

**I don't know, are you sure we want to do this?
Sea level adaptation decisions under uncertainty**

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Key Points:

- Decisions on adaptation measures need to take careful account of all uncertainties
 - Modeling local sea level rise is essential
 - Decisions under uncertainty do not have to be based on precise costs

12 **Abstract**

13 Sea level rise has serious consequences for harbor infrastructure, storm drains and sewer
 14 systems, and many other issues. Adapting to sea level rise requires comparing different
 15 possible adaptation strategies, comparing the cost of different actions (including no ac-
 16 tion), and assessing at what point in time the chosen strategy should be implemented. All
 17 these decisions must be made under considerable uncertainty—in the amount of sea level
 18 rise, in the cost of adaptation actions, and in the cost of no action. We develop two illus-
 19 trative examples: for Bergen on Norway’s west coast and for Esbjerg on the west coast of
 20 Denmark. Different components of uncertainty are visualized. We show that failing to take
 21 uncertainty into account can result in the median projected damage costs being an order of
 22 magnitude smaller.

23 **1 Introduction**

24 The potential impact of climate change on local sea level, yielding effects such as
 25 frequent flooding, inundation and backflow of storm drainage and sewer systems, destruc-
 26 tive erosion and contamination of wetlands and other habitats, requires city planners to
 27 make decisions in the presence of substantial uncertainty.

28 As adaptation decision-making is an ongoing process of weighing and choosing
 29 which measures should be taken at which moment in time [Hallegatte *et al.*, 2012], adap-
 30 tive planning methods need to support decisions in the short term, while considering long-
 31 term developments. Challenges of adaptation decision-making under uncertainty relate
 32 to the incorporation of spatial, inter-temporal and flexibility aspects of adaptation priori-
 33 ties [Fankhauser and Soare, 2013], and the linkage with specific characteristics of sectors
 34 and contexts [Bisaro *et al.*, 2016; Hinkel and Bisaro, 2016]. Several economic decision
 35 support tools and methods exist for adaptation assessment under uncertainty [e.g. Chamb-
 36 wera *et al.*, 2014; Wilby and Dessai, 2010; Walker *et al.*, 2013]. However, Watkiss *et al.*
 37 [2015] conclude that these tools are very resource intensive and complex in the context of
 38 long-term adaptation investment decisions and call for the development of “light touch”
 39 approaches to better support local adaptation making.

40 In this paper we employ light touch decision tools to demonstrate the importance of
 41 combining projections of sea level rise and flood damages alongside a detailed quantifi-
 42 cation of both hydrologic and economic uncertainties in the context of real-life decision-
 43 problems experienced by stakeholders and authorities in two northern European cities,
 44 Bergen in Norway and Esbjerg in Denmark, see Figure 1. Based on communications with
 45 local end-users we highlight the value of taking into account uncertainty through two sim-
 46 plified and complementary case studies, where in the first one planners want to know how
 47 early they should implement costly adaptation measures, whereas in the second case the
 48 aim is to highlight the risk of flooding in coastal areas, e.g. in order to prioritize future
 49 adaptation actions and investments. In both cases we show that embracing the uncertain-
 50 ties derived from economic and hydrologic models is absolutely crucial in order to answer
 51 the question of “are we sure we want to do this?”

53 The Norwegian city of Bergen is the capital of Hordaland County. The city center is
 54 located on Byfjorden, and is surrounded by mountains. It has the largest port in Norway,
 55 both in terms of freight and passengers. The historic harbor area, Bryggen, is the only
 56 Hanseatic trade center remaining in its original style, and has been declared a UNESCO
 57 World Heritage site¹. Bryggen is regularly flooded at extreme tides, and it is feared that
 58 as sea levels rise, floods will become a major problem in other parts of Bergen as well
 59 [Grieg Foundation, 2009].

¹ See <http://whc.unesco.org/en/list/59>.



52 **Figure 1.** Terrain maps of central Bergen, Norway (left) and Esbjerg, Denmark (right).

60 The municipality of Bergen has, in cooperation with private actors, analyzed sev-
 61 eral possible adaptation measures against sea level rise. The measures range from an outer
 62 barrier that would protect the entire metropolitan area to various protection measures of
 63 limited areas in the inner harbor [Grieg Foundation, 2009]. While the viability of the con-
 64 structions and the associated construction costs have been carefully analyzed, the optimal
 65 timing of potential adaptation measures and the effects of the associated uncertainties have
 66 yet to be investigated. We perform such an analysis where we consider uncertainty in pro-
 67 jected sea level rise, damage costs, and the effect of sea level rise on changes in damage
 68 costs.

69 Esbjerg, on the southwest coast of Jutland, is the fifth-largest city in Denmark and
 70 the largest urban area in the region. The city hosts one of the largest harbours in Den-
 71 mark, which serves as a focal point for offshore activities in the North Sea, including the
 72 continued development of offshore wind power and extensive activities related to the ex-
 73 traction of oil and gas. As a result critical infrastructures and commercial buildings fig-
 74 ure prominently in the coastal zone. Esbjerg is frequently subject to substantial storms
 75 and storm surges, causing severe flooding of the harbour and the city. The highest since
 76 records began in 1874 was recorded in 1981, where the harbour was completely flooded
 77 and the water level reached 433 cm above the norm, causing massive economic losses.
 78 More generally, storm surges causing water levels in Esbjerg to rise to between 2 and 3
 79 metres have quadrupled over the last four decades according to local records, whereas half
 80 of the most severe events have taken place since 1975.

