# Notes on the Larson et al. (2016) shoreline change model

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# Larson et al. (2016)

The Larson et al. (2016) model was developed to simulate cross-shore transport and the resulting profile response at **decadal** time scales. In contrast with other models described before (e.g. Vitousek et al. (2017), Yates et al. (2009)), the Larson et al. (2016) model explicitly accounts for morphological changes in diverse beach profile features (see Fig. 1 and Table 1). Since the Yates et al. (2009) model neglects geological feedbacks on shoreline evolution, the simulation of morphological features dynamics may be crucial to overcome current limitations on long-time scale shoreline change predictions.

The Larson et al. (2016) model employs equations of both sediment transport and sediment volume conservation based on physics and empirical observations. The modules of the model describes: 1) cliff/dune erosion and overwash, 2) bar-berm material exchange, and 3) wind-blown sand transport. The application of the Larson et al. (2016) model to diverse study sites is described in a companion paper (Palalane et al., 2016). Note: Although not explicitly addressed, the Larson et al. (2016) model includes an approach to simulate the effects of sea level rise.

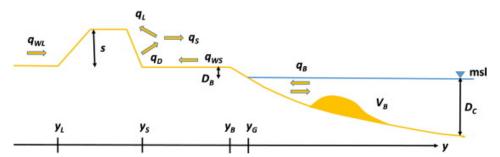


Figure 1: Sketch to calculate cross-shore transport and beach profile evolution (Larson et al., 2016).

#### Transport equations

• Dune erosion and overwash

$$q_D = 4C_s \frac{(R - Z_D)^s}{T} \tag{1}$$

The cross-shore transport rate  $q_D$  is related to the runup height R, the distance from the mean water level to the dune foot  $Z_D$ , the wave period T and an empirical coefficient  $C_s$ .

Where

$$R = 0.158\sqrt{H_0 L_0} \tag{2}$$

The backwash  $q_s$  and overwash  $q_L$  transports are derived from the cross-shore transport rate  $q_D$ :

$$q_s = \frac{q_D}{(1+\alpha)} \tag{3}$$

Parameter	Description	Units
	Sediment transport	
$q_{wL}$	Constant wind-blown transport on the shoreward side	$m^3 s^{-1} m^{-1}$
$q_L$	Overwash transport	$m^3 s^{-1} m^{-1}$
$q_D$	Cross-shore transport rate	$m^3 s^{-1} m^{-1}$
$q_s$	Seaward transport resulting from erosion of the dune (backwash)	$m^3 s^{-1} m^{-1}$
$q_{ws}$	Constant wind-blown transport towards the dune	$m^3 s^{-1} m^{-1}$
$q_B$	Berm-bar transport	$m^3 s^{-1} m^{-1}$
	Morphological locations	
$y_L$	Location of landward end of the dune/barrier	m
$y_s$	Location of seaward end of the dune/barrier (dune foot)	m
$y_B$	Location of berm crest (represents the shoreline position)	m
$y_G$	Location of still-water shoreline	m
$D_C$	Depth of closure	m
	Morphological features	
S	Dune height	m
$D_B$	Berm crest height	m
$V_B$	Bar volume	$m^{3}m^{-1}$

Table 1: List of symbols on Figure 1.

$$q_L = \frac{\alpha q_D}{(1+\alpha)} \tag{4}$$

where

$$\alpha = \frac{1}{A} \left( \frac{R - Z_D}{s} - 1 \right) \tag{5}$$

## • Bar-berm material exchange

$$\frac{dV_B}{dt} = \lambda(V_{BE} - V_B) \tag{6}$$

The change in the bar volume  $V_B$  is taken to be proportional to the deviation from its equilibrium value  $V_{BE}$ . If  $V_{BE}$  and  $\lambda$  are constants, we obtain the solution:

$$V_B = V_{BE} - (V_{BE} - V_{B0})e^{-\lambda t} (7)$$

where  $V_{B0}$  is the bar volume at t=0. The bar volume at equilibrium  $V_{BE}$  is given by:

$$\frac{V_{BE}}{L_0^2} = C_B \left(\frac{H_0}{wT}\right)^{\frac{4}{3}} \frac{H_0}{L_0} \tag{8}$$

where w is the sediment fall speed and  $C_B$  a dimensionless coefficient. The parameter  $\lambda$  is the rate coefficient that depends on the wave and sediment properties:

$$\lambda = \lambda_0 \left(\frac{H_0}{wT}\right)^m \tag{9}$$

where  $\lambda_0$  and m are coefficients with values to be calibrated against data.

#### • Wind-blown sand

In all the study sites where Larson et al. (2016) model was applied (Palalane et al., 2016), constant values of wind-blown transport towards the dune on the seaward  $(q_{WS})$  and landward  $(q_{WL})$  sides were addopted.

### Sediment conservation equations for morphological features

• Bar evolution

$$\frac{dV_B}{dt} = q_B \tag{10}$$

• Berm evolution

$$\frac{dy_B}{dt} = \frac{1}{D_B + D_C} \left( -q_{ws} - q_B + q_s - \frac{dQ_L}{dx} \right) \tag{11}$$

where  $Q_L$  is the longshore sand transport.

• Dune evolution

$$\frac{dy_s}{dt} = \frac{-q_D + q_{WS}}{s} \tag{12}$$

$$\frac{dy_L}{dt} = \frac{-q_L - q_{WL}}{s} \tag{13}$$

where is  $y_G$ ?

If the beach profile has a fixed shape from the berm crest to the depth of closure, the following relationship holds:

$$y_G = y_B + \frac{D_B}{\tan \beta_f} = y_B + L_B \tag{14}$$

where  $L_B$  is the horizontal distance from the berm crest to the shoreline (the foreshore length), and  $tan\beta_f$  is the foreshore slope.

# References

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