How will climate change impact Spanish dryland biological soil crust growth & diversity?

Introduction

During the last century, changes in temperature and precipitation have impacted ecosystem structure and processes worldwide (Montoya & Raffaelli, 2010). These changes in climate have been attributed to the global migration of >2,000 species by 16.9 km/decade towards the poles (Chen *et al.*, 2011), and threaten one in six species with extinction (Urban, 2015). Fully understanding how biotic communities will respond to changes in temperature and precipitation, is crucial to improving current forecasts of the consequences of climate change (Zhou *et al.*, 2012), particularly in ecosystems where rainfall is already scarce and unpredictable such as drylands (Maestre *et al.*, 2012b; Huang *et al.*, 2016). Drylands cover ~40% of the Earth's surface and are home to over 38% of the human population (Reynolds *et al.*, 2007). However, there remain significant unknowns regarding dryland functioning in the context of future change, including the limited understanding of how biological soil crusts (BSCs) will respond to climate change (Belnap & Weber, 2013; Young *et al.*, 2016), as BSCs have only been a topic of research since the 1960s (eg. Friedmann *et al.*, 1967).

BSCs are the result of a close association between soil particles and cyanobacteria, eukaryotic algae, microfungi, lichens, and bryophytes which live within, or immediately on top of, the uppermost millimetres of the soil surface (Belnap *et al.*, 2003; Figure 1). Due to their poikilohydric nature, BSC organisms are uniquely adapted to limited moisture and low nutrient conditions and respond rapidly to pulsed environmental conditions, characteristic of dryland environments (Lange, 2003). For example, Feng *et al.* (2014) find cyanobacteria, lichen and moss crusts require negligible amounts of moisture to function (0.38 ± 0.17 mm, 0.92 ± 0.06 mm, and 2.10 ± 0.02 mm, respectively). Consequently, BSCs are the dominant functional vegetation unit in some of the harshest habitats in the world, comprising up to 70% of the living cover in dryland environments (Belnap & Lange, 2003). However, due to a soil pH of ≥7 and high potential evapotranspiration cyanobacteria, mosses and lichens are the dominant crust organisms in varying amounts (Belnap, 2006; Figure 1).

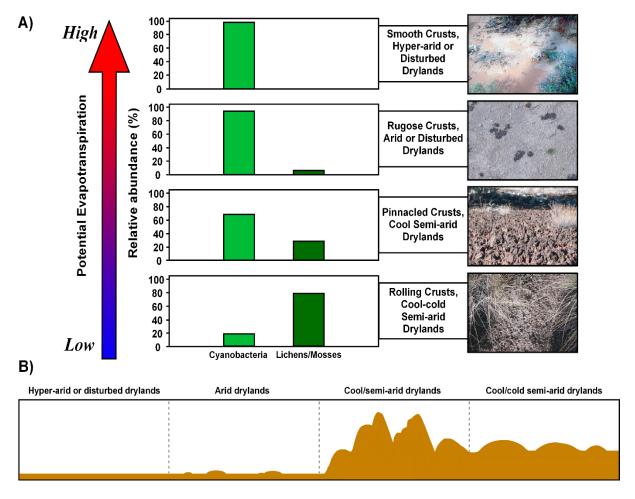


Figure 1 | A) Potential evapotranspiration and associated BSC community composition (with corresponding image of crusts on the right). **B)** Conceptual cross-sectional diagram of the four main soil crust types and their distinctive micro-topographies. Created in Serif (v. X6) using data and by adapting graphics in Belnap (2006).

BSCs are widely considered to be important ecosystem engineers, aggregating and stabilising soil surfaces, playing a dominant role in numerous ecosystem functions (Belnap & Lange, 2003; Figure 2). For example, recent reviews by Elbert *et al.* (2009; 2012) suggest annual global BSC net uptake of nitrogen (N) to be 45-49 Tg and 3.6-3.9 Pg for carbon (C), with ~1 Pg of this fixed by BSCs in arid and semi-arid regions. They may also buffer dryland ecosystems against the negative impacts of climate change (Delago-Baquerizo *et al.*, 2016; Figure 2). This is illustrated by Maestre *et al.* (2015) who find the micro-topography of semi-arid BSCs mitigate the warming-induced physiological stress on Gram-negative bacterial communities.

