data resulted in a UVP predicted flux profile that broadly fit the expected trap observed flux profiles.

Particle flux profiles, predicted from the above fit particle size distribution, varied between casts between the base of the photic zone and 500 m (Figure 5A-5B). To examine the rate of change of flux and to identify regions and time points where flux appeared to increase with depth, we examined the rate of change of flux. This rate of change was fifth root transformed to normalize the data and to allow us to focus on the cases where flux attenuation varied about zero, since we were interested in identifying factors that related to whether flux was positive or negative. Between 250 m and 500 m particle flux appeared to increase on some, but not all, casts, while attenuating slowly on the other casts (Figure 5C). Below 500 m, there were not enough casts to measure variability between casts.

The general additive model that we used to examine the how the of change of flux between 250 m and 500 m varied with depth, study day and hour found that depth (*p* = 0.061) and hour of the day (*p* = 0.196) did not statistically associate with the fifth root transformed rate of change of flux while day of study did (*p* = 0.019) (R2 = 0.264) (Figure S6). In particular, there were generally increases in flux over this region towards the beginning and end of the sampling period and decreases in flux nearer to day 10. A general additive model that looked only at the relationship between study day and rate of change of flux (fifth root transformed) in this region suggested that day accounted for 14% of the variance in this value, as determined by adjusted R­­2 (*p* = 0.040). If the fifth root transformation was not applied to the rate of change of flux, there was a statistically significant relationship between depth and rate of change (p = 0.001), but not study day (*p* = 0.062) or hour (*p* = 0.719) (R2 = 0.341). This pattern indicated that, without the transformation, any temporal signal is swamped by the substantial changes in rate of change in depth, with shallower depths losing flux faster than deeper ones.

