

Grenoble passive seismic

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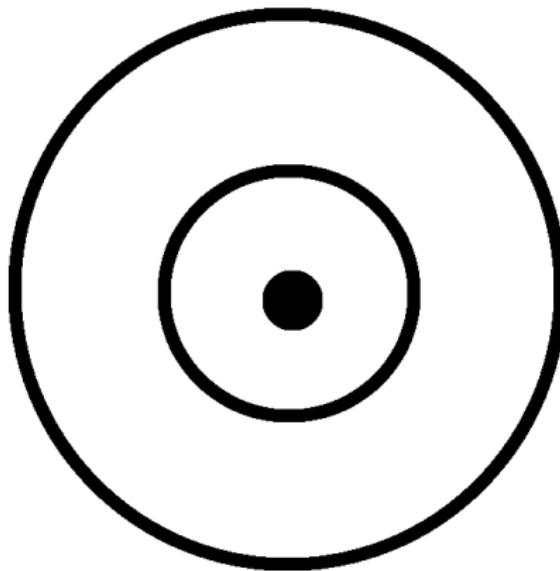
Spring 2019

Outline

- Data
- Methods
- Implementation
- Observations

Data

Circular arrays



three of these, center in all, 3 component data

Methods

Main idea

seismic event → group dispersion → integral → phase dispersion

Main idea

interferometry → group dispersion → integral → phase dispersion

Inspiration

Campillo et al, 1996

Shapiro et al, 1997

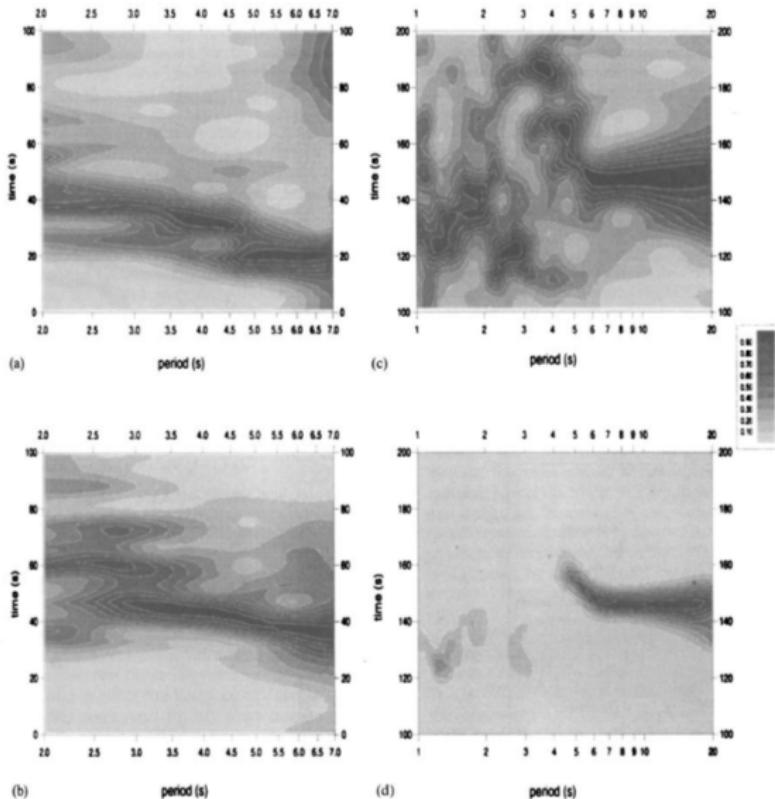
Brenguier et al, 2007

**Surface-wave tomography on valleys and volcanoes from
noise records and group velocities**

(MX volcanic belt & Piton de la Fournaise)

* they use all 3 component data

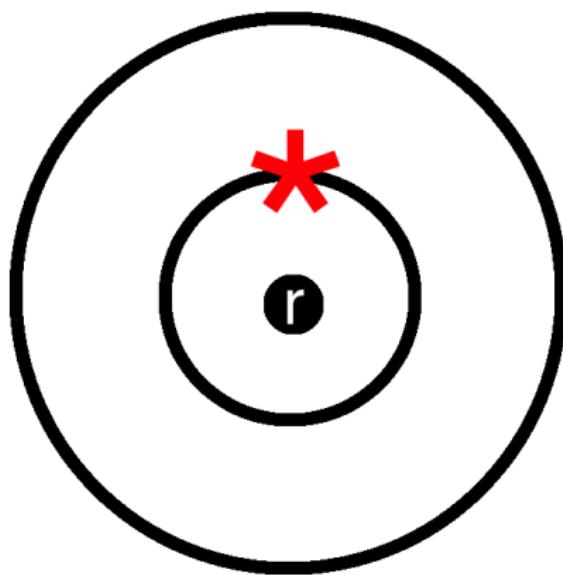
Inspiration



idea is so old, it's in black and white (Shapiro et. al, 1997)

