

**I don't know, are you sure we want to do this?
Sea level adaptation decisions under uncertainty**

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Key Points:

- Decisions on adaptation measures need to take careful account of all uncertainties
 - Modeling local sea level rise is essential
 - Decisions under uncertainty do not have to be based on precise costs

12 **Abstract**

13 Sea level rise has serious consequences for harbor infrastructure, storm drains and sewer
 14 systems, and many other issues. Adapting to sea level rise requires comparing different
 15 possible adaptation strategies, comparing the cost of different actions (including no ac-
 16 tion), and assessing at what point in time the chosen strategy should be implemented. All
 17 these decisions must be made under considerable uncertainty—in the amount of sea level
 18 rise, in the cost of adaptation actions, and in the cost of no action. We develop two illus-
 19 trative examples: for Bergen on Norway’s west coast and for Esbjerg on the west coast of
 20 Denmark. Different components of uncertainty are visualized. We show that failing to take
 21 uncertainty into account can increase the projected damage costs by an order of magni-
 22 tude.

23 **1 Introduction**

24 The potential impact of climate change on local sea level, yielding effects such as
 25 frequent flooding, inundation and backflow of storm drainage and sewer systems, destruc-
 26 tive erosion and contamination of wetlands and other habitats, requires city planners to
 27 make decisions in the presence of substantial uncertainty.

28 As adaptation decision-making is an ongoing process of weighing and choosing
 29 which measures should be taken at which moment in time [Hallegatte *et al.*, 2012] adap-
 30 tive planning methods need to support decisions in the short term, while considering long-
 31 term developments. Challenges of adaptation decision-making under uncertainty relate
 32 to the incorporation of spatial, inter-temporal and flexibility aspects of adaptation priori-
 33 ties [Fankhauser and Soare, 2013], and the linkage with specific characteristics of sectors
 34 and contexts [Bisaro *et al.*, 2016; Hinkel and Bisaro, 2016]. Several economic decision
 35 support tools and methods exist for adaptation assessment under uncertainty [e.g. Chamb-
 36 wera *et al.*, 2014; Wilby and Dessai, 2010; Walker *et al.*, 2013]. However, Watkiss *et al.*
 37 [2015] conclude that these tools are very resource intensive and complex in the context of
 38 long-term adaptation investment decisions and call for the development of “light touch”
 39 approaches to better support local adaptation making.

40 In this paper we employ light touch decision tools to demonstrate the importance of
 41 combining projections of sea level rise and flood damages alongside a detailed quantifi-
 42 cation of both hydrologic and economic uncertainties in the context of real-life decision-
 43 problems experienced by stakeholders and authorities in two northern European cities,
 44 Bergen in Norway and Esbjerg in Denmark, see Figure 1. Based on communications with
 45 local end-users we highlight the value of taking into account uncertainty through two sim-
 46 plified and complementary case studies, where in the first one planners want to know how
 47 early they should implement costly adaptation measures, whereas in the second case the
 48 aim is to highlight the risk of flooding in coastal areas, e.g. in order to prioritize future
 49 adaptation actions and investments. In both cases we show that embracing the uncertain-
 50 ties derived from economic and hydrologic models is absolutely crucial in order to answer
 51 the question of “are we sure we want to do this?”.

52 The Norwegian city of Bergen is the capital of Hordaland County. The city center is
 53 located on Byfjorden, and is surrounded by mountains. It has the largest port in Norway,
 54 both in terms of freight and passengers. The historic harbor area, Bryggen, is the only
 55 Hanseatic trade center remaining in its original style, and has been declared a UNESCO
 56 World Heritage site¹. Bryggen is regularly flooded at extreme tides, and it is feared that
 57 as sea levels rise, floods will become a major problem in other parts of Bergen as well
 58 [Grieg Foundation, 2009].

59 ¹ See <http://whc.unesco.org/en/list/59>.



52 **Figure 1.** Terrain maps of central Bergen, Norway (left) and Esbjerg, Denmark (right).

60 The municipality of Bergen has, in cooperation with private actors, analyzed sev-
 61 eral possible adaptation measures against sea level rise. The measures range from an outer
 62 barrier that would protect the entire metropolitan area to various protection measures of
 63 limited areas in the inner harbor [Grieg Foundation, 2009]. While the viability of the con-
 64 structions and the associated construction costs have been carefully analyzed, the optimal
 65 timing of potential adaptation measures and the effects of the associated uncertainties have
 66 yet to be investigated. We perform such an analysis where we consider uncertainty in pro-
 67 jected sea level rise, damage costs and the effects of sea level rise on changes in damage
 68 costs.

69 Esbjerg, on the southwest coast of Jutland, is the fifth-largest city in Denmark and
 70 the largest urban area in the region. The city hosts one of the largest harbours in Den-
 71 mark, which serves as a focal point for offshore activities in the North Sea, including the
 72 continued development of offshore wind power and extensive activities related to the ex-
 73 traction of oil and gas. As a result critical infrastructures and commercial buildings figure
 74 prominently in the coastal zone. Esbjerg is frequently subject to substantial storms and
 75 storm surges, causing severe flooding of the harbour and the city. The highest ever since
 76 1874 was recorded in 1981, where the harbour was completely flooded and the water level
 77 reached a record-high 433 cm above the norm, causing massive economic losses. More
 78 generally, storm surges causing water levels in Esbjerg to rise to between 2 and 3 metres
 79 have quadrupled over the last four decades according to local records, whereas half of the
 80 most severe events have taken place since 1975.

81 As in the case of Bergen, sea level rise caused by climate change is expected to
 82 compound these risks, alongside parallel threats caused by increased risks of pluvial flood-
 83 ing and rising ground water levels in Esbjerg. The municipality recently adopted its cli-
 84 mate adaptation plan, which in its first phase is aimed at identifying present and future
 85 flood-prone areas, e.g. to avoid urban development into such areas, to limit damages to
 86 buildings of high societal or cultural value, and to pave the way for implementing cost-
 87 effective adaptation measures in the second phase of the plan.