81 As in the case of Bergen, sea level rise caused by climate change is expected to
 82 compound these risks, alongside parallel threats caused by increased risks of pluvial flood-
 83 ing and rising ground water levels in Esbjerg. The municipality recently adopted its cli-
 84 mate adaptation plan, which in its first phase is aimed at identifying present and future
 85 flood-prone areas, e.g. to avoid urban development into such areas, to limit damages to
 86 buildings of high societal or cultural value, and to pave the way for implementing cost-
 87 effective adaptation measures in the second phase of the plan.

88 The remainder of the paper is organized as follows. Section 2 describes our ap-
 89 proach to projecting sea level. In sections 3 and 4 we describe the type of decision prob-
 90 lems that we are going to attack. We apply these tools to sea level projections for Bergen,
 91 Norway, and for Esbjerg, Denmark in section 5. In section 6 we demonstrate the conse-

92 consequences of ignoring the uncertainty in the projections, and the paper is closed with a sum-
93 mary and discussion in section 7.

94 2 Sea level projections

95 We project local sea level changes by modeling two processes, the relationship be-
96 tween global temperature and global sea level, and the relationship between global sea
97 level and local sea level.

98 2.1 Global sea level

99 Most climate models do not explicitly provide sea level as an output of the calcu-
100 lations. Rather, the IPCC AR5 report [Stocker *et al.*, 2013, ch. 13] combines the heat ex-
101 pansion of the ocean with temperature forced models for glacial melt, Greenland ice melt,
102 and Antarctic ice melt and with land rise due to rebound from the last ice age and other
103 tectonic effects. Judging from the supplementary material to Stocker *et al.* [2013, ch. 13],
104 the uncertainty assessment is only based on the spread of the ensemble of temperature
105 projections, not on the additional uncertainty in the ice models used.

106 We will instead use the empirical approach of Rahmstorf and collaborators [Rahm-
107 storf, 2007; Rahmstorf *et al.*, 2011], employing the statistical modeling of Bolin *et al.*
108 [2014] to relate global annual mean temperature anomalies to global mean sea level anom-
109 alies. We then apply the estimated historical relationship to projected temperatures from the
110 CMIP5 experiment [Taylor *et al.*, 2012] to obtain projected global annual mean sea level,
111 taking into account the uncertainty in the statistical model as well as the spread of the
112 temperature projection ensemble (see subsection 2.3). For the i'th temperature projection
113 T_t^i we estimate the corresponding global mean sea level as

$$114 H_t^{gl,i} = \int_{t_0}^t \hat{a}(T_u^i - \hat{T}_0) du + \varsigma_t,$$

115 where \hat{a} and \hat{T}_0 are regression parameters of observed global sea level rise on observed
116 global temperature and ς_t the integrated time series regression error.

117 2.2 Local sea level

118 In order to get from global sea level projections to local ones, it is important to note
119 that sea level rise is not uniform over the globe. Glacial and land ice melting affect the
120 local sea level differently depending on where the melted ice is located. Another major
121 effect in Fennoscandia is the land rise due to isostatic rebound from the glaciers of the
122 last ice age. Again, we will use historical data to relate global sea level to isostatically
123 corrected local sea level using a time series regression model. The local sea level projec-
124 tions are then obtained by first relating projected temperature to global sea level, and then
125 relating the global sea level to the local one. Each climate model temperature projection
126 yields a different local sea level projection. The local sea level projection based on the i'th
127 climate model for years beyond 2000 is estimated as

$$128 H_t^{loc,i} + \gamma(t - 2000) = \hat{b}H_t^{gl,i} + \varepsilon_t,$$

129 where γ is the annual land rise rate, t denotes year, \hat{b} is the regression coefficient relating
130 global to local sea level and the ε_t are Gaussian errors..

131 2.3 Uncertainty assessment

132 Following the approach of Guttorm *et al.* [2014] we assess the uncertainty in the lo-
133 cal sea level projections taking into account the variability between the climate projections

134 used, the uncertainties in the regressions of global mean temperature on global mean sea
 135 level and of global on local sea level. We express the sea level projection uncertainty in
 136 terms of a confidence band that is simultaneously of the intended level for all projection
 137 years. This allows us, for example, to get a confidence band for the years when a given
 138 sea level rise is obtained.

139 2.4 Limitations of the sea level projections

140 The main assumption is using historical relationships in statistical projections of
 141 the type used in this paper is that there is no major change in how temperature relates
 142 to sea level, globally and locally. Among the factors that may invalidate this approach
 143 are changes in water storage on land (in essence removing water from the oceans), ex-
 144 cessive siphoning of groundwater (resulting in land subsidence), changes in the rates of
 145 glacial and land ice melt, and changes in Earth's gravitational field due to transfer of mass
 146 from land ice to ocean water. For example, the rate of ice melt on Greenland may sud-
 147 denly increase substantially due to intense warming of both air and sea water [Bamber
 148 and Aspinall, 2013]. A recent paper [Jevrejeva *et al.*, 2016] indicates that the upper tails
 149 of sea level rise may be substantially higher when taking into account expert assessment
 150 of land ice melting. Our current climate models are not able to resolve the ice processes
 151 sufficiently to include such so called tipping points into the projections. Also, the IPCC
 152 scenarios [*van Vuuren et al.*, 2011] do not include changes in water usage (cf. Wada *et al.*
 153 [2012]).

