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FOURIER IMAGE SYNTHESIS AND SLOPE SPECTRUM ANALYSIS OF  
DEEPWATER, WIND-WAVE SCENES VIEWED AT BREWSTER'S ANGLE

BY

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(1978)

A thesis submitted in partial fulfillment  
of the requirements for the degree of  
Master of Science in the Center for  
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DEEPWATER, WIND-WAVE SCENES VIEWED AT BREWSTER'S ANGLE

BY

JAN A. NORTH  
CAPTAIN USAF

Submitted to the  
Center for Imaging Science  
in partial fulfillment of the requirements  
for the Master of Science Degree  
at the Rochester Institute of Technology

ABSTRACT

A semi-empirical model for the Fourier synthesis of deepwater, wind-wave scenes has been constructed for the analysis of water-wave slope spectra. The main simplifying assumptions of this model are 1) fully-developed wind-wave surfaces are quasi-homogeneous, quasi-stationary and are therefore treatable by Fourier methods, 2) the subsurface is both optically and mechanically deep, and 3) the small range of spectral wave components defines a fetch-limited, small-amplitude condition. A nonlinear transformation of wave slope to reflected and refracted radiance in both horizontal and vertical polarizations was effected under the special conditions of Brewster-angle viewing under clear skies at a spectral wavelength of 460 nanometers. Seventy two syntheses were varied with respect to six distinct solar positions, four distinct wind directions, and three distinct wind velocities.

The synthetic wave scenes were analyzed via the forward Fourier transformation and their radiance magnitude spectra were compared with the original slope magnitude spectra of the initial synthesis in order to estimate the effects of the nonlinear radiance transformation on the recovery of the wave slope spectrum from imagery. Within the boundaries of this study, it was determined that 1) the limited results of Chapman and Irani [1981] have been generally verified, 2) the existence of an optimal imaging geometry for slope spectrum estimation is indicated, and 3) the presence of sub-resolution wave slopes creates a significant effect on wave slope spectra derived from imagery.

Theses. (hd)

**DEDICATION**

**To Kathy, Hollie and Janson**

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## 1.0 INTRODUCTION

The intent of this study is twofold:

- 1) The first goal is to construct a reasonable linear model to describe the interaction of wind-driven water waves with the two primary sources of natural illumination: the down-welling sun and sky radiance (of which a fraction is reflected in the direction of an overhead observer) and the up-welling back-scattered subsurface radiance (of which a fraction is refracted in the direction of an overhead observer). This linear model will be used to create synthetic scenes of wind-driven water waves and derive their radiance magnitude spectra.
- 2) The second goal is to analyze the spatial radiance distribution of the synthetic wave scenes with respect to sun position, wind direction, wind velocity, polarization, and sub-resolution slope variance in order to compare the magnitudes of error that arise from the treatment of the radiance image magnitude spectrum as a linearizable reconstruction of the slope magnitude spectrum of the imaged water surface.

This study is essentially an amplification and extension of

the work of Chapman and Irani [1981].

In the 1950s, the two unrelated fields of oceanography and photographic science both experienced a minor renaissance with the introduction of the concepts of stochastic processes and power spectral analysis that were first applied in communications engineering [Blackman & Tukey, 1958]. Power spectral analysis provided a powerful analytical tool to photographic science for evaluating image quality; likewise, power spectral analysis provided a linear, statistical mechanism to oceanography for evaluating the large number of deterministic theories concerning the development, propagation, and decay of water waves [Kinsman, 1965]. An excerpt of a letter from the oceanographer, Dr. Walter Munk, to Drs. Blackman and Tukey gives an indication of beneficial results that were obtained with the new methods:

"...we were able to discover in the general wave record a very weak low-frequency peak which would surely have escaped our attention without spectral analysis. This peak, it turns out, is almost certainly due to a swell from the Indian Ocean, 10,000 miles distant. Physical dimensions are: 1 mm high, a kilometer long." [Blackman & Tukey, 1958]

Oceanographers had previously developed a number of photographic techniques to capture and optically analyze large amounts of high-resolution, two-dimensional (2D) spatial

information about water-wave surface structure [Hulbert, 1934; Barber, 1949, 1954; Cox & Munk, 1954a,b; Schooley, 1954; Pierson, 1960]. It was not until the advent of the digital computer and the implementation of the digital Fast Fourier Transform (FFT) that oceanography was able to efficiently analyze the wave scene into its spectral components and derive the directional 2D spectrum for elevation and slope [Kinsman, 1965; Pierson & Stacy, 1973]. The current study is a continuation of existing research which analyzes digital images of water waves through digital simulations.

The current study builds upon the work of Chapman and Irani [1981]. They, in turn, extended the work of Stilwell [1969; Stilwell & Pilon, 1974] and Kasevich [1975; Kasevich et al., 1971, 1972] in an attempt to quantify the parametric nonlinearities that may exist between the slope magnitude spectrum of the water surface and the radiance magnitude spectrum of the water surface image - as a linearizable reconstruction of the slope magnitude spectrum of the actual water surface. The current study reconstructs the synthetic model of Chapman and Irani with some selected modifications and enhancements:

- 1) no sensor or path radiance model is introduced as the

intent of this study is to estimate only the effects of wave and illumination geometry on the nonlinear transformation of slope to radiance,

2) a correlation model is introduced to provide a conservative spatial filter for synthesizing the effects of sub-resolution wave-slope variance on the image spatial radiance distribution,

3) a subsurface upwelling refracted radiance model is introduced to provide an additional nonlinear element to the analysis and to more adequately model the slope spectra obtained from visible images,

4) the zenith angle of observation is set at the Brewster angle (relative to mean sea surface) in order to define the extent to which the nonlinear transformation is affected by vertically polarized radiance added by the variance of wave slopes, and

5) the analysis of normalized difference spectra is introduced in order to provide a more direct correspondence with the analysis of slope variance as calculated from integrated power spectra.

The remainder of the Introduction is organized into four sub-chapters:

Chapter 1.1 introduces the topic of wind-generated waves, with emphasis on the small nonlinearities that must be considered.

Chapter 1.2 is an overview of the development of realistic syntheses of water-wave scenes, with emphasis on synthetic radiance imagery.

Chapter 1.3 is an overview of the development of image analysis of water-wave scenes, with emphasis on wave-slope spectrum analysis.

Chapter 1.4 is a review of the pertinent research which precedes the current study, with emphasis on the development of the first-order theory. Extracts from the introductory analysis of Kasevich [1975] and Stilwell [1969] are presented, followed by an overview of both the simulation experiment of Chapman and Irani [1981] and the experimental design of the current study.

### 1.1 Introduction to Wind-Driven Water Waves

Fully three quarters of the planet is covered with a layer of water; except for the rare instances of sustained windlessness where the surface is smooth and specular, the water surface is disturbed by the addition of directional friction energy applied by wind passing over the surface. (Tides, Coriolis-force waves, earthquake-generated waves, and other very-long-period waves are not considered here.) Up to some maximum wind velocity, the disturbed surface remains analytic (infinitely differentiable) and can be described as a quasi-stationary, pseudo-Gaussian process [Kinsman,1965] with an approximate power elevation spectrum.

Figure 1.1:1 illustrates Dr. Kinsman's relative estimation of the power elevation spectrum describing the energy contained in the surface waves of the oceans. Note that energy,  $L^2$ , is proportional to  $|\text{elevation}|^2$ .

The nonlinearities of real water-wave surfaces can be attributed to four dominant factors [Kinsman,1965; LeMehaute,1976]:

First, wind energy is added to the water in a directional

and time-variant manner. Cox and Munk [1954a,b] found the wave-slope distribution to be nearly Gaussian, as would be expected for a continuous wideband spectrum. However, the cross-wind distribution of slopes is symmetric and the along-wind distribution is slightly skewed so that the most probable wave slope is directed upwind, a result they attributed to directional wind stress.