88 The remainder of the paper is organized as follows. Section 2 describes our ap-
 89 proach to projecting sea level. In sections 3 and 4 we describe the type of decision prob-
 90 lems that we are going to attack. We apply these tools to sea level projections for Bergen,
 91 Norway, and for Esbjerg, Denmark in section 5. In section 6 we demonstrate the conse-

92 consequences of ignoring the uncertainty in the projections, and the paper is closed with a sum-
 93 mary and discussion in section 7..

94 2 Sea level projections

95 We project local sea level changes by modeling two processes, the relationship be-
 96 tween global temperature and global sea level, and the relationship between global sea
 97 level and local sea level.

98 2.1 Global sea level

99 Most climate models do not explicitly provide sea level as an output of the calcu-
 100 lations. Rather, the IPCC AR5 report [Stocker *et al.*, 2013, ch. 13] combines the heat ex-
 101 pansion of the ocean with temperature forced models for glacial melt, Greenland ice melt,
 102 and Antarctic ice melt and with land rise due to rebound from the last ice age. Judging
 103 from the supplementary material to Stocker *et al.* [2013, ch. 13], the uncertainty assess-
 104 ment is only based on the spread of the ensemble of temperature projections, not on the
 105 additional uncertainty in the ice models used.

106 We will instead use the empirical approach of Rahmstorf and collaborators [Rahm-
 107 storf, 2007; Rahmstorf *et al.*, 2011], employing the statistical modeling of Bolin *et al.*
 108 [2014] to relate global annual mean temperature anomalies to global mean sea level anom-
 109 alies.

110 We then apply the estimated historical relationship to projected temperatures from
 111 the CMIP5 experiment [Taylor *et al.*, 2012] to obtain projected global annual mean sea
 112 level, taking into account the uncertainty in the statistical model as well as the spread of
 113 the temperature projection ensemble (see subsection 2.3). For the i 'th temperature projec-
 114 tion T_t^i we estimate the corresponding global mean sea level as

$$115 H_t^{gl,i} = \int_{t_0}^t \hat{a}(T_u^i - \hat{T}_0) du + \varsigma_t,$$

116 where \hat{a} and \hat{T}_0 are regression parameters of observed global sea level rise on observed
 117 global temperature and ς_t the integrated time series regression error.

118 2.2 Local sea level

119 In order to get from global sea level projections to local ones, it is important to note
 120 that sea level rise is not uniform over the globe. Glacial and land ice melting affect the
 121 local sea level differently depending on where the melted ice is located. Another major
 122 effect in Fennoscandia is the land rise due to isostatic rebound from the glaciers of the
 123 last ice age. Again, we will use historical data to relate global sea level to isostatically
 124 corrected local sea level using a time series regression model. The local sea level projec-
 125 tions are then obtained by first relating projected temperature to global sea level, and then
 126 relating the global sea level to the local one. Each climate model temperature projection
 127 yields a different local sea level projection. The local sea level projection based on the i 'th
 128 climate model for years beyond 2000 is estimated as

$$129 H_t^{loc,i} + \gamma(t - 2000) = \hat{b}H_t^{gl,i} + \varepsilon_t,$$

130 where γ is the annual land rise rate, t denotes year, \hat{b} is the regression coefficient relating
 131 global to local sea level and the ε_t are Gaussian errors..

132 **2.3 Uncertainty assessment**

133 Following the approach of *Guttorp et al.* [2014] we assess the uncertainty in the lo-
 134 cal sea level projections taking into account the variability between the climate projections
 135 used, the uncertainties in the regressions of global mean temperature on global mean sea
 136 level and of global on local sea level. We express the sea level projection uncertainty in
 137 terms of a confidence band that is simultaneously of the intended level for all projection
 138 years. This allows us, for example, to get a confidence band for the years when a given
 139 sea level rise is obtained.

140 **2.4 Limitations of the sea level projections**

141 The main assumption is using historical relationships in statistical projections of
 142 the type used in this paper is that there is no major change in how temperature relates
 143 to sea level, globally and locally. Among the factors that may invalidate this approach
 144 are changes in water storage on land (in essence removing water from the oceans), ex-
 145 cessive siphoning of groundwater (resulting in land subsidence), changes in the rates of
 146 glacial and land ice melt, and changes in Earth's gravitational field due to transfer of mass
 147 from land ice to ocean water. For example, the rate of ice melt on Greenland may sud-
 148 denly increase substantially due to intense warming of both air and sea water [*Bamber*
 149 and *Aspinall*, 2013]. A recent paper [*Jevrejeva et al.*, 2016] indicates that the upper tails
 150 of sea level rise may be substantially higher when taking into account expert assessment
 151 of land ice melting. Our current climate models are not able to resolve the ice processes
 152 sufficiently to include such so called tipping points into the projections. Also, the IPCC
 153 scenarios [*van Vuuren et al.*, 2011] do not include changes in water usage (cf. *Wada et al.*
 154 [2012]).

155 **3 Timing of adaptation measures**

156 There is an increasing need for more detailed economic analysis, including simple
 157 methods and tools for assessment of options, especially since as Downing (2012) recog-
 158 nizes adaptation is moving from theory to practice and practitioners try to deal with how
 159 to begin adapting. This leads to an increasing need and interest in the appraisal of options.

160 Several economic decision support tools and methods exist for adaptation assessment
 161 under uncertainty. Robust decision-making approaches are able to better incorporate un-
 162 certainty and a broad range of climate scenarios to capture as much of the uncertainty on
 163 future climates as possible (Dittrich et al. 2016). These approaches can be classified ac-
 164 cording to a science-first or policy-first approach. The former has a “predict-then-act” foundation,
 165 which starts with for example probabilistic climate projections and impact as-
 166 sessments, not linked to any specific adaptation choices (Jones et al., 2014). The latter
 167 starts out with the formulated adaptation plans and not impacts and their functioning is
 168 tested against different future projections (Dittrich et al. 2016).

169 We take a policy-first approach, in which we test a current adaptation plan against
 170 different sea level and damage projections and the inclusion of different sources of un-
 171 certainty. We focus on what this implies for the timing of adaptation measures and the
 172 implications of including uncertainty. In particular, we employ a probabilistic extension of
 173 the framework described by [Fankhauser et al., 1999] in which we obtain a probabilistic
 174 distribution for the net present value damage in a given year for various adaptation op-
 175 tions. The probabilistic distribution is constructed by considering uncertainty in the local
 176 sea level projections, in the annual damage costs, and in the effect of changes in sea level
 177 on the annual damage costs.