154 3 Timing of adaptation measures

155 There is an increasing need for more detailed economic analysis, including simple
 156 methods and tools for assessment of options, especially since Downing (2012) recog-
 157 nizes, adaptation is moving from theory to practice, and practitioners try to deal with how
 158 to begin adapting. This leads to an increasing need and interest in the appraisal of options.

159 Several economic decision support tools and methods exist for adaptation assess-
 160 ment under uncertainty. Robust decision-making approaches are able to better incorporate
 161 uncertainty and a broad range of climate scenarios to capture as much of the uncertainty
 162 on future climates as possible (Dittrich *et al.* 2016). These approaches can be classified
 163 according to a science-first or policy-first approach. The former has a "predict-then-act"
 164 foundation, which starts with climate projections and impact assessments, not linked to
 165 any specific adaptation choices (Jones *et al.*, 2014). The latter starts out with the formu-
 166 lated adaptation plans and not impacts, and their functioning is tested against different
 167 future projections (Dittrich *et al.* 2016).

168 We take a policy-first approach, in which we test a current adaptation plan against
 169 different sea level and damage projections and the inclusion of different sources of un-
 170 certainty. We focus on what this implies for the timing of adaptation measures and the
 171 implications of including uncertainty. In particular, we employ a probabilistic extension of
 172 the framework described by [Fankhauser *et al.*, 1999] in which we obtain a probabilistic
 173 distribution for the net present value damage in a given year for various adaptation op-
 174 tions. The probabilistic distribution is constructed by considering uncertainty in the local
 175 sea level projections, in the annual damage costs, and in the effect of changes in sea level
 176 on the annual damage costs.

177 3.1 Annual damage costs

178 We model the distribution of annual damage, F_{d,t_0} , for the year $t_0 = 2015$ by the
 179 three parameter Burr distribution [Burr, 1942] with density

$$180 f_{d,t_0}(x) = \frac{\alpha\gamma(x/\theta)^\gamma}{x[1 + (x/\theta)^\gamma]^{\alpha+1}} \quad (1)$$

for $x > 0$, where α and γ are shape parameters with $\alpha, \gamma > 0$, and $\theta > 0$ is a scale parameter. The Burr distribution has a heavy upper tail and is commonly used to model damage loss, see e.g. *Klugman et al.* [2012]. The parameters of the distribution are estimated using historical data for annual storm surge damage. Data prior to 2015 are adjusted to the 2015 level using the consumer price index. After adjustment, we assume stationarity over the period and independence between years.

Under a constant sea level, we can obtain a sample trajectory $\{d_{t_1}, d_{t_2}, \dots, d_{t_{85}}\}$ of future annual damages for $t_1 = 2016, \dots, t_{85} = 2100$ by drawing 85 i.i.d. values from the estimated distribution \hat{F}_{d,t_0} . By repeating this process J times, we obtain an empirical damage distribution for each future year t_i given by

$$\hat{F}_{d,t_i}(x) = \frac{1}{J} \sum_{j=1}^J \mathbb{1} \left\{ \frac{d_{t_i}^{(j)}}{\prod_{l \leq i} (1 + r_{t_l})} \leq x \right\}$$

for $i = 1, \dots, 85$, where r_{t_i} is the discount rate for year t_i . Alternatively, we obtain an empirical distribution of the total damage over the period 2016 – 2100 by considering

$$d_{\text{total}}^{(j)} = \sum_{i=1}^{85} \frac{d_{t_i}^{(j)}}{\prod_{l \leq i} (1 + r_{t_l})}$$

and similarly for the cumulative damage.

3.2 Effect of changes in sea level

We assume that changes in sea level have a multiplicative effect on the annual damage cost. That is, for a sea level anomaly s_{t_i} in year $t_i > 2015$ compared to the 2015 level, the annual damage cost becomes

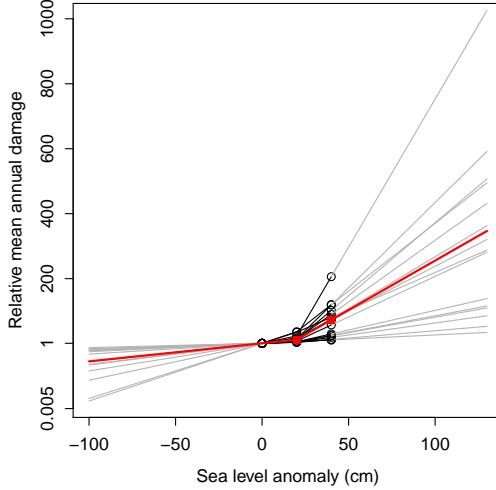
$$g(s_{t_i} | \boldsymbol{\beta}) d_{t_i},$$

where $g(\cdot | \boldsymbol{\beta})$ is a monotonic positive function with parameter vector $\boldsymbol{\beta}$ such that $g(s | \boldsymbol{\beta}) > 1$ for $s > 0$ and $g(s | \boldsymbol{\beta}) < 1$ for $s < 0$. *Halleygate et al.* [2013] estimate a similar effect function valid in 2050 for $s \in \{0, 20, 40\}$ for 136 coastal cities. Here, we use their results for 15 European cities: Amsterdam, Athens, Barcelona, Dublin, Glasgow, Hamburg, Helsinki, Copenhagen, Lisbon, London, Marseilles, Naples, Porto, Rotterdam and Stockholm. To obtain a city-specific effect function for a large range of sea level anomalies we employ a linear extrapolation as shown in Figure 2. We then obtain a sample of effect functions $\{g(\cdot | \boldsymbol{\beta}^{(j)})\}_{j=1}^J$ by sampling with replacement from this ensemble of trajectories with all 15 ensemble members considered equally probable.

