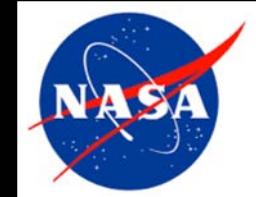


# First Estimate of Upper Tropospheric NO<sub>2</sub> from TROPOMI

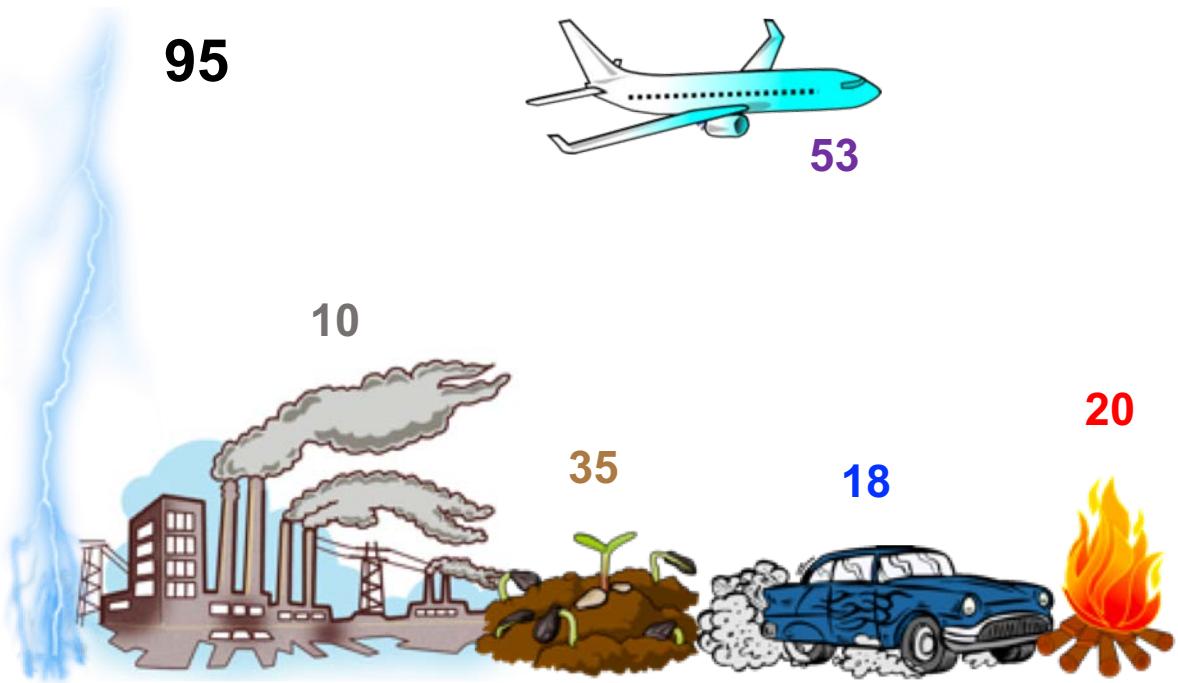


Eloise A. Marais, Nana Wei, Henk Eskes, Sungyeon Choi, Joanna Joiner, Alexander Cede

# Global Relevance of the Upper Troposphere (~8-12 km)

$\text{NO}_x$  ( $\text{NO} + \text{NO}_2$ ) is very efficient at forming ozone in the upper troposphere where ozone is a potent greenhouse gas

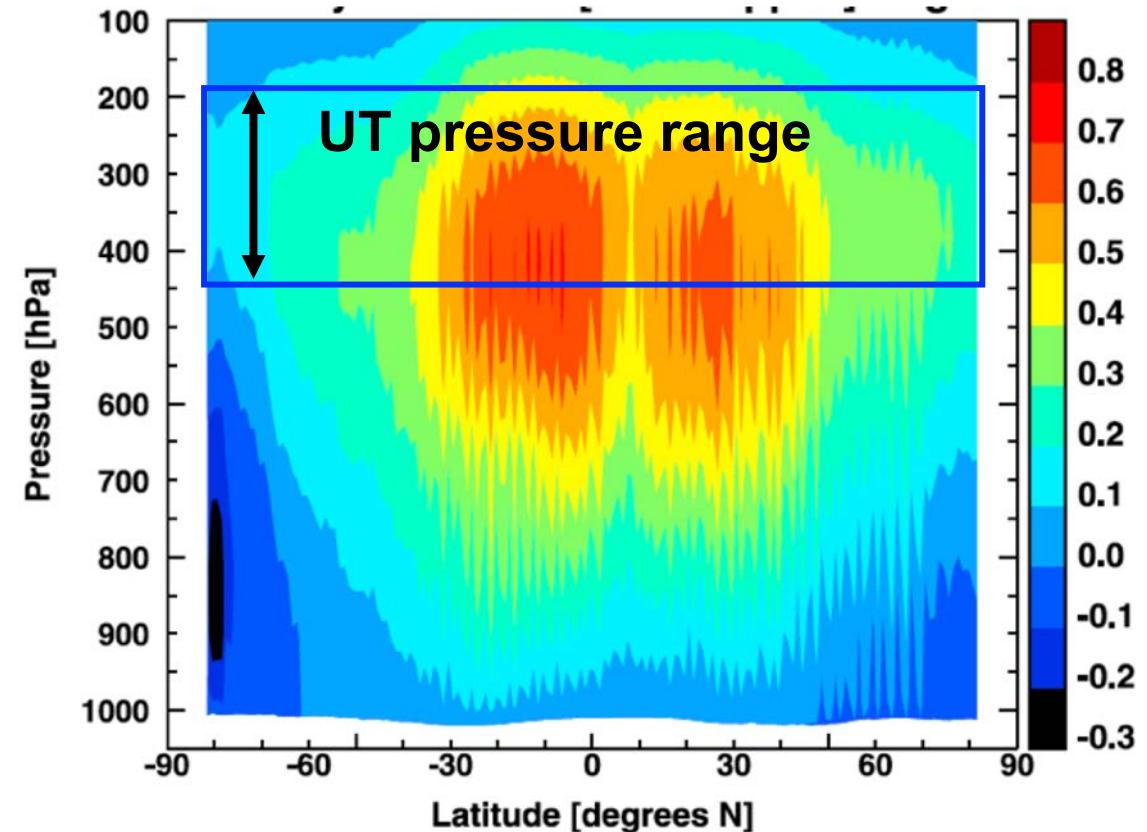
## OPE (molecules $\text{O}_3$ /molecule $\text{NO}_x$ ) for individual $\text{NO}_x$ sources



[adapted from Dahlmann et al., 2011]

Longer  $\text{NO}_x$  lifetime at higher altitude → greater OPE

## Sensitivity of outgoing longwave radiation to variations in ozone concentrations

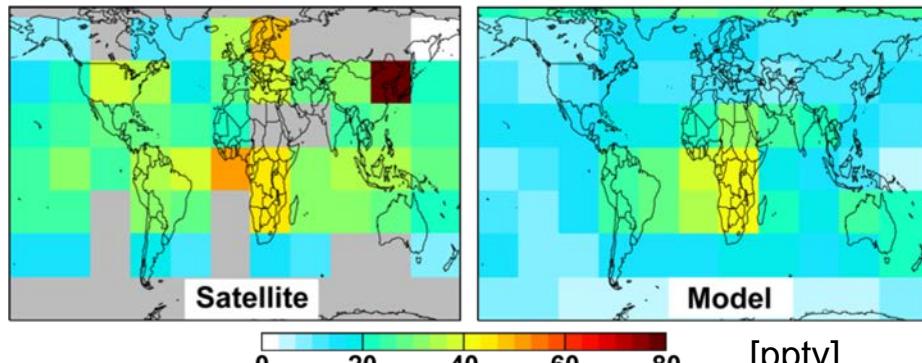


[Aghedo et al., 2011]

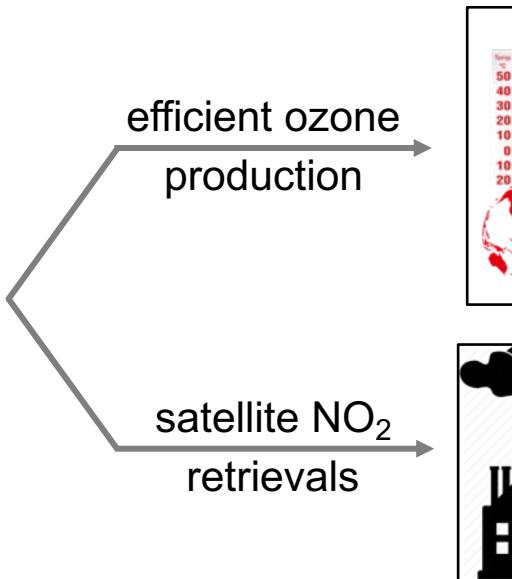
# Errors in the UT Impact the Whole Atmosphere

Small uncertainties in NO<sub>x</sub> in the UT induce large climate and air quality errors

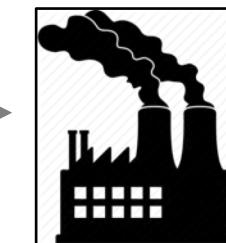
Model NO<sub>2</sub> too low across the UT in the northern midlatitudes



[Marais et al., 2018]