Idea 1

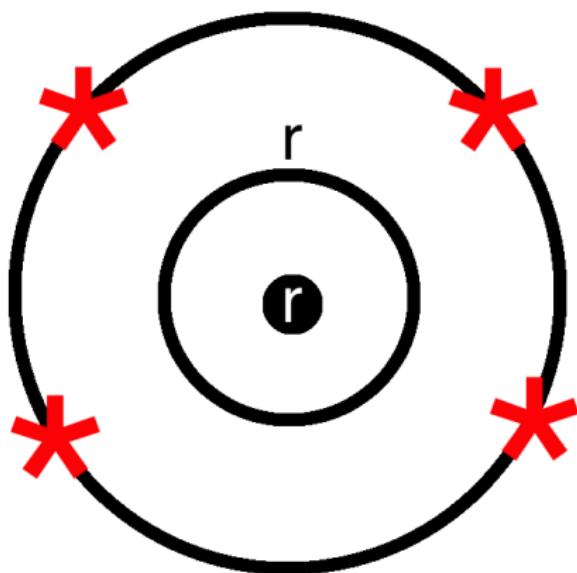
Interferometry



do many source-receiver pairs

Idea 2 - part 1

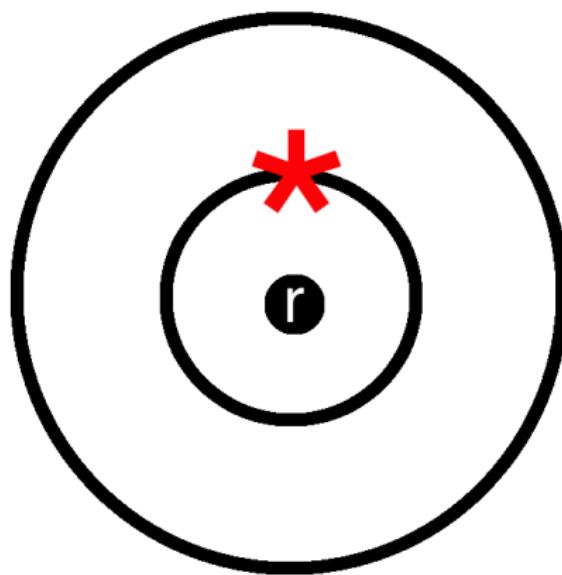
Interferometry - twice



many outer virtual sources

Idea 2 - part 2

