

Constrained Circulation Ensemble

D. Kennedy^{1,2}, D. M. Lawrence¹, I. R. Simpson¹

¹Climate and Global Dynamics Laboratory, NSF-NCAR

²Earth Research Institute, University of California Santa Barbara

Key Points:

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9 **Abstract**

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11 **Plain Language Summary**

12 Enter your Plain Language Summary here or delete this section. Here are instructions
 13 on writing a Plain Language Summary: <https://www.agu.org/Share-and-Advocate/Share/Community/Plain-language-summary>
 14

15 **1 Introduction**

16 The summer of 2020 was the driest of the meteorological record within the Southwest
 17 United States (Mankin et al., 2021; Hoell et al., 2022), capping off an extended twenty
 18 year megadrought (Williams et al., 2022). Models incorrectly predict positive vapor pres-
 19 sure trends in semi-arid regions (Simpson et al., 2024).

20 **2 Methods**

21 Our study focused on the ‘four corners’ regions of the western United States, which
 22 is defined as the region encompassing four adjoining states: Colorado, Utah, New Mex-
 23 ico, and Arizona. This region was chosen due to the spatial footprint of the 2020 drought
 24 (Supp Figure 1), as well as its significance for water management (Mankin et al., 2021)
 25 Our goal was to create a large sample of simulations exhibiting extreme drought con-
 26 ditions, while still allowing for local land-atmosphere coupling, in order to investigate
 27 the influences of antecedent conditions and climate change.

28 We ran multiple ensembles of simulations using CESM2, each consisting of 30 mem-
 29 bers, with varied initial conditions. For each simulation, winds were nudged to the ERA5
 30 reanalysis for the year 2020. Nudging occurred four times per day in each grid cell out-
 31 side the box bounded by 28°N, 47°N, 130°W, 96°W. This serves to constrain the large-
 32 scale circulation, imposing meteorological drought conditions for every ensemble mem-
 33 ber. Inside the bounding box, the atmosphere evolves freely, allowing for full land-atmosphere
 34 coupling, which is necessary for studying the effects of initial conditions. Sea surface tem-
 35 peratures (SSTs) were imposed from an ERA5 reanalysis product, and all other forcing
 36 (e.g. CO₂, aerosols) followed the standard conventions for transient simulations as in the
 37 CESM2-LE. The simulations were initialized in April, using random initial conditions
 38 from the CESM2-LE for all the requisite land and atmosphere states (e.g. temperature,
 39 soil moisture, humidity).

40 Our primary ensemble was based in 2020, mirroring the extreme 2020 Southwest
 41 U.S. drought. We also simulated this drought during pre-industrial and future conditions
 42 to investigate the influence of climate change. All of the ensembles were nudged to the
 43 2020 winds, but initial conditions and other climate forcings were drawn from the ap-
 44 propriate time period mirroring the CESM2-LE, effectively porting the 2020 circulation
 45 to these two alternative climates. SSTs were anomaly-based, computing anomalies from
 46 the CESM2-LE and applying them to the ERA5 reanalysis SSTs used in the 2020 en-
 47 semble. Initial conditions were drawn at random from the CESM2-LE from either April
 48 1850 or April 2090, and all other forcing (e.g. CO₂, aerosols) followed the standard con-
 49 ventions for transient simulations as in the CESM2-LE, which follows the SSP3-7.0 emis-
 50 sions scenario.

51 We have compared our ensembles to the existing CESM2-LE (Rodgers et al., 2021),
 52 which consists of 100 transient simulations (1850-2100) of the fully coupled configura-
 53 tion of CESM2. This ensemble is useful, because it provides a large number of model re-
 54 alizations, sampling across a range of internal climate variability. For the sake of com-

55 parison, we have utilized the same model source code and utilized the same forcing where
 56 applicable. The two key differences between our ensembles and the CESM2-LE are that
 57 1) we impose the 2020 drought circulation via nudging and 2) we forced the simulations
 58 with observed SSTs (in lieu of dynamically coupled SSTs).

Table 1. Ensemble descriptions

Name	Winds	Climate forcing	SSTs	Initial Conditions
Control	ERA5-2020	2020 (SSP3-7.0)	ERA5-2020	CESM2-LE (2020)
Pre-industrial	ERA5-2020	1850	ERA5-2020 + CESM2-LE anomalies (1850)	CESM2-LE (1850)
Future	ERA5-2020	2090 (SSP3-7.0)	ERA5-2020 + CESM2-LE anomalies (2020)	CESM2-LE (2090)

59 **3 Outline and in flux**

60 **3.1 Key points**

- 61 • CCE can reproduce the 2020 drought
- 62 • Soil water initialization influences drought severity via variables (soil moisture, runoff,
 63 temperature extremes), but does not exacerbate the meteorological drought, it-
 64 self
- 65 • Warming affects this drought
 - 66 – ET reduced in both free-running and constrained droughts
 - 67 – P reduced in CCE, but mixed signal in free-running CESM2-LE
- 68 • 2090 CCE is much drier than the CESM2-LE

69 **3.2 Some numbers I'd like**

- 70 • 2020 JAS Precip (obs or CCE?) sits at the XXth percentile of the CESM2 large
 71 ensemble.