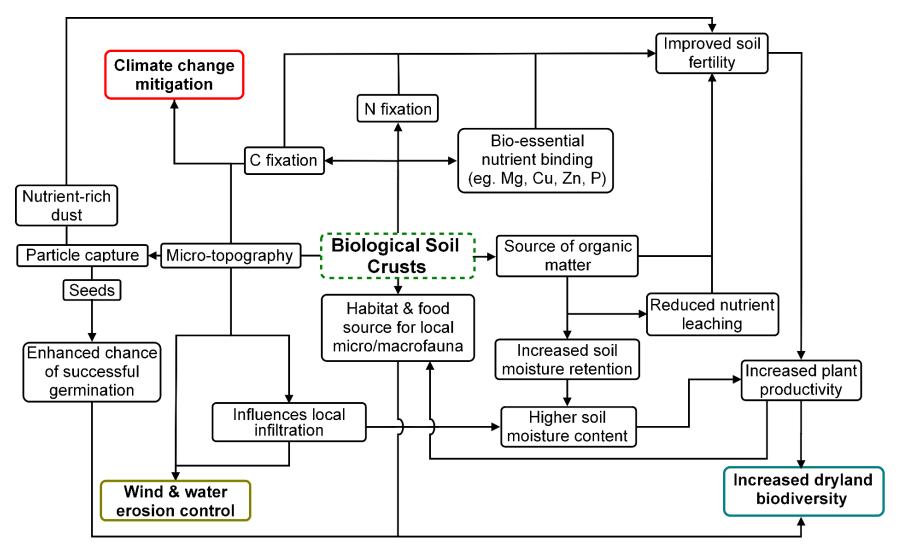


Figure 2 | Simplified conceptual model of BSC ecological roles and functions. Created in Serif (X6) using data from Elbert *et al.* (2009; 2012), Belnap (2006a), Harper & Belnap (2001), Williams *et al.* (2012), Hawkes (2004), Eldridge *et al.* (2000), Neher *et al.* (2009), Housman *et al.* (2006). See Belnap (2003) and Viles (2008) for full reviews.

However, like all biota, their reliance on ambient environmental conditions to function also makes BSCs vulnerable to climate change (Maestre et al., 2012a; Dettweiler-Robinson et al., 2013). Climate change is expected to negatively impact soil crust lichens and mosses (Tuba et al., 2011). This has been observed in the Succulent Karoo, Southern Africa, and on the Colorado Plateau, USA where warming and altered precipitation frequencies have negatively impacted the functioning and mortality of moss and lichen dominated BSCs (Zelikova et al., 2012; see Figure 3). Additionally, reductions in BSC cyanobacteria have also been reported with change in precipitation patterns (Johnson et al., 2012). Ferrenberg et al. (2015) and Steven et al. (2015) find climate change is as detrimental to BSC functioning and survival as chronic physical disturbances either by livestock/human trampling or off-road vehicles. Alterations in BSC community composition, due to environmental change, may considerably impact their ability to provide certain ecosystem services. For example, rougher surfaces trap dust more effectively, provide more habitat for soil invertebrates and increase soil moisture retention in comparison to flatter, cyanobacteria dominated crusts (Belnap, 2006; Viles, 2008; Figure 1).

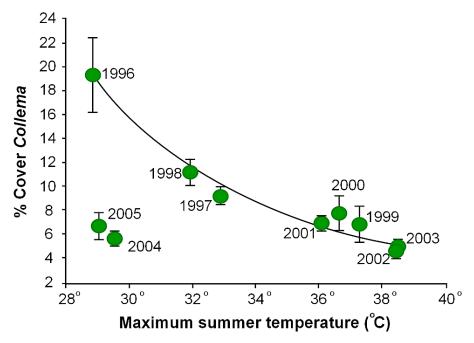


Figure 3 | Decline of *Collema* lichen cover with increasing average maximum temperature of the previous June in the Canyonlands National Park, SE Utah ($r^2 = 0.97$). Created using data from Belnap *et al.* (n.d) and Belnap *et al.* (2006b).

