

eReefs optical and biogeochemical model.



Coastal Environmental Modelling Team

- Physical oceanography



- Sediment dynamics



- Biogeochemistry



- Data assimilation



- Scientific computing



BGC state variables:

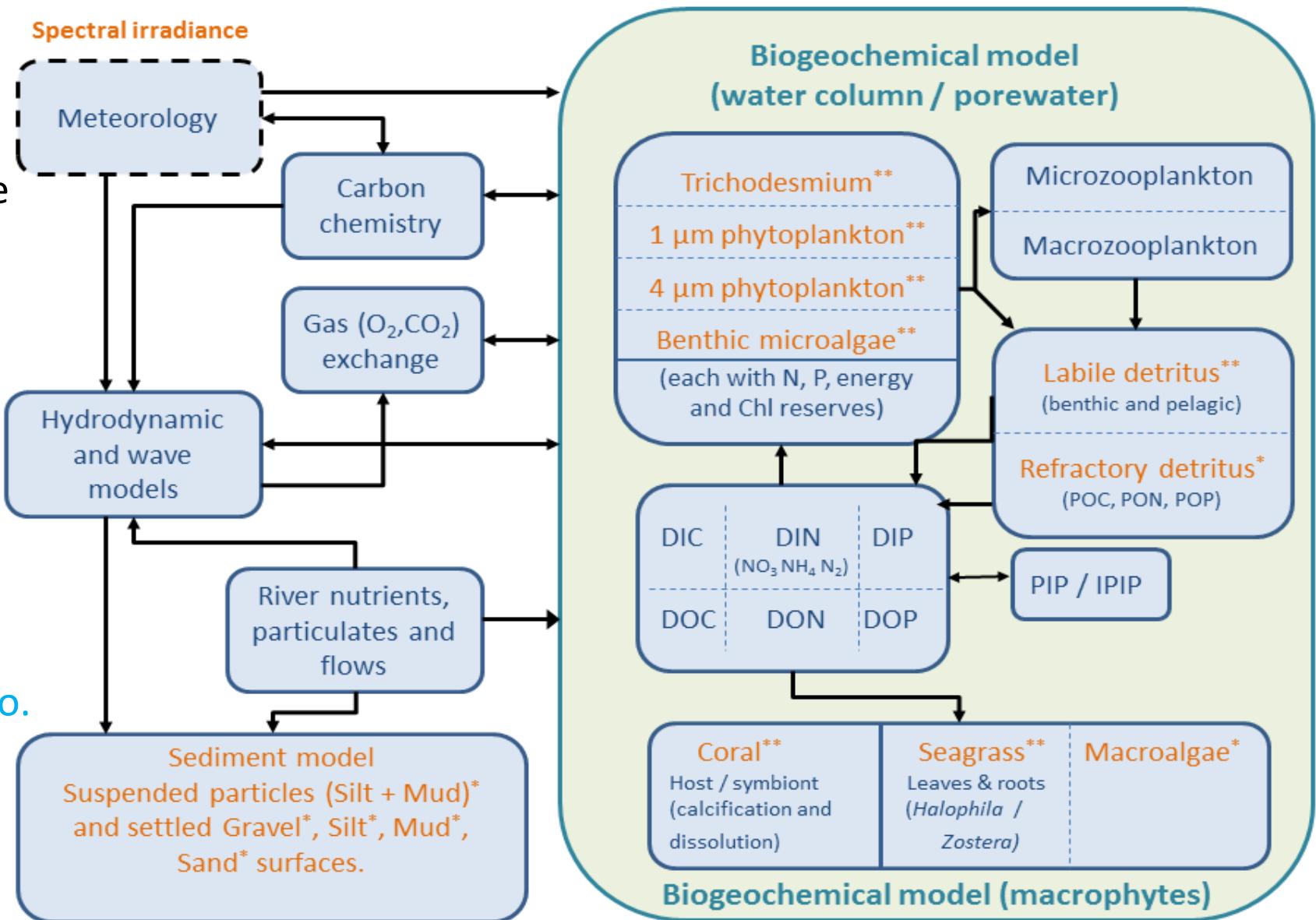
- 10 dissolved
- 22 living particulate
- 11 non-living part.
- 6 epibenthic.

Parameters:

- approx 100.

Full description:

[https://research.csiro.
au/reefs/models/](https://research.csiro.au/reefs/models/)



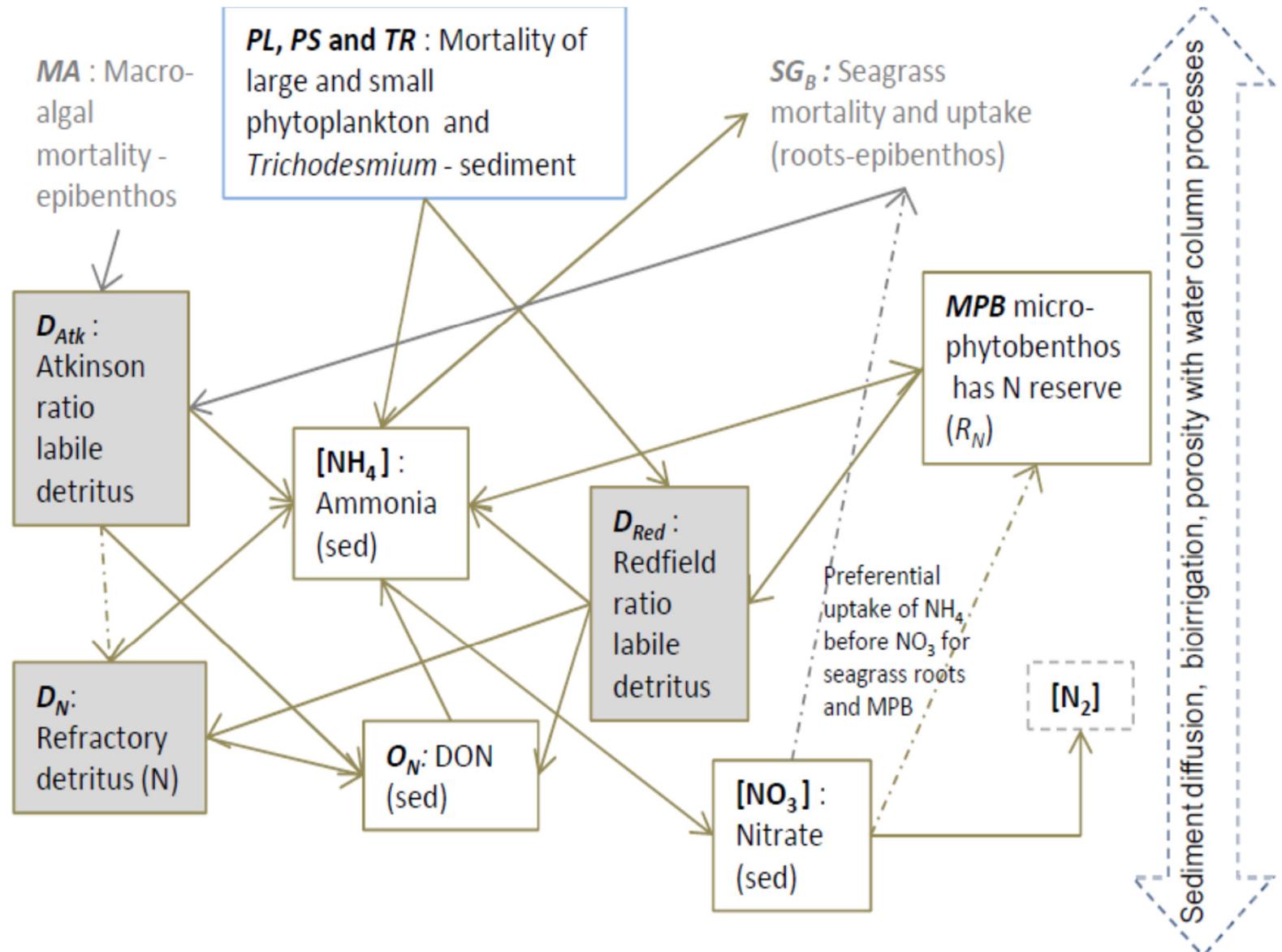
Organic nitrogen

Detrital processes include:

- Mortality
- Detrital decomposition
- Denitrification
- Nitrification

Other N processes

- Atmospheric deposition
- Preferential NH_4^+ uptake.
- N fixation.



