

Comets: Formation and early evolution

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Comet formation

- *For sure:*
- *Time & place:* the solar nebula, beyond the snow line!
- *Identity:* icy planetesimals!
- *Under debate:*
- *Mechanism:* pebble swarm or slow accretion?
- *Material:* pre-solar grains or condensates?

Properties: **Porosity**

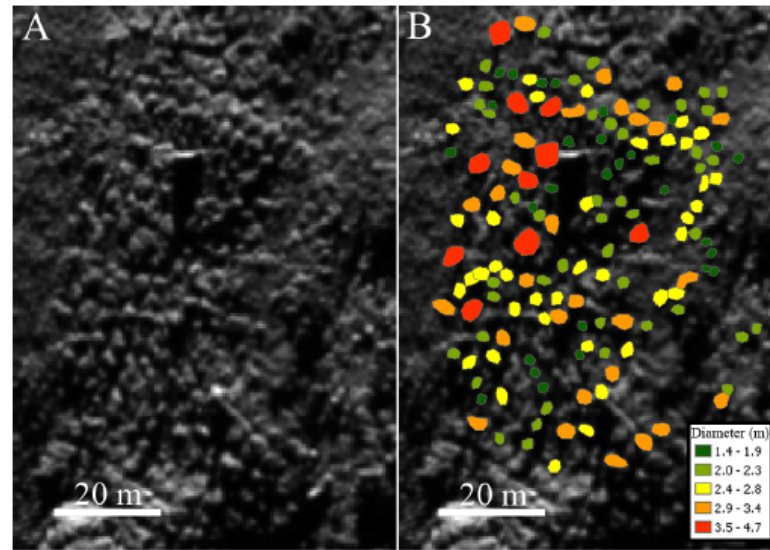
- ***67P/Churyumov-Gerasimenko (Rosetta)***
- ***Density:*** 0.533 ± 0.006 g/cc (RSI+OSIRIS)
- ***Dust/Gas Ratio:*** 4 ± 2 (Giada+OSIRIS)
This is very large; even 6 has been suggested
- ***Porosity = density / compact density*** \Rightarrow
72–74% (Pätzold et al. 2016)

This is mainly micro-porosity (no large voids)

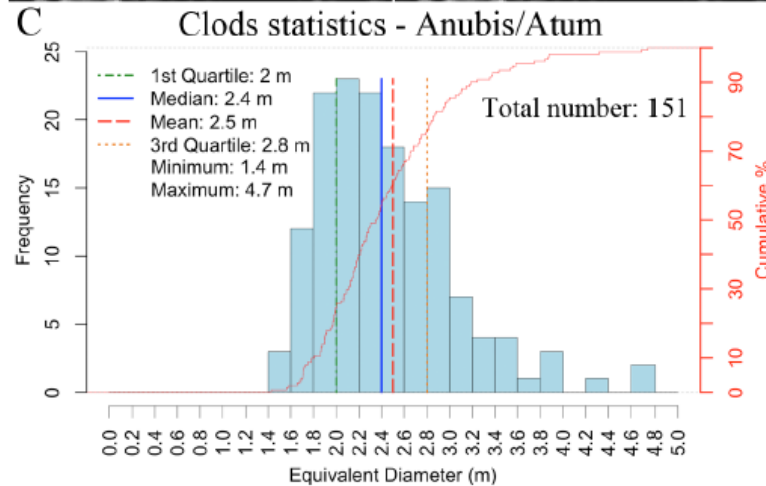
Properties: Being pristine

- ***Risks of modification:*** *Radiogenic heating* (^{26}Al , ^{40}K); *Collisions at high speed*
- ^{26}Al heating may purge km-sized comets of super-volatiles, contrary to observations
- But this is *strongly dependent on the time of formation!*
- *Low-velocity accretion* is required to avoid compaction yielding too high density

Davidsson et al. (2016)



“goosebumps”