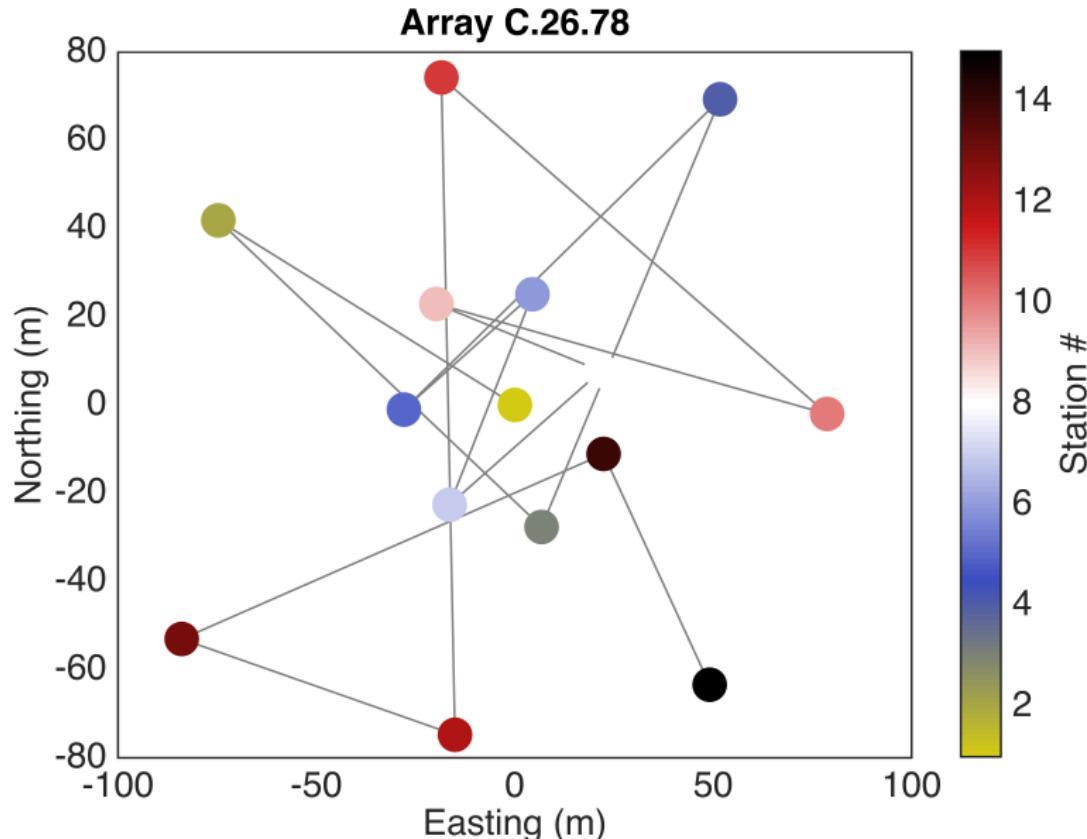
Interferometry - twice



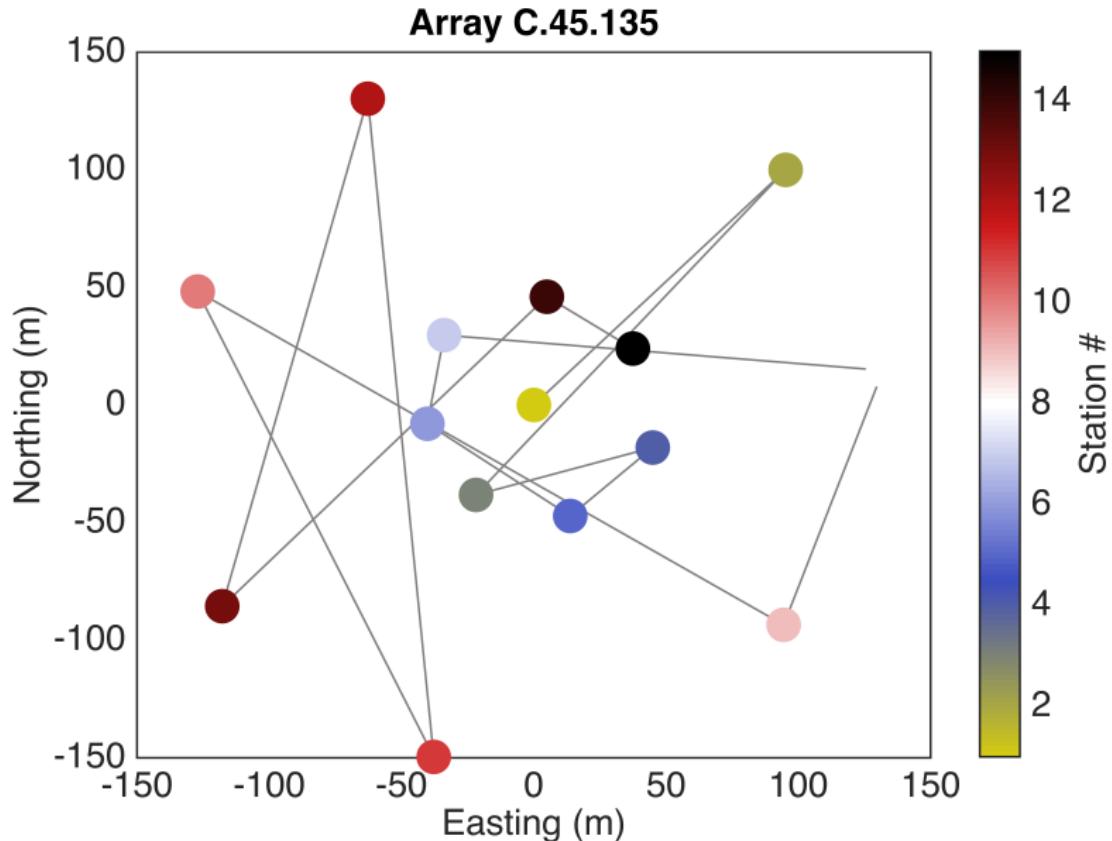
focus on one source-receiver, then do Idea 1

Implementation

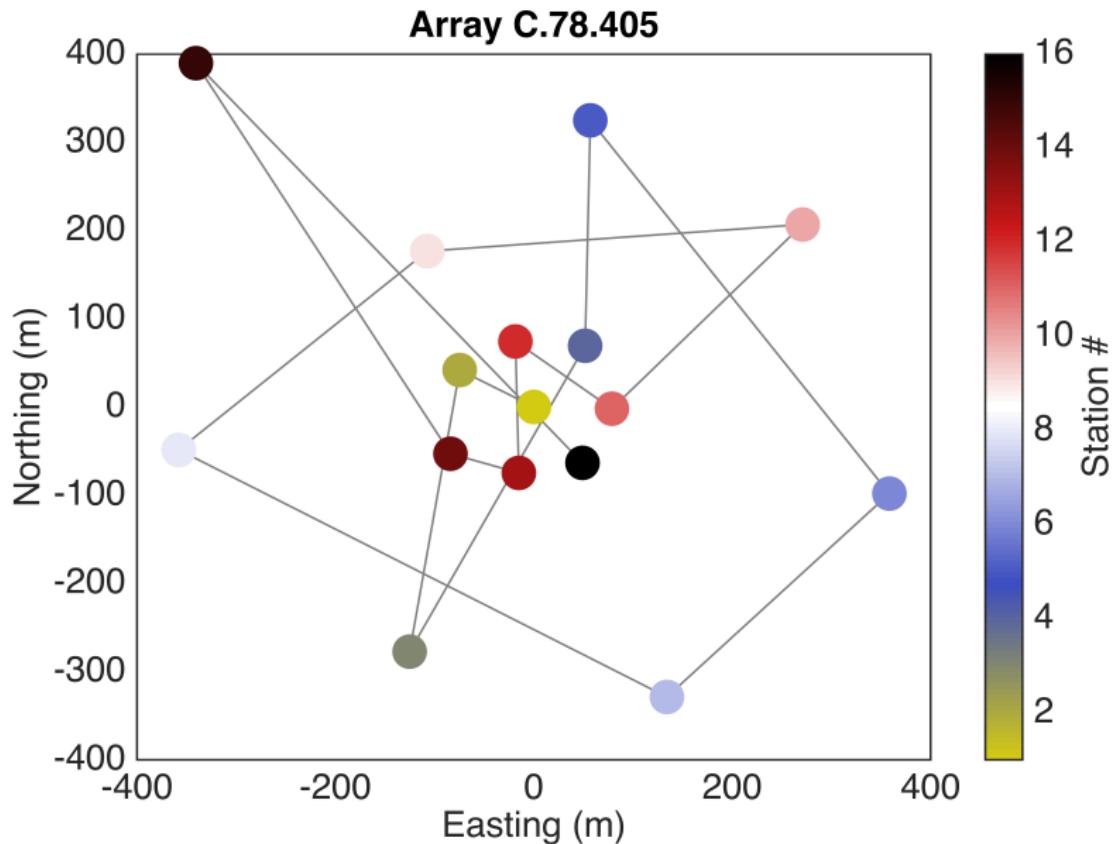
Array “small”



