LETTER

Correcting datasets leads to more homogeneous early–twentieth–century sea surface warming

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Existing estimates of sea surface temperatures (SSTs) indicate that, during the early twentieth century, the North Atlantic and northeast Pacific oceans warmed by twice the global average, whereas the northwest Pacific Ocean cooled by an amount equal to the global average[1–4]. Such a heterogeneous pattern suggests first-order contributions from regional variations in forcing or in ocean–atmosphere heat fluxes[5,6]. These older SST estimates are, however, derived from measurements of water temperatures in ship-board buckets, and must be corrected for substantial biases[7–9]. Here we show that correcting for offsets among groups of bucket measurements leads to SST variations that correlate better with nearby land temperatures and are more homogeneous in their pattern of warming. Offsets are identified by systematically comparing nearby SST observations among different groups[10]. Correcting for offsets in German measurements decreases warming rates in the North Atlantic, whereas correcting for Japanese measurement offsets leads to increased and more uniform warming in the North Pacific. Japanese measurement offsets in the 1930s primarily result from records having been truncated to whole degrees Celsius when the records were digitized in the 1960s. These findings underscore the fact that historical SST records reflect both physical and social dimensions in data collection, and suggest that further opportunities exist for improving the accuracy of historical SST records[9,11].

According to recent estimates from the National Oceanic and Atmospheric Administration (NOAA)[1], global average SST warmed by 0.43 °C between 1908 and 1941. Whereas the North Atlantic warmed by 0.82 °C, the North Pacific showed a bimodal structure, with the northwest Pacific cooling by −0.39 °C and the northeast Pacific warming by 1.02 °C. Other gridded SST products give similarly disparate SST trends for the early twentieth century (Table 1 and Extended Data Fig. 1), and together these SST estimates suggest that internal modes of variability strongly contributed to early-twentieth-century climate change. Specifically, the Atlantic Multidecadal Oscillation and the Pacific Decadal Oscillation have been suggested to account for regional variations as well as more than half of the global warming between 1908 and 1941 (refs [6,12]). Model simulations of the atmospheric and oceanic response to prescribed radiative forcing do not, however, reproduce either the magnitude[13,14] or the pattern[5,15] of the early-twentieth-century warming seen in observations (Extended Data Fig. 1e). Difficulty in reproducing observations has been suggested to arise from deficiencies in how radiative forcing is prescribed[16] or from model limitations in representing internal climate variability[17,18].

Another possibility is that observational estimates of SST changes contain undetected biases, for which there are some precedents. Difficulty in simulating a slowdown in global warming between 1997 and 2012 was partly reconciled by revising SST estimates[19], amongst other considerations[20]. In another study[21], a jump in global temperature by 0.3 °C in 1945 was attributed to offsets between engine-room intake and bucket SST estimates.

The four major SST products covering the early twentieth century each rely upon the International Comprehensive Ocean-Atmosphere Data Set (ICOADS)[22], whose latest release is 3.0. It is estimated that 94% of observations between 1908 and 1941 were from buckets (Fig. 1). Bucket measurements of SST are biased by evaporative, sensible and solar heat fluxes that depend on a range of factors, including weather, ship deck height and bucket type[7]. For example, a canvas bucket left on deck for three minutes under typical wind and other weather conditions can give water temperatures that are approximately 0.5 °C cooler than a wooden bucket measured using the same protocol[7,9].

Previous corrections for bucket-measurement biases have involved assumptions that these biases change smoothly in space or time[1,4].

Table 1 | Early-twentieth-century SST trends

	ICOADSa	ICOADSb	ERSST5	COBESST2	HadISST2	HadSST3
Global trend	0.51 ± 0.03	0.56 ± 0.10	0.43	0.48	0.44 ± 0.04	0.47 ± 0.03
N Atlantic trend	0.85 ± 0.03	0.66 ± 0.11	0.82	0.79	0.74 ± 0.04	0.71 ± 0.03
N Pacific trend	0.31 ± 0.03	0.56 ± 0.11	0.37	0.39	0.32 ± 0.04	0.37 ± 0.03
NW Pacific trend	−0.35 ± 0.04	−0.02 ± 0.11	−0.39	−0.14	−0.34 ± 0.06	−0.30 ± 0.04
NE Pacific trend	0.86 ± 0.04	1.03 ± 0.12	1.02	0.85	0.85 ± 0.06	0.94 ± 0.04
East Asia trend	−0.06 ± 0.04	0.34 ± 0.11	−0.31	0.04	−0.09 ± 0.04	0.00 ± 0.04
Eastern US trend	0.92 ± 0.05	0.65 ± 0.13	0.75	0.82	0.61 ± 0.05	0.56 ± 0.05
PDO trend	1.82 ± 0.19	1.53 ± 0.20	2.03 ± 0.15	1.64 ± 0.11	2.02 ± 0.12	1.98 ± 0.18
East Asia air-temperature correlation	0.67	0.85	0.53	0.68	0.69	0.71
Eastern US air-temperature correlation	0.65	0.70	0.67	0.56	0.72	0.76
Correction pattern correlation	−0.49	−0.10	−0.49	−0.43	−0.45	−0.42