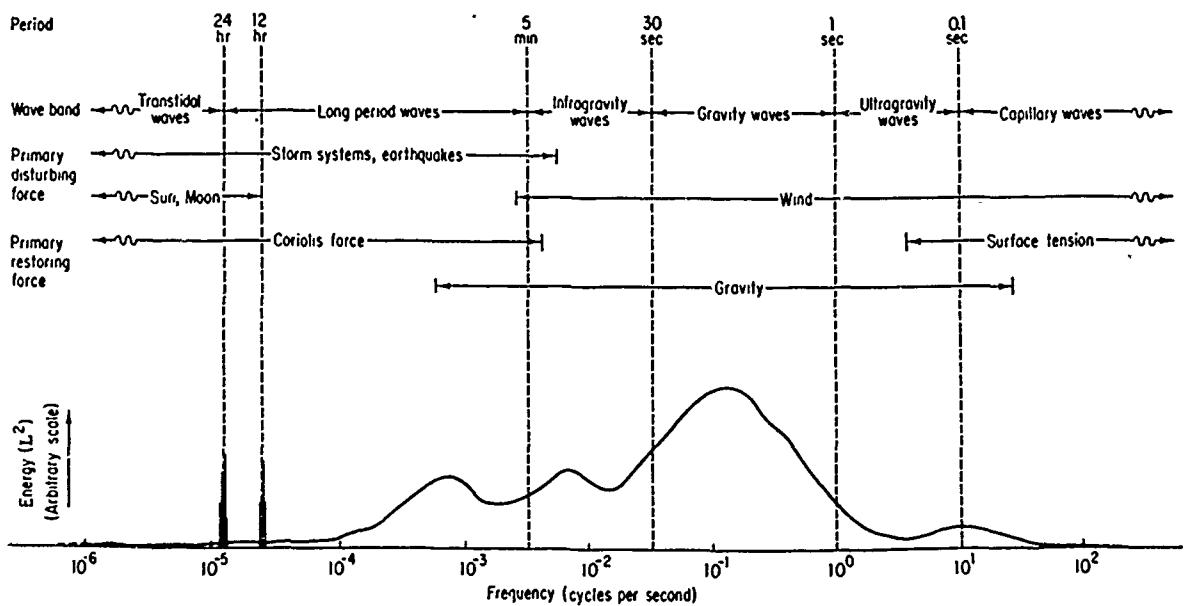


FIGURE 1.1:1. Schematic representation of the energy contained in the surface waves of the oceans - in fact, a guess at the power spectrum. (from Kinsman, 1965)

Second, both gravitational and molecular forces act on material water waves to inhibit random generation of very-large-elevation waves and, in the derivative, random generation of very steep wave slopes. Also, the natural

propagation state of a high-amplitude, narrowband water wave is cycloidal rather than sinusoidal, i.e. with elongated flat troughs and shorter peaked crests. This cycloidal form diminishes as the sea state increases and spectral wave-component energy is transferred over a wider band of wave-numbers, but it does not disappear. Therefore, the wave surface is pseudo-Gaussian in that the elevation and slope distributions exhibit a slight kurtosis (peakedness) and a truncation of the large-variance components. Cox and Munk correlated their results with a Gram-Charlier representation of a Gaussian distribution to account for this slight skewness and kurtosis.

Figure 1.1:2 illustrates the cross-wind and along-wind distributions of wave slopes (from Cox & Munk, 1954b).

Note that the curves are normalized to equal areas.

Third, there is a nonlinear transfer of wave energy from the high-frequency components of the wave spectrum to the low-frequency components as a function of time, fetch, and sea state. As previously mentioned, there is some maximum wind velocity beyond which the water surface acquires excessive energy, which leads to instability and a consequent loss of analytic structure. A familiar manifestation is white-

capping of the waves. Up to this point of fully-developed sea state, the effect of wind energy is to generate small-amplitude capillary waves which tend to become damped by surface tension and which impart energy to the lower-frequency waves upon which they are riding. In 1847, Stokes [1880] defined the maximum wave slope which each wave component could attain before it became unstable; each spectral component transferred energy to lower frequencies until saturation; if the lower regime was saturated, whitecapping resulted. Within the range of gravity waves, the maximum wave slope is approximately 1/7. For capillary waves, the maximum wave slope can approach 1/1. Chapter 2.2 will illustrate the fact that the greater part of the slope variance from a fully-developed wind-wave surface is due to capillary waves.

Fourth, water waves refract in the presence of subsurface obstacles just as light refract in the presence of media with differing refractive indices. Refractive nonlinearities appear only when the ratio of wavelength to depth becomes large. As long as the deepwater condition is satisfied (i.e. the longest wavelength under consideration is small relative to depth) the wave surface remains linear with respect to subsurface refraction.

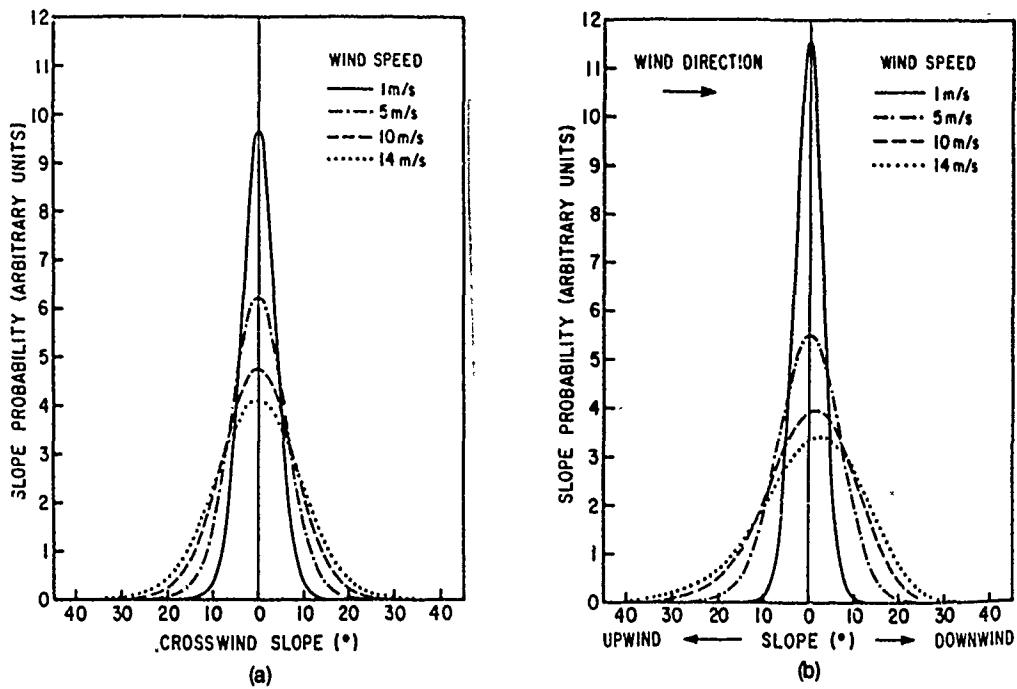


FIGURE 1.1:2. Facet slope-angle ( $\beta$ ) distribution probability for a wind-roughened sea, computed for various wind speeds by means of the Cox & Munk probability function. (from Sidran, 1981)

In sum, a linear model for a synthetic water surface would most accurately describe infinite-fetch, fully developed wind waves propagating over deep water and with steady homogeneous winds blowing in a constant direction at velocities well below the point of spectral saturation. The predominant discrepancy of a linear model is that the lack of directional skew allows a 180-degree ambiguity in the direction of wind propagation.

## 1.2 Image Synthesis of Water-Wave Scenes

### An Overview

The subject of image synthesis is best introduced by Upson:

"Within computer graphics and animation, two major sub-disciplines have developed: graphics for entertainment or artistic purposes, and graphics for scientific or technical purposes. Computer scientists working on entertainment applications have emphasized the correct visual representation of natural phenomena and have developed ad hoc physical simulations to accomplish this. Computational scientists working in the physical sciences have devoted their efforts to the underlying physics for simulating phenomena, with little emphasis on the visual representation.

In recent years these two approaches have begun to reach the limits of their ability to work without each other. The realism required in entertainment animation is beyond that obtainable without physically realistic models. Similarly, numerical simulations in the physical sciences are complex to the point of being incomprehensible without visual representations." [Upson, 1987]

Progress in the image synthesis of natural phenomena is occurring within three overlapping lines of research:

- 1) geometric models which provide a more accurate model of natural texture and shape,
- 2) illumination models which provide a more accurate model of the physical interaction between light and matter, and
- 3) temporal models which provide a more accurate model of the equations of motion for natural systems.

Most research devoted to the realistic synthesis of water-wave scenes has occurred within the last decade as part of the general interest in rendering of complex natural phenomena by computer. Early artistic renderings concentrated on generating low oblique scenes with realistic surface texture but with little attention to correct illumination effects. Early scientific renderings concentrated on the generation of near-nadir scenes with correct radiometry in order to simulate remotely sensed oceanographic imagery having low spatial resolution.

