

1 **Highly dynamic gamma-ray emissions are common in**
2 **tropical thunderclouds**

3
4 M. Marisaldi^{1,14*}, N. Østgaard^{1***}, A. Mezentsev^{1***}, T. Lang², J. E. Grove³, D. Shy³, G. M. Heymsfield⁴,
5 P. Krehbiel⁵, R. J. Thomas⁵, M. Stanley⁵, D. Sarria¹, C. Schultz², R. Blakeslee², M. G. Quick², H.
6 Christian⁶, I. Adams⁴, R. Kroodsma⁴, N. Lehtinen¹, K. Ullaland¹, S. Yang¹, B. Hasan Qureshi¹, J.
7 Søndergaard¹, B. Husa¹, D. Walker⁶, M. Bateman⁶, D. Mach⁷, S. Cummer⁸, M. Pazos⁹, Y. Pu⁸, P.
8 Bitzer⁶, M. Fullekrug¹⁰, M. Cohen¹¹, J. Montanya¹², C. Younes¹³, O. van der Velde¹², J. A. Roncancio¹²,
9 J. A. Lopez¹², M. Urbani¹², A. Santos¹³

10
11
12
13 ¹ Department of Physics and Technology, University of Bergen, Norway

14 ² NASA Marshall Space Flight Center, Huntsville, USA

15 ³ U.S. Naval Research Laboratory, Washington DC, USA

16 ⁴ NASA Goddard Space Flight Center, Greenbelt, USA.

17 ⁵ New Mexico Institute of Mining and Technology, USA

18 ⁶ Department of Atmospheric Science, Earth System Science Center, University of Alabama in
19 Huntsville, Huntsville, USA

20 ⁷ Universities Space Research Association, Huntsville, AL, 35805, USA

21 ⁸ Duke University, USA

22 ⁹ Instituto de Ciencias de la Atmosfera y Cambio Climatico, UNAM, Mexico

23 ¹⁰ University of Bath, UK

24 ¹¹ Georgia Institute of Technology, USA

25 ¹² Polytechnic University of Catalonia, Spain

26 ¹³ Universidad Nacional de Colombia, Colombia

27 ¹⁴ Astrophysics and Space Science Observatory, National Institute for Astrophysics, Bologna, Italy.

28
29
30
31
32
33
34
35
36
37
38
39
40 *E-mail: Martino.Marisaldi@uib.no, orcid.org/0000-0002-4000-3789

41 **E-mail: Nikolai.Ostgaard@uib.no, orcid.org/0000-0002-2572-7033

42 ***E-mail: Andrey.Mezentsev@uib.no, orcid.org/0000-0002-3471-7267

43
44
45
46
47

48 Thunderstorms emit fluxes of gamma rays known as gamma-ray glows¹⁻², sporadically
49 observed by aircraft^{1,3-7}, balloons⁸⁻¹¹ and from ground¹²⁻¹⁸. Observations report increased
50 gamma-ray emissions by tens of percent up to two orders of magnitude above the
51 background, sometimes abruptly terminated by lightning discharges^{1,3-5}. Glows are
52 produced by the acceleration of energetic electrons in high electric field regions within
53 thunderclouds⁸, and contribute to charge dissipation³. Glows had been considered as
54 quasi-stationary phenomena^{3,5,12}, with durations up to a few tens of seconds and spatial
55 scales up to 10-20 kilometers. However, no measurement of the full extension in space
56 and time of a gamma-ray glow region and their occurring frequency have been reported
57 so far.

58 Here we show that tropical thunderclouds over ocean and coastal regions commonly emit
59 gamma rays for hours over areas up to a few thousands of square kilometers. Emission
60 is associated with deep convective cores; it is not uniform and continuous but shows
61 characteristic timescales of 1-10 seconds and even sub-second for individual glows. The
62 dynamics of gamma-glowing thunderclouds starkly contradicts the quasi-stationary
63 picture of glows, but rather resembles that of a huge gamma-glowing «boiling pot» both
64 in pattern and behavior.

65

66

67 During the Airborne Lightning Observatory for FEGS and TGFs (ALOFT) aircraft campaign
68 over Caribbean and Central America in the summer of 2023 the simplistic picture of localized
69 and uniformly gamma-ray glowing thunderclouds was overridden. Due to a novel mission
70 concept, where one-second resolution data were downlinked during flight, gamma-ray
71 glowing clouds could be identified in real time, and the pilot was instructed to return to the
72 same location until the thundercloud stopped glowing. ALOFT detected more than 500 1-to-
73 10-second-long individual gamma-ray glows during 9 of the 10 flights, all over ocean or
74 coastal regions (see Extended Data 10), showing that thunderclouds can emit gamma rays for
75 hours and over huge regions.

76

77 Figure 1a shows the flight path on 24th July 2023 above a very active mesoscale convective
78 system over southern Campeche Bay. Glows with fluxes up to ~12 times the background
79 level (on one-second time scales) were detected during almost all the three hours spent on
80 target (Figure 1b). Glows were detected repeatedly following consecutive passages over the
81 same thundercloud system (see crossing trajectories and black dots in Figure 1a). The overall
82 region including the cyan and magenta trajectories in Figure 1a was glowing for at least the 3
83 hours we stayed on target, starting at 05:15 UT. We do not know whether the thundercloud
84 was glowing before we arrived and after we had to leave. The glowing region corresponds to
85 an area of more than 9000 km².

86

87 Our glow observations show significant temporal variability down to the 10-millisecond time
88 scale, overriding the previous picture of quasi-steady smoothly varying glows. Figure 2
89 shows successively zoomed-in observations of the gamma-ray count rate during an 8-minute-
90 long pass and glow sequence, labelled G1 and highlighted in magenta in Figure 1 (~100 km
91 aircraft flight distance). Figures 2c and 2e show the individual few-second-long glows (< 2
92 km aircraft flight distance). Figure 2f shows a “glow burst” with brightening by a factor ~3
93 over 20 ms, followed by a gradual flux reduction to the previous level on the same timescale.
94 On a 1-ms timescale, this corresponds to an enhancement above the background by a factor

~30. A duration of few tens of milliseconds is three orders of magnitude lower than the typical duration of glows reported from previous observations (tens of seconds to minutes) yet is much longer than that of Terrestrial Gamma-ray Flashes (TGF) (~10 - ~100 μ s), which are bright very-short transients usually associated with lightning discharges¹⁹⁻²². Transients tens of milliseconds long have been presented²³ and have been related to positron clouds. The glow burst presented here does not show any evidence for an enhanced positron annihilation line at 511 keV (see spectral analysis in Methods 4c and Extended Data 9b) and therefore must be considered as a new phenomenon. Figure 2d shows another glow that evolved into repetitive behavior termed a Flickering Gamma Flash (FGF)²⁴, in which the intensity is modulated in a succession of millisecond-duration spikes, separated by tens of milliseconds. Both the occurrence of FGFs and glow bursts lasting only few tens of milliseconds contradict the canonical picture of glows as quasi-static phenomena well separated from the very short transient TGFs, rather suggesting a more complex scenario linking these two phenomena.