Trends are averaged over non-grey areas as in Fig. 3a, with the North Atlantic and North Pacific defined as poleward of 20° N; the northwest Pacific between 120° E and 180° E, and 25° N and 45° N; and the northeast Pacific between 120° W and 160° W, and 20° N and 60° N. East Asia and Eastern US regions are shown in Extended Data Fig. 5c. All trends are between 1908 and 1941, and are in units of °C per 34 years, with uncertainties reported at the 2 s.d. level. Each reported SST trend uncertainty, derived from dataset ensembles where available, includes contributions from bucket corrections, but only ICOADSb also accounts for groupwise offsets. Cross-correlations (Pearson's r) are reported between interannual air temperatures[26] and SSTs from East Asia and the Eastern US between 1908 and 1941, and between the spatial patterns of trends in SSTs (Fig. 3a and Extended Data Fig. 1a–d) and groupwise corrections (Fig. 3b).

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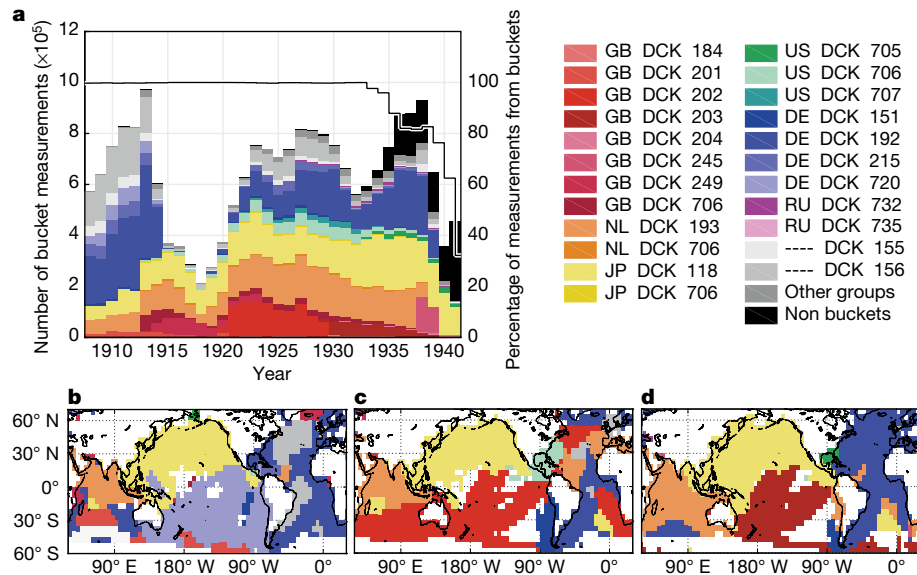
Fig. 1 | Groups of bucket SST measurements. a, Left-hand *y*-axis: number of bucket SST measurements from individual groups identified by country and deck information in ICOADS3.0. Country name abbreviations are: DE, Germany; GB, Great Britain; JP, Japan; NL, The Netherlands; RU, Russia; US, United States; and —, missing. Groups having fewer than 100,000 measurements are labelled as 'other groups'. Decks 118 and 762 are combined into 'JP DCK 118' because they are both Japanese Kobe Collection decks. Right-hand *y*-axis: percentage of measurements that have come from buckets, showing that nearly all observations before 1935 are from buckets (black line). **b–d**, Maps indicating nations that contribute the most observations within 5° × 5° grids for the periods 1908–1918 (**b**), 1919–1928 (**c**) and 1929–1941 (**d**). White grid boxes have fewer than three years of data.

HadSST3, for example, represents transitions between wooden buckets and less-insulated canvas buckets by using globally uniform, linear weights[4]. Another method assumes that differences between SST and night-time marine air temperatures remain similar to an average for 1971–2000 (ref. [1]). Simple assumptions regarding bucket corrections are commensurate with the limited metadata available for determining observational characteristics[2,4,7], but are questionable on the basis that ICOADS contains observations from a wide variety of ships that sailed for different purposes and sampled the ocean unevenly[22] (Fig. 1b–d).

