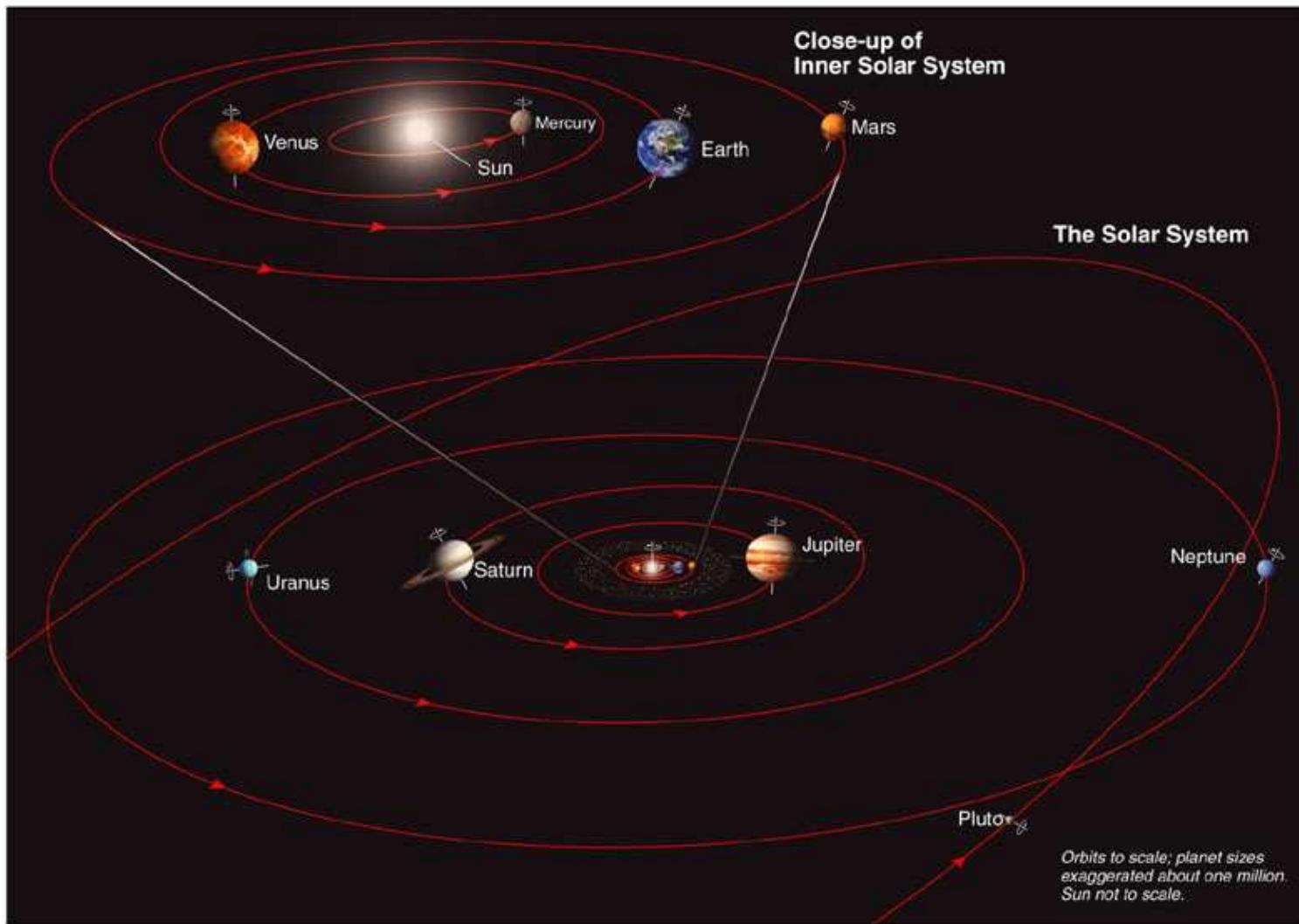


Solar System overview

- 1) inventory
- 2) spin/orbit/shape
- 3) heated by the Sun
- 4) how do we find out



Inventory

1 star
(99.9% of M)
8 planets
(99.9% of L)

- Terrestrial:

Mercury
Venus
Earth
Mars

- Giant:

Jupiter
Saturn
Uranus
Neptune

Lots of small bodies
incl. dwarf planets

Ceres
Pluto
Eris

Inventory (cont'd)

Many moons & rings

Mercury: 0

Venus: 0

Earth: 1 (1700km)

Mars: 2 (~10km)

Jupiter: 63 + rings

Saturn: 60 + rings

Uranus: 27 + rings

Neptune: 13 + rings

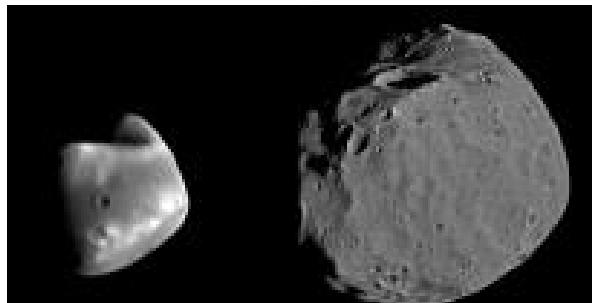
Even among dwarf planets, asteroids, Kuiper belt objects, and comets. E.g.,

Pluto: 3

Eris: 1

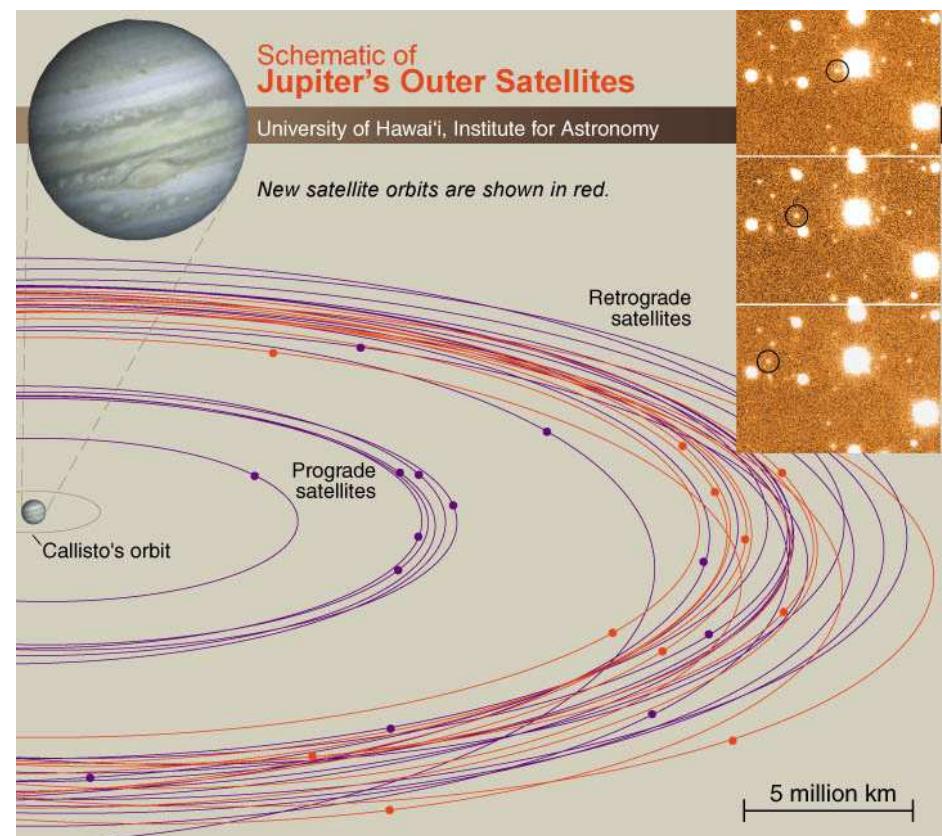
Moons of Mars:

Deimos & Phobos, ~10km

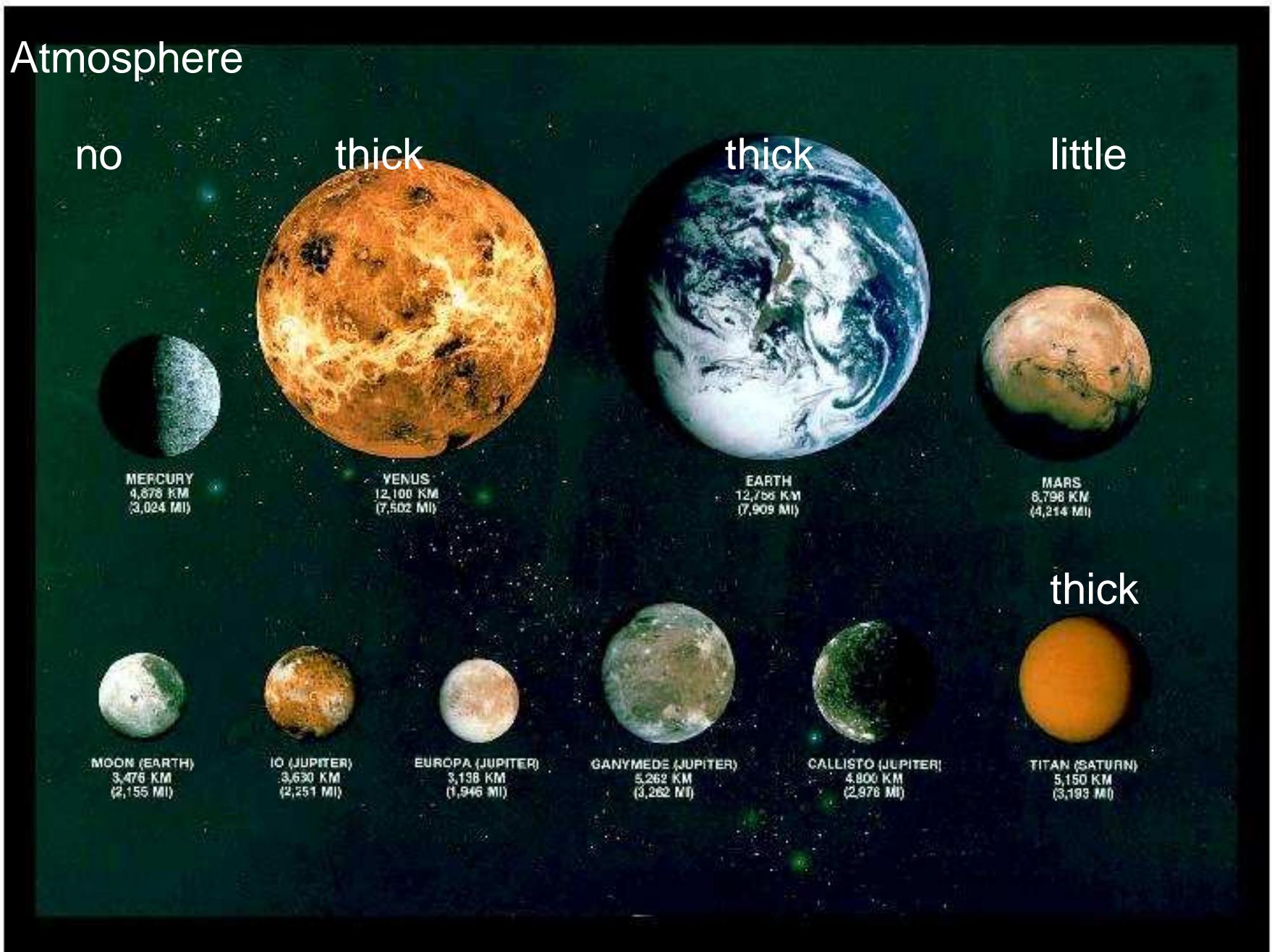


Moons of Jupiter

4 Galilean satellites (Ganymede, Callisto, Io & Europa),
~ 10^3 km (close to Jupiter, likely primordial)



Atmosphere



Inventory (cont'd)

