## An Overview of Functional Linear Models

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#### Functional linear models

A linear model can become functional in one or both of two ways:

- The response variable *y* with argument *t* is functional.
- One or more of the covariates x is functional.

#### FLM: Some scenarios

- Chapter 13: functional response with categorical covariates (*functional analysis of variance*) or scalar covariates (*functional multiple regression*).
- Chapter 14: functional response depends on functional covariates concurrently (concurrent model).
- Chapter 15: scalar response and functional covariates.

## The average Canadian weather data

- 35 Canadian weather stations selected to cover the country.
- Daily temperatures (0.1 degrees Celsius) and precipitations (0.1 mm) averaged over the years 1960 to 1994. (Feb 29th combined with Feb. 28th).
- Canada divided into Atlantic, Continental, Pacific and Arctic weather zones.

## 35 Canadian weather stations



#### Variables in Canadian weather data

- Response variable: a functional response Prec(t)
  measuring average precipitation over time t or a
  scalar response Prectot measuring the total yearly
  precipitation averaged over the years.
- Categorical covariate: climate groups referring to the Atlantic, Continental, Pacific and Arctic weather zones.
- Functional covariate: average temperature Temp(t) measured over time t.

## Problem setup

- Scientific question: How does the shape of the mean annual precipitation profile depend on which climate zone the station is in?
- Climate zones indexed by g = 1, 2, 3, 4.
- Weather station in each zone indexed by  $m = 1, ..., N_a$ .

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#### Model

The model for the mth precipitation function in the gth zone is

$$\mathsf{Prec}_{mg}(t) = \mu(t) + \alpha_g(t) + \epsilon_{mg}(t).$$

- $\bullet$   $\mu(t)$ : the grand mean profile across all 35 weather stations
- $\alpha_g(t)$ : the *effect functions* representing departures from the grand mean specific to climate zones.
- $\epsilon_{mg}(t)$ : residual effect at the weather station.
- Zero sum constraint:

$$\sum_g \alpha_g(t) = 0 \text{ for all } t.$$

## Some interesting questions

- Functional contrasts: "there are no differences in mid-summer between the climate zones"
- Multiple comparisons: "Over which time intervals are there significant differences between climate zones?"

## Problem setup

- Scientific question: Does the total amount of precipitation depend on specific features of the temperature profile of a weather station?
- Response variable:

$$Prectot_i = \int_0^{365} Prec_i(t) dt,$$

for i = 1, ..., 35.

#### Model

The model for the *i*th total precipitation is

$$\mathsf{Prectot}_i = lpha + \int_0^{365} \mathsf{Temp}_i(t) eta(t) dt + \epsilon_i.$$

- $\alpha$ : the intercept adjusting for the origin of the precipitation variable.
- $\beta(t)$ : the regression coefficient function and the functional parameter of interest.
- $\epsilon_i$ : residuals.

# Comparison with multiple linear regression

- Multiple linear regression:
  - *n* observations:  $(y_i, \mathbf{x}_i), i = 1, \dots, n$
  - each  $\mathbf{x}_i$  is a *p*-vector:  $\mathbf{x}_i = (x_{1i}, \dots, x_{pi})'$ .
  - the p-vector parameter  $\beta$  of coefficients is estimable if n > p.
- Functional linear regression:
  - n observations:  $(y_i, x_i), i = 1, ..., n$
  - each  $x_i$  is a function with dimension=  $\infty$ .
  - how to make sure the coefficient function  $\beta(t)$  is estimable?

#### A FPCA solution

• Karhunen-Loéve decomposition of the covariance function K(s,t) of x(t):

$$K(s,t) = \sum_{k=1}^{\infty} \lambda_k e_k(s) e_k(t),$$

where  $\{\lambda_k, e_k(t)\}$  are eigenvalues and eigenfunctions of the linear operator  $f(t) \mapsto \int K(s, t) f(s) ds$  associated with K.

• x(t) can be expressed as  $x(t) = \sum_{k=1}^{\infty} \theta_k e_k(t)$ , where  $\theta_k = \int x e_k$ .

#### A FPCA solution: The truncation

- truncate the series at a small K: so x(t) is approximated by  $x(t) \approx \sum_{k=1}^{K} \theta_k e_k(t)$
- expand  $\beta(t)$  using the same truncated basis functions:  $\beta(t) = \sum_{k=1}^{K} \beta_k e_k(t)$ .
- essentially a multiple linear regression problem again.
- Drawback: the basis system  $\{e_k(\cdot): k=1,\ldots,K\}$  may not be an effective one for representing  $\beta(t)$ .

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## A regularization solution

- add a roughness penalty on  $\beta(t)$ ,
- essentially cut down the dimension of  $\beta$  by enforcing certain smoothness constraint on  $\beta$ .

## Problem setup

- Scientific question: How does a precipitation profile depend on the associated temperature profile?
- There are several versions of models that are relevant under different assumptions on the influence range of the temperature profile.

#### The concurrent model

- Assumption: precipitation at time t depends only on the temperature at time t.
- The concurrent model for the ith precipitation function is

$$\operatorname{Prec}_{i}(t) = \alpha(t) + \operatorname{Temp}_{i}(t)\beta(t) + \epsilon_{i}(t).$$

- A special case of the *varying coefficient models* in Hastie and Tibshirani (1993).
- Should we use regularization to force  $\beta$  to be smooth in t?

#### The annual or total model

- Assumption: temperature influence on precipitation Prec(t) at a fixed time point t extends over the whole year.
- The annual model for the *i*th precipitation function is

$$\mathsf{Prec}_i(t) = lpha(t) + \int_0^{365} \mathsf{Temp}_i(s) eta(s,t) ds + \epsilon_i(t).$$

- The value  $\beta(s,t)$  determines the impact of temperature at time s on precipitation at time t.
- We need roughness penalties for variation in both s and t.

#### The short-term feed-forward model

- Assumption: precipitation Prec(t) at a fixed time point t only depends on temperatures over a small time interval back.
- Example: falling temperature before precipitation.
- The corresponding model for the *i*th precipitation function is

$$\mathsf{Prec}_i(t) = lpha(t) + \int_{t-\delta}^t \mathsf{Temp}_i(s) eta(s,t) ds + \epsilon_i(t).$$

- $\delta$ : a pre-specified time lag over which we use temperature information.
- $\beta(s,t)$  is only defined on a complicated trapezoidal domain:  $t \in [0,365], t-\delta \leq s \leq t$ .
- More care taken for non-periodic data on the boundary of the domain.

#### The local influence model

- Assumption: precipitation Prec(t) at a fixed time point t only depends on temperatures over a t-dependent set  $\Omega_t$ .
- Example: thunderstorm in summer comes with falling temperature before it and rising temperature afterwards.
- The corresponding model for the ith precipitation function is

$$\mathsf{Prec}_i(t) = lpha(t) + \int_{\Omega_t} \mathsf{Temp}_i(s) eta(s,t) ds + \epsilon_i(t).$$

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