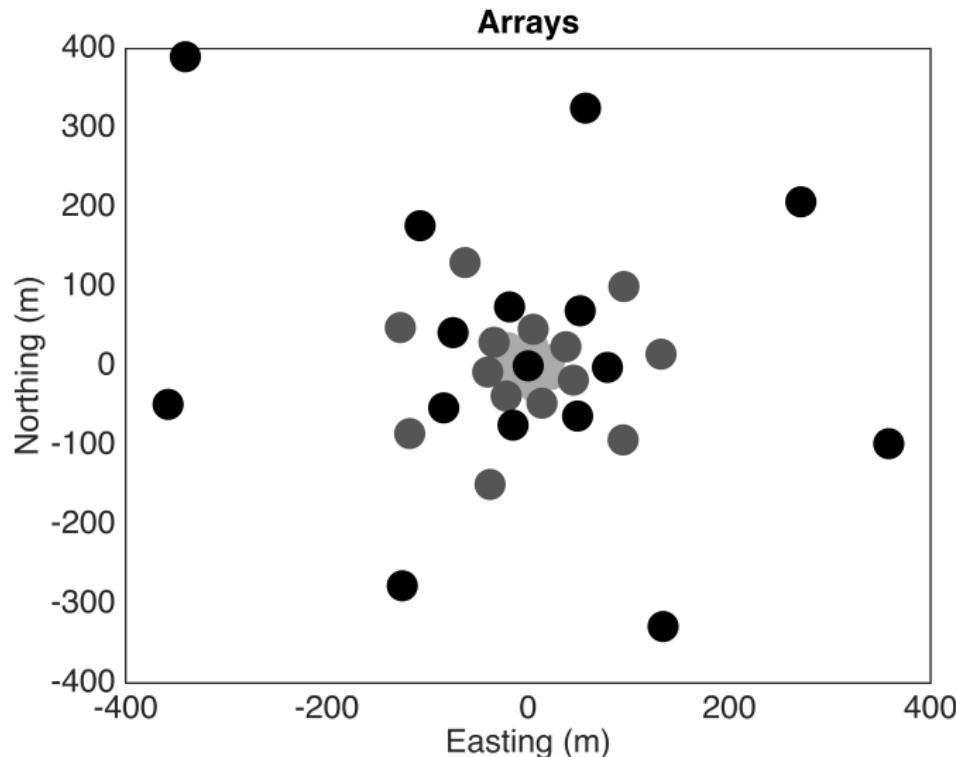
Array “medium”



Array “large”

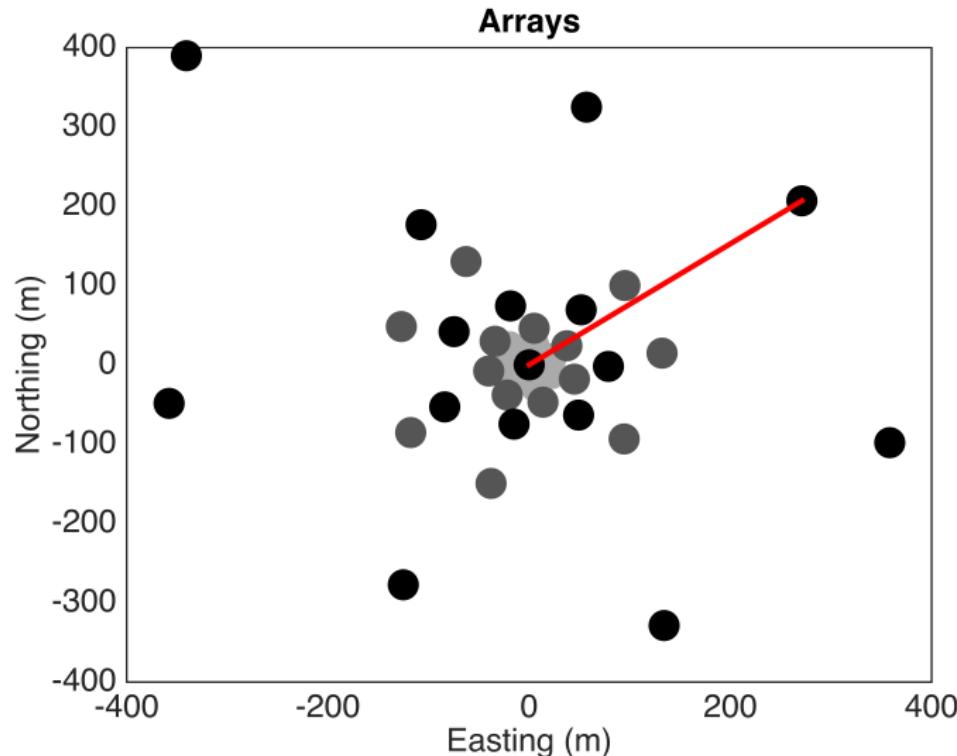


All arrays



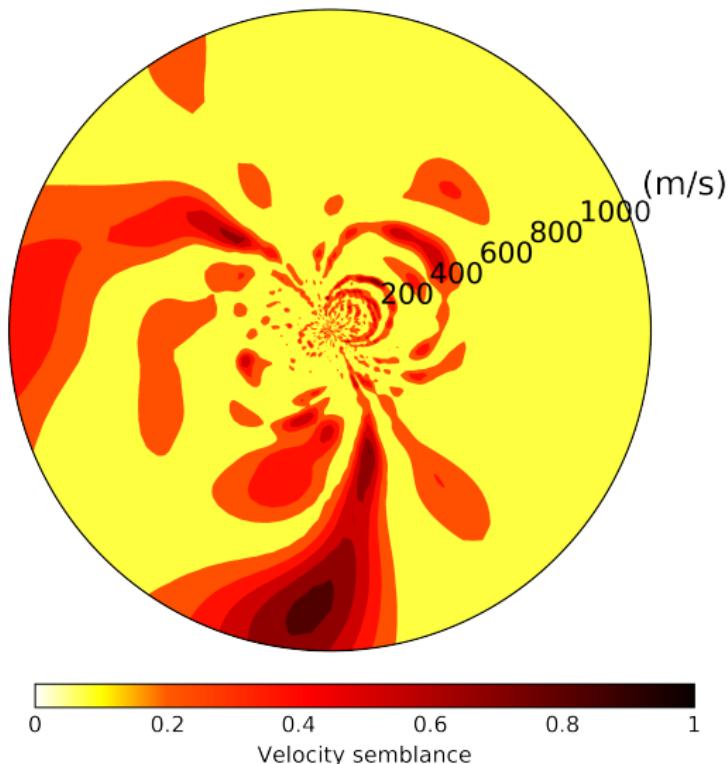
it's a white walker signal (spiral)