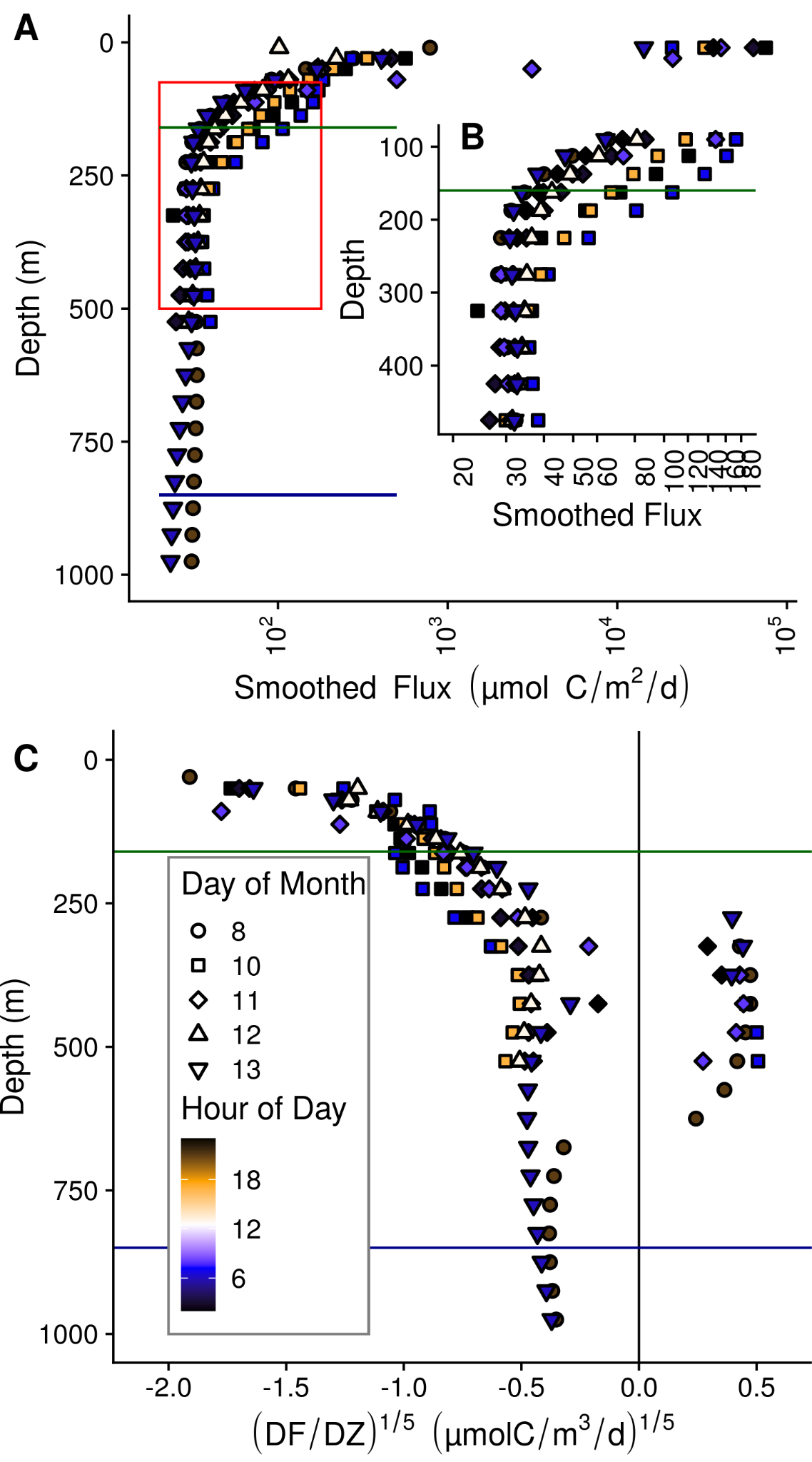


Figure 5. Within and between day variability in UVP predicted particle flux at ETNP station P2. All profiles are depth binned with higher resolution towards the surface (methods). **A.** Flux profiles in the top 1000m of the water column. **B.** A more detailed depiction of the area enclosed by the rectangle in **A**. -- **C.** The rate of change of flux, divided by the rate in change in depth. We show the fifth root of these values in order to highlight differences between values close to zero.

## ETNP particle dynamics differ from those seen at an oxic site

The oxic site, P16 Station 100, was characterized by a more gradually sloping pycnocline, and an oxygen minimum at 500 m of 19.7 μM, which is hypoxic (Figure S1B). There was no working fluorescence sensor on that cruise, but data from World Ocean Atlas (Boyer et al. 2018) suggest that the photic zone is characterized by a single fluorescence peak with a maximum at 110m and which disappeared at 200m (Figure S1C). Turbidity followed chlorophyll concentration and did not have a peak in the mesopelagic (Figure S1D), unlike the ODZ site. There was a salinity peak at 150m (Figure S1B).