Figure 3 shows how the glow episodes of a typical pass (in this case the cyan-highlighted episode G2 of Figure 1) are related to the storm structure, as determined by the planar cloud-top brightness temperature (Figure 3b) and the vertical reflectivity profile (Figure 3c). Together, the observations clearly indicate three convective cores producing high radar reflectivity, each 20-25 km in horizontal extent, with cloud tops extending up to ~16.5 km altitude. This implies large liquid and ice contents as well as strong convective updrafts at high altitudes in the cores. The upward-pointing electric field vectors (Figure 3c) are consistent with the charge structure of thunderclouds typically having an upper positive charge layer, which is supported also by ground observations for the flight on 29.07.2023 over Florida (see Methods 3b and Extended Data 5 and 6).

The above correlations pertain to episode G2 of the 24.07.2023 observations, but similar correlations are obtained for episode G1 and overflights of other storms, as seen in Extended Data 2 for an earlier 06.07.2023 Campeche Bay storm, and Extended Data 4 for the 29.07.2023 Florida storm. In each case the gamma-ray glow activity is remarkably temporally aligned with the overpass of successive cores, with an overall modulation in maximum intensity qualitatively consistent with the cloud-top altitude. Gamma-ray episodes over cores consisted of several (up to few tens) individual glows, often partially overlapping. The temporal profile of individual glows is typically asymmetric, with a slow rise lasting a few seconds and a fast decay (<1 s). This scenario is a common pattern observed throughout the campaign (9 of 10 flights), see Methods 3. Each glow episode (a passage over a convective core, encompassing many individual glows) is also accompanied by an increase in the local rate of electrical discharges detected by the Electric Field Change Meter (EFCM) (see Extended Data 2 and 4, and Methods 2).

To summarize, we can state with certainty that a thundercloud region can glow for hours and extend over several thousands of square kilometers. The longest continuous enhancement previously reported is 40 minutes²⁵ but this was only a 2% increase above background over a 1-minute time bin, not even close or comparable to the glows we report here with increases of 10-30 times the background. We can also state that each glowing episode typically lasts for about 2 minutes (20-25 km aircraft flight distance). For the individual glows (1 to ~10 s duration), however, we cannot disentangle the spatial and temporal variability components from gamma-ray observations, therefore we cannot establish the physical extent of the single source regions, whether they cover the full extent of the convective cores or if they are more compact regions within the cores. Monte Carlo simulations suggest an effective range for glow observations of 2 to ~4.5 km, depending on production altitude (see Methods 4a and

145 Extended Data 7). Therefore, active regions must be located within the cores. Typical rise
146 time (2-4 seconds) and fall time (<1 second) of individual glows are too short compared to
147 the characteristic times expected from aircraft motion towards a stationary source (see
148 Methods 4b and Extended Data 8). Therefore, the temporal dynamics of individual glows is
149 most likely dominated by the intrinsic time variability of the gamma-ray emission, which
150 reflects the variability of the underlying electric field. We expect the rise time of glows to be
151 associated with electric field increase due to convection-driven charge build-up, and the
152 shorter fall time to be associated with electric field reduction due to charge removal following
153 discharge processes.

154

155 In most cases, a glow flux reduction does not bring the observed flux down to the background
156 level, but rather to an intermediate level immediately followed by a new rise phase (Figure 2c
157 and 2e). We cannot state whether this is due to a partial discharge followed by a recharging of
158 the same active region, or by the superposition of fluxes coming from nearby active regions
159 within the same core. The overall time profile of glow episodes, with the brightest glows
160 typically observed passing over the center of the convective cores (see Figure 3) suggests that
161 most of the gamma rays do not come from a large-scale, core-sized active region, but rather
162 from one or more compact regions concentrated towards the center of the cores. For the sub-
163 second glow bursts with only tens of millisecond rise and fall times we can with high
164 confidence state they are due to temporal variability of the source region.

165

166 The observations suggest what we call ‘the boiling pot analogy’. Gamma-ray glowing regions
167 are tightly associated to strong convective cores, with several individual cores evolving
168 during the development and mature phases of tropical thunderstorms. Local gamma-ray
169 emission follows the lifetime of the associated core, but particularly with large multicell
170 storms the overall gamma-ray glowing region can last for hours and extend over thousands of
171 square kilometers, tightly matching the most active regions of the storm.

172

173 So far, modeling efforts have been based on simplistic assumptions on the extent of the glow
174 region and electric field configurations²⁶⁻²⁷, namely uniform and constant electric field
175 extended over large (several square km) regions. Our observations prompt the need to go
176 beyond this simplistic static representation of glows. Inhomogeneous charge structure and its
177 temporal dynamics must be considered in realistic simulations. The pervasiveness of gamma-
178 ray emission over large time and spatial scales in ocean and coastal tropical thunderclouds
179 has a broad range of implications, barely touched on in previous studies, spanning from their
180 effects on cloud discharge³, alteration of the local environment due to isotope production²⁸,
181 and potential effect on cloud chemistry and dynamics.

182

183

184 REFERENCES

185

186

- 187 1. Parks, G.K., et al., *X-ray enhancements detected during thunderstorm and lightning activities*.
188 Geophysical Research Letters, **8**(11): p. 1176-1179 (1981).
- 189 2. Dwyer, J.R., D.M. Smith, and S.A. Cummer, *High-Energy Atmospheric Physics: Terrestrial
190 Gamma-Ray Flashes and Related Phenomena*. Space Science Reviews, **173**(1-4): p. 133-196
191 (2012).
- 192 3. Kelley, N.A., et al., *Relativistic electron avalanches as a thunderstorm discharge competing
193 with lightning*. Nat Commun, **6**: p. 7845 (2015).
- 194 4. Ostgaard, N., et al., *Gamma Ray Glow Observations at 20-km Altitude*. Journal of
195 Geophysical Research-Atmospheres, **124**(13): p. 7236-7254 (2019).