More comprehensive bucket corrections are made possible by calculating offsets between groups of SST measurements that are nearby in space and time[10]. Specifically, we difference bucket SST measurements
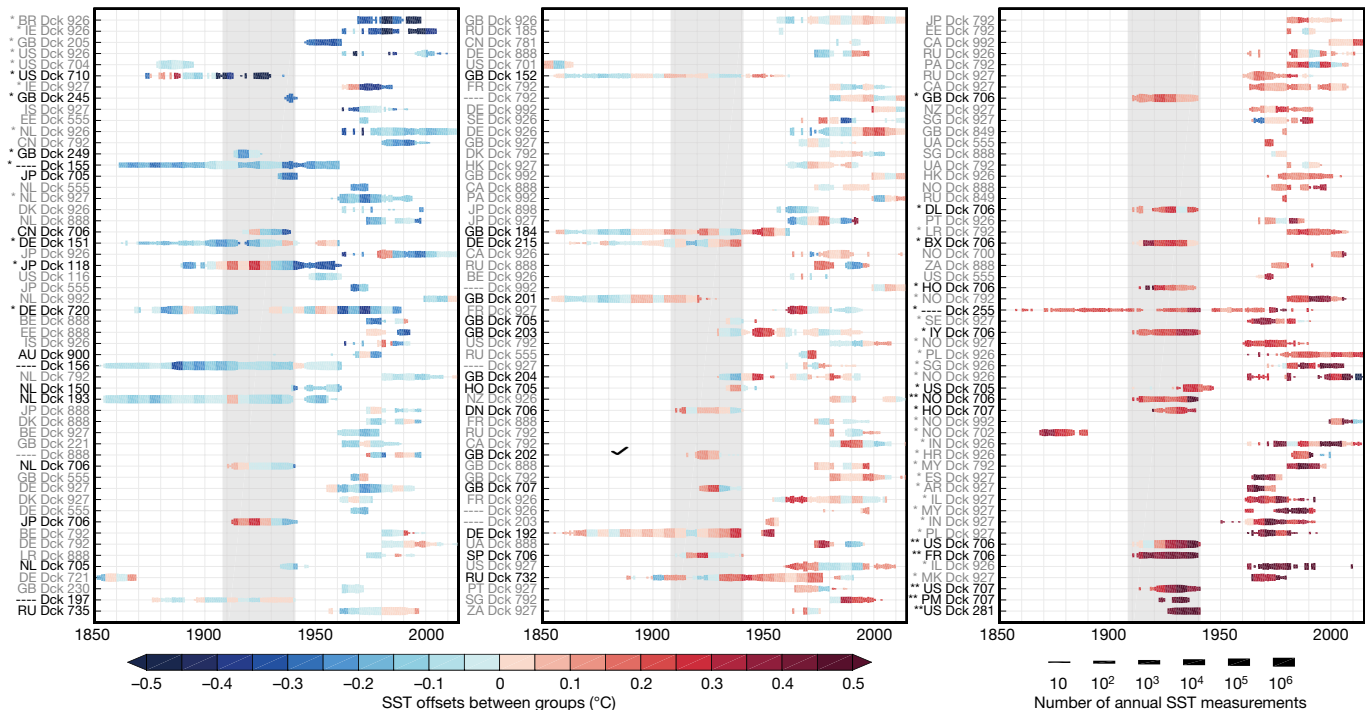