~ 10^5 known small objects in the

- **Asteroid belt**
(Ceres ~300 km)

- **Kuiper belt**
(Eris, Pluto, Sedna, Quaoar, ~1000 km)

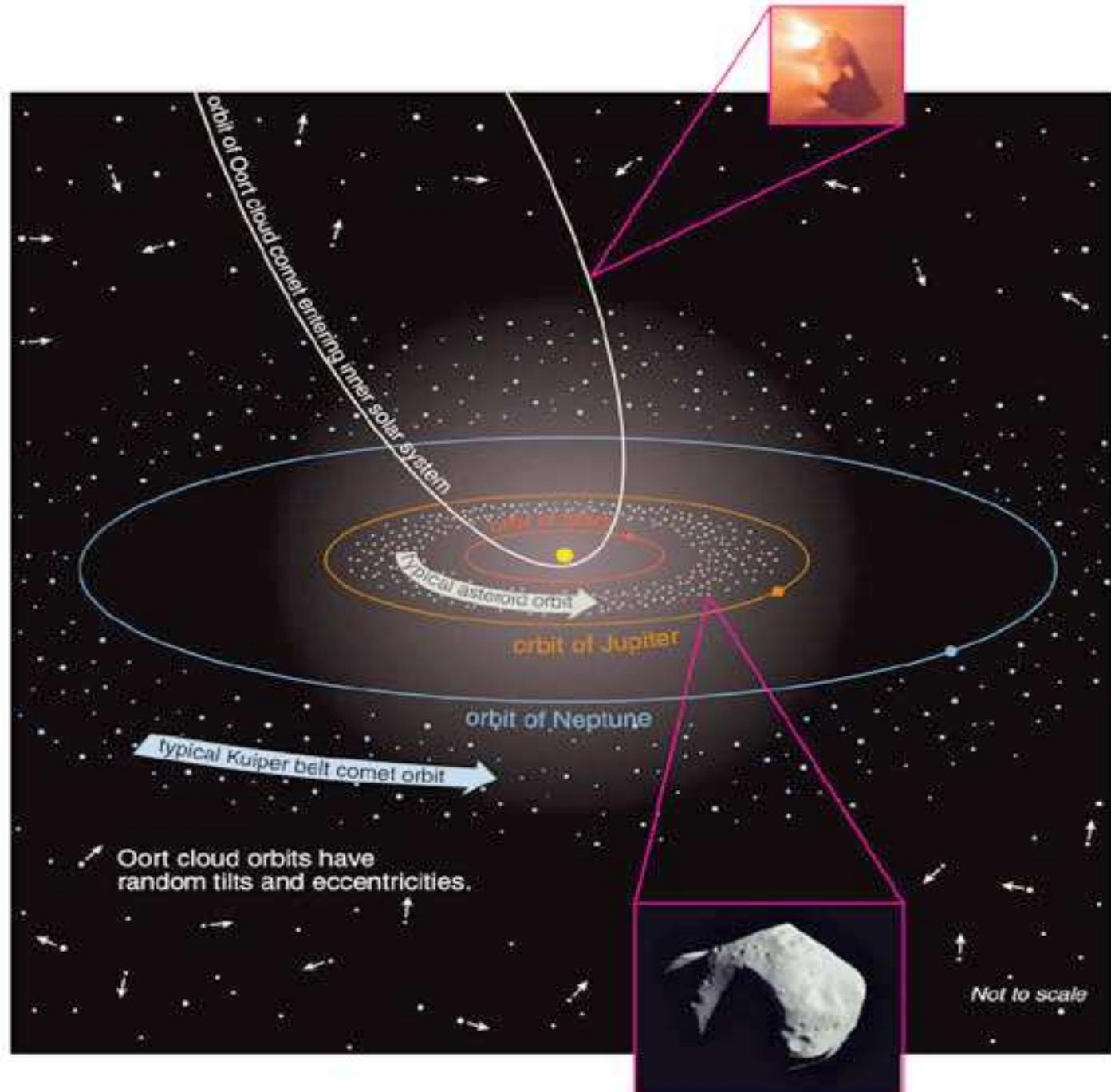
Estimated: ~ 10^{12} comets in the

- **Oort cloud**
(~ 10^4 AU)

Associated:

- **zodiacal dust**

(fire-works on the sky:
comets & meteorites)



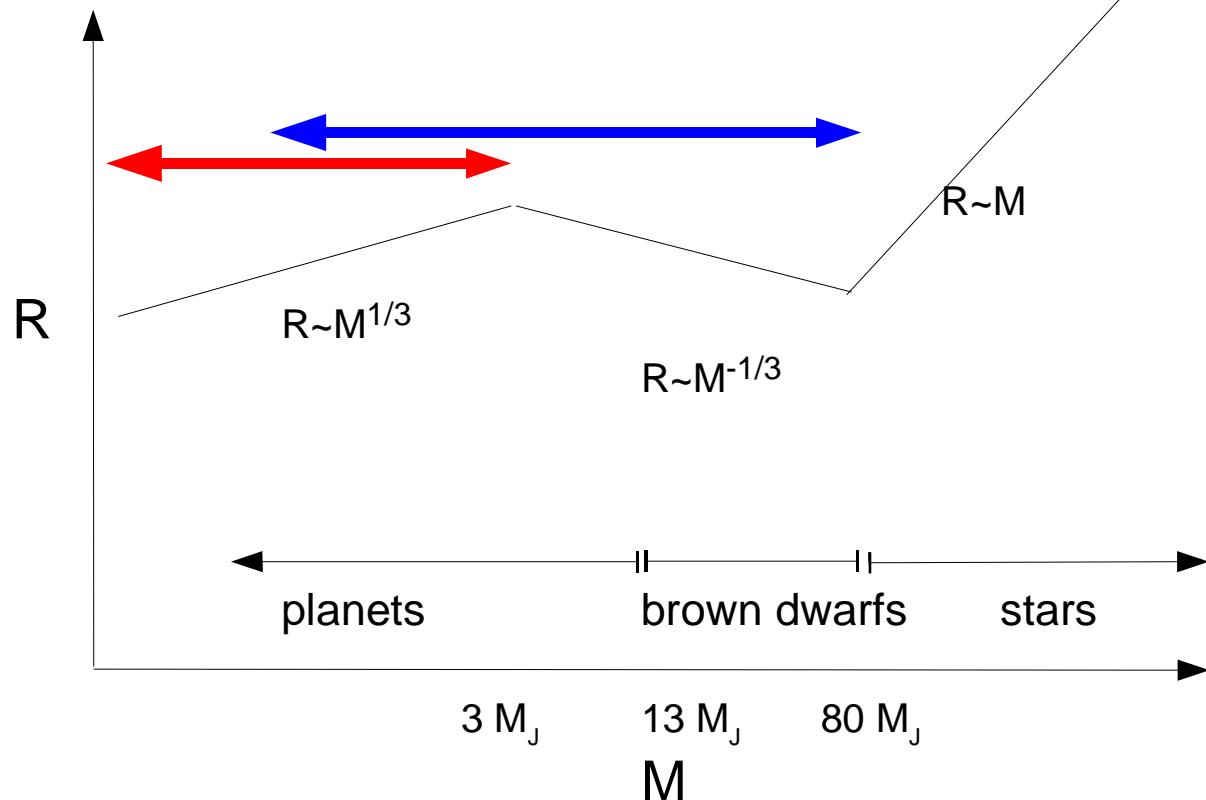
What are planets?

IAU (for solar system):

Orbits Sun, massive enough to be round and to have cleared its neighbourhood.

More general:

- 1) no nuclear fusion (not even deuterium): $T_c < 10^6 \text{ K}$
- 2) pressure provided by **electron degeneracy** and/or **Coulomb force**
($\rho \sim h/p \sim d$) ($d \sim \text{atomic radius}$)
- 3) can be solid or gaseous (with solid cores) --- *similar density*



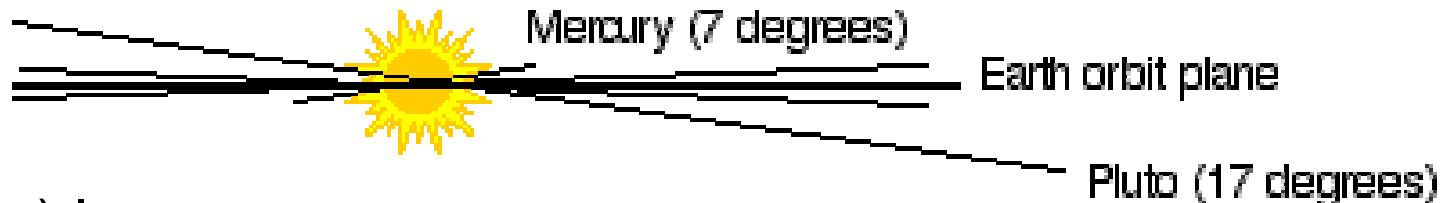
Mass & Mean ρ		
	M_J	[g/cm ³]
Jupiter	1.0	1.33
Saturn	0.3	0.77
Neptune	0.05	1.67
Uranus	0.04	1.24
Earth	0.003	5.52
Venus	0.002	5.25
Mars	0.0003	3.93
Mercury	0.0002	5.43

Orbits

inclination: largely coplanar (history)

direction: all the same

eccentricity: a few percent (except for Mercury)



Titus-Bode ([fitting](#)) law (1766)

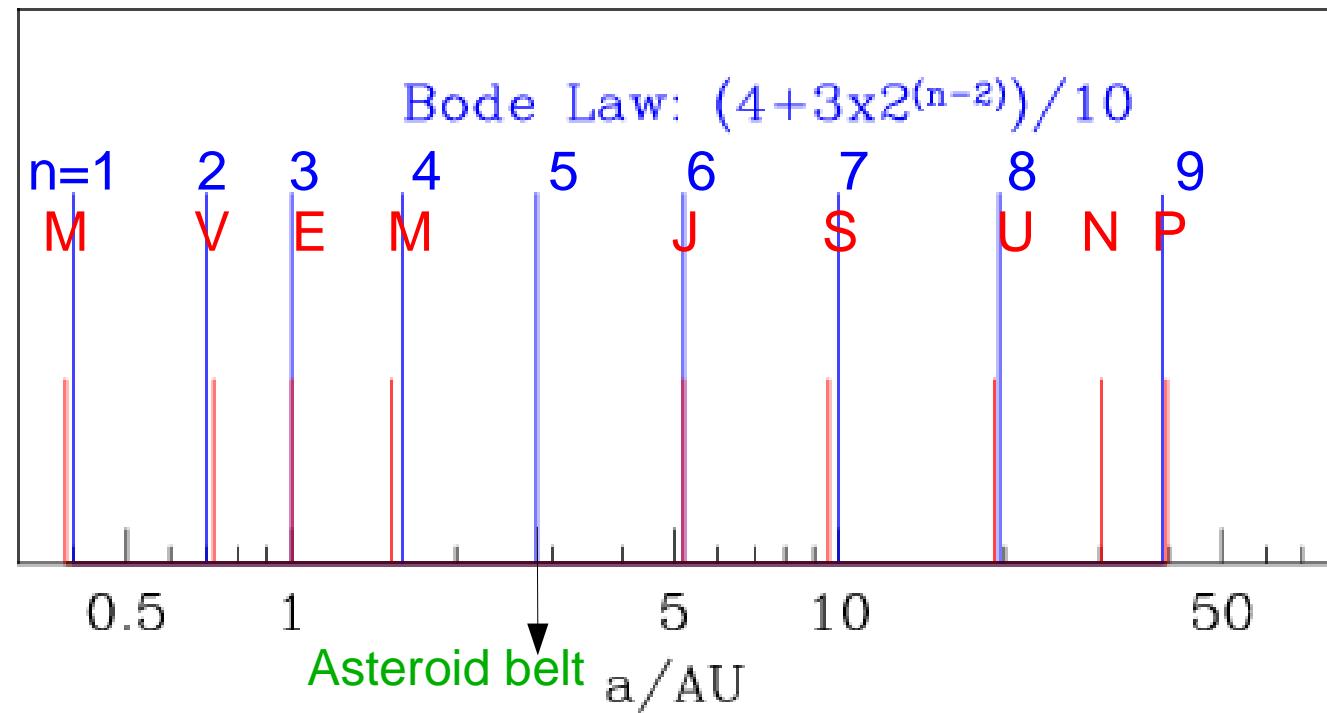
planetary orbits appear to (almost) satisfy a single relation

'Predict' the existence of the asteroid belt (1801: Ceres discovered)

coincidence or something deeper?