Increased temperature, and reductions in precipitation, have been observed across Spain since the 1960s (Gosling *et al.*, 2011). Furthermore, there is a high confidence that the drylands of the Iberian Peninsula are likely (>66% probability) to be severely

impacted by climate change (Field *et al.*, 2012; Kovats *et al.*, 2014). Despite this, few studies have elucidated the effects of changing ambient temperature and moisture on the BSCs of dryland Spain (see Maestre *et al.*, 2011; Figure 4), especially compared to the USA and Australia (Belnap & Weber, 2013). This has led to a limited understanding of their potential responses to climate change. This essay aims to: (i) outline the role temperature and moisture availability play in BSC growth; (ii) summarise the projected future climate change scenarios for Spain (up to 2100); and (iii) draw together this information to discuss the potential impacts these changes may have on Spanish dryland BSC growth and diversity.

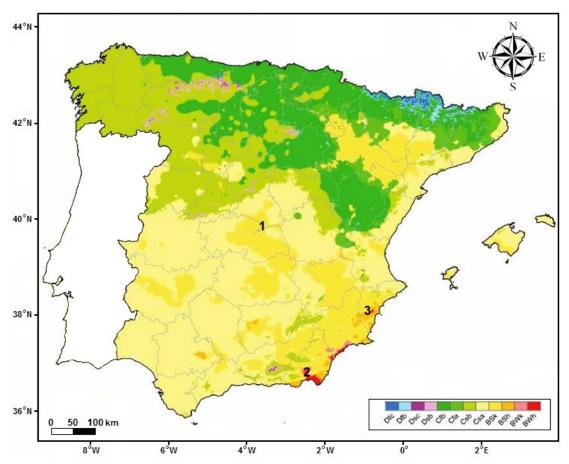


Figure 4 Climate of Spain according to temperature and precipitation data from 1971-2000, using the Köppen-Geiger Climate Classification, ie. wetter continental climates (Group D: blue-dark green) to Mediterranean and to arid climates (Group B: light green-red). Adapted from AEMET (2011). Numbers denote general study areas on the interactions between climate and dryland BSCs (eg. Escolar *et al.*, 2012 = 1; de Guevara *et al.*, 2014 = 2; Maestre *et al.*, 2015 = 3).

BSCs, Temperature & Moisture

Biological activity in deserts is primarily determined by the amount and availability of moisture (Noy-Meir, 1973). Similarly, low and unpredictable precipitation means desert BSCs are predominantly dry and inactive (Lange *et al.*, 1994). However, BSCs

do have an optimal level of hydration. High volumes of moisture increases diffusional resistance, decreasing CO₂ availability, whereas too little water inhibits basic cellular functioning (Grote *et al.*, 2010). Lange *et al.* (1997; 1998) observed that BSCs can obtain this moisture either directly from liquid water (eg. cyanobionts associated with lichens such as *Collema tenax*) or through dew and/or fog inputs (eg. phycobionts, associated with green algal lichens such as *Psora cerebriformis*).

Following adequate hydration, temperature plays a dominant role in the regulation of CO₂ exchange in BSCs. It is well known that photosynthesis is sensitive to inhibition by moderate heat stress (Berry & Björkman, 1980). Whilst photosynthesis has been observed in BSCs between 0-50°C, optimum temperatures for photosynthesis range from 10-28°C (Lange, 2003). This is largely attributed to alterations to key components of the Calvin cycle. As temperatures increases, the primary carboxylating enzyme Ribulose-1,5-bisphosphate carboxylase/oxygenase (RuBisCo), increases affinity for O₂ and the activase aggregates and can no longer activate RuBisCo, thereby reducing carboxylation efficiency (Crafts-Brandner & Salvucci, 2000; Parry *et al.*, 2008).

BSCs immediately lose some CO₂ after wetting through non-metabolic pathways (Farrar & Smith, 1976), such as osmotic shock and cell membrane leakage (Kieft *et al.*, 1987; Oliver *et al.*, 2000), followed by 'resaturation respiration' (Smith & Molesworth, 1973). This respiration is hypothesised to be associated with an increased energy demand to repair damages associated with desiccation. In general, desiccation events damage lipids, nucleic acids and proteins, attributed to excessive oxidative stress from the release of reactive oxygen species (ROS), such as hydroxyl radical and singlet oxygen, exacerbated by periods of intense light and UV irradiation (Holzinger & Karsten, 2013). For example, field experiments on the Colorado Plateau by Belnap *et al.* (2008) shows prolonged exposure to UV radiation leads to reductions in BSC quantum yield, chlorophyll *a*, UV-protective pigments, nitrogenase activity and exopolysaccharides.