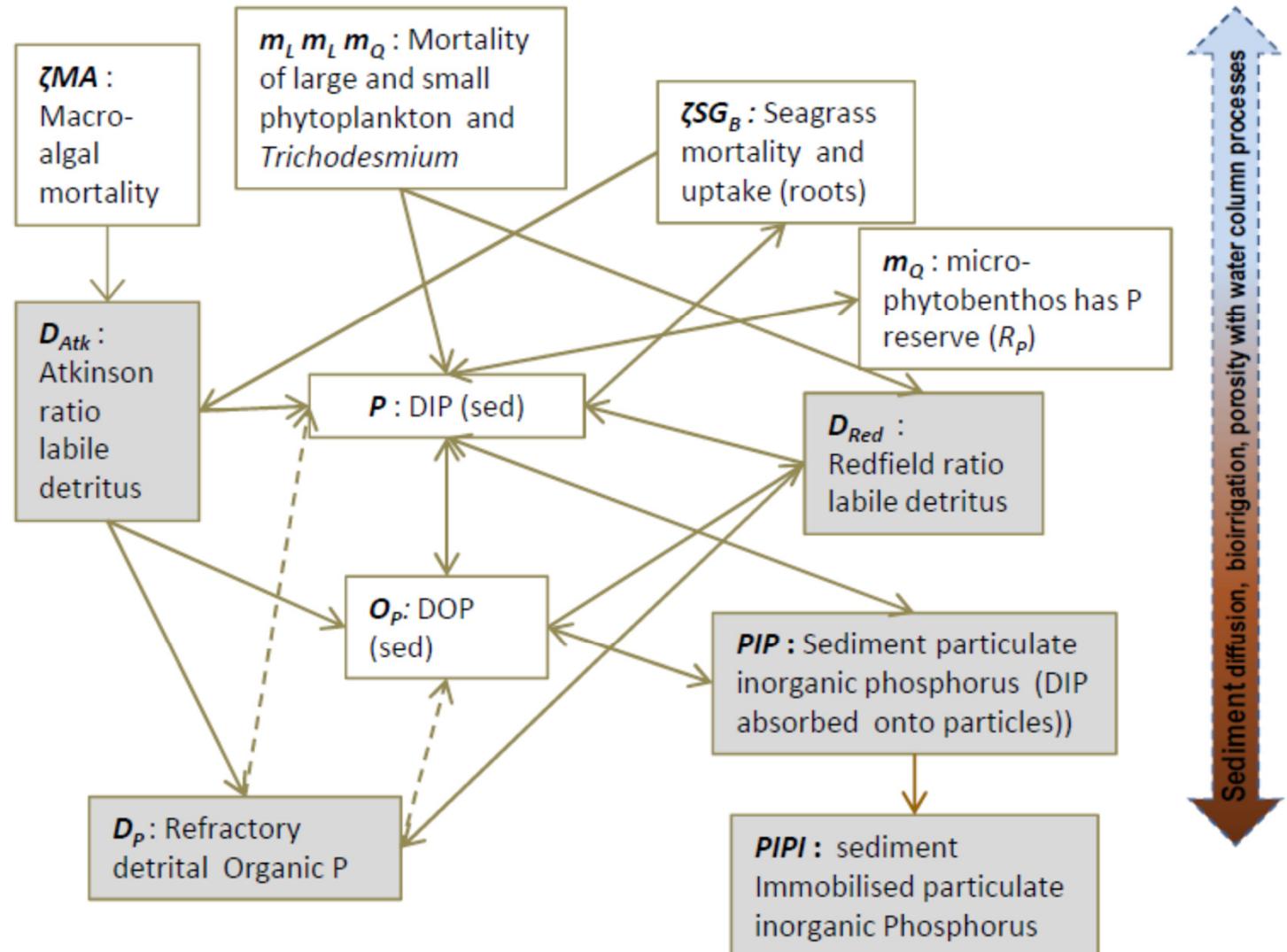
Organic phosphorus

Detrital processes include:

- Mortality
- Detrital decomposition

Other P processes

- Phosphorus adsorption / desorption.



Documentation of the biogeochemical model.

- Precise description of the model equations and parameter values are given in Appendix B. available on website.
- 132 pages, 50 Tables, 282 equations.
- New developments for eReefs published in Ecological modelling, Limnology and Oceanography, Environmental Modelling and Software:

Appendix B: CSIRO Environmental Modelling Suite: Scientific description of the optical, carbon chemistry and biogeochemical models

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January 28, 2016

A biophysical representation of seagrass growth for application in a complex shallow-water biogeochemical model *Ecol. Mod.* **325**: 13-27.
Remote-sensing reflectance and true colour produced by a coupled hydrodynamic, optical, sediment, biogeochemical model of the Great Barrier Reef,
Australia: comparison with satellite data. *Env. Model. Software* **78**: 79-96.

The exposure of the Great Barrier Reef to ocean acidification. *Nature Communications* **7**, 10732.

The interacting effects of photosynthesis, calcification and water circulation on carbon chemistry variability on a coral reef flat *Ecol. Mod.* **284**, 19-34.

A physiological model for the marine cyanobacteria, *Trichodesmium*. In Piantadosi, J., Anderssen, R.S. and Boland J. (eds) MODSIM2013, 20th International Congress on Modelling and Simulation. Modelling and Simulation Society of Australia and New Zealand, December 2013, pp. 1652-1658.

A dynamic model of the cellular carbon to chlorophyll ratio applied to a batch culture and a continental shelf ecosystem. *Limnol. Oceanogr.* **58**, 1215-1226.
The interchangeability of autotrophic and heterotrophic nitrogen sources in Scleractinian coral symbiotic relationships. *Ecol. Model.* **250**, 183-194.

- Appendix B gives detail explanation of the parameters and state variables in easy to access tables (p108-134).

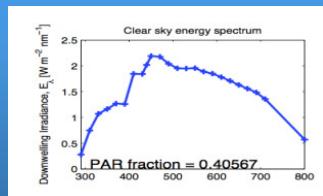
Description	Name in code	Symbol	Value	Units
Reference temperature	Tref	T_{ref}	20.000000	Deg C
Temperature coefficient for rate parameters	Q10	Q_{10}	2.000000	none
Nominal rate of TKE dissipation in water column	TKEeps	ϵ	0.000001	$m^2 s^{-3}$
Atmospheric CO2	xco2_in_air	pCO_2	396.480000	ppmv
Concentration of dissolved N2	N2	$[N_2]_{gas}$	2000.000000	$mg N m^{-3}$
DOC-specific absorption of CDOM 443 nm	acdom443star	$k_{CDOM,443}$	0.000130	$m^2 mg C^{-1}$

Phytoplankton chlorophyll nc_iV [mg m⁻³] Concentration of the chlorophyll *a* pigment of the population. The four phytoplankton classes have two pigments, a chlorophyll *a*-based pigment and an accessory pigment. As the pigment concentration adjusts to optimise photosynthesis, including the presence of the accessory pigment, the intracellular content, c_iV , represents only the chlorophyll *a*-based pigment. As the model does not distinguish between mono-vinyl and di-vinyl forms of chlorophyll, this c_i represents either form, depending on the phytoplankton type.

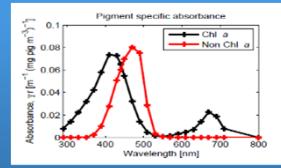
Optical model

Baird et al 2016. Environmental Modelling and Software 78: 79-96

Atmospheric forcing uses spectrally-resolved energy distribution of sun light.



IOPs = adsorption ($a_{T,\lambda}$) and backscattering ($b_{T,\lambda}$) of plankton, CDOM, particles.



