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天然河道水流运动一般都属于三维流动，运动要素即沿程变化，又沿水深和河宽方向变化。由于三维水流运动比较复杂，河流数值模拟常用的一种简化方法是将运动要素沿水深方向平均，把三维问题转化为平面二维问题。本节基于一定条件将三维流动的雷诺平均运动微分方程简化为平面二维浅水方程。

图 2.1: 水位基准示意图

2.3.1 浅水假设和水深平均积分法则

浅水假设

三维流动的雷诺平均运动微分方程如式(2.1a)-(2.1d)所示。

$$\frac{\partial \bar{u}_x}{\partial x} + \frac{\partial \bar{u}_y}{\partial y} + \frac{\partial \bar{u}_z}{\partial z} = 0 \quad (2.1a)$$

$$\frac{\partial \bar{u}_x}{\partial t} + \frac{\partial(\bar{u}_x \bar{u}_x)}{\partial x} + \frac{\partial(\bar{u}_x \bar{u}_y)}{\partial y} + \frac{\partial(\bar{u}_x \bar{u}_z)}{\partial z} = \bar{f}_x - \frac{1}{\rho} \frac{\partial \bar{p}}{\partial x} + \nu_t \left(\frac{\partial^2 \bar{u}_x}{\partial x^2} + \frac{\partial^2 \bar{u}_x}{\partial y^2} + \frac{\partial^2 \bar{u}_x}{\partial z^2} \right) \quad (2.1b)$$

$$\frac{\partial \bar{u}_y}{\partial t} + \frac{\partial(\bar{u}_y \bar{u}_x)}{\partial x} + \frac{\partial(\bar{u}_y \bar{u}_y)}{\partial y} + \frac{\partial(\bar{u}_y \bar{u}_z)}{\partial z} = \bar{f}_y - \frac{1}{\rho} \frac{\partial \bar{p}}{\partial y} + \nu_t \left(\frac{\partial^2 \bar{u}_y}{\partial x^2} + \frac{\partial^2 \bar{u}_y}{\partial y^2} + \frac{\partial^2 \bar{u}_y}{\partial z^2} \right) \quad (2.1c)$$

$$\frac{\partial \bar{u}_z}{\partial t} + \frac{\partial(\bar{u}_z \bar{u}_x)}{\partial x} + \frac{\partial(\bar{u}_z \bar{u}_y)}{\partial y} + \frac{\partial(\bar{u}_z \bar{u}_z)}{\partial z} = \bar{f}_z - \frac{1}{\rho} \frac{\partial \bar{p}}{\partial z} + \nu_t \left(\frac{\partial^2 \bar{u}_z}{\partial x^2} + \frac{\partial^2 \bar{u}_z}{\partial y^2} + \frac{\partial^2 \bar{u}_z}{\partial z^2} \right) \quad (2.1d)$$

在河道、湖泊或水库水流中，水平尺度一般远大于垂向尺度。如果垂向加速度与重力加速度相比很小，则可以忽略垂向加速度，流速等水力参数沿垂向的变化常采用其垂向平均值，并假定沿水深方向的动水压强分布符合静水压强分布。三维流动的运动微分方程可简化为：

$$\frac{\partial \bar{u}_x}{\partial x} + \frac{\partial \bar{u}_y}{\partial y} + \frac{\partial \bar{u}_z}{\partial z} = 0 \quad (2.2a)$$

$$\frac{\partial \bar{u}_x}{\partial t} + \frac{\partial(\bar{u}_x \bar{u}_x)}{\partial x} + \frac{\partial(\bar{u}_x \bar{u}_y)}{\partial y} + \frac{\partial(\bar{u}_x \bar{u}_z)}{\partial z} = -\frac{1}{\rho} \frac{\partial \bar{p}}{\partial x} + \nu_t \left(\frac{\partial^2 \bar{u}_x}{\partial x^2} + \frac{\partial^2 \bar{u}_x}{\partial y^2} + \frac{\partial^2 \bar{u}_x}{\partial z^2} \right) \quad (2.2b)$$