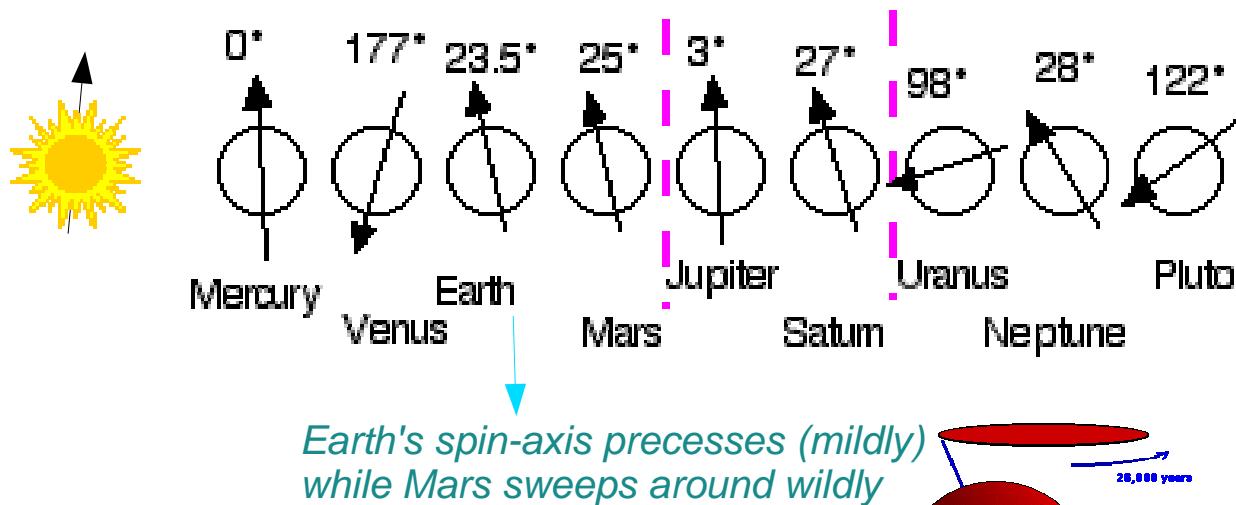
other systems?

*computer simulations
indicate that planets are as
maximally packed as allowed
by stability*



Spin (obliquity)

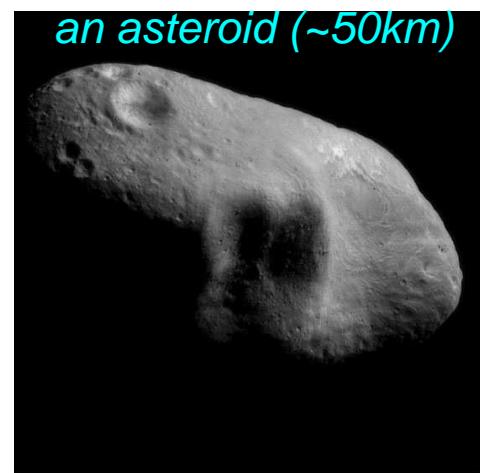
smaller planets:
almost random, affected by
impacts and giant planets



Shape --- *the bigger the rounder*

All gaseous planets are spherical.

Large rocky objects are rather spherical. Smaller ones are less so.



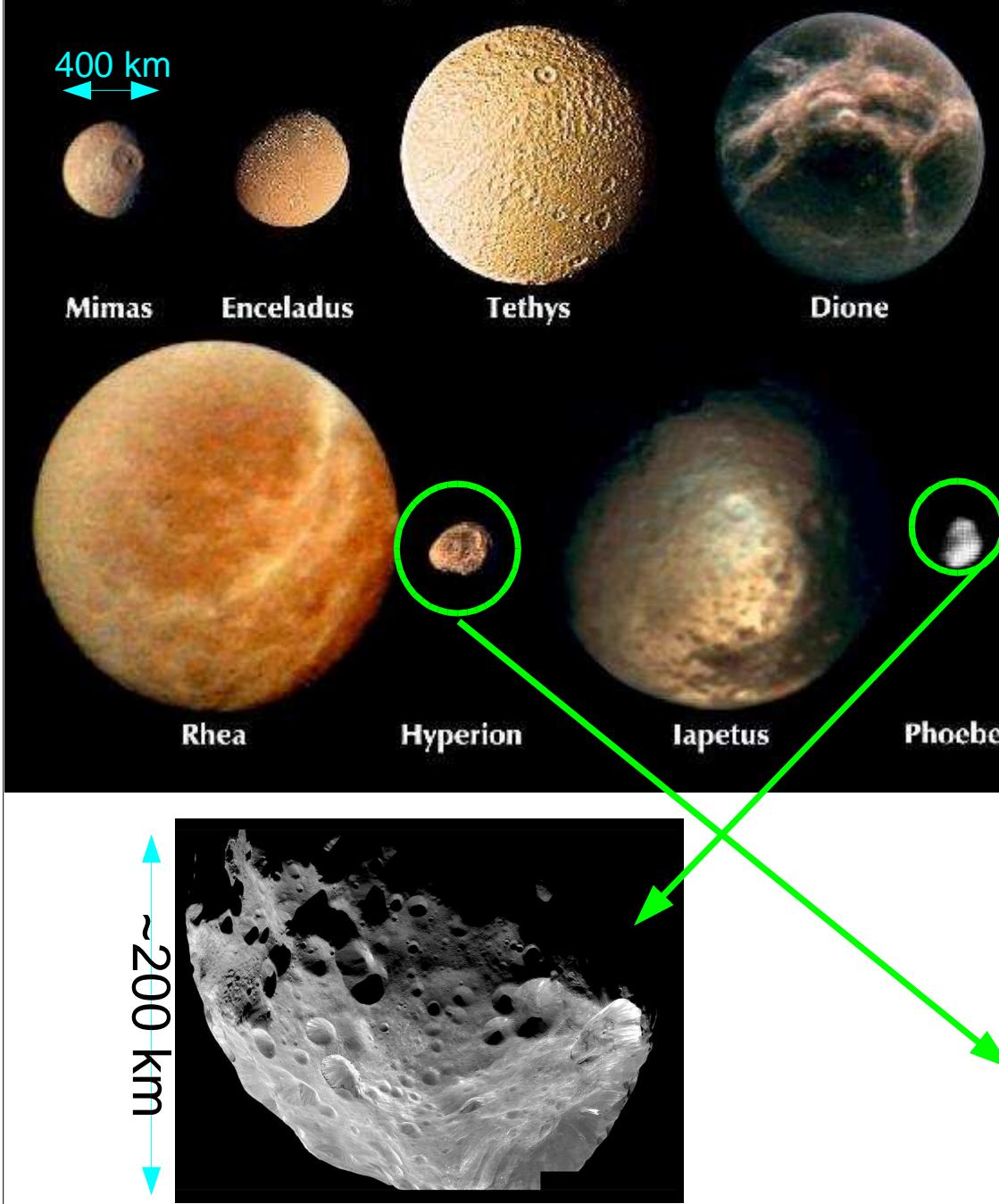
	h (km)	R (km)	$g = GM/R^2$ (m/s ²)
Earth	8	6400	9.8
Mars	24	3400	3.7

scaling: highest mountain on Earth ~8 km (on Mars ~ 24 km) $h * g \sim \text{constant}$

rough estimate: irregular body has mountain $h \sim R \implies R \sim 240$ km

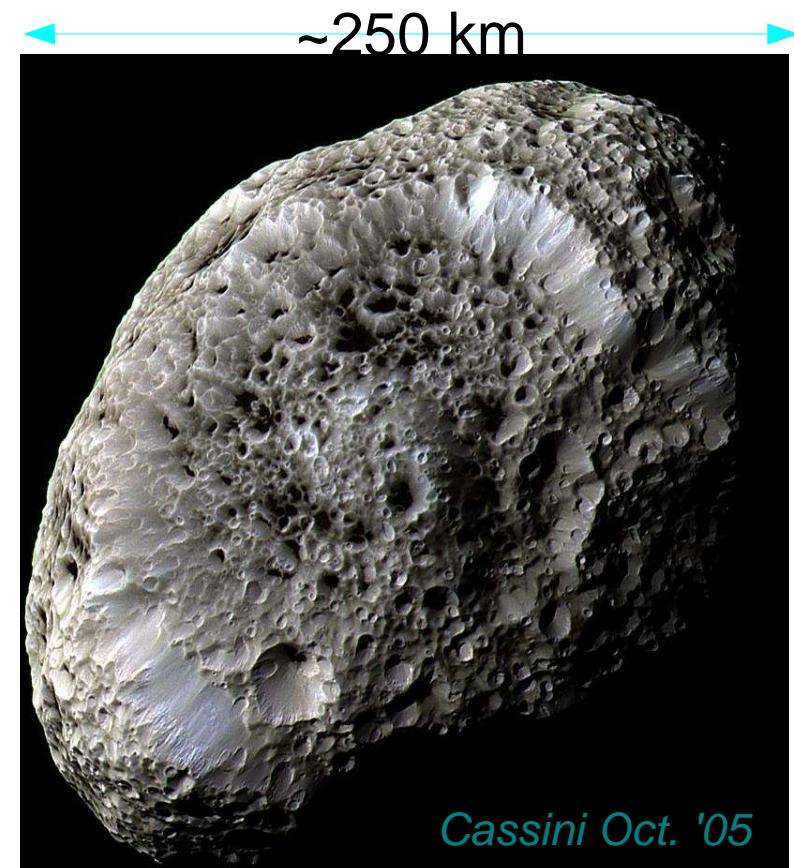
thus: objects with $R > 240$ km are approximately spherical

Saturn's Eight Major Icy Satellites



The bigger the rounder

	$\Delta R/R$	$g=GM/R^2$
Earth	8/6400	9.8 m/s ²
Mars	24/3400	3.7 m/s ²
Hyperion: 150/250		~0.4 m/s ²



Passively Heated by the Sun --- *the further the cooler*

Typically we observe objects in reflected light, however, all objects emit re-processed thermal radiation which is observable at longer wavelengths.