Whitted [1980] created one of the earliest attempts at rendering water waves in his film "The Compleat Angler". The ripples on a flat pool surface were generated by 'bump mapping', where the surface normal was perturbed by a single sinusoidal function. Bump mapping was first introduced by Blinn [1978]. Information International used a similar technique with cycloidal waveforms to create a leader for Pyramid Films [1981]. Perlin [1985] amplified this technique by using 20 superimposed cycloidal waveforms, each radiating from randomly selected centers. The characteristic limitation of bump mapping is that the actual surface is flat; there is no change in surface height to correspond to the perturbation of a local normal vector.

Schacter [1980] developed a scalar texture model for realtime image generation of waves, using sums of three or fewer long-crested, narrowband noise waveforms. This stochastic geometric model was intended for realtime image synthesis of a variety of textured surfaces or random fields with no regard to the physics of surface-light interaction. Norton et al.[1982] employed frequency-limited (or 'clamped') analytical functions for the same purpose.

Max [1981] used a vectorized ray-tracing algorithm on a Cray-1 supercomputer to model optical reflection on water waves for use in a color animation film "Carla's Island". The wave-surface height field was modeled by superimposed traveling sine waves. Each ray was constrained to reflect off the water at most two times before it proceeded to other scene elements. However, no effort was made to accurately model the attenuated spectral radiance of the synthetic scene; a geometric rule-of-thumb approach was used to compute each pixel based on a weighted average of 'sun color' and 'sky color' that were input from an RGB (Red-Green-Blue) translation table.

In 1985, Ts'o and Barsky [1987] applied the Stokes

model of the sea surface and generated height fields by fitting beta-splines with parameters that yielded symmetric waves. This model simulated directional refraction of water waves due to interaction with the ocean floor. Peachey [1986] similarly implemented the Stokes model as a height field but used quadric surfaces to create asymmetry. This synthesis also used a particle-system model to simulate the effects of surf spray on a beach but without the effect of directional refraction. Peachey gave some thought to surface reflection of skylight but ultimately assumed a Lambertian approximation for rough surfaces in the calculation of pixel brightness. Fournier and Reeves [1986] reviewed the work of the previous authors and produced a Gerstner wave model that incorporated a wind field and a stochastic element to perturb the wave model in addition to the previously noted features. Even in this final product, reflection is simulated by a one-level ray trace with selective filtering of the RGB color 'spectrum'.

Fractal and sequential Markov-chain synthesis techniques have been developed to simulate the texture of many natural surfaces including ocean waves [Mandelbrot, 1977, 1988; Gagalowicz & Ma, 1985; Monne et al., 1981]. In general, these techniques are considered to produce poor synthetic

water-wave images. The main objection to the surfaces generated by these processes is that they are not continuously differentiable. The surface normal vector at a point can only be approximated via a neighborhood average. However, it is interesting to note (in the context of wave surfaces) that the mathematician Karl Weierstrass developed the first fractal function by describing an infinite series of superimposed harmonic sinusoids:

$$f(t) = (1)\cos(w*t)+(1/a)\cos(b*w*t)+(1/a^2)\cos(b^2*w*t)+\dots ,$$

to yield a series representation of a continuous yet non-differentiable function. In this particular case, the fractal dimension  $d = \ln(a)/\ln(b)$ .

Mastin et al.[1987] recognized the inadequacy of the fractal model for ocean scenes and instead treated water surfaces as filtered white-noise processes. For his filter, Mastin used a modified Pierson-Moskowitz, temporal-domain, elevation power spectrum for fully-developed, wind-driven seas. A white-noise surface was filtered in the frequency domain and then inverse-Fourier-transformed to create the synthesized wave image. The surfaces were animated by varying the phase of the frequency components between each

frame, based on the dispersion relation between phase velocity and wave frequency. Mastin employed a simple illumination model to render the scenes but indicated that "incorporation of Fresnel's law into the reflectance model would probably enhance image quality."

Borrego and Machado [1985] used essentially the same filter technique as Mastin but with a Pierson-Neumann spectrum. They wrote their images to photographic film and then analyzed the film via the optical Fourier transformation and optical cross-correlation in order to compare synthetic results with real photographs of the sea. They did not use any illumination model - density values on film corresponded to surface height - since their analysis was only a first-order approximation.

Attempts to combine a water-wave surface model with correct radiometry are fewer in number:

Chapman and Irani [1981] attempted to quantify the error magnitudes associated with linearizing the relationship between the spatial frequency spectrum of an image and the slope spectrum of the actual water surface. Their radiometric model, though incomplete, is the most comprehen-

sive found within this survey. It will be described in detail in the Methods section as their combined model provides the point of departure for the current study.

Wilf and Manor [1984] applied the Chapman and Irani model to a restricted case: the simulation of water-wave images in the far infrared. They simplified their model with the assumption of a Forward Looking Infra-Red (FLIR) sensor having limited spatial resolution and sensitivity. Schwartz and von [unpublished] and Masuda et al. [1988] also limited their models to the far infrared with the intent of simulating the angular effect of emissivity on sea temperature variation due to wind-roughened water.

It is apparent from this overview that the field of image synthesis, with emphasis on rendering of water waves, is less than a decade old. The integration of radiometry within these syntheses has occurred infrequently and for a limited objective in each instance.

### 1.3 Image Analysis of Water-Wave Scenes

#### An Overview

In 1925, Schumacher made oblique stereo-pairs from a ship with the intent of measuring the variability of wave heights. The utility of this method was severely limited due to many factors: 1) the camera baseline was restricted to the length of the ship; 2) waves in the foreground obstructed waves in the background; 3) backsides of waves were not visible; and 4) there was a lack of 'ground' control on open seas for height determination - the errors are especially pronounced from an oblique perspective [Pos, 1988].

In 1933, Hulbert [1934] made polarized and unpolarized oblique photographs of sun glitter on sea waves to measure the polarization of light at sea with respect to surface roughness, sun angle, and weather conditions. Because the widths of glitter patterns correlate to the maximum slope of the sea surface, Hulbert was able to demonstrate that waves in the North Atlantic varied from 15 degrees inclination when winds were blowing at 3 knots up to 25 degrees inclination at 18 knots.

Sawyer [1949] mounted a Sonne strip camera on a fast

low-flying airplane to photograph narrow strips of the sea surface and measure the directional spectrum of the waves. The field of view was too narrow to capture significant amounts of surface data orthogonal to the flightline. The accuracy of the spectral estimate decreased with the angle from the flightline.

Also in 1949, Barber [1949,1954] analyzed single photographs of sea surfaces to determine wave direction but he was unable to determine the two-dimensional (2D) spectrum due to the computing limitations of the time. At this point, it became apparent that an essential requirement for future analysis of water-wave surfaces was to have near-vertical, high-resolution, high-contrast images covering large areas. The intent was to photographically capture significant information about the largest range of spatial frequency components without perspective distortion or hidden surface detail. It also became apparent that the analysis of large amounts of spatial information required the power of a digital computer.

In 1951, Cox and Munk [1954a,b] measured the wave-slope distributions of the sea surface from aerial photographs of sun-glitter patterns. They computed the distribution from

the measured variation of radiance within a glitter pattern instead of computing maxima from the pattern boundaries as done by Hulbert. Four cameras were flown from a single airplane at altitudes of 2000 feet, with two used as imagers and two used as radiometers. Their image analysis was quite sophisticated; it accounted for sun diameter, angular reflectivity, lens falloff, film sensitivity, exposure calibration, and ultimately provided a first-order relationship between film density and directional wave-slope probability.

In 1953, Schooley [1954] performed a simplified version of the Cox and Munk experiment by taking flash photographs of a river surface from a 45-foot bridge elevation at night. The main limitation of his experiment was the probable inhomogeneity of the water surface due to limited fetch and the presence of surrounding obstacles.

In 1954, medium-altitude (3000 ft) stereophotography was employed by Marks and Ronne [1955] to generate stereo-pairs of sea surfaces. Two airplanes carried radio-synchronized cameras and a surface ship acted as 'ground' control in the photographs. Elevations were photogrammetrically measured at discrete points and the sampled elevation array was then auto-correlated (the sampling distance determined

the desired spectral resolution). This experiment marks the first recorded use of a digital computer to calculate the directional 2D spectra of water waves. The work of Cote et al. [1960] enhanced this basic technique. More recent stereophotogrammetric efforts include Holthuijsen [1983a,b] and Pos et al. [1988]. Elements of this later work include methods to render the water opaque so that a more exact calculation of the height field can be made.