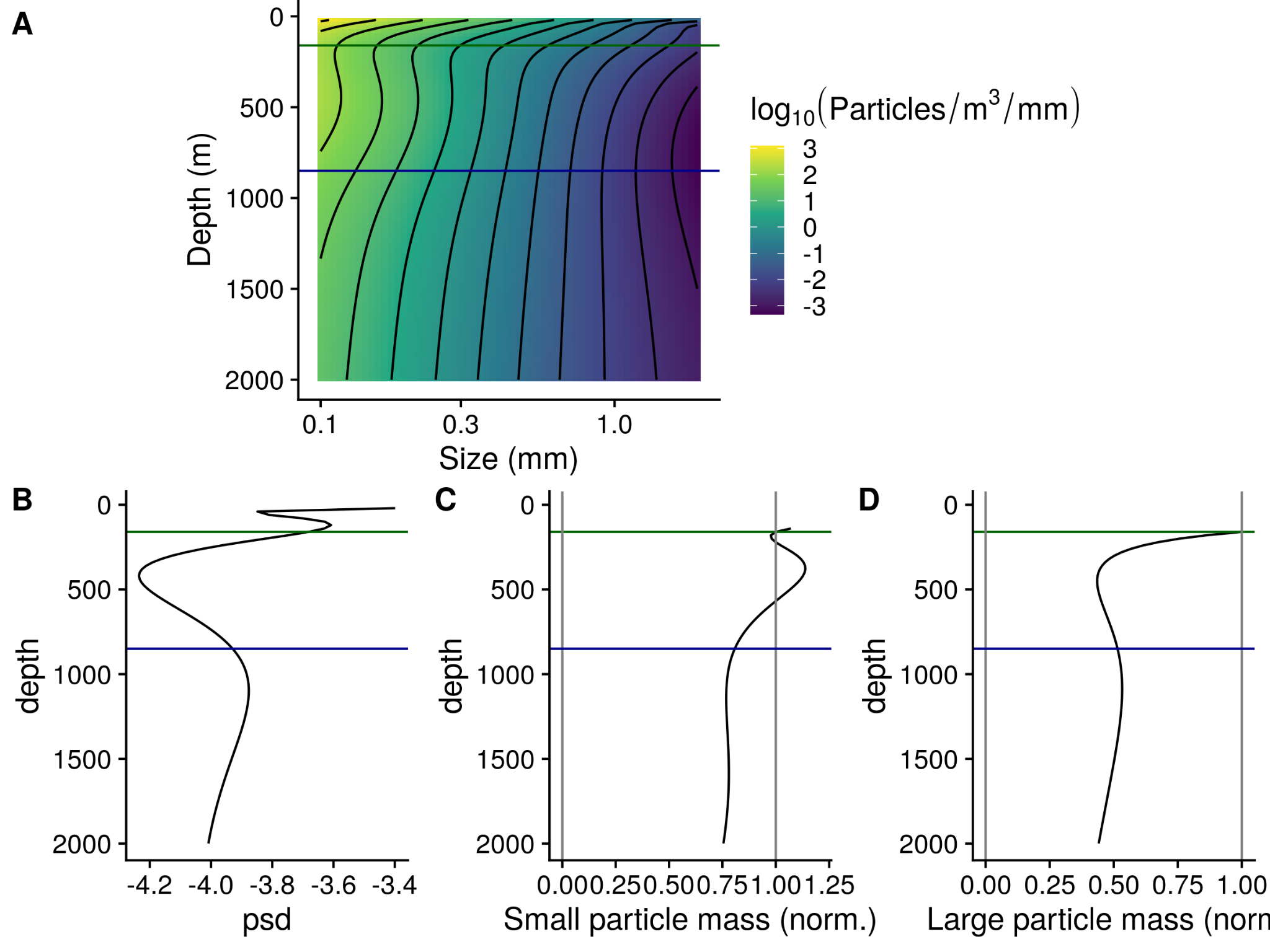
Particle numbers were higher, between the base of the photic zone through 1000 m at the ETNP ODZ site, than at the same-latitude, oxygenic P16 station 100 (Figure S7A). Particle size distributions were similar between the two sites above 500 m, being characterized by overlapping confidence intervals generated by a general additive model. From 500 m to 1000 m, particle size distributions were flatter at the ETNP site, being characterized by a smaller proportion of small particles, relative to large ones (Figure S7B).

Small particles (100 m - 500 m) at the ETNP ODZ site were about two orders of magnitude more common than large particles (>= 500 m) (Figure S8). Large particle numbers appeared to attenuate more quickly than small particles, and more generally follow a power law decrease, while small particles appeared to increase around 500 m. Flux was predicted to be predominantly from small, rather than large particles, at all depths except the very surface. The particle size distribution, calculated only on large particles, was more variable between depths than calculated for small particles. Data from the oxic P16 station 100 suggested more particles, steeper particle size distribution, and more flux at this station than at the ETNP station. They also suggested that differences between large and small particles, with respect to number, flux and size distribution that were broadly similar to the ones seen at ETNP Station P2.

In contrast to the anoxic station, at the oxic station, flux always decreased with depth (Figure S9A+B).

## Smoothed and averaged data

Highly smoothed particle data suggested that particle size, averaged across all casts, followed a pattern in which the abundance of small particles increased between the oxycline and 500 m (Figure 6A), which corresponded with steepening of the particle size distribution (Figure 6A), an increase in small particle biomass (Figure 6B), but not of large particle biomass (Figure 6C). Deeper in the ODZ, the small particle number, particle size distribution slope, and biomass of small particles declined. At the oxic site, particle size distributions generally steepened with depth, while both small and large particle estimated biomass followed a power law decrease with depth (Figure S10).