Blackbody temperature for a non-self-luminous spherical body at distance a away from the Sun (with albedo A -- reflectivity)

$$L_{\text{abs}} = (1-A) \frac{\pi R_p^2}{4\pi a^2} 4\pi R_s^2 \sigma T_s^4; \quad L_{\text{em}} = 4\pi R_p^2 \sigma T_p^4$$

If $L_{\text{abs}} = L_{\text{em}}$, then

$$T_p = \left(\frac{R_o}{2a} \right)^{1/2} T_s (1-A)^{1/4}$$

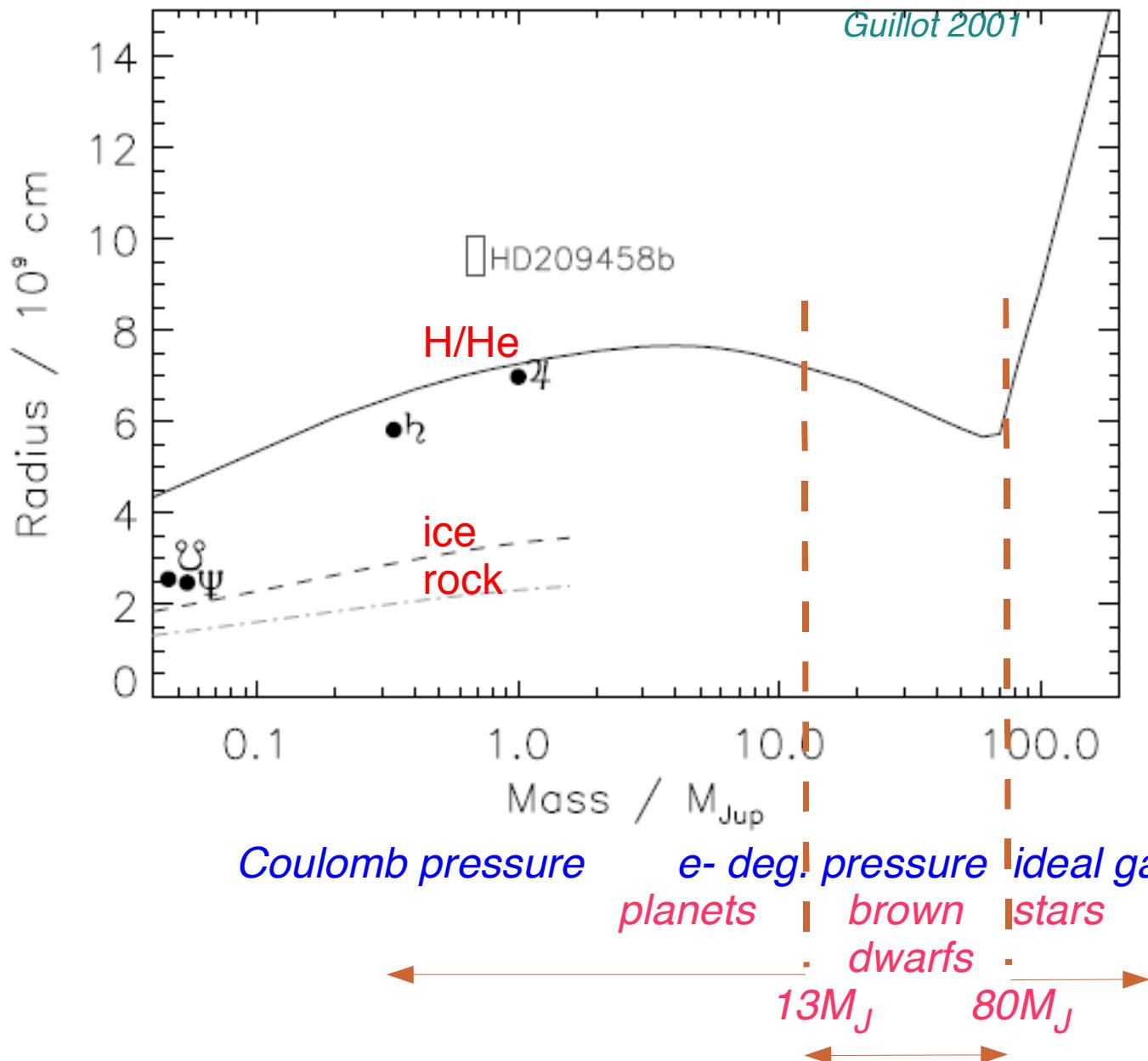
	a (AU)	A	T_{pred} (K)	T_{act} (K)	
Mercury	0.4	0.06	422 K	100-725	(?)
Venus	0.7	0.77	230K	733	(?)
Earth	1	0.30	255K	288	(?)
Mars	1.5	0.25	218K	223	good
Jupiter	5	0.51	113K	125	(?)
Saturn	9	0.47	83K	95	(?)
Uranus	19	0.51	60K	60	good
Neptune	30	0.62	40K	60	(?)
Comet at	5000	0.51	3.4K		

Giant Planets



made mostly of H, He and H-compounds, no solid surface
99.5% planet mass, 99.8% solar system angular momentum

Giant planets border stars



Equation of state determines mass-radius relation

Ideal gas: $P \propto \rho T$
+Fusion: $T_{\text{core}} \sim \text{Const}$
 $\rightarrow R \propto M$

el. degeneracy: $P \propto \rho^{5/3}$
 $\rightarrow R \propto M^{-1/3}$

Coulomb: $\rho \sim \text{Const}$
 $\rightarrow R \propto M^{1/3}$

Working definition:
Brown-dwarfs are 'failed' stars that cannot ignite hydrogen (but can burn deuterium); hence $M < 80 M_J$ ($0.08 M_\odot$)

Planets are formed in disks around stars.
Planets cannot burn deuterium ($10^6 K$); hence $M < 13 M_J$

Are planets just gas balls like stars? Probably not.

Jupiter & Saturn: largely degenerate H & He, mean $\rho = 1.3$ & 0.7 g/cm^3

- hydrogen metallic (conductive) below certain depth (?)

- core: solid, heavy metal + ices

Jupiter's core: $< 10 M_E$ (or 0?); Saturn's core: $\sim 13 M_E$ (15% of mass)

Uranus & Neptune: largely ices (H_2O , CH_4 , NH_3), mean $\rho = 1.2$ & 1.7 g/cm^3

- relatively thin gaseous H & He envelope

- mostly icy + rocky core

Why do we care about the solid cores?

Formation of giant planets likely starts with a solid core – unlike stars

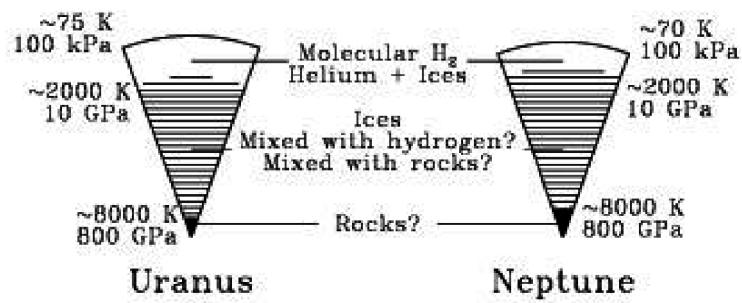
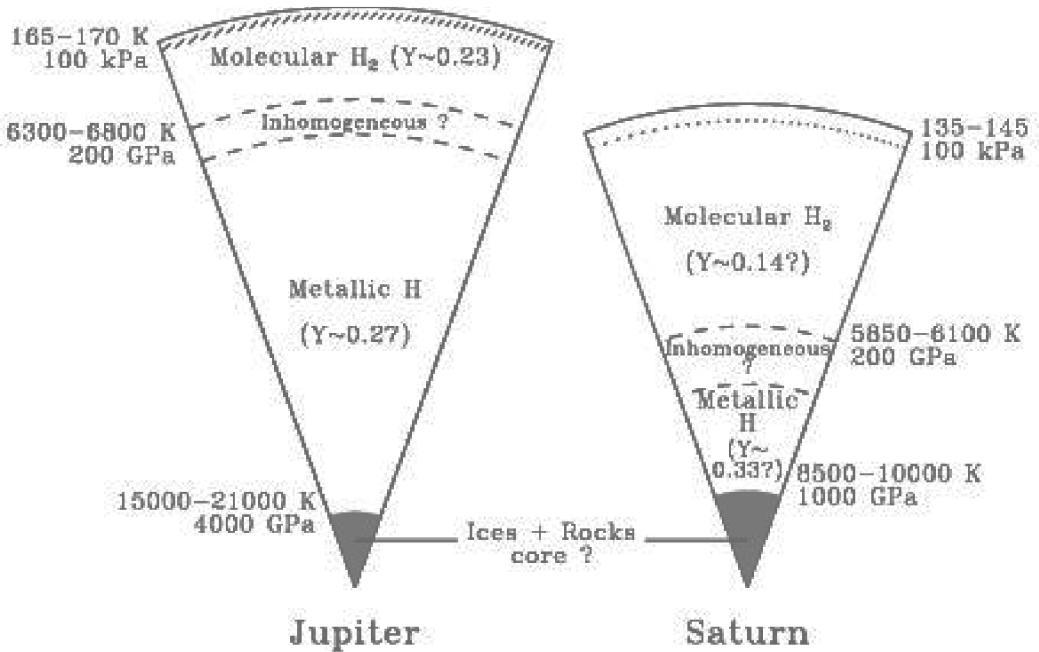
How do we figure out about the cores? Spin it!

core: a high density central region

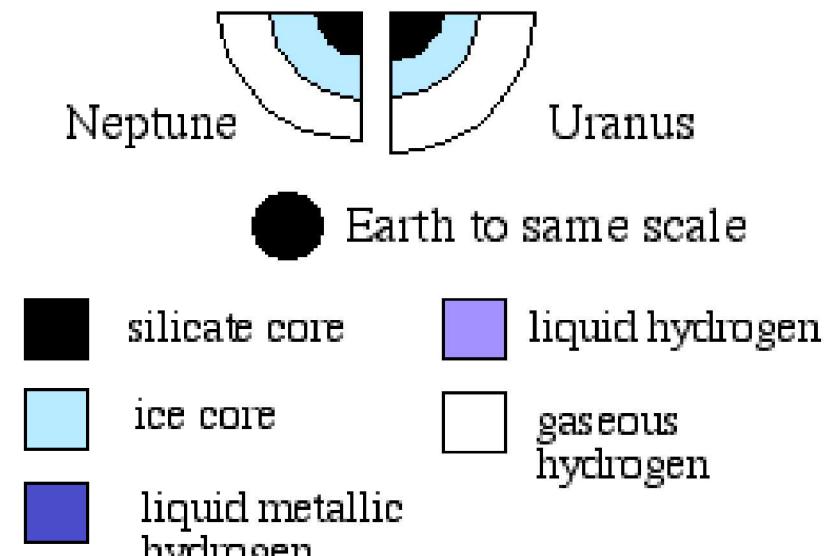
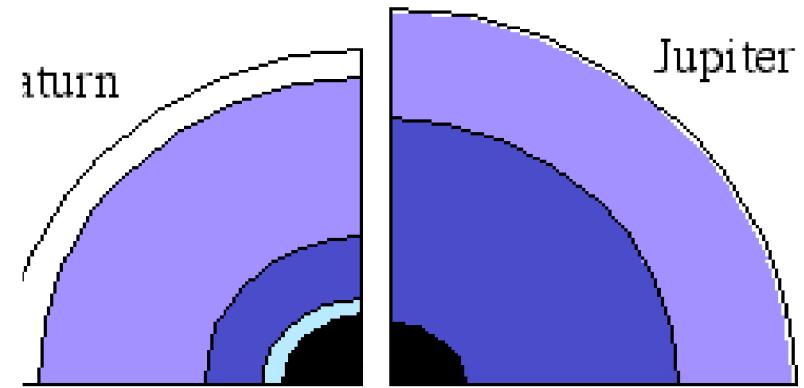
spherical body: gravitational potential is independent of density profile
but when the planet rotates, its oblateness depends on $\rho=\rho(r)$

$$\Phi(\theta) = -\frac{GM}{r} \left[1 - \left(\frac{R}{r} \right)^2 J_2 P_2(\cos \theta) - \left(\frac{R}{r} \right)^4 J_4 P_4(\cos \theta) - \dots \right]$$





Cores of giant planet are likely primordial and do not form by gravitational settling.
Did Jupiter melt part of its core?