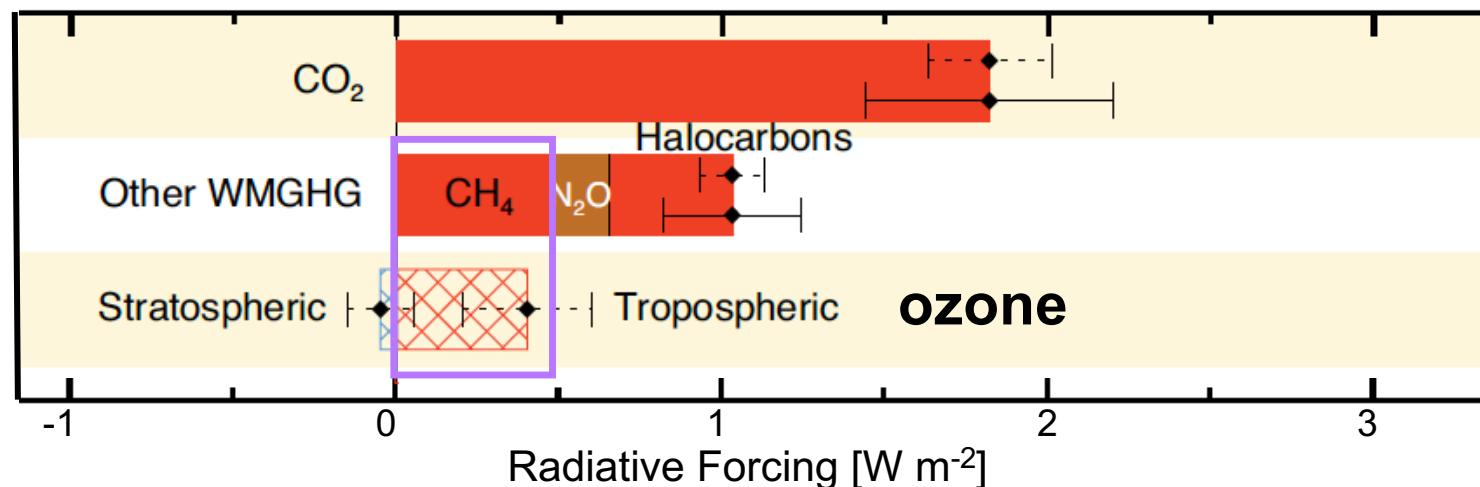


Errors in tropospheric ozone radiative forcing



Uncertain global air quality constraints from satellites

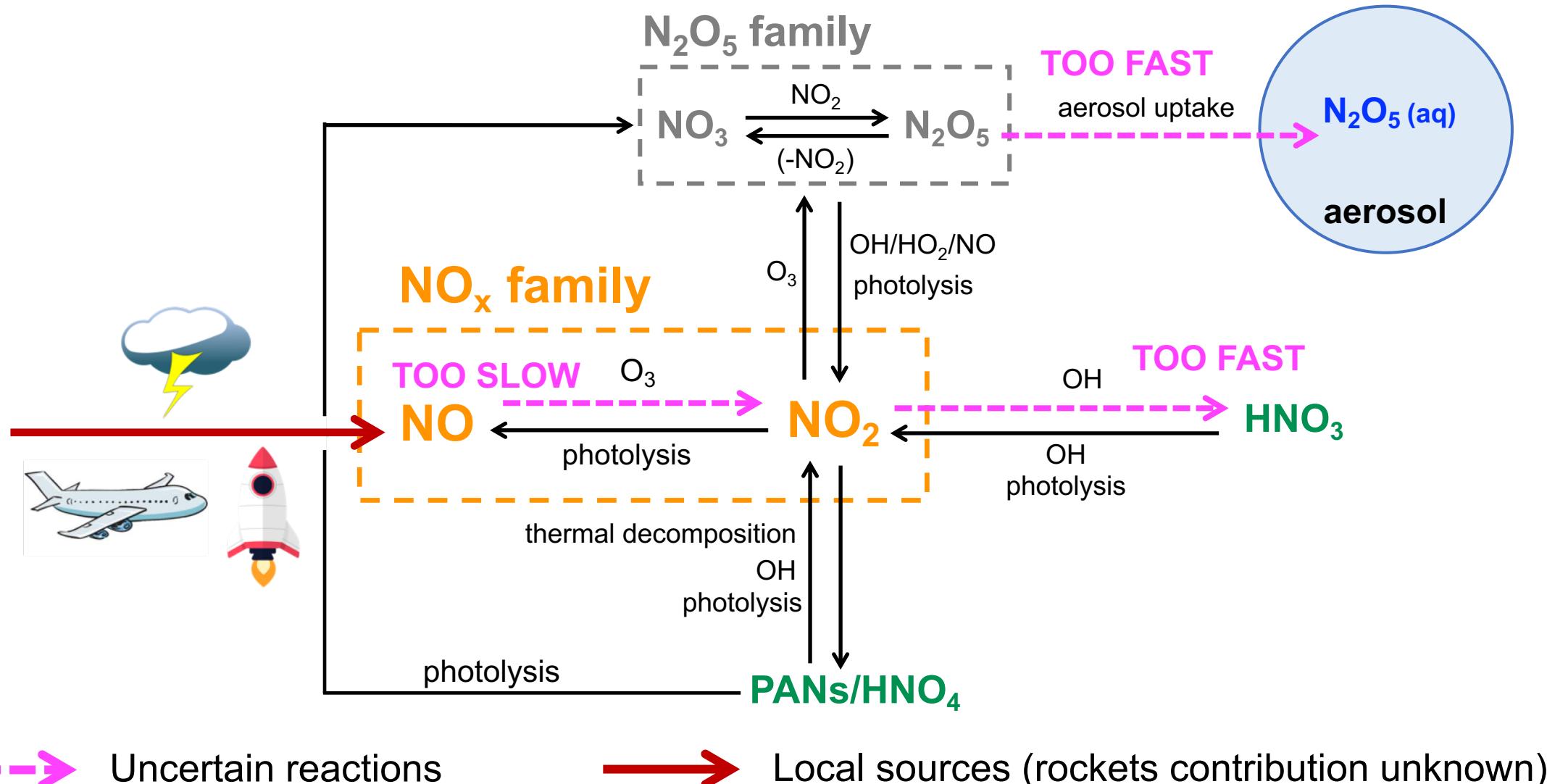
Tropospheric ozone and methane have near-equal climate impacts



[IPCC AR5, 2013]

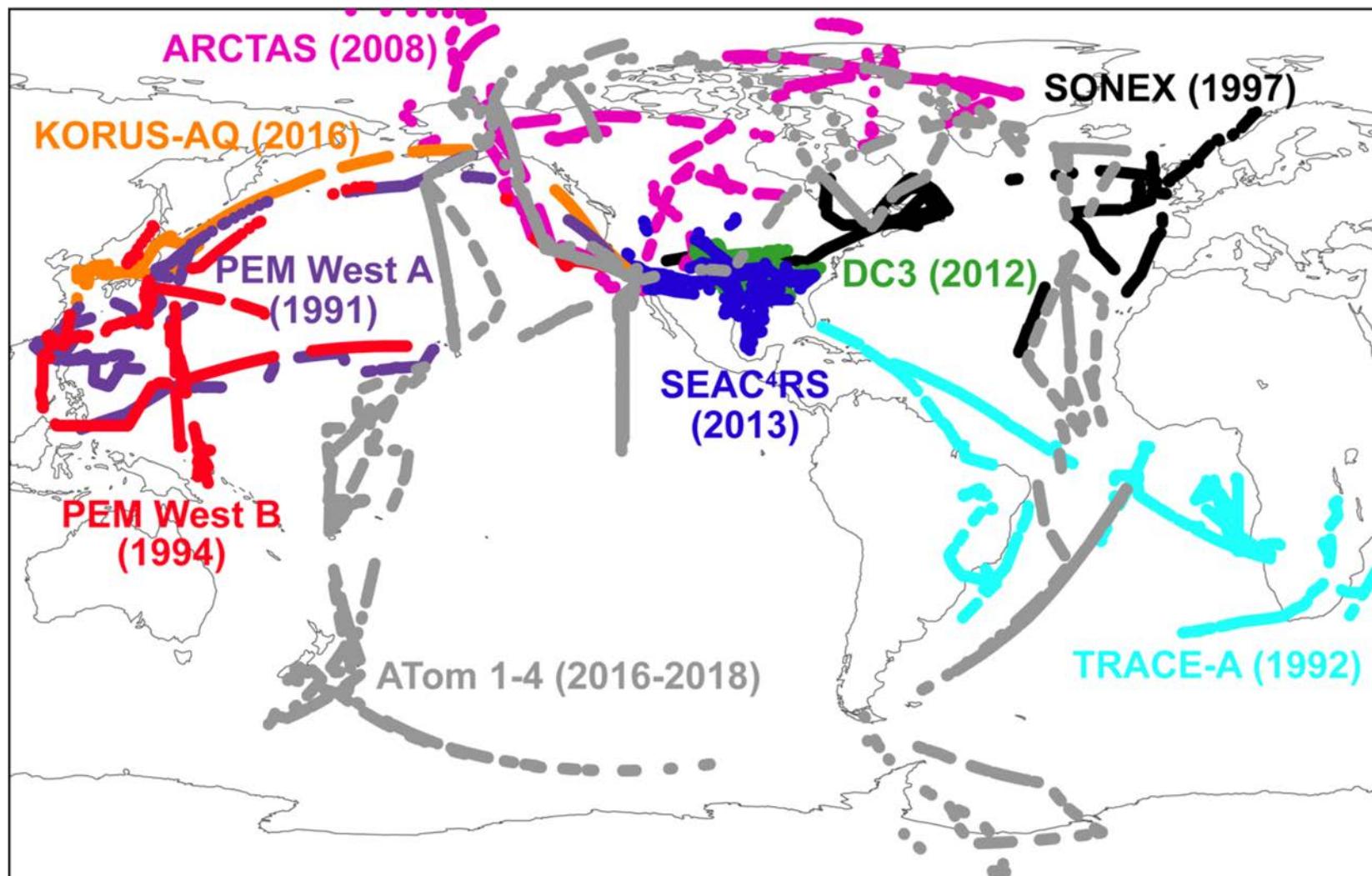
# Large Uncertainties in Chemistry of Reactive Nitrogen

Uncertainties in  $\text{NO}_x$  chemistry have implications for relative ( $\text{NO}:\text{NO}_2$ ) and total  $\text{NO}_x$  abundances and  $\text{NO}_x$  lifetimes.



