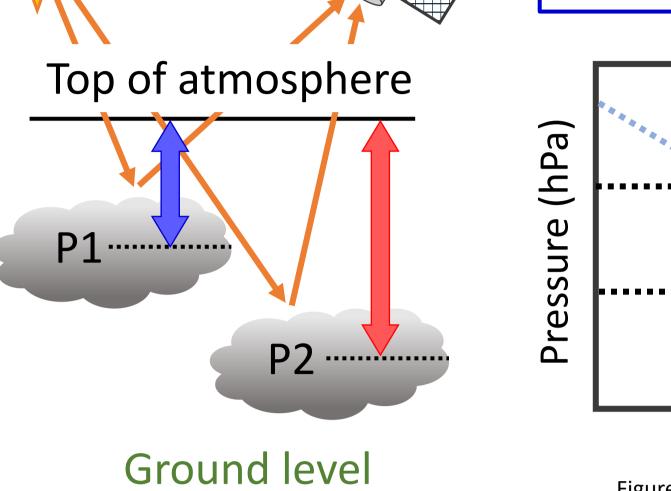
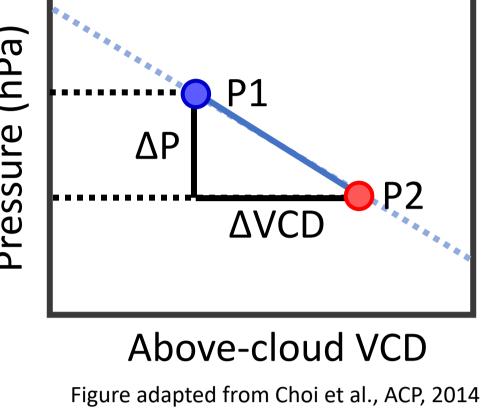


# Challenging our understanding of upper tropospheric NOx

Robert G. Ryan<sup>1</sup>, Eloise A. Marais<sup>1</sup>, Nana Wei<sup>1</sup>

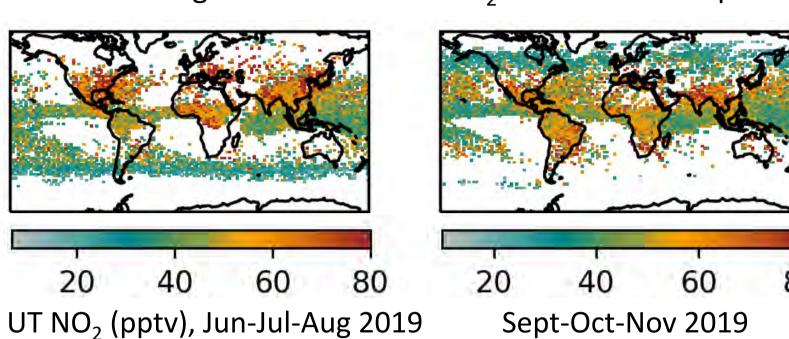
# Cloud-slicing of satellite data Top of atmosphere