Let further $\{s_{t_1}^{(j)}, \dots, s_{t_{85}}^{(j)}\}_{j=1}^J$ denote a sample of projections for annual sea level anomalies compared to the 2015 value. An empirical damage distribution for the future year t_i that accounts for uncertainty in damage, sea level rise and its effect on the damage is then given by

$$\hat{F}_{d,t_i}^s(x) = \frac{1}{J} \sum_{j=1}^J \mathbb{1} \left\{ \frac{g(s_{t_i}^{(j)} | \boldsymbol{\beta}^{(j)}) d_{t_i}^{(j)}}{\prod_{l \leq i} (1 + r_{t_l})} \leq x \right\}. \quad (2)$$

The distribution in (2) describes the projected damage distribution with no adaptation measures. In addition, we can incorporate an adaptation measure of cost C that protects against K cm of increased sea level from year t_k onward. This results in a damage



210 **Figure 2.** Relative change in mean annual damage as a function of sea level rise for 15 European cities as
211 estimated by *Hallegatte et al.* [2013] (black circles) with linearly extrapolated values indicated by gray lines.
212 The median change and the corresponding extrapolation are indicated in red.

221 distribution given by

$$222 \quad \hat{F}_{d,t_i}^{s,a_k}(x) = \begin{cases} \frac{1}{J} \sum_{j=1}^J \mathbb{1} \left\{ \frac{g(s_{t_i}^{(j)} | \beta^{(j)}) d_{t_i}^{(j)}}{\prod_{l \leq i} (1+r_{t_l})} \leq x \right\}, & t_i < t_k \\ \frac{1}{J} \sum_{j=1}^J \mathbb{1} \left\{ \frac{g(s_{t_i}^{(j)} - K | \beta^{(j)}) d_{t_i}^{(j)} + C}{\prod_{l \leq i} (1+r_{t_l})} \leq x \right\}, & t_i = t_k \\ \frac{1}{J} \sum_{j=1}^J \mathbb{1} \left\{ \frac{g(s_{t_i}^{(j)} - K | \beta^{(j)}) d_{t_i}^{(j)}}{\prod_{l \leq i} (1+r_{t_l})} \leq x \right\}, & t_i > t_k. \end{cases}$$

223 **3.3 Limitations of the decision framework**

224 The main limitation of this light touch decision framework is that we have significantly simplified the assessment of the effect of sea level rise on the damage costs. In
225 particular, the linear extrapolation of the results reported in *Hallegatte et al.* [2013] might
226 provide a conservative estimate of the effect of extreme sea level rise. However, with only
227 two data points, extrapolation approaches such as a power law or exponential growth seem
228 difficult to justify.

229 Alternatively, a modeling framework similar to that of *Hallegatte et al.* [2013] could
230 be applied directly to a larger range of potential changes in sea level. The elements of
231 such a framework might include an appropriate social discount rate, valuing environmental
232 goods in monetary terms, incorporate socio-economic assumptions and long-term policy
233 goals of decision makers, as well as that climate change is often not the only driver that
234 decision makers should consider, therefore costs and benefits should be studied in a wider
235 context (Dittrich et al., 2016).

236 Our framework simplifies the cost and effect of an adaptation option during con-
237 struction in that we assume no effect until the construction is finished with all the con-
238 struction cost falling in the last year of the construction. Especially for larger construc-
239 tions, these assumptions might need to be modified. Additionally, we have not specifically
240 accounted for potential changes in storm surge patterns.

242 4 Danish decision framework

243 In 2014 the municipality of Esbjerg adopted its climate adaptation plan ([ref](#)), which
 244 will cover the period from 2014-2026, and which aims to reduce the risk of flooding
 245 caused by storm surges, heavy rainfall and rising ground water levels, respectively. In the
 246 following (and based on interviews with end-users from the municipality) we will focus
 247 on the harbour of Esbjerg and the nearby coastal areas, and we will only consider the hazard
 248 of coastal flooding, though evidently in some areas the flood risk is compounded.

249 The initial scoping of the climate adaptation plan in Esbjerg includes a preliminary
 250 value and risk mapping, considering critical infrastructure and buildings of high cultural
 251 and societal value as identified by the municipality, while informed by spatial floods maps
 252 for different scenarios corresponding to each of the three different kinds of hydrological
 253 events (sea level rise/storm surges, pluvial flooding and rising ground water levels). In
 254 terms of coastal floods the mapping considers only one type of storm surge, corresponding
 255 to a 20-year return event (based on historical storm surge statistics), and increased sea
 256 level due to climate change was not accounted for. We will extend these flood maps to
 257 also deal with 100-year return events and quantiles of the projected sea level rise.

258 The risk for any given map area as the probability of, e.g., a certain flood depth,
 259 is derived from a hydrological flood model (not including urban drainage system) times
 260 a valuation of the consequences, which—similar to the Bergen case—is essentially a loss
 261 function associating the flood depth with a measure of cost.