# $\text{NO}_x$ Measurements are Limited in Time and Space

Spatial distribution of data from NASA DC8 aircraft campaigns in the upper troposphere

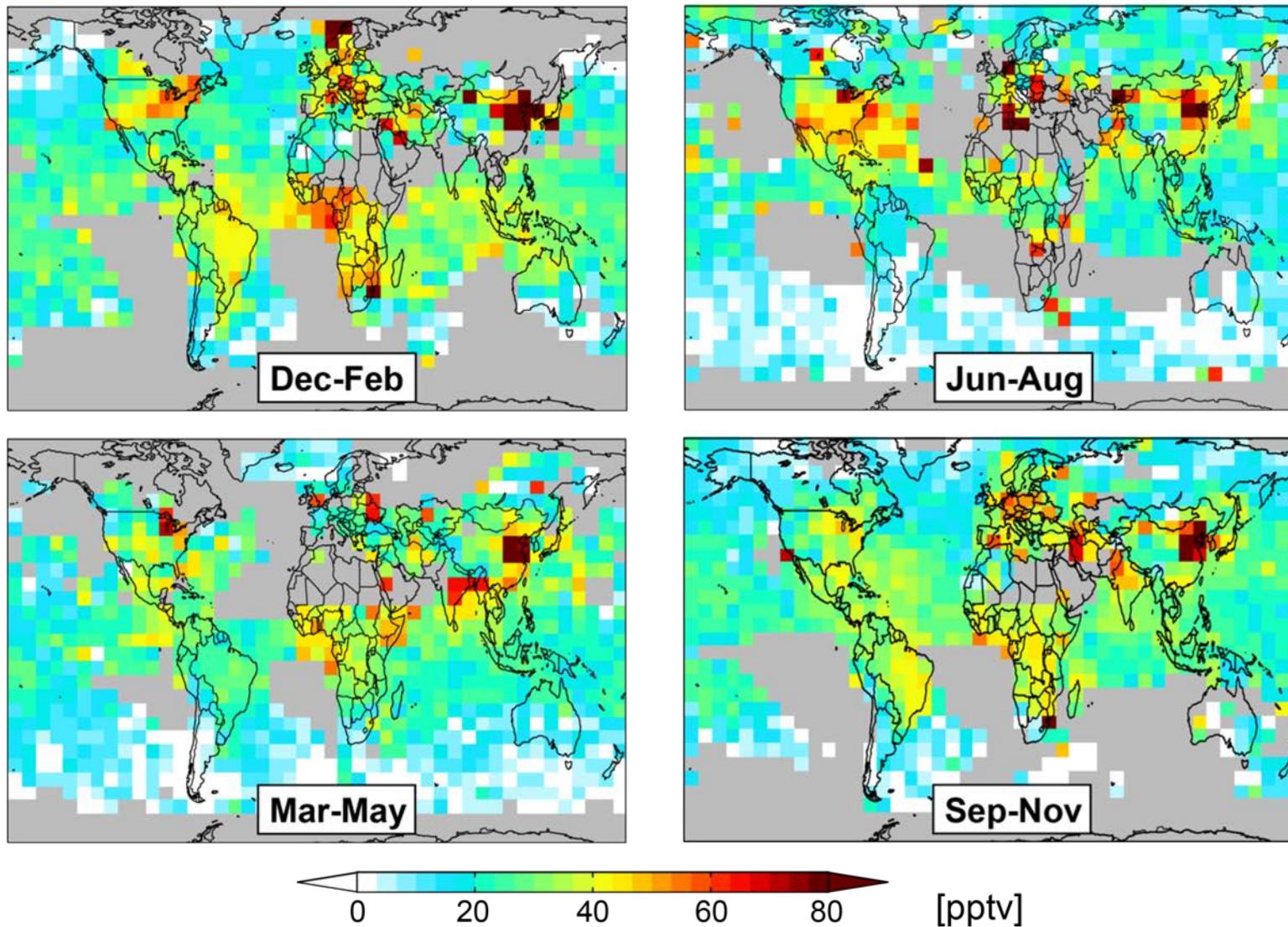


$\text{NO}_2$  in situ instruments are also susceptible to positive biases in the upper troposphere

# New UT NO<sub>2</sub> products from the Ozone Monitoring Instrument (OMI)

Near global spatial coverage of seasonal mean UT NO<sub>2</sub> at  $5^\circ \times 8^\circ$  (latitude  $\times$  longitude)

NASA OMI upper troposphere NO<sub>2</sub> (2005-2007)



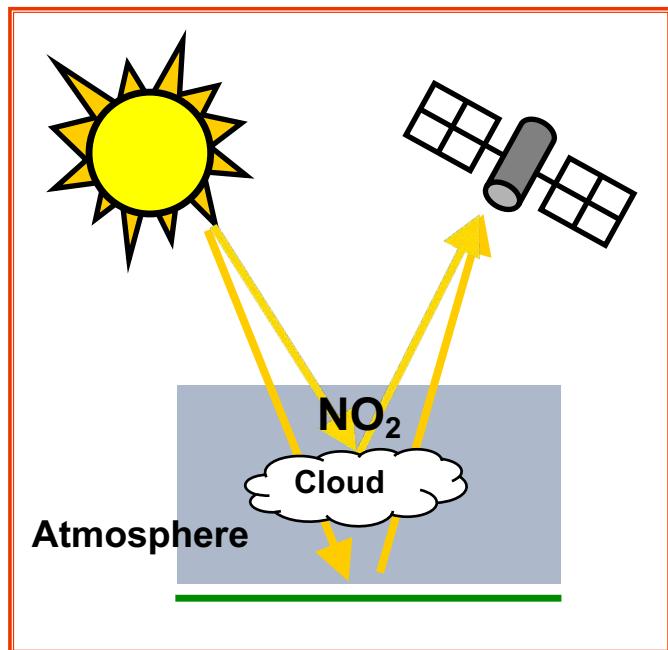
[Marais et al., 2018;  
Choi et al., 2014]

# UT NO<sub>2</sub> Obtained with the Innovative Cloud-Slicing Technique

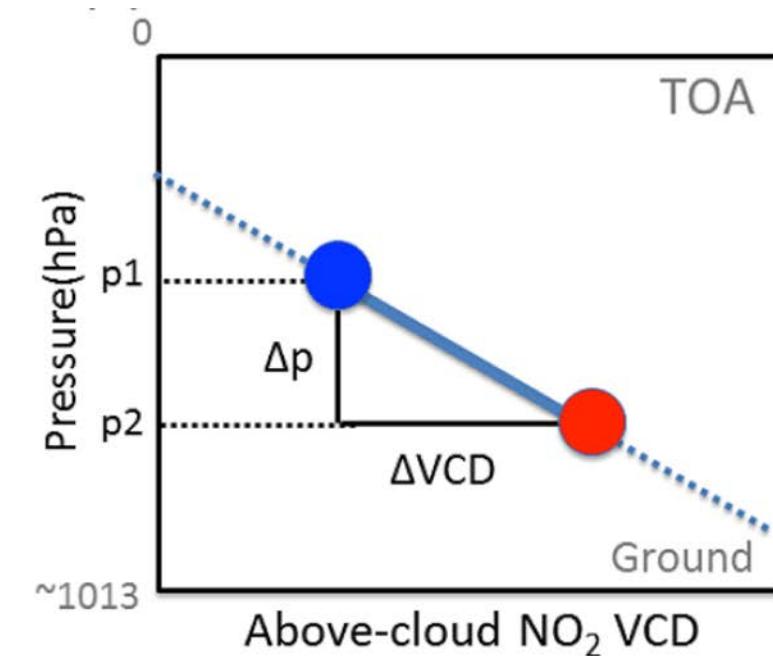
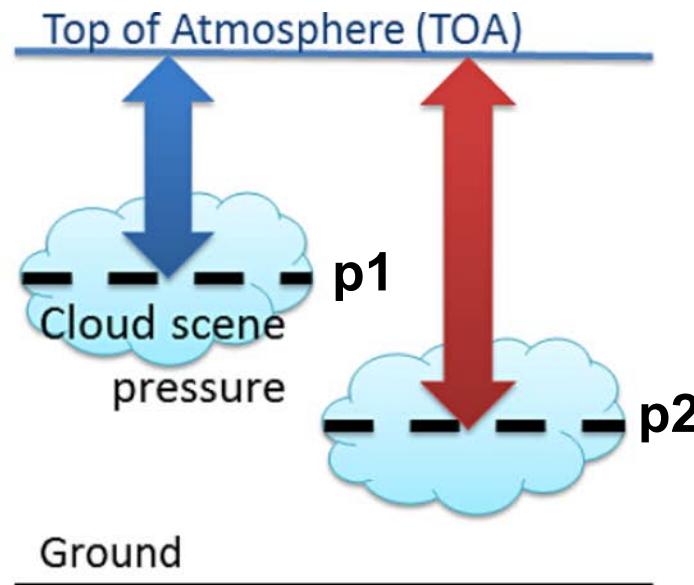
First applied by Ziemke et al. [2001] to TOMS ozone

Retrieve partial NO<sub>2</sub> columns over cloudy scenes at different heights

## APPROACH



## Use cloud height variability to derive partial columns



[adapted from Choi et al., 2014]

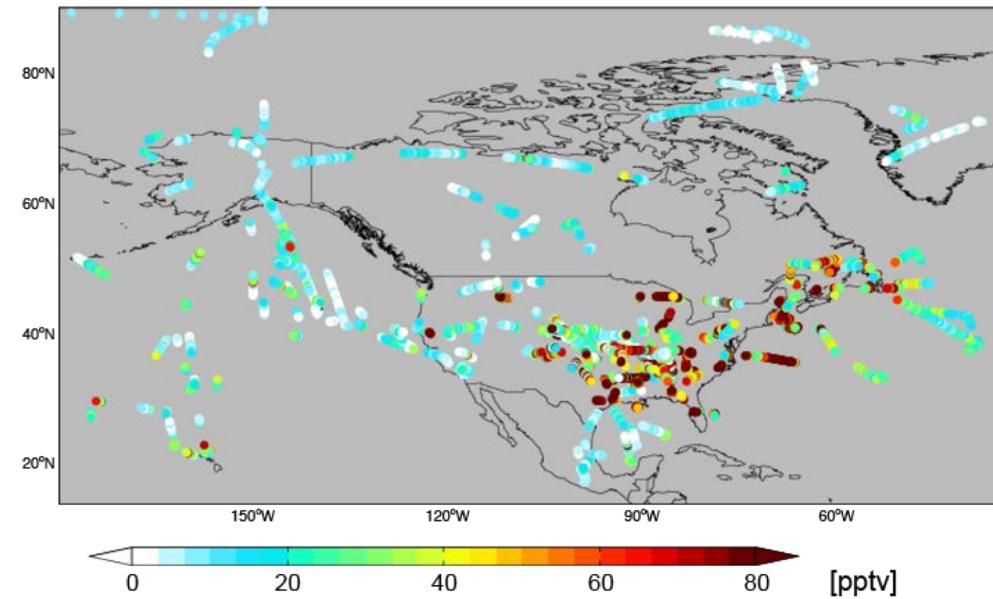
NO<sub>2</sub> volume mixing ratio (VMR) between clouds at p1 and p2

$$\text{NO}_2 \text{ VMR} = \frac{\Delta \text{VCD}}{\Delta p} \times \frac{k_B g}{R_{\text{air}}}$$