During the 1950s and early 1960s, Longuet-Higgins [1952-1962; Cartwright & Longuet-Higgins, 1956] elaborated on the results of Cox and Munk to formulate the statistical theory of patterns, paths, number, frequency, and distributions of specular reflection points on randomly moving surfaces. Stilwell [1969; Stilwell & Pilon, 1974] correlated the statistics of sea-surface images to the wave-slope statistics of the actual sea surface. Under the assumptions of uniform sky radiance, optimized viewing geometry, and small surface slopes, Stilwell derived the relationship between the film transmittance of an imaged surface point and the range component of the wave slope at that surface point. He further demonstrated a linearizable relationship between the spatial image spectrum and the imaged surface slope spectrum. Kasevich [1975] extended Stilwell's model

to second order to develop an optimization criterion for the relationship. Chapman and Irani [1981] took this work one step further by applying a synthetic model and executing a limited quantification of the error magnitudes associated with the parametric dependence of this linear model. This work will be further detailed in Chapter 1.4.

Sheres [1980] developed a novel technique for remotely sensing surface-flow velocities based on imagery of monochromatic wavetrains of known frequency (such as those generated by a motor boat) propagating over the region of interest. His work demonstrated that the wavelength and direction of two different wavetrains generated all the required information to calculate surface flows. Gotwols and Irani [1980] developed a similar technique to determine the phase velocity of short gravity waves.

Exotic sensors using LASER [Palm et al., 1977; Schau, 1978; Abshire & McGarry, 1987] and LIDAR [Weinman, 1988] have been used to extract directional spectra and surface backscatter data at higher wavenumbers (i.e. the capillary wave regime). Synthetic Aperture Radar (SAR) imagery has been used to estimate spectra, phase velocities, and propagation directions at lower wavenumbers (i.e. the gravity wave re-

gime) [Monaldo & Lyzenga, 1986; Monaldo & Kasevich, 1982; Carlson, 1984]. Also, Long Wave Infrared (LWIR) sensors have been used to calculate spatial spectra of ocean-surface temperature [Saunders, 1967, 1968; McLeish, 1970].

Lybanon [1985] reported on the implementation of an automated image-analysis system by the U.S. Naval Ocean Research & Development Activity (NORDA). The Interactive Digital Satellite Image Processing System, or IDSIPS, can automatically derive the sea-surface slope statistics from sun-glitter images through analysis of the imaging geometry. As a late example of the practical application of water-surface spectra determination, Fisher [1986] analyzed four sun-glitter images taken from the space shuttle Challenger (STS-41G) to locate acoustically important oceanographic features in support of hydro-acoustical sensor placement.

It is apparent from this overview that, in six decades, the methods of water-surface imaging have moved from surface ships to airplanes to satellites and spacecraft. Likewise, the measurement and analysis of the resulting imagery have undergone a corresponding increase in computational power and sophistication.

#### 1.4 Review of Pertinent Literature

A comprehensive review would include the work of Stilwell and Pilon [1974], and Kasevich et al. [1971,1972]. The emphasis of this earlier work is on the analysis of coherent optical processing techniques as applied to photographic emulsions of wave scenes. Kasevich [1975] provides a general introduction to the first-order theory subsequent to the development of his approximate geometric-optics second-order theory to estimate the optimum viewing geometry for the obtainment of reasonable spectra. Stilwell [1969] provides additional development of image analysis with respect to the law of Malus subsequent to performing an optical analysis to derive directional energy spectra. Only the first-order theory is reviewed here; a review of the second-order theory is beyond the scope of this study since the theory of Kasevich assumes simplified approximations for both Fresnel reflectivity and sky radiance distributions. However, the most general results of the second-order theory can be compared, with caution, to the results of this study. Any second-order theory that is developed for this geometric problem loses definiteness because the spatial distributions for natural radiance are independent; no general solution can be specified [Stilwell,1969]. This is the prime motiva-

tion for the simulation and analysis of geometric effects through empirical models.

Review of Kasevich [1975]

The essential requirement for the determination of slope spectra from wave images is to have the spatial modulation of the image be proportional to the wave profile. Kasevich uses the example of a transparency with film exposure, E, defined over its linear region by

$$E(y) = [f_0(y)^{-\gamma/2}] * [1+f(y)/f_0(y)]^{(-\gamma/2)} \quad [1.4:1]$$

where

$$E(y) = f_0(y) + f(y), \quad [1.4:2]$$

and

$f_0(y)$  = the mean exposure on film,

such that

$f_0(y)$  = the mean exposure on film,

and

$f(y)$  = the exposure modulation due to scattering of radiance from specular wave-slope facets.

This example is given for a two-dimensional case.

If  $f_0(y) \gg f(y)$ , then Equation [1.4:1] can be expanded

in a binomial series to yield the approximation

$$E(y) \tilde{=} [f_0(y)^{-\gamma/2}] * [1 - (\gamma/2) * f(y)/f_0(y)]. \quad [1.4:3]$$

The estimation of the slope spectrum from the forward Fourier transform of the image requires that  $f(y)$  be linear with respect to the wave slope  $dz/dy$ , where  $z$  is the surface elevation. This condition can only be approximately satisfied because of the nonlinearity of 1) the spatial radiance distributions found in nature, 2) the Fresnel reflectivity variation with respect to incidence angle, and 3) the refraction of upwelling subsurface radiance in the direction of the observer.

#### Review of Stilwell [1969]

For a small wave-slope angle  $\beta$ , the small-angle approximation is:

$$\beta = \text{ATAN}(dz/dy) \tilde{=} dz/dy, \quad [1.4:4]$$

where  $\beta$  is the fundamental parameter for extracting wave-slope spectra from imagery. The law of Malus defines the radiance observed at azimuth angle  $\theta$  as a simple function of Fresnel reflectivity and incident radiance:

$$L_o(\theta, \beta, \omega) = L(\mu) * R(\omega), \quad [1.4:5]$$

where

$L_o(\theta, \beta, \omega)$  = the observed reflected radiance,

$L(\mu)$  = the incident radiance to be reflected,

$R(\omega)$  = the Fresnel reflectivity (for any arbitrary polarization),

$\theta$  = the zenith angle of observation,

$\beta$  = the slope of the reflecting surface,

$\mu$  = the zenith angle of the incident radiance,

and

$\omega$  = the angle of incidence,

such that

$$\omega = \theta - \beta \quad [1.4:6]$$

and

$$\mu = \omega - \beta. \quad [1.4:7]$$

Figure 1.4:1 illustrates the angular relations in two dimensions.

The variation of observed radiance with respect to a change of surface slope  $d\beta$  at some point is

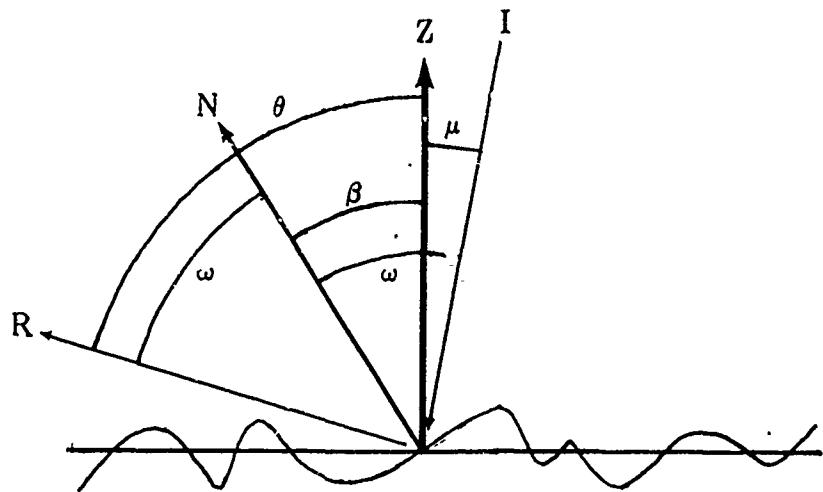


FIGURE 1.4:1. Angular relations for the 2D case.

$$\begin{aligned} \frac{dL(\theta)}{d\beta} &= \frac{dL(\mu)}{d\mu} * R(\omega) * \frac{d\mu}{d\beta} \\ &+ L(\mu) * \frac{dR(\omega)}{d\omega} * \frac{d\omega}{d\beta}. \end{aligned} \quad [1.4:8]$$

If  $\beta$  is a small angle and

$$\omega \approx \mu, \quad [1.4:9]$$

then

$$\begin{aligned} \frac{dL(\theta)}{d\beta} &= [L'(\mu) * R(\omega) \\ &+ L(\mu) * R'(\omega)] * \frac{d\omega}{d\beta}, \end{aligned} \quad [1.4:10]$$

where the prime ('') denotes the first derivative with respect to the argument. With the assumption that the water surface remains analytic, the small-angle linear approxima-

tion holds in the Fourier transform for wave-slope angles ( $\beta$ ) up to at least 30 degrees [Stilwell, 1969].