$$\frac{\partial \bar{u}_y}{\partial t} + \frac{\partial(\bar{u}_y \bar{u}_x)}{\partial x} + \frac{\partial(\bar{u}_y \bar{u}_y)}{\partial y} + \frac{\partial(\bar{u}_y \bar{u}_z)}{\partial z} = -\frac{1}{\rho} \frac{\partial \bar{p}}{\partial y} + \nu_t \left(\frac{\partial^2 \bar{u}_y}{\partial x^2} + \frac{\partial^2 \bar{u}_y}{\partial y^2} + \frac{\partial^2 \bar{u}_y}{\partial z^2} \right) \quad (2.2c)$$

$$\frac{\partial \bar{p}}{\partial z} = -\rho g \quad (2.2d)$$

水深积分平均法则

将式(2.2)沿水深积分平均，即可得到沿水深平均的平面二维流动的基本方程。在沿水深积分平均过程中，采用以下定义和公式：

(1) 定义水深为

$$H = \zeta - z_0 \quad (2.3)$$

式中， H 为水深， $\zeta = \zeta(x, y, t)$ 、 $z_0 = z_0(x, y, t)$ 分别为某一基准面下的水面高程和河床高程（见图）

(2) 定义沿水深平均流速 U_i 为

$$U_i = \frac{1}{H} \int_{z_0}^{\zeta} \bar{u}_i dz \quad (2.4)$$

式中，下标 i 取 1, 2 和 3 分别对应 x , y 和 z 方向的速度分量。

(3) 莱布尼兹公式

$$\frac{\partial}{\partial x_i} \int_a^b f dz = \int_a^b \frac{\partial f}{\partial x_i} dz + f|_b \frac{\partial b}{\partial x_i} - f|_a \frac{\partial a}{\partial x_i} \quad (2.5)$$

式中， $x_i = x, y, z$ 。 a 、 b 和 f 都是 x_i 的函数。

(4) 自由表面及河床底部运动学条件为：

$$\overline{u_z}|_{z=\zeta} = \frac{D\zeta}{Dt} = \frac{\partial\zeta}{\partial t} + \overline{u_x}|_{z=\zeta} \frac{\partial\zeta}{\partial x} + \overline{u_y}|_{z=\zeta} \frac{\partial\zeta}{\partial y} \quad (2.6)$$

$$\overline{u_z}|_{z=z_0} = \frac{Dz_0}{Dt} = \frac{\partial z_0}{\partial t} + \overline{u_x}|_{z=z_0} \frac{\partial z_0}{\partial x} + \overline{u_y}|_{z=z_0} \frac{\partial z_0}{\partial y} \quad (2.7)$$

2.3.2 沿水深平均的连续性方程

采用上述定义和公式对连续性方程(2.2a)沿水深积分平均得：

$$\int_{z_0}^{\zeta} \left(\frac{\partial \overline{u_x}}{\partial x} + \frac{\partial \overline{u_y}}{\partial y} + \frac{\partial \overline{u_z}}{\partial z} \right) dz = 0 \quad (2.8)$$

根据式(2.5)，式(2.8)中前两项分别可写成

$$\int_{z_0}^{\zeta} \frac{\partial \overline{u_x}}{\partial x} dz = \frac{\partial}{\partial x} \int_{z_0}^{\zeta} \overline{u_x} dz - \overline{u_x}|_{z=\zeta} \frac{\partial\zeta}{\partial x} + \overline{u_x}|_{z=z_0} \frac{\partial z_0}{\partial x} \quad (2.9)$$

$$\int_{z_0}^{\zeta} \frac{\partial \overline{u_y}}{\partial y} dz = \frac{\partial}{\partial y} \int_{z_0}^{\zeta} \overline{u_y} dz - \overline{u_y}|_{z=\zeta} \frac{\partial\zeta}{\partial y} + \overline{u_y}|_{z=z_0} \frac{\partial z_0}{\partial y} \quad (2.10)$$

式(2.8)中最后一项

$$\int_{z_0}^{\zeta} \frac{\partial \overline{u_z}}{\partial z} dz = \overline{u_z}|_{z=\zeta} - \overline{u_z}|_{z=z_0} \quad (2.11)$$

