

2D reductions of the rotating Boussinesq equations

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1 Preliminaries

The rotating Boussinesq equations are

$$D_t \mathbf{u} + 2\boldsymbol{\Omega} \times \mathbf{u} - b\hat{\mathbf{z}} + \boldsymbol{\nabla} p = F\mathbf{u}, \quad (1)$$

$$D_t b + wN^2 = Mb, \quad (2)$$

$$\boldsymbol{\nabla} \cdot \mathbf{u} = 0, \quad (3)$$

where $D_t \stackrel{\text{def}}{=} \partial_t + \mathbf{u} \cdot \boldsymbol{\nabla}$ is the material derivative,

$$2\boldsymbol{\Omega} \stackrel{\text{def}}{=} \underbrace{2\Omega \cos \phi}_{\stackrel{\text{def}}{=} \check{f}} \hat{\mathbf{y}} + \underbrace{2\Omega \sin \phi}_{\stackrel{\text{def}}{=} f} \hat{\mathbf{z}}, \quad (4)$$

is the axis around which the Earth rotates, and the F and M are operators that represent dissipative frictional processes and diffusive mixing processes, respectively. If dissipation and diffusion are due to isotropic molecular processes, then $F = \nu \Delta$ and $M = \kappa \Delta$, where $\Delta = \partial_x^2 + \partial_y^2 + \partial_z^2$ is the three-dimensional Laplacian.

2 Boussinesq equations non-linearized around two-dimensional flow in x, z

The Boussinesq equations with $\partial_y = 0$ and non-linearized around the two-dimensional flow

$$\mathbf{U} = U(z, t) \hat{\mathbf{x}} + V(x, z, t) \hat{\mathbf{y}} \quad (5)$$

become

$$D_t u + w(\check{f} + U_z) - fv + p_x = Fu, \quad (6)$$

$$D_t v + u(f + V_x) + wV_z = Fv, \quad (7)$$

$$D_t w - \check{f}u - b + p_z = Fw, \quad (8)$$

$$D_t b + wN^2 = Mb, \quad (9)$$

$$\boldsymbol{\nabla} \cdot \mathbf{u} = 0, \quad (10)$$

where the material derivative is now

$$D_t \stackrel{\text{def}}{=} \partial_t + (u + U) \partial_x + w \partial_z. \quad (11)$$

In addition to the terms describing advection by the ‘mean’ background flow, there are three qualitatively new refraction terms wU_z , wV_z , and uV_x that appear in the momentum equations.

3 Hydrostatic Boussinesq equations linearized around two-dimensional flow in x, y

The traditional hydrostatic Boussinesq equations linearized around a two-dimensional mean flow are formed by using the hydrostatic approximation in the vertical component of (1), assuming that $\check{f} = 0$ and constant $f = f_0$, and using the map $\mathbf{u} \mapsto \mathbf{U}(x, y, t) + \mathbf{u}(x, y, z, t)$, so that (1)–(3) become

$$u_t + \mathbf{U} \cdot \nabla u + \mathbf{u} \cdot \nabla U - f_0 v + p_x = F_u u, \quad (12)$$

$$v_t + \mathbf{U} \cdot \nabla v + \mathbf{u} \cdot \nabla V + f_0 u + p_y = F_u v, \quad (13)$$

$$p_z = b, \quad (14)$$

$$b_t + \mathbf{U} \cdot \nabla b + w N^2 = M b, \quad (15)$$

$$u_x + v_y + w_z = 0. \quad (16)$$

A natural model for the two-dimensional flow

$$\mathbf{U}(x, y, t) = -\psi_y \hat{\mathbf{x}} + \psi_x \hat{\mathbf{y}} \quad (17)$$

is that solves the two-dimensional vorticity equation,

$$\triangle_h \psi_t + \mathbf{J}(\psi, \triangle_h \psi) = F_\psi (\triangle_h \psi), \quad (18)$$

Equations (12)–(16) describe the advection and refraction of waves by a two-dimensional flow with $\mathbf{U}_z = \psi_z = 0$ and thus no buoyancy field. The linearization neglects the complications of nonlinear wave dynamics and permits a two-dimensionalization of (12)–(16) by projection onto vertical modes. (12)–(16) do not have a cascade to small-scales, in general, and therefore permit inviscid numerical solutions.

3.1 The vertical mode decomposition

We restrict attention to waves with simple vertical structure by projecting (12)–(16) onto the hydrostatic vertical modes $\phi_n(z)$ that solve the eigenproblem

$$\frac{f_0^2}{N^2} \phi_{nzz} + \lambda_n^{-2} \phi_n = 0, \quad \text{with} \quad \phi_n = 0 \quad \text{at} \quad z = -H, 0. \quad (19)$$

Note that the derivative h_{nz} satisfies $h_{nz} = -\lambda_n^2 L h_{nz}$. The modal amplitudes of the independent variables A, \mathbf{u}, b, p are defined by their weighted projection onto ϕ_n or its derivative ϕ_{nz} , with

$$\Phi_n \stackrel{\text{def}}{=} \int_{-H}^0 \Phi \phi_{nz} dz \quad \text{for} \quad \Phi = (A, u, v, p), \quad (20)$$

and

$$b_n \stackrel{\text{def}}{=} \int_{-H}^0 b \phi_n dz \quad \text{and} \quad w_n \stackrel{\text{def}}{=} \int_{-H}^0 \frac{N^2}{\lambda_n^2 f_0^2} w \phi_n dz. \quad (21)$$

We assume A, \mathbf{u}, b , and p satisfy free-slip, rigid-lid homogeneous boundary conditions with $A_z = u_z = v_z = p_z = 0$ and $w = b = 0$ at $z = -H, 0$.

