

The goal of this problem set is to practice applying conservation laws and force budget techniques to find useful information about a glacier. Make sure to describe and motivate any necessary physical assumptions in your solution writeup.

Problem 1. [7 pts] In the previous problem set, you produced a conservation equation for an idealized glacier in equilibrium with its climate:

$$\int_0^L \alpha(h - E) dx = 0, \quad (1)$$

where L is the glacier length, h its surface elevation, E its equilibrium line altitude, and x the along-flow direction. The surface elevation $h(x) = b_0 - sx + H$, where b is the bed elevation at $x = 0$, s is a constant linear slope, and H is the ice thickness.

- (a) Using Equation 1 and the approach we demonstrated in class, show the steps to find an approximation of glacier length change resulting from equilibrium line elevation change, $\frac{\partial L}{\partial E}$. You can assume perfectly plastic deformation, as we did in class, so the driving stress τ_{dx} must be balanced by a constant yield stress τ_0 .
- (b) Now, consider the Alpine glacier Vernagtferner, shown in profile in Figure 1. Estimate the length change that would result from a change in equilibrium line altitude for this glacier.
- (c) Download two datasets from the internet:
 - Davaze & Rabatel dataset “Annual glacier equilibrium-line altitude (ELA) for 239 glaciers located in the European Alps quantified from optical remote sensing data from 2000 to 2016” [[link](#), also on Canvas]. Extract the ELA history for Vernagtferner (you may do this manually or by coding; let me know which you do);
 - Vernagtferner frontal variation, “FoG.FVobs_489.csv” from the World Glacier Monitoring Service [[link](#)].

Compute the true glacier length change observed per change in equilibrium line altitude during the 2000-2016 period.

- (d) Is the equation we derived in class a suitable approximation for Vernagtferner? Why or why not?

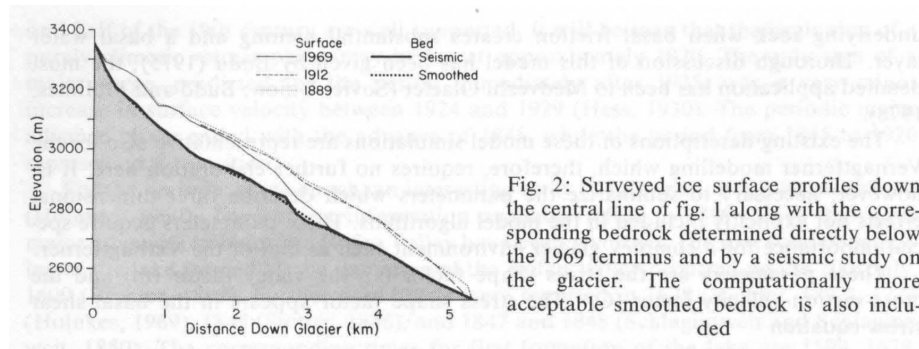


Figure 1: Observed profile of Vernagtferner surface and bed elevation.¹

Problem 2. [5 pts] Consider a laterally symmetric ice stream with a thickness of $H = 1100$ m thick, half-width of $W = 15$ km, a density of $\rho = 910 \text{ kg m}^{-3}$, and an along-flow slope of $\frac{\partial h}{\partial x} = 1.3 \times 10^{-3}$. The shear stress has been measured at one glacier margin to be 1.7×10^5 Pa and is assumed to be constant with ice thickness. Assume that the ice stream is rectangular in shape and assign a coordinate system with x pointing along flow.

- Compute the average along-flow gravitational driving stress of this ice stream.
- Compute the width-averaged lateral drag from the margins (van der Veen Section 4.4).
- What fraction of the driving stress is balanced by lateral drag in this case? What contribution to the force balance do you expect from basal drag?

Problem 3. [3 pts] In general, solving systems of PDEs is hard. We therefore make use of simplifying approximations wherever possible. For each of the following, outline the assumptions that must be satisfied and provide an example of a glaciological setting where the approximation **is not** suitable.

- Shallow Ice Approximation controlled by basal drag
- Shallow Ice Approximation controlled by lateral drag
- Shallow Shelf Approximation

¹P.D. Kruss and I.N. Smith (1982) "Numerical modelling of the Vernagtferner and its fluctuations", *Zeitschr. Gletscherkunde & Glazialgeologie* 18(1):93-106.