Energy budget for giant planets

$$\text{Absorb solar flux: } (1-A)4\pi R_o^2 \sigma T_o^4 \times \frac{\pi R_p^2}{4\pi a^2}$$

$$\text{Emit blackbody flux: } 4\pi R_p^2 \sigma T_p^4$$

$$T_p = (1-A)^{1/4} \left(\frac{R_o}{2a} \right)^{1/2} T_o$$

	Jupiter	Saturn	Uranus	Neptune
passive T_p	113K	83K	60K	48K
actual T_p	130K	95K	59K	59K
$L_{\text{total}}/L_{\text{received}}$	1.7	1.8	1.0	2.6

3 sources of planetary intrinsic luminosity: primordial + settling + radio-active

Jupiter: **primordial heat** + He settling relative to H
(very long thermal time-scale: $\sim 10^9$ yrs)

Saturn: primordial heat + He settling relative to H

Uranus: no additional source required

Neptune: Do require add'l source; but so similar to Uranus, so why?

--- *what about gravitational contraction? No, already shrunk*

--- *terrestrial planets: radio-active elements*

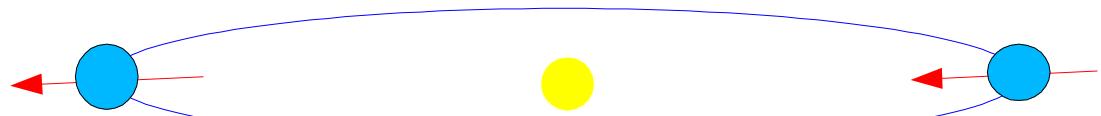
--- *how much energy can you gain by separating H & He?*

Other cool points?

1) magnetic fields: all 4 have appreciable B fields, Jovian aurorae,
Jupiter's magnetic influence extends past Saturn orbit
generation of these fields -- primordial or dynamo?

2) seasons:

Uranus: 97.92° inclined relative to orbit, very weird seasons!



3) rings & satellites: all 4 have rings and many satellites

rings: sandy or icy dust and some boulders, 2.5 planet radii (~Roche radius)

-- $H/R \sim 10^{-6}$ (*a razor blade?*)

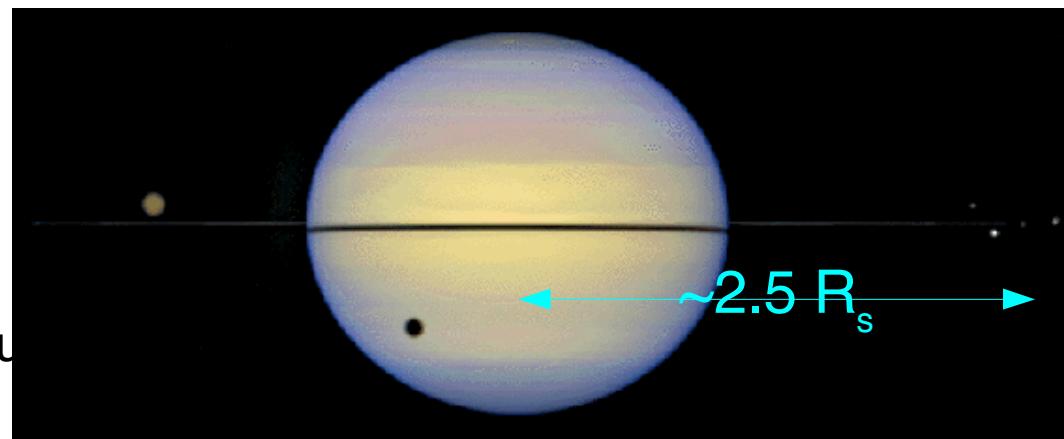
--- gaps: shepherding moons

--- origin: tidally disrupted satellites
or primordial?

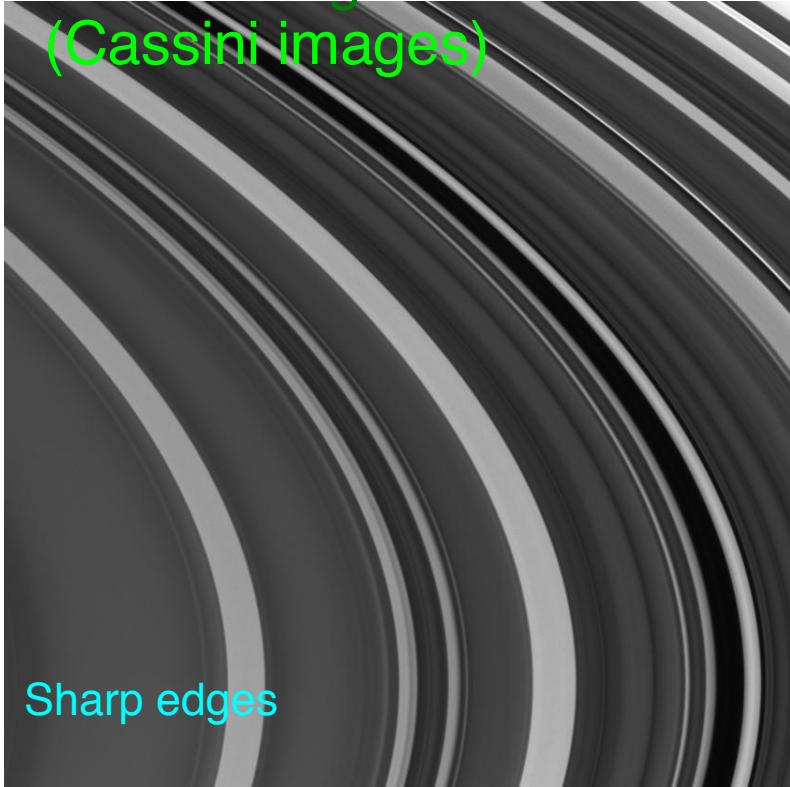
Satellites: worlds of their own
captured (Phoebe) or formed in-situ

Europa (@J): cracky surface
underground H_2O ocean

Titan (@S): smoggy atmosphere
surface H-compound ocean?



Saturn's rings
(Cassini images)