# Consistency with Aircraft Observations of NO<sub>2</sub>

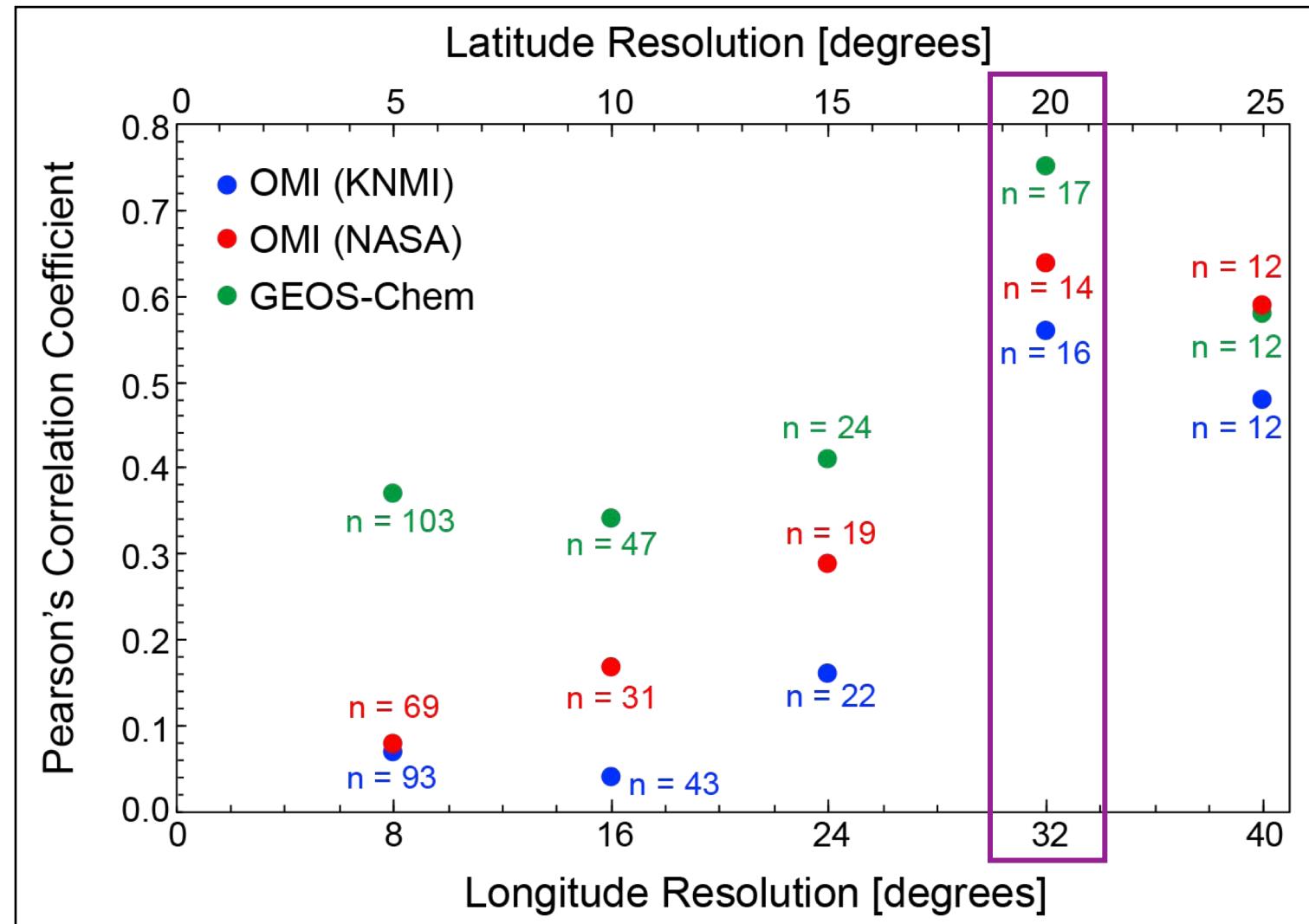
Agreement in spatial distribution between (reliable) aircraft and OMI UT NO<sub>2</sub> observations, but at coarse scales (seasonal, 20° latitude (2000 km) × 32° longitude (3200 km))

NASA DC8 NO<sub>2</sub> in Spring-Summer



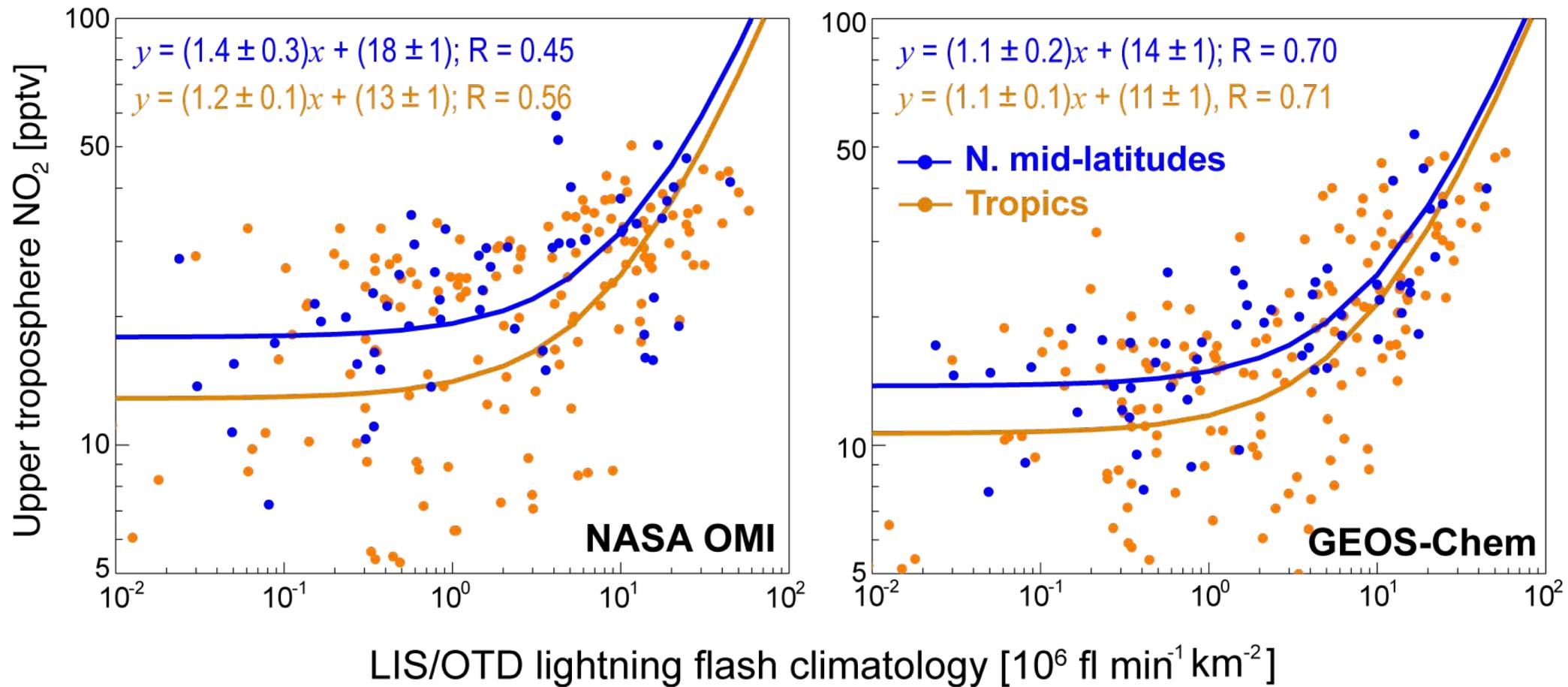
NO<sub>2</sub> obtained for multiple campaigns  
from Ron Cohen's TD-LIF instrument

[Marais et al., 2018]