- 196 5. Kochkin, P., et al., *In-Flight Observation of Gamma Ray Glows by ILDAS*. Journal of
197 Geophysical Research-Atmospheres, **122**(23): p. 12801-12811 (2017).
- 198 6. McCarthy, M. and G.K. Parks, *FURTHER OBSERVATIONS OF X-RAYS INSIDE*
199 *THUNDERSTORMS*. Geophysical Research Letters, **12**(6): p. 393-396 (1985).
- 200 7. McCarthy, M.P. and G.K. Parks, *ON THE MODULATION OF X-RAY FLUXES IN*
201 *THUNDERSTORMS*. Journal of Geophysical Research-Atmospheres, **97**(D5): p. 5857-5864
202 (1992).
- 203 8. Eack, K.B. and W.H. Beasley, *Long-duration X-ray emissions observed in thunderstorms*.
204 Journal of Geophysical Research: Atmospheres, **120**(14): p. 6887-6897 (2015).
- 205 9. Eack, K.B., et al., *X-ray pulses observed above a mesoscale convective system*. Geophysical
206 Research Letters, **23**(21): p. 2915-2918 (1996).
- 207 10. Eack, K.B., et al., *Initial results from simultaneous observation of X rays and electric fields in*
208 *a thunderstorm*. Journal of Geophysical Research-Atmospheres, **101**(D23): p. 29637-29640
209 (1996).
- 210 11. Eack, K.B., et al., *Gamma-ray emissions observed in a thunderstorm anvil*. Geophysical
211 Research Letters, **27**(2): p. 185-188 (2000).
- 212 12. Wada, Y., et al., *Catalog of gamma-ray glows during four winter seasons in Japan*. Physical
213 Review Research, **3**(4) (2021).
- 214 13. Chilingarian, A., *Thunderstorm ground enhancements-Model and relation to lightning*
215 *flashes*. Journal of Atmospheric and Solar-Terrestrial Physics, **107**: p. 68-76 (2014).
- 216 14. Chilingarian, A., B. Mailyan, and L. Vanyan, *Recovering of the energy spectra of electrons*
217 *and gamma rays coming from the thunderclouds*. Atmospheric Research, **114-115**: p. 1-16
218 (2012).
- 219 15. Chilingarian, A., G. Hovsepyan, and A. Hovhannisyan, *Particle bursts from thunderclouds:*
220 *Natural particle accelerators above our heads*. Physical Review D, **83**(6) (2011).
- 221 16. Chilingarian, A., G. Hovsepyan, and L. Vanyan, *On the origin of the particle fluxes from the*
222 *thunderclouds: Energy spectra analysis*. EPL (Europhysics Letters), **106**(5) (2014).
- 223 17. Tsuchiya, H., et al., *Hardening and Termination of Long-Duration γ Rays Detected Prior to*
224 *Lightning*. Physical Review Letters, **111**(1) (2013).
- 225 18. Tsuchiya, H., et al., *Long-duration ray emissions from 2007 and 2008 winter thunderstorms*.
226 Journal of Geophysical Research, **116**(D9) (2011).
- 227 19. Smith, D.M., et al., *Terrestrial gamma-ray flashes observed up to 20 MeV*. Science, **07**(5712):
228 p. 1085-8 (2005).
- 229 20. Marisaldi, M., et al., *Detection of terrestrial gamma ray flashes up to 40 MeV by the AGILE*
230 *satellite*. Journal of Geophysical Research: Space Physics, **115**(A3) (2010).
- 231 21. Briggs, M.S., et al., *First results on terrestrial gamma ray flashes from the Fermi Gamma-ray*
232 *Burst Monitor*. Journal of Geophysical Research: Space Physics, **115**(A7) (2010).
- 233 22. Østgaard, N., et al., *First 10 Months of TGF Observations by ASIM*. Journal of Geophysical
234 Research-Atmospheres, **124**(24): p. 14024-14036 (2019).
- 235 23. Dwyer, J.R., et al., *Positron clouds within thunderstorms*. Journal of Plasma Physics, **81**(4)
236 (2015).
- 237 24. Østgaard, N.; Mezentsev, A.; Marisaldi, M.; Grove, J. E.; M. Quick, M.; et al., *Flickering*
238 *Gamma-Ray Flashes, the Missing Link between Gamma Glows and TGFs*. Nature, this issue
239 (2024).
- 240 25. Tsuchiya, H., et al., *Observation of thundercloud-related gamma rays and neutrons in Tibet*.
241 Physical Review D, **85**(9), (2012).
- 242 26. Sarria, D., et al., *Library of Simulated Gamma-Ray Glows and Application to Previous*
243 *Airborne Observations*. Journal of Geophysical Research-Atmospheres, **128**(9) (2023).
- 244 27. Wada, Y., et al., *Negative Excursion of Surface Electric Fields During Gamma-Ray Glows in*
245 *Winter Thunderstorms*. Journal of Geophysical Research: Atmospheres, **128**(21): p.
246 e2023JD039354 (2023).
- 247 28. Enoto, T., et al., *Photonuclear reactions triggered by lightning discharge*. Nature, **551**(7681):
248 p. 481-484 (2017).

250 **FIGURE LEGENDS**
251

252 **Fig. 1. Map and gamma-ray sequence for the flight on 24.07.2023.**

253 a) Flight trajectory (grey) overlaid on GOES-18 infra-red temperatures, along with marked
254 gamma-ray fluxes exceeding 25% above the background level (black dots). The magenta
255 trajectory corresponds to the glow episodes plotted in Figure 2b, also marked as G1 in panel
256 b. The cyan trajectory corresponds to the glow episodes plotted in Figure 3, also marked as
257 G2 in panel b. For both trajectories the direction of motion of the plane is from left to right.
258 The brightness temperature cloud image corresponds to time 07:24:251 UT, close to the
259 maximum emission in the cyan region. b) BGO count rate. Time on target: 05:00 UT – 08:20
260 UT.

261
262 **Fig. 2. Gamma-ray glow temporal variability.**

263 a) BGO (black) and iSTORM (red) count rates (1s bin) for 24.07.23. b) zoomed view (100 ms
264 bin) for the 10-minutes time interval highlighted in black in panel a, corresponding to time
265 interval G1 highlighted in magenta in Figure 1. c, e) zoomed view (10 ms bin) for the 8-seconds
266 time intervals highlighted in black in panel b. d, f) zoomed view (1 ms bin) for the 1-second
267 time intervals highlighted in black in panel c and e, respectively.

268
269
270 **Fig. 3. Sequence of glows with simultaneous electric field and cloud characterization data.**
271 Time interval G2 on 24.07.2023 highlighted in cyan in Figure 1. a) BGO count rate. b)
272 Brightness Temperature (TB) from the Advanced Microwave Precipitation Radiometer
273 (AMPR) 85.5 GHz Channel A. Sweep Index 0 (49) corresponds to scan direction 45 degrees
274 from Nadir towards starboard (port). c) vertical radar reflectivity profile from the X-band radar
275 (EXRAD). Black vectors are the E_{xz} component of the electric field from the Lightning
276 Instrument Package (LIP), where x-axis is the direction of motion and z-axis is the upward-
277 pointing vertical direction. AMPR, EXRAD and LIP (see Method 2) data are not shown if roll
278 angle is larger than 2 degrees (aircraft maneuvering).

279
280

281 **Acknowledgements**
282

283 This work made use of data from UIB-BGO, iSTORM, EXRAD, AMPR, LIP, EFCM, from the ALOFT
284 instrument suite, VHF data from the Central Florida Lightning Mapping Array (CFLMA), and ABI data
285 from GOES-16 and GOES-18.