将式(2.9)，(2.10)和(2.11)代入(2.8)，并利用自由表面及河床底部运动学条件式(2.6)和(2.7)，可得

$$\begin{aligned} & \int_{z_0}^{\zeta} \left(\frac{\partial \overline{u_x}}{\partial x} + \frac{\partial \overline{u_y}}{\partial y} + \frac{\partial \overline{u_z}}{\partial z} \right) dz \\ &= \frac{\partial}{\partial x} \int_{z_0}^{\zeta} \overline{u_x} dz - \overline{u_x}|_{z=\zeta} \frac{\partial\zeta}{\partial x} + \overline{u_x}|_{z=z_0} \frac{\partial z_0}{\partial x} + \\ & \quad \frac{\partial}{\partial y} \int_{z_0}^{\zeta} \overline{u_y} dz - \overline{u_y}|_{z=\zeta} \frac{\partial\zeta}{\partial y} + \overline{u_y}|_{z=z_0} \frac{\partial z_0}{\partial y} + \\ & \quad \overline{u_z}|_{z=\zeta} - \overline{u_z}|_{z=z_0} \\ &= \frac{\partial HU_x}{\partial x} + \frac{\partial HU_y}{\partial y} + \frac{\partial\zeta}{\partial t} - \frac{\partial z_0}{\partial t} = 0 \end{aligned}$$

最后得

$$\frac{\partial H}{\partial t} + \frac{\partial HU_x}{\partial x} + \frac{\partial HU_y}{\partial y} = 0 \quad (2.12)$$

也可以写成

$$\frac{\partial H}{\partial t} + \nabla \cdot (H\mathbf{U}) = 0 \quad (2.13)$$

2.3.3 沿水深平均的运动方程

以 x 方向为例，式(2.2b)沿水深积分为

$$\int_{z_0}^{\zeta} \left[\frac{\partial \overline{u_x}}{\partial t} + \frac{\partial(\overline{u_x u_x})}{\partial x} + \frac{\partial(\overline{u_x u_y})}{\partial y} + \frac{\partial(\overline{u_x u_z})}{\partial z} + \frac{1}{\rho} \frac{\partial \bar{p}}{\partial x} - \nu_t \left(\frac{\partial^2 \overline{u_x}}{\partial x^2} + \frac{\partial^2 \overline{u_x}}{\partial y^2} + \frac{\partial^2 \overline{u_x}}{\partial z^2} \right) \right] dz = 0 \quad (2.14)$$

式(2.14)中包含了非恒定流项积分、对流项积分、压力项积分和阻力项积分。接下来分项讨论。

(1) 非恒定流项积分

$$\begin{aligned} \int_{z_0}^{\zeta} \frac{\partial \bar{u}_x}{\partial t} dz &= \frac{\partial}{\partial t} \int_{z_0}^{\zeta} \bar{u}_x dz - \bar{u}_x|_{z=\zeta} \frac{\partial \zeta}{\partial t} + \bar{u}_x|_{z=z_0} \frac{\partial z_0}{\partial t} \\ &= \frac{\partial H U_x}{\partial t} - \bar{u}_x|_{z=\zeta} \frac{\partial \zeta}{\partial t} + \bar{u}_x|_{z=z_0} \frac{\partial z_0}{\partial t} \end{aligned} \quad (2.15)$$

(2) 对流项积分

首先将时均流速按照式(2.16)进行分解

$$\begin{aligned} \bar{u}_x &= U_x + \Delta \bar{u}_x \\ \bar{u}_y &= U_y + \Delta \bar{u}_y \end{aligned} \quad (2.16)$$

式中, $\Delta \bar{u}_x$ 和 $\Delta \bar{u}_y$ 分别为 x 和 y 的时均流速与垂线平均流速的差值。

对流项中第一项的积分

$$\int_{z_0}^{\zeta} \frac{\partial \bar{u}_x \bar{u}_x}{\partial x} dz = \frac{\partial}{\partial x} \int_{z_0}^{\zeta} \bar{u}_x \bar{u}_x dz - \bar{u}_x \bar{u}_x|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \bar{u}_x \bar{u}_x|_{z=z_0} \frac{\partial z_0}{\partial x} \quad (2.17)$$

式中

$$\begin{aligned} \int_{z_0}^{\zeta} \bar{u}_x \bar{u}_x dz &= \int_{z_0}^{\zeta} (U_x + \Delta \bar{u}_x)(U_x + \Delta \bar{u}_x) dz \\ &= \int_{z_0}^{\zeta} (U_x U_x + 2U_x \Delta \bar{u}_x + \Delta \bar{u}_x \Delta \bar{u}_x) dz \\ &= H U_x U_x + \int_{z_0}^{\zeta} \Delta \bar{u}_x \Delta \bar{u}_x dz \\ &= \beta_{xx} H U_x U_x \end{aligned}$$