Fig. 2 | Relative offsets between groups of bucket SSTs. Groups for which fixed effects differ significantly from zero are indicated by * (*P* < 0.05). The 46 out of 162 groups that contribute data between 1908 and 1941 are indicated in black, and those remaining significant after a Bonferroni correction (*P* < 0.05/46) are indicated by **. Shading indicates the sum of fixed and five-yearly effects (regional effects are not shown). Bar widths indicate the number of SST measurements contributed by each group for each year. Abbreviated country names correspond to those in Fig. 1 and are listed in Supplementary Table 1.
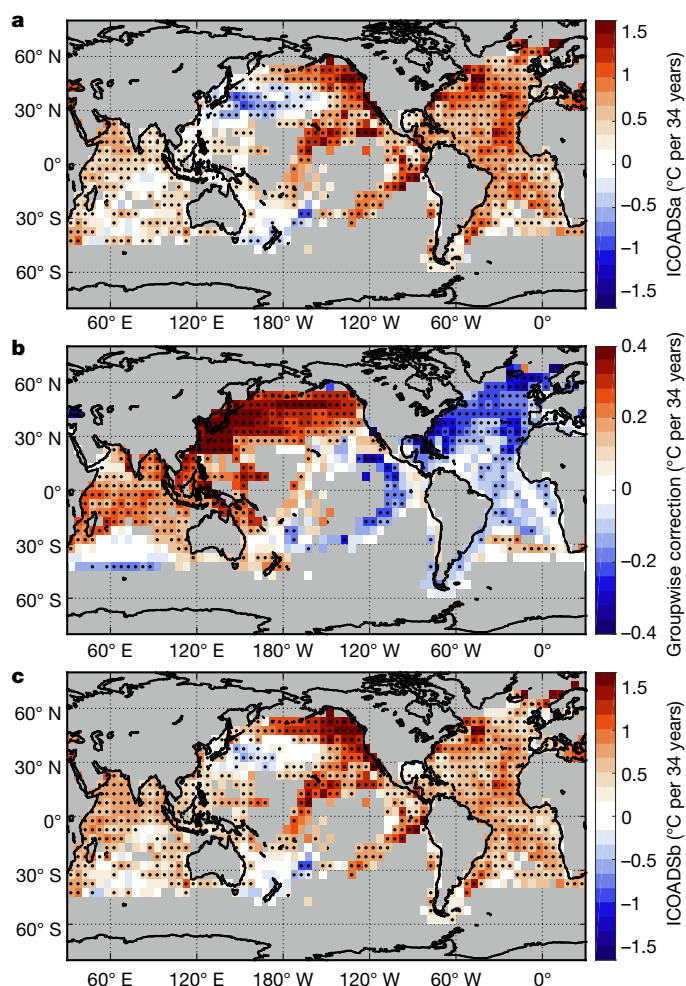
Fig. 3 | Maps of SST trends and corrections between 1908 and 1941.
a, SST trends in ICOADSa are similar to patterns found in existing SST estimates (Extended Data Fig. 1). **b**, Trends associated with the corrections for groupwise offsets. Note that panel **b** is plotted on a different colour scale. **c**, SST trends in ICOADSb after applying groupwise corrections. Areas in grey are inadequately sampled for purposes of calculating trends (see Methods). Dots indicate significant trends ($P < 0.05$). In ICOADSb (**c**), 77% of boxes show statistically significant warming, whereas only 2% show significant cooling. By contrast, in ICOADSa (**a**), 6% of boxes indicate significant cooling.

Fig. 4 | Diverging estimates of regional temperature variations.
a, **b**, Annual SST anomalies from different datasets in the North Pacific (**a**) and North Atlantic (**b**) oceans. Anomalies are relative to the 1920–1929 mean of each SST estimate. ICOADSb shows greater warming in the North Pacific and less warming in the North Atlantic relative to previous estimates. Uncertainties associated with ICOADSb (blue shading, 2 s.d.) are for annual average SSTs for each sub-basin, and are an order of magnitude larger than those reported for HadSST3 (red shading). Note that those uncertainties included in HadSST3 are mostly removed when computing the anomaly. The discrepancy in annual average SST uncertainties is larger than the discrepancy for trends (Table 1).

that come from distinct groups of ships within 300 km and 2 days of one another, giving a dataset of 6.1×10^6 SST comparisons between 1908 and 1941. Groups are designated according to nations and 'decks', with the latter term inherited from the fact that marine observations were stored using decks of punch cards. SST differences are analysed using a linear-mixed-effects (LME) methodology after accounting for climatological effects associated with location, day of year, and hour of day[10]. Mean offsets between groups of SST measurements range from $-0.3\,°C$ to $+0.6\,°C$ (Supplementary Table 1). Of the 46 nation–deck groups that contribute SST observations between 1908 and 1941, 21 have significant offsets ($P < 0.05$), and 6 remain significant after applying a Bonferroni correction[23] for multiple hypothesis testing ($P < 0.05/n$, $n = 46$; Fig. 2).