178 **3.1 Annual damage costs**

179 We model the distribution of annual damage, F_{d,t_0} , for the year $t_0 = 2015$ by the
 180 three parameter Burr distribution [Burr, 1942] with density

$$181 \quad f_{d,t_0}(x) = \frac{\alpha\gamma(x/\theta)^\gamma}{x[1+(x/\theta)^\gamma]^{\alpha+1}} \quad (1)$$

182 for $x > 0$, where α and γ are shape parameters with $\alpha, \gamma > 0$, and $\theta > 0$ is a scale parameter.
 183 The Burr distribution has a heavy upper tail and is commonly used to model damage
 184 loss, see e.g. *Klugman et al.* [2012]. The parameters of the distribution are estimated using
 185 historical data for annual storm surge damage. Data prior to 2015 are adjusted to the 2015
 186 level using the consumer price index. After adjustment, we assume stationarity over the
 187 period and independence between years.

188 Under a constant sea level, we can obtain a sample trajectory $\{d_{t_1}, d_{t_2}, \dots, d_{t_{85}}\}$ of
 189 future annual damages for $t_1 = 2016, \dots, t_{85} = 2100$ by drawing 85 i.i.d. values from
 190 the estimated distribution \hat{F}_{d,t_0} . By repeating this process J times, we obtain an empirical
 191 damage distribution for each future year t_i given by

$$192 \quad \hat{F}_{d,t_i}(x) = \frac{1}{J} \sum_{j=1}^J \mathbb{I}\left\{ \frac{d_{t_i}^{(j)}}{\prod_{l \leq i} (1+r_{t_l})} \leq x \right\}$$

193 for $i = 1, \dots, 85$, where r_{t_i} is the discount rate for year t_i . Alternatively, we obtain an
 194 empirical distribution of the total damage over the period 2016 – 2100 by considering

$$195 \quad d_{\text{total}}^{(j)} = \sum_{i=1}^{85} \frac{d_{t_i}^{(j)}}{\prod_{l \leq i} (1+r_{t_l})}$$

196 and similarly for the cumulative damage.

197 **3.2 Effect of changes in sea level**

198 We assume that changes in sea level have a multiplicative effect on the annual dam-
 199 age cost. That is, for a sea level anomaly s_{t_i} in year $t_i > 2015$ compared to the 2015
 200 level, the annual damage cost becomes

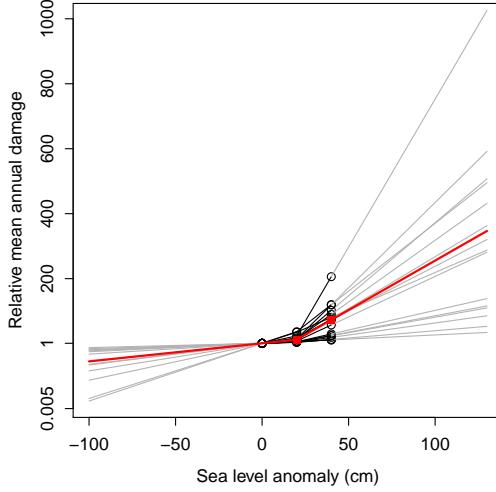
$$201 \quad g(s_{t_i} | \beta) d_{t_i},$$

202 where $g(\cdot | \beta)$ is a monotonic positive function with parameter vector β such that $g(s | \beta) >$
 203 1 for $s > 0$ and $g(s | \beta) < 1$ for $s < 0$. *Hallegatte et al.* [2013] estimate a similar ef-
 204 fect function valid in 2050 for $s \in \{0, 20, 40\}$ for 136 coastal cities. Here, we use their
 205 results for 15 European cities: Amsterdam, Athens, Barcelona, Dublin, Glasgow, Hamburg,
 206 Helsinki, Copenhagen, Lisbon, London, Marseilles, Naples, Porto, Rotterdam and Stock-
 207 holm. To obtain a city-specific effect function for a large range of sea level anomalies we
 208 employ a linear extrapolation as shown in Figure 2. We then obtain a sample of effect
 209 functions $\{g(\cdot | \beta^{(j)})\}_{j=1}^J$ by sampling with replacement from this ensemble of trajectories
 210 with all 15 ensemble members considered equally probable.

214 Let further $\{s_{t_1}^{(j)}, \dots, s_{t_{85}}^{(j)}\}_{j=1}^J$ denote a sample of projections for annual sea level
 215 anomalies compared to the 2015 value. An empirical damage distribution for the future
 216 year t_i that accounts for uncertainty in damage, sea level rise and its effect on the damage
 217 is then given by

$$218 \quad \hat{F}_{d,t_i}^s(x) = \frac{1}{J} \sum_{j=1}^J \mathbb{I}\left\{ \frac{g(s_{t_i}^{(j)} | \beta^{(j)}) d_{t_i}^{(j)}}{\prod_{l \leq i} (1+r_{t_l})} \leq x \right\}. \quad (2)$$

219 The distribution in (2) describes the projected damage distribution with no adap-
 220 tation measures. In addition, we can incorporate an adaptation measure of cost C that pro-
 221 tects against K cm of increased sea level from year t_k onward. This results in a damage



211 **Figure 2.** Relative change in mean annual damage as a function of sea level rise for 15 European cities as
212 estimated by *Hallegatte et al.* [2013] (black circles) with linearly extrapolated values indicated by gray lines.
213 The median change and the corresponding extrapolation are indicated in red.