Vertical attenuation:

$$K_\lambda = \frac{a_{T,\lambda}}{\cos \theta_{sw}} \sqrt{1 + (g_i + g_{ii} \cos \theta_{sw}) \frac{b_{T,\lambda}}{a_{T,\lambda}}}$$

Reflectance is the ratio of backscatter to attenuation:

$$u_\lambda = \int_0^z \frac{w_{\lambda,z'} b_{b,\lambda,z'}}{a_{\lambda,z'} + b_{b,\lambda,z'}} dz'$$

Where w is weighted by optical depth

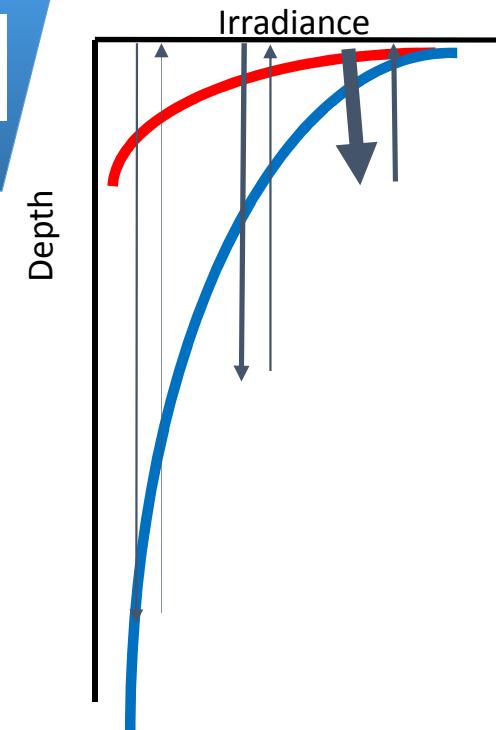
$$w_{\lambda,z} = \frac{1}{z_1 - z_0} \int_{z_0}^{z_1} \exp(-2K_{\lambda,z'}) dz'$$

Reflectance is function of

$$r_{rs,\lambda} = g_0 u_\lambda + g_1 u_\lambda^2$$

through the sea-air interface

$$R_{rs,\lambda} = \frac{0.52 r_{rs,\lambda}}{1 - 1.7 r_{rs,\lambda}}$$

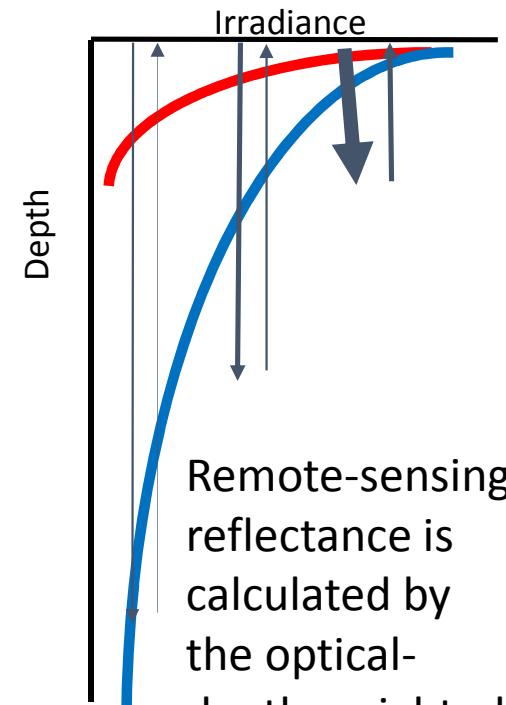
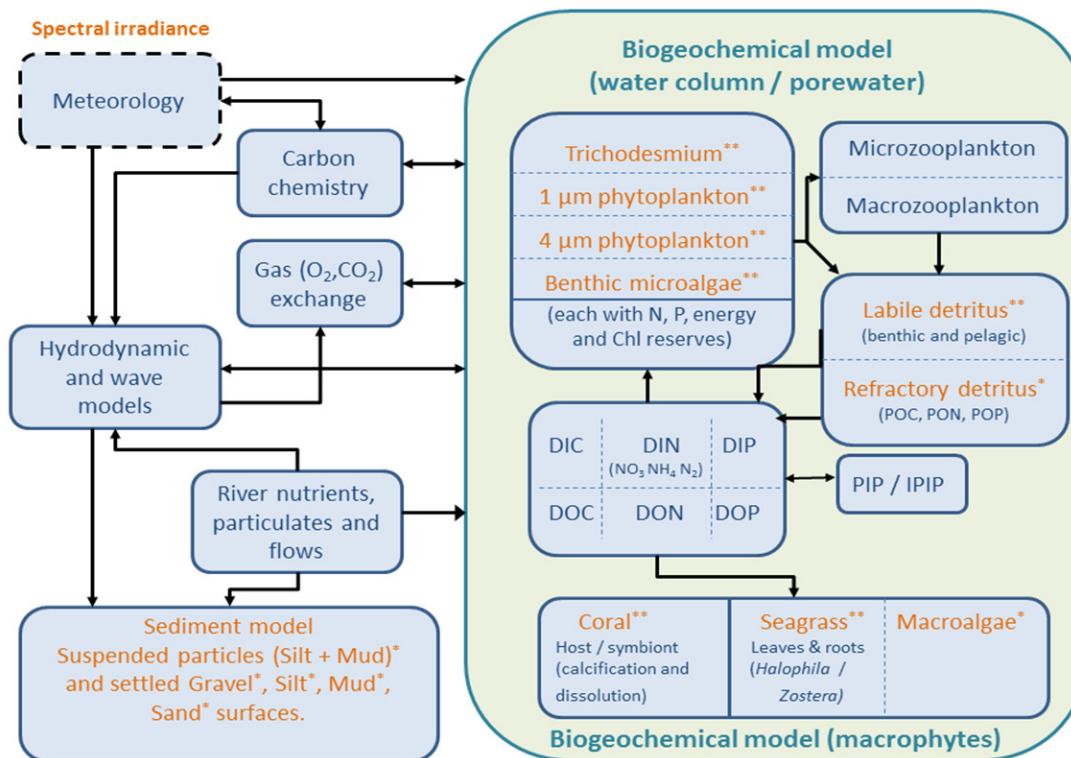


In other words:

Remote-sensing reflectance is calculated by the optical-depth weighted reflectance and bottom reflectance

24 optically – active components

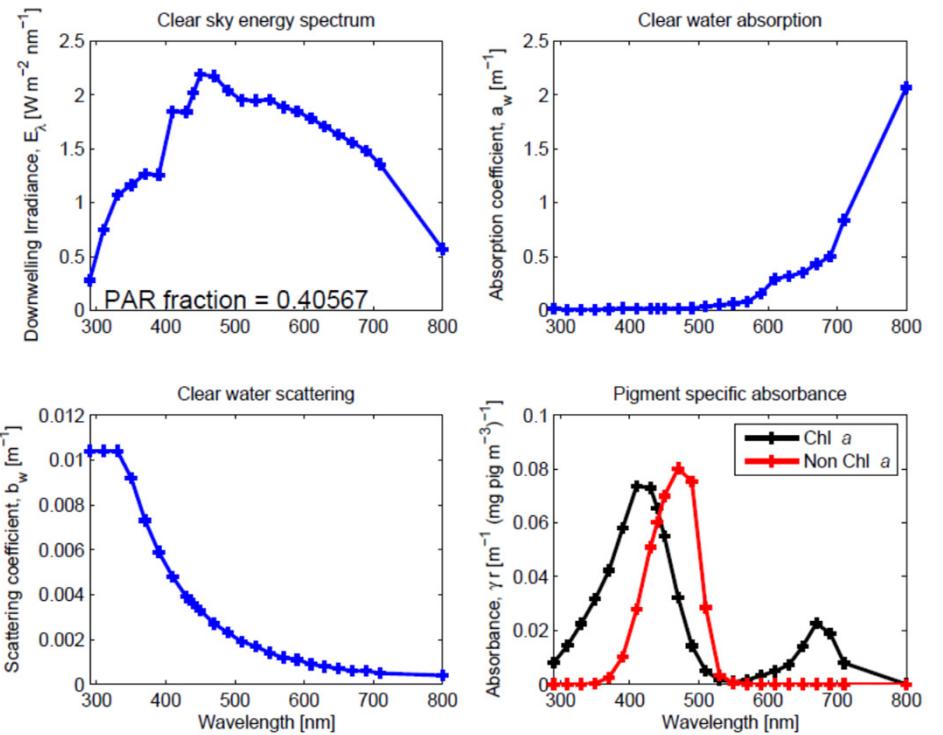
- Clear water
- 4 types of phytoplankton with two pigment types each
- CDOM
- Macroalgae
- 2 types of seagrass
- Coral skeletons
- Zooxanthellae
- sediment x 6
- Detritus x 3