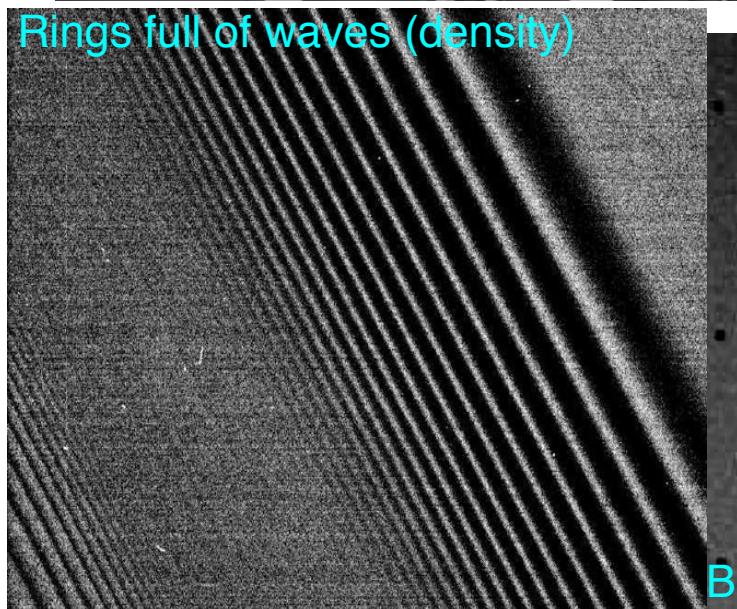


Sharp edges

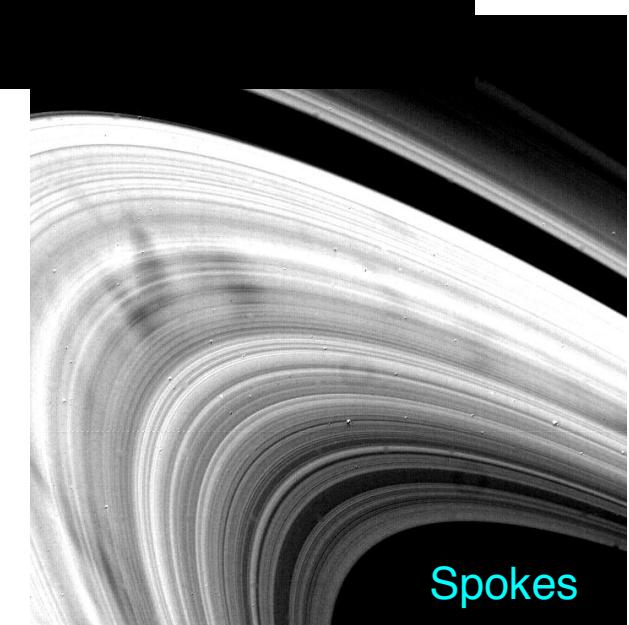
Prometheus shepherding



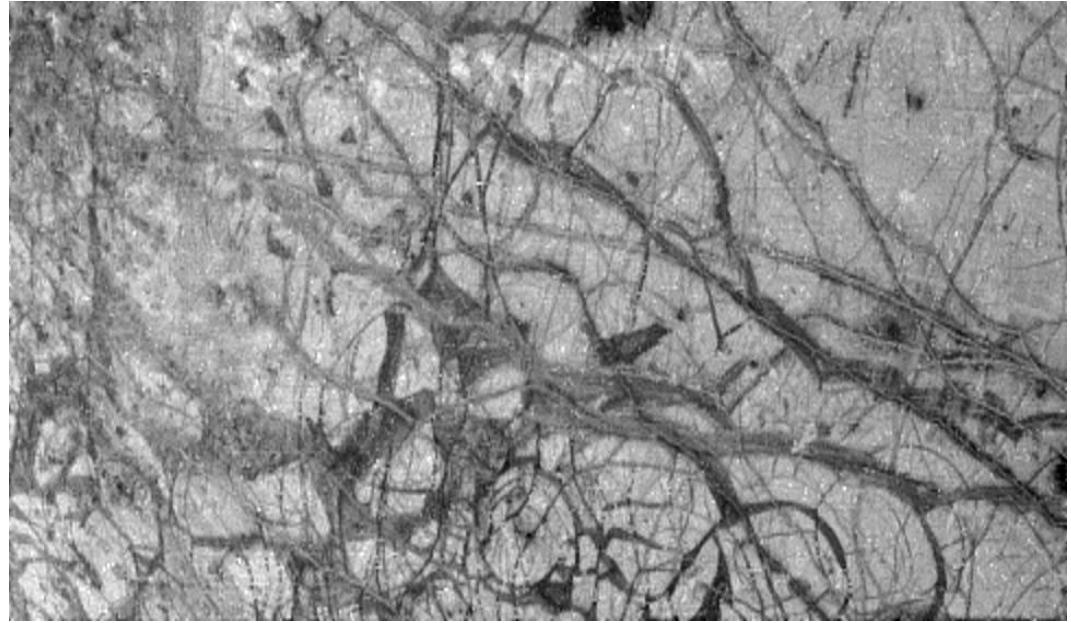
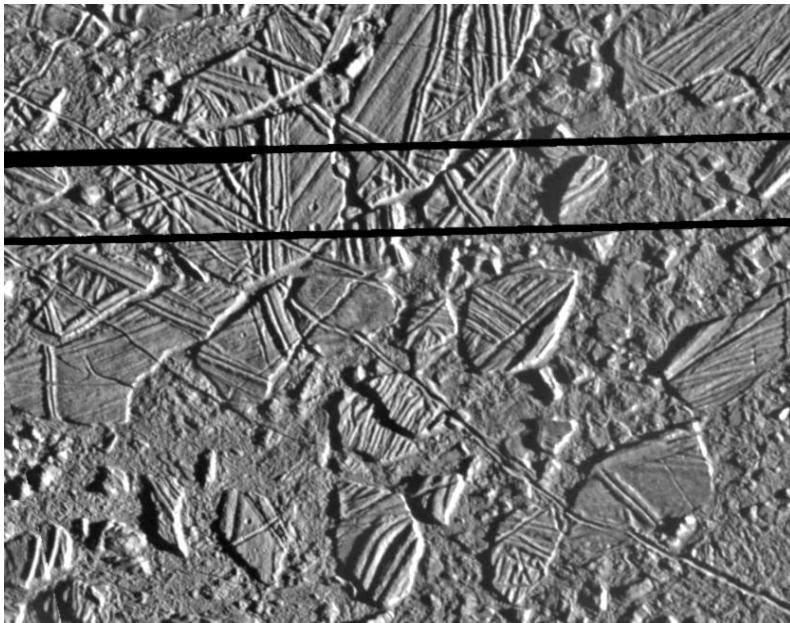
Rings full of waves (density)



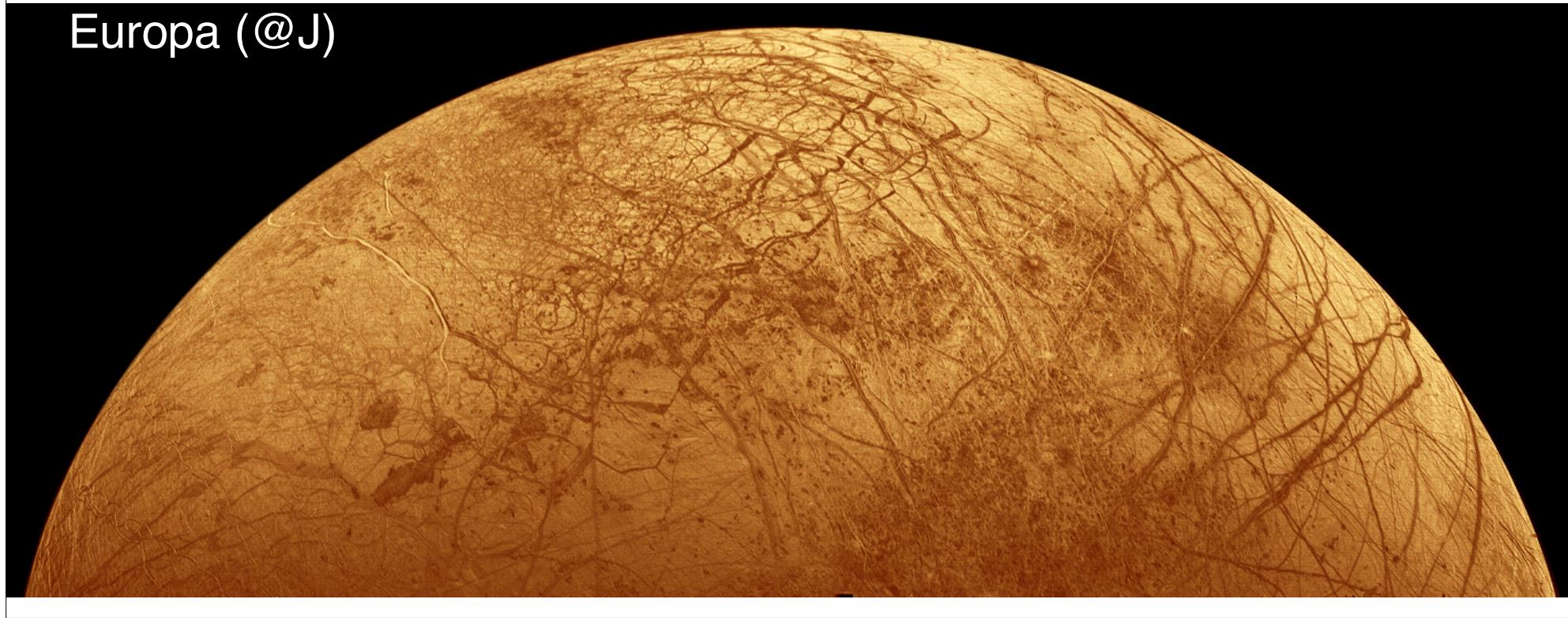
Braided ring



Spokes

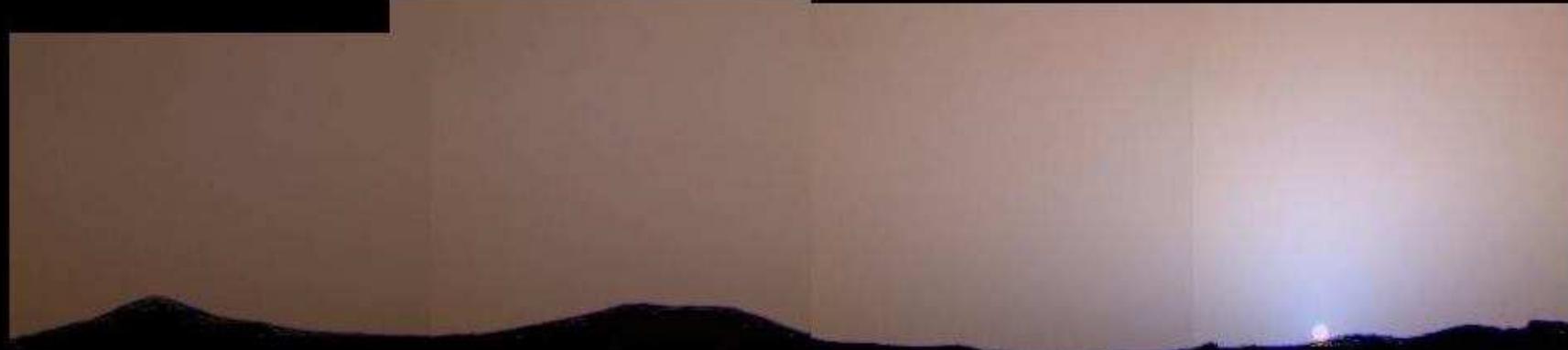


Europa (@J)



Planetary Atmospheres

- 1) Densities, temperatures
- 2) Origin of terrestrial planet atmospheres
- 3) Optics: colour, clouds
- 4) What happened to Venus?



Passively Heated by the Sun --- *the further the cooler*

Typically we observe objects in reflected light, however, all objects emit re-processed thermal radiation which is observable at longer wavelengths.

Blackbody temperature for a non-self-luminous spherical body at distance a away from the Sun (with albedo A -- reflectivity)

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If $L_{\text{abs}} = L_{\text{em}}$, then

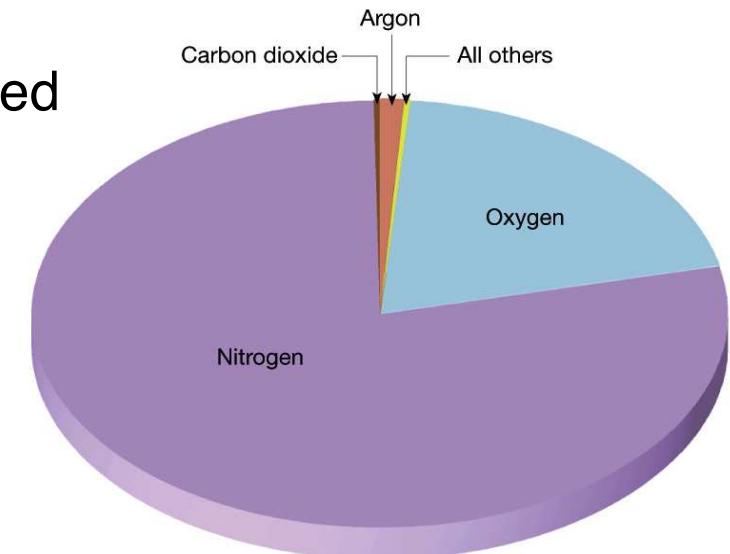
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Neptune	30	0.62	40K	60	(?)
Comet at	5000	0.51	3.4K		

Atmospheres: Terrestrial Planets

	Atm. Composition	surface pressure/T	
Mercury	--	< 10 ⁻¹² bar	100-725 K
Venus	97% CO ₂ , 3% N ₂	92 bar	733 K (460°C)
Earth	78% N ₂ , 21% O ₂ , 1% Ar	1 bar	288 K (15°C)
Mars	95% CO ₂ , 3% N ₂ , 1.6% Ar	0.006 bar	223 K (-50°C)
Titan (@S)	95% N ₂ , few% CH ₄ , Ar	1.5 bar	93 K (-180°C)

Most atmospheres are reasonably well-mixed
(no molecular weight separation)