286 The ALOFT campaign and the UIB-BGO instrument was supported by European Research Council
287 under the European Union's Seventh Framework Programme (FP7/2007–2013)/ERC grant agreement
288 no. 320839 and the Research Council of Norway under contracts 223252/F50 (CoE) and contract
289 325582. Some of the simulations were performed on resources provided by UNINETT Sigma2—the
290 National Infrastructure for High Performance Computing and Data Storage in Norway, under project
291 no. NN9526K. Work at NRL on ALOFT is supported by the Office of Naval Research 6.1 funds. In
292 addition, D. Shy is supported by the U.S. Naval Research Laboratory's Jerome and Isabella Karle
293 Fellowship. The New Mexico Tech authors acknowledge NSF Grants 1720600 and 2214044, Robert
294 Brown of the University of Central Florida for setting up and operating the CFLMA stations, and
295 William Rison for providing the data. The participation of J.A.R, J.L., M.U., O.vdV. and J.M. and the
296 fielding of instruments on San Andrés island was supported by projects EQC2021-006957-P and
297 PID2022-136348NB-C33 by the government of Spain (MCIN/AEI/10.13039/501100011033) and the
298 European Regional Development Fund “ERDF - A way of making Europe” by the European Union.
299 M.F. was sponsored by the Royal Society (UK) grant NMG/R1/180252 and the Natural Environment
300 Research Council (UK) under grants NE/L012669 /1 and NE/H024921/1. S.A.C. and Y.P. were partly
301 supported by the National Science Foundation Dynamic and Physical Meteorology program through
302 Grant AGS-2026304. The authors thank the NASA ER-2 Project Team at NASA Armstrong Flight
303 Research Center, for supporting the ALOFT campaign, and the MacDill Air Force Base for hosting.
304 Significant financial and logistical support for ALOFT was provided by the NASA Earth Science
305 Division. We thank the governments of Mexico, Bahamas, Colombia, Belize, Guatemala, Honduras,
306 Nicaragua, Panama, Dominican Republic, Costa Rica, El Salvador, Haiti, Turks & Caicos, Jamaica,
307 and the Cayman Islands for approving ER-2 overflights in support of ALOFT. GOES MDS sectors in
308 support of ALOFT were enabled by NOAA. The FEGS and EFCM team acknowledge the work of Scott
309 Podgorny, David Corredor, and Mike Stewart. For AMPR we thank Douglas Huie (UAH), Corey Amiot
310 (NASA Postdoctoral Program), and Eric Cantrell (NASA MSFC). We acknowledge the High-altitude
311 Radar group at NASA Goddard that provided data including Matthew McLinden, Lihua Li, Peter
312 Pantina, and Charles Helms.

313
314 **Author Contributions**
315

316 M.M, N.O. and A.M. have led this study. The ALOFT campaign was led by N.O., T.L. M.M. and C.S.
317 The UIB-BGO instrument was provided by N.O., M.M., K.U., S.Y., B.H.Q, J.S and B.H. and analyzed
318 by A.M., N.O., M.M., D.S., N.L. The iSTORM data were provided and analyzed by J.E.G., D.S. and
319 D. W. The AMPR data were provide by T.L. The EXRAD data were provided by G.H. The LIP data
320 were provided by C.S, D.M. and M.B. The EFCM data were provided by H.C. and R.B. The CFLMA
321 data was analyzed by P.K., R.T., and M.S. Additional radar and radiometric data for the entire campaign
322 were provided by I.A. and R.K. Additional ground radio stations were operated for the entire campaign
323 by S.A.C., Y.P., M.P., P.B., M.F., M.C., J.M, C.Y., O.vdV., J.A.R, J.A.L., M.U. and A.S.

324
325 **Competing Interests** The authors declare that they have no competing financial interests.
326

327 **Correspondence** Correspondence and requests for materials should be addressed to M. Marisaldi
328 (Martino.Marisaldi@uib.no), N. Østgaard (Nikolai.Ostgaard@uib.no) and A. Mezentsev
329 (Andrey.Mezentsev@uib.no).
330
331
332

333 **METHODS**

334

335 **1. THE ALOFT FLIGHT CAMPAIGN**

336

337 The Airborne Lightning Observatory for FEGS* and TGFs (ALOFT) is an aircraft campaign
338 conducted during the month of July 2023 with a NASA ER-2 aircraft based at the McDill Air
339 Force Base in Tampa, Florida, USA. A total of ten flights (3-8 hours each) at 20 km altitude
340 were performed. Each flight spent 3-4 hours above active thunderstorms. Extended Data 10
341 shows an overview of the flight dates, durations, target locations, and total number of
342 observed glow episodes and gamma-ray glows.

343

344 The science goals of ALOFT were the observation of Terrestrial Gamma-ray Flashes (TGF)
345 and gamma-ray glows from thunderstorms and the study of their connection. The airborne
346 scientific payload consists of five independent gamma-ray detectors, 30 photometers, three
347 electric field sensors, two radars and two passive radiometers. In addition, nine ground-based
348 radio receivers were operated during the campaign, covering Very Low Frequencies (VLF),
349 Low Frequencies (LF) and Very High Frequencies (VHF), including two interferometers and
350 one Lightning Mapping Array (LMA). The instruments used in this study are described in
351 Methods Section 2.

352

353 A key feature of the mission was the near real-time downlink of one-second time resolution
354 gamma-ray count rates. This enabled the immediate identification of gamma-ray glowing
355 thunderclouds, and the pilot was instructed to return on the same location as long as gamma-
356 ray glows were detected.

357

358

359 **2. INSTRUMENTATION**

360

361 ***The Bismuth-Germanate instrument from the University of Bergen (UIB-BGO).***

362 This instrument consists of four independent gamma-ray detectors; one Bismuth Germanate
363 (BGO) detector, with three independent BGO crystals with Photomultiplier Tube (PMT)
364 readout, and three LYSO detectors. All detectors have different geometric areas ranging from
365 0.09 cm² to 225 cm², in order to provide four orders of magnitude dynamic range in flux. In
366 fact, modelling work²⁹ showed that the flux from a typical TGF observed at 20 km altitude
367 can vary by 4 orders of magnitude, depending on the radial distance from the source to the
368 aircraft. The three BGO/PMT detectors are similar in design and readout architecture to one
369 of the four High Energy Detector modules of the Modular X- and Gamma-ray Sensor
370 (MXGS)³⁰ onboard the Atmosphere Space Interaction Monitor (ASIM) on the International
371 Space Station (ISS). The three BGO/PMT detectors have a total geometrical area of 225 cm²,
372 are sensitive in the energy range 300 keV - >30 MeV, and have a 28 ns time-tagging
373 accuracy. In this work we use only data from the BGO detectors. Additional information on
374 UIB-BGO can be found in the companion paper²⁴.