Energy demand also increases as BSCs undertake in desiccation tolerance strategies which may include: (i) synthesising scytonemin, mycosporine-like amino acid, and carotenoid-xanthophyll UV-protective pigments to prevent ROS damage (Belnap *et al.*, 2008), (ii) the accumulation of compatible solutes including amino acids, various betaines and trehalose for osmoprotection (Kempf & Bremer, 1998), and (iii) in some

BSC organisms the production of exopolysaccharides that slow the speed of water loss (eg. observed in cyanobacteria by Rajeev *et al.*, 2013).

To reach net C gain, and therefore ensure long-term survival, growth and functioning, BSCs must remain hydrated long enough to overcome this initial C loss, repair and increase their resilience to next desiccation event (Belnap *et al.*, 2004). Consequently, Büdel *et al.* (2009) and Coe *et al.* (2012) suggest that rain frequency and the subsequent length of hydration period, largely determines BSC carbon balances and consequently survival, development and composition. This has been illustrated by Li *et al.* (2014) on *Bryum argenteum* from the Tengger Desert, China, where a lower desiccation time (one week in comparison to a year) increased the effective quantum yield of photosystem II (PSII) following rehydration, with one year treatments requiring a rehydration time >5 times longer to recover. Beyond enzymatic reactions, higher temperatures also influence rates of evapotranspiration and the overall hydration period (Grote *et al.*, 2010). For example, in BSC *Nostoc flagelliforme* assemblages, Zhao *et al.* (2008) observe that variations in the rate of photosynthetic recovery following desiccation are caused by different water absorption rates at higher temperatures.

Thus whilst higher temperatures may lead to an increase in photosynthetic activity, ultimately the duration of active photosynthesis will be significantly reduced because of more rapid rates of desiccation (Lange *et al.*, 2006). Observed in the field by Zedda *et al.* (2010; 2011) who find the richness and spatial coverage of lichen BSCs, across the BIOTA transect in Southern Africa, are positively correlated to lower temperatures and increased water availability. Studies by Feng *et al.* (2014) and Grote *et al.* (2010; Figure 5) find the optimum temperature and moisture content for maximum metabolic activity length and efficiency. These optimal environmental conditions are between 20 and 30°C for cyanobacteria dominated crusts and 15°C to 20°C in lichens and mosses, at an intermediate relative water content (40-60%).

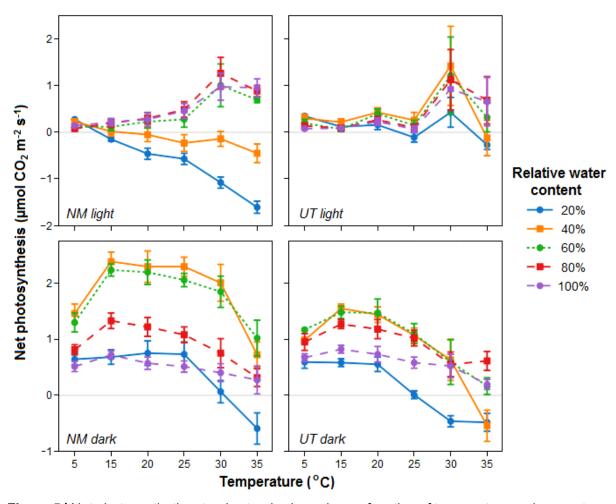


Figure 5 | Net photosynthetic rates (± standard error), as a function of temperature and percentage soil water content of light (early successional crust: cyanobacteria dominant) and dark (later successional crust: cyanobacteria/lichen and moss dominant) BSCs on the Colorado Plateau, Utah (*UT*) and at the Jornada Experimental Range, in the Chihuahuan Desert, New Mexico (*NM*), USA. Created in Plotly (v. 2.0) using data in Grote *et al.* (2010).