262 5 Case studies

263 5.1 Data

264 The historical global mean temperature series is obtained from *Hansen et al.* [2001].
 265 Climate projections of global mean temperature are from the fifth climate model intercom-
 266 parison project, CMIP5 [*Taylor et al.*, 2012]. The global mean sea level series is obtained
 267 from *Church and White* [2011]. We use local tide gauge data from the Permanent Service
 268 for Mean Sea Level, UK, which is the worldwide repository for national sea level data.
 269 Glacial isostatic adjustment for Bergen is obtained from *Simpson et al.* [2014], and for Es-
 270 bjergr in personal communication from Peter Thejll at the Danish Meteorological Institute.

271 The Bergen monthly series is missing data for 62 months, including all of the years
 272 1942–43. To deal with occasional short stretches of missing data (at most one or two
 273 months) we use median polish replacement [*Mosteller and Tukey*, 1977] and then com-
 274 pute annual averages. For the years 1942-43, we use the average difference between
 275 Bergen and the average of all other Norwegian stations in 1940 and 1943 to estimate val-
 276 ues for 1941 and 1942, using the average of all other Norwegian stations corrected by the
 277 average difference.

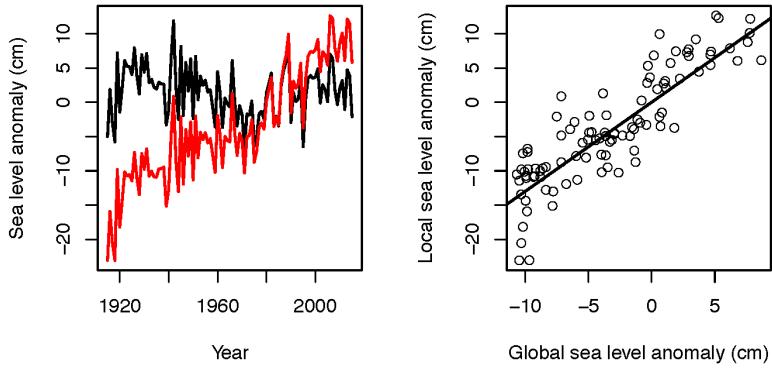
278 The Esbjerg monthly series is missing data for 19 months. Here, too, we use median
 279 polish to fill in missing data and then compute annual averages.

280 Annual damage costs for the Bergen case study are obtained from the Norwegian
 281 Natural Perils Pool (NPP; data are available at <https://www.finansnorge.no/statistikk/skadeforsikring/Naturskadestatistikk-NASK/>). The NPP data are available for
 282 the period 1980-2015 and are aggregated to a county level. For improved parameter es-
 283 timation, we include the data from Rogaland county which is the county directly south
 284 of Hordaland and with similar characteristics. We use a discount rate of 4% for the first
 285 40 years of the analysis, a rate of 3% for 40 to 75 years into the future and a rate of 2%
 286 beyond 75 years (cf. Section 5.8 of *Norwegian Ministry of Finance* [2012]).

288 Storm surge data for Esbjerg are obtained from the Danish Coastal Authority [*Sorensen*
 289 *et al.*, 2013].

290 **5.2 Sea level rise in Bergen and Esbjerg**

291 Figure 3 shows uncorrected and corrected Bergen sea level data, and the relationship
 292 between the corrected Bergen data and the global sea level data. The glacial iso-
 293 static adjustment is 0.26 (standard error 0.07) cm/yr. The time series regression uses an
 294 ARMA(1,1)-model [Box and Jenkins, 1970], with AR parameter 0.82 (0.13), and MA pa-
 295 rameter -0.61 (0.17). The regression slope is 1.30 (0.12).



296 **Figure 3.** The left figure shows raw (black) and gia-corrected (red) sea level data from Bergen. The right
 297 figure relates the gia-corrected Bergen sea level to the global sea level series of *Church and White* [2011]. The
 298 straight line is the time series regression line.

303 For the relationship between global annual mean temperature and global annual
 304 mean sea level rise we use the results from *Bolin et al.* [2014]. The left panel of figure 4
 305 shows the simultaneous 90 % confidence region for Bergen sea level rise relative to 1999
 306 under scenario RCP 8.5, which is the scenario Norwegian authorities recommend for plan-
 307 ning purposes.

