

Reduced equations for Boussinesq internal waves

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1 The YBJ equation

The two prognostic variables in the YBJ equation are wave field amplitude A and quasi-geostrophic potential vorticity q . Quasi-geostrophic potential vorticity is related to the quasi-geostrophic streamfunction through the elliptic equation

$$q = \left(\underbrace{\partial_x^2 + \partial_y^2}_{\stackrel{\text{def}}{=} \Delta_h} + \underbrace{\partial_z \frac{f_0^2}{N^2} \partial_z}_{\stackrel{\text{def}}{=} L} \right) \psi, \quad (1)$$

where the streamfunction is related to the velocity and buoyancy fields via $\mathbf{U} = -\psi_y \hat{\mathbf{x}} + \psi_x \hat{\mathbf{y}}$ and $B = f_0 \psi_z$. The quasi-geostrophic potential vorticity solves the advection equation

$$q_t + \mathbf{J}(\psi, q) = F_u \Delta_h \psi + M_b L \psi. \quad (2)$$

The wave field amplitude represents the complexified horizontal velocity via

$$\tilde{u} + i\tilde{v} = e^{-if_0 t} L A, \quad (3)$$

and solves the YBJ equation

$$L A_t + \frac{if_0}{2} \Delta_h A + \mathbf{J}(\psi, L A) + L A \left(\frac{i}{2} \Delta_h \psi + \beta y \right) = F_A L A. \quad (4)$$

1.1 In x, y with barotropic $\psi(x, y, t)$

1.2 In x, z with stationary $\psi(x, z)$

A The vertical mode decomposition

The hydrostatic vertical modes $\phi_n(z)$ solve the eigenproblem

$$\frac{f_0^2}{N^2} \phi_{nzz} + \lambda_n^{-2} \phi_n = 0, \quad \text{with} \quad \phi_n = 0 \quad \text{at} \quad z = -H, 0. \quad (5)$$

Note that the derivative h_{nz} satisfies $h_{nz} = -\lambda_n^2 L h_{nz}$. The amplitudes of certain quantities are determined by their weighted projection onto ϕ_n or its derivative ϕ_{nz} . For example,

$$u_n \stackrel{\text{def}}{=} \int_{-H}^0 \Phi \phi_{nz} dz, \quad b_n \stackrel{\text{def}}{=} \int_{-H}^0 b \phi_n dz \quad \text{and} \quad w_n \stackrel{\text{def}}{=} \int_{-H}^0 \frac{N^2}{\lambda_n^2 f_0^2} w \phi_n dz. \quad (6)$$

The proper choice depends on the equation set and chosen boundary conditions.

Some trickiness is associated with the friction operator F . We want to assume, for example, that we can define an operator F_n such that

$$F_{nu}(u_n) \approx \int_{-H}^0 \phi_{nz} F_u(u) dz, \quad (7)$$

for example. This is not always possible, however. Notice that if $F = \nu \Delta$, then

$$\int_{-H}^0 \phi_{nz} \nu \Delta u dz = \nu \Delta_h u_n + \frac{\kappa_n^2}{f_0^2} \int_{-H}^0 N^2 \phi_n u_z dz, \quad (8)$$

and different modes are therefore coupled by the rightmost term. When N is constant, however, this becomes

$$\int_{-H}^0 \phi_{nz} \nu \Delta u dz = \nu \Delta_h u_n + \nu \left(\frac{n\pi}{H} \right)^2 u_n, \quad (9)$$

and therefore $F_n = \nu \left(\Delta_h + \left(\frac{n\pi}{H} \right)^2 \right)$. A similar conundrum is associated with the diffusion operator M .

A The non-hydrostatic Boussinesq equations

The rotating Boussinesq equations are

$$D_t \mathbf{u} + 2\boldsymbol{\Omega} \times \mathbf{u} - b \hat{\mathbf{z}} + \nabla p = F \mathbf{u}, \quad (10)$$

$$D_t b + w N^2 = M b, \quad (11)$$

$$\nabla \cdot \mathbf{u} = 0, \quad (12)$$

where $D_t \stackrel{\text{def}}{=} \partial_t + \mathbf{u} \cdot \nabla$ is the material derivative,

$$2\boldsymbol{\Omega} \stackrel{\text{def}}{=} \underbrace{2\Omega \cos \phi}_{\stackrel{\text{def}}{=} \check{f}} \hat{\mathbf{y}} + \underbrace{2\Omega \sin \phi}_{\stackrel{\text{def}}{=} f} \hat{\mathbf{z}}, \quad (13)$$

is the axis around which the Earth rotates, and the F and M are operators that represent dissipative frictional processes and diffusive mixing processes, respectively. If dissipation and diffusion are due to isotropic molecular processes, then $F = \nu \Delta$ and $M = \kappa \Delta$, where $\Delta = \partial_x^2 + \partial_y^2 + \partial_z^2$ is the three-dimensional Laplacian.

B ‘Wave operator form’ of the Boussinesq equations

The component-wise rotating inviscid Boussinesq equations are

$$u_t - f v + \check{f} w + p_x = -\mathbf{u} \cdot \nabla u, \quad (14)$$

$$v_t + f u + p_y = -\mathbf{u} \cdot \nabla v, \quad (15)$$

$$w_t - \check{f} u - b + p_z = -\mathbf{u} \cdot \nabla w, \quad (16)$$

$$b_t + w N^2 = -\mathbf{u} \cdot \nabla b, \quad (17)$$

$$\nabla \cdot \mathbf{u} = 0. \quad (18)$$

We first form the ‘oscillation equation’ with the combination $\partial_t(16) + (17)$:

$$(\partial_t^2 + N^2) w - \check{f}u_t + p_{zt} = -\partial_t(\mathbf{u} \cdot \nabla w) - \mathbf{u} \cdot \nabla b. \quad (19)$$

The ‘divergence equation’ follows from $-\partial_x(14) - \partial_y(15)$ and using $u_x + v_y = -w_z$,

$$w_{zt} + f\omega - uf_y - \check{f}w_x - \Delta_h p = \partial_x(\mathbf{u} \cdot \nabla u) + \partial_y(\mathbf{u} \cdot \nabla v). \quad (20)$$

The vertical vorticity equation is obtained from $\partial_x(15) - \partial_y(14)$,

$$\omega_t - fw_z + vf_y = -\nabla_\perp \cdot (\mathbf{u} \cdot \nabla) \mathbf{u}. \quad (21)$$

Yes! In the penultimate step we calculate $\partial_z \partial_t(20) - f \partial_z(21)$, yielding

$$\left[(\partial_t^2 + f^2) \partial_z^2 - \check{f} \partial_x \partial_z \partial_t \right] w - \Delta_h p_{zt} = f_y \partial_z (u_t + fv) + \partial_z (\partial_t \nabla + f \nabla_\perp) \cdot (\mathbf{u} \cdot \nabla) \mathbf{u}_h. \quad (22)$$

Finally, the combination $\Delta_h(19) + (22)$ yields the wave operator form,

$$\begin{aligned} \left[\Delta \partial_t^2 + f^2 \partial_z + N^2 \Delta_h \right] w = & \check{f} \partial_t (\Delta_h u + w_{xz}) + u_t \check{f}_{yy} + 2u_{yt} \check{f}_y + f_y \partial_z (u_t + fv) \\ & + \partial_z (\partial_t \nabla + f \nabla_\perp) \cdot (\mathbf{u} \cdot \nabla) \mathbf{u} - \Delta_h (\mathbf{u} \cdot \nabla b) - \Delta \partial_t (\mathbf{u} \cdot \nabla w). \end{aligned} \quad (23)$$