**Figure 6.** **A.** GAM smoothed bin-size and volume particle numbers at each particle size class, from the one cast that traversed the top 2000m of the water column, collected on January 13 beginning at 06:13. **B.** Particle size distributions. **C-D**. Estimated biomass of **(C)** small and **(D)** large particles.

## Particle number dynamics differ from model expectations

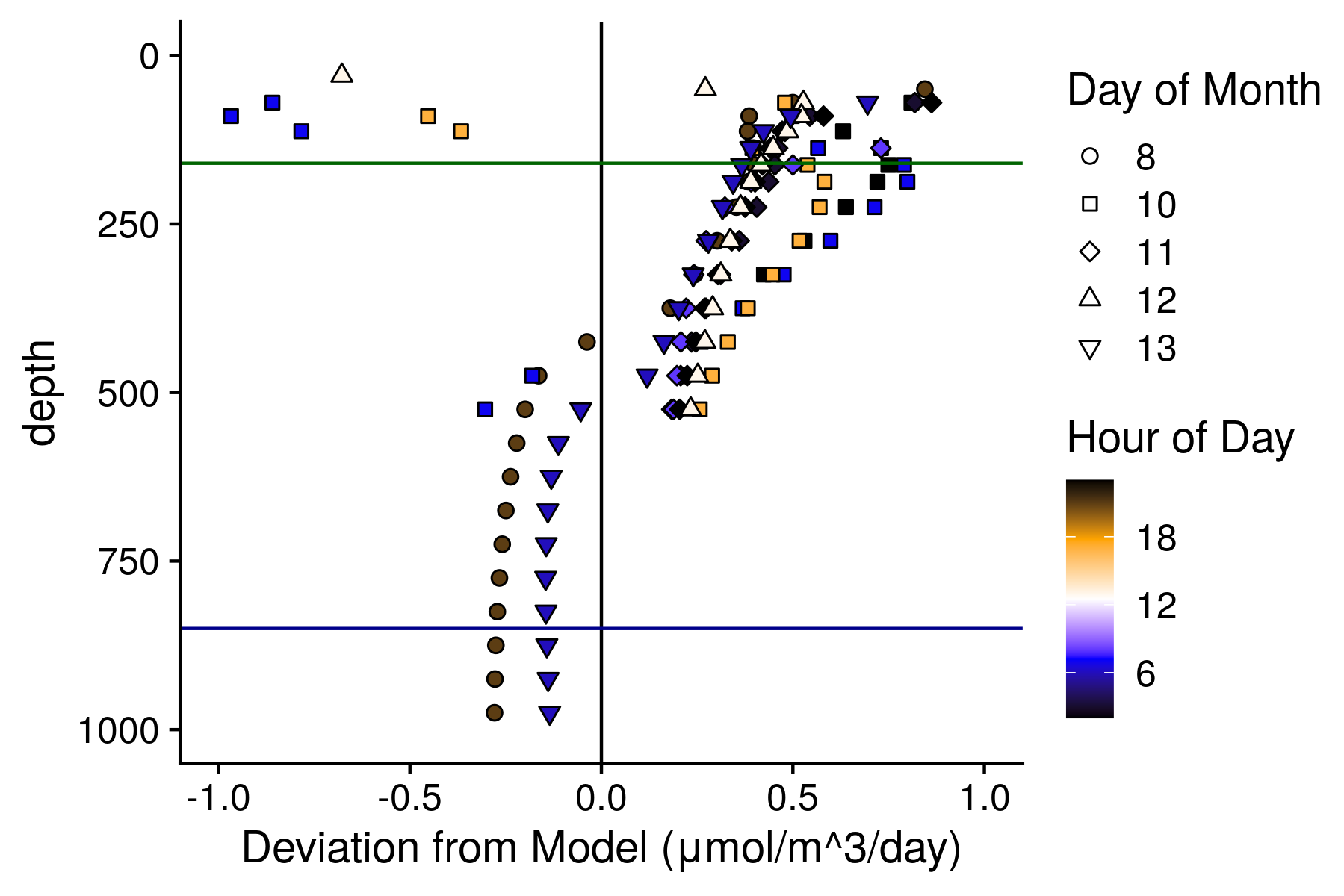
The modified particle remineralization and sinking model predicted particle size distributions at each depth from the particle size distribution one depth-bin shallower and the calculated flux attenuation between the two depths. We found that the observed particle size distributions usually deviated from model expectations (Figure S11). In the model, remineralization rates are optimized s, to ensure that the total predicted flux at each depth matches the observed total flux. However, the predicted size spectrum will diverge from the observed spectrum if the assumptions of the model (i.e., sinking and remineralization are the only particle transformations) are violated. The difference between the observed and predicted flux of *small particles* (), normalized to depth, therefore serves as a metric of observed deviations from the size distribution expected from sinking and remineralization alone. We call this value *Deviation from Model* (DFM).