Remote-sensing reflectance is calculated by the optical-depth weighted reflectance and bottom reflectance

Optical model

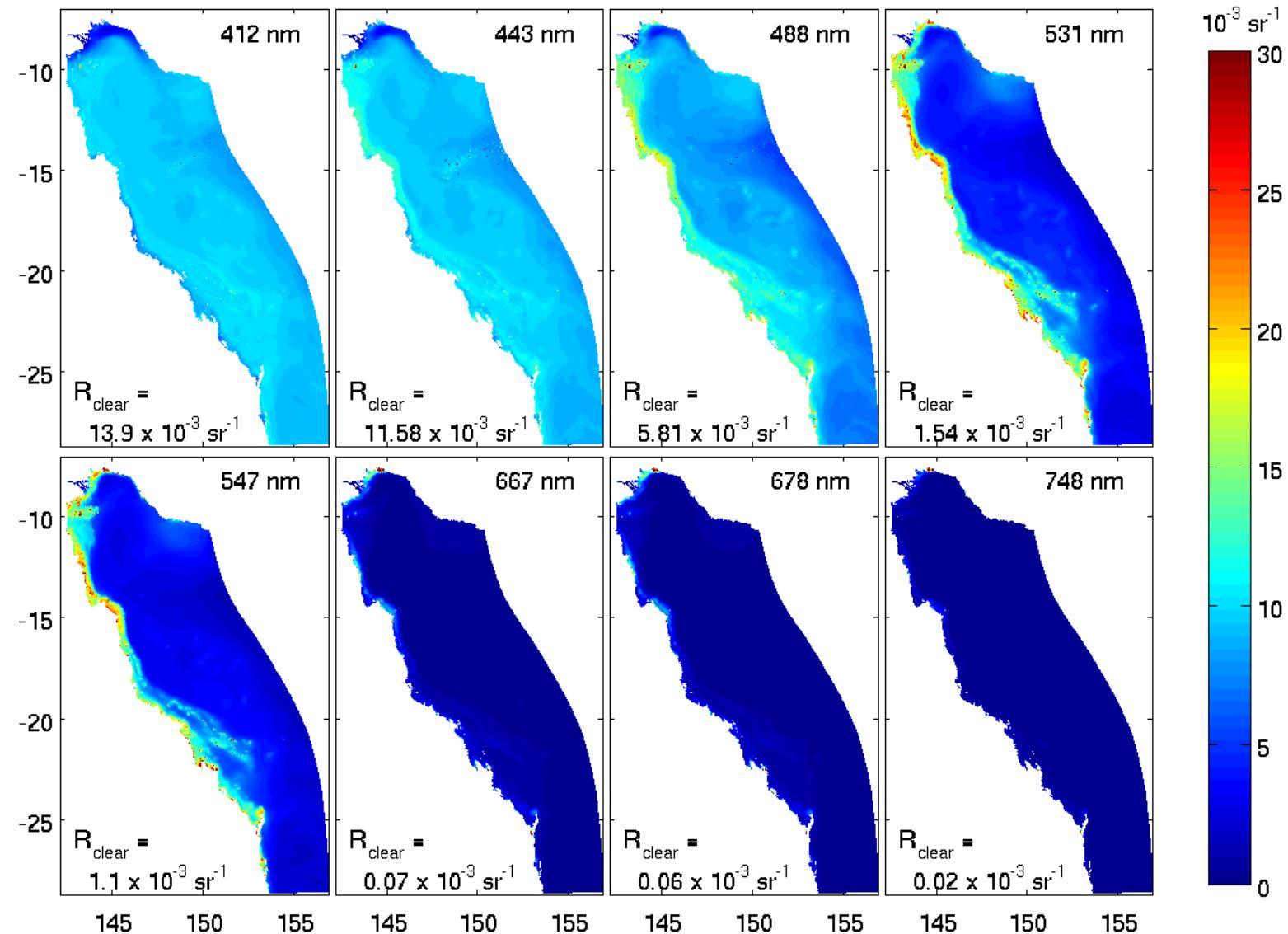
- During integration 23 wavelengths for the forcing of photosynthesis. Wavelength more resolved at pigment peak (430,440, 450 resolved).
- At output steps, light calculated at the MODIS 11 bands.
- Spectral-resolution is important because:
 - Photosynthetic plant response varies with spectral distribution.
 - Parameterisation of absorption characteristics more precise when done at a waveband.
 - Comparison with observation more direct if wavelength-specific.



- Mean simulated remote-sensing reflectance for 2013 at the 8 MODIS ocean colour bands.