其中

$$\beta_{xx} = 1 + \frac{\int_{z_0}^{\zeta} \Delta \bar{u}_x \Delta \bar{u}_x dz}{H U_x U_x} \quad (2.18)$$

是由于流速沿垂线分布不均匀而引入的修正系数, 类似于水力学中的动量修正系数。 β_{xx} 的取值一般在 1.02 与 1.05 之间, 可近似取为 1.0。因此

$$\int_{z_0}^{\zeta} \frac{\partial \bar{u}_x \bar{u}_x}{\partial x} dz = \frac{\partial H U_x U_x}{\partial x} - \bar{u}_x \bar{u}_x|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \bar{u}_x \bar{u}_x|_{z=z_0} \frac{\partial z_0}{\partial x} \quad (2.19)$$

同理, 对流项的第二项可写为

$$\int_{z_0}^{\zeta} \frac{\partial \bar{u}_x \bar{u}_y}{\partial y} dz = \frac{\partial H U_x U_y}{\partial y} - \bar{u}_x \bar{u}_y|_{z=\zeta} \frac{\partial \zeta}{\partial y} + \bar{u}_x \bar{u}_y|_{z=z_0} \frac{\partial z_0}{\partial y} \quad (2.20)$$

对流项的第三项

$$\int_{z_0}^{\zeta} \frac{\partial \bar{u}_x \bar{u}_z}{\partial z} dz = \bar{u}_x \bar{u}_z|_{z=\zeta} - \bar{u}_x \bar{u}_z|_{z=z_0} \quad (2.21)$$

将非恒定流项和对流项积分相加, 并利用自由表面和河床底部运动学条件可得:

$$\int_{z_0}^{\zeta} \left[\frac{\partial \bar{u}_x}{\partial t} + \frac{\partial (\bar{u}_x \bar{u}_x)}{\partial x} + \frac{\partial (\bar{u}_x \bar{u}_y)}{\partial y} + \frac{\partial (\bar{u}_x \bar{u}_z)}{\partial z} \right] dz = \frac{\partial H U_x}{\partial t} + \frac{\partial H U_x U_x}{\partial x} + \frac{\partial H U_x U_y}{\partial y} \quad (2.22)$$

(3) 压力项积分

$$\begin{aligned} \int_{z_0}^{\zeta} \frac{\partial \bar{p}}{\partial x} dz &= \frac{\partial}{\partial x} \int_{z_0}^{\zeta} \bar{p} dz - \bar{p}|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \bar{p}|_{z=z_0} \frac{\partial z_0}{\partial x} \\ &= \frac{\partial}{\partial x} \int_{z_0}^{\zeta} \rho g (\zeta - z) dz - \rho g (\zeta - z)|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \rho g (\zeta - z)|_{z=z_0} \frac{\partial z_0}{\partial x} \\ &= \rho g H \frac{\partial H}{\partial x} + \rho g H \frac{\partial z_0}{\partial x} = \rho g H \frac{\partial \zeta}{\partial x} \end{aligned} \quad (2.23)$$

(4) 阻力项积分

$$\begin{aligned}
\int_{z_0}^{\zeta} \frac{\partial^2 \bar{u}_x}{\partial x^2} dz &= \int_{z_0}^{\zeta} \frac{\partial}{\partial x} \left(\frac{\partial \bar{u}_x}{\partial x} \right) dz = \frac{\partial}{\partial x} \int_{z_0}^{\zeta} \frac{\partial \bar{u}_x}{\partial x} dz - \frac{\partial \bar{u}_x}{\partial x} \Big|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \frac{\partial \bar{u}_x}{\partial x} \Big|_{z=z_0} \frac{\partial z_0}{\partial x} \\
&= \frac{\partial}{\partial x} \left(\frac{\partial}{\partial x} \int_{z_0}^{\zeta} \bar{u}_x dz - \bar{u}_x|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \bar{u}_x|_{z=z_0} \frac{\partial z_0}{\partial x} \right) - \frac{\partial \bar{u}_x}{\partial x} \Big|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \frac{\partial \bar{u}_x}{\partial x} \Big|_{z=z_0} \frac{\partial z_0}{\partial x} \\
&= \frac{\partial^2 H U_x}{\partial x^2} + \frac{\partial}{\partial x} \left(-\bar{u}_x|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \bar{u}_x|_{z=z_0} \frac{\partial z_0}{\partial x} \right) - \frac{\partial \bar{u}_x}{\partial x} \Big|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \frac{\partial \bar{u}_x}{\partial x} \Big|_{z=z_0} \frac{\partial z_0}{\partial x}
\end{aligned} \quad (2.24)$$