The presence of systematic offsets between groups of measurements, combined with changes in the distribution of these groups over time (Fig. 1), is liable to introduce spurious SST trends. To diagnose these trends, we first construct a bucket SST dataset that is corrected for biases common to all groups of bucket SSTs following the same approach used for HadSST3 (refs [4,7]). This reference dataset, ICOADSa, is then further corrected for offsets between groups of ships to obtain our best estimate of SST trends, ICOADSb (see Methods). Comparing
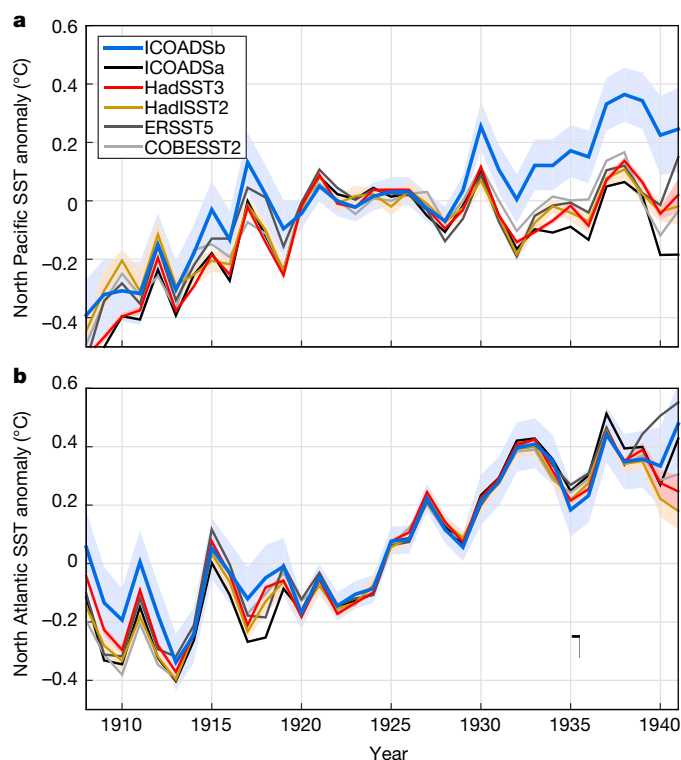
the trends in ICOADSa (Fig. 3a) against the difference in trends between ICOADSa and ICOADSb between 1908 and 1941 (Fig. 3b) shows a spatial anticorrelation (Pearson's r-value) of -0.49 ± 0.03. All uncertainties are reported with two standard deviations (s.d.) unless otherwise noted. Similar anticorrelations between our groupwise trend corrections and baseline SST trends are found for each of four major SST products (Table 1), indicating that some of the structure in early-twentieth-century SST trends reflects offsets between nations and decks.

The corrections included in ICOADSb result in a more homogeneous warming pattern (Extended Data Fig. 1f). North Pacific trends change from $0.31 \pm 0.03\,°C$ per 34 years in ICOADSa to $0.56 \pm 0.11\,°C$ per 34 years in ICOADSb, and North Atlantic trends from $0.85 \pm 0.03\,°C$ per 34 years to $0.66 \pm 0.11\,°C$ per 34 years (Fig. 4). Results are also temporally more homogeneous, with ICOADSa and other SST estimates indicating a slight cooling between 1920 and 1941 in the North Pacific, but ICOADSb showing a continuous warming trend throughout the early twentieth century (Fig. 4a). Furthermore, whereas ICOADSa indicates that 6% of the 5° grid boxes for which SST trends are computed contain significant cooling ($P < 0.05$), ICOADSb indicates that only 2% of grid boxes contain significant cooling (Fig. 3). Our focus is on the 1908–1941 interval for consistency with a previous study[12], but results are similar if neighbouring starting and ending years are chosen instead (Supplementary Table 2).

The prominent revision to Pacific SST trends relates to a change in offsets identified in the Japanese Kobe Collection, where a mean offset of $0.07 \pm 0.12\,°C$ between 1908 and 1930 drops to $-0.28 \pm 0.13\,°C$ between 1935 and 1941 in the North Pacific (Extended Data Fig. 2c).

One indication of why offsets change comes from considering the distribution of trailing digits in the Kobe Collection records, which are distributed across decks 118, 119 and 762. Before 1932, Kobe Collection records come from deck 762, and the distribution of trailing digits in reported SSTs is consistent with measurements that were originally recorded in whole-degrees Fahrenheit, then converted to Celsius, and finally rounded to tenths-of-a-degree (Extended Data Fig. 3a). Such unit conversion and rounding increases noise but results in negligible systematic offsets[24]. Between 1933 and 1941, however, more than 99% of Kobe Collection records come from deck 118 and are archived in whole-degrees Celsius. Examination of the reference manual for deck 118 reveals that digitization of these observations involved truncation[25] (Extended Data Fig. 3d).

If, before truncation, deck 118 contained the same distribution of trailing digits as found in deck 762, truncation would be expected to introduce a cold offset of $-0.46\,°C$. We find an offset of $-0.35 \pm 0.07\,°C$, with the smaller magnitude possibly reflecting the presence of additional offsets between decks. Cold offsets identified in the Kobe Collection during the 1930s are robust to whether or not decks 762 and 118 are treated as distinct groups in our analysis, because these decks are well separated in time and our LME model allows offsets to vary temporally (see Extended Data Figs. 2d and 4 for a sensitivity analysis). Kobe Collection deck 119 is also truncated and found to have a cold offset, but there is little consequence for early-twentieth-century trends because deck 119 spans 1951 to 1961. The major biases identified in decks 118 and 119 are associated with data-management practices and are not addressed in physical models used for correcting bucket temperatures. Furthermore, truncation biases in SSTs could, in principle, be identified by comparing against night-time marine air temperatures, but in the case of decks 118 and 119 this other indicator of surface temperature was also truncated.