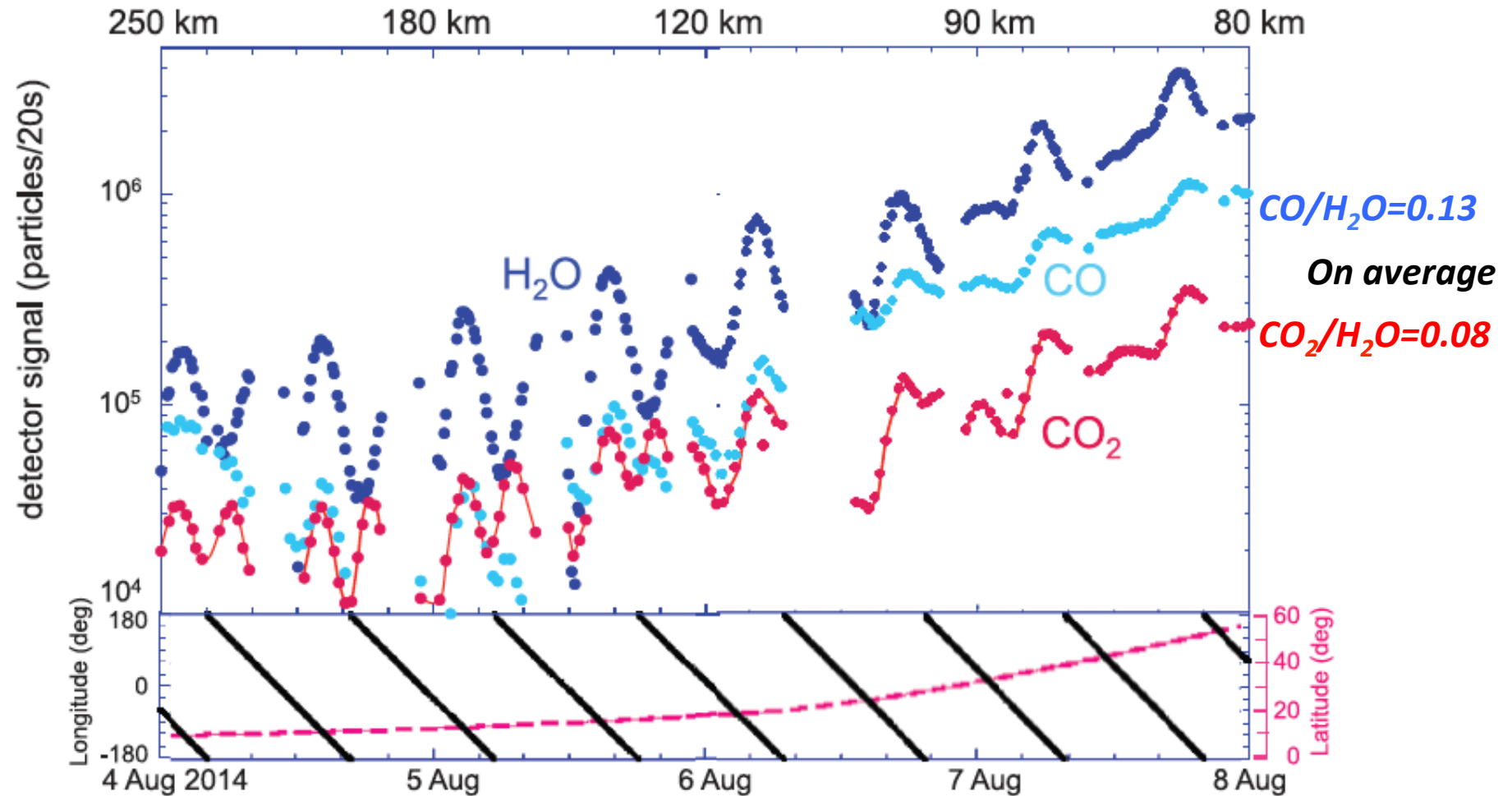


Possibly, original structural units: $St \approx 1$

Properties: **Composition**

- ***67P Rosetta measurements***
- *Main contributors so far:*
- *Volatiles:* ROSINA orbiting mass spectrometer
- *Refractories:* VIRTIS near-IR spectrometry of the nucleus surface (mainly organics)
- *Bulk:* DGR from production rates of solids and gases measured by different instruments

Hässig et al. (2015) **ROSINA**



Both CO and CO_2 are major, oxidized volatiles

Minor gaseous species

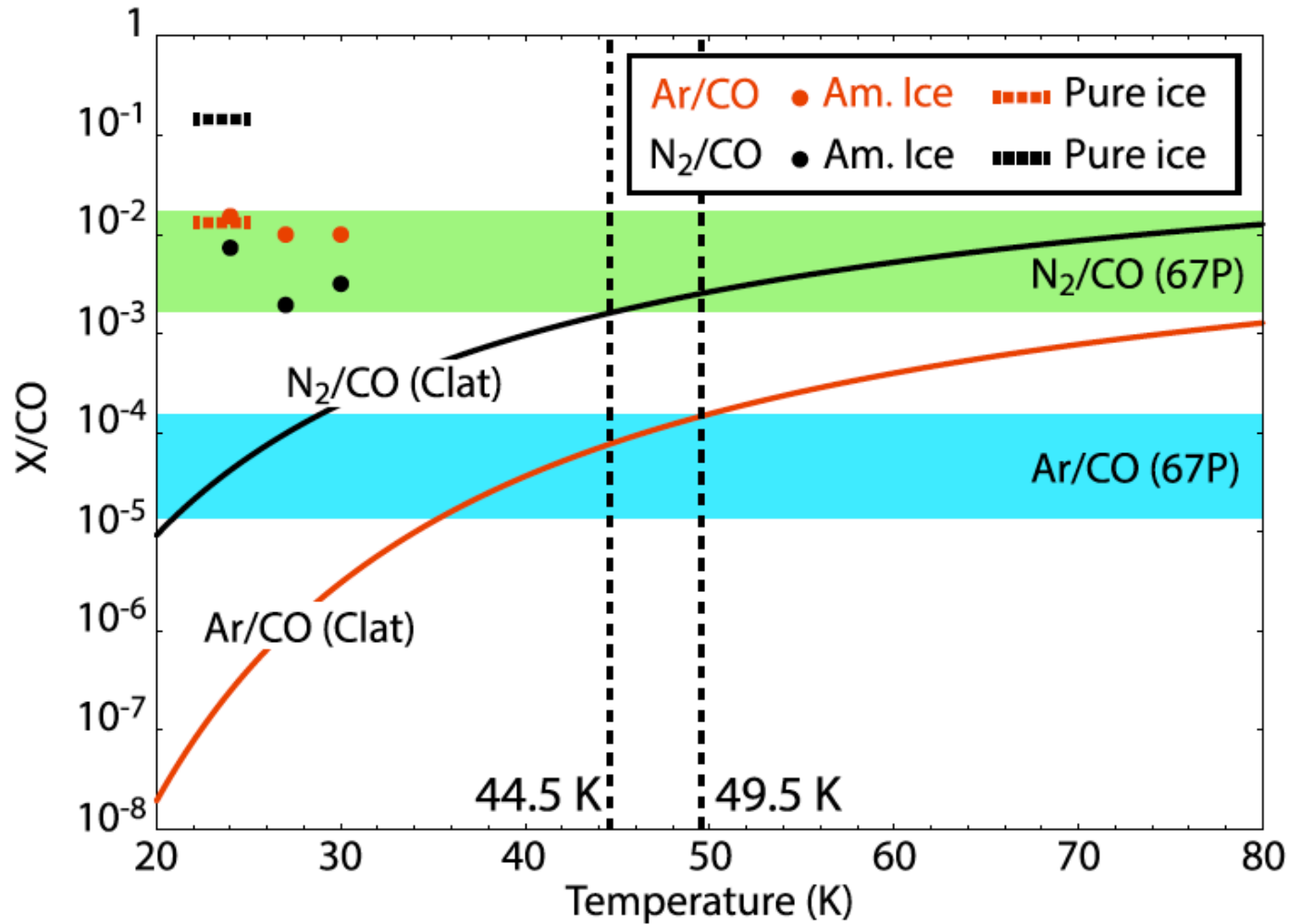
- ***Data from ROSINA in 67P:***

First detections!