类似地

$$\int_{z_0}^{\zeta} \frac{\partial^2 \bar{u}_x}{\partial y^2} dz = \frac{\partial^2 H U_x}{\partial y^2} + \frac{\partial}{\partial y} \left(-\bar{u}_x|_{z=\zeta} \frac{\partial \zeta}{\partial y} + \bar{u}_x|_{z=z_0} \frac{\partial z_0}{\partial y} \right) - \frac{\partial \bar{u}_x}{\partial y} \Big|_{z=\zeta} \frac{\partial \zeta}{\partial y} + \frac{\partial \bar{u}_x}{\partial y} \Big|_{z=z_0} \frac{\partial z_0}{\partial y} \quad (2.25)$$

$$\int_{z_0}^{\zeta} \frac{\partial^2 \bar{u}_x}{\partial y^2} dz = \frac{\partial \bar{u}_x}{\partial z} \Big|_{z=\zeta} - \frac{\partial \bar{u}_x}{\partial z} \Big|_{z=z_0} \quad (2.26)$$

将式(2.24)至(2.26)相加, 得

$$\begin{aligned}
&\int_{z_0}^{\zeta} \nu_t \left(\frac{\partial^2 \bar{u}_x}{\partial x^2} + \frac{\partial^2 \bar{u}_x}{\partial y^2} + \frac{\partial^2 \bar{u}_x}{\partial z^2} \right) dz = \\
&\nu_t \left(\frac{\partial^2 H U_x}{\partial x^2} + \frac{\partial^2 H U_x}{\partial y^2} \right) \\
&- \nu_t \left[\frac{\partial}{\partial x} \left(\frac{\partial \zeta}{\partial x} \bar{u}_x|_{z=\zeta} \right) + \frac{\partial}{\partial y} \left(\frac{\partial \zeta}{\partial y} \bar{u}_x|_{z=\zeta} \right) + \frac{\partial \bar{u}_x}{\partial x} \Big|_{z=\zeta} \frac{\partial \zeta}{\partial x} + \frac{\partial \bar{u}_x}{\partial y} \Big|_{z=\zeta} \frac{\partial \zeta}{\partial y} - \frac{\partial \bar{u}_x}{\partial z} \Big|_{z=\zeta} \right] \\
&+ \nu_t \left[\frac{\partial}{\partial x} \left(\frac{\partial z_0}{\partial x} \bar{u}_x|_{z=z_0} \right) + \frac{\partial}{\partial y} \left(\frac{\partial z_0}{\partial y} \bar{u}_x|_{z=z_0} \right) + \frac{\partial \bar{u}_x}{\partial x} \Big|_{z=z_0} \frac{\partial z_0}{\partial x} + \frac{\partial \bar{u}_x}{\partial y} \Big|_{z=z_0} \frac{\partial z_0}{\partial y} - \frac{\partial \bar{u}_x}{\partial z} \Big|_{z=z_0} \right]
\end{aligned} \quad (2.27)$$

式(2.27)中右边后两项分别为由底部床面阻力和自由表面风阻力引起的阻力项, 通常可以式(2.28)表示:

$$g \frac{n^2 U_x \sqrt{U_x^2 + U_y^2}}{H^{1/3}} - C_w \frac{\rho_a}{\rho} \omega^2 \cos \beta \quad (2.28)$$