North Atlantic warming in ICOADSb is revised downwards by $-0.20 \pm 0.11\,°C$ between 1908 and 1941, primarily as a consequence of the revision of SSTs from deck 156 and German deck 192 (Fig. 4b and Extended Data Fig. 2b). Deck 156, a group without country information, is biased relatively cold by $-0.22 \pm 0.12\,°C$ during 1908–1941 in the North Atlantic (Extended Data Fig. 2c). Deck 156 contributes 41% of North Atlantic observations between 1908 and 1912, but this contribution drops to less than 1% during World War I and returns only to 7% thereafter. Deck 156 corrections predominantly warm SST estimates earlier in the twentieth century, decreasing the warming trend by $-0.05 \pm 0.03\,°C$ per 34 years. Another major correction involves making German deck 192 cooler by $0.33 \pm 0.13\,°C$ between 1920 and 1941 (Extended Data Fig. 2c), thereby decreasing estimates of North Atlantic warming by $-0.08 \pm 0.03\,°C$ per 34 years. A possible clue to the origins of this offset is that German deck 192 shows an increasing portion of SSTs that are reported in whole-degrees Celsius after the 1920s (Extended Data Fig. 3c), although no major modification of the data is indicated by available documentation.

A test of whether our groupwise SST corrections are skilful is made by comparing ICOADSa and ICOADSb against land-based air-temperature records. We compare SSTs from the northwest Pacific and northwest Atlantic against adjacent land temperatures from CRUTEM4 (ref. [26]), because these regions show large corrections near a network of land-temperature stations. The Pearson cross-correlation for inter-annual temperature variability in the northwest Pacific between 1908 and 1941 increases from 0.67 when using ICOADSa to 0.85 when using ICOADSb. Furthermore, the 1908–1941 trend increases from $-0.06 \pm 0.04\,°C$ per 34 years in ICOADSa to $0.34 \pm 0.11\,°C$ per 34 years in ICOADSb, bringing it into agreement with the $0.33\,°C$ per 34 years trend in CRUTEM4 (Table 1 and Extended Data Fig. 5). Improved agreement for the northwest Atlantic is more modest. Northwest Atlantic correlation increases from 0.65 to 0.70 from ICOADSa to ICOADSb, and trends decrease from $0.92 \pm 0.05\,°C$ per 34 years to $0.65 \pm 0.13\,°C$ per 34 years, but this is still significantly higher than the CRUTEM4 trend of $0.32\,°C$ per 34 years.

The Pacific Decadal Oscillation[18] (PDO) shifts towards an increasingly positive phase over 1908–1941 (ref. [27]). The magnitude of the PDO trend is smaller in ICOADSb than in ICOADSa (Table 1), but both are consistent with trends found in North Pacific sea level pressure fields[28] (Extended Data Fig. 6).

Equally important to changed SST trends is the greater uncertainty estimated to accompany these trends (Extended Data Fig. 7). The global average SST trend in HadSST3 between 1908 and 1941 has a reported 2 s.d. uncertainty of $0.03\,°C$ per 34 years, whereas accounting for groupwise offsets in ICOADSb reveals an uncertainty of $0.10\,°C$ per 34 years. Larger uncertainties in ICOADSb reflect the fact that averaging repeated observations within a group does not decrease systematic groupwise errors. One implication is that the expected correspondence between observed and simulated trends should be revised downwards. These results have implications for attribution of early-twentieth-century warming and extreme events[5,27,29]. Another implication is that the scope for further improvement of regional temperature estimates associated with better diagnosing and correcting for groupwise biases[9] is greater than previously recognized.

Finally, we briefly explore the implications of our results for model–data mismatches during the early twentieth century. Differences in rates of warming in the North Atlantic and North Pacific reduce from $0.54 \pm 0.03\,°C$ per 34 years in ICOADSa to $0.10 \pm 0.07\,°C$ per 34 years in ICOADSb. These revised interbasin trend differences are consistent with that of $0.00 \pm 0.40\,°C$ per 34 years found in the early-twentieth-century simulations from the Fifth Climate Model Intercomparison Project[30] (CMIP5; Extended Data Fig. 1e). But we note that the global-average rate of SST warming in ICOADSb is $0.56 \pm 0.10\,°C$ per 34 years, and that the same domain in the CMIP5 ensemble warms by only $0.19 \pm 0.17\,°C$ per 34 years—a discrepancy in warming rates noted previously for other SST estimates[27]. The model–data mismatch in rates of overall warming highlights the importance of continuing to investigate forcing, sensitivity and internal variability of the climate system along with corrections to historical SST estimates.