Review of Chapman and Irani [1981]

The intent of the work of Chapman and Irani was to quantify the parametric dependencies of errors inherent in linearly relating a wave-slope magnitude spectrum to the corresponding radiance image magnitude spectrum. Their approach was to simulate radiance images of sea-like surfaces in two dimensions. Their simulation utilized models of the sea surface, the sky radiance distribution, and the non-linear transfer function that transforms surface slope to radiance. They determined that the two-dimensional simulation was computationally expensive so only a small number of geometries were investigated. However, they initially used a simpler one-dimensional model of a  $\pm 15$ -degree sinusoidal surface propagating in a single direction along the sensor field in order to survey various imaging geometries for potential synthesis and analysis in two dimensions.

Figure 1.4:2 illustrates the two-dimensional methodology for image synthesis and error analysis used by Chapman and Irani. This methodology is described in detail in Chapters 2.1 through 2.8.

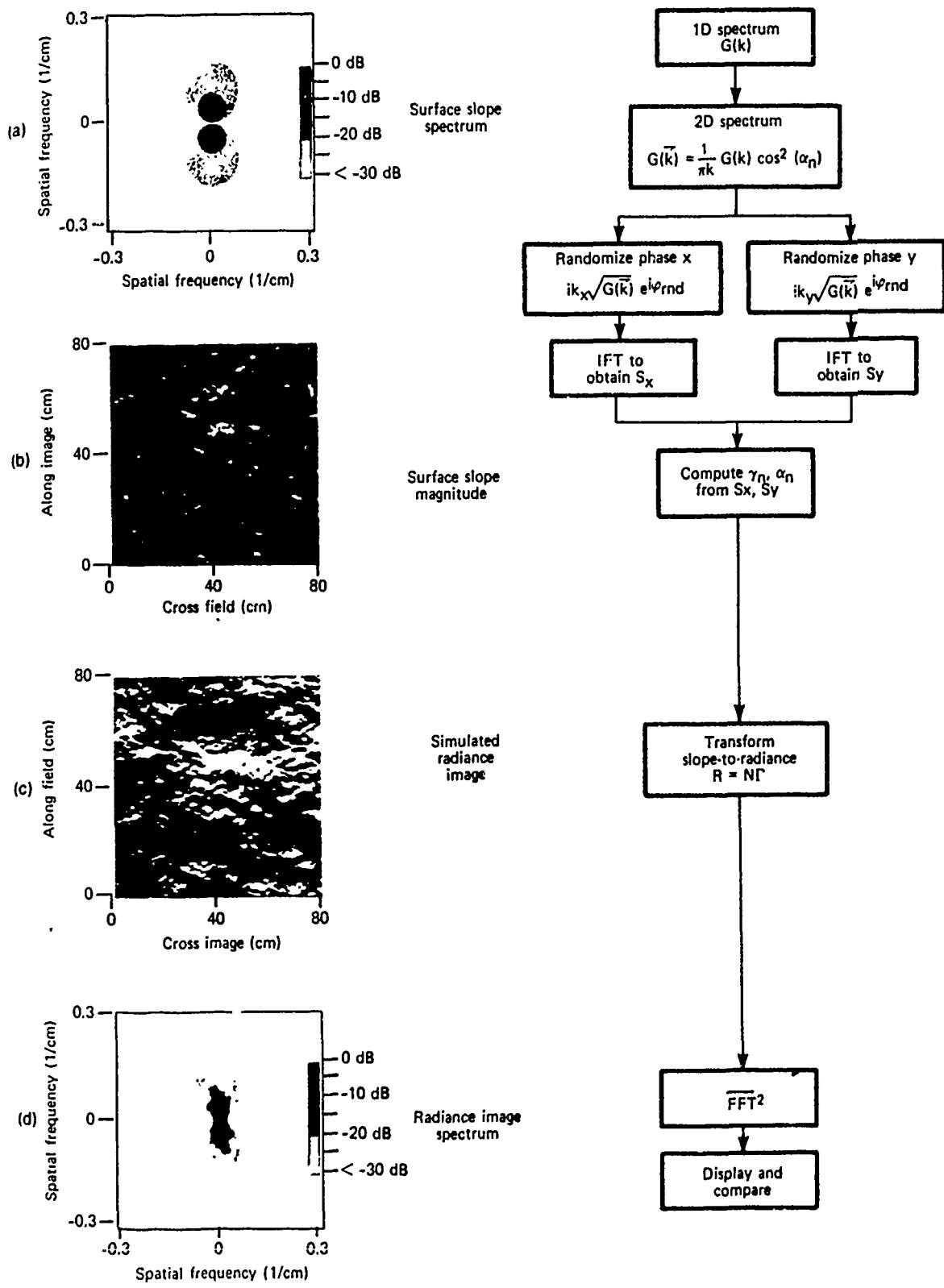


FIGURE 1.4:2. Methodology of error estimation in two dimensions. (from Chapman & Irani, 1981)

Chapman and Irani used two figures of merit derived from signal analysis to describe the nonlinearity of the slope-to-radiance transformation. The first figure of merit is the ratio of the fundamental component to the DC component which they termed the contrast of the wave image. The second figure of merit is the square root of the power in the upper harmonics divided by the power in the fundamental, which they termed the Total Harmonic Distortion (THD) of the wave image. An optimal imaging geometry for measuring slope statistics would maximize contrast and minimize THD, based upon their terminology. These figures of merit were computed for the results of their one-dimensional model. To quote Chapman and Irani:

"...these figures of merit do not represent realistic contrasts or distortion statistics for sea surface imaging because the distribution of slopes over a sinusoid differs substantially from the approximately Gaussian distribution of sea surface slopes." [Chapman & Irani, 1981]

Their initial results indicated that the figures of merit showed a substantial, asymmetric dependence on slope azimuth for certain geometries and may misrepresent the utility of a particular geometry.

Chapman and Irani then applied their two-dimensional synthesis and analysis to six selected geometries with the

intent of evaluating the azimuthal dependence of their figures of merit. Their two-dimensional results indicated that 1) nonlinearity is largely independent of wavenumber, 2) nonlinearity remains relatively low (error < 3 dB) over a relatively wide range of azimuth (100-120 degrees) around the sensor azimuth, and 3) nonlinearities increase within 30-40 degrees of the cross-field direction.

#### Experimental Design of the Current Study

As earlier stated, the current study is essentially an amplification and extension of Chapman and Irani as developed from a proposed statement of work:

"An exhaustive evaluation of our radiance modulation model would involve: 1) simulating the geometry of a surface area; 2) transforming that surface geometry into an image using a particular set of optical conditions; 3) estimating the slope spectrum of the simulated surface and the spectrum of the image; 4) repeating steps 1-3 for all other possible imaging geometries; and 5) comparing each image spectrum with the corresponding slope spectrum to obtain estimates of the error arising from nonlinearity. This, in fact, is impractical because there are a large number of parameters to be varied, and, for any set of these parameters, the error computation involves a large number of operations." [Chapman & Irani, 1981]

The intent of this study is to develop and execute a test design which provides a minimal yet optimal set of parameters to both verify the limited results of Chapman and

Irani and to complete a more detailed parametric surface exploration.

For their two-dimensional analysis, the experimental design of Chapman and Irani was of the order 1x3x2 factors relative to the current study, which is of the order 3x4x6 factors:

Wind Friction Velocity - 12.0, 36.0, and 60.0 cm/sec versus 12.0 cm/sec. These three velocities define the envelope of the linear wind-wave model.

Wind Azimuth - +90:-90, +45:-135, 0:180, and -45:135 degrees (with ambiguity) versus +90:-90, +30:-120, and 0:180 degrees (with ambiguity).

Sun Position - (0,0), (0,53.2), (45,53.2), (90,53.2), (135,53.2), and (180,53.2), where the coordinates are (azimuth angle, zenith angle) in degrees, versus (90,45) and (135,45).