375

376 ***The in-Situ Thunderstorm Observer for Radiation Mechanisms (iSTORM),***

377 iSTORM is a gamma-ray spectrometer designed and built by the US Naval Research
378 Laboratory for the measurement of bright, fast transients in the energy range ~300 keV - >5
379 MeV). iSTORM consists of an array of 32 1-inch diameter CeBr₃ scintillating crystals with
380 Silicon Photomultiplier (SiPM) readout, with temporal resolution <1μs, for a total

* FEGS = Fly's Eye GLM Simulator, GLM=Geostationary Lightning Mapper

381 geometrical area of 157 cm². The large area provides high sensitivity, while the high
382 segmentation and fast scintillation decay time provides robustness towards potential detector
383 polarization from bright TGFs. Additional information on iSTORM can be found in the
384 companion paper²⁴.

385

386 ***The Advanced Microwave Precipitation Radiometer (AMPR)***

387 The Advanced Microwave Precipitation Radiometer (AMPR) is a cross-track scanning
388 radiometer that passively measures total power at 10.7 GHz, 19.35 GHz, 37.1 GHz, and 85.5
389 GHz³¹. The instrument uses a rotating splash plate mirror to focus microwave energy into its
390 feedhorns. This causes the polarization of the incoming signal to vary sinusoidally as a
391 function of scan angle. AMPR has two orthogonally polarized channels per frequency, so it is
392 possible to retrieve fully vertical and fully horizontal polarizations by deconvolving the
393 orthogonal measurements at each frequency. At ER-2 altitudes, AMPR's instantaneous field-
394 of-view (IFOV) resolution varies between 0.6-2.8 km, with the finest resolution associated
395 with 85.5 GHz and the coarsest resolution associated with 10.7 GHz. AMPR's swath is ~38
396 km wide when flying at 20 km altitude on the ER-2. AMPR completes a cross-track scan in
397 ~2.5 seconds, and after 5 scans the instrument collects measurements of two blackbodies – a
398 heated hot load and an air-cooled cold load – to enable calibrated brightness temperature
399 (TB) retrievals.

400 AMPR TBs are first retrieved using a simple two-point linear method that compares scene
401 radiometer counts to the radiometer counts from the hot load (typically ~320 K) and from the
402 cold load (which can reach a minimum temperature of ~240 K during a flight). With most
403 meteorological scenes, this methodology is capable of producing well-calibrated TBs suitable
404 for quantitative geophysical retrievals³¹. However, during ALOFT it was found that, due to
405 the large amount of ice scattering observed in these storms, a linear assumption was not
406 suitable for the 85-GHz channels, as the scene targets were so radiometrically cold that
407 receiver response entered a nonlinear regime. Thus, AMPR 85-GHz TBs were corrected for
408 receiver nonlinearity using a simple gain adjustment factor which became stronger as scene
409 counts went lower than cold load counts, and thus prevented negative TB retrievals. This
410 correction also reproduced 85-GHz TBs that were within approximately +/- 10 K of previous
411 AMPR observations of intense convection³²⁻³³ and that remained highly correlated with 85-
412 GHz scene counts (Pearson correlation coefficient > 0.9). The corrected 85-GHz TBs are thus
413 suitable for qualitative identification of significant precipitation cores and the presence of ice
414 scattering.

415

416 ***The ER-2 X-band Radar (EXRAD)***

417 The ER-2 X-band radar (EXRAD) is an X-band (9.6 GHz) precipitation radar and
418 scatterometer that flies in the ER-2 nose and it is used for measuring both the three-
419 dimensional precipitation structure and surface winds. It has a fixed nadir beam and a conical
420 scanning beam that is tilted approximately 30 degrees off nadir. EXRAD measures radar
421 reflectivity, Doppler velocity, and Doppler spectral width. It was first flown in 2014. EXRAD
422 uses a 0.66-m diameter slotted waveguide antenna with a single linear polarization for the
423 scanning beam plus a second 0.66-m linear polarized slotted waveguide antenna for a nadir
424 beam. This ~ 3.4 degree beam provides a beamwidth of approximately 0.6 km at 10 km
425 altitude, and 1.2 km at the surface. EXRAD calibration is performed using the ocean surface,
426 and current Doppler processing includes non-uniform beam filling (NUBF) corrections for
427 artifacts caused by the aircraft motion. Calibration accuracy is better than 1 dBZ for
428 reflectivity and <1 ms⁻¹ for Doppler. More details on EXRAD hardware and processing can
429 be found in³⁴.

430

431

432 ***The Lightning Instrument Package (LIP)***

433 The Lightning Instrument Package (LIP) consists of seven rotating vane electric field mills
 434 installed on the NASA ER-2 aircraft to observe the total vector electric field and vector
 435 electric field (E_x , E_y , E_z) generated by cloud charging and lightning discharges³⁵⁻³⁸. LIP
 436 measures the vector electric field in the atmosphere and the charge induced on the aircraft
 437 using the processing and calibration technique in³⁹. The individual laboratory dynamic range
 438 for the mills is $\pm 1.75 \text{ V m}^{-1}$ to 920 kV m^{-1} . Properly calibrated on the aircraft, LIP can
 439 reliably measure fields lower than 1 V m^{-1} and as high as 512 kV m^{-1} with 0.1 s temporal
 440 resolution. Effects on the derived vertical electric field due to aircraft charging are generally
 441 5% or less, with a maximum upper error of 10%³⁶. Both aircraft relative and Earth relative
 442 frameworks are determined, and the total electric field magnitude is then computed from the
 443 vector components.

444

445 ***The Electric Field Change Meter (EFCM)***

446 The Electric Field Change Meter (EFCM) is a two-channel (fast and slow) antenna that
 447 measures the derivative of the electric field impulse produced by lightning. The fast channel
 448 is designed to isolate the radiative component of the lightning discharge field while the slow
 449 channel is optimized to observe the electrostatic field component. The EFCM has multiple
 450 sensitivity ranges that are selectable during flight and samples with 16-bit resolution. The
 451 EFCM trigger rate shown in Extended Data 2 and 4 is clipped at a maximum of ten triggers in
 452 ten seconds. This is because an EFCM trigger implies acquisition of one-second data around
 453 the trigger time. Such rate is therefore the maximum trigger rate possible, implying nearly
 454 continuous data recording. The real discharge rate may be much higher. Additional
 455 information on EFCM can be found in the companion paper²⁴.

456

457 ***The Central Florida Lightning Mapping Array (CFLMA)***

458 The Central Florida Lightning Mapping Array (CFLMA) consisted of 6 widely-spaced VHF
 459 stations that receive the VHF radiation produced by lightning in a 6 MHz bandwidth between
 460 60 and 66 MHz (U.S. TV Channel 3). Each station detects the peak radiation event above a
 461 local threshold value in successive 80 microsecond time windows with 40 ns time resolution
 462 and utilizes the arrival times at the different sites to locate the sources of impulsive radiation
 463 sources in 3 spatial dimensions and time, as well as the VHF source powers. Owing to
 464 logistical and sensitivity issues, a small number of events were detected per flash, but these
 465 were sufficient both for obtaining the basic location of individual flashes and for determining
 466 the overall electric charge structure of the storms. Important high-power events were well-
 467 detected and located by the network.