Online content

Any methods, additional references, Nature Research reporting summaries, source data, statements of data availability and associated accession codes are available at https://doi.org/10.1038/s41586-019-1349-2.

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METHODS

Identification of bucket measurements. To identify bucket measurements, we follow the same procedure used for HadSST3 (ref. [4]), using World Meteorological Organization Report Number 47 (WMO47) and ICOADS metadata. Prior to 1941, all SST measurements are assumed to be from buckets unless explicitly recorded otherwise. Analysis of the amplitude of the diurnal cycle in SST before 1941 supports unidentified records as being overwhelmingly from buckets[31]. From 1941 onwards, if the method of measurement is missing in both WMO47 and ICOADS metadata, SST measurements are assumed to come from buckets if the associated nations are reported to have at least 95% of their ships making bucket measurements in WMO47.

The nationality corresponding to each bucket SST measurement is determined using ICOADS country-code information and WMO47. If these metadata are unavailable, the first three letters of the ICOADS identification code are matched with international call signs[32], or the first two letters for decks 705, 706 and 707 (ref. [33]). Decks that may include substantial amounts of measurements not coming from buckets (decks 740, 780 and 874) are discarded, but this has little influence because these decks together contribute only 249 measurements between 1908 and 1941 out of a total of 20.5 million measurements.

Bucket corrections. The same methodology used for HadSST3 (ref. [4]) is applied to ICOADSa and ICOADSb to correct for biases common to all groups (see Extended Data Fig. 8 for comparison and Supplementary Table 3 for details). Corrections common to all groups are made using wooden- and canvas-bucket models[7] run at $5° \times 5°$ for individual climatological months. Bucket models are driven by the 1973–2002 monthly climatology of SST, 10-m air temperature, wind speed and specific humidity from the National Oceanography Centre (NOC) version 2.0 surface flux and meteorological dataset[34] and an insolation climatology from ERA-interim reanalysis[35].

Additional corrections are applied to groups of bucket observations in ICOADSb that are determined using an LME model:

$$\delta T = X\alpha + Z_y\beta_y + Z_r\beta_r + \beta_\sigma \qquad (1)$$

The vector of temperature differences, δT, is determined from proximal pairs of bucket SST observations that come from ships within 300 km and 2 days of one another that are associated with different nationalities and deck assignments. All bucket SST data identified in ICOADS3.0 between 1850 and 2014 are analysed, yielding 17.8 million paired SST differences (Extended Data Fig. 9). The 1908–1941 period contains a subset of 6.1 million SST differences. SST differences are adjusted for climatological effects associated with location, day of year, and hour of day[10].

SST differences contained in δT are represented as a 'fixed-effect' term describing offsets between groups, α, and random effects describing temporal variations (five-year blocks), β_y, and regional variations (17 sub-basin regions), β_r. Matrices X, Z_y and Z_r specify, respectively, common pairs of groups, five-year blocks, and region. β_σ is the residual, and estimates are derived using an expectation-maximization procedure[36]. Groupwise SST corrections are applied in ICOADSb by removing estimated offsets from each SST measurement according to group, year and region.

Equation (1) is run at two levels, one for determining international offsets and one, a more detailed level, for determining interdeck offsets within nations. Each level of offsets is constrained to equal zero when summed across all paired measurements and all years. The groupwise corrections applied to ICOADSb thus adjust for offsets between groups but do not alter the average across all data. A detailed description of the LME design and implementation, along with the sensitivity of results to plausible variants, is available in a methods paper[10]. In an update to ref. [10], the analysis presented here uses international call signs for identifying nationality, thereby allowing us to increase the number of groups for which more than 5,000 SSTs are compared from 96 to 162 (Extended Data Fig. 10). Only groups associated with at least 5,000 pairs of SST observations are retained.

Trend estimates. Regional trends are the average of local trend estimates resolved at $5° \times 5°$ grid boxes at monthly resolution. Monthly errors are represented as the sum of four different components, $e \approx N(0, \sigma_o^2 + \sigma_s^2 + \sigma_b^2 + \sigma_g^2)$. Terms represent the uncertainty due to errors associated with individual observations, σ_o; partial sampling of each grid box, σ_s; HadSST3-type bucket-adjustment errors, σ_b; and errors common to individual groups of SST measurements, σ_g. The first three terms are assumed to follow those reported for HadSST3 (ref. [4]), and the last is inferred through the LME model.

A 1,000-member ensemble of SST observations is generated to represent both the random and the systematic components of uncertainty in ICOADSb. In addition to uncertainties that are equivalent to those in HadSST3 (ref. [37]), groupwise bucket errors are included according to the results from our LME model. In particular, intergroup offset terms are drawn from a multivariate normal distribution that represents uncertainties associated with fixed group effects as well as random five-yearly and regional effects[10].