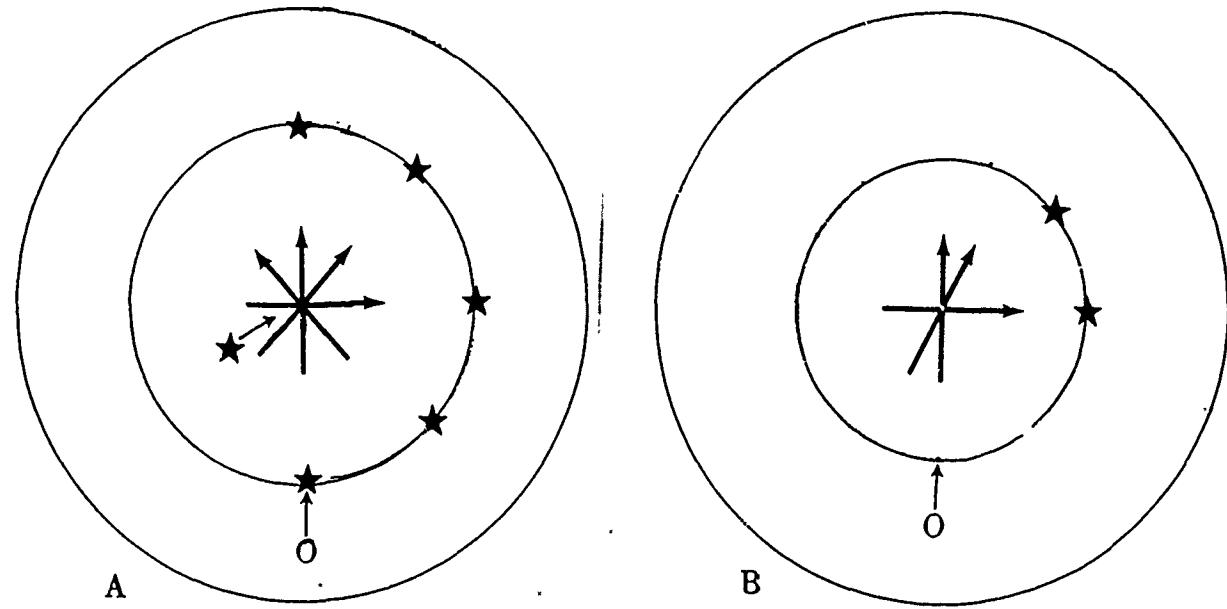
The experimental design of the current study spans the complete range of azimuthal geometries in increments of 45 degrees.

The zenith angle of observation for the current study model is set at the mean Brewster angle (53.2 degrees at 460 nm) (versus 60.0 degrees at 800 nm in the Chapman and Irani model) so that the effect of vertically polarized radiance due to wave-slope deviations from the mean surface can be compared with the observed horizontally polarized radiance. Only the horizontally polarized radiance was computed in the Chapman and Irani study.

In addition, the effect of upwelling subsurface radiance is incorporated into the current study model in order to estimate the effect that refracted radiance may have on the nonlinearity of the reconstructed slope spectrum.

Also, the effect of sub-resolution wave slopes is incorporated into the current study model in order to account for the spatially filtered wave-slope spectrum that exists beyond the maximum sampling frequency of the model.

Figure 1.4:3 illustrates the differences between the two studies, based on the combinations of synthetic imaging geometries that are analyzed in two dimensions.



FIGURES 1.4;3(a) & (b). 2D imaging geometries used in the current study (a) versus the Chapman and Irani study (b).

In summation, the experimental design for the current study requires the analysis of 72 distinct synthetic imaging combinations versus 6 combinations in the previous study. The analysis of these synthetic results will provide more complete information on the geometric effects of slope spectrum reconstruction than the simplified one-dimensional analysis in the previous work of Chapman and Irani [1981].

## 2.0 METHODS

### Overview

The current study model uses a two-dimensional Pierson-Stacy elevation power spectrum to create a synthetic wind-wave-slope surface that conforms to the results of a number of statistical studies [Cox & Munk, 1954a,b; Longuet-Higgins, 1957; Pierson & Stacy, 1973]. Downwelling skydome (solar and skylight) radiance for clear skies and upwelling subsurface radiance for clear, deep water are both synthetically created from analytic models. These synthetic radiance distributions are decomposed into their polarized components.

The synthetic radiance image is generated by first calculating the surface-slope distribution of the synthesized water surface and then applying the law of Malus to linearly map the radiance distributions, the polarized Fresnel reflection coefficients, and the surface projection coefficients to the surface-slope distributions to compute the reflected and refracted polarized radiance from each sample area in the direction of observation.

Figure 2.0:1 illustrates the methodology for image synthesis and error analysis used in the current study.

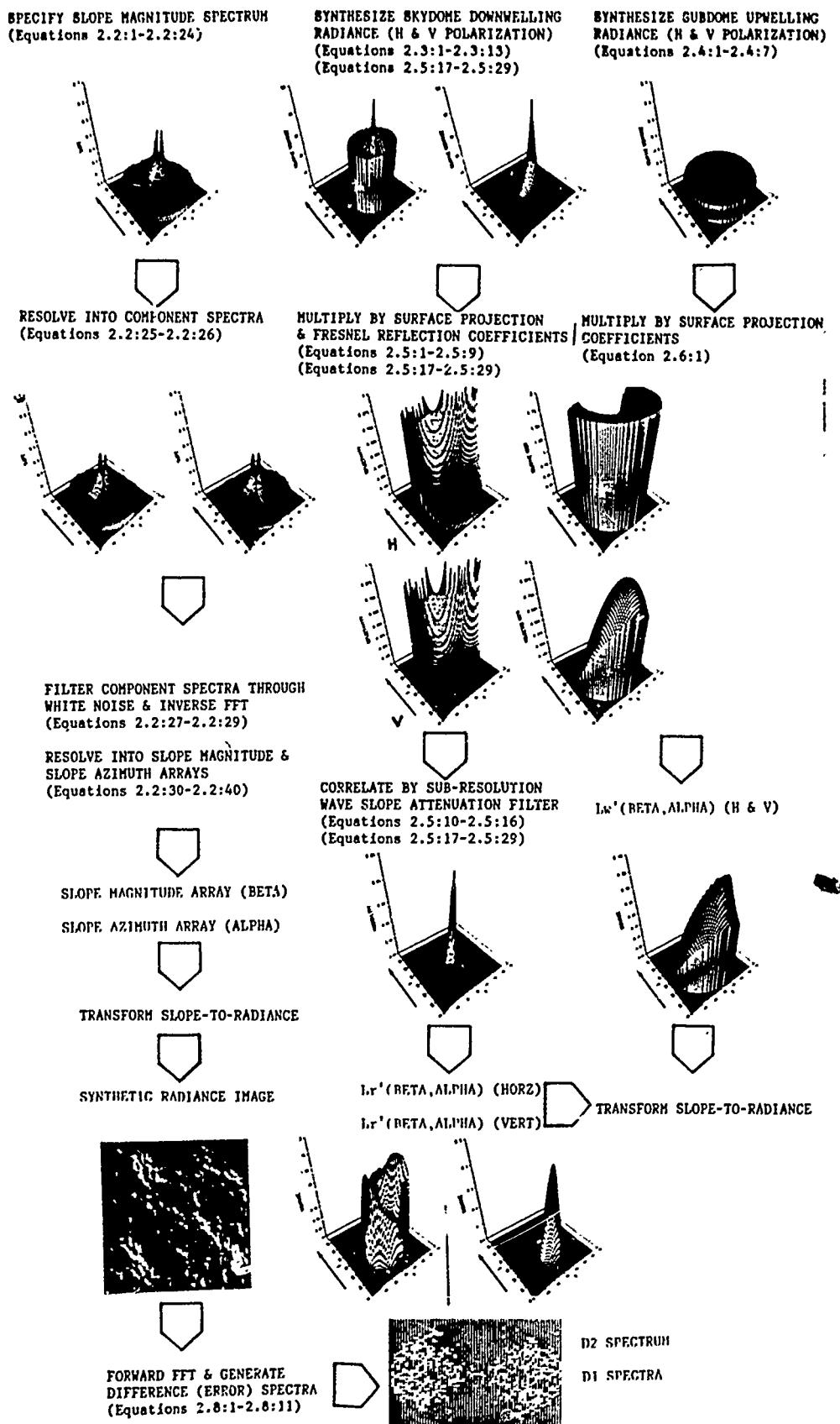


FIGURE 2.0:1. Schematic methodology of the current study.

Chapman and Irani [1981], in their error computations of the linear estimation of surface-slope spectra from image spectra, assumed that the nonlinearity of the slope-to-radiance transfer function was the principal source of error. They did note other possible sources of error but chose not to incorporate them in their study. The current study attempts to include those sources of variation which are directly or indirectly attributable to the geometric effect of water surfaces on the modulation of observed radiance:

- 1) Perspective Distortion
- 2) Variation of the Sensor Response Function
- 3) Spatial and Spectral Filtering

No sensor model is implied in this synthesis other than the existence of a spatial Nyquist frequency and a geometric location for a viewpoint. Chapman and Irani stated "we assume digital processing techniques are used so that perspective distortion and trends in the response function over the field of measurement can be removed prior to obtaining the image spectra" and "we assume that the optical modulation transfer function of the camera adequately filters the surface radiance spectrum and thereby eliminates aliasing errors." This synthesis does, however, incorporate a corre-

lation model to account for the spatially filtered, high-frequency wave-slope spectrum that exists beyond the maximum spatial sampling frequency.