468

469 ***Advanced Baseline Imager (ABI)***

470 This work makes use of public data from the Advanced Baseline Imager (ABI) instruments
 471 onboard the GOES-16 and GOES-18 geostationary satellites. ABI provides radiometric
 472 measurements in 16 spectral bands from the visible ($0.47 \mu\text{m}$) to the infrared ($13.3 \mu\text{m}$), with
 473 spatial resolution from 0.5 km (visible) to $\sim 2 \text{ km}$ (infrared). The temporal resolution depends
 474 on the ABI operational mode. For seven flights out of ten, the ALOFT mission was granted a
 475 dedicated Mesoscale Domain Sector (MDS) which provides scans of a $1000 \times 1000 \text{ km}^2$ box
 476 every 60 seconds. For the three flights over Florida, we made use of the Continental US
 477 (CONUS) scans, which are provided every five minutes. In this work we use radiance
 478 measurements in the ABI band 13 (infrared, $10.3 \mu\text{m}$) which are then converted to brightness
 479 temperatures by applying the Planck function and a spectral bandpass correction according to
 480 the GOES-R Series Product Definition and Users' Guide⁴⁰.

481
482
483 **3. DURATION AND EXTENSION OF GAMMA-RAY GLOW ACTIVITY**
484

485 In this section we present information on two other flights: 06.07.2023 over Campeche Bay,
486 and 29.07.2027 over Florida. In addition to the very active storm on 24.07.2023 the other two
487 case studies (shown here), as well as the 6 other storms all show that the novel findings of
488 this paper (hour-long durations, large spatial extension of gamma-ray activity, significant
489 variability on second to millisecond timescales) are common features also in less severe
490 storms.

491
492 For all case studies we present time-based color-coded maps and gamma-ray sequences
493 (Extended Data 1 and 3) which provide a visualization of gamma-ray data alternative to that
494 presented in Figure 1 in the main text. Here the reader can match a specific glow episode with
495 a position on the map based on the highlight color. In addition, we provide gamma-ray count
496 rate with simultaneous electric field and cloud characterization data for selected time
497 intervals (Extended Data 2 and 4). These figures show the same observation pattern presented
498 in Figure 3 in the main text: glow episodes, each of them comprising several individual
499 glows, are observed during the overpass of convective cores with large radar reflectivity,
500 implying large liquid and ice contents as well as strong convective updrafts at high altitudes.
501 In all cases, the upward-pointing electric field vectors are consistent with the presence of an
502 upper positive charge layer. In addition, Extended Data 2 and 4 show EFCM trigger rate
503 which peaks during the passage over gamma-ray active convective cores. EFCM data were
504 not available for the flight on 24.07.2023.

505
506 In the following we report additional information for the case-study flights on 06.07 and
507 29.07.

509 **a) Flight on 06.07.2023**
510

511 The target area showed developing multicell convection that evolved into a MCS at the time
512 of overflight. The area is Campeche Bay, same as for the flight on 24.07.2023 discussed in
513 the main text. This storm represents a very clean case illustrating the large spatial extension
514 and long duration of a gamma-ray glowing thundercloud region. Extended Data 1 shows six
515 consecutive passes over the same core region, each associated with glow episodes (Extended
516 Data 2). The total time span of these passes is 40 minutes, which represents a lower limit of
517 the duration of the glow activity. The overall spatial extension of the glowing region is about
518 600 km².

519
520 **b) Flight on 29.07.2023: comparison with lightning activity and storm charge
521 structure.**

522
523 The flight over east-central Florida on July 29 produced observations similar to those
524 obtained over the Central America coastal areas, but included important observations of VHF
525 radiation from the 3-D Central Florida Lightning Mapping Array (CFLMA) over the area.
526 The Florida storm cells were more localized than those of the other flights and produced
527 relatively well-separated glows, as seen in Extended Data 3 and 4. Extended Data 3 shows
528 seven consecutive passes over the same ~400 km² core region, each of them resulting in
529 significant glow episodes, for a total duration of about one hour.

530
531 The strongest episode occurred at 20:30:30 UT (Extended Data 3b), and is shown in detail in
532 Extended Data 5, which compares the glow sequence with LMA observations of the lightning
533 activity. The episode consisted of three glows, the last one much dimmer than the others.
534 Here we focus on the first two glows, each of which was rapidly quenched at the time of an
535 intracloud lightning flash in the storm. The glows were noticeably different from each other,
536 with the first glow lasting 8 seconds and being detected as the ER-2 entered the periphery of
537 the electrically active storm from the northeast (top horizontal arrow in panel 8c). In contrast,
538 and by sheer chance, the second glow occurred 20 seconds later when the ER-2 was directly
539 above the central part of the storm and the flash that terminated the glow. Instead of
540 increasing linearly with time, as in the first glow, the second glow grew exponentially for 4
541 seconds to twice the count rate of the first glow (namely to ~20,000 per second), at which
542 point a highly energetic intracloud flash was initiated that quenched the glow (red sources in
543 the panels of Extended Data 5, with the ER-2 location indicated by the bottom horizontal
544 arrow). That the quenching flash was energetic is indicated by two of the initial LMA
545 sources having extremely strong VHF powers of 53.9 and 52.9 dBW, or ~200 kW. (The
546 quenching flash for the first glow had moderate source powers, and corresponded to the
547 light green sources in the same plan area and altitude range as the subsequent red sources in
548 Extended Data 5c.)
549

550 Extended Data 6 shows the lightning-inferred charge structure around the time of the 20:30
551 glows, showing that the storm had a normal-polarity tripolar structure, with upper positive
552 charge between ~13-15 km and mid-level negative charge at ~9-12 km, and confirming that
553 the first glow was observed on the periphery of both charge regions, while the second glow
554 was observed directly above the core. The polarity is inferred by virtue of intracloud flashes
555 producing higher-power RF-noisy negative polarity breakdown through the storm's upper
556 positive charge, and less noisy positive breakdown within the storm's mid-level charge⁴¹.
557
558

559 **4. MONTE CARLO SIMULATIONS**

560

561 Using Monte Carlo simulations, we can evaluate certain theoretical characteristics of gamma-
562 ray glows. It is anticipated that all observed spectra originate from the Relativistic Runaway
563 Electron Avalanche (RREA) process. We utilized the GEANT4 software⁴², which enables the
564 simulation of photon, electron, and positron propagation in the atmosphere. The simulation
565 initiates with a point source of upward-pointing electrons, with a classical RREA energy
566 spectrum up to 40 MeV at the source⁴³. The simulations follow the following steps: 1)
567 Propagation, scattering, and absorption of RREA electrons in the atmosphere, along with the
568 generation and propagation of secondary Bremsstrahlung photons from various altitudes
569 (ranging from 8 to 16 km). 2) The recording of photons at an altitude of 20 km, which
570 corresponds to the altitude of the aircraft. 3) These point source simulations are extended by
571 random sampling to generate any extended source.