Arrays line



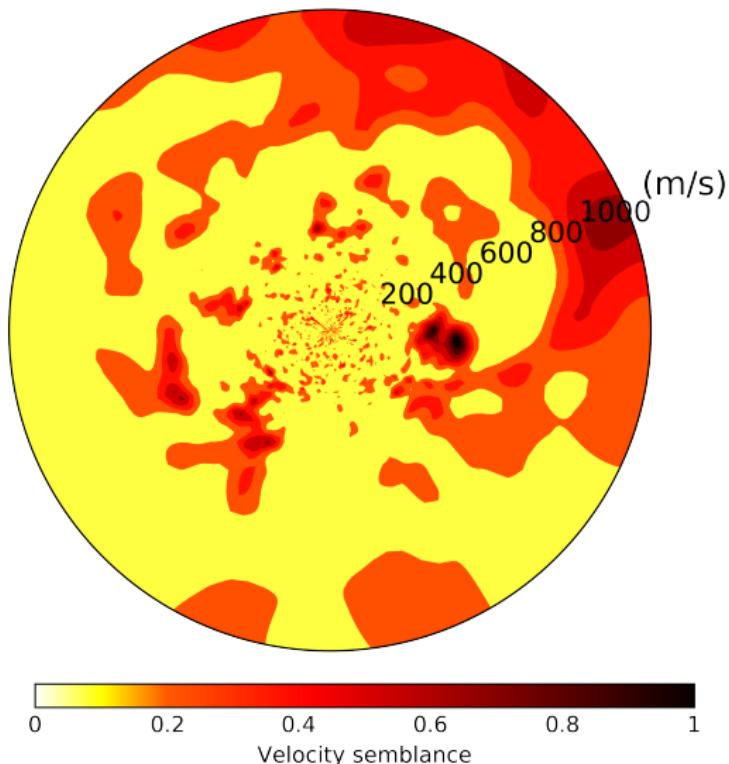
Noise sources

Beamformer C.26.78



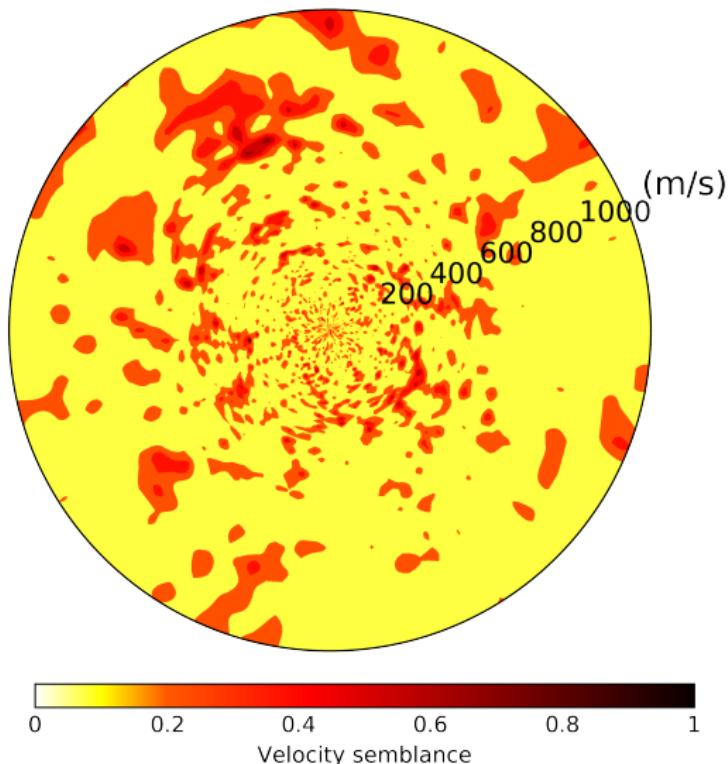
Noise sources

Beamformer C.45.135