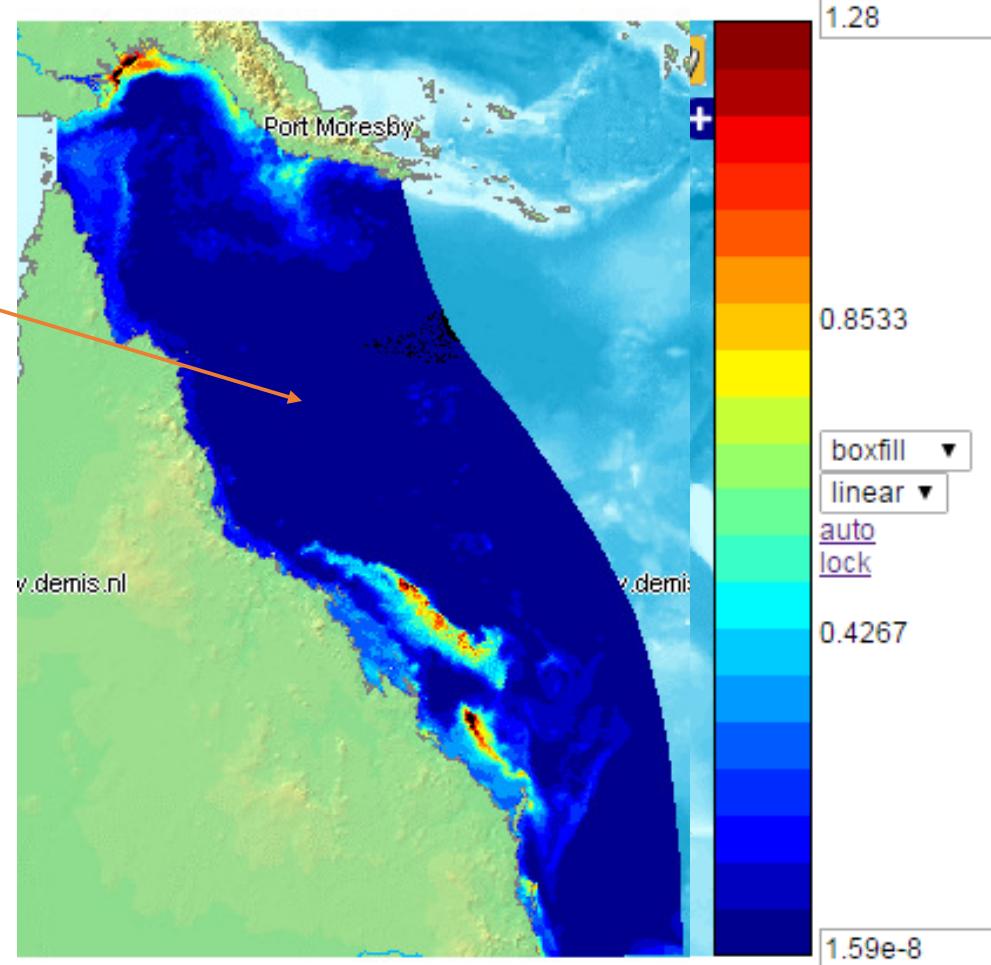
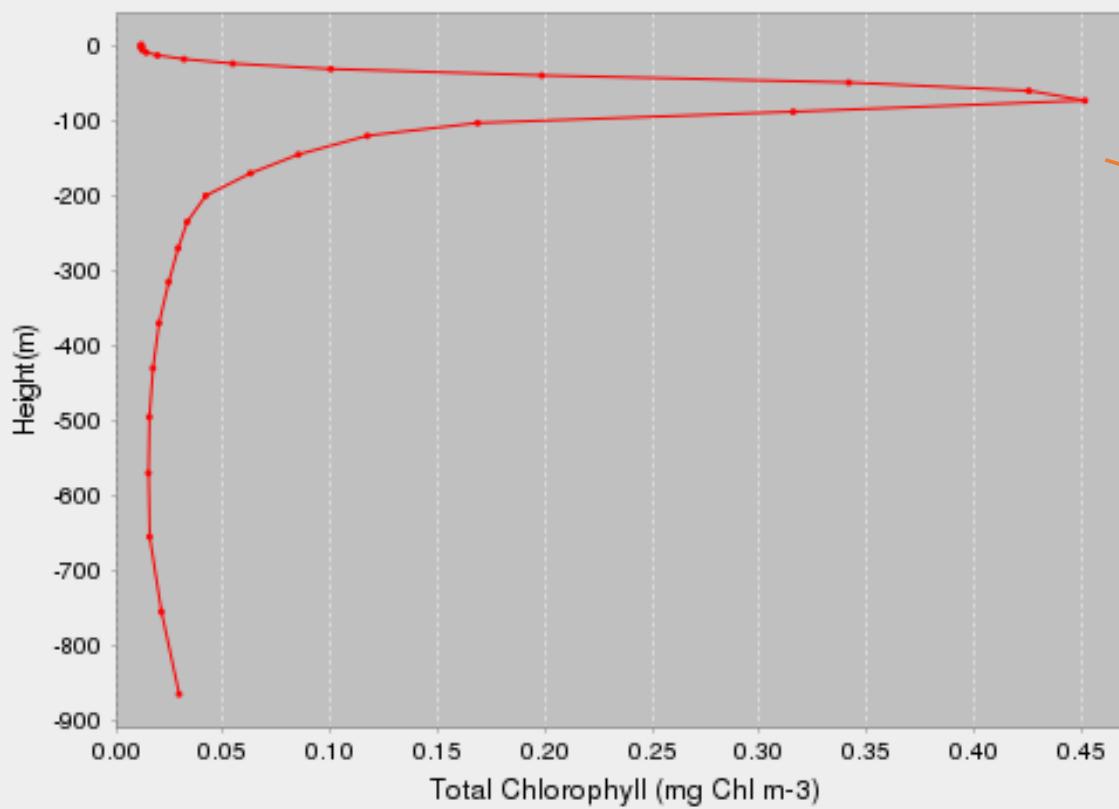
- More reflectance at shorter wavelengths.

- On a relatively clear day 1 million plus data points per wavelength, up to twice a day



Lets look at chlorophyll through on-line data.

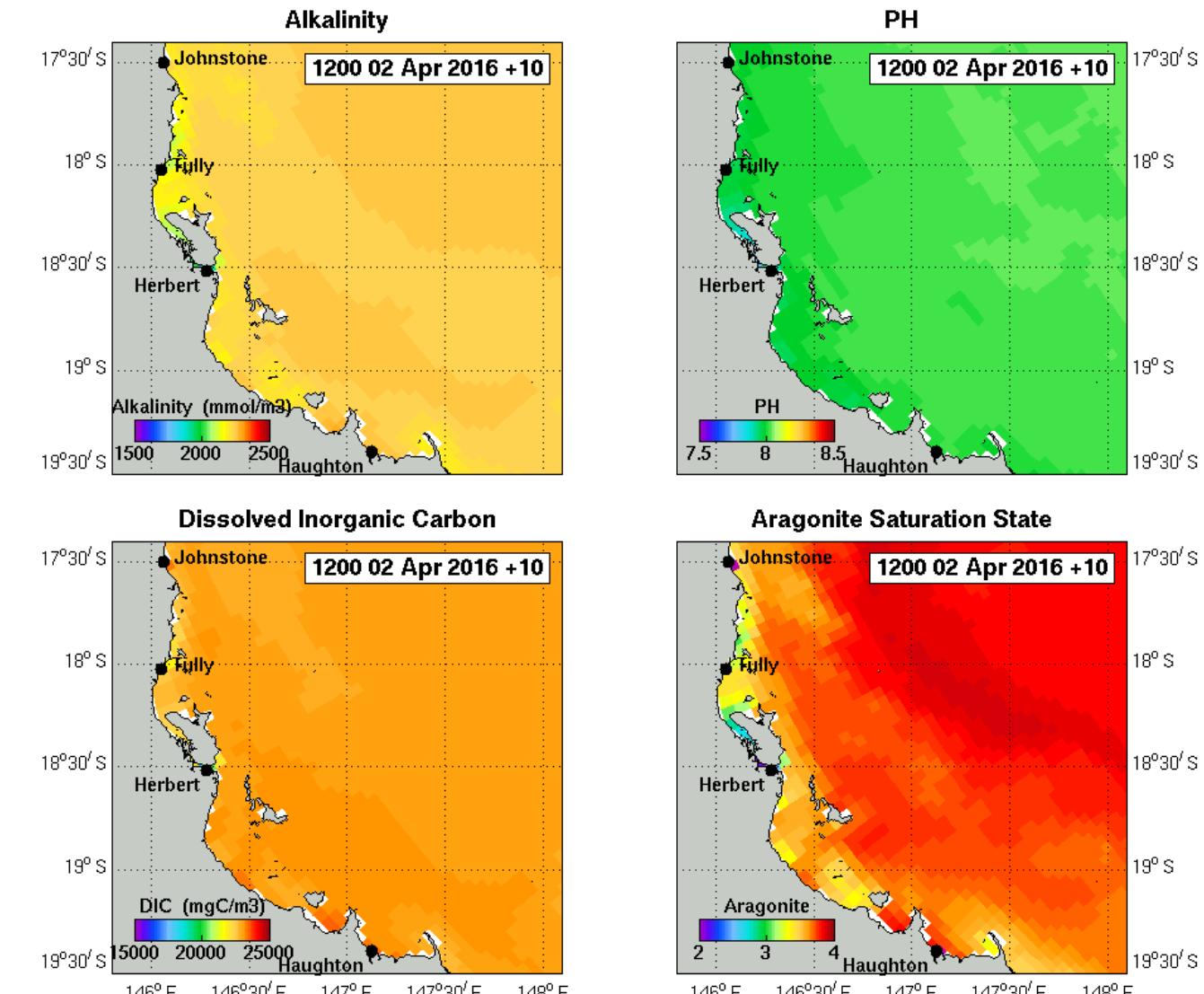
Profile of Total Chlorophyll at 148.8495844355469E, 17.555808432499997S at 2014-12-01T02:00:00.000Z



Carbon chemistry.