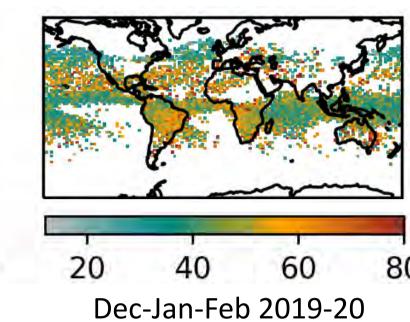


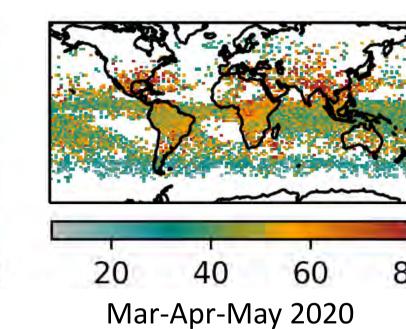


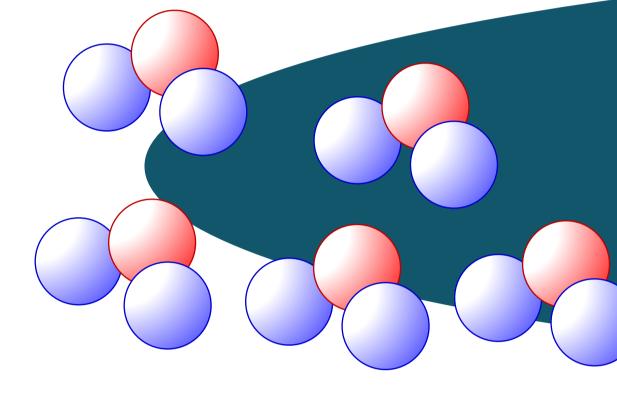
### Importance of the upper troposphere (UT)

- Upper tropospheric (~450-180 hPa) nitrogen oxides (NO + NO<sub>2</sub>  $\equiv$  NO<sub>x</sub>) are especially long-lived. This allows UT NO<sub>x</sub> to play a key role in the chemistry of tropospheric  $O_3$ , an important greenhouse gas.
- UT NO<sub>x</sub> has natural (lightning) and anthropogenic (aircraft) sources • UT reactive nitrogen chemistry is controlled by cycling between NO<sub>x</sub> and its reservoir compounds (HNO<sub>4</sub>, HNO<sub>3</sub>, PAN, &
- PPN). Cycling rates in this cold, low-pressure part of the atmosphere are uncertain. • UT NO<sub>x</sub> can be measured in-situ by aircraft, but these measurements are subject to interferences
- Cloud-slicing total columns of NO<sub>2</sub> from TROPOMI provides near-global coverage of NO<sub>2</sub> in the UT in each season