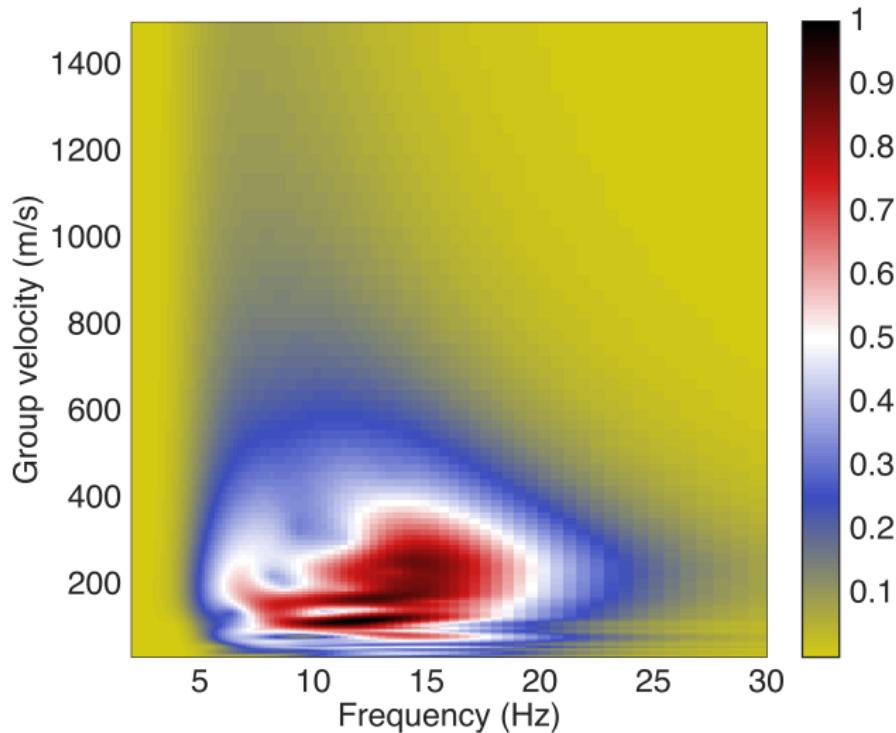
Noise sources

Beamformer C.78.405

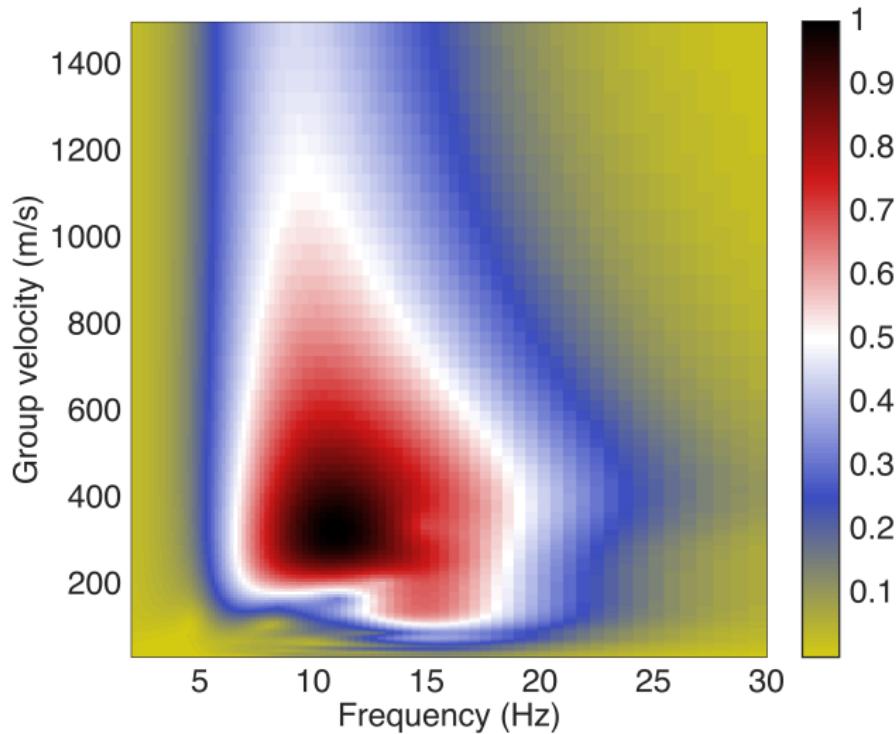


On y vamos!

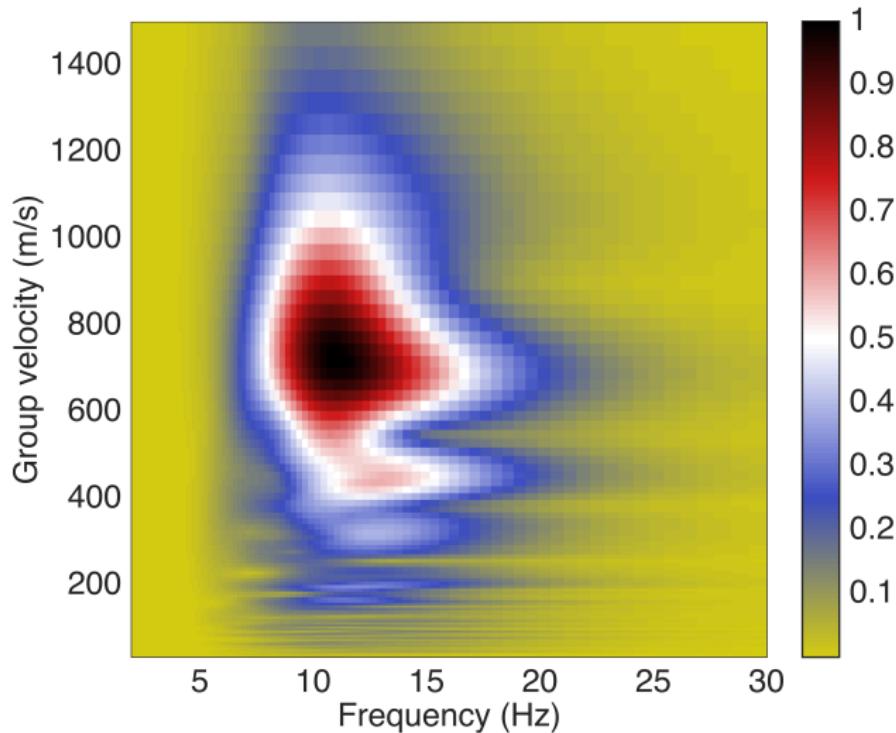
Group velocity “small”