(Eqn. 4)

In the above equation is the distance, in meters, between the current depth bin and the previous depth bin, whose particle size distribution is fed into the predictive model.

DFM was positive between the photic zone and 500 m, meaning that less small flux attenuated than would be expected from the *PRiSM* model in this region (Figure 7). There was some variability in the DFM parameter between casts. A general additive model (GAM) that showed that the variability in DFM was statistically significantly related to depth (*p* < 10^-5), day of the study (p = 0.002), but not to hour of the day (*p* = 0.051), with these factors together explaining 41.6% of the variance, as measured by R2. DFM was highest shallower in the water column (Figure S12A), highest near day 10 and lower at the beginning and end of the study (Figure S12B). A GAM that only explored the effect of depth accounted for 27% of the variance. Comparing a GAM that accounted for study day and depth to one that only accounted for depth effects showed an increase in R2 value of 11%, suggesting that study day accounts for 11% of the variance, after accounting for depth. Comparing the model that accounts for depth, day and hour to one that only accounts for depth and day, suggests that hour of the day, while not statistically significant, could explains an additional 3% of the variance.

Below 500 m, DFM was negative. There were only two casts that reached below 500m at this station, and so an analysis of the dynamics of DFM in this region is not possible. At P16 Station 100, DFM was positive between the base of the photic zone and 350m and negative below 350 m (Figure S9C).



**Figure 7.** Quantification of non remineralization and sinking like processes. *Deviation from Model* indicates the difference between the observed small particle flux, and the flux that would be estimated if particles from the size distribution in the depth bin above remineralized and sank only following the PRiSM model. Values are normalized to the change in depth and are in units of μ mol Carbon m-3d-1.

# Discussion

## Diel migrators spend time in the ODZ core

Organisms of all sizes appear to migrate into the core of the ODZ at our site. Most migrators appear to leave the surface at dawn, spend the day in the top 500m of the ODZ and return to the surface at dusk (Figure 2A), while others show the opposite pattern, leaving the surface at dusk and returning at dawn (Figure 2B). Diel migration is prevalent throughout the oceans (Cisewski et al., 2010; Hays, 2003; Heywood, 1996; Jiang et al., 2007; Rabindranath et al., 2011; Sainmont et al., 2014; Yang et al., 2019), including at other ODZ sites (Antezana, 2009; Kiko et al., 2020; Riquelme-Bugueño et al., 2020), some of which are highly anoxic sites with secondary, anoxic deep chlorophyll maxima like this one (Hidalgo et al., 2005), and much of the ETNP ODZ (Herrera et al., 2019). Sampling efforts elsewhere in the ETNP suggest that many of these diel migrators are euphausiids and fish (Maas et al., 2014; Wishner et al., 2013), and that diel migrators are primarily 2-5 mm in size (Wishner et al., 2013). Krill in the Humboldt current ODZ migrate to the surface at night (Riquelme-Bugueño et al., 2020), as seen for some organisms at our site (Figure 1B). The presence of organisms that appear and disappear just above the base of the photic zone, in the region of the deeper anoxic fluorescence peak region, but absence of a tell-tale signature of mass migration before or after they appear (Figure 1C) may suggest that these organisms migrate at different times of the day to this deep region, rather than all at once. Another possibility is that they pass through our station at this depth in mid-day, but do not migrate to depth at this location, but rather at another location.