- $Q(\text{O}_2)/Q(\text{H}_2\text{O}) = 3.8\%$ (Bieler et al. 2015)
- [In retrospect, $Q(\text{O}_2)/Q(\text{H}_2\text{O}) = 3.7\%$ in 1P/Halley]
- $Q(\text{N}_2)/Q(\text{CO}) \approx 1\%$ (Rubin et al. 2015)
- $Q(\text{Ar})/Q(\text{N}_2) \approx 1\%$ (Balsiger et al. 2015)

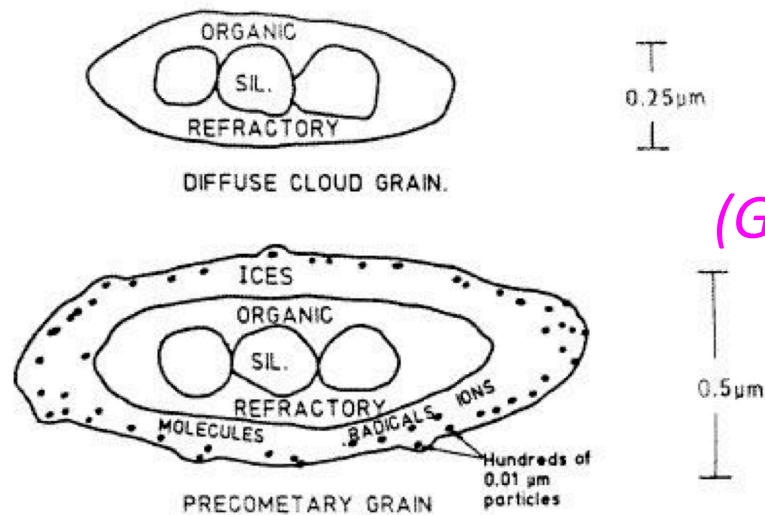
We cannot extrapolate to other comets!

Mousis et al. (2016)



Favored structure: crystalline clathrate

Cometary ice evolution



(Greenberg & Hage 1990)

Mousis et al.:

- (1) Pre-solar core-mantle grains with likely amorphous ice
- (2) Ice evaporation due to heating in the solar nebula
- (3) Ice re-condensation at low T in opaque mid-plane layer
- (4) Clathrate formation upon further cooling (?)

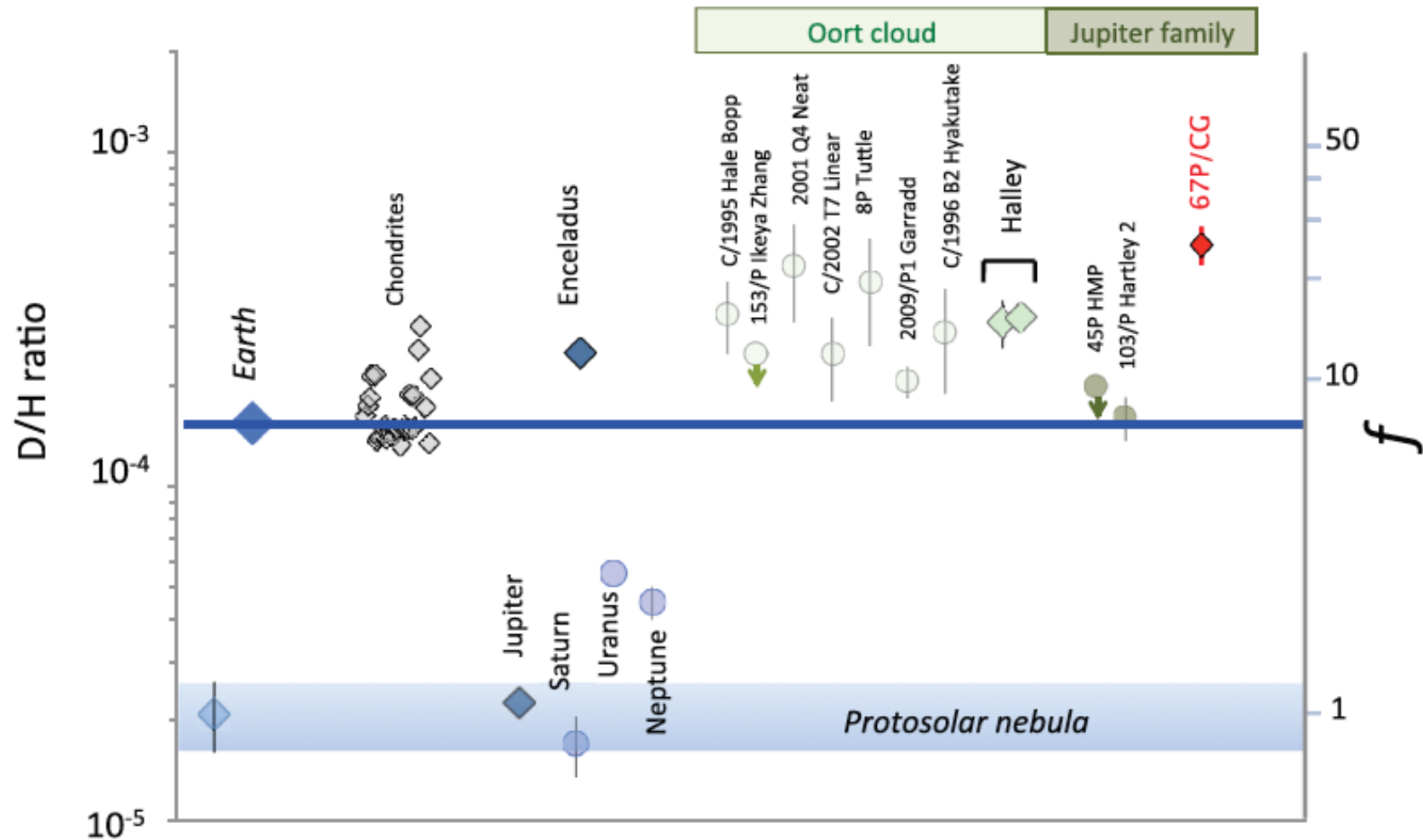
All this depends on heliocentric distance!