Projected Climate Change Scenarios for Spain & Potential Impacts on BSCs

Currently, Spain is largely arid with an annual average precipitation of 200-500 mm across mid and south-eastern Spain between 1971 and 2000 (AEMET, 2011), with days exceeding 40°C most summers in southern Spain (de Castro *et al.*, 2005; AEMET, 2014). Representative concentration pathways (RCPs) have replaced the SRES emission scenarios of previous IPCC reports: the low (RCP2.6; van Vuuren *et al.*, 2011), medium–low (RCP4.5; Thomson *et al.*, 2011), medium–high (RCP6; Masui *et al.*, 2011), and high (RCP8.5; Riahi *et al.*, 2011) pathways are plausible trajectories of global greenhouse gas emissions based on potential ecological, physical and socioeconomic developments (see Moss *et al.*, 2010; Meinhausen *et al.*, 2011; Figure 6).

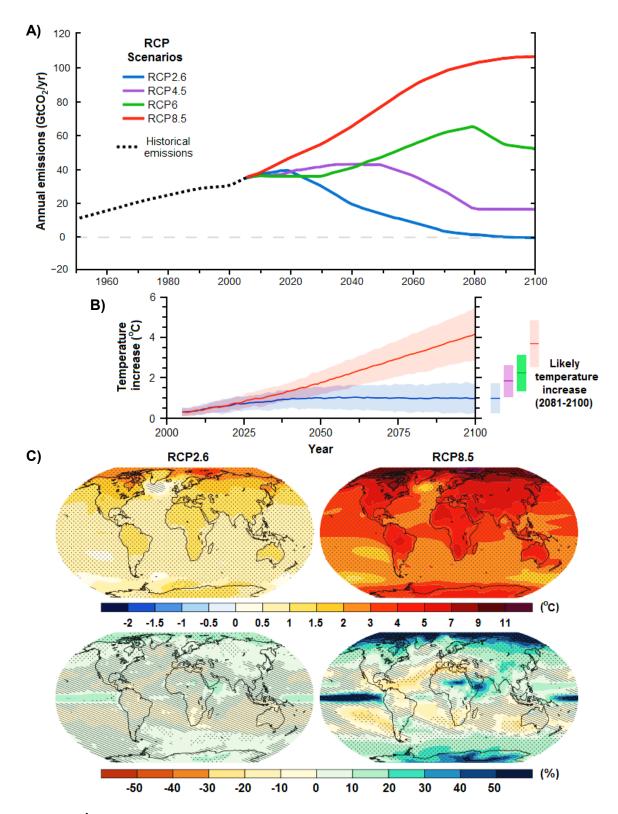
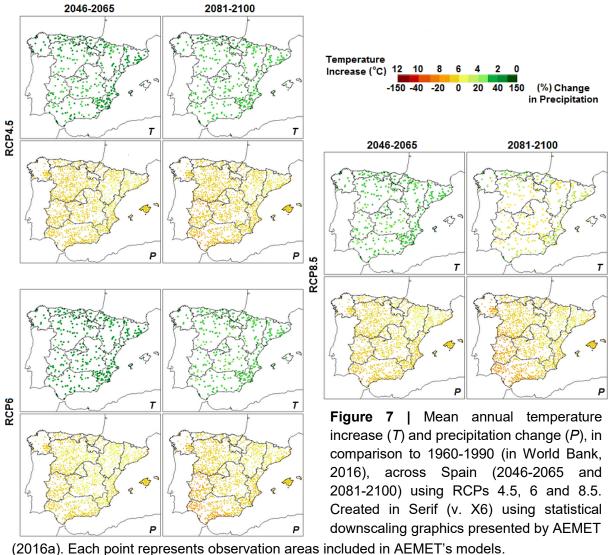


Figure 6 | A) Historical (1950-2005) and annual greenhouse gas emissions from RCPs (2005-2100). **B)** Global average surface temperature change between 2006-2100 (relative to 1986-2005) and *likely* (>66% probability) temperature increase range between 2081-2100. **C)** Global surface temperature rise and precipitation change projections (in 2081-2100 compared to 1986-2005) using the RCP 2.6 and RCP 8.5 pathways under 32 and 39 models, respectively (many these of these models and for RCPs 4.5 and 6 are available on the WorldClim (v 1.4) database). Created in Plotly (v. 2.0) and Serif PagePlus (v. X6) using data from van Vuuren *et al.* (2011), Thomson *et al.* (2011), Masui *et al.* (2011) and Stocker *et al.* (2013).