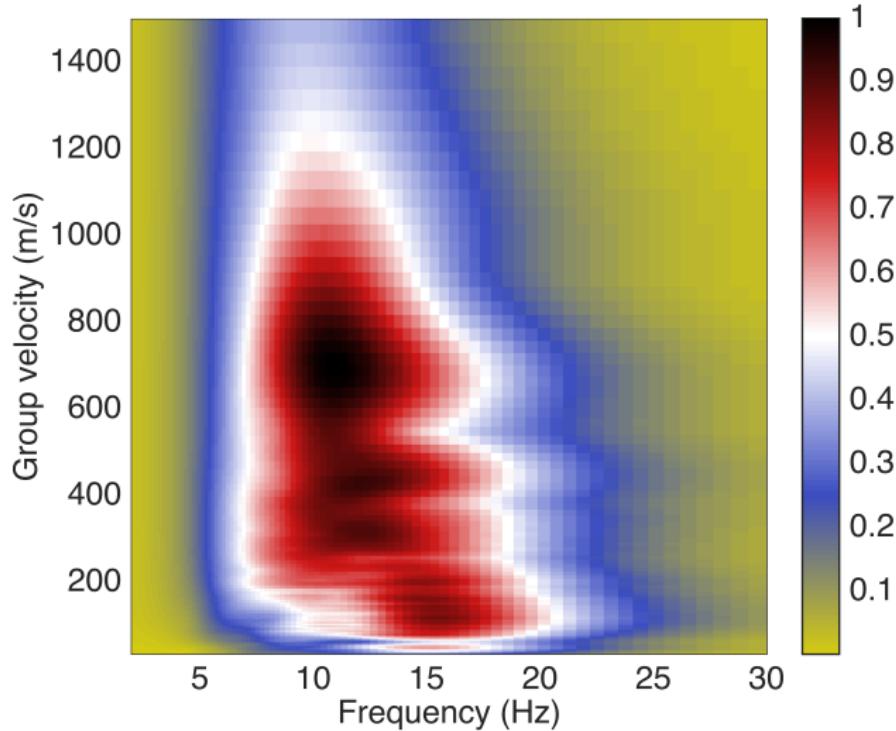
Group velocity “medium”



Group velocity “large”



Group velocity - all together



Group to phase velocity

Let v_g , v_p , ω and k denote group velocity, phase velocity, angular frequency and wavenumber respectively. We have,

$$v_g = \partial_k \omega,$$

$$v_p = \frac{\omega}{k}.$$

Let s denote slowness,

$$v_g = \partial_k(k v_p) \quad \text{or} \quad s_g = \partial_\omega(\omega s_p)$$

and so

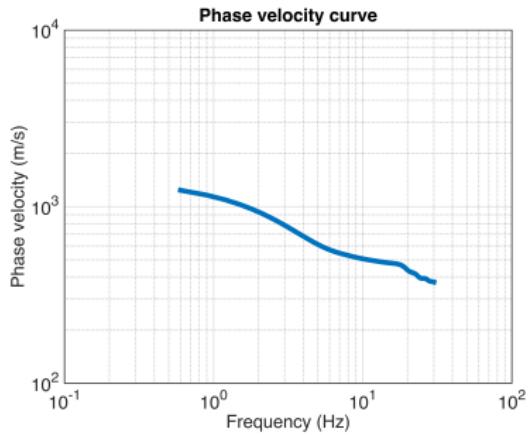
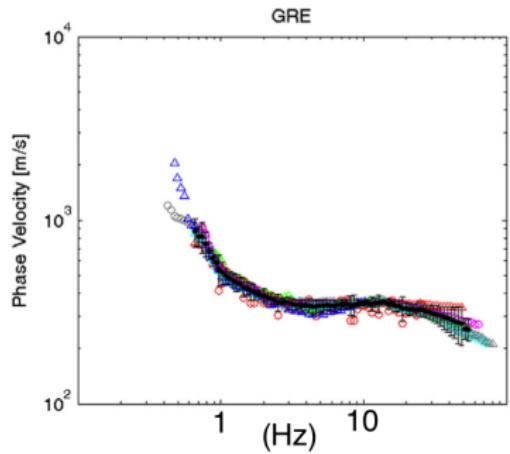
$$v_g = k \partial_k v_p + v_p \quad \text{or} \quad s_g = \omega \partial_\omega s_p + s_p.$$

Solving these differential equations we have,

$$v_p(\omega) = \frac{1}{k} \left(k_o v_p^o + \int_{k_o}^k v_g(k) dk \right) \quad \text{or}$$

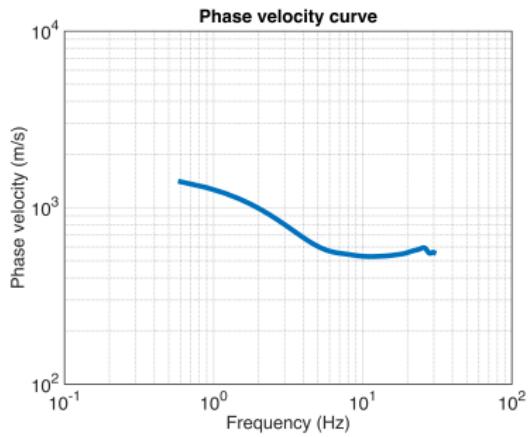
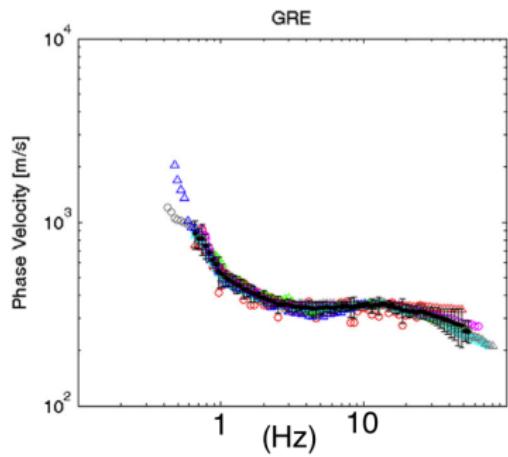
$$s_p(\omega) = \frac{1}{\omega} \left(\omega_o s_p^o + \int_{\omega_o}^\omega s_g(\omega) d\omega \right).$$