4) Upwelling Radiant Energy

5) Scattered Radiant Energy

No atmospheric model is implied in this synthesis. However, a model for subsurface absorption and scattering is incorporated, providing radiance variation due to refraction of upwelling surface radiance by surface waves. The model of Chapman and Irani assumed that the surface is viewed 15-60 degrees from nadir so that surface reflection predominates. Also, they assumed that their camera was viewing through clear air and sensing in the near-infrared region so that the effects of path radiance and refracted subsurface radiance could be ignored.

6) Wave Shadows

The work of both Saunders [1967] and Goodell [1971] was reviewed in regard to potential models which compensate for wave shadowing or "hiding". The current linear model, as developed, adequately incorporates the effect described by Goodell for the established zenith angle of observation, the mean Brewster angle. To quote Goodell:

"Facets with small tilts, although more probable, reflect only small portions of the sky. Facets with greater tilts, although less probable, reflect much larger sky portions from higher up. Eventually the [G]aussian slope distribution allows only vanishingly small contributions from wave slopes so that the average contribution seems to come from sky elevations of [approximately 30 degrees, for an oblique observation looking below the horizon]. Thus, masking need not be postulated in order to explain the apparent 15-degree sea slope effect." [Goodell, 1971]

Figure 2.0:2 illustrates the effect modeled by Goodell.

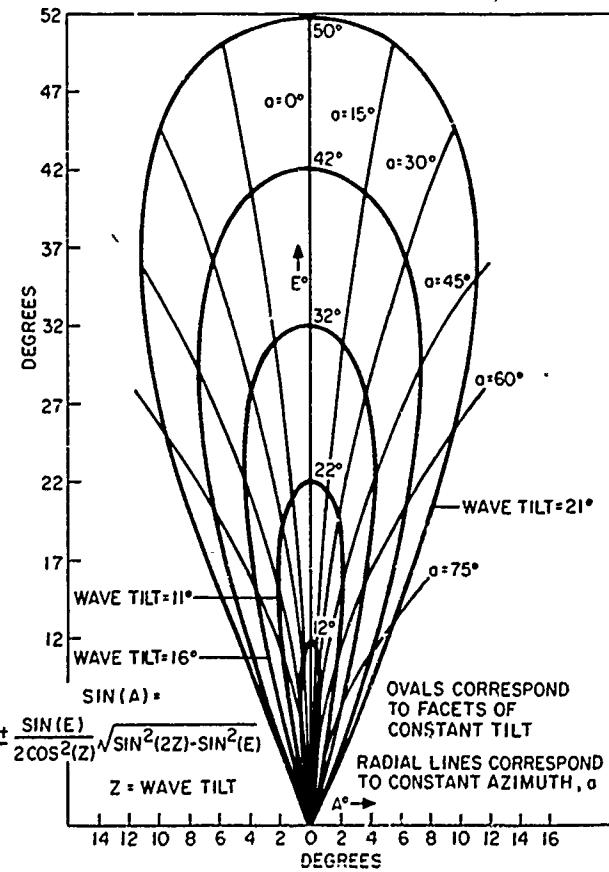


FIGURE 2.0:2. Sky portions reflected to an observer on the ocean surface, looking at the horizon. (from Goodell, 1971)

Although the effect is apparently reversed for near-zenith-viewing imagery, the modeling approach is the same for any angle of observation. In Figure 2.0:2, the slope distribution is mapped to sky coordinates; in the current study model, skydome radiance is mapped to slope coordinates.

It is intuitive that the effects of wave shadowing diminish as the observation point approaches zenith. Because the majority of wave slopes are less than 30 degrees, the effect of wave shadowing on Brewster-angle-viewing imagery (53.2 degrees from zenith) is assumed to be insignificant.

#### 7) Multiple Reflections

Only an elaborate Monte Carlo simulation or multiple-ray-tracing model could account for multiple reflections, simulations which are beyond the scope of this linear-model study. The assumption for this synthesis is that the probability of multiple reflections in near-nadir imagery is negligible.

#### Definition of Model Terms

The following list is a brief glossary of the terms used in the synthetic image model. These terms are formally defined in Chapters 2.1 through 2.8.

alpha = azimuth angle to (wind-wave) surface normal

beta = zenith angle to (wind-wave) surface normal

D<sub>0</sub> = scalar squared difference (error), integrated over both wavenumber (k) and azimuth (phi)

D<sub>1</sub>(k) = 1D difference (error) spectrum with respect to wavenumber (k), integrated over azimuth (phi)

D<sub>1</sub>(phi) = 1D difference (error) spectrum with respect to azimuth (phi), integrated over wavenumber (k)

D<sub>2</sub>(k<sub>x</sub>, k<sub>y</sub>) = 2D difference (error) spectrum

H = horizontally polarized fraction of radiance

H<sub>0</sub>(x,y) = horizontally polarized, reflected & refracted synthetic radiance image

H<sub>1</sub>(x,y) = horizontally polarized, reflected synthetic radiance image

H<sub>3</sub>(x,y) = H<sub>1</sub>(x,y) without the incorporation of sub-resolution wave-slope model

I = 3D incident radiance vector

k = dummy variable for scalar wavenumber

k<sub>x,y</sub> = scalar wavenumber relative to spatial coordinate axes

$K$  = 2D vector wavenumber =  $(k_x, k_y)$   
 $L_g$  = reflected radiance due to sun glitter  
 $L_p$  = atmospheric path radiance  
 $L_r$  = total reflected sky (+ sun) radiance =  $L_s + L_g$   
 $L_{ref}$  = reference Lsky radiance at zenith  
 $L_s$  = reflected radiance due to skylight  
 $L_{sky}$  = downwelling radiance from sky (including sun)  
 $L_u$  = upwelling radiance above water surface =  $L_r + L_w$   
 $L_w$  = upwelling subsurface refracted radiance above surface  
 $L_0(x, y)$  = total (unpolarized) reflected & refracted synthetic radiance image  
 $L_1(x, y)$  = total (unpolarized) reflected synthetic radiance image  
 $m_{x,y}(x, y)$  = (wind-wave) slope component arrays  
 $\mu$  = zenith angle to reflecting sky element  
 $\mu_0$  = angle from sun to sky element  
 $M_{x,y}(k_x, k_y)$  = (wind-wave) slope-component magnitude spectra  
 $\nu$  = azimuth angle to reflecting sky element  
 $\nu_0$  = angle between vector parallel to direction of polarization and projection of vector to surface plane  
 $N(k_x, k_y)$  = synthetic image radiance magnitude spectrum  
 $N$  = 3D surface normal vector  
 $\omega$  = angle of incidence

phi = azimuth angle to point of observation (180 deg)  
= dummy variable for azimuth angle

phi<sub>0</sub> = azimuth angle to sun

psi = multiplicative depolarization factor

P<sub>2D</sub>(k<sub>x</sub>,k<sub>y</sub>) = (wind-wave) elevation power spectrum

Pr(beta,alpha) = (wind-wave) slope probability distribution

R(omega) = Fresnel reflectivity (arbitrary polarization)

R<sub>0</sub> = unpolarized Fresnel reflectivity at normal incidence

R = 3D reflected radiance vector

S<sub>2D</sub>(k<sub>x</sub>,k<sub>y</sub>) = (wind-wave) slope power spectrum

S<sub>x,y</sub>(k<sub>x</sub>,k<sub>y</sub>) = (wind-wave) slope-component power spectra

theta = zenith angle to point of observation (53.2 deg)

theta<sub>0</sub> = zenith angle to sun

T<sub>a</sub> = atmospheric transmittance

T<sub>w</sub> = transmittance through water

v<sub>fric</sub> = friction velocity of wind at water surface

v<sub>1950</sub> = velocity of wind at height z=1950 cm above surface

V = vertically polarized fraction of radiance

WGN(k<sub>x</sub>,k<sub>y</sub>) = white-noise spectrum

wind\_az = azimuth angle to dominant wind direction (with  
180-degree ambiguity)

x,y = spatial coordinates

z = height above mean water surface

## 2.1 Radiometric Model for Image Synthesis

The total radiance reaching an overhead radiometric sensor be can defined by,

$$L_{tot} = (L_u * T_a) + L_p, \quad [2.1:1]$$

where

$L_u$  = the spectral upwelling radiance just above the water surface in direction of the sensor,

$T_a$  = the transmittance of the intervening atmosphere between the water surface and the sensor,

and

$L_p$  = the spectral path radiance due to light scattered in the direction of the sensor by the intervening atmosphere.