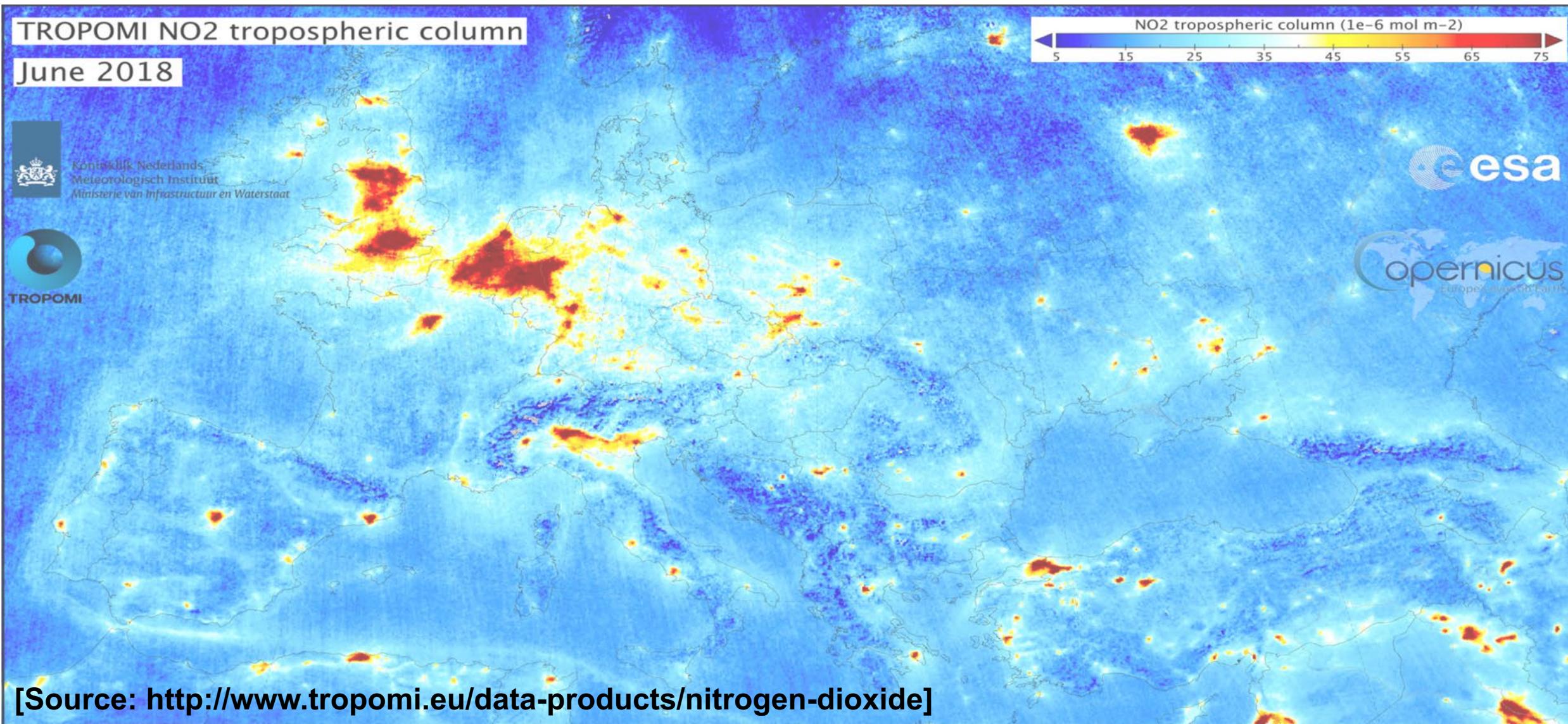
# Used OMI UT NO<sub>2</sub> to Better Constraint Global Lightning NO<sub>x</sub>

Log-log relationship between UT NO<sub>2</sub> from OMI and GEOS-Chem and satellite observations of lightning flashes in the **northern midlatitudes** and **tropics**



Similar slope in northern midlatitudes and tropics supports similar lightning NO<sub>x</sub> production rates

# Can We Do Better with the High-Resolution TROPOMI?

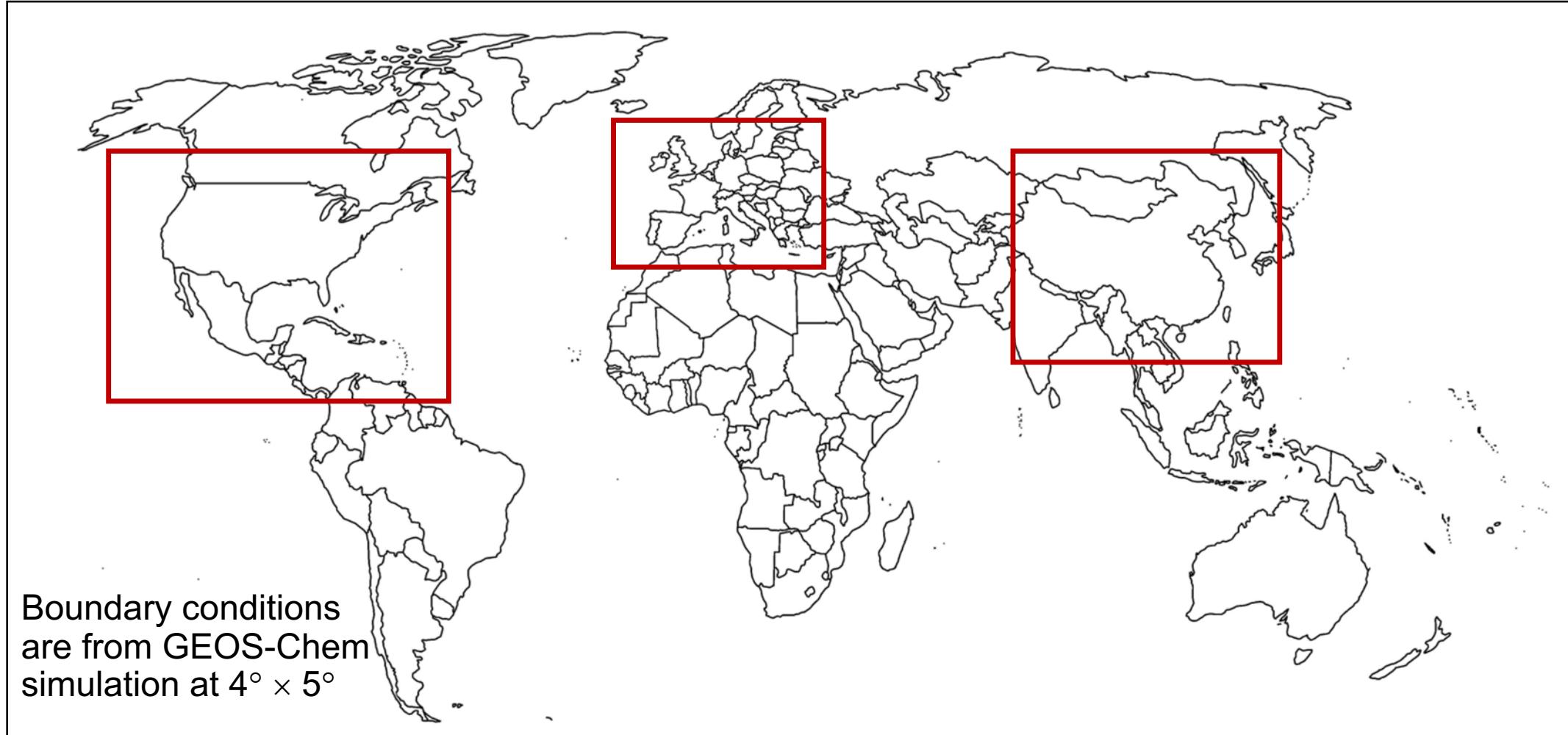


Nadir spatial resolutions in km (along × across): **13 × 24 (OMI)**; **5.6 × 3.5 (TROPOMI)**

# Synthetic Cloud-Slicing Experiment with GEOS-Chem

Determine whether UT NO<sub>2</sub> obtained with cloud-slicing reproduces UT NO<sub>2</sub> obtained by averaging across the UT (450-180 hPa)

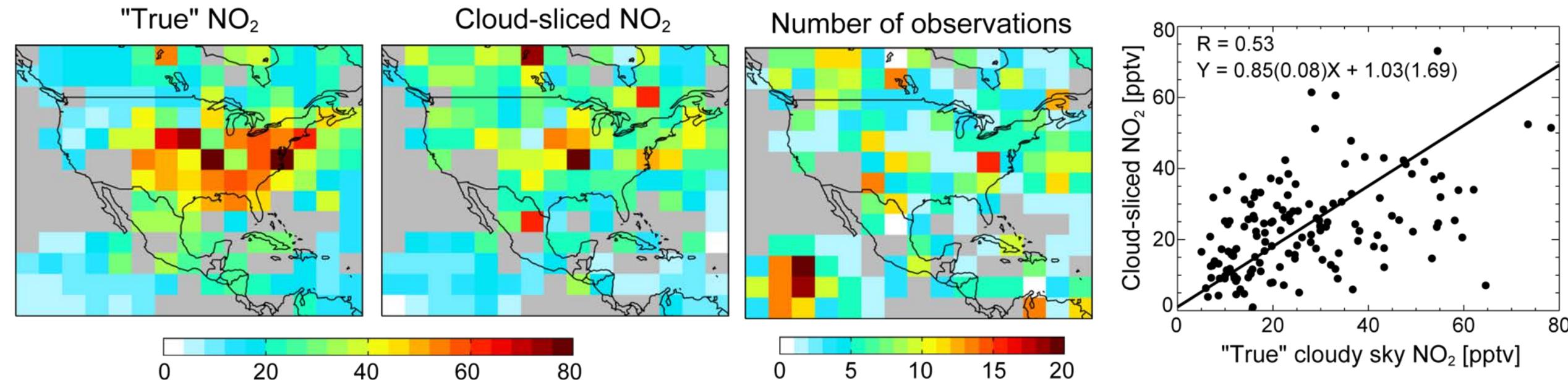
GEOS-Chem nested domains at  $0.25^\circ \times 0.3125^\circ$  driven with NASA GEOS-FP meteorology



# Synthetic Cloud-Slicing Experiment with GEOS-Chem

Compare “true” and cloud-sliced UT NO<sub>2</sub> from GEOS-Chem over North America in June-August

Daily cloud-sliced NO<sub>2</sub> obtained at  $4^\circ \times 5^\circ$  from daily nested grid GEOS-Chem partial columns sampled around the satellite overpass