222 distribution given by

$$223 \quad \hat{F}_{d,t_i}^{s,a_k}(x) = \begin{cases} \frac{1}{J} \sum_{j=1}^J \mathbb{1} \left\{ \frac{g(s_{t_i}^{(j)} | \beta^{(j)}) d_{t_i}^{(j)}}{\prod_{l \leq i} (1+r_{t_l})} \leq x \right\}, & t_i < t_k \\ \frac{1}{J} \sum_{j=1}^J \mathbb{1} \left\{ \frac{g(s_{t_i}^{(j)} - K | \beta^{(j)}) d_{t_i}^{(j)} + C}{\prod_{l \leq i} (1+r_{t_l})} \leq x \right\}, & t_i = t_k \\ \frac{1}{J} \sum_{j=1}^J \mathbb{1} \left\{ \frac{g(s_{t_i}^{(j)} - K | \beta^{(j)}) d_{t_i}^{(j)}}{\prod_{l \leq i} (1+r_{t_l})} \leq x \right\}, & t_i > t_k. \end{cases}$$

224 **3.3 Limitations of the decision framework**

225 The main limitation of this light touch decision framework is that we have significantly simplified the assessment of the effect of sea level rise on the damage costs. In
226 particular, the linear extrapolation of the results reported in *Hallegatte et al.* [2013] might
227 provide a conservative estimate of the effect of extreme sea level rise. However, with only
228 two data points, extrapolation approaches such as a power law or exponential growth seem
229 unfeasible.

231 Alternatively, a modeling framework similar to that of *Hallegatte et al.* [2013] could
232 be applied directly to a larger range of potential changes in sea level. The elements of
233 such a framework might include an appropriate social discount rate, valuing environmental
234 goods in monetary terms, incorporate socio-economic assumptions and long-term policy
235 goals of decision makers, as well as that climate change is often not the only driver that
236 decision makers should consider, therefore costs and benefits should be studied in a wider
237 context (Dittrich et al., 2016).

238 Our framework simplifies the cost and effect of an adaptation option during con-
239 struction in that we assume no effect until the construction is finished with all the con-
240 struction cost falling in the last year of the construction. Especially for larger construc-
241 tions, these assumptions might need to be modified. Additionally, we have not specifically
242 accounted for potential changes in storm surge patterns.

243 **4 Danish decision framework**

244 In 2014 the municipality of Esbjerg adopted its climate adaptation plan ([ref](#)), which
 245 will cover the period from 2014-2026, and which aims to reduce the risk of flooding
 246 caused by storm surges, heavy rainfall and rising ground water levels, respectively. In the
 247 following (and based on interviews with end-users from the municipality) we will focus
 248 on the harbour of Esbjerg and the nearby coastal areas, and we will only consider the hazard
 249 of coastal flooding, though evidently in some areas the flood risk is compounded.

250 The initial scoping of the climate adaptation plan in Esbjerg includes a preliminary
 251 value and risk mapping, considering critical infrastructure and buildings of high cultural
 252 and societal value as identified by the municipality, while informed by spatial floods maps
 253 for different scenarios corresponding to each of the three different kinds of hydrological
 254 events (sea level rise/storm surges, pluvial flooding and rising ground water levels). In
 255 terms of coastal floods the mapping considers only one type of storm surge, corresponding
 256 to a 20-year return event (based on historical storm surge statistics), and increased sea
 257 level due to climate change was not accounted for. We will extend these flood maps to
 258 also deal with 100-year return events and quantiles of the projected sea level rise.

259 The risk for any given map area as the probability of, e.g., a certain flood depth,
 260 is derived from a hydrological flood model (not including urban drainage system) times
 261 a valuation of the consequences, which—similar to the Bergen case—is essentially a loss
 262 function associating the flood depth with a measure of cost.

263 **5 Case studies**264 **5.1 Data**

265 The historical global mean temperature series is obtained from *Hansen et al.* [2001].
 266 Climate projections of global mean temperature are from the fifth climate model intercom-
 267 parison project, CMIP5 [*Taylor et al.*, 2012]. The global mean sea level series is obtained
 268 from *Church and White* [2011]. We use local tide gauge data from the Permanent Service
 269 for Mean Sea Level, UK, which is the worldwide repository for national sea level data.
 270 Glacial isostatic adjustment for Bergen is obtained from *Simpson et al.* [2014], and for Es-
 271 bjergr in personal communication from Peter Thejll at the Danish Meteorological Institute.

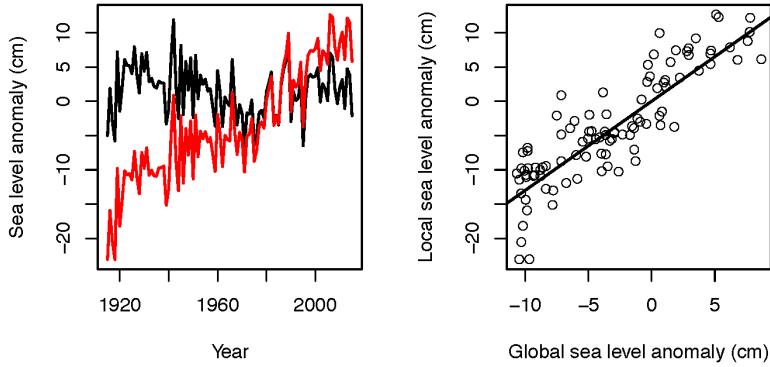
272 The Bergen monthly series is missing data for 62 months, including all of the years
 273 1942–43. To deal with occasional short stretches of missing data (at most one or two
 274 months) we use median polish replacement [*Mosteller and Tukey*, 1977] and then com-
 275 pute annual averages. For the years 1942-43, we use the average difference between
 276 Bergen and the average of all other Norwegian stations in 1940 and 1943 to estimate val-
 277 ues for 1941 and 1942, using the average of all other Norwegian stations corrected by the
 278 average difference.

279 The Esbjerg monthly series is missing data for 19 months. Here, too, we use median
 280 polish to fill in missing data and then compute annual averages.

281 Annual damage costs for the Bergen case study are obtained from the Norwegian
 282 Natural Perils Pool (NPP; data are available at [https://www.finansnorge.no/statistikk/
 283 skadeforsikring/Naturskadestatistikk-NASK/](https://www.finansnorge.no/statistikk/skadeforsikring/Naturskadestatistikk-NASK/)). The NPP data are available for
 284 the period 1980-2015 and are aggregated to a county level. For improved parameter es-
 285 timation, we include the data from Rogaland county which is the county directly south
 286 of Hordaland and with similar characteristics. We use a discount rate of 4% for the first
 287 40 years of the analysis, a rate of 3% for 40 to 75 years into the future and a rate of 2%
 288 beyond 75 years (cf. Section 5.8 of *Norwegian Ministry of Finance* [2012]).