572 **a) Attenuation with radial distance**

573
574 Here we estimate how far away from the aircraft nadir can gamma-ray glows be detected by
575 ALOFT. We can safely identify dim glows which exhibit a minimum count rate increase of
576 25% above the background level (using one-second time bin). Conversely, the brightest
577 glows exhibit a count rate up to ~12 times the background (using one-second time bin). We
578 therefore detect glows with maximum intensity spanning a dynamic range of ~50. This
579 observed dynamic range includes both contributions from the intrinsic glow intensity

variability, and absorption due to the source location with respect to the aircraft. Extended Data 7 shows the expected number of photons per unit area at 20 km altitude for sources at different altitudes and radial distance from the aircraft, normalized to 1 at radial distance 0. The dashed horizontal line marks the detection threshold (0.02, following from a dynamic range of 50). Assuming that the brightest glows are detected from sources close to the nadir, we can see that a source as bright would be detectable at 25% above the background level up to 2 – ~4.5 km radial distance, depending on the source altitude. Considering that glows are detected while passing over convective cores extending at high altitudes, it is likely that the sources have a high altitude as well (see discussion of the cloud charge structure in Methods 3b) therefore the detection range is likely lower than ~3.5 km.

b) Time profile due to aircraft motion

Here we estimate the observed time profile of a stationary gamma-ray glow due only to the aircraft motion. Extended Data 8 shows the relative number of counts at aircraft altitude assuming the aircraft is flying towards an extended source (7.5 km radius disk) at 16 km altitude for different radial offsets. The rise time from 2% (detection threshold assuming the dynamic range 50 discussed in Method 4a) to 95% of the maximum is >18 seconds. This is way longer than the typically observed rise times (2-4 s) suggesting that the intrinsic time variability of the source dominates the observed time profile. This rise time further increases with increasing radial offset from the source. The relative intensity sharply decreases when the radial offset exceeds the physical extension of the glow region, in agreement with the attenuation with radial distance discussed above in Method 4a.

c) Glow spectral analysis

Here we show that the energy spectra of a sample of the detected gamma-ray glows can be well modeled by the RREA process. Extended Data 9a presents energy spectra and lightcurves for four cases detected during the flight on 06.07.2023. It also includes the best fit results from a RREA model, which closely matches the observed glow spectra. The fitting procedure involved the following steps:

1. Start with an upward electron source at 15 km altitude, with a typical RREA electron energy spectrum.
2. Propagating these electrons through the atmosphere and recording bremsstrahlung photons at 20 km altitude (only a negligible fraction of electrons reach this altitude).
3. Applying the response matrix of the BGO instrument to simulate the spectrum and subsequently comparing it with the observed data.

For the best fit, we utilized an averaged spectrum derived from all the observations shown. A key parameter optimized in this modeling was the extent of the RREA region. The best fit was achieved with a source region larger than 8 km spatial extent, determined using a maximum likelihood approach. Although the approach is potentially mixing sources with different characteristics, it shows that the RREA process is a good explanation for the observed gamma-ray glow spectra.

With the same procedure we also performed spectral analysis of the glow burst detected on 24.07.2024 at 06:44:59.182 UTC and presented in Figure 2f. Extended Data 9b presents the energy spectrum for this event. The spectrum can be well fit with a RREA model and does not show any evidence of enhanced features at the 511 keV positron annihilation line. Therefore, we conclude that this event is intrinsically different from the positron bursts presented in²³.

629

630

631 **REFERENCES**

632

- 633 29. Hansen, R.S., et al., *How simulated fluence of photons from terrestrial gamma ray flashes at*
634 *aircraft and balloon altitudes depends on initial parameters.* Journal of Geophysical
635 Research: Space Physics, **118**(5): p. 2333-2339 (2013).
- 636 30. Østgaard, N., et al., *The Modular X- and Gamma-Ray Sensor (MXGS) of the ASIM Payload*
637 *on the International Space Station.* Space Science Reviews, **215**(2) (2019).
- 638 31. Amiot, C.G., et al., *Dual-Polarization Deconvolution and Geophysical Retrievals from the*
639 *Advanced Microwave Precipitation Radiometer during OLYMPEX/RADEX.* Journal of
640 Atmospheric and Oceanic Technology, **38**(3): p. 607-628 (2021).
- 641 32. Leppert, K.D. and D.J. Cecil, *Signatures of Hydrometeor Species from Airborne Passive*
642 *Microwave Data for Frequencies 10–183 GHz.* Journal of Applied Meteorology and
643 Climatology, **54**(6): p. 1313-1334 (2015).
- 644 33. Battaglia, A., et al., *Using a multiwavelength suite of microwave instruments to investigate the*
645 *microphysical structure of deep convective cores.* Journal of Geophysical Research:
646 Atmospheres, **121**(16): p. 9356-9381 (2016).
- 647 34. Heymsfield, G.M.L.L.M.W.M.C.H.S.G., *NASA Airborne Radars for Studying Clouds and*
648 *Precipitation.* Advances in Weather Radar, **1**: p. 231-282 (2024).
- 649 35. Bateman, M.G., et al., *A Low-Noise, Microprocessor-Controlled, Internally Digitizing*
650 *Rotating-Vane Electric Field Mill for Airborne Platforms.* Journal of Atmospheric and
651 Oceanic Technology, **24**(7): p. 1245-1255 (2007).
- 652 36. Mach, D.M., et al., *Electric fields, conductivity, and estimated currents from aircraft*
653 *overflights of electrified clouds.* Journal of Geophysical Research: Atmospheres, **114**(D10)
654 (2009).
- 655 37. Mach, D.M., et al., *Comparisons of total currents based on storm location, polarity, and flash*
656 *rates derived from high-altitude aircraft overflights.* Journal of Geophysical Research:
657 Atmospheres, **115**(D3) (2010).
- 658 38. Schultz, C.J., et al., *Remote Sensing of Electric Fields Observed Within Winter Precipitation*
659 *During the 2020 Investigation of Microphysics and Precipitation for Atlantic Coast-*
660 *Threatening Snowstorms (IMPACTS) Field Campaign.* Journal of Geophysical Research:
661 Atmospheres, **126**(16) (2021).
- 662 39. Mach, D.M. and W.J. Koshak, *General Matrix Inversion Technique for the Calibration of*
663 *Electric Field Sensor Arrays on Aircraft Platforms.* Journal of Atmospheric and Oceanic
664 Technology, **24**(9): p. 1576-1587 (2007).
- 665 40. GOES R SERIES PRODUCT DEFINITION AND USERS' GUIDE (PUG). 2019. p. 27-28.
- 666 41. Thomas, R.J., et al., *Observations of VHF source powers radiated by lightning.* Geophysical
667 Research Letters, **28**(1): p. 143-146 (2001).
- 668 42. Agostinelli, S., et al., *Geant4—a simulation toolkit.* Nuclear Instruments and Methods in
669 Physics Research Section A: Accelerators, Spectrometers, Detectors and Associated
670 Equipment, **506**(3): p. 250-303 (2003).
- 671 43. Dwyer, J.R., *A fundamental limit on electric fields in air.* Geophysical Research Letters,
672 **30**(20) (2003).