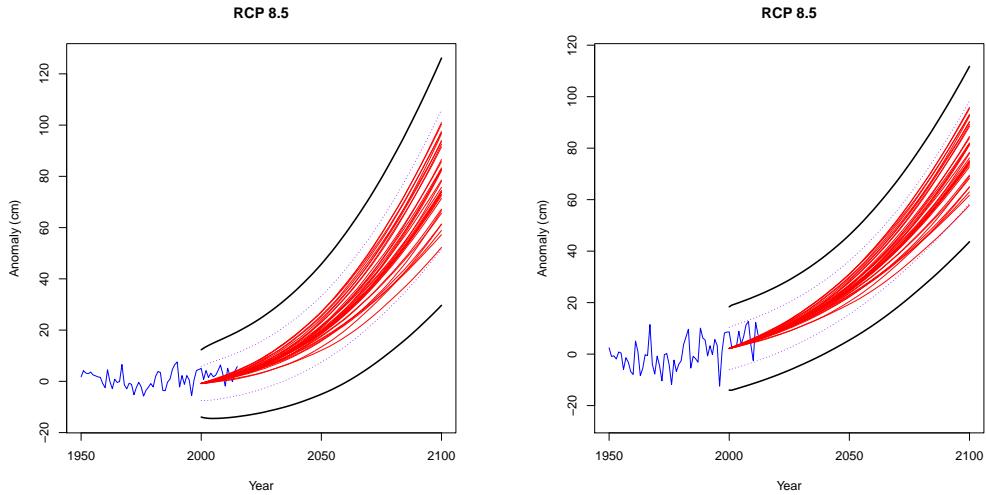
308 For Esbjerg, the glacial isostatic adjustment is 0.06 (0.03) cm/yr. The time series re-
 309 gression model relating glacial isostatic adjustment to global sea level is an MA(1) model with pa-
 310 rameter 0.17 (0.09). The regression slope is 1.02 (0.06). The right panel of figure 4 shows
 311 the simultaneous 90% confidence region for sea level rise relative to 1999 under scenario
 312 RCP 8.5.

313 **5.3 Timing of adaptation measures in Bergen**

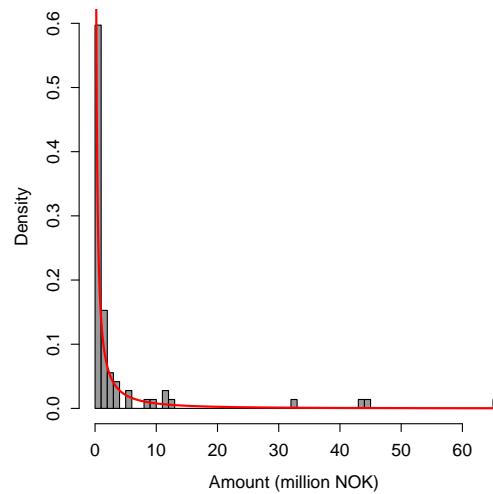
316 Figure 5 shows the histogram of observed annual damage costs for Bergen and the
 317 associated Burr distribution. The parameter estimates are $\hat{\alpha} = 7.84$ (3.6), $\hat{\gamma} = 0.40$ (0.04)
 318 and $\hat{\theta} = 0.007$ (0.01). *Grieg Foundation* [2009] discuss several different adaptation options
 319 for Bergen. In Figure 6 we consider the optimal timing of an adaptation option that in-
 320 cludes two inner barriers at Vågen and Damgårdssundet, that is, one on each side of cen-
 321 tral Bergen. The combined construction cost of the two barriers for a protection against 75
 322 cm sea level rise is estimated at 1.13 billion NOK (2015 level)².

323 Applying the methodology from section 3 we find that the optimal time of building
 324 the barriers is in 2046 (Figure 6), and that by the year 2100 this decision will on average

² 100 NOK is about 11 EUR.

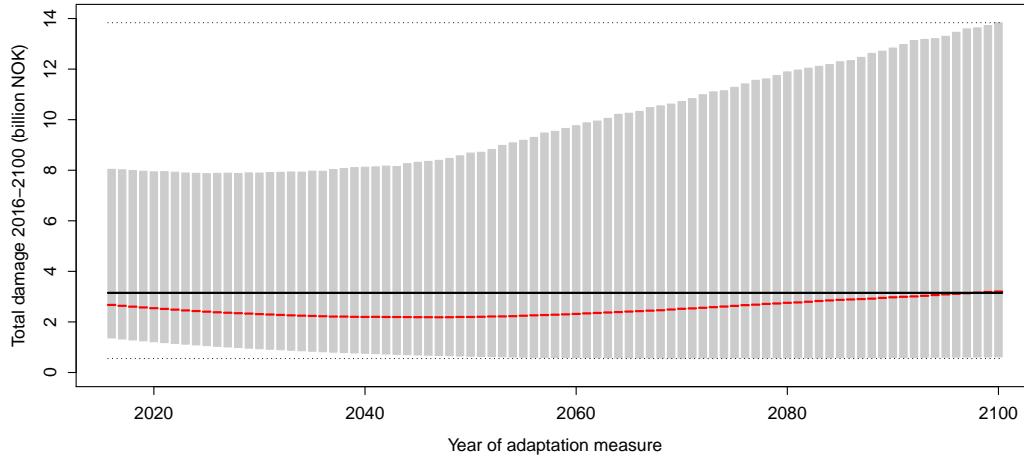


299 **Figure 4.** Simultaneous 90% confidence set (thick black lines) for Bergen (left) and Esbjerg(right)
 300 sea level projections for the years 2000-2100 using RCP8.5. The sea level data are shown in blue and end in
 301 2015. The thin red lines are the projections without uncertainty based on each of the climate models. The
 302 dashed purple lines connect pointwise confidence intervals for each year.

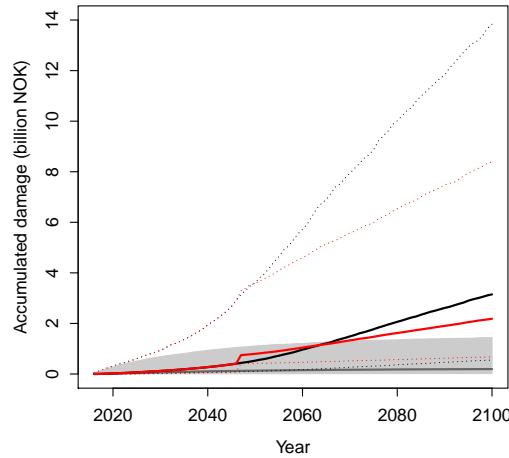


314 **Figure 5.** Estimated distribution of annual damage costs in Bergen for 2015 (red) based on observed annual
 315 damage in Hordaland and Rogaland counties 1980-2015 (gray bars).