Properties: D/H ratio

- We express this in units of VSMOW (likely terrestrial value)
- **In 67P: $D/H \approx 3$** (ROSINA - Altwegg et al. 2015)
- **In 103P: $D/H \approx 1$** (Herschel – Hartogh et al. 2011)
- **In Halley, Hyakutake & Hale-Bopp: $D/H \approx 2$**
- Both *Jupiter Family* and *Long-period* comets show a spread, roughly from 1 to 3

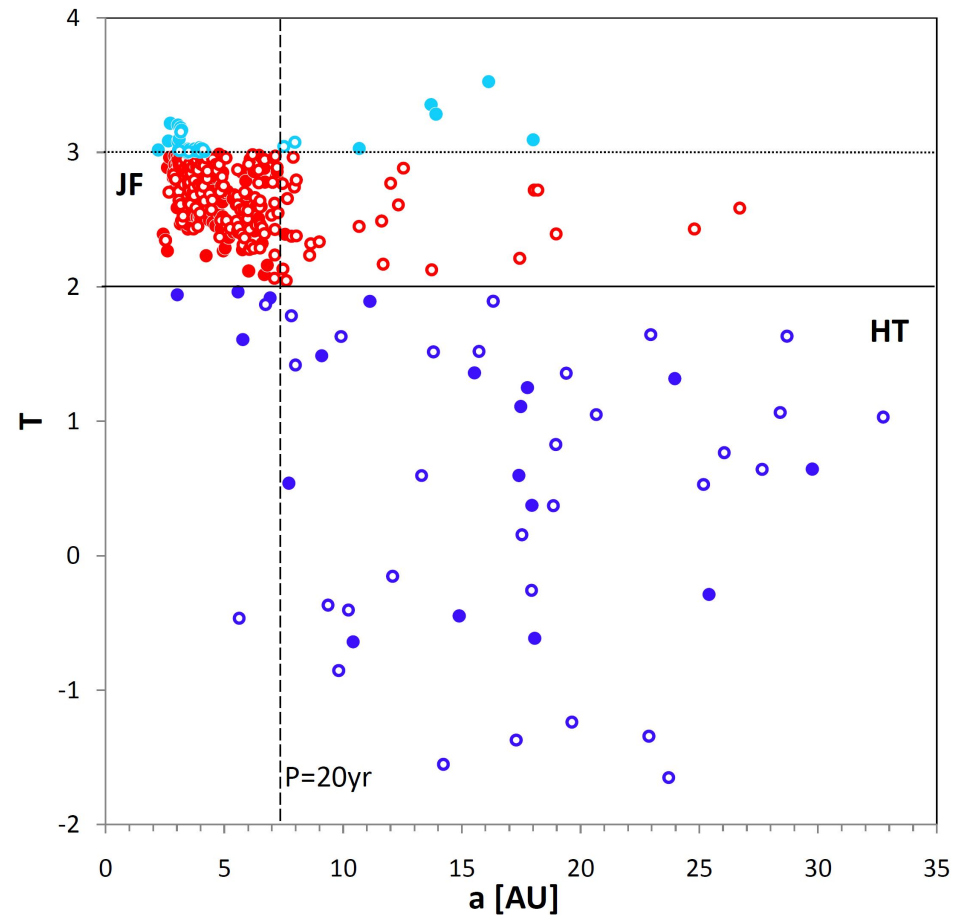
Comets of different types are similar

Altwegg et al. (2015)



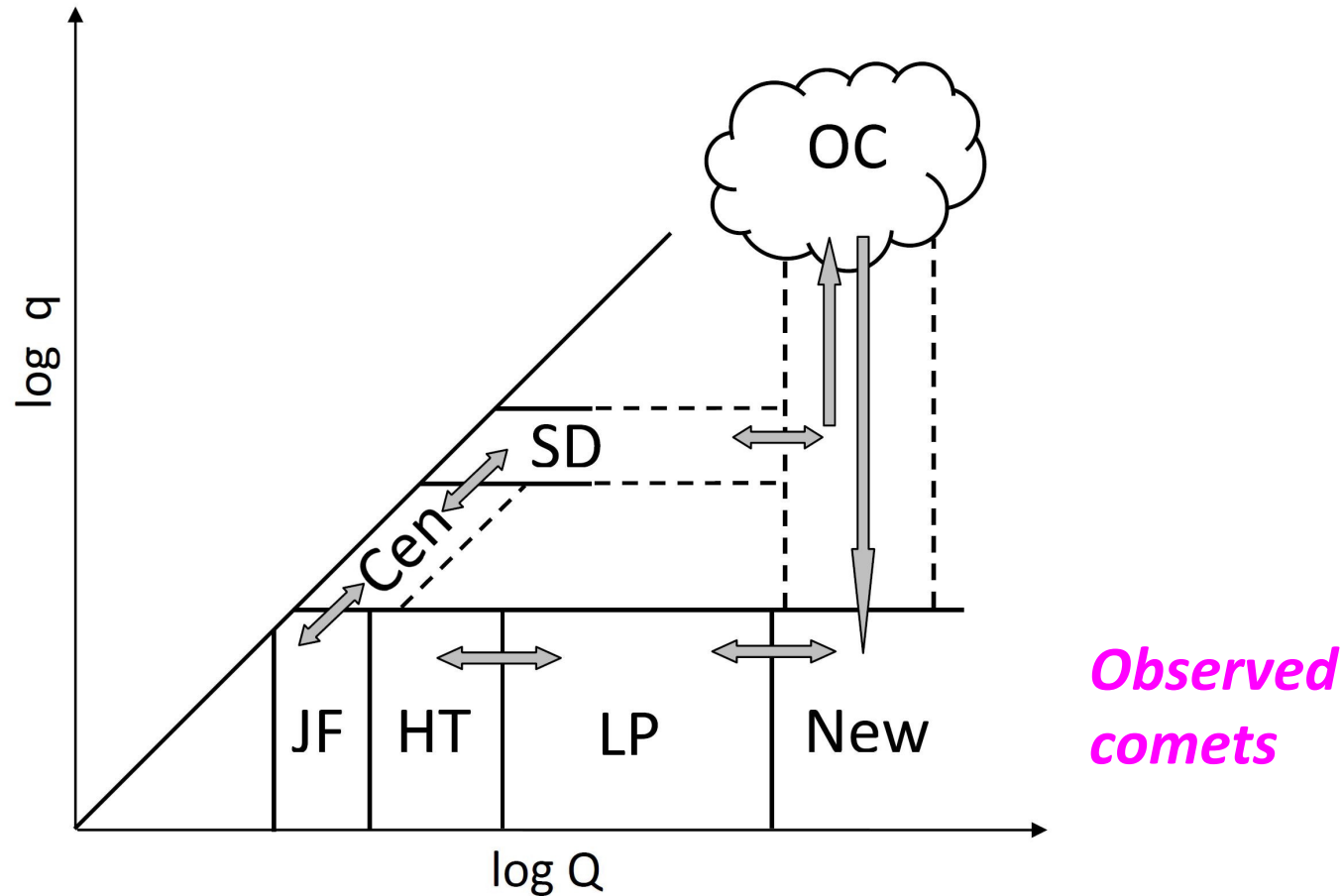
Earth's water was delivered as carbonaceous chondrites