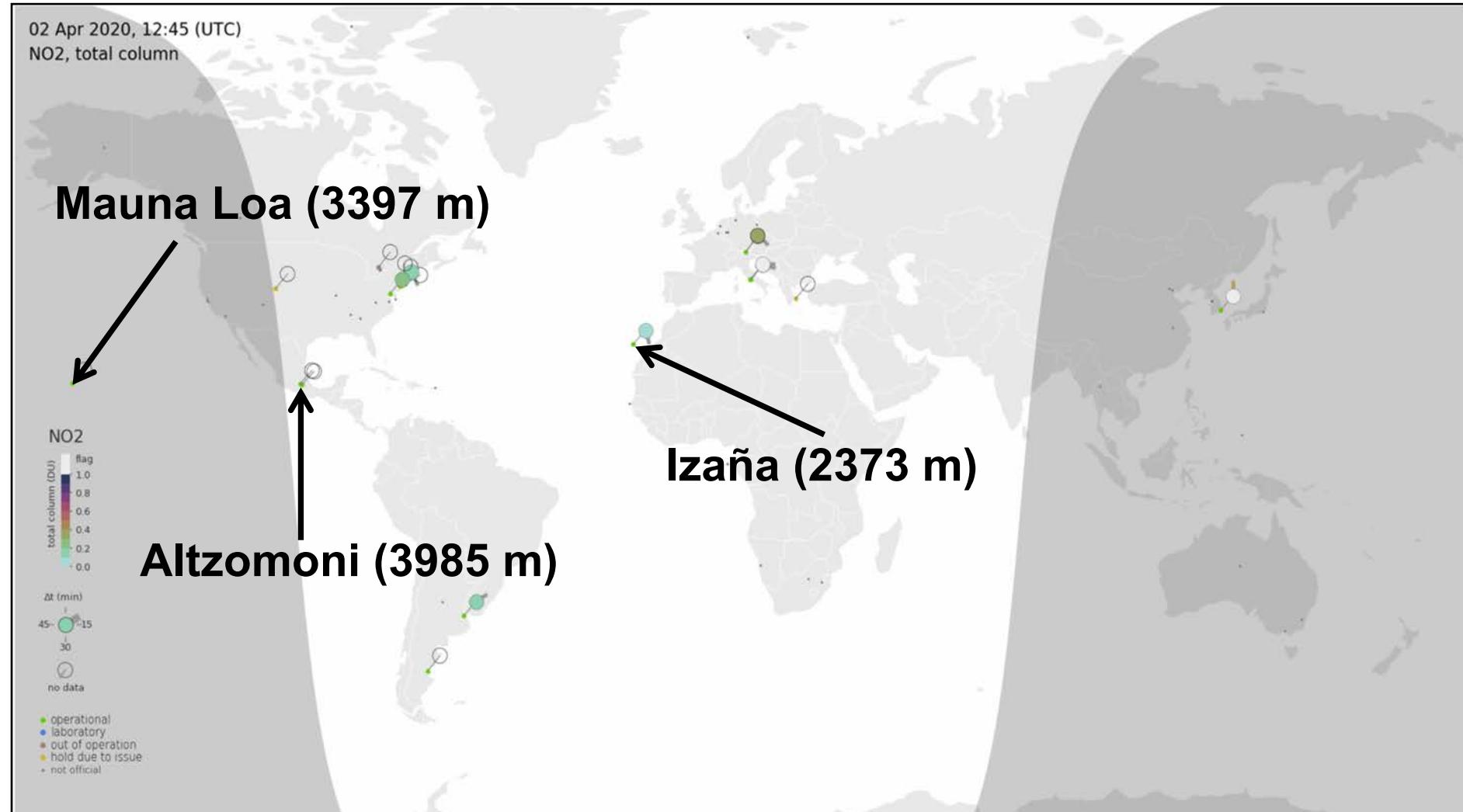
## Sampling conditions informed by Choi et al. [2014] approach:

Satellite overpass (12h00-15h00), dynamic cloud top height range (standard deviation > 35 hPa, range > 160 hPa), and UT NO<sub>2</sub> with relative error < 100%, < 200 pptv, significantly greater than zero.

Initial results are encouraging ( $R > 0.5$ ). Data density is an issue. Simulating more years.

# Evaluate TROPOMI total column NO<sub>2</sub> against Pandora

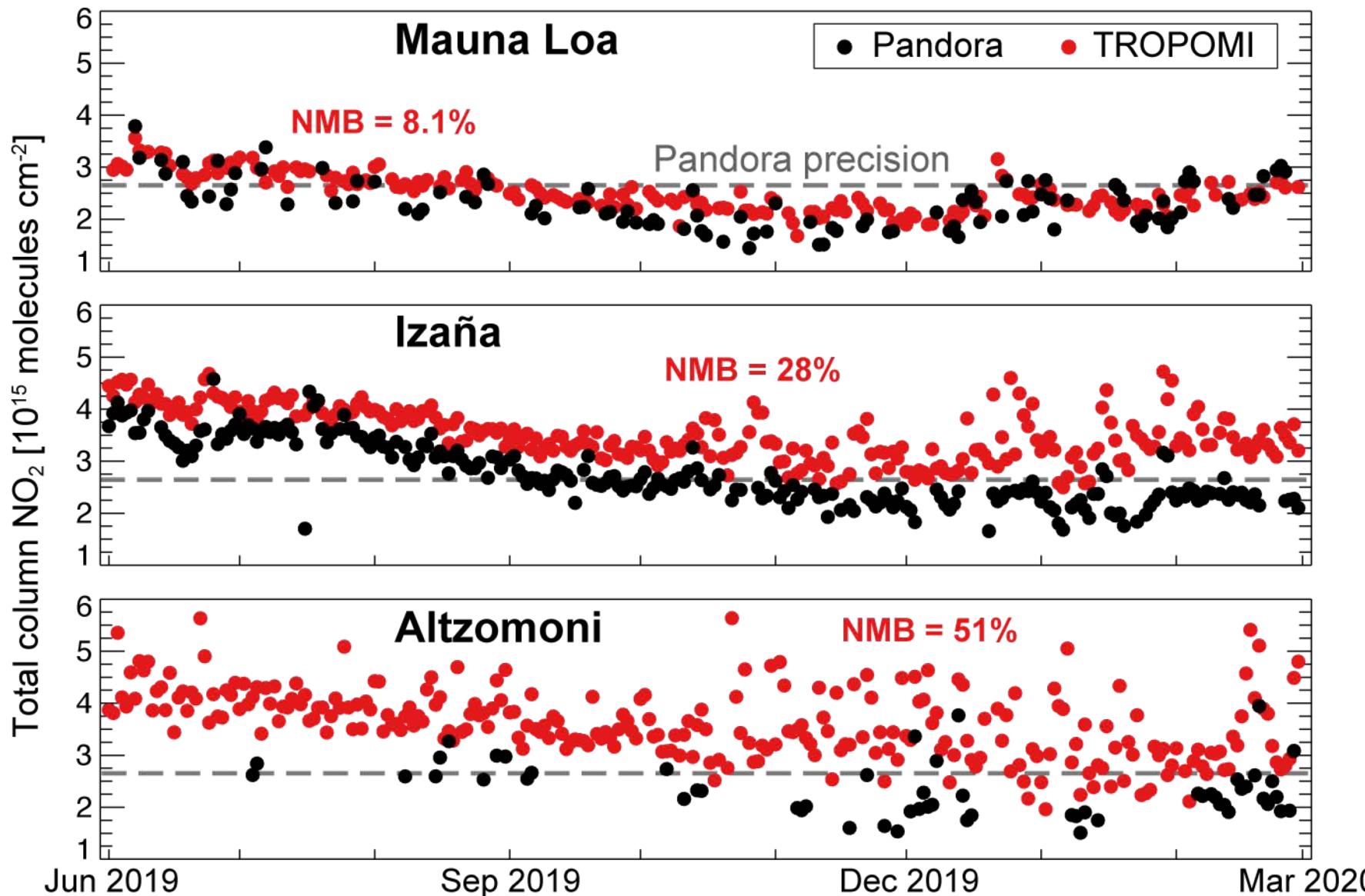
Global Pandora network, indicating locations of high-altitude sites (large relative contribution from the UT) used to evaluate TROPOMI



[Source: <https://www.pandoria-global-network.org/>]

# Evaluate TROPOMI total column NO<sub>2</sub> against Pandora

Pandora and TROPOMI total column NO<sub>2</sub> time series at the 3 measurement stations



Sample Pandora  $\pm$  30 min of TROPOMI overpass and TROPOMI within 20 km of Pandora site

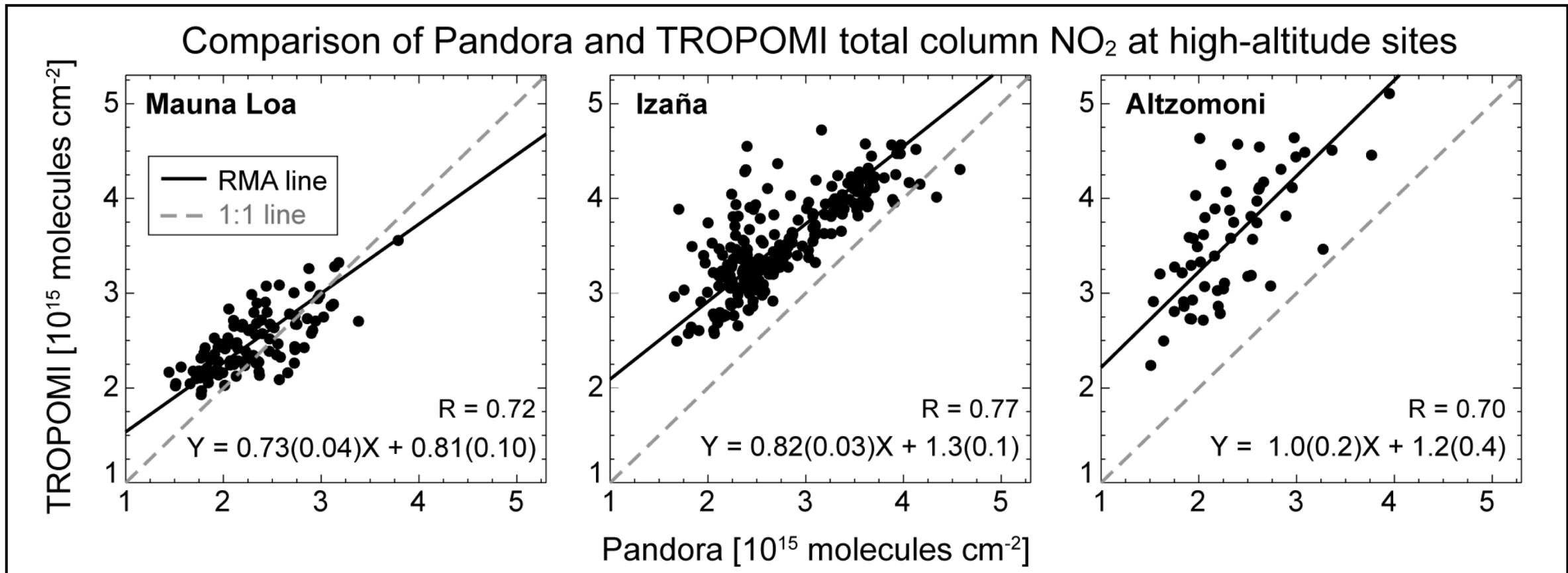
Focus is on cloudy scenes, so use a TROPOMI qa\_value threshold of 0.45