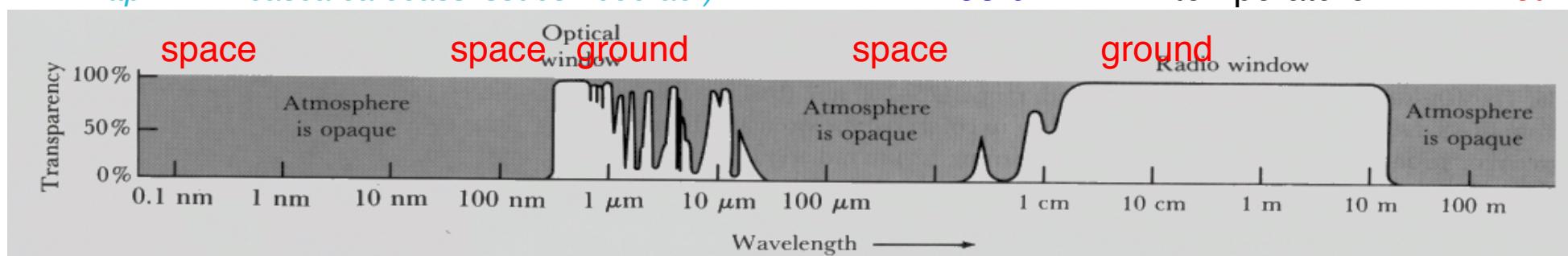
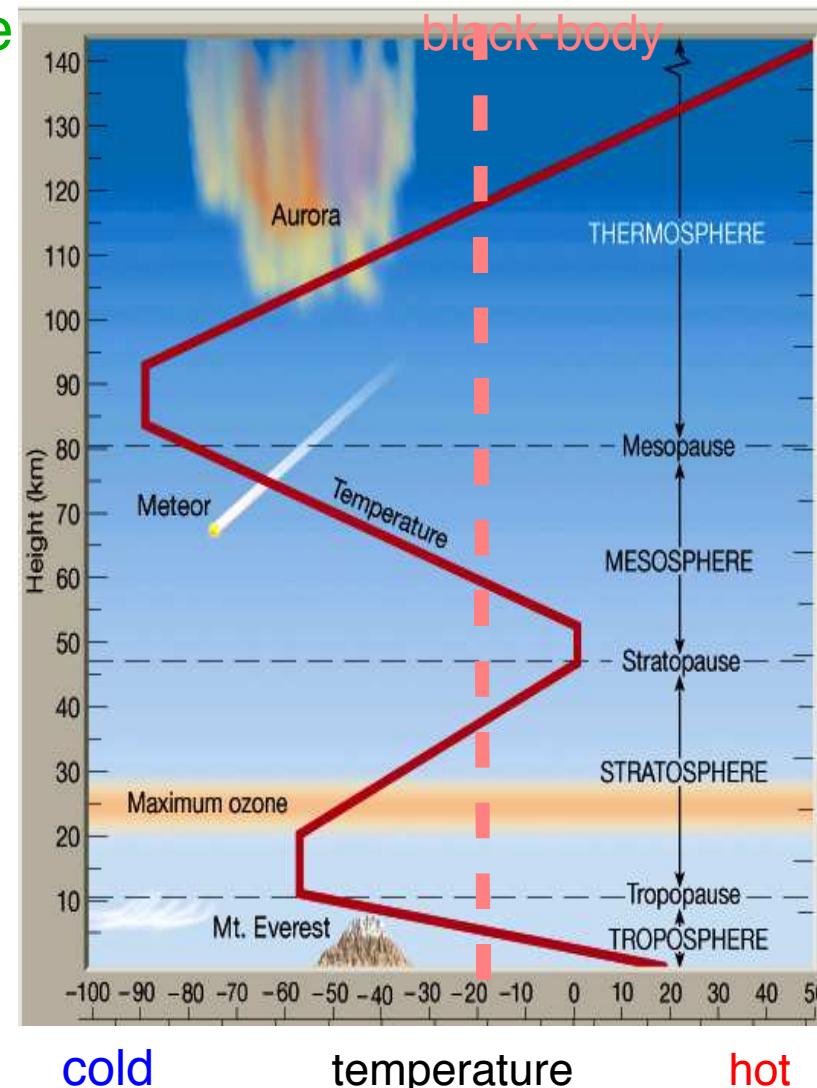
Earth's atmospheric composition

From http://www.ux1.eiu.edu/~cfjps/1400/atmos_origin.html

Density & Temperature of our atmosphere

- 1) Temperature largely **isothermal**; density decreases exponentially, $H \sim 8 \text{ km}$
 Three local departures (T maxima)
 - Thermosphere absorbs X rays ($\sim 2000 \text{ K}$)
 - Stratosphere absorbs UV (O_3)
 - Ground absorbs whatever passes
- 2) Atmosphere largely transparent in optical, but opaque in infrared --> green-house effect
 - Troposphere heated by ground--> turbulent --> twinkling stars, planes fly @ $\sim 10\text{km}$
 - Astronomical observations: overcome turbulence & avoid absorption

(for Canadian Arctic site-testing, see
http://www.hia-iba.nrc-cnrc.gc.ca/atrv/inuksuit_e.html,
<http://www.casca.ca/ecass/issues/2006-ae/>)



Atmospheric optics: I) Why is the sky blue on Earth? Rayleigh scattering

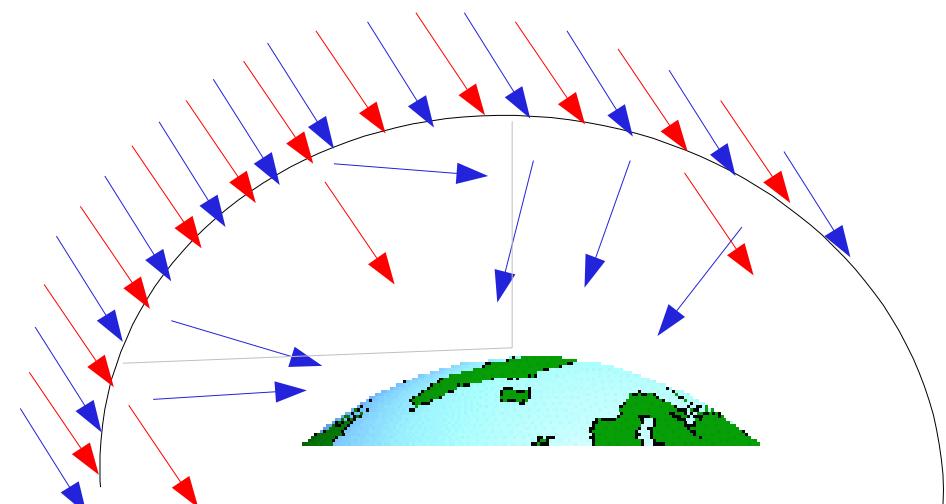
air molecules & other constituents
(N_2 , O_2 , H_2O droplets, dust...) all have sizes smaller than optical λ , and they preferentially scatter short- λ photons:
 $\sigma \sim 1/\lambda^4$

Earth: *sky is blue (--> ocean blue)*
sunset is red (reddened)
horizon whiter than zenith
Fall/Winter sky dark blue
UV is diffuse

Moon: *sky is black*

Mars: *sky is reddish yellow*

fine-dust (1-10 μm) Mie scattering --> white
iron oxide mineral absorption in the blue --> reddish



	x	y	sRGB pixel color
Sun above atmosphere	0.3259	0.3379	#fff3ea
5770 K blackbody (a Sun approximation)	0.3287	0.3397	#ffff1e6
Illuminant B ("direct sunlight")	0.3840	0.3516	#ffbfaa

Mars Pathfinder true-color picture of Martian noon



Atmospheric optics: II) Clouds

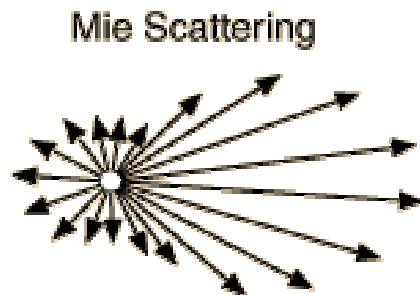
What are clouds?

How do they form?

Why are clouds white?

Aggregates of water or ice droplets suspended in air
In troposphere: low clouds-- water; high clouds-- ice
100% hum. + condensation nuclei (dust, cosmic-rays)
e.g., rising air that cools (--> humidity increases)

Water droplet colorless, solar light white
Mie scattering (droplets size $r \sim 10\mu\text{m} > \lambda$),
nearly geometric optics, no λ dependence
(at sunset, cloud is red)
soap foam: geometric scattering, also no λ dep.



Why don't clouds fall from the sky?

Tiny droplets, fall slowly; updraft mixing?

Fall and evaporate and form new ones?

Electrically charged clouds?



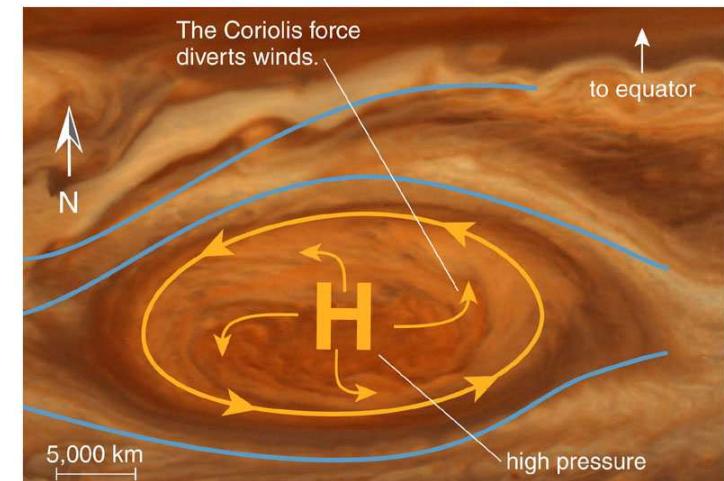
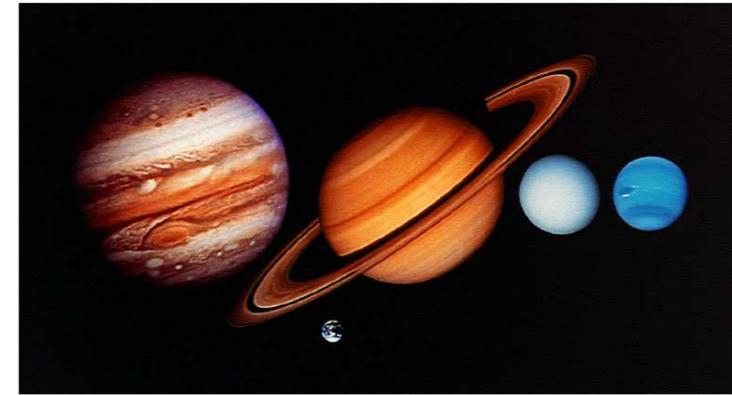
Intermezzo: Gas giant atmospheres