Jupiter Family vs Halley Types



$$T = \frac{a_J}{a} + 2\sqrt{\frac{a}{a_J}(1 - e^2)} \cos i$$

Roadmap of comet dynamics

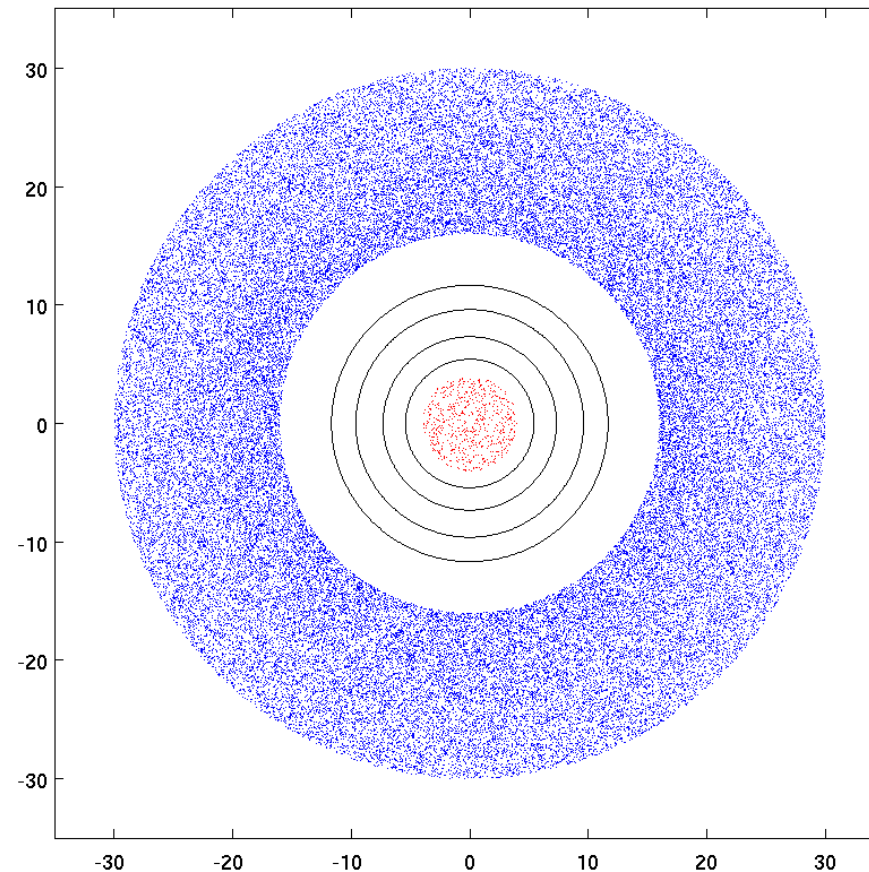


Source regions:

Jupiter Family: Scattered Disk

Halley Types: Oort Cloud

After dispersal of the solar nebula

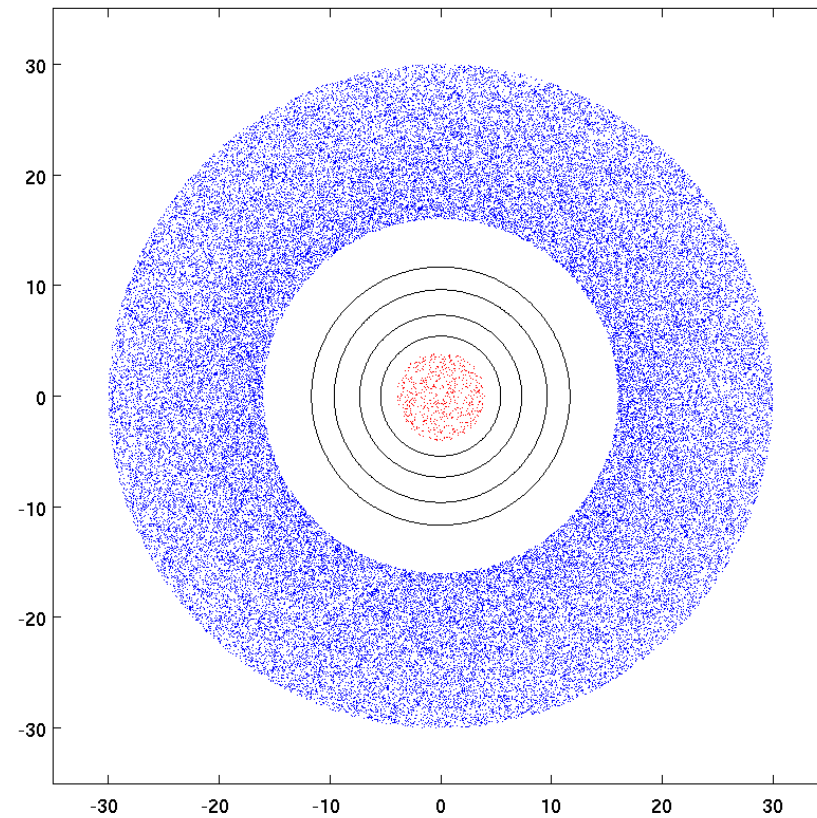


*Rocky
planetesimals*

*Icy
planetesimals*

Where did the GP stray planetesimals go?