72 **3.3 Supp figures I'll need**

- 73 • CESM2-LE JAS Precip histogram with obs and/or CCE

74 **4 Results and Discussion**

75 **5 Conclusions**

76 **Open Research Section**

77 This section MUST contain a statement that describes where the data supporting
 78 the conclusions can be obtained. Data cannot be listed as "Available from authors" or
 79 stored solely in supporting information. Citations to archived data should be included
 80 in your reference list. Wiley will publish it as a separate section on the paper's page. Ex-
 81 amples and complete information are here: <https://www.agu.org/Publish with AGU/Publish/Author Resources/Data for Authors>

83 **Acknowledgments**

84 Enter acknowledgments here. This section is to acknowledge funding, thank colleagues,
 85 enter any secondary affiliations, and so on.

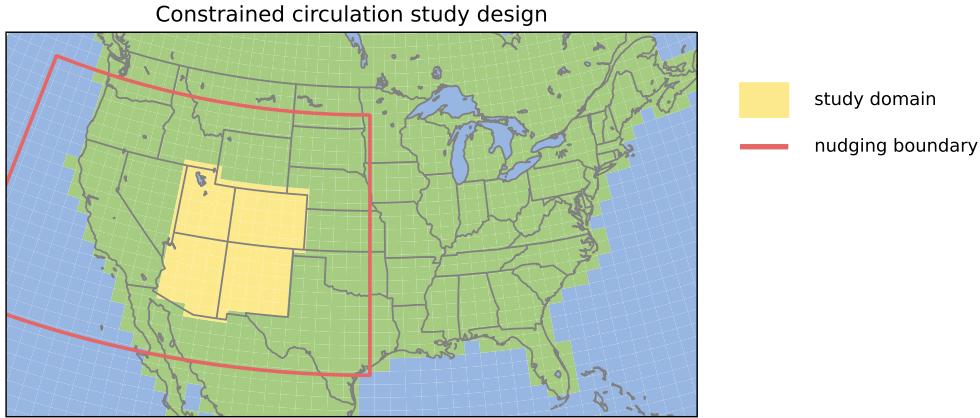


Figure 1. Map of the study domain and nudging boundary. All analyses were based in the ‘four corners’ region of the western United States, yellow area. Model simulations were nudged to reanalysis winds outside of the red box, in order to induce the observed large-scale circulation, while allowing for local land-atmosphere coupling.