289 5.2 Sea level rise in Bergen and Esbjerg

290 Figure 3 shows uncorrected and corrected Bergen sea level data, and the relationship
 291 between the corrected Bergen data and the global sea level data. The glacial isostatic ad-
 292 justment is 0.26 (0.07) cm/yr. The time series regression uses an ARMA(1,1)-model [Box
 293 and Jenkins, 1970], with AR parameter 0.82 (0.13), and MA parameter -0.61 (0.17). The
 294 regression slope is 1.30 (0.12).



295 **Figure 3.** The left figure shows raw (black) and glia-corrected (red) sea level data from Bergen, The right
 296 figure relates the glia-corrected Bergen sea level to the global sea level series of *Church and White* [2011]. The
 297 straight line is the time series regression line.

302 For the relationship between global annual mean temperature and global annual
 303 mean sea level rise we use the results from *Bolin et al.* [2014]. The left panel of figure 4
 304 shows the simultaneous 90 % confidence region for Bergen sea level rise relative to 1999
 305 under scenario RCP 8.5, which is the scenario Norwegian authorities recommend for plan-
 306 ning purposes.

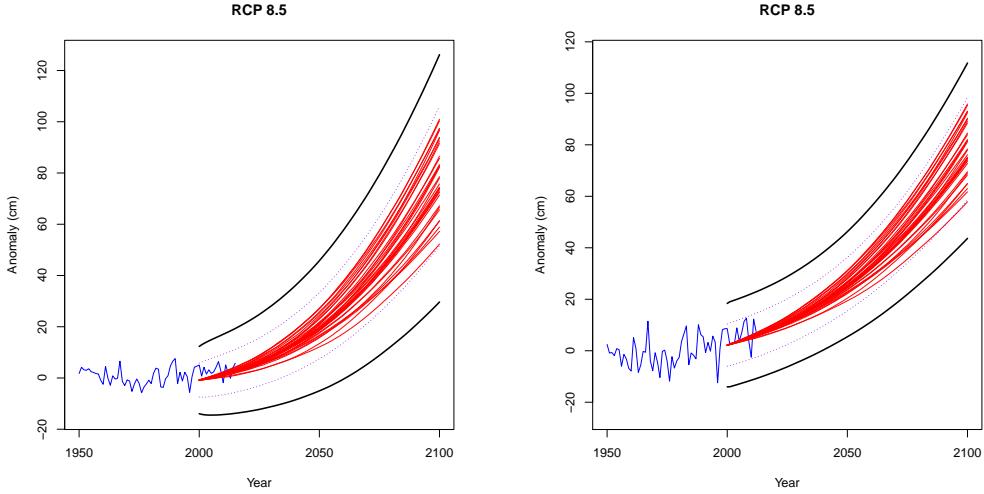
307 For Esbjerg, the glacial isostatic adjustment is 0.06 (0.03) cm/yr. The time series re-
 308 gression model relating glia-corrected local to global sea level is an MA(1) model with pa-
 309 rameter 0.17 (0.09). The regression slope is 1.02 (0.06). The right panel of figure 4 shows
 310 the simultaneous 90% confidence region for sea level rise relative to 1999 under scenario
 311 RCP 8.5.

312 5.3 Timing of adaptation measures in Bergen

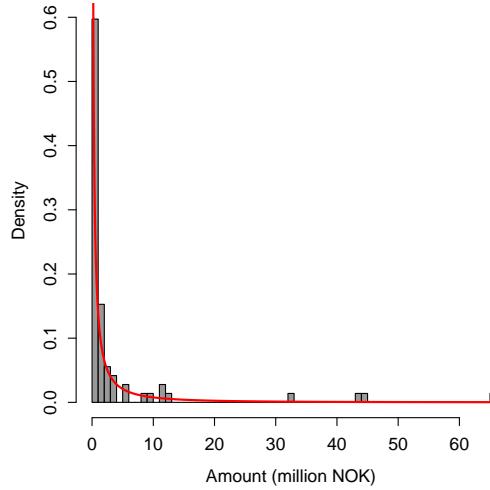
315 Figure 5 shows the histogram of observed annual damage costs for Bergen and the
 316 associated Burr distribution. The parameter estimates are $\hat{\alpha} = 7.84$ (3.6), $\hat{\gamma} = 0.40$ (0.04)
 317 and $\hat{\theta} = 0.007$ (0.01). *Grieg Foundation* [2009] discuss several different adaptation options
 318 for Bergen. In Figure 6 we consider the optimal timing of an adaptation option that in-
 319 cludes two inner barriers at Vågen and Damgårdssundet, that is, one on each side of cen-
 320 tral Bergen. The combined construction cost of the two barriers for a protection against 75
 321 cm sea level rise is estimated at 1.13 billion NOK (2015 level).

331 5.4 Identifying flood-prone areas

332 Table 5.4 contains the total projected storm surges for Esbjerg, corresponding to the
 333 20-year and 100-year historical surges, with and without sea level rise. We see that the
 334 historical maximum 433 cm is very likely to be exceeded by 2100.