These RCPs have been utilised by the Agencia Estatal de Meteorología (AEMET) to model potential climate change scenarios across Spain (Figure 7). They find similar trends to the IPCC (C in Figure 6) and the downscaled EURO-CORDEX model for Europe (Jacob *et al.*, 2014): annual average surface temperature may rise by up to 7°C and precipitation may reduce by >30% in south-eastern Spain by 2081-2100 under RCP8.5, with increases in temperature of ~2-3°C and reductions in annual precipitation of 10% for all RCPs by 2065. However, AEMET do not include RCP2.6 in their projections as they consider that pathway to be "absolutely unrealistic" (AEMET, 2016, pers. comm., 23 March), by projecting comparatively negligible changes to present-day climatic conditions, with a ~0.5°C rise by 2081-2100 over the Mediterranean, as a result of radical, rapid mitigation activities to reduce greenhouse gas emissions by 50% over the next twenty years to stabilise present-day climatic conditions (van Vuuren *et al.*, 2011).



Spanish BSCs will undoubtedly respond differently to the varied temperature and precipitation regimes of each RCP, with increased uncertainty in the possible range of each RCP (Table 1). This uncertainty is illustrated by studies on the Colorado Plateau, which suggest lichens and mosses can generally adapt to solely increased temperatures over a number of years (Zelikova *et al.*, 2012; Reed *et al.*, 2012). Johnson *et al.* (2012) suggest a 2°C rise, similar to those at RCP4.5 and 6 by 2065, may be insufficient to prompt a significant response from BSCs. In contrast, Escolar *et al.* (2012) find a 2.4°C increase reduced lichen coverage by 40-70% in the dryland BSCs of Aranjuez, Central Spain, with mosses displaying a slight increase (0.3-7%). This observation has recently been supported by Escolar *et al.* (2015) who find a 2-3°C warming increased soil respiration at Aranjuez, and Maestre *et al.* (2015) who report a 2°C rise over 53 months reduced BSC cover by 45% in Sax, SE Spain. The temperature increase associated with RCP8.5 over a long period exceeds thresholds for optimum photosynthetic activity (Lange, 2003; Figure 5), and will undoubtedly contribute to considerable reductions in BSC growth and diversity (Table 1).

Table 1 | Summarised likely trends in Spanish BSC productivity and cover. Based on reported moisture and temperature thresholds, and the extent of projected climatic changes in temperature and precipitation, influencing activity and desiccation period frequency and length under each RCP (Figure 7; Figure 8; Table 2). Created in Serif (v. X6).

RCP	Lichens		Mosses		Cyanobacteria			
- KCP	2046-2065	2081-2100	2046-2065	2081-2100	2046-2065	2081-2100		
2.6	_	_	_	_	_	_		
4.5	$\hat{\mathbb{T}}$	$\hat{\mathbb{T}}$	-\ [↑]	$\hat{\mathbb{T}}$	\Diamond	$\hat{\mathbf{t}}$		
6	$\hat{\mathbb{T}}$	∜/♣	- \û	↑/ ♣	\bigcirc	- \û		
8.5	+	↓	↓	↓	∜/↓	₽ / ↓		
	↑ Small increase			— Negligible change				
	∏	Small (hollow) and large (solid) decrease			/ Uncertainty in future trend, either outcome is possible			

Projected reductions in annual precipitation (Figure 7) will also negatively impact BSC growth and diversity. This is largely in response to prolonged unfavourable environmental conditions for metabolic activity (Figure 5) instigating continued negative C balances, and reducing the capacity to produce or repair chlorophyll *a*

and/or protective pigments (Belnap *et al.*, 2004). Indeed, a combined 2-3°C experimental warming and rainfall exclusion conducted by Maestre *et al.* (2013), in Sorbas, SE Spain and Aranjuez, finds a reduction in photosynthetic activity (~3.5 times lower) and in lichen and moss cover of >40% over 4 years. Seasonal temperatures and precipitation regimes (Table 2), as opposed to annual trends, under each RCP further elucidates the primary stresses and potential drivers of mortality in Spanish BSCs.