How long did this last?



Did the Nice Model instability happen early or late?

Timing of the NM instability

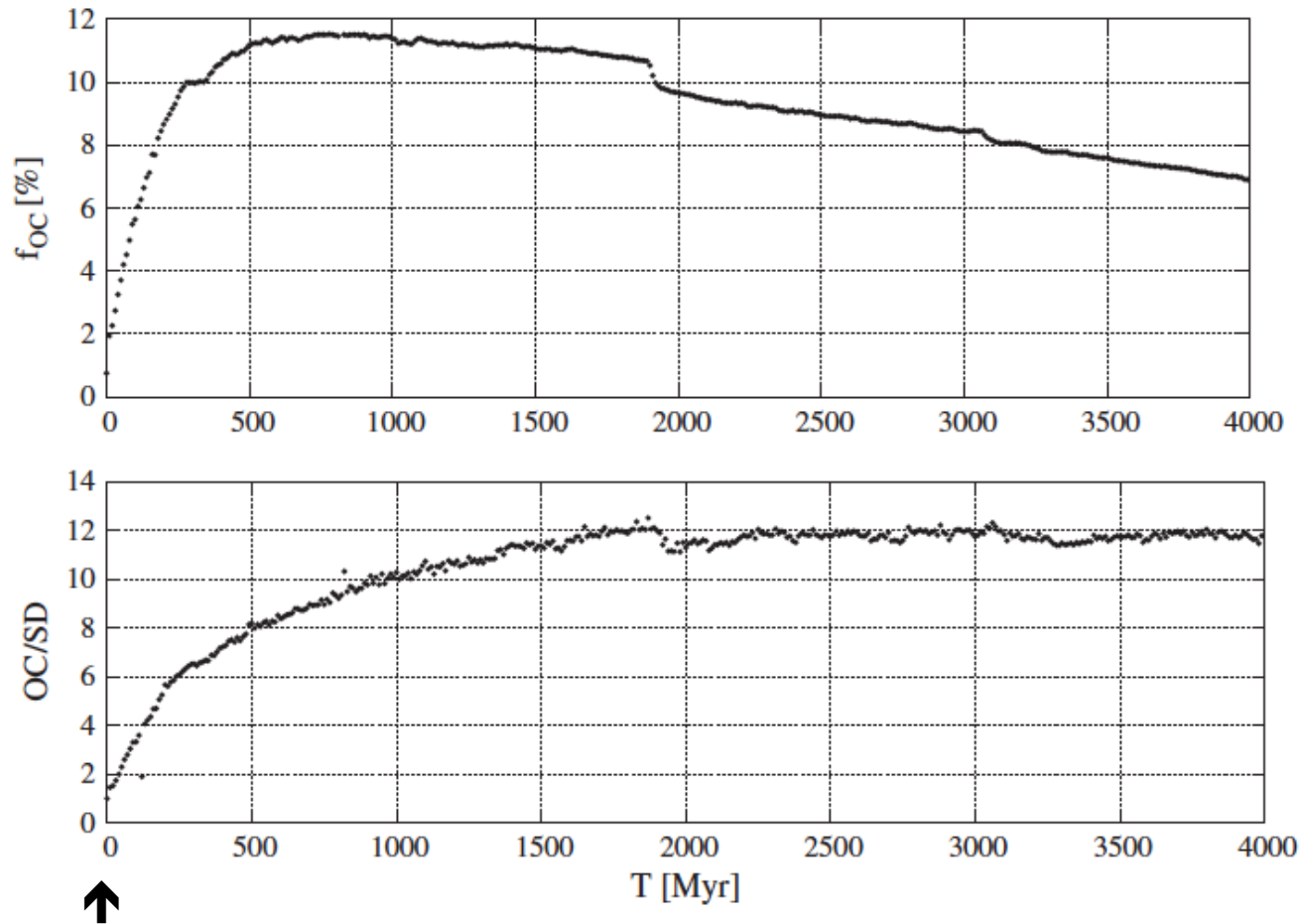
- Traditionally, associated with the start of the Late Heavy Bombardment 4.1-4.2 Gyr ago
- But the NM does not depend on this timing, and the existence of the LHB is not certain
- Fine tuning is needed to explain the TP orbits with late instability (Kaib & Chambers 2016)
- Jumping Jupiter scenario is anyway needed to explain MB asteroid orbits (Toliou et al 2016)

Very early or late instability: Open issue

Consequences of the instability

- ***Migration of ice giants*** through the disk of icy planetesimals \Rightarrow Dispersal by grav. scattering
- ***Outward scattering \Rightarrow Scattered Disk***
subject to energy diffusion
- ***External grav. torques \Rightarrow Decoupling*** from planetary influence; storage into high-q orbits
- Thus, ***formation of an Oort Cloud***

Brasser & Morbidelli (2013)



End of
migration

Currently, $M_{OC} \approx 7\%$ of $M(\text{disk})$
 M_{SD} is 12 times less

Nice Model inferences

- *All comets stem from the icy planetesimal disk*
- Gradients in chemical composition or D/H ratio will be seen as scatter of similar extent among JF and LP/HT (SD and OC) comets
- This seems to agree with observations