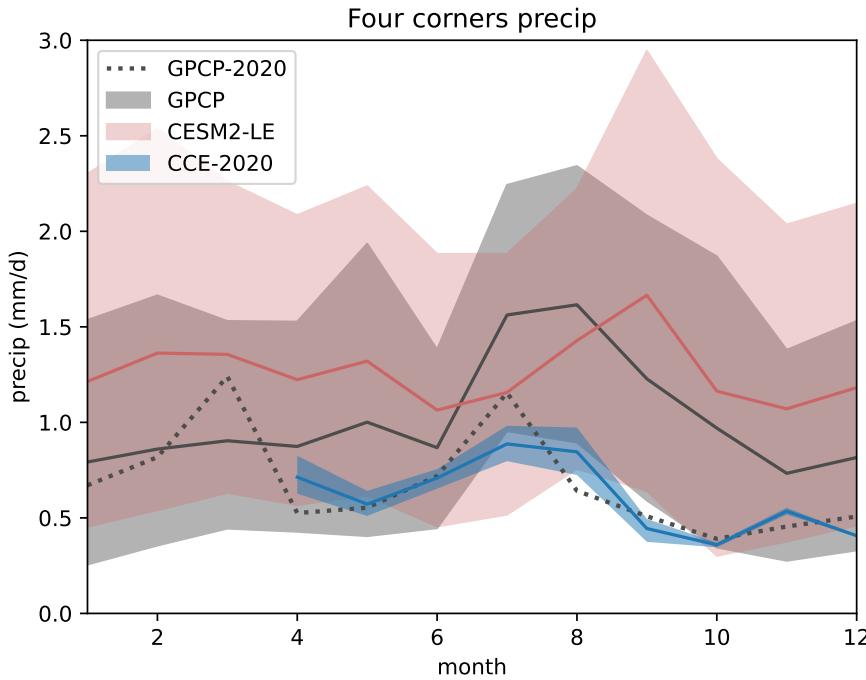


Figure 2. Precipitation climatologies for the four corners region from reanalysis (GPCP, 1981–2020) and a model large ensemble (CESM2-LE, 1981–2020) alongside our constrained circulation ensemble (CCE-2020). For each of the datasets, shading spans the 5th to 95th percentiles (across years and/or ensemble members), and the solid line tracks the mean. The dotted black line shows the reanalysis precipitation for 2020. The CESM2-LE shows a small high bias in mean and variance relative to reanalysis precipitation, but does not effectively capture the timing or magnitude of the summer monsoon. When nudged to 2020 winds, CESM2 can reproduce the 2020 drought, even though it falls outside the large ensemble envelope. N.b. will combine this with Figure 1.

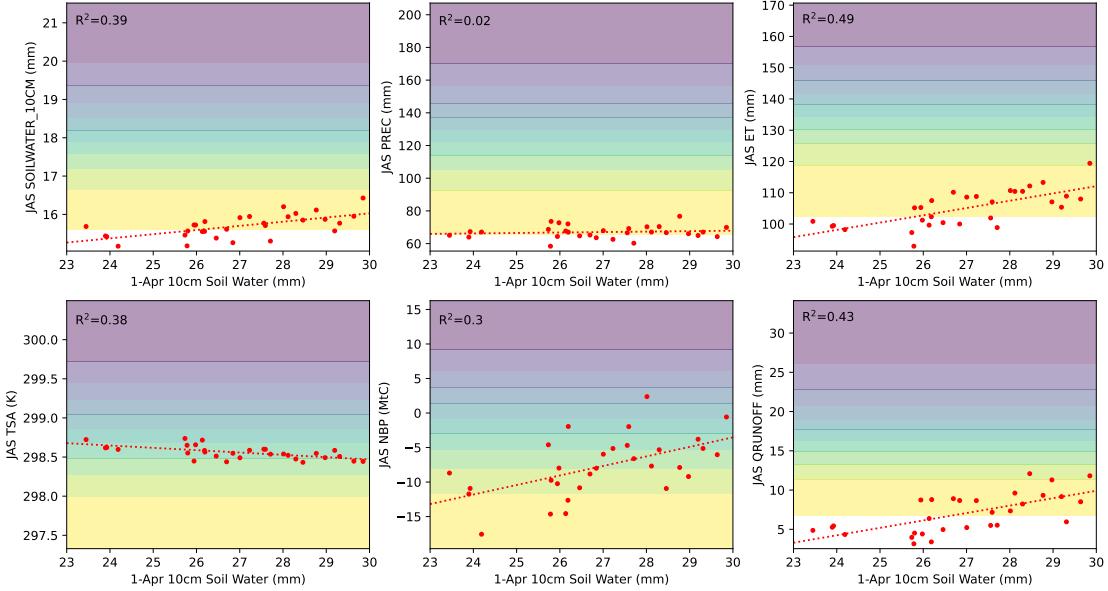


Figure 3. The influence of soil water initialization on various summer (JAS) drought indicators in the CCE, shown in the context of the variance across the CESM2-LE. In each case, the 10cm soil water content at initialization is plotted (red dots) against the summer drought indicator, with linear regression where appropriate (red lines). The colored stripes span the CESM2-LE deciles for the given drought indicator, truncated at the 1st and 99th percentiles. Dots in the white background area, exist below the 1st percentile of the CESM2-LE. Initial soil water content influences many drought indicators, but not the severity of the meteorological drought itself, i.e. summer precipitation.