Table 2 | Mean changes in seasonal surface temperature and precipitation across the dryland regions of Spain in Figure 4, compared to average historical (1960-1990) environmental conditions presented in World Bank (2016). Temperature results are an average of the individual points of data presented in each seasonal graph by AEMET (2016a), with precipitation estimated as a range of the most common points in AEMET (2016a). Each season is defined as: Spring (March-May), Summer (June-August), Autumn (September-November) and Winter (December-February). See Figure 8 for an example of these seasonal graphs by AEMET (2016a).

	Season	Tempera	ture (°C)	Precipitation (mm)		
	Spring	11.8		53.5		
	Summer	20).9	23.3		
Historical	Autumn	14	.7	58.5		
	Winter	7.1		66.4		
RCP		Temperature	increase (°C)	Precipitation change (%)		
KOF		2046-2065	2081-2100	2046-2065	2081-2100	
	Spring	1.92	2.77	-10 - +10	-10 – +10	
4.5	Summer	2.98	3.67	-10 - +10	-10 – +10	
4.5	Autumn	2.68	3.22	- 20 – 0	- 20 – 0	
	Winter	2.00	2.58	-1 0 – 0	-10 – 0	
	Spring	2.01	2.97	-10 – 0	-20 – 0	
•	Summer	3.89	4.49	-10 - +10	-2010	
6	Autumn	2.50	3.31	- 10 – 0	- 20 – 0	
	Winter	1.89	2.36	0 – +10	-20 – 0	
	Spring	4.11	4.84	-10 – -20	-20 – 0	
0.5	Summer	4.07	6.61	-30 - +20	-40 – +30	
8.5	Autumn	3.78	6.05	- 20 – 0	- 40 – 0	
	Winter	4.10	4.52	-10 – 0	-10 – -20	

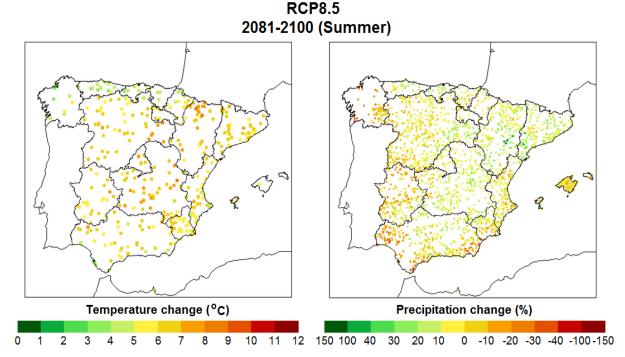


Figure 8 | Projected trends in summer temperature and precipitation under RCP8.5, used to calculate average seasonal temperature statistics and estimate seasonal rainfall change in Table 1. Created in Serif (v. X6) with graphics from AEMET (2016b) and AEMET (2016c).

Belnap *et al.* (2007) find scytonemin and canthoxanthin concentrations did not increase without 119 mm of precipitation over 6 months, with high concentrations of echinenone and β-carotene only maintained >98 mm, in BSCs of the Mojave Desert. Based on this data, Spanish BSC recovery periods will be seasonally restricted under some RCPs, or even for most/all of the year under climate change associated with higher CO₂ concentrations (Table 2). Furthermore, seasonal temperature increases may exacerbate this issue, by increasing evapotranspiration rates, resulting in wasted BSC activations. For example, Johnson *et al.* (2012) find a 2-3°C warming and an increased frequency of small (1.2 mm) rainfall events reduced BSC biomass (on average 78% less chlorophyll *a*) in Moab, Utah, as this water had evaporated within the ~30 minutes it took for the BSCs to become metabolically active. This is likely to occur in the summer under each RCP (Table 2), and even during the other seasons under some RCPs, exacerbating the reductions in precipitation (eg. RCP8.5; Table 2).