Data close to estimated Pandora precision

TROPOMI normalized mean bias (NMB) for coincident observations is 8.1-51%

# Evaluate TROPOMI total column NO<sub>2</sub> against Pandora

Scatterplots of TROPOMI versus Pandora total column NO<sub>2</sub> at the 3 stations



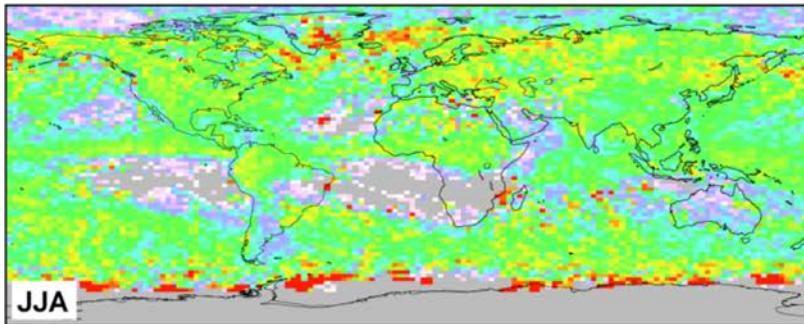
Positive intercept supports systematic bias in TROPOMI total column

TROPOMI validation documents show a slight underestimate (10%) in the stratospheric column

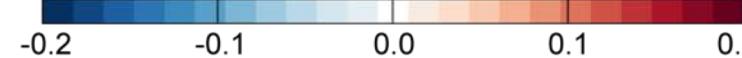
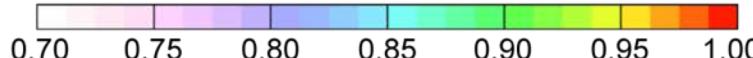
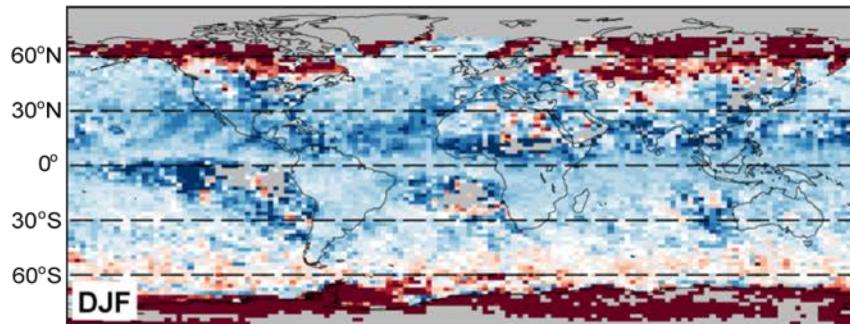
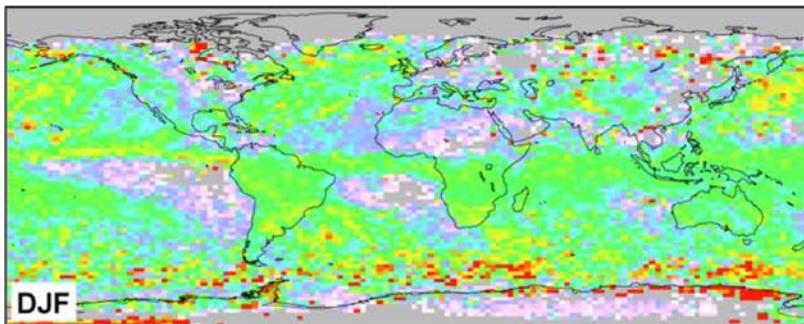
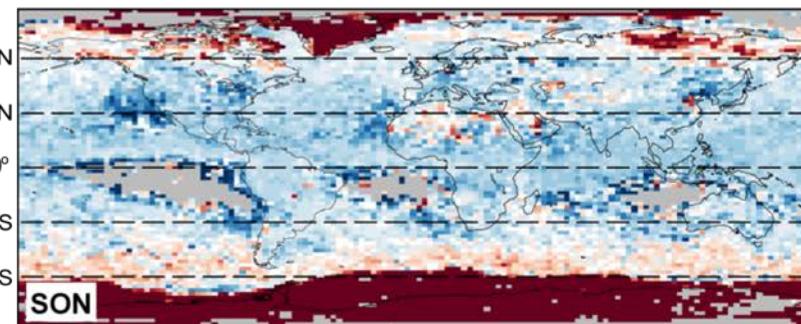
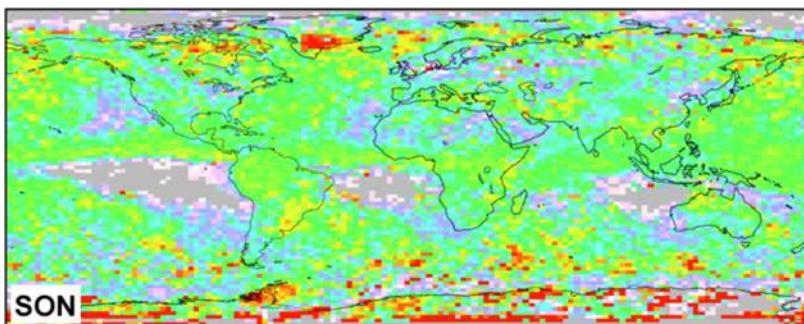
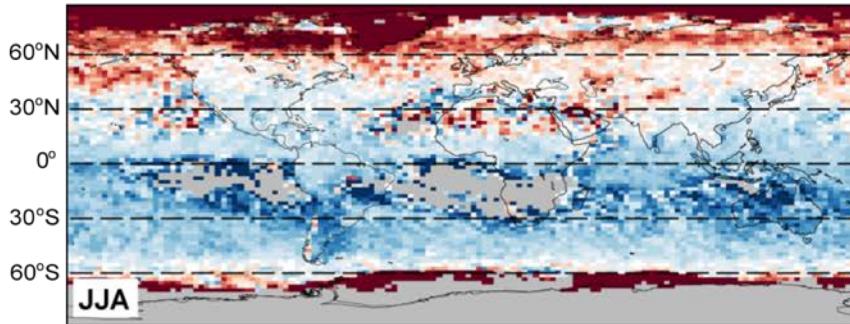
# Evaluate TROPOMI Cloud Fractions

Seasonal mean FRESCO cloud fraction and the difference between FRESCRO and DLR-OCRA

FRESCO cloud fraction



FRESCO minus DLR-OCRA cloud fraction



Only consistent data versions are used from June 2019 to February 2020

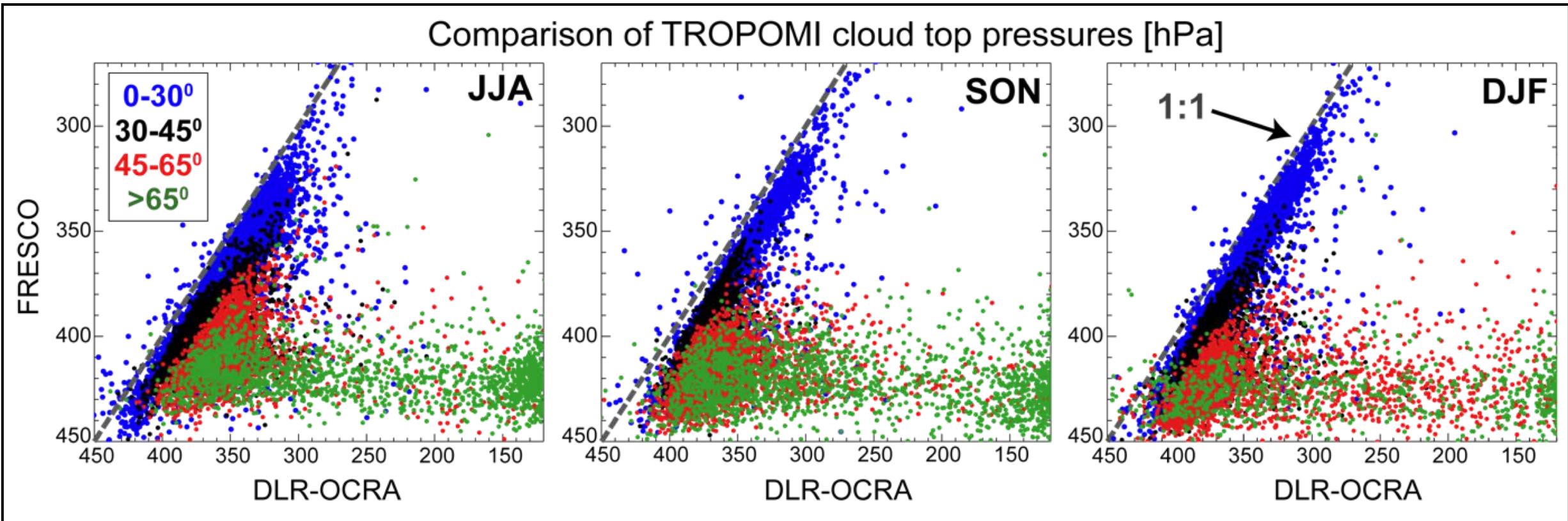
**FRESCO:** cloud product used in the TROPOMI NO<sub>2</sub> column retrieval

**DLR-OCRA:** Official TROPOMI cloud product (\*OFFL\_L2\_CLOUD\*)

Compare scenes with FRESCO cloud fraction > 0.7

# Evaluate TROPOMI Cloud Top Heights

Scatterplots of seasonal mean of TROPOMI cloud products at  $2^\circ \times 2.5^\circ$  colored by latitude bands