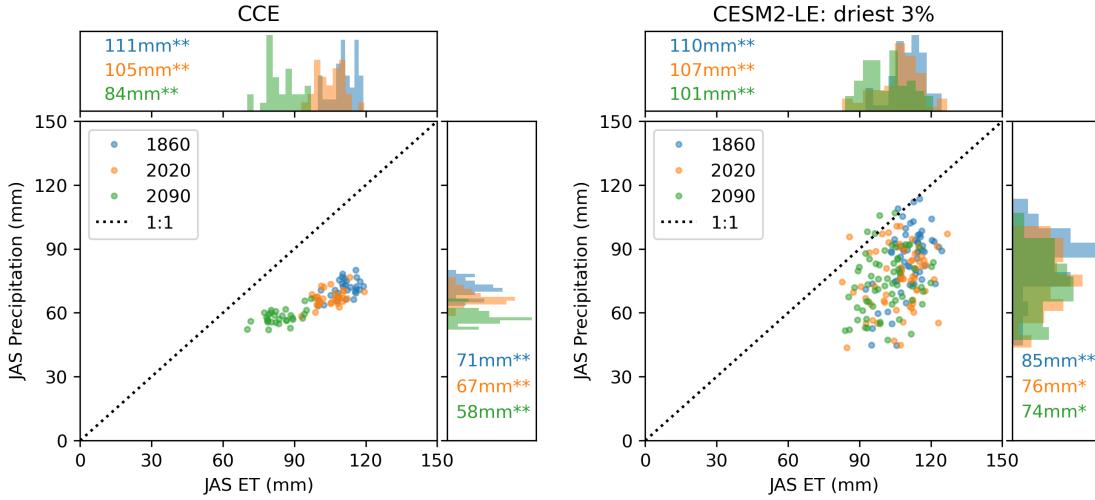


Figure 4. Summer evapotranspiration vs. precipitation in the CCE (a) and a dry subset of the CESM2-LE (d), alongside histograms of ET (b,e) and precipitation (c,f) across three time periods. The CESM2-LE is subset to the 30 driest ensemble-years for each period from 1000 total ensemble-years (100 members, 10 years around the given year) based on JAS 10cm soil water content. Stars indicate significant differences in the mean relative to the two other time periods, via t-test at the $p < 0.05$ level. Forced responses tend to reduce ET and precipitation in both the CCE and the dry subset of the CESM2-LE.

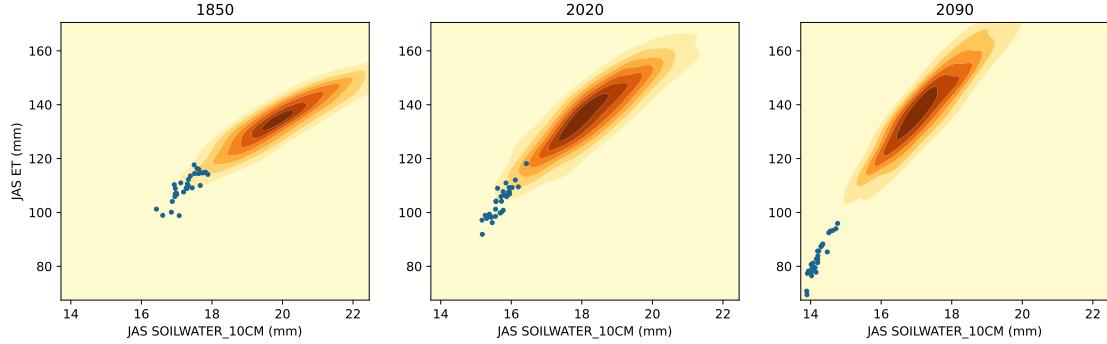


Figure 5. Summer evapotranspiration vs. 10cm soil water content from the CCE (dots) and the CESM2-LE (scatter density contours) across the three time periods. The relationship in the CCE closely follows the CESM2-LE pattern. The CCE exists at the dry extreme of the CESM2-LE, or even beyond, in the 2090 ensemble.

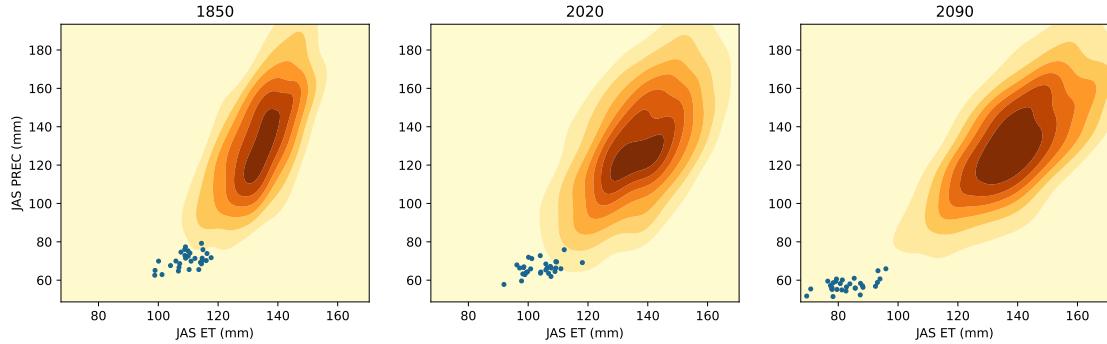


Figure 6. Summer precipitation vs. evapotranspiration from the CCE (dots) and the CESM2-LE (scatter density contours) across the three time periods. The location of the CCE data comports with the CESM2-LE pattern, but the slopes do not appear to match. N.b. will combine this with Figure 5.

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