The linearized Boussinesq equations (12)–(16) are processed in similar fashion. We project (12) and (13) onto ϕ_{nz} . We assume we can write, approximately

$$F_{nu}(u_n) \approx \int_{-H}^0 \phi_{nz} F_u(u) dz. \quad (22)$$

This approximation does not hold for all linear operators F_u . Notice that if $F = \nu \Delta$, then

$$\int_{-H}^0 \phi_{nz} \nu \Delta u dz = \nu \Delta_h u_n + \frac{\kappa_n^2}{f_0^2} \int_{-H}^0 N^2 \phi_n u_z dz, \quad (23)$$

and different modes are therefore coupled by the rightmost term. When N is constant, however, this becomes

$$\int_{-H}^0 \phi_{nz} \nu \Delta u dz = \nu \Delta_h u_n + \nu \left(\frac{n\pi}{H}\right)^2 u_n, \quad (24)$$

and the modes separate. With the notation in (22), the horizontal momentum equations

$$u_{nt} - f_0 v_n + p_{nx} = -\mathbf{U} \cdot \nabla u_n - \mathbf{u}_n \cdot \nabla U + F_{nu} u_n, \quad (25)$$

$$v_{nt} + f_0 u_n + p_{ny} = -\mathbf{U} \cdot \nabla v_n - \mathbf{u}_n \cdot \nabla V + F_{nv} v_n. \quad (26)$$

We next combine (14)–(16) by projecting (16) onto ϕ_{nz} , integrating by parts once, and using (19) to yield $w_n = -u_{nx} - v_{ny}$. We then use $p_z = b$ to combine (14) and (15) and project the result onto ϕ_n . Similar to (22), we use the notation

$$M_n p_n \approx - \int_{-H}^0 \phi_n M p_z dz. \quad (27)$$

Similar to (22), this is not true when N is not constant. When N is constant we have $M_n p_n = \kappa \Delta_h p_n + \kappa \left(\frac{n\pi}{H}\right)^2 p_n$. Finally, integrating by parts and using $w_n = -u_{nx} - v_{ny}$ transforms (15) into

$$p_{nt} + \left(\frac{f_0}{\kappa_n}\right)^2 (u_{nx} + v_{ny}) = -\mathbf{U} \cdot \nabla p_n + M_n p_n. \quad (28)$$

The three equations (25)–(28) describe the evolution of hydrostatic, vertical mode- n waves in a two-dimensional flow $\mathbf{U} = U \hat{\mathbf{x}} + V \hat{\mathbf{y}}$ with $\mathbf{U}_z = 0$. The parameter f_0/κ_n is the phase speed of a linear wave with mode- n vertical structure.

A ‘Wave operator form’ of the Boussinesq equations

The component-wise rotating inviscid Boussinesq equations are

$$u_t - f v + \check{f} w + p_x = -\mathbf{u} \cdot \nabla u, \quad (29)$$

$$v_t + f u + p_y = -\mathbf{u} \cdot \nabla v, \quad (30)$$

$$w_t - \check{f} u - b + p_z = -\mathbf{u} \cdot \nabla w, \quad (31)$$

$$b_t + w N^2 = -\mathbf{u} \cdot \nabla b, \quad (32)$$

$$\nabla \cdot \mathbf{u} = 0. \quad (33)$$

We first form the ‘oscillation equation’ with the combination $\partial_t(31) + (32)$:

$$(\partial_t^2 + N^2) w - \check{f}u_t + p_{zt} = -\partial_t(\mathbf{u} \cdot \nabla w) - \mathbf{u} \cdot \nabla b. \quad (34)$$

The ‘divergence equation’ follows from $-\partial_x(29) - \partial_y(30)$ and using $u_x + v_y = -w_z$,

$$w_{zt} + f\omega - uf_y - \check{f}w_x - \Delta_h p = \partial_x(\mathbf{u} \cdot \nabla u) + \partial_y(\mathbf{u} \cdot \nabla v). \quad (35)$$

The vertical vorticity equation is obtained from $\partial_x(30) - \partial_y(29)$,

$$\omega_t - fw_z + vf_y = -\nabla_\perp \cdot (\mathbf{u} \cdot \nabla) \mathbf{u}. \quad (36)$$

Yes! In the penultimate step we calculate $\partial_z \partial_t(35) - f \partial_z(36)$, yielding

$$\left[(\partial_t^2 + f^2) \partial_z^2 - \check{f} \partial_x \partial_z \partial_t \right] w - \Delta_h p_{zt} = f_y \partial_z (u_t + fv) + \partial_z (\partial_t \nabla + f \nabla_\perp) \cdot (\mathbf{u} \cdot \nabla) \mathbf{u}_h. \quad (37)$$

Finally, the combination $\Delta_h(34) + (37)$ yields the wave operator form,

$$\begin{aligned} \left[\Delta \partial_t^2 + f^2 \partial_z + N^2 \Delta_h \right] w = & \check{f} \partial_t (\Delta_h u + w_{xz}) + u_t \check{f}_{yy} + 2u_{yt} \check{f}_y + f_y \partial_z (u_t + fv) \\ & + \partial_z (\partial_t \nabla + f \nabla_\perp) \cdot (\mathbf{u} \cdot \nabla) \mathbf{u} - \Delta_h (\mathbf{u} \cdot \nabla b) - \Delta \partial_t (\mathbf{u} \cdot \nabla w). \end{aligned} \quad (38)$$