The organisms that appear between 500 m and 1000 m (Figure 2E) have acoustic signatures that resemble those of jellyfish (Kaartvedt et al., 2007). That they appear in horizontal bands that do not appear to trend upwards over time suggests that they are traveling through our site at progressively shallower depths over the course of the day, but that individual swarms are not themselves moving upward at this station. This suggests that any vertical migration carried out by these organisms happens elsewhere or occurs more slowly than the advection seen at this site. That they appear at different depths at different times of the day suggest that these organisms have some sort of vertical migration pattern. Future work may consider more highly resolved spatial and temporal monitoring of this phenomenon. Indeed molecular surveys have found evidence of both Cnidarians and Ctenophores both within and below the ETSP ODZ near Chile (Parris et al., 2014).

## Flux is lower at this site than previous measurements in the ETNP

Flux at our site was lower at all depths, ranging from 10-100 10-100 uM/m^2/day than seen in previous measurements by traps at other, more productive, ODZ sites, ranging from 1-10 mM/m^2/day (Hartnett & Devol, 2003; Van Mooy et al., 2002).

## The flux to size relationship is typical of other sites

The exponent of the particle size to flux relationship that we saw at our site (2.00) is of a similar magnitude to, but slightly smaller than, those seen by other studies that compare UVP flux to trap flux (Guidi et al., 2008; Kiko et al., 2020). Differences in the size-flux relationship could indicate that this relationship truly varies between sites, or that imprecision in flux measurements leads to differences in these values between studies. The single fit and relationship that we carried out does not account for variation in the size to flux ratio across time and depth, nor does it account for differences in particles of different origin. In practice, this value could change over depth and time. More highly depth and space and time resolved flux and particle size measurements might allow one to explore the spatiotemporal variability of this relationship. Image analysis and gel traps (McDonnell & Buesseler, 2010, 2012) provide the opportunity to explore particle size to flux relationships and how they vary between particle types in more depth.

## Remineralization rates of all particles decrease in the ODZ, but disaggregation does not

Particle size profiles, particle size distribution slopes, and estimated biovolume, averaged across all casts and smoothed, are all similar to the predictions made by Weber and Bianchi’s (2020)(2020) “Model 1”. (Figure 5), and therefore our hypothesis **HA1**. This suggests that the low oxygen at this site decreases the particle remineralization rate of all particles, including small ones. It does not support the Weber-Bianchi Model 2 (**HA2**) in which disaggregation is suppressed in the ODZ, nor their Model 3 (**HA3**) in which only the very large particles’ remineralization is slowed. The data at the oxic site, generally conformed to Weber and Bianchi’s “Model 0” (2020), which was their prediction for particle distributions at oxic sites. However, one difference was that the observed particle size distribution, while essentially constant from the base of the photic zone through 1000 m, appeared to steepen between 1000 m and 2000 m, suggesting an increase in the abundance of small particles, relative to Model 0. This could indicate increased disaggregation in this region or horizontal transport of small particles through advection in this region. One possible source of horizontal transport is colloids in a deep iron plume (Homoky et al., 2021; P. J. Lam et al., 2020).