Nocturnal moistening by fog or dewfall heavily influences the photosynthetic activity and distribution patterns of lichen BSCs in Mediterranean drylands such as the Negev, Israel (Veste *et al.*, 2001), with some evidence of dew playing an important role in PSII activity in cyanobacterial crusts in the Hopq Desert, China (Rao *et al.*, 2009). Dewfall contributes greatly to the local moisture budget in SE Spain, although only 0.17 ± 0.10

mm/night between 2007 and 2010, this represented a 94% contribution to the local water balance during dry periods, despite being shorter in comparison to winter and autumn dewfalls (Uclés et al., 2014). The distribution and growth of Teloschistes lacunosus (del Prado & Sancho, 2007), in SE Spain are highly dependent on dewfalls, with Diploschistes diacapsis only active during these periods of prolonged dewfall (Pintado et al., 2010). Similar results have been observed by Maestre et al. (2013) and de Guevara et al. (2014) in Central and SE Spain, where net CO2 uptake in BSCs occurred predominantly during late autumn and winter, during these periods of increased moisture availability. Wilske et al. (2008) highlight that dew may also be important in initiating a 'pre-activated' state in BSCs during the transition from a long summer anabiosis towards the first rains later in the year. Maestre et al. (2013) find a 2-3°C warming substantially reduces the duration of suitable conditions for the formation and residence time of dewfall in Spain. Therefore increased temperatures will undoubtedly exacerbate the C deficit of Spanish BSCs by further reducing moisture availability, contributing to increased stress and mortality in BSC organisms (Table 1). For example, in South Africa, Maphangwa et al. (2012) find significant reductions in the quantum yield of BSC-forming lichens (up to 46%) with a 2.1-3.8°C warming and a 30.1-31.9% reduction in dew and fog inputs.

Overall, the decrease in length and frequency of recovery periods, associated with higher temperatures and reduced moisture availability, will likely lead to a reduction in BSC lichens and mosses. The most extreme environmental conditions may even instigate high stress or mortality in cyanobacteria, and thus a complete lack of BSC cover, under some RCPs (Table 1). For example, mosses may be more adaptable, compared to lichens, to the reduced precipitation and increases in temperature (Bjerke et al., 2011). Therefore, climate change under RCP4.5 and RCP6 by 2065, may either negatively impact their growth, or their adaptability may be able to negate these unfavourable conditions, as previously highlighted by Escolar et al. (2012). However, this is unlikely to continue long-term (by 2100), as environmental conditions deteriorate (Table 1). Similarly, cyanobacteria can tolerate more extreme environmental conditions (Figure 5), especially as species such as *Microcoleus* use their filamentous sheaths to glide to the soil surface upon wetting, avoiding UV radiation and temperature damage during desiccation periods (Mazor et al., 1996). Zelikova et al. (2012) find cyanobacteria dominate BSC assemblages as lichen and moss cover

reduces, leading to potential increases in cyanobacterial crusts under lower RCPs (Table 1). These unique coping mechanisms may also negate the impacts of climate change under RCP6 by 2100. However, under RCP8.5 these coping mechanisms may only mitigate the impacts of climate change, as even the optimum environmental thresholds for cyanobacteria are exceeded (Figure 5). This may lead to a complete lack of BSC cover, with widespread mortality amongst all BSC organisms (Table 1).

Additionally, the increased CO₂ concentrations driving these changes exacerbates the uncertainties involved in these estimations. Early experimental work by Tuba *et al.* (1998) on lichens and mosses, and more recently by Lane *et al.* (2013) on cyanobacteria, find increased ambient CO₂ concentrations may result in enhanced growth and functioning. However, in contrast, Wertin *et al.* (2012) find this is primarily mediated by ambient moisture availability, with a 10 year PCR survey of cyanobacteria small-subunit rRNA genes, by Steven *et al.* (2012), suggesting increased CO₂ may reduce cyanobacterial abundance and diversity.

Conclusions

Despite some uncertainties in projected climate change scenarios and the potential impacts of these changes, it is clear that continued unprecedented greenhouse gas emissions, and associated changes in temperature and moisture availability, will drastically impact BSC growth and diversity (Table 1). Reductions in BSC assemblage diversity, and the length and duration of their metabolically active periods, will also impact local ecosystem processes. Either through alterations in BSC microtopography with cyanobacterial crusts dominating (Figure 1), to a complete lack of ecosystem services should BSC cover reduce to negligible amounts, or disappear entirely, which is a possibility under RCP8.5. Further research is essential across Spain, and globally, to fully elucidate the response of dryland BSCs, and their surrounding ecosystem, to projected changes in change.

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