

Normal modes of the Earth

For long-period waves (> 100 s period), self-gravity effects become important and we can write the linearized equations of motion as

$$\rho_0 \partial_t^2 \underline{u}(r, t) = \nabla \cdot \underline{\tau}'(r, t) - \underbrace{\rho_0 \nabla \phi'(r, t) - \rho' \nabla \phi_0(r, t)}_{\text{gravity effects}}$$

$$\text{with } \begin{cases} \nabla^2 \phi' = 4\pi G \rho' & \text{compressible} \\ \nabla^2 \phi' = 0 & \text{incompressible } (\rho' = 0) \end{cases}$$

where we have introduced the gravitational potential $\phi(r)$ and the linearized forms

$$\rho(r, t) = \rho_0(r) + \rho'(r, t)$$

$$\phi(r, t) = \phi_0(r) + \phi'(r, t)$$

$$\tau(r, t) = \tau_0(r) + \tau'(r, t)$$

with small perturbations in density ρ , potential ϕ , and stress τ

In general, the gravitational potential is governed by Poisson's equation

$$\nabla^2 \phi(r, t) = 4\pi G \rho(r, t)$$

SNREI (spherically symmetric, Non-Rotating, Elastic Isotropic) Earth

In frequency domain:

$$-\rho_0 \omega^2 \underline{u}(r, \omega) = \underline{\nabla} \cdot \underline{\underline{\tau}}'(r, \omega) - \rho_0 \underline{\nabla} \phi'(r, \omega) - \rho' \underline{\nabla} \phi_0(r, \omega)$$

We separate time & space and radial & horizontal components

$$\underline{u}(r, \omega) = \underbrace{U_{lm}(r, \omega) \underline{P}_{lm}(\vartheta, \phi) + V_{lm}(r, \omega) \underline{B}_{lm}(\vartheta, \phi)}_{\text{spheroidal}} + \underbrace{W_{lm}(r, \omega) \underline{C}_{lm}(\vartheta, \phi)}_{\text{toroidal}}$$

$$\phi(r, \omega) = P_{lm}(r, \omega) Y_{lm}(\vartheta, \phi)$$

where $Y_{lm}(\vartheta, \phi)$: spherical harmonics (angular degree l , azimuthal order m)

$\underline{P}_{lm}, \underline{B}_{lm}, \underline{C}_{lm}$: vector spherical harmonics

Assuming small perturbation, i.e., $\underline{\underline{\tau}} = \underline{\underline{\tau}}_0 + \underline{\underline{\tau}}'$ with respect to hydrostatic pressure $\underline{\underline{\tau}}_0 = -\rho_0 \underline{\underline{I}}$, and density $\rho = \rho_0 + \rho'$ with $\rho' = -\underline{\nabla} \cdot (\rho_0 \underline{u})$ leads to a system

$$\left\{ \begin{array}{l} \frac{1}{r^2} \frac{d}{dr} [r^2 (1+2\mu) \dot{U} + \dots] = 0 \\ \frac{1}{r^2} \frac{d}{dr} [\mu r^2 (\dot{V} - \dots)] = 0 \end{array} \right\} \text{spheroidal}$$

$$\left\{ \begin{array}{l} \frac{1}{r^2} \frac{d}{dr} [\mu r^2 (\dot{W} - \dots)] = 0 \\ \ddot{P} + \frac{2}{r} \dot{P} - \dots = -4\pi G \rho U - \dots \end{array} \right\} \text{toroidal} \leftarrow \text{decoupled}$$

For each (l, m) , we find a discrete set of eigenfrequencies ω_{lm} associated with eigenfunctions U_{lm}, V_{lm} and W_{lm} .

Properties of SNREI:

- toroidal-spheroidal decoupling:

W_{lm} (toroidal) independent of U_{lm} & V_{lm} (spheroidal)

and so are eigenfrequencies

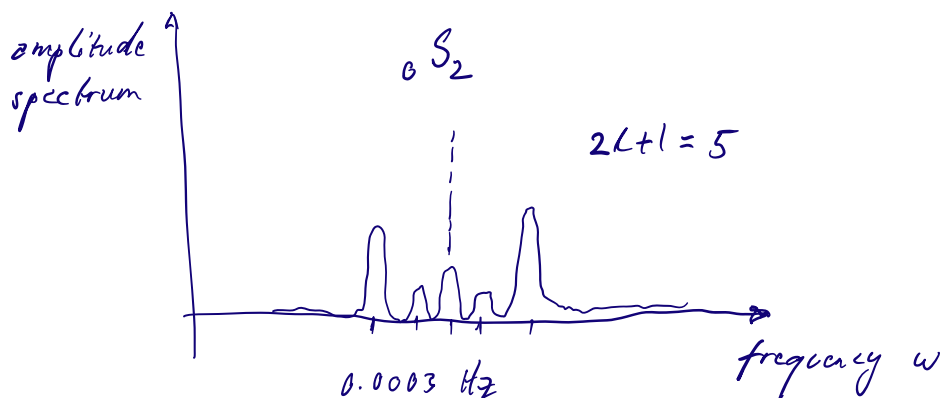
ω_{nlm}^T (toroidal) with n = overtone number

ω_{nlm}^S (spheroidal)

$n=0$ fundamental mode
 $n=1$ 1. overtone

- degeneracy: independent of m ($-l \leq m \leq l$), i.e.,
all $2l+1$ modes are the same of angular
degree l : ${}_n T_l = W_{lm}^n(r) C_{lm}(\vartheta, \varphi)$ multiplet

Rotation, ellipticity and 3D heterogeneity remove the degeneracy
 \Rightarrow for each (n, l) there are $2l+1$ singlets with different
frequencies



Green's tensor & synthetic seismograms:

Modes $s_k(r)$, $k=(n, l, m)$ and eigenfrequency ω_k form a complete basis such that we can write the Green's tensor as

$$\underline{G}(r, r'; t) = \sum_k \frac{1}{\omega_k} \underline{s}_k(r) \underline{s}_k(r') \sin(\omega_k t)$$

"mode summation"

splitting functions:

The contribution of an isolated multiplet to the displacement in the perturbed Earth is given by

$$u(t) = \text{Re} \left\{ \underline{r}^T \cdot e^{i \frac{H}{\omega} t} \cdot \underline{s} e^{i \omega t} \right\}$$

with $\underline{r} = \{ \underline{v} \cdot \underline{u}_m \}$, n entries, $\underline{s} = \{ \underline{M} : \nabla \underline{u}_m \}$, and eigenfrequency ω of degenerate multiplet.

The splitting matrix $e^{i \frac{H}{\omega} t}$ governs multiplet splitting:

$$H_{ij} = \omega \sum_{s,t} c_{st} \int_{\Omega} Y_{li}^* Y_{st} Y_{lj} d\Omega$$

"splitting coefficients"

with corresponding splitting function $f = \sum_{s=0, 2, \dots}^{2L} \sum_{t=-s}^s c_{st} Y_{st}$