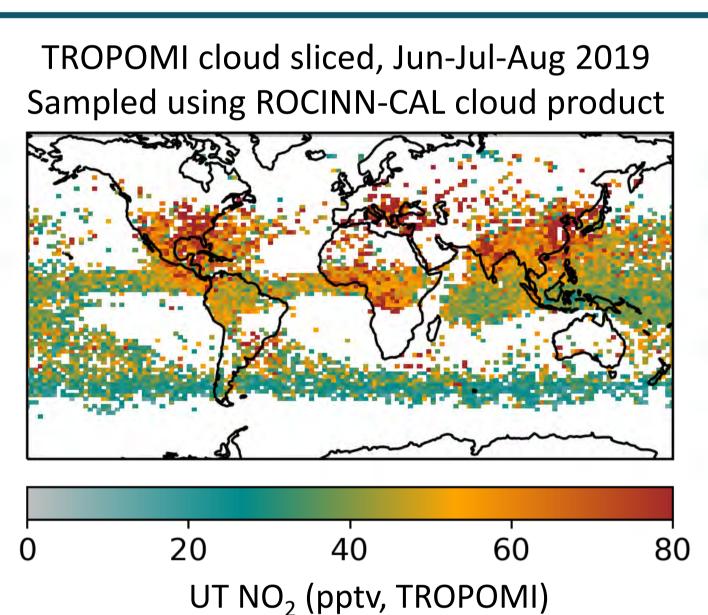


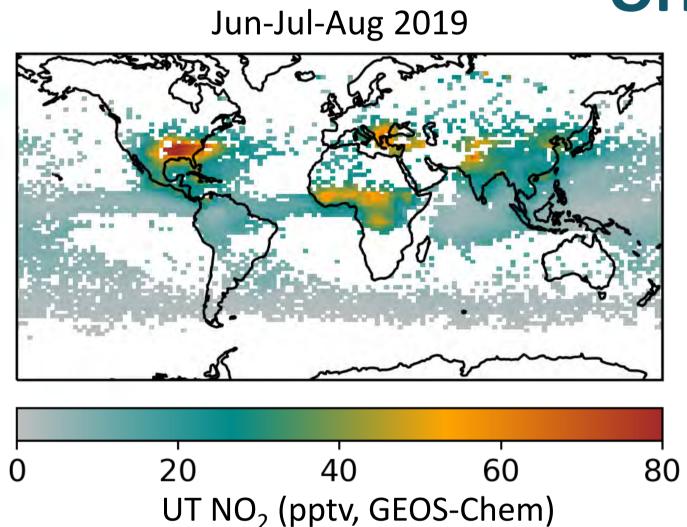


## WHY DO MODELS UNDERESTIMATE UPPER

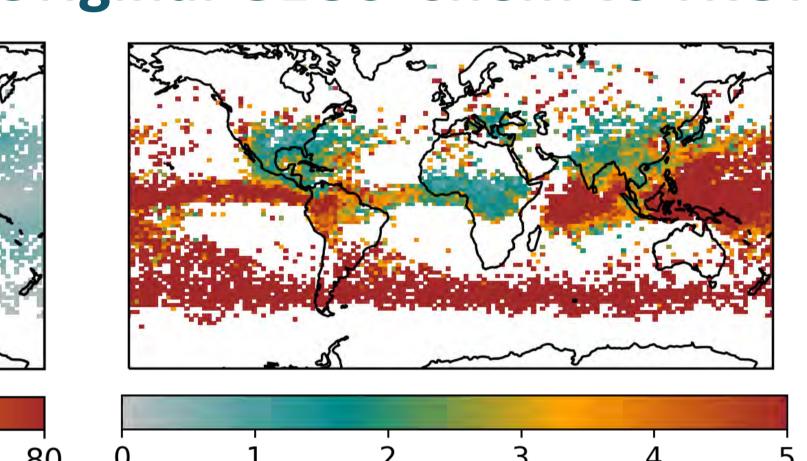
TROPOSPHERIC NO<sub>2</sub>?