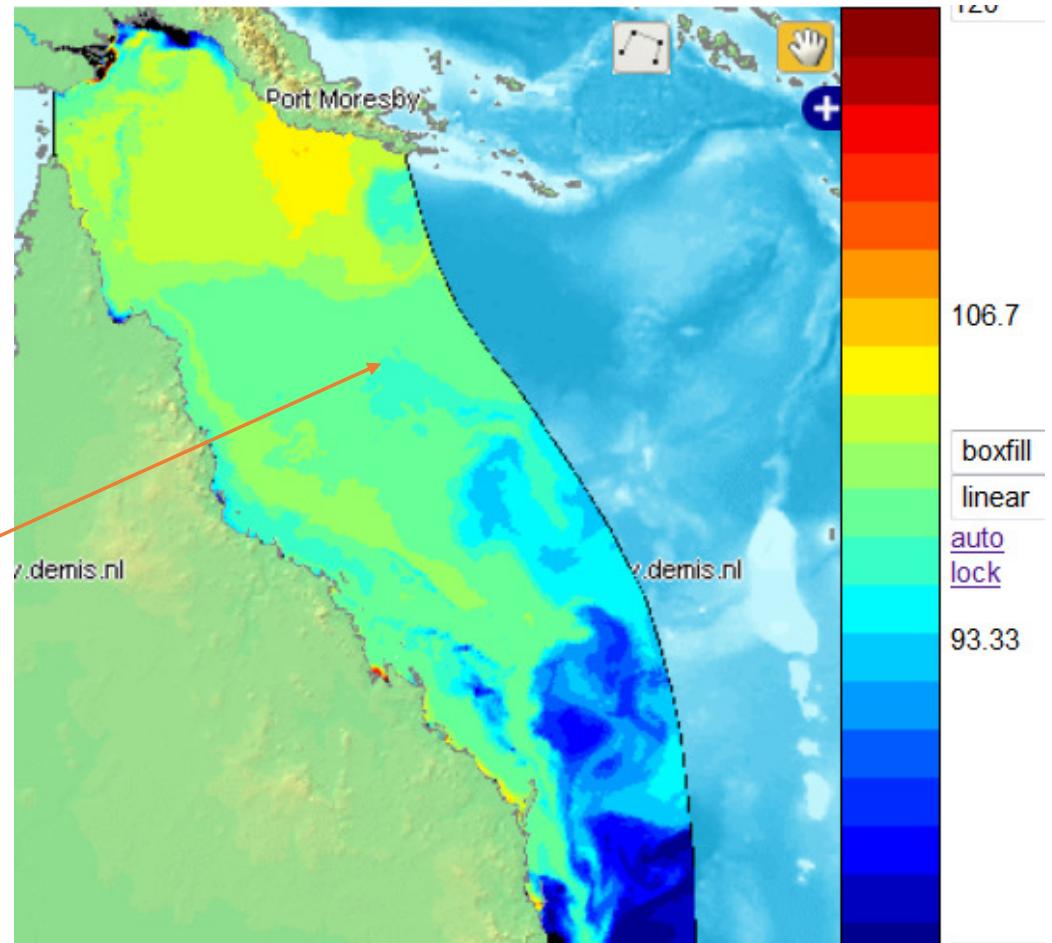
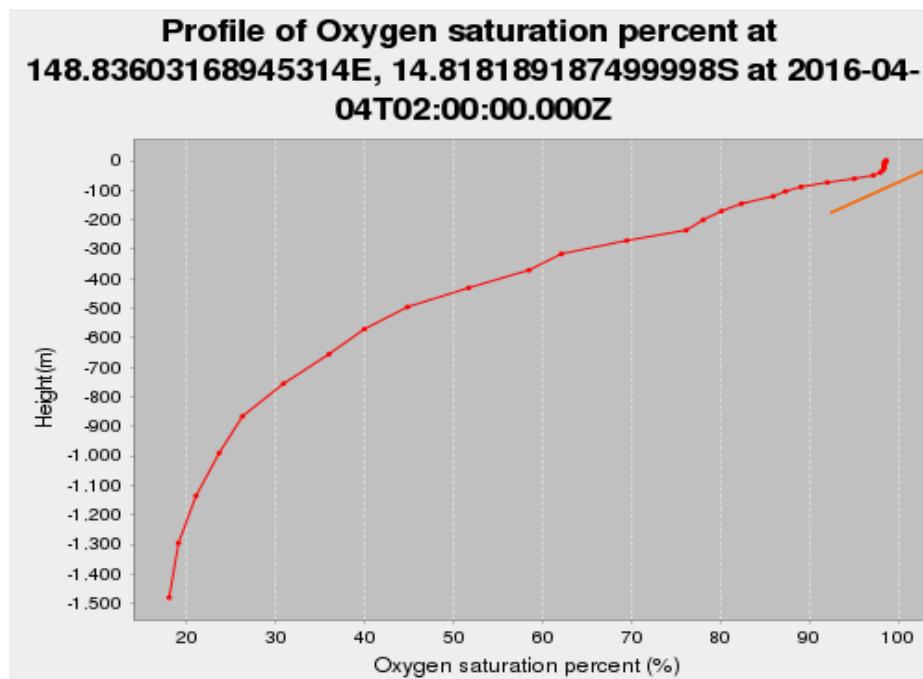
- Model uses OCMIP carbon chemistry routines to calculate pH, aragonite saturation, and air-sea fluxes from the total alkalinity, A_T , and dissolved inorganic carbon concentration, C_T , as well as temperature, salinity and wind speed.

GBR NEAR REAL-TIME BIOGEOCHEMICAL MODELLING



Oxygen dynamics.

- Model uses saturation state calculations as well as wind speed to determine oxygen concentrations, and percent saturation.



Seagrass model

- Two species model – *Zostera*-like and *Halophila*-like.
- Nutrient uptake from multiple layers of sediments.
- New formulation of relationship between % cover and benthic biomass.
- Translocation between roots and leaves.

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A biophysical representation of seagrass growth for application in a complex shallow-water biogeochemical model