All 4 have deep atmospheres with mostly H₂ & He

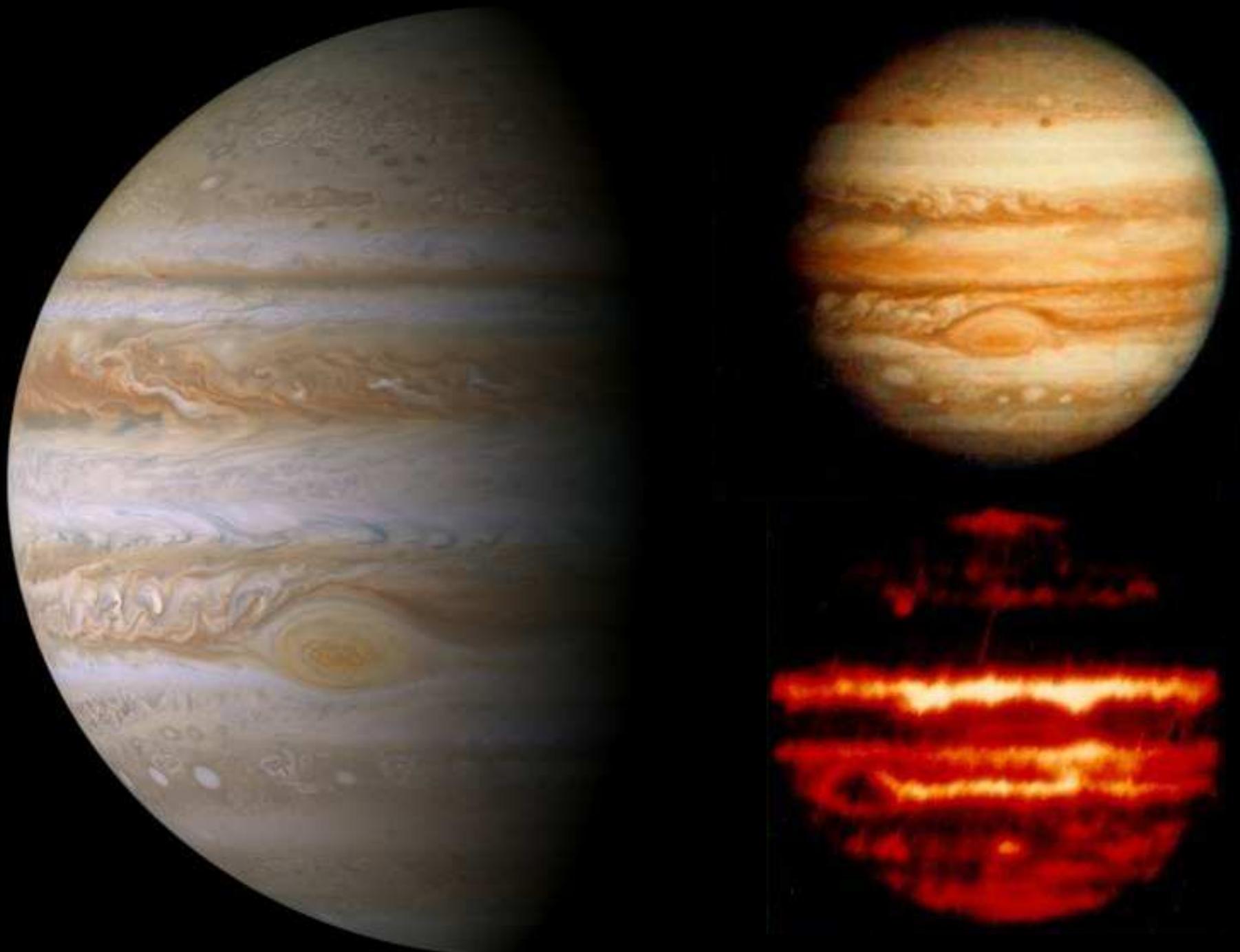
(fractions in % by volume, not by mass)

	J	S	U	N	Sun
H	88	97	83	74	86
He	11	3	15	25	14
CH ₄	0.2		2	1	

0.02 NH₃ helium settling no helium settling



- 1) Trace gases condense into **clouds** at diff. temperature
Clouds are also passive tracers of local wind pattern
- 2) Jupiter, Saturn & Neptune have strong **zonal winds**
(up to 500 m/s)
zonal winds driven by solar irradiation,
a combination of cold pole-- hot equator pressure gradient & Coriolis force:
great red-spot of Jupiter: a giant anti-cyclonic vortex, surprisingly long-lived
cyclone: $2 \mathbf{V} \times \boldsymbol{\Omega} = - \nabla P/\rho$; **tornado**: $\nabla^2 \mathbf{V} = - \nabla P/\rho$
- 3) Uranus: uniquely bland & sedate (no internal heat flux, obliquity 97 deg)



Origin of Earth's atmosphere

Our (& Venusian) atmosphere cannot be primordial

- 1) N₂, CO₂, H₂O are not condensed at 1AU from Sun,
O₂ does not naturally occur
- 2) Earth too low in mass to accrete gas directly
- 3) Gas is unlikely to have been trapped in solids and dragged
to Earth, since noble gases (Ne, Kr, Xe) are heavily depleted
relative to solar abundance.
- 4) New-born Earth molten and hot (10³K)
--> most gases can escape thermally.

Some relief only in that in the early bombardment period (~ 700 Myr)
water can be brought in by comets & asteroids.

(Note: D/H ratio in comets ~2 higher than ocean, so these cannot do it alone)

Origin of Earth's atmosphere (cont'd)

Our atmosphere is obtained gradually: volcanic outgassing & invaders

1 st atmosphere thermal escape	2 nd atmosphere outgassing/accretion	3 rd atmosphere absorbing CO₂
H & He(?)	CO ₂ /NH ₃ outgassed H ₂ O accreted/outgassed (<i>solid crust/ocean, 3.5Gyrs ago</i>)	most H ₂ O liquid CO ₂ got locked in O ₂ produced
P: ?	~ 100 bar (like Venus!)	~ 1 bar
T: ~10 ³ K	0°C < T < 100°C	~ 15°C

sinks of CO₂: sedimentary rock via H₂O, life (carbon) via photon-synthesis

sources of CO₂: volcanic outgassing (+human activities)

sinks of H₂O: subducting plates

sources of H₂O: outgassing, comets/asteroids?

Currently sensitive balance reached, mild green-house

run-away green-house: too much CO₂, H₂O can all disappear -->

--> sink disappears as well while outgassing produces yet more CO₂

Venus: divergent evolution from Earth

	a(AU)	mass(M_E)	spin	atm. Pressure	T	tectonics	ocean
Earth:	1	1	1 day	1bar	288 K	Yes	Yes
Venus:	0.7	0.8	243 day	92bar	770 K	No	No

- 1) 97% CO₂ in the atmosphere, ~ 700K, *no CO₂ sink due to dryness*
- 2) Why so dry? high D/H ratio indicates past large H₂O reserve
Green-house runaway and H₂O photo-evaporated
- 3) Cratering no older than ~0.8 Gyr --> tectonics stopped recently

**A planet is a nonlinear system.
Strongly divergent evolution can occur.**

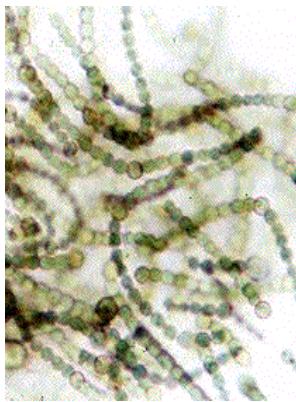
Cause & Effect?

- 1) Slightly closer to the Sun and got torched?
Or formation site had naturally less H₂O?
- 2) Too much CO₂ to start with and H₂O never condensed
(But: Initial Earth atm. ~100 bar, mostly CO₂ --> would require *fine tuning*?)



The Story for Mars: 2nd atmosphere gradually lost, no outgassing (tectonics)

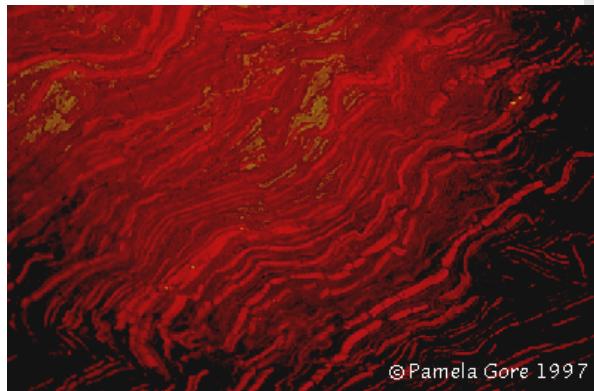
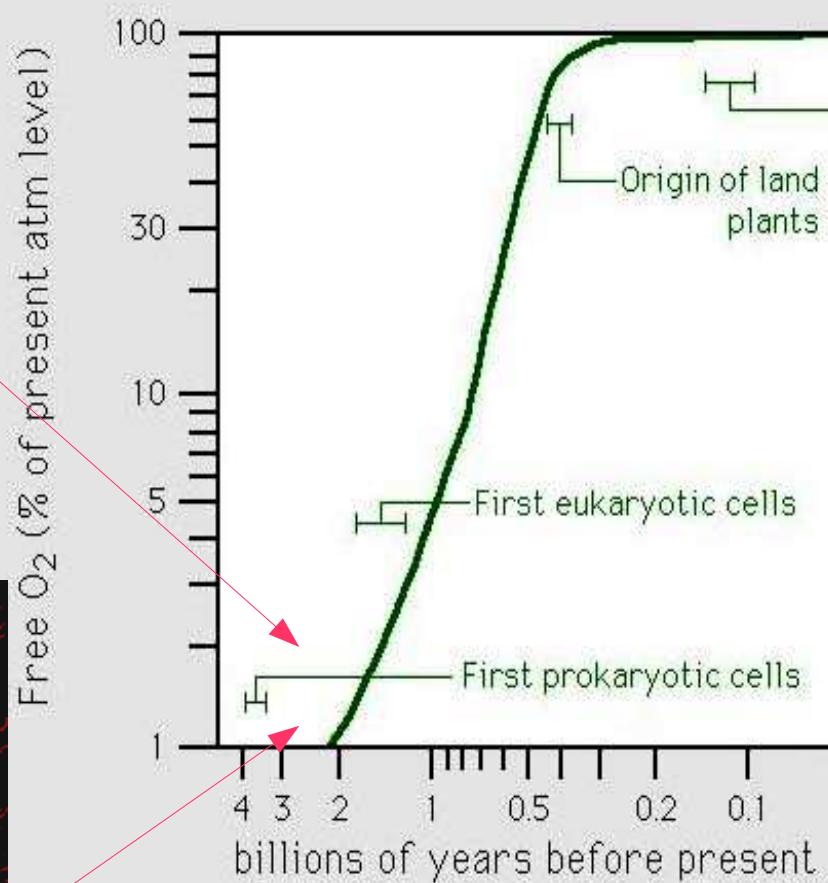
Origin of O₂ on Earth: photosynthesis; CO₂ + H₂O + hν → O₂ + carbo-hydrate



Archean Life: blue-green 'algae'
or cyanobacteria (3.5-2.2 BYA)
anaerobic

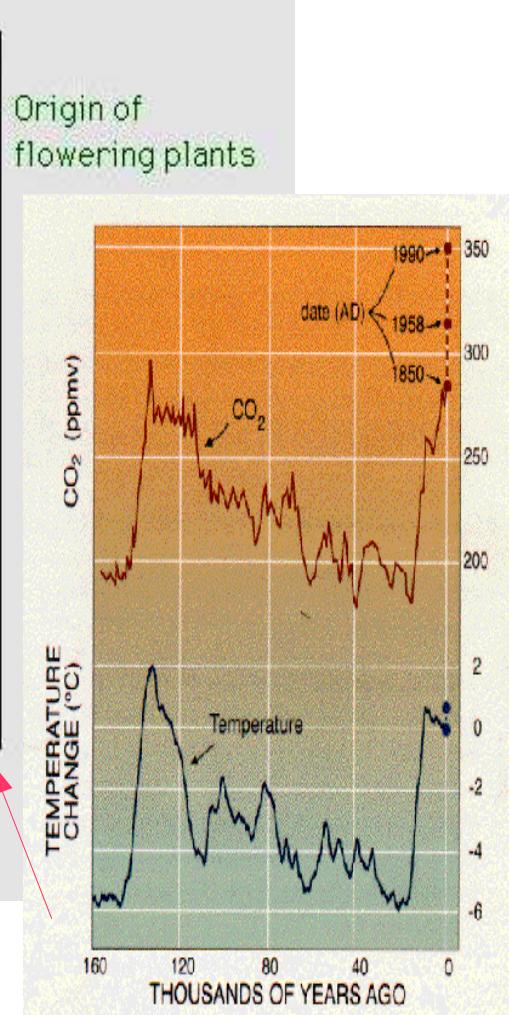


From <http://www.clas.ufl.edu/users/mrosenme/Oceanography>



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Red-banded un-oxidized iron-rich rocks, pre-cambrian, ~2.5BYA
<http://www.dc.peachnet.edu/~pgore/geology/geo102/precamb.htm>



CO₂ and atm. T correlation
(April 1989, Scientific American)