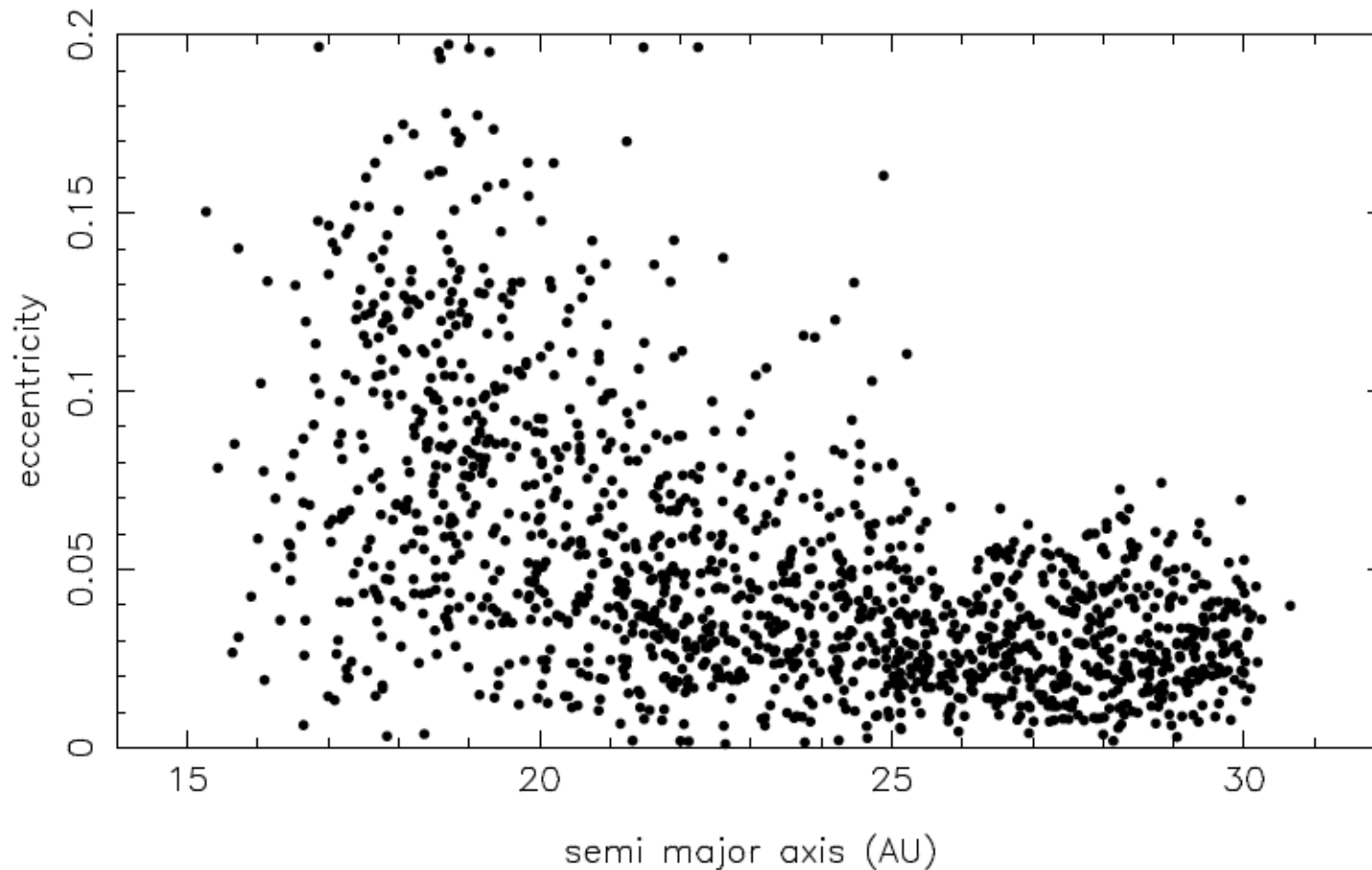
Collisional evolution of comets?!

- *Davis & Farinella (1996)*: Collisional evolution in the Edgeworth-Kuiper Belt \Rightarrow comet-size objects are ***collisional fragments***
- *Stern & Weissman (2001)*: During Oort Cloud formation out of planetesimals from the giant planet zone, the ***objects were destroyed***
- *Charnoz & Morbidelli (2003)*: OC comets may be primordial, while ***SD objects are collisional fragments***

What about the Nice Model?

- *Morbidelli & Rickman (2015)*
- We model the *collisional evolution of comets in a Nice Model scenario* from the beginning until the present: the *pre-instability disk* (400 My); the *disk dispersal stage* (SD formation, 350 My); the *SD residence time* (~4 Gy)

Excitation of the primordial disk



Disk state after 300 My according to Levison et al. (2011), self-excited by 1000 Pluto-size objects

Model parameters

- ***Disk population***
- SD formation $\Rightarrow 2 \times 10^{11}$ *objects with $D > 2.3$ km at the start of disk dispersal* (Brasser & Morbidelli 2013)
- We divide this disk into three radial zones
- ***Collisional break-up condition***
- Specific disruption energy $Q^*(R)$ *from Benz & Asphaug (1999)* for “strong ice”

Pre-instability disruptions

Number of catastrophic collisions per 400 My for a target with $R = 2$ km (as appropriate for 67P)