## Zooplankton likely transport organic matter into the ODZ core

Predicted flux levels sometimes increase between 275 m and 625 m, and at all other times attenuate very slowly in this region. The EK60 data suggest the diel migration of all sizes of organisms to this region, agreeing with previous analysis of copepods collected with nets (Wishner et al., 2020). Taken together, the concurrent increase in flux with diel migration suggests that zooplankton are transporting organic matter. That the rate of change in flux with depth varies between days suggests some day to day variability in this transport. That does not vary statistically significantly between day and night suggests that any diel release of particles is relatively small compared to the particles already present in situ. Indeed, it suggests that particle sinking is slow enough that any particles that are transported to depth during the day are retained at night. Furthermore, nocturnal migrators are likely playing a role in carbon transport which may smooth out any diel signal. Another possibility given that the magnitude of the day-to-day variability in apparent particle flux is small is that the zooplankton themselves, which likely make up about 5% of what the UVP counts as particles, may be driving this apparent pattern and that particle flux itself does actually not vary. Zooplankton are known to also congregate at the lower boundaries of ODZs (Wishner et al., 2020) and high urea concentrations in the lower oxycline of the ETNP has been suggested to be due to these zooplankton (Widner et al., 2018). Beam attenuation indicates a third peak in the oxycline below the ODZ. We do not see this congregation in the EK60 data. However, it is possible that small organisms do congregate here, but are not detected by the EK60’s 12000 and 20000 kHZ signals, which do not penetrate to 1000 in our data. The EK60 data do however suggest that larger, krill to fish sized organisms are not abundant in the lower oxycline. Alternatively, the beam attenuation signal could have a non-zooplankton source, such as in-situ formation or horizontal advection of small particles.

## Zooplankton likely disaggregate particles in the ODZ core

The observation that there is greater flux by small particles (< 500 μm) than would be predicted by remineralization and sinking alone (Figure 7), between the photic zone and 500 m suggests that some process is disaggregating large particles into smaller ones. That this apparent disaggregation corresponds with the region where migratory organisms are found suggests that some of these organisms, likely small animals such as copepods and euphausiids (Herrera et al., 2019; Maas et al., 2014), may break down particles (Dilling & Alldredge, 2000; Goldthwait et al., 2005). While other processes such as horizontal advection of water containing small particles (Inthorn, 2005) could be responsible for this increase in small particles in principle, there is no reason to expect horizontal differences at this site, which is at the core of the ODZ and far from shore.

Other deviations from model assumptions could also explain the increase in small particles relative to model predictions. For instance, if the model’s assumed relationships between size, flux, sinking speed and biomass are not all accurate, particle dynamics might differ from our predictions. In particular, small particles might break down more slowly than large ones, or sink more quickly for their size than expected, as has been seen elsewhere (McDonnell & Buesseler, 2010). Our model assumes a spherical particle drag profile, such that the particle sinking speed fractal dimension () is one less than the particle size fractal dimension () (Cram et al., 2018; Guidi et al., 2008), and that these two values sum to the particle flux fractal dimension. If any of these assumptions do not hold, the magnitude of the values may differ.

In contrast to the upper ODZ core, there is an apparent flattening of the particle size distribution below 500 m, beyond the expected effects generated by particle remineralization. This could suggest aggregation processes (Burd & Jackson, 2009). Indeed, aggregation could be occurring throughout the ODZ core, but only exceed disaggregation in the lower ODZ region. Alternatively, in this region, processes resembling Weber and Bianchi’s (2020). Model 3, corresponding to H3B, in which large particles remineralize slower than larger ones could also be occurring. Like aggregation, this process could be occurring through the OMZ but is overwhelmed by the effects of disaggregation above 500 m.

## Water mass changes may affect particle flux and size changes

The observation that particle flux begins to attenuate below 500 m more quickly than it does between the base of the photic zone and 500m could be explained in part by a shift in water mass from NEPIW to AAIW that occurs at this depth (Figure S2). The AAIW is suggested to have higher oxygen concentration than overlying waters though much of the ODZ (Evans et al., 2020). However, measurements taken at this station in 2012 observed zero oxygen though 800m with the highly sensitive STOX electrode, suggesting that oxygen, if present, is below 2nM, even in the AAIW water mass (Tiano et al., 2014). It is conceivable that the AAIW water mass harbored trace oxygen during our study, but not in 2012. It is also possible that the AAIW has different particle size and abundance characteristics due to its having advected from different geographic regions than the overlying water. In any case, the NEPIW to AAIW transition coincides with the lower limit of the depth to which vertically migrating zooplankton travel (Figure 2), and so we are not able to deconvolve the effects of water mass changes from that of changes in zooplankton effects on particle characteristics.