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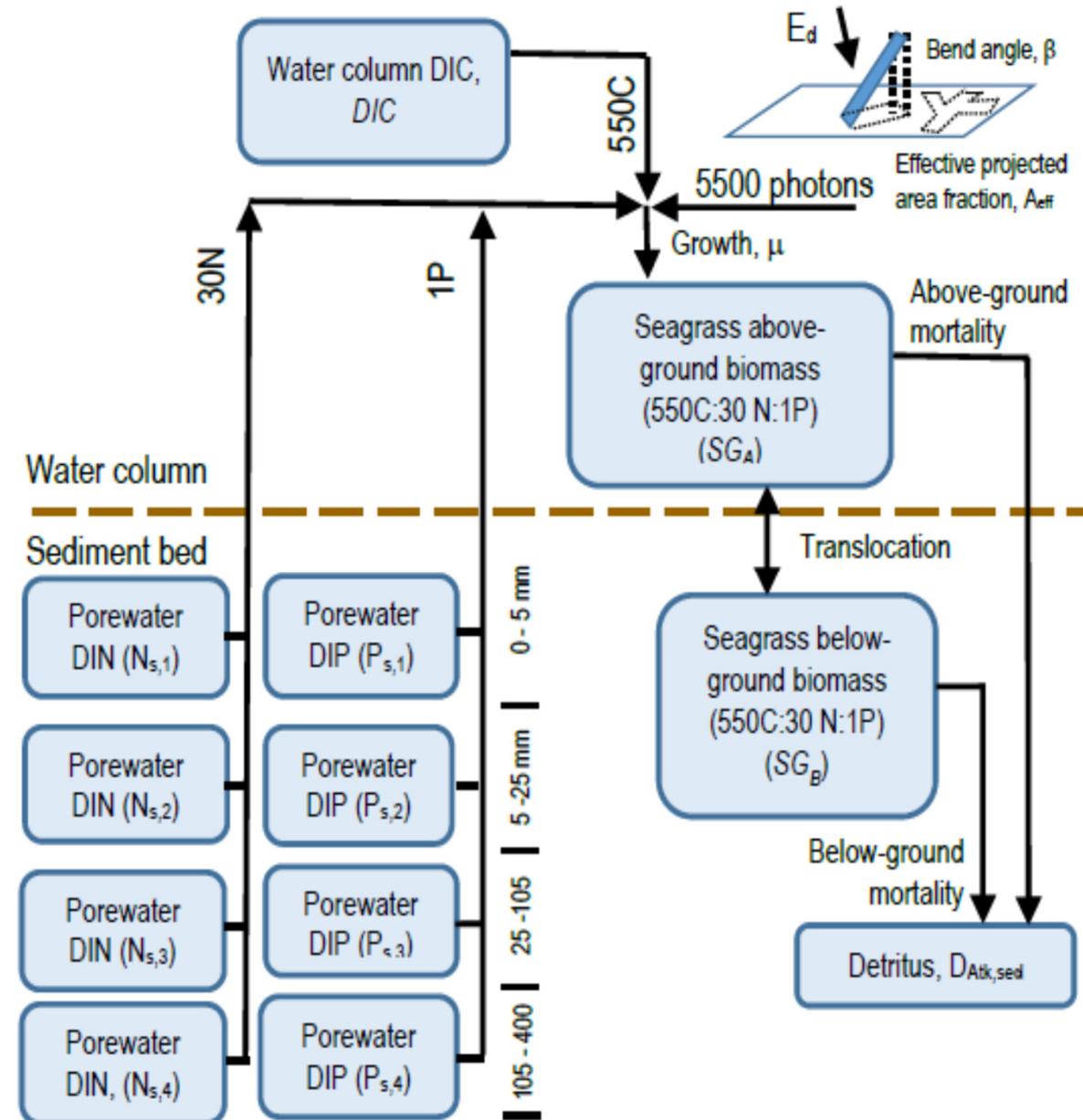
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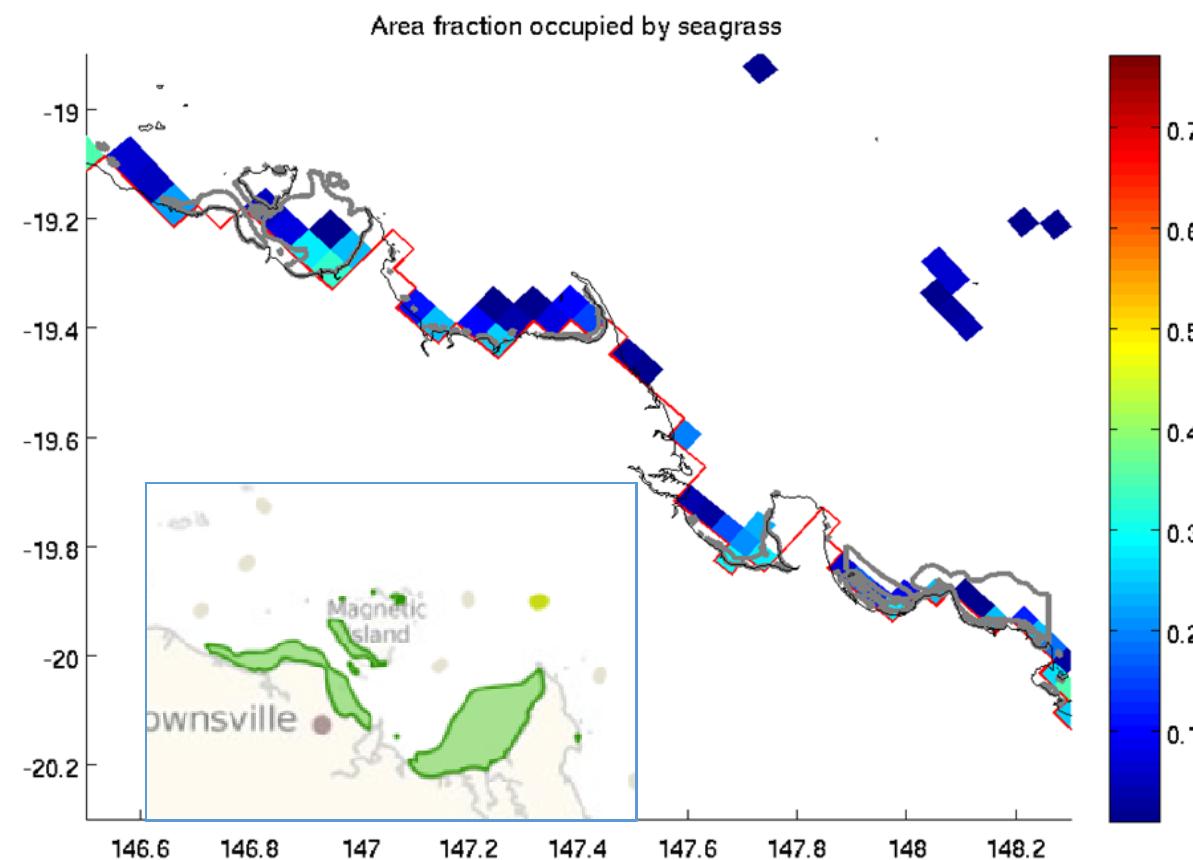
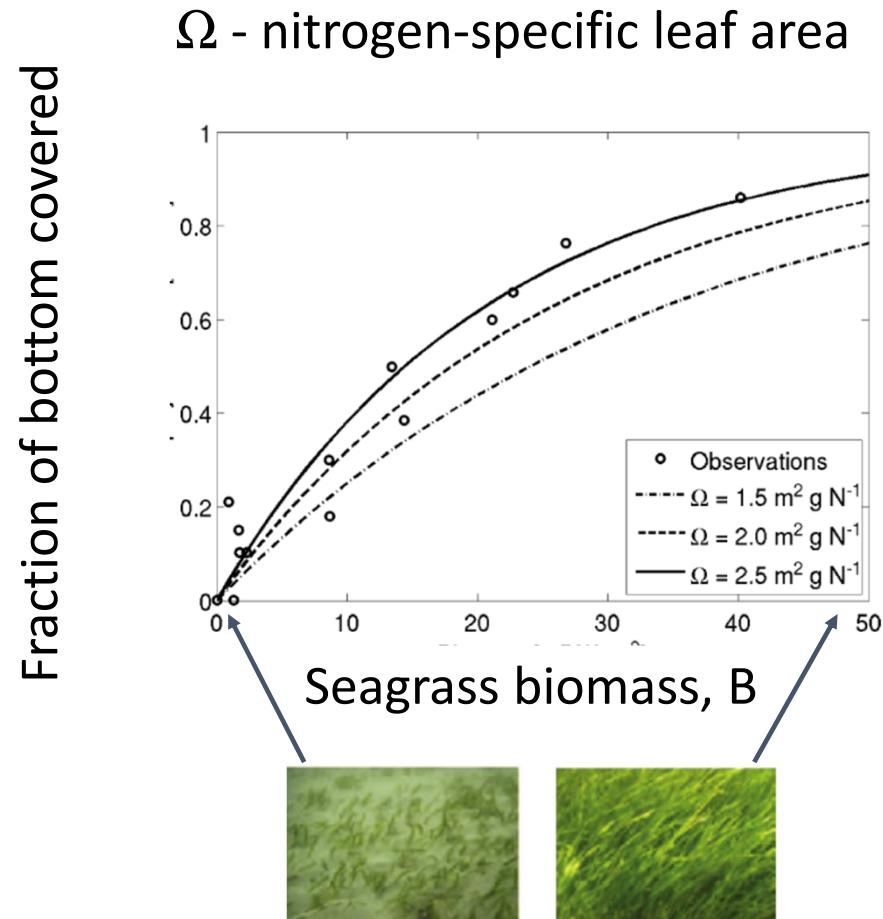
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^e Centre for Tidal Water & Aquatic Ecosystem Research, James Cook University, Townsville, Australia