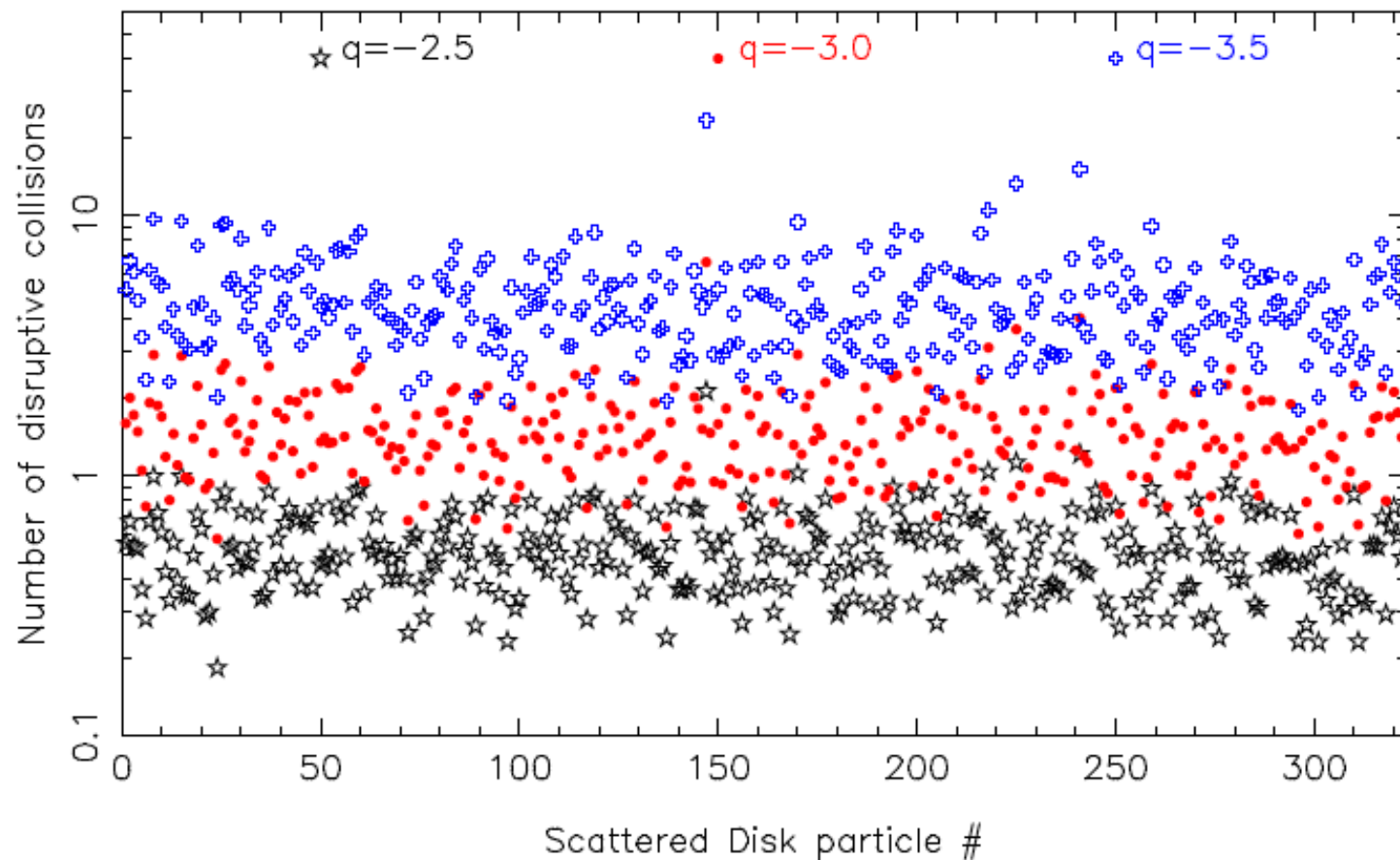
$q \backslash \text{targetzone}$	I	II	III
−2.5	58.0 (51.2)	28.7 (20.7)	12.3 (9.6)
−3.0	94.5 (75.0)	39.7 (23.7)	12.1 (7.9)
−3.5	190.6 (137.7)	70.2 (35.3)	15.4 (8.2)

q is the power law index of the assumed differential size frequency distribution

*Numbers in parentheses refer to the dynamical state after 100 My
(lower eccentricities and velocities)*

***TZ III may be favored for 67P due to its high D/H ratio.
However, the chance to avoid disruption is $< 10^{-4}$!***

Dispersal stage disruptions



For a steep SFD, comets do not survive

The standard case (-3.0) is borderline

With a shallow SFD, comets tend to survive to a large degree

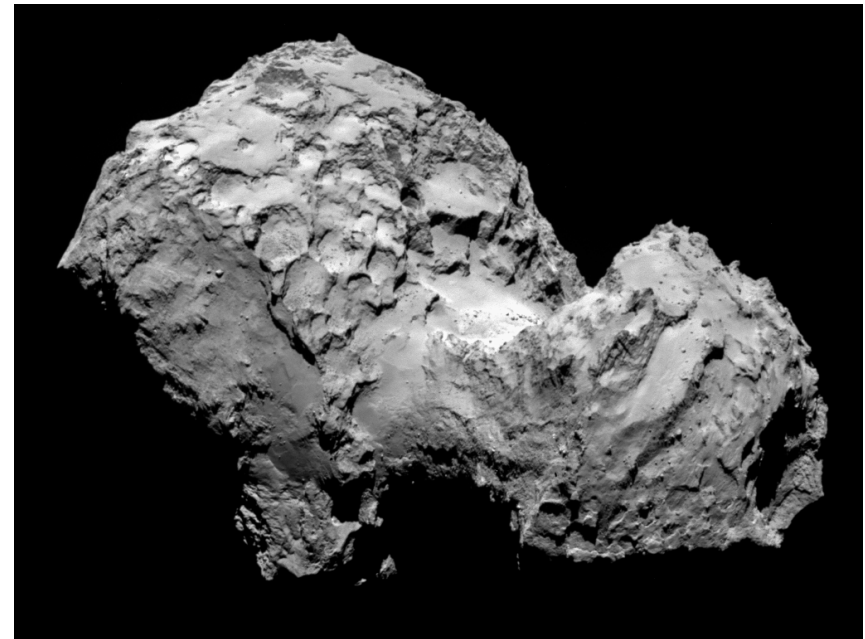
Collisional dilemma (?)

- A low-mass disk does not save comets from collisional destruction (Rickman et al. 2015)
- And **67P does appear to be primordial!** (highly porous, full of volatiles) ...

– *Could the disk be slim enough to save the comets?* **No**

– *Can comets stay primordial in spite of collisions?* **Yes**

Comet 67P (OSIRIS)



An early instability would help...

Evidence for late instability?

- *Marty et al.* (2016) find *Ar abundance in 67P* consistent with a late veneer for the Earth's atmospheric argon, implying a comet input of $3 \times 10^{18} - 6 \times 10^{20} \text{ kg}$. Rickman et al. (2017) find $3 \times 10^{19} \text{ kg}$.
- Abundant *water in lunar apatites* (*Greenwood et al. 2011*) may be difficult to explain, unless the mare forming projectiles were water rich.

There are still major, open issues!