We look at the role of chemistry and lightning

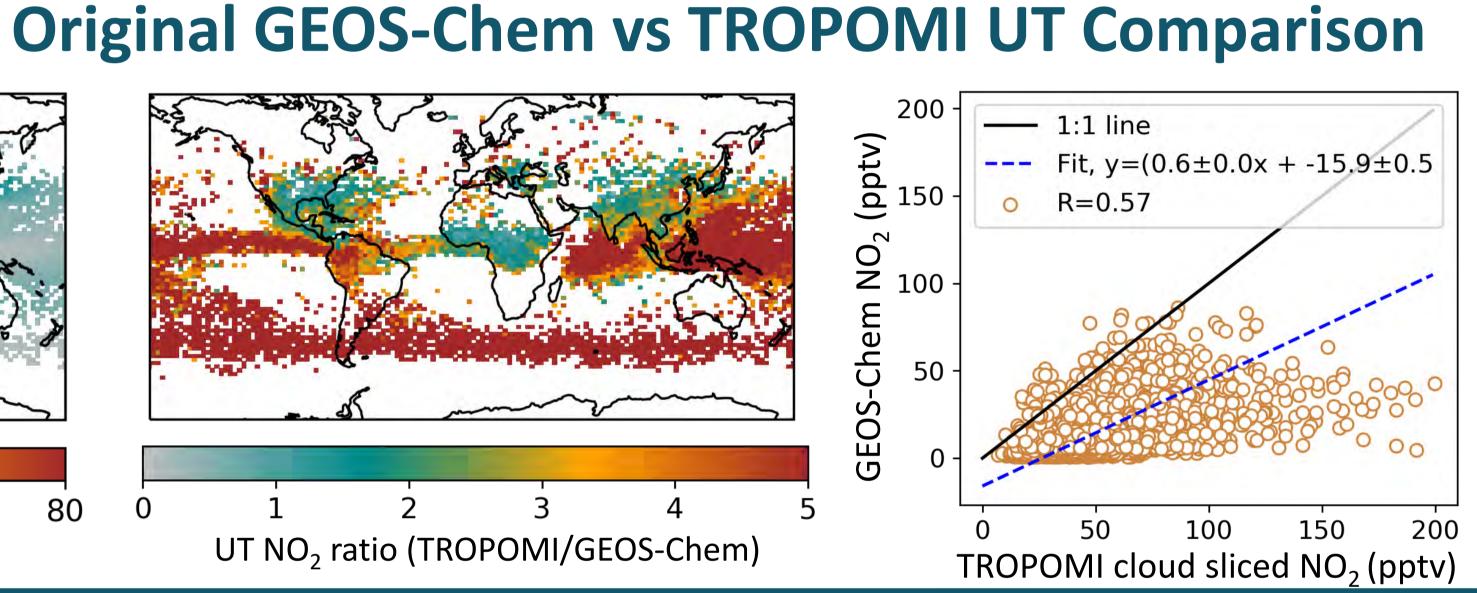




**GEOS-Chem results** 



UT NO<sub>2</sub> ratio (TROPOMI/GEOS-Chem)



### Zonal average plots, Improving UT chemistry scheme Jun-Jul-Aug 2019 Reaction rates under UT conditions have been updated based on aircraft observations of reactive nitrogen: +21.0% • Speed up NO + $O_3 \rightarrow NO_2 + O_2$ NO<sub>2</sub> Slow down $NO_2 + OH \rightarrow HNO_3$ Slow down $NO_2 + HO_2 \rightarrow HNO_4$ Add sinks (OH oxidation and photolysis) for peroxypropionyl nitrate (PPN) -7.0% NO Key results: UT PPN decreased by 64 % (hPa) • UT O<sub>3</sub> unchanged UT NO<sub>2</sub> increased by 21 %, slight improvement in GEOS-Chem vs TROPOMI -1.0% comparison slope (0.6 $\rightarrow$ 0.7). - 1:1 line (pptv) --- Fit, $y=(0.7\pm0.0x + -18.6\pm0.5)$ o R=0.57 Jun-Jul-Aug 2019 +9.0% PAN GEOS-Chem 50 -64.0% 150 50 TROPOMI cloud sliced NO<sub>2</sub> (pptv) Latitude (°)

**Summary** GEOS-Chem underestimates TROPOMI cloud sliced UT NO<sub>2</sub> by ~40 %. We account for 21% of this by updating the chemistry scheme, especially for the reservoir compound PPN. Large discrepancies between TROPOMI and GEOS-Chem coincide with high lightning radiance, which is inversely related to flash rate. We are currently experimenting with LNO, parameterization by radiance rather than flash rate.

#### Improving lightning NO<sub>x</sub> (LNO<sub>x</sub>) scheme Suggested update to per-flash LNO<sub>x</sub> parameterization based on the latest literature: N. America Eurasia 150 ± 95 342 ± 198 Current July Lightning NO emissions mol/fl mol/fl (x10<sup>-11</sup> kg m<sup>-2</sup> s<sup>-1</sup>) in GEOS-Chem 360 ± 200 155 ± 95 200 ± 110 mol/fl mol/fl mol/fl 2.5 3.0 3.5 4.0 2.0 ex midlat. land value in Buscela (2019) Inverse relationship $(y \approx 1 \times 10^{-6}/x)$ between lightning flash rate (x) and flash radiance (y)This may allow us to parameterize the amount of LNO<sub>x</sub> based on the radiance not Lightning Image Sensor on International Space Station purely the flash rate density Lightning radiance (J m<sup>-2</sup> sr<sup>-1</sup> µm<sup>-1</sup> flash<sup>-1</sup>) Flash rate density (10<sup>-5</sup> flashes km<sup>-2</sup> s<sup>-1</sup>) (pptv) 150 — 1:1 line Fit, $y=(1.0\pm0.0x + -14.8\pm0.5)$ o R=0.18 O N -Chem 100 Adjusted July Lightning NO emissions (x10<sup>-11</sup> kg m<sup>-2</sup> s<sup>-1</sup>) in GEOS-Chem 2.5 3.0 3.5 2.0 In this experiment we target the underestimation of LNO, over the $^{ m O}$ oceans by parameterizing LNO<sub>x</sub> over the oceans using radiance 100 150 rather than flash rate. This still leaves significant remote continental TROPOMI cloud sliced NO<sub>2</sub> (pptv) areas of UT NO<sub>2</sub> underestimated in GEOS-Chem.