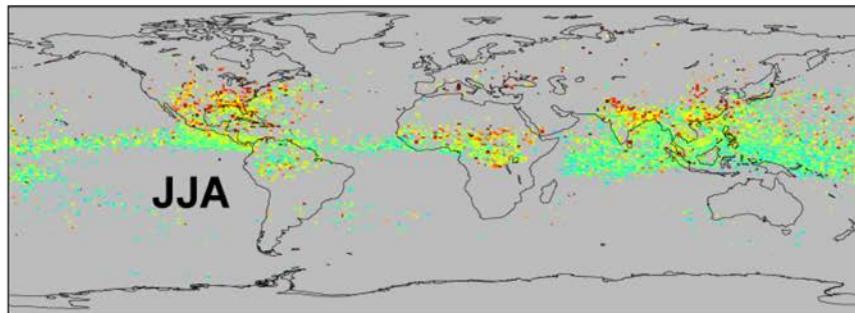
Compare scenes with FRESCO cloud top pressure range of 180 to 450 hPa

DLR-OCRA systematically lower than FRESCO and discrepancy increases with latitude

At poles ( $>65^\circ$  N/S) DLR-OCRA ranges from <150-450 hPa, whereas FRESCO range is 400-450 hPa

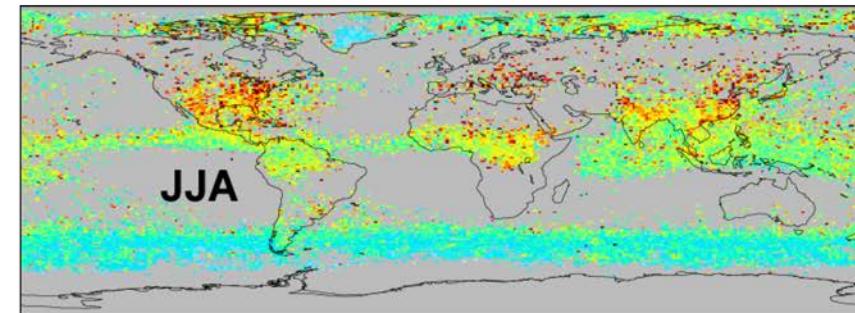
# Seasonal Mean TROPOMI Cloud-sliced UT NO<sub>2</sub>

TROPOMI-FRESCO UT NO<sub>2</sub>

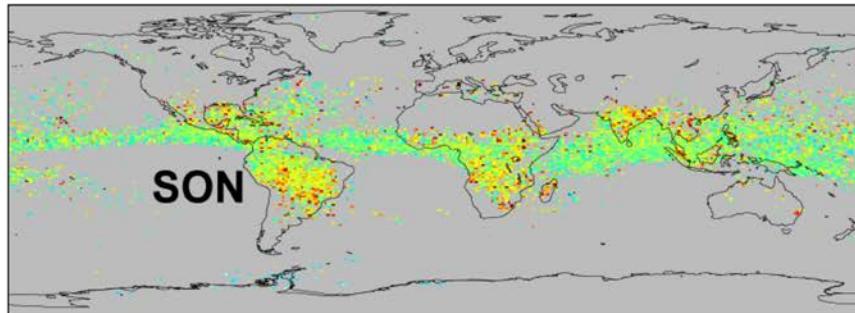


Data are  
at  $1^\circ \times 1^\circ$

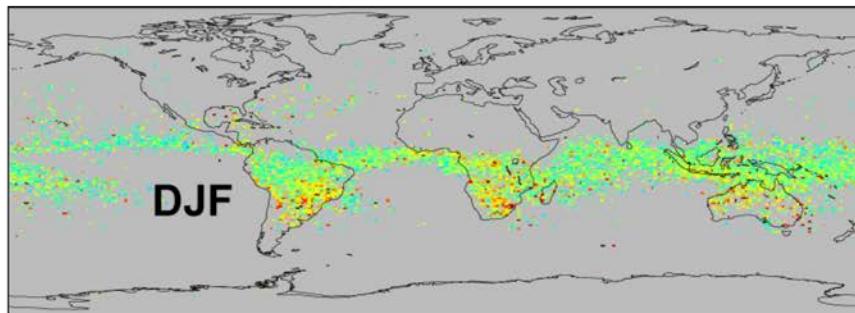
TROPOMI-DLR-OCRA UT NO<sub>2</sub>



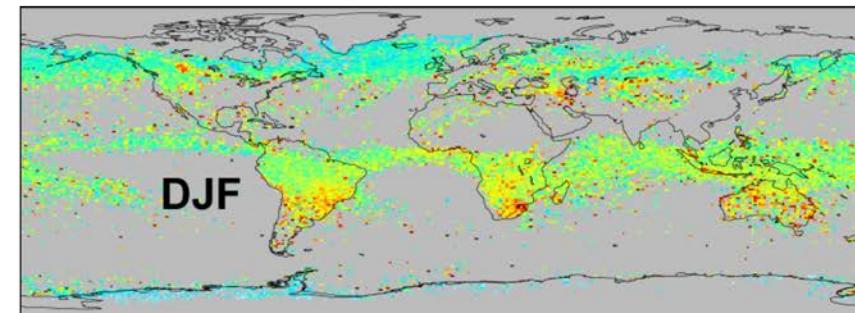
SON



SON



DJF

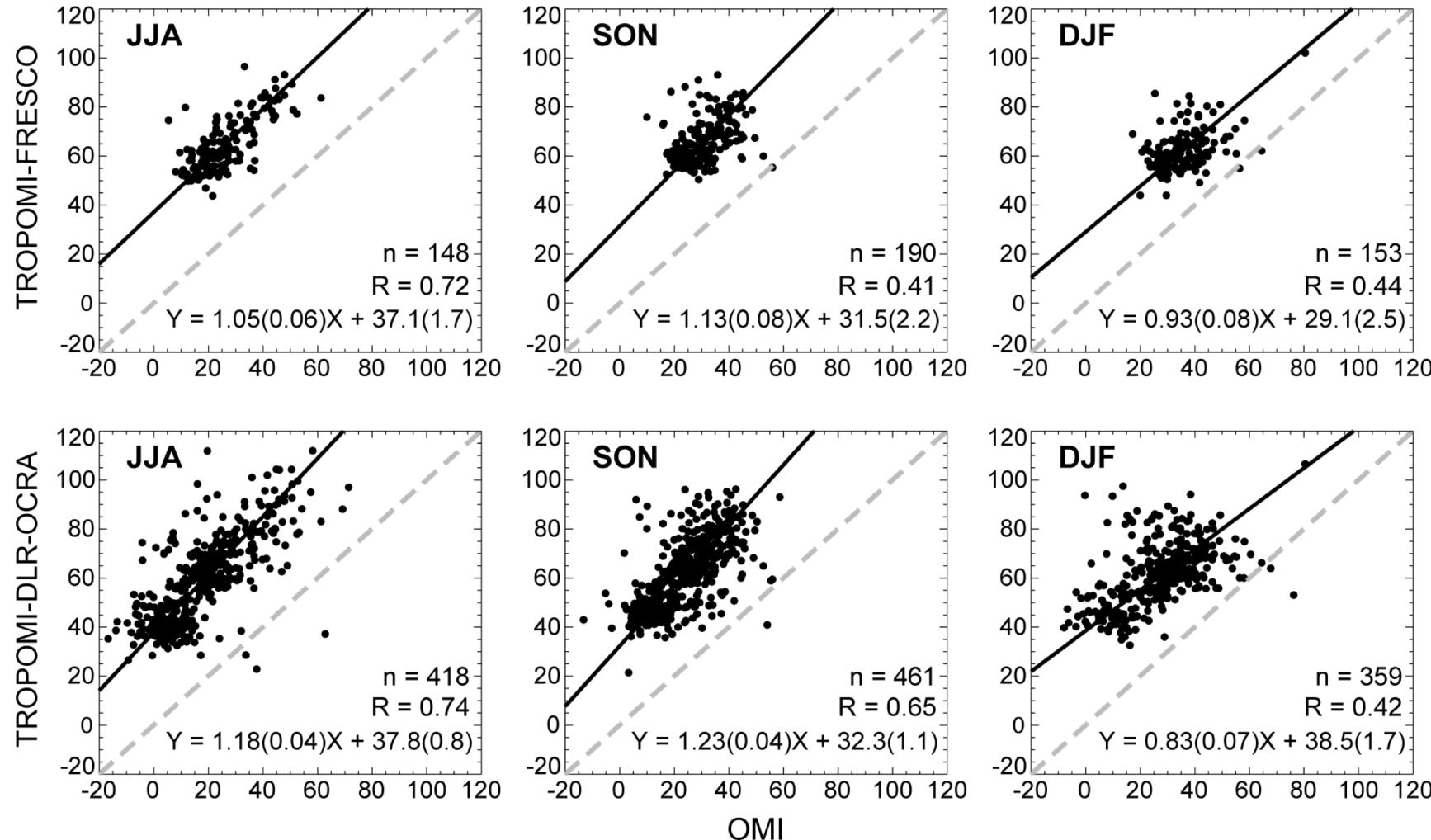


Results consistent ( $R > 0.6$ ; slope  $\sim 1$ ) for coincident scenes (mostly the tropics)

# TROPOMI and OMI UT NO<sub>2</sub> Comparison

TROPOMI UT NO<sub>2</sub> obtained at 1° × 1° and gridded to the NASA product resolution (8° × 5°)

Comparison of TROPOMI and OMI seasonal mean UT NO<sub>2</sub> [pptv]



Greater data density and better R with DLR-OCRA. TROPOMI bias consistent with Pandora comparison

## Concluding Remarks

The cloud-sliced approach induces a small low bias in UT NO<sub>2</sub> variance, but this may be impacted by sampling density.

TROPOMI is biased high in the free troposphere compared to Pandora measurements. The cause is not yet apparent.

Large differences in cloud fractions and cloud top heights from different products beyond the tropics and subtropics.

Using DLR-OCRA product yields greater global coverage than using the FRESCO product.

Spatial consistency between OMI and TROPOMI UT NO<sub>2</sub> products

## What's Next

Evaluate GEOS-Chem synthetic cloud-slicing over Europe and China

Independent assessment against reliable aircraft NO<sub>2</sub> observations from NASA DC8 campaigns