325 save about 1/3 of the median damage costs without adaptation (Figure 7). At the high end
 326 (or the 95th percentile of the damage distribution) adaptation saves even more, about 43%.



327 **Figure 6.** Projected total damage costs in Bergen for the time period 2016-2100 as a function of the timing
 328 of an adaptation measure consisting of the construction of two inner barriers. The median projection under
 329 each adaptation scenario is indicated in red with gray bars denoting the 90% projection intervals. The median
 330 projected total damage cost under no action is shown with a black line with the corresponding 90% projection
 331 interval indicated by dotted lines.



332 **Figure 7.** Median projected cumulative damage costs in Bergen under constant sea level (gray line), under
 333 sea level rise according to RCP 8.5 with no adaptation action (black line) and with the construction of two
 334 inner barriers in 2047 (red line). The shaded gray area denotes the 90% projection interval under constant sea
 335 level. Dotted lines indicate the 90% projection intervals with sea level rise according to RCP 8.5.

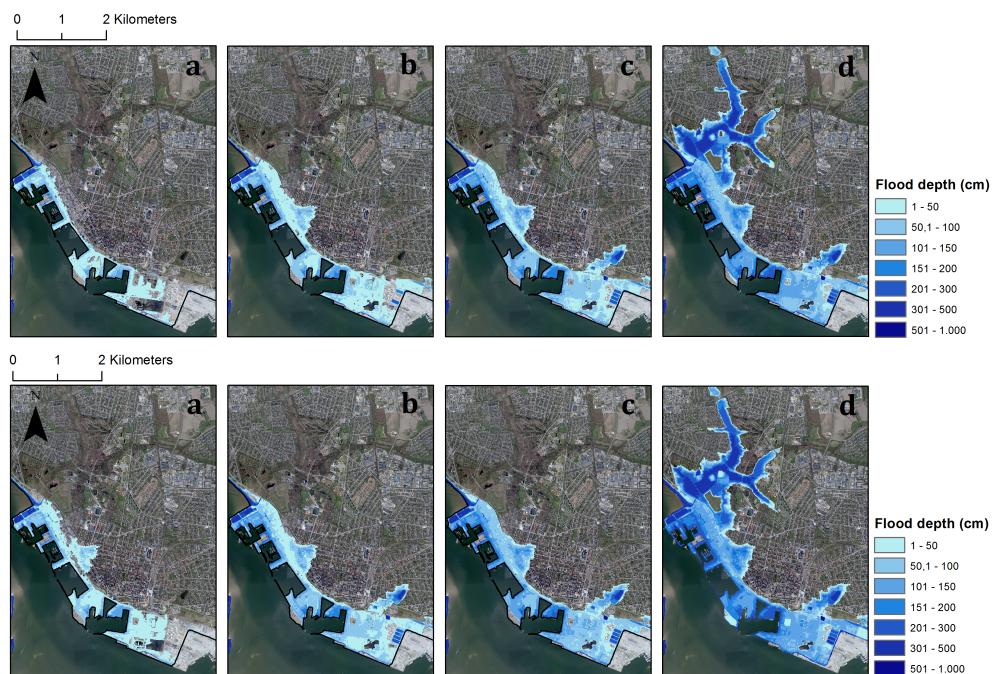
Storm surge	No sea level rise	5th percentile	RCP 8.5 median	95th percentile
RP 20	388 cm	431 cm	464 cm	500 cm
RP100	405 cm	448 cm	481 cm	517 cm

340 **Table 1.** Storm surge water levels [Sorensen *et al.*, 2013] with a return period of 20 years (RP20) and 100
 341 years (RP100) with no sea level rise and with sea level rise corresponding to RCP 8.5 (5th percentile, median,
 342 and 95th percentile).

336 **5.4 Identifying flood-prone areas**

337 Table 1 contains the total projected storm surges for Esbjerg, corresponding to the
 338 20-year and 100-year historical surges, with and without sea level rise. We see that using
 339 our projections the historical maximum 433 cm is almost certain to be exceeded by 2100.

346 Figure 8 shows flooding maps corresponding to the entries in Table 1. These maps
 can be used to further develop adaptation policies for Esbjerg.

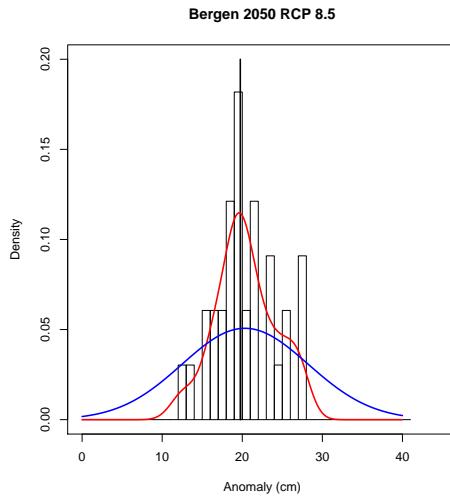


343 **Figure 8.** Flood extents and depths for the city of Esbjerg in year 2100 during storm surges with a 20 year
 344 return period (upper row) and a 100 year return period (lower row) with (a) no sea level rise and with sea level
 345 rise corresponding to RCP 85 (5th percentile (b), median (c) and 95th percentile (d)).