To compute trends, monthly anomalies are averaged to annual values, and years with fewer than three months of data are discarded. Empty monthly 5° grid boxes are infilled by averaging neighbouring grid boxes that are within 10°. Trends are reported only if SST estimates are present (or have been infilled) for both the first and the last five years, and data cover at least 26 of the years between 1908 and 1941. Trends are computed using standard linear least squares. The same procedure is followed for computing trends from the ensemble of realizations in order to estimate uncertainties.

The average trend uncertainty between 1908 and 1941 across sampled grid boxes is 0.12 °C per 34 years for $\sqrt{\sigma_o^2 + \sigma_s^2}$, 0.01 °C per 34 years for σ_b, and 0.06 °C per 34 years for σ_g. When taking spatial averages, σ_o and σ_s are independent across boxes, whereas σ_b is globally systematic and σ_g is partially systematic. The contributions of uncorrelated terms to uncertainties in the global mean trend are essentially negligible for σ_o and σ_s, remain at 0.01 °C per 34 years for σ_b, and become 0.05 °C per 34 years for σ_g. Groupwise errors are thus expected to dominate the uncertainties associated with large-scale SST trends.

Comparison with other datasets. There are a number of notable differences between the SST datasets considered here. ICOADSa, ICOADSb, and ERSST5 (ref. [1]) are based on ICOADS3.0 (ref. [22]), whereas COBESST2 (ref. [2]), HadISST2 (ref. [3]) and HadSST3 (ref. [4]) are based on ICOADS2.5 (ref. [38]). ICOADSa and ICOADSb use only bucket SST measurements, which are estimated to account for 94% of all observations in ICOADS3.0 between 1908 and 1941. Other datasets also make use of engine-room intake, buoy and drifter observations, which become more common after 1941. Finally, ERSST5, COBESST2 and HadISST2 infill monthly grid boxes without data, whereas ICOADSa, ICOADSb and HadSST3 leave these boxes unfilled.

Coastal near-surface air temperatures from CRUTEM4 (ref. [26]) are used to check the validity of our groupwise corrections. We choose to compare ICOADSa and ICOADSb with CRUTEM4 near the east coasts of Asia and North America because these two regions experience the largest adjustments in trends and because of the availability of relatively dense station and bucket data. In each area, regional SST and land air temperature time series are computed using only those grid boxes that contain both types of measurement (Extended Data Fig. 5c).

Pacific Decadal Oscillation. An SST index for the Pacific Decadal Oscillation, SST-PDO, is obtained by projecting annual-average SSTs poleward of 20° N in the Pacific onto a normalized PDO pattern, where the PDO pattern is obtained by regressing SST onto a standard National Centers for Environmental Information (NCEI) PDO index[18] over the years 1948–2014 and is then normalized to have zero mean and a range of one over the North Pacific. The same method is applied to sea level pressure (SLP) using the NOAA 20th Century Reanalysis[28], yielding SLP-PDO. SST-PDO is regressed against SLP-PDO over the periods 1908–1941 and 1948–2010 for sensitivities in units of °C mb^{-1}. Uncertainty in PDO trends is estimated by randomly perturbing PDO indices using error estimates of projections in individual years. When estimating the sensitivity of the SST-PDO index against the SLP-PDO index, the same random seeding is used to draw realizations of both indices for all SST products.

Data availability

All datasets used in this study are publicly available as follows: ICOADS3.0 (https://rda.ucar.edu/datasets/ds548.0/), HadSST3 and a 100-member ensemble (https://www.metoffice.gov.uk/hadobs/hadsst3/data/download.html), HadISST2 and a 10-member ensemble (https://www.metoffice.gov.uk/hadobs/hadisst/data/hadisst2/), COBESST2 (https://www.esrl.noaa.gov/psd/data/gridded/data.cobe2.html), ERSST5 (https://www.esrl.noaa.gov/psd/data/gridded/data.noaa.ersst.v5.html), CRUTEM4 (https://crudata.uea.ac.uk/cru/data/temperature/), ERA-Interim Reanalysis (https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era-interim), NOAA 20th Century Reanalysis (https://www.esrl.noaa.gov/psd/data/20thC_Rean/), and NOCv2.0 surface flux (https://rda.ucar.edu/datasets/ds260.3/). ICOADSa and ICOADSb, together with archived versions of all other SST datasets used in this study, are posted at https://doi.org/10.7910/DVN/DXJIGA.

Code availability

Code allowing the full reproduction of our results is posted on Github at https://github.com/duochanatharvard/Homogeneous_early_20th_century_warming.

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