Phase velocity - Idea 1



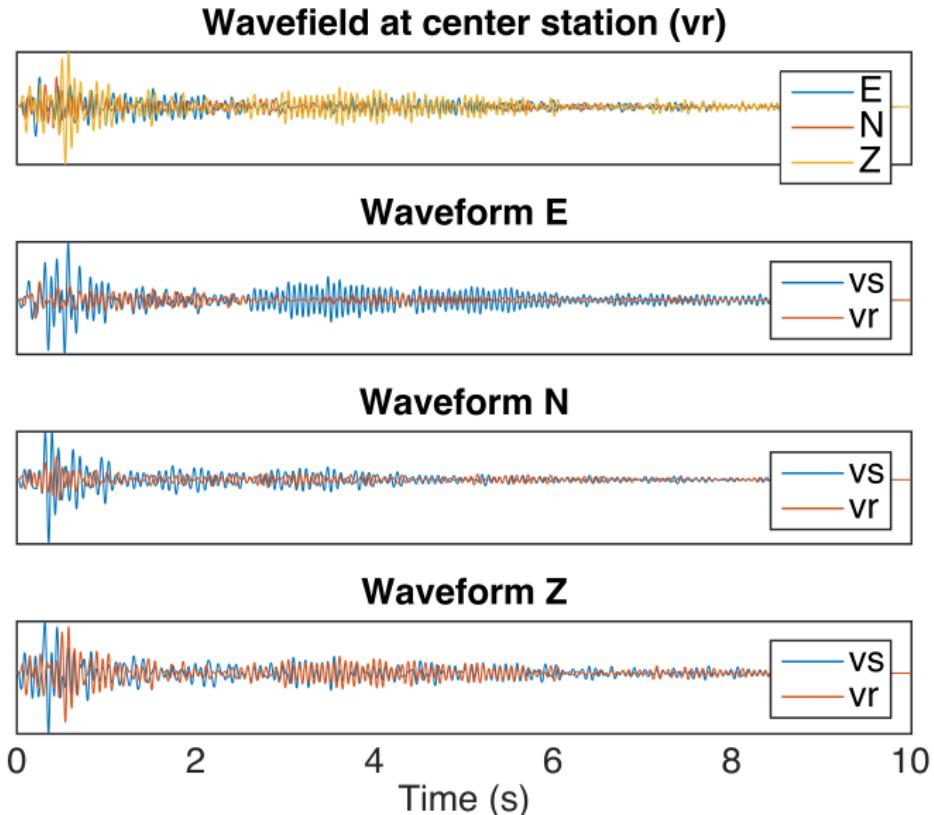
Garofalo et al, 2016

Phase velocity - Idea 2



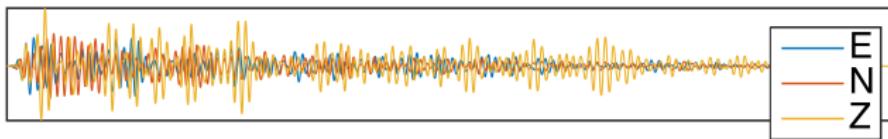
Garofalo et al, 2016

Shot gather - Idea 1

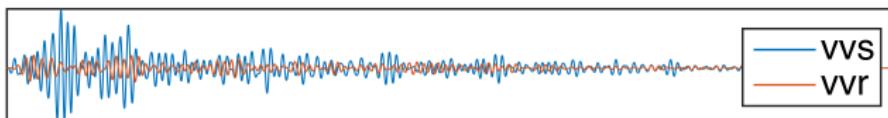


Shot gather - Idea 2

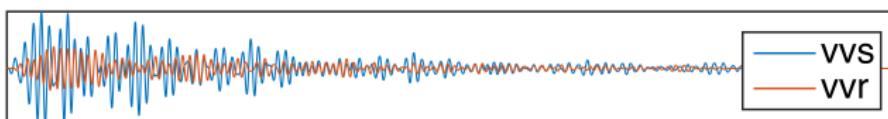
Wavefield at center station (vvr)



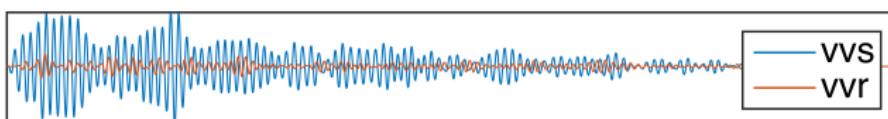
Waveform E



Waveform N



Waveform Z



0 2 4 6 8 10
Time (s)

Observations

Interesting

- This method overshoots (or undershoots) values for v_p
- Concavities of v_p are switched
- Interferometry-inception enhances contrasts in v_p
- Sensitive knobs to turn:

- Frequencies for bandpass:

data can get aliased and turn into wiggly cones

- Time windows for xcorrelations:

group dispersion semblance can become skewed and useless

- Choosing s_p^o :

phase velocities can result increasing in frequency