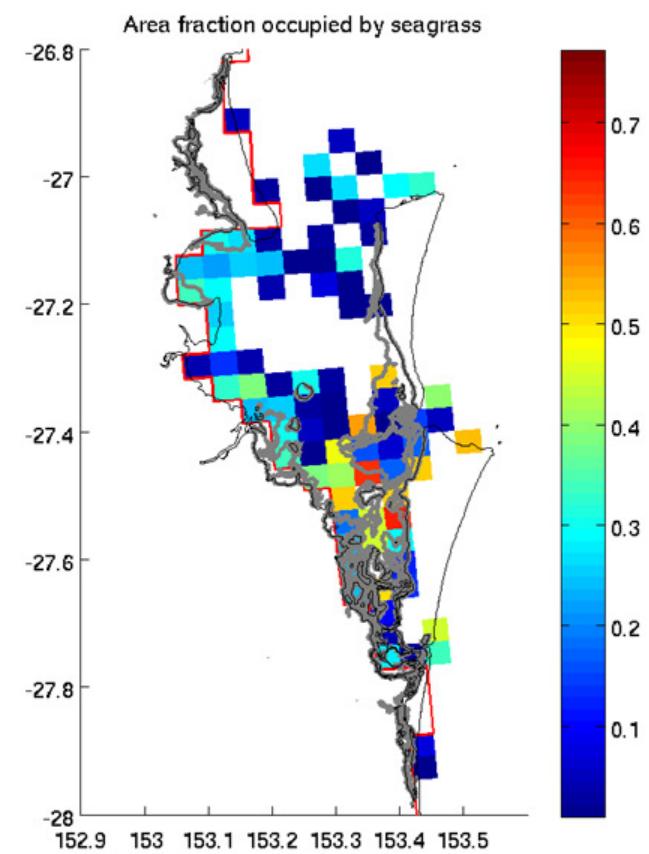
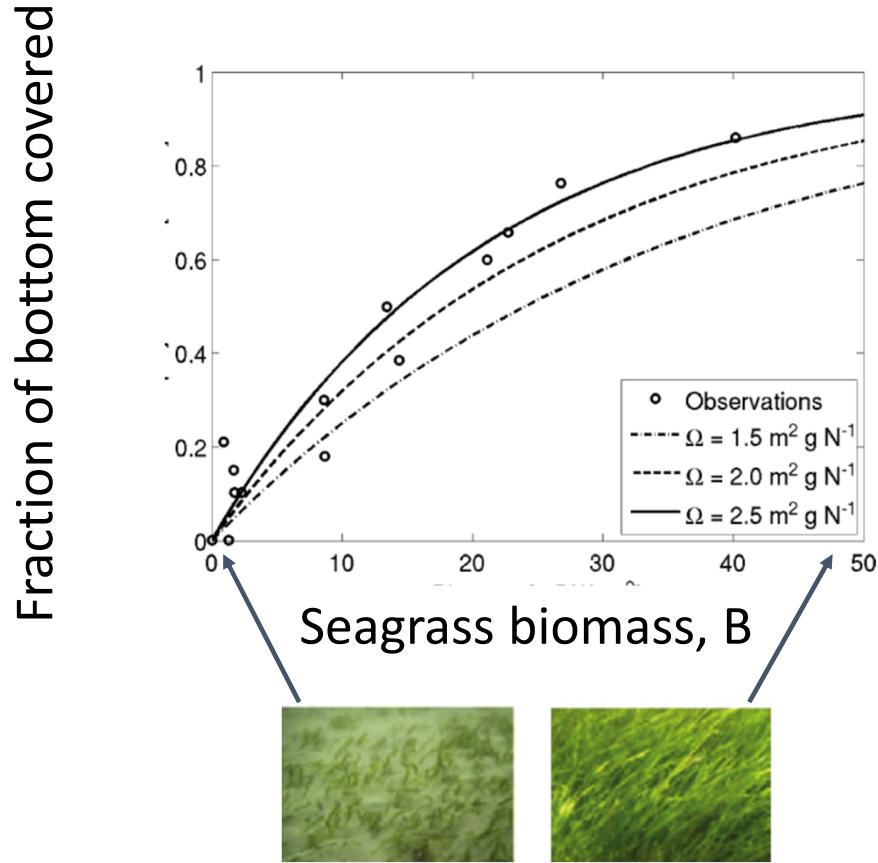


Seagrass behaviour - % of the bottom covered.



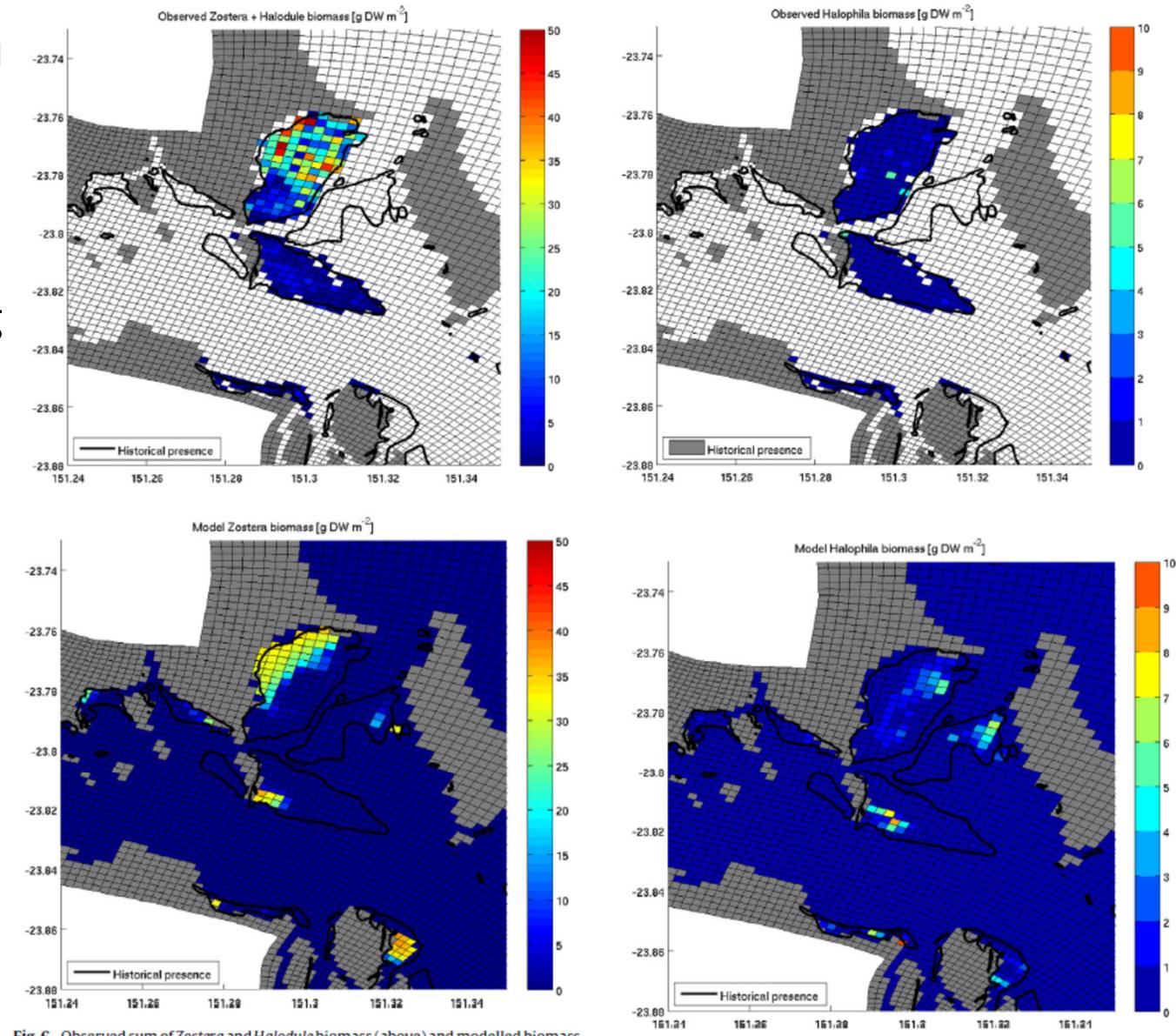
Seagrass behaviour - % of the bottom covered.

Ω - nitrogen-specific leaf area



Seagrass behaviour in Gladstone Harbour.

- Spectrally-resolved daily-varying light surface light field
- Seagrass depth varies with changing tides.
- Application of laboratory and field observation of compensation scale irradiance to determine respiration rates.



Thank you to JCU team for obs.

Coral model

- Coral biomass represented by three variables, host biomass, zooxanthellae biomass and chlorophyll content of zooxanthellae.
- Processes include:
 - coral host growth
 - zooxanthellae growth
 - zooxanthellae chlorophyll synthesis and photosynthesis
 - organic matter uptake (host)
 - inorganic matter uptake (zooxanthellae),
 - translocation of organic matter from zooxanthellae to host.
 - calcification (alkalinity and DIC uptake).