348 6 The value of including uncertainty

349 6.1 Sea level projections

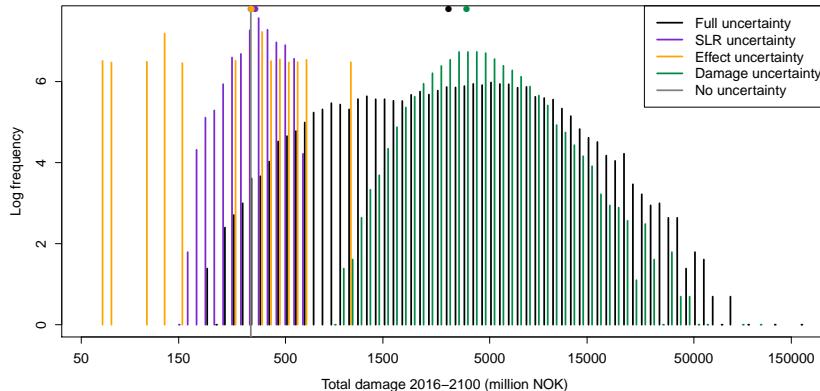
350 In many cases sea level rise projections are given as a single number for each sce-
 351 nario, usually the mean or median of the ensemble of projections from different climate
 352 models (e.g. *Mote et al.* [2008]). Sometimes the spread of the ensemble is used to as-
 353 sess the uncertainty in the projections (e.g., the Norwegian Environmental Agency rec-
 354 ommends using the upper ensemble value for RCP 8.5 as the basis for planning decision,
 355 pers. comm. from Even Nilsson, Norwegian Mapping Authority). In our analysis there are
 356 two more sources of uncertainty, namely the two regression models. Figure 9 shows the
 357 single number (vertical black line), the ensemble spread (histogram), the uncertainty in-
 358 cluding only the global model (red) and the full uncertainty (blue) for Bergen projections
 359 of sea level rise relative to 1999 under RCP 8.5. We see that the ensemble range is about
 360 16 cm, whereas the overall uncertainty range is about 40 cm.



361 **Figure 9.** 2050 Bergen sea level projections with uncertainty due to different sources for RCP 8.5. The
 362 black vertical line is the median projection (with no uncertainty), while the gray histogram corresponds
 363 to the spread of the climate models, the red curve adds the uncertainty due to the relation between global
 364 temperature and global sea level, and the blue line that due to downscaling global sea level to Bergen.

365 6.2 Damage costs

366 A simplistic analysis of projected damage costs at the year 2100, not taking into
 367 account the uncertainty, would use the median historical damage cost multiplied by the
 368 median damage effect factor at 2100 at the median sea level rise projected for 2100. This
 369 yields a damage cost of NOK 338 million (the grey vertical line in Figure 10). Similar
 370 results obtain when allowing sea level or effect factor to vary, holding the other quantities
 371 at the median (yellow and purple dots on top of Figure 10). However, allowing only the
 372 damage cost to vary yields a median cost of 3.85 billion (green dot on top of Figure 10).
 373 The appropriate uncertainty analysis for our model should draw each of sea level, effect
 374 factor and damage cost at random from their 2100 distributions. This corresponds to a
 375 median cost of 3.15 billion NOK, over 9 times higher than the simplistic value. Over 99%
 376 of the costs in our simulation are higher than the simplistic median.



377 **Figure 10.** Simulated distribution of total damage cost for 2100 without adaptation on log-log scale (black
 378 histogram). We also show the distributions of costs varying only one aspect of the uncertainty (sea level rise
 379 in blue, effect multiplier in yellow, and damage cost in green), holding the other two at their median values.
 380 The grey vertical line is the result of holding all three factors at their median value. The median of each
 381 distribution is shown as a dot on top of the figure.

382 7 Conclusions and discussion

383 Our case studies demonstrate that it is possible to take uncertainty into account in
 384 deciding when and where to implement adaptation measures, even if one uses a light
 385 touch decision-making approach. If one fails to do so, bad scenarios, such as 95th per-
 386 centiles, can be a order of magnitude worse than what the planners are expecting. It is
 387 likely worthwhile to be pessimistic in the planning and in the projections.

388 We consider our case studies proofs of concept, which will be first steps in a se-
 389 quence of interactions with local planners and other end-users. This has to be an itera-
 390 tive and interactive process, as the decision framework provides actionable information
 391 to decision-makers, who will then make their own decisions. These decision can then be
 392 incorporated into the current adaptation strategy. Further simulation studies allow a con-
 393 tinued loop to identify potential vulnerabilities of the approaches across a wide range of
 394 possible futures.

395 Down the line we plan to develop a flexible and easy-to-use tool kit for decision-
 396 making under uncertainty regarding sea level rise. An initial step in this direction is the
 397 software used in this analysis, which is publicly available and using free software.

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 402 (ClimateXL). The source code for the analysis is implemented in the statistical program-
 403 ming language R (*R Core Team [2016]*) and is available on GitHub at <http://github.com/eSACP/SeaLevelDecisions/Code>.

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