673

674

675

676

677 **DATA AVAILABILITY**

678

679 All the data used in this study are available on ZENODO Repository, doi:
680 10.5281/zenodo.12531291. The description of the formats is uploaded in the file:
681 Data_description.pdf

682

683 **CODE AVAILABILITY**

684

685 References to software tools and codes for this study are given in the file:
686 Data_description.pdf, which is uploaded to the ZENODO Repository, doi:
687 10.5281/zenodo.12531291.

688

689 EXTENDED DATA FIGURES AND TABLES
690
691

692 **Extended Data 1.**

693 **Time-based color-coded map and gamma-ray sequence for the flight on 06/07/2023
694 (Campeche Bay): 06.07.2023 04:25 UT – 07:05 UT.**

695 a) Flight trajectory (grey) overlaid on GOES-18 infra-red temperatures, along with marked
696 gamma-ray fluxes exceeding 25% above the background level. The brightness temperature
697 cloud image corresponds to time 05:00:30 UT. b) BGO count rate. Gamma-ray fluxes
698 exceeding 25% above the background level are color-coded based on time from 04:25:00 UT,
699 correspondingly in panels a and b.

700
701 **Extended Data 2**

702 **Gamma-ray count rate with simultaneous electric field and cloud characterization data
703 for flight on 06.07.2023, 04:57 – 05:20 UT.**

704 a) BGO count rate (black, left vertical axis) and EFCM trigger rate (red, right vertical axis). b)
705 Brightness Temperature (TB) from AMPR 85.5 GHz Channel A. Sweep Index 0 (49)
706 corresponds to scan direction 45 degrees from Nadir towards starboard (port). C) vertical radar
707 reflectivity profile from EXRAD. Black vectors are the E_{xz} component of the electric field from
708 LIP, where x-axis is the direction of motion and z-axis is the upward-pointing vertical direction.
709 AMPR, EXRAD and LIP data are not shown if roll angle is larger than 2 degrees (aircraft
710 maneuvering).

711
712 **Extended Data 3**

713 **Time-based color-coded map and gamma-ray sequence for the flight on 29/07/2023
714 (Florida, east coast): 29.07.2023 19:55 UT – 21:05 UT.**

715 a) Flight trajectory (grey) overlaid on GOES-16 infra-red temperatures, along with marked
716 gamma-ray fluxes exceeding 25% above the background level. The brightness temperature
717 cloud image corresponds to time 20:31:17 UT. b) BGO count rate. Gamma-ray fluxes
718 exceeding 25% above the background level are color-coded based on time from 19:55:00 UT,
719 correspondingly in panels a and b.

720
721 **Extended Data 4**

722 **Gamma-ray count rate with simultaneous electric field and cloud characterization data
723 for flight on 29.07.2023, 20:28 – 20:50 UT.**

724 a) BGO count rate (black, left vertical axis) and EFCM trigger rate (red, right vertical axis).
725 EFCM data were lost between 20:45:44 and 20:56:00 due to a system failure. b) Brightness
726 Temperature (TB) from AMPR 85.5 GHz Channel A. Sweep Index 0 (49) corresponds to scan
727 direction 45 degrees from Nadir towards starboard (port). c) vertical radar reflectivity profile
728 from EXRAD. Black vectors are the E_{xz} component of the electric field from LIP, where x-axis
729 is the direction of motion and z-axis is the upward-pointing vertical direction. AMPR, EXRAD
730 and LIP data are not shown if roll angle is larger than 2 degrees (aircraft maneuvering).

731
732
733
734
735
736
737
738

739 **Extended Data 5**

740 **LMA observations for the energetic two-pulse glow event at 20:30 UT on 29.07.2023.**

741 a) Temporal comparison of the 1-second averaged BGO count rate (blue line and red + signs)
742 with the occurrence of lightning in the storm (square dots, colored by time), c) Plan view of
743 the LMA sources and ER-2 trajectory (blue diagonal line from northeast to southwest),
744 showing the location of the ER-2 at the time of the glows (horizontal arrows) relative to the
745 LMA sources (black lines show coastal and inland water boundaries around Kennedy Space
746 Center, green squares show locations of two LMA stations), b,d) North-south and east-west
747 vertical projections of the LMA sources and the spatial extent of the glows and the height of
748 the ER-2 relative to the lightning, e) histogram of the source altitudes.

749 **Extended Data 6**

750 **Lightning-inferred charge structure around the time of the 20:30 glows.**

751 Same as Extended Data 5, except showing the typical tri-polar structure (main mid-level
752 negative (blue) and upper positive charges (red-orange), and weaker lower positive charge)
753 over a five minute interval around the time of the glow sequence.

754 **Extended Data 7**

755 **Gamma-ray glow detectability.**

756 Expected number of photons per unit area at 20 km altitude (Arbitrary Units, normalized to 1
757 at radial distance 0) for sources at different altitudes and radial distance from the aircraft.
758 Assuming a maximum intensity of ~12 times the background for glows detected at radial
759 distance 0 (the maximum observed intensity) the dashed horizontal line marks the detection
760 threshold for gamma-ray glows 25% above the background.

761 **Extended Data 8**

762 **Count rate variability due to aircraft motion.**

763 Relative number of counts at aircraft altitude assuming the aircraft flying toward an extended
764 source (7.5 km radius disk) at 16 km altitude for different radial offsets.

765 **Extended Data 9**

766 **Gamma-ray glow and glow burst spectra.**

767 a) Background-subtracted differential energy spectra for four time intervals including several
768 glows detected on 06.07.2023. Error bars are one standard deviation. The insets show the
769 BGO count rates. The corresponding time intervals (start time – stop time, UTC) are: glow 4
770 05:00:49 – 05:03:07, glow 5 05:07:09 – 05:09:20, glow 7 05:23:35 – 05:24:13, glow 10
771 06:51:13 – 06:52:33. b) Background-subtracted differential energy spectra for the glow burst
772 detected on 24.07.2024 at 06:44:59.182 UTC. Error bars are three standard deviations.

773 **Extended Data 10.**

774 **ALOFT flights summary table.**