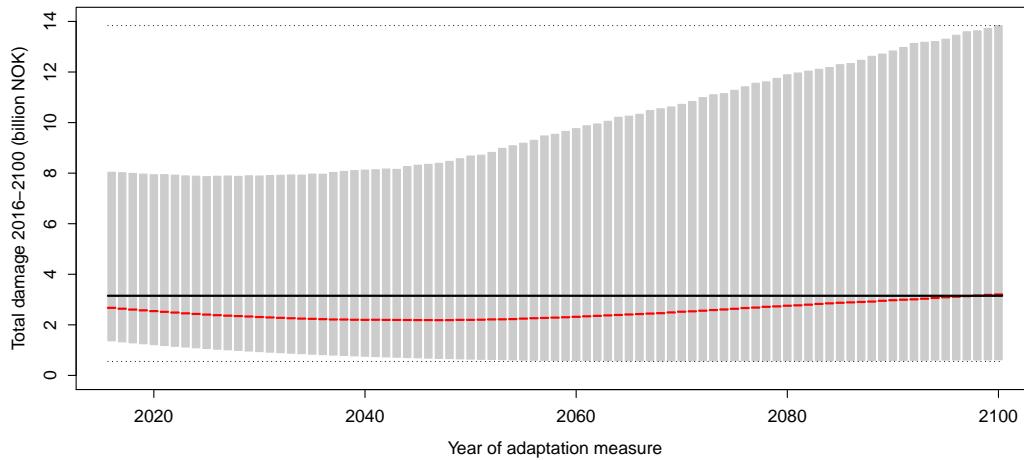
298 **Figure 4.** Simultaneous 90% confidence set (thick black lines) for Bergen (left) and Esbjerg(right) sea
299 level projections for the years 2000-2100 using RCP8.5. The sea level data are shown in blue and end in
300 2015. The thin red lines are the projections without uncertainty based on each of the climate models. The
301 dashed purple lines connect pointwise confidence intervals for each year.



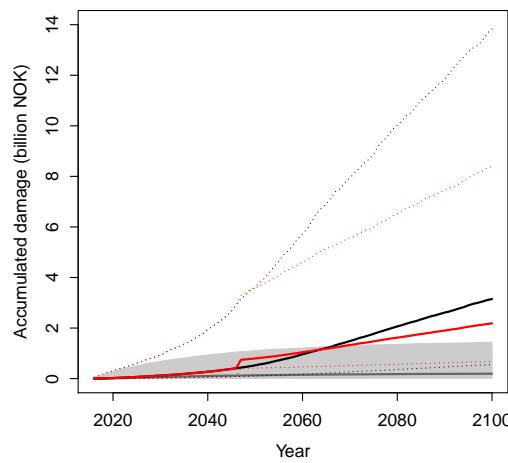
313 **Figure 5.** Estimated distribution of annual damage costs in Bergen for 2015 (red) based on observed annual
314 damage in Hordaland and Rogaland counties 1980-2015 (gray bars).

Storm surge	No sea level rise	RCP 8.5		
		5th percentile	median	95th percentile
RP 20	388 cm	431 cm	464 cm	500 cm
RP100	405 cm	448 cm	481 cm	517 cm

335 **Table 1.** Storm surge water levels [Sorensen *et al.*, 2013] with a return period of 20 years (RP20) and 100
336 years (RP100) with no sea level rise and with sea level rise corresponding to RCP 8.5 (5th percentile, median,
337 and 95th percentile).

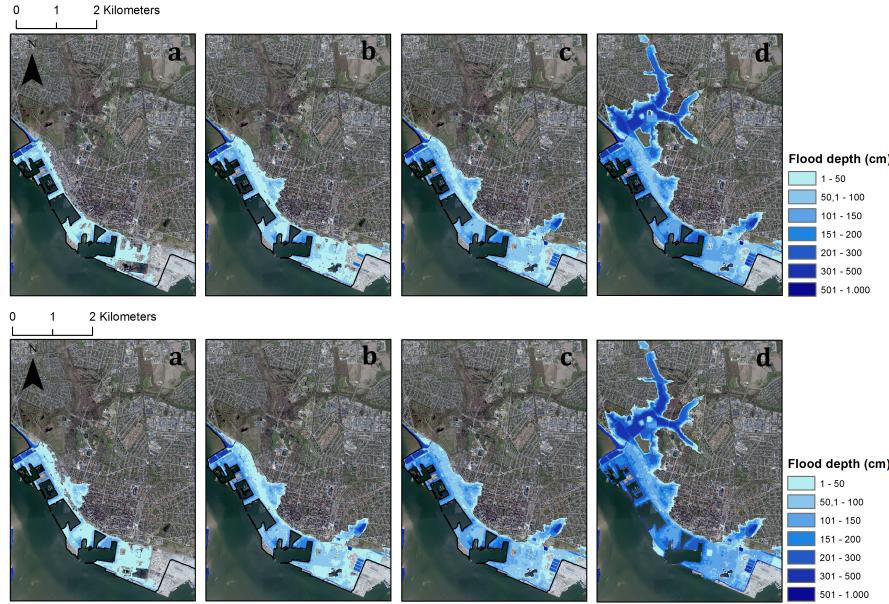


322 **Figure 6.** Projected total damage costs in Bergen for the time period 2016–2100 as a function of the timing
 323 of an adaptation measure consisting of the construction of two inner barriers. The median projection under
 324 each adaptation scenario is indicated in red with gray bars denoting the 90% projection intervals. The median
 325 projected total damage cost under no action is shown with a black line with the corresponding 90% projection
 326 interval indicated by dotted lines.



327 **Figure 7.** Median projected cumulative damage costs in Bergen under constant sea level (gray line), under
 328 sea level rise according to RCP 8.5 with no adaptation action (black line) and with the construction of two
 329 inner barriers in 2047 (red line). The shaded gray area denotes the 90% projection interval under constant sea
 330 level. Dotted lines indicate the 90% projection intervals with sea level rise according to RCP 8.5.

341 Figure 8 shows flooding maps corresponding to the entries in Table 5.4. These maps
 can be used to further develop adaptation policies for Esbjerg.



338 **Figure 8.** Flood extents and depths for the city of Esbjerg in year 2100 during storm surges with a 20 year
 339 return period (upper row) and a 100 year return period (lower row) with (a) no sea level rise and with sea level
 340 rise corresponding to RCP 85 (5th percentile (b), median (c) and 90th percentile (d)).