式中, C_w 为无因此风应力系数, ρ_a 为空气密度, ω 为风速, β 为风向与 x 方向的夹角。最后, x 方向的运动方程为

$$\begin{aligned}
&\frac{\partial H U_x}{\partial t} + \frac{\partial H U_x U_x}{\partial x} + \frac{\partial H U_x U_y}{\partial y} = \\
&- g H \frac{\partial \zeta}{\partial x} - g \frac{n^2 U_x \sqrt{U_x^2 + U_y^2}}{H^{1/3}} + \nu_t \left(\frac{\partial^2 H U_x}{\partial x^2} + \frac{\partial^2 H U_x}{\partial y^2} \right) + C_w \frac{\rho_a}{\rho} \omega^2 \cos \beta
\end{aligned} \quad (2.29)$$

同理可得, y 方向运动方程为

$$\begin{aligned}
&\frac{\partial H U_y}{\partial t} + \frac{\partial H U_x U_y}{\partial x} + \frac{\partial H U_y U_y}{\partial y} = \\
&- g H \frac{\partial \zeta}{\partial y} - g \frac{n^2 U_y \sqrt{U_x^2 + U_y^2}}{H^{1/3}} + \nu_t \left(\frac{\partial^2 H U_y}{\partial x^2} + \frac{\partial^2 H U_y}{\partial y^2} \right) + C_w \frac{\rho_a}{\rho} \omega^2 \sin \beta
\end{aligned} \quad (2.30)$$

当自由表面风应力影响较小时, 风应力项可以忽略。此外, 当模拟区域尺度较大, 还要考虑地球自转的影响, 可在方程(2.29)和(2.30)右边分别加入科氏力。

$$\begin{aligned}
f_x &= 2\omega \sin \varphi U_x \\
f_y &= 2\omega \sin \varphi U_y
\end{aligned} \quad (2.31)$$

式中, ω 为地球自转角速度, φ 为模拟区域所处纬度。

根据推导过程中所采用的假定条件, 在使用二维浅水方程时应注意以下问题:

1. 方程推导中引用了牛顿流体所满足的本构关系式, 因此上述方程只适用于牛顿流体, 对类似高含沙水流的非牛顿流体不适用。
2. 方程推导中对流体做了均质不可压的假设, 因此上述方程只能在含沙量较小的情况下近似使用, 当含沙量较大时, 应考虑密度变化的影响。
3. 在垂向积分过程中, 略去流速等水力参数沿垂直方向的变化, 并假定沿水深方向的动水压强分布符合静水压强分布。因此所研究问题的水平尺度应远大于垂向尺度, 流速等水力参数沿垂直方向的变化较之沿水平方向的变化要小得多。

2.3.4 二维平面浅水方程不同形式

2.4 一维非恒定流基本控制方程

2.4.1 一维连续性方程

如图??所示, 在明槽非恒定流中, 沿水流流动方向取长为 dx 的微小流段。流段进口 1-1 断面流量为 Q , 在 Δt 时段内, 从 1-1 断面进入流段的液体质量为 $\rho Q \Delta t$ 。流段出口 2-2 断面流量为 $Q + \frac{\partial Q}{\partial x} dx$, 在 Δt 时段内, 从 2-2 断面流出的液体质量为 $\rho Q \Delta t + \rho \frac{\partial Q}{\partial x} dx \Delta t$ 。该时段内进出此流段的液体质量差为 $-\rho \frac{\partial Q}{\partial x} dx \Delta t$ 。

时段内的质量差表现为流段内的槽蓄量变化。在起始时刻, 流段内的槽蓄量为 $\rho \bar{A} dx$ 。而经过 Δt 时段后, 流段内的槽蓄量为 $\rho \left(\bar{A} + \frac{\partial \bar{A}}{\partial t} \Delta t \right) dx$ 。 Δt 时段内, 流段内的槽蓄量变化量为 $\rho \frac{\partial \bar{A}}{\partial t} \Delta t dx$ 。其中 \bar{A} 为微小流段的平均过水面积。当流段内过水断面面积变化较小时, 可直接用 1-1 断面段面积 A 来替代 \bar{A} 。

因此, 根据质量守恒原理, 进出该流段的液体质量差等于流段内槽蓄量改变量, 即

$$-\rho \frac{\partial Q}{\partial x} dx \Delta t = \rho \frac{\partial A}{\partial t} dx \Delta t$$

化简后得到明槽一维非恒定流连续性方程

$$\frac{\partial A}{\partial t} + \frac{\partial Q}{\partial x} = 0 \quad (2.32)$$

2.4.2 一维运动方程

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