The change in water mass between 13CW and NEPIW, around 250 m, in contrast, does not appear to correspond to any apparent changes in particle flux or size, even though the 13CW is more anoxic through much of the ODZ than the NEPIW (Evans et al., 2020; Larsen et al., 2016). Thus, we would argue that any trace oxygen effects are likely to be very subtle, and that active transport differences above and below 500 m likely have a larger effect.

## Revisiting Hypotheses

Our data provide the first combined look at the relationship between particle size, particle flux, and migratory organisms. Our data support only the first of the three Weber-Bianchi hypotheses (Weber and Bianchi 2020), that particles of all sizes remineralize more slowly in ODZs **HA1**. It *did not support* the hypotheses that disaggregation is lower in the ODZ **HA2**, *nor* that the largest particles degraded particularly slowly in ODZs **HA3**.

While we hypothesized temporal variability in particle size, number, and flux (**HB1)**, we only observed statistically significant temporal variability in flux, and this variability was relatively small, compared to the total amount of flux. Flux was, in general, highest in the day-time, suggesting transport by diel migrators, but this variability was not very strong. This suggests that particle sinking speed may be slow enough that the diel migratory patterns of zooplankton occur on a shorter time-scale than that at which particles sink out of different water column layers. Future measurements of particle sinking speeds in this region would help to constrain these time scales. That the decreased flux attenuation happened more in the day, and primarily in the DVM region supported our hypothesis **HB2** that changes in flux corresponded with migrator abundances.

Our data supported our hypothesis that particle size patterns cannot be explained by remineralization and sinking alone (**HB3)**. Indeed, the abundance of small particles above 500m, suggested and important role for disaggregation is shaping the particle spectrum in the upper ODZ.

## Opportunities for future directions

Due to the relatively small effect sizes of day to day and within day variability in particle fux and disaggregation, we would argue that particle size measurements can be generally thought of as representative of similar environments over a time scale of at least a week. Therefore, we advocate for applying these sorts of analysis and measurements to other parts of the ETNP or other ODZ regions, even ones where there are not time-series data. Expanded spatial analysis of particle size spectra in ODZs would allow the community to confirm whether Weber and Bianchi’s (2020) model **(HA1)**, that particles of all sizes break down more slowly in ODZs, applies elsewhere. Similarly, a clear next step is to apply our disaggregation model to other ocean regions, perhaps using particle size data are already collected by other groups (Guidi et al., 2008; Kiko et al., 2017, 2020).

# Conclusions

If OMZs expand in response to the changing climate, larger areas of the ocean are likely to resemble this environment, which is oligotrophic and with an oxygen deficient zone spanning most of the mesopelegic zone. Previous models have suggested that OMZs are site of efficient carbon transfer to the deep ocean, and our data appear to support this contention. Indeed, the mechanism of efficient transfer appears to be slowing of particle remineralization, presumably from decreased microbial metabolisms, with zooplankton playing an important role in both active particle transport and particle disaggregation.

Our data could potentially be used in conjunction with mechanistic models (e.g. Weber & Bianchi, 2020) to constrain the relative carbon oxidation rate by the two metabolic processes, which is currently poorly understood (Bristow, 2018). Furthermore, it appears that diel migratory organisms both disaggregate particles and transport carbon throughout the top 500 m of the water column. Day to day and within day variability in organic matter transport was evident, though overall patterns in particle size, flux and disaggregation appeared to be consistent over the course of the time-series. The change in particle abundance and size between 500m and the bottom of the ODZ has implications for the free-living microbes living in this region. These microbes are likely particularly organic matter starved. Our data highlights the heterogeneous nature of the ETNP ODZ with depth and indicates that more detailed sampling should be performed for rate and microbial measurements to properly extrapolate to the entire ODZ.

# Acknowledgements

The authors thank the captain and crew of the RV Sikuliaq for making field collection possible. The authors also thank Gabrelle Rocap, Curtis Deutch for assistance in the field and valuable insight. Jacquelyn Burchfield provided valuable insights about the mathematical underpinnings of the particle remineralization model.

Funding for this project was provided by NSF Grant Number DEB-1542240, as well as startup funds to JAC and CAF provided by University of Maryland Center for Environmental Science. The McDonnell laboratory acknowledges support from NSF-OCE 1654663.

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