342

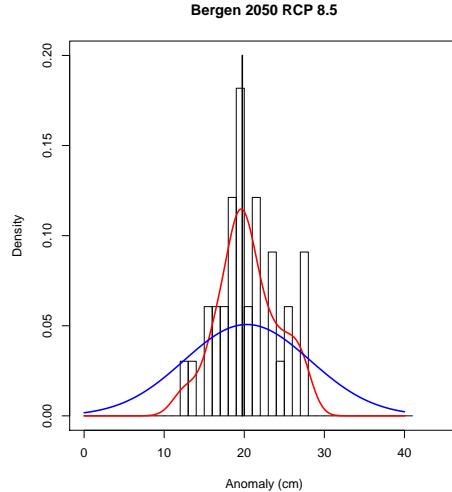
343 6 The value of including uncertainty

344 6.1 Sea level projections

345 In many cases sea level rise projections are given as a single number for each sce-
 346 nario, usually the mean or median of the ensemble of projections from different climate
 347 models (e.g. *Mote et al. [2008]*). Sometimes the spread of the ensemble is used to as-
 348 sess the uncertainty in the projections (e.g., the Norwegian Environmental Agency rec-
 349 commends using the upper ensemble value for RCP 8.5 as the basis for planning decision,
 350 pers. comm. from Even Nilsson, Norwegian Mapping Authority). In our analysis there are
 351 two more sources of uncertainty, namely the two regression models. Figure 9 shows the
 352 single number (vertical black line), the ensemble spread (histogram), the uncertainty in-
 353 cluding only the global model (red) and the full uncertainty (blue) for Bergen projections
 354 of sea level rise relative to 1999 under RCP 8.5. We see that the ensemble range is about
 355 16 cm, whereas the overall uncertainty range is about 40 cm.

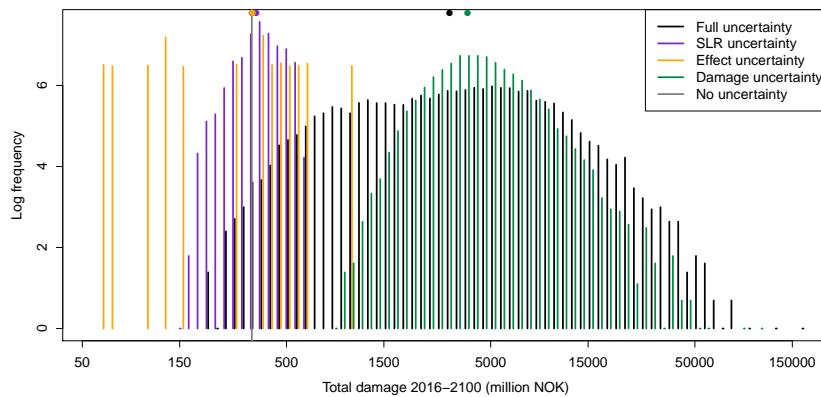
360 6.2 Damage costs

361 A simplistic analysis of projected damage costs at the year 2100, not taking into
 362 account the uncertainty, would use the median historical damage cost multiplied by the
 363 median damage effect factor at 2100 at the median sea level rise projected for 2100. This
 364 yields a damage cost of NOK 335 million (the grey vertical line in Figure 10). Similar
 365 results obtain when allowing sea level or effect factor to vary, holding the other quantities



356 **Figure 9.** 2050 Bergen sea level projections with uncertainty due to different sources for RCP 8.5. The
 357 black vertical line is the median projection (with no uncertainty), while the gray histogram corresponds
 358 to the spread of the climate models, the red curve adds the uncertainty due to the relation between global
 359 temperature and global sea level, and the blue line that due to downscaling global sea level to Bergen.

366 at the median (yellow and purple dots on top of Figure 10). However, allowing only the
 367 damage cost to vary yields a median cost of 3.84 billion (green dot on top of Figure 10).
 368 The appropriate uncertainty analysis for our model should draw each of sea level, effect
 369 factor and damage cost at random from their 2100 distributions. This corresponds to a
 370 median cost of 3.00 billion NOK, a factor of 9 higher than the simplistic value. Over 99%
 371 of the costs in our model are higher than the simplistic median.



372 **Figure 10.** Simulated distribution of total damage cost for 2100 without adaptation on log-log scale (black
 373 histogram). We also show the distributions of costs varying only one aspect of the uncertainty (sea level rise
 374 in blue, effect multiplier in yellow, and damage cost in green), holding the other two at their median values.
 375 The grey vertical line is the result of holding all three factors at their median value. The median of each
 376 distribution is shown as a dot on top of the figure.

377 **7 Conclusions and discussion**

378 Important to continue an iterative and interactive with end-users, as the decision
 379 framework provides actionable information to decision-makers, who will then make their
 380 own decisions. These decision are then incorporated into the current adaptation strategy.
 381 Continued loop to identify potential vulnerabilities of the approaches across a wide-range
 382 of possible futures.

383 Take home message

- 384
- It is possible to account for uncertainty even if you do a light touch decision-making
 385 approach – Explain what it implies if you don't account and if you do
 - Proof of concept – connect in a probabilistic way (normally done in a discrete way)
 386 – consider uncertainty in both ways
 - Improve the accessibility and demonstrating the general applicability of policy first
 387 approaches – towards more generic toolkit
- 388

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 394 (ClimateXL). The source code for the analysis is implemented in the statistical program-
 395 ming language R (*R Core Team [2016]*) and is available on GitHub at <http://github.com/eSACP/SeaLevelDecisions/Code>.
 396

397 **References**

- 398 Bamber, J., and W. Aspinall (2013), An expert judgement assessment of future sea level
 399 rise from the ice sheets, *Nature Climate Change*, 3, 424–427.
- 400 Bisaro, A., R. Swart, and J. Hinkel (2016), Frontiers of solution-oriented adaptation re-
 401 search, *Regional Environmental Change*, 16(1), 123–136.
- 402 Bolin, D., P. Guttorm, A. Januzzi, M. Novak, H. Podschwit, L. Richardson, C. Sowder,
 403 A. Särkkä, and A. Zimmerman (2014), Statistical prediction of global sea level from
 404 global temperature, *Statistica Sinica*, 25, 351–367.
- 405 Box, G., and G. Jenkins (1970), *Time series analysis: Forecasting and control*, Holden-
 406 Day.
- 407 Burr, I. W. (1942), Cumulative frequency functions, *The Annals of Mathematical Statistics*,
 408 13(2), 215–232.
- 409 Chambwera, M., G. Heal, C. Dubeux, S. Hallegatte, L. Leclerc, A. Markandya, B. Mc-
 410 Carl, R. Mechler, and J. Neumann (2014), *Climate Change 2014: Impacts, Adaptation,*
 411 *and Vulnerability. Part A: Global and Sectoral Aspects. Contribution of Working Group II*
 412 *to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, chap.
 413 Economics of adaptation, pp. 945–977, Cambridge University Press, NY, USA.
- 414 Church, J., and N. White (2011), Sea-level rise from the late 19th to the early 21st cen-
 415 tury, *Surveys in Geophysics*, 32, 585–602.
- 416 Fankhauser, S., and R. Soare (2013), An economic approach to adaptation: illustrations
 417 from europe, *Climatic Change*, 118(2), 367–379.
- 418 Fankhauser, S., J. Smith, and R. S. J. Tol (1999), Weathering climate change: some simple
 419 rules to guide adaptation decisions, *Ecological Economics*, 30, 67–78.
- 420 Grieg Foundation, v. (2009), Regional havstigning – prosjektrapport, *Tech. rep.*, Grieg
 421 Foundation, Visjon Vest and G0 Rieber Fondene, Bergen.
- 422 Guttorm, P., D. Bolin, A. Januzzi, D. Jones, M. Novak, H. Podschwit, L. Richardson,
 423 A. Särkkä, C. Sowder, and A. Zimmerman (2014), Assessing the uncertainty in pro-

- 424 jecting local mean sea level from global temperature, *Journal of Applied Meteorology*
 425 and *Climatology*, 53, 2163–2170.
- 426 Hallegatte, S., A. Shah, R. Lempert, C. Brown, and S. Gill (2012), *Investment Decision*
 427 *Making Under Deep Uncertainty Application to Climate Change*, The World Bank,
 428 Washington, D.C.
- 429 Hallegatte, S., C. Green, R. J. Nicholls, and J. Corfee-Morlot (2013), Future flood losses
 430 in major coastal cities, *Nature Climate Change*, 3(9), 802–806.
- 431 Hansen, J., R. Rued, M. Sato, M. Imhoff, W. Lawrence, D. Easterling, T. Peterson, and
 432 T. Karl (2001), A closer look at United States and global surface temperature change, *J.*
 433 *Geophys. Res.*, 106, 23,947–23,963.
- 434 Hinkel, J., and A. Bisaro (2016), Methodological choices in solution-oriented adaptation
 435 research: a diagnostic framework, *Regional Environmental Change*, 16(1), 7–20.
- 436 Jevrejeva, S., L. P. Jackson, R. E. M. Riva, A. Grinsted, and J. C. Moore (2016), Coastal
 437 sea level rise with warming above 2°c, *Proceedings of the National Academy of Sci-*
 438 *ences*, doi:10.1073/pnas.1605312113.
- 439 Klugman, S. A., H. H. Panjer, and G. E. Willmot (2012), *Loss models: from data to deci-*
 440 *sions*, vol. 715, 3rd ed., John Wiley & Sons, Hoboken, NJ.
- 441 Mosteller, F., and J. W. Tukey (1977), *Data Analysis and Regression*, Addison-Wesley,
 442 Reading, MA.
- 443 Mote, P., A. Petersen, S. Reeder, H. Shipman, and L. W. Binder (2008), Sea level rise in
 444 the coastal waters of Washington State, *Tech. rep.*, University of Washington Climate
 445 Impacts Group.
- 446 Norwegian Ministry of Finance (2012), Cost-benefit analysis, Official Norwegian Reports
 447 NOU 2012: 16.
- 448 R Core Team (2016), *R: A Language and Environment for Statistical Computing*, R Foun-
 449 dation for Statistical Computing, Vienna, Austria.
- 450 Rahmstorf, S. (2007), A semi-empirical approach to projecting future sea-level rise, *Sci-*
 451 *ence*, 315, 368–370.
- 452 Rahmstorf, S., M. Perrette, and M. Vermeer (2011), Testing the robustness of semi-
 453 empirical sea level projections, *Climate Dynamics*, 39, 861–875.
- 454 Simpson, M. J. R., K. Breili, and H. P. Kierulf (2014), Estimates of twenty-first century
 455 sea-level changes for norway, *Climate Dynamics*, pp. 1405–1424.
- 456 Sorensen, C., H. T. Madsen, and S. B. Knudsen (2013), Hojvandsstatistikker 2012, *Tech.*
 457 *rep.*, Kystdirektoratet, Lemvig.
- 458 Stocker, T., D. Qin, G.-K. Plattner, M. Tignor, S. K. Allen, J. Boschung, A. Nauels,
 459 Y. Xia, V. Bex, and P. Midgley (2013), *Climate Change 2013: The Physical Science*
 460 *Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergov-*
 461 *ernmental Pabel on Climate Chang*, Cambridge University Press.
- 462 Taylor, K., R. Stouffer, and G. Meehl (2012), An overview of CMIP5 and the experiment
 463 design, *Bull. Amer. Meteor. Soc.*, 93, 485–498.
- 464 van Vuuren, D. P., J. Edmonds, M. Kainuma, K. Riahi, A. Thomson, K. Hibbard, G. C.
 465 Hurtt, T. Kram, V. Krey, J.-F. Lamarque, T. Masui, M. Meinshausen, N. Nakicenovic,
 466 S. J. Smith, and S. K. Rose (2011), The representative concentration pathways: an
 467 overview, *Clim. Change*, 109, 5–31.
- 468 Wada, Y., L. P. H. van Beek, F. C. Sperna Weiland, B. F. Chao, Y.-H. Wu, and M. F. P.
 469 Bierkens (2012), Past and future contribution of global groundwater depletion to sea-
 470 level rise, *Geophysical Research Letters*, 39, doi:10.1029/2012GL051230.
- 471 Walker, W. E., M. Haasnoot, and J. H. Kwakkel (2013), Adapt or perish: a review of
 472 planning approaches for adaptation under deep uncertainty, *Sustainability*, 5(3), 955–
 473 979.
- 474 Watkiss, P., A. Hunt, W. Blyth, and J. Dyszynski (2015), The use of new economic de-
 475 cision support tools for adaptation assessment: A review of methods and applications,
 476 towards guidance on applicability, *Climatic Change*, 132(3), 401–416.

- 477 Wilby, R. L., and S. Dessai (2010), Robust adaptation to climate change, *Weather*, 65(7),
478 180–185.