In general,  $L_p$  is much larger than  $L_u$  across the visible and infrared spectrum. The estimation of of  $L_p$  must be accurate within 1% in order to estimate  $L_u$  within 10%. This is known as the 'atmospheric correction problem' and is a well-defined subject of research. However, this synthesis does not include an atmospheric model since it is the variation of  $L_u$  with respect to several independent geometries that is the object of study. Therefore, this synthetic model will assume a perfectly non-absorbing, non-scattering atmosphere

and will set  $T_a = 1.0$  and  $L_p = 0.0$ . This simplification reduces the radiometric equation to:

$$L_{tot} = (L_u * 1.0) + 0.0 = L_u \quad [2.1.2]$$

for the duration of this study. More sophisticated models may vary  $T_a$  and  $L_p$  and may include a sensor model for spectral/spatial/temporal attenuation and distortion as well. An alternative perspective on this study is to consider that the model creates a synthetic radiant object rather than an image since no part of an imaging chain (including the intervening atmosphere) is modeled; we are interested only in the spatial effects of the water surface on the reflected/refracted radiance in the direction of a hypothetical distant sensor.

The above surface radiance  $L_u$  can be decomposed into three terms,

$$L_u = L_g + L_s + L_w, \quad [2.1:3]$$

where

$L_g$  = the above-surface radiance due to reflected sunlight (commonly known as sun glitter),

$L_s$  = the above-surface radiance due to reflected skylight,

and

$L_w$  = the above-surface radiance due to subsurface upwelling radiance propagating across the water surface.

Most oceanographic imaging has been directed at estimating  $L_w$ , as this term conveys information about water color and the factors which affect water color (non-aqueous constituents and water temperature, to name a few). For this reason, the large radiance due to glitter  $L_g$  has been generally avoided by the judicious selection of a minimizing geometry. However, reflected skylight  $L_s$  cannot be geometrically avoided; instead,  $L_s$  has usually been considered to be constant by assuming that a uniformly radiant skydome reflecting off a uniformly rough water surface yields an approximately Lambertian condition. This is one of the few instances where poor spatial resolution can be used to advantage. The reflected skylight term is then usually collected with the path radiance term:

$$\begin{aligned} L_{tot} &= L_w * T_a + (L_s * T_a + L_g * T_a + L_p) \\ &= L_w * T_a + (L_s * T_a + 0.0 + L_p) \\ &= L_w * T_a + (L_p') \end{aligned} \quad [2.1:4]$$

Thus, estimation of  $L_w$  has typically been reduced to cal-

culating  $T_a$  and  $L_p'$  and applying a linear atmospheric correction.

This synthetic model does not assume that  $L_w$  is the pre-eminent datum to collect: for information about the surface geometry, it is important to establish a relation between  $L_u$  and surface slope. In this context, the reflected radiance terms,  $L_g$  and  $L_s$ , are re-collected under the common term,

$$L_r = L_g + L_s \quad [2.1:5]$$

where

$L_r$  - the above-surface spectral radiance due to the reflected skydome, since, as the surface varies its slope, the surface normal vector points to the skydome independent of the position of the sun on the skydome.

In this study, an optimal imaging geometry provides for the accurate estimation of the slope magnitude spectrum over the broadest range of possible water-surface slope spectra. The general approach to this estimation is to maximize the spatial contrast of  $L_r$  and either minimize or correct for the spatial variation of  $L_w$ .

## 2.2 Fourier Synthesis of Wind-Wave Surfaces

The geometric model for Fourier synthesis of wave-slope surfaces closely follows the model of Chapman and Irani [1981] with one exception: the often substantial slope variance that exists above the spatial Nyquist frequency is calculated and the radiometric effect of this sub-resolution slope variance is introduced within the illumination models.

### Specification of a 1D Elevation Power Spectrum

The first step is to specify a 1D wave elevation power spectrum. Several spectra were previously cited: Pierson-Neumann [Borrego & Machado, 1985; Pierson-Moskowitz [Mastin et al., 1987], and modified Pierson-Stacy [Chapman & Irani, 1981]. For this study, the original 1D Pierson-Stacy power spectrum is employed for two reasons: 1) the spectral scale is spatial wavenumber [radians per centimeter or just /cm] and 2) Pierson and Stacy [1973] went into great detail in their report to attempt to correlate their spectrum with the empirical results of other researchers, including the Cox and Munk [1954a,b] photographic study.

The equations of the 1D wave elevation power spectrum  $P(k)$  come directly from Pierson and Stacy [1973].  $P(k)$  is

defined over five spectral regions:

$P_1$  = the gravity wave-gravity equilibrium spectral range,

$$P_1(k) = (a/2*k^3)*\text{EXP}(-b*g^2/(v_{1950}^4*k^2))$$

for  $0 < k < k_1$ , [2.2:1]

$P_2$  = the Kitaigorodskii spectral range,

$$P_2(k) = a/(2*\text{SQRT}(k_1)*k^{2.5})$$

for  $k_1 < k < k_2$ , [2.2:2]

$P_3$  = the Leykin-Rosenberg spectral range,

$$P_3(k) = (a*d)/(2*(k_3^p)*(k^{3-p}))$$

for  $k_2 < k < k_3$ , [2.2:3]

$P_4$  = the capillary spectral range,

$$P_4(k) = (a*d)/(2*k^3)$$

for  $k_3 < k < k_{nu}$ , [2.2:4]

and

$P_5$  = the Cox viscous cut-off range,

$$P_5(k) = (E/(nu*g))*v_{fric}^3*k_{max}^6/k^9$$

for  $k_{nu} < k < \text{infinity}$ , [2.2:5]

where

$$d = (1.274 + (0.0268*v_{fric}) + (6.03E-5*v_{fric}^2))^2, \quad [2.2:6]$$

$$k_1 = (k_2*v_{min}^2)/v_{fric}^2, \quad [2.2:7]$$

$$k_{nu} = (0.5756*SQRT(v_{fric})*k_{max})/(d^{1/6}), \quad [2.2:8]$$

and

$$p = LOG(d/(v_{fric}/v_{min}))/LOG(k_3/k_2), \quad [2.2:9]$$

such that

a = 0.0081 = the Phillips constant [unitless],

b = 0.74 [unitless],

E/(nu\*g) = 1.473E-4 [unitless],

g = 980.0 = the acceleration of gravity [cm/sec<sup>2</sup>],

k<sub>2</sub> = 0.359 [rad/cm],

k<sub>3</sub> = 0.942 [rad/cm],

k<sub>max</sub> = 3.63 [rad/cm]

v<sub>min</sub> = 12.0 [cm/sec] = the minimum friction velocity defined by the Pierson-Stacy model,

v<sub>1950</sub> = the wind velocity at 1950 cm above water level [cm/sec], and

v<sub>fric</sub> = the wind friction velocity at the water surface (i.e. z = 0 cm).

v<sub>fric</sub> and the wind velocity v [cm/sec] at height z [cm] have an empirical correspondence under the assumption of a neutrally stratified atmosphere:

$$v(z) = (v_{fric}/0.4) * \ln(z/z_0) \quad [2.2:10]$$

and

$$z_0 = (0.684/v_{fric}) + (4.28E-5*v_{fric}^2) - 0.0443 \quad [2.2:11]$$

Values for  $v_{fric}$ ,  $v_{1950}$ , as well as  $v_{250}$ ,  $v_{1000}$ , and  $v_{1250}$  are interpolated from values tabulated by Pierson and Stacy. (Refer to V\_DATA.DAT input data file in Appendix III.)

Figure 2.2:1 graphs  $k*P(k)$  versus  $\log_{10}(k)$  for friction velocities of 12.0, 36.0, and 60.0 cm/sec, to illustrate of the effect of wind velocity on the character of the elevation power spectrum. Note that most of the power is located in the gravity wavenumber region, i.e. below  $k = 0.002/\text{cm}$ . The curve for 12.0 cm/sec has a maximum value of 30.0  $\text{cm}^2$ .

#### Specification of a 2D Directional Spreading Function

The 2D power spectrum is created via a directional spreading factor  $D(k, \phi)$  that attenuates the 1D spectrum at the azimuth angle  $\phi$  from the upwind/downwind directions:

$$P_{2D}(k, \phi) = P(k) * D(k, \phi) \quad [2.2:12]$$

The equations of the 1D wave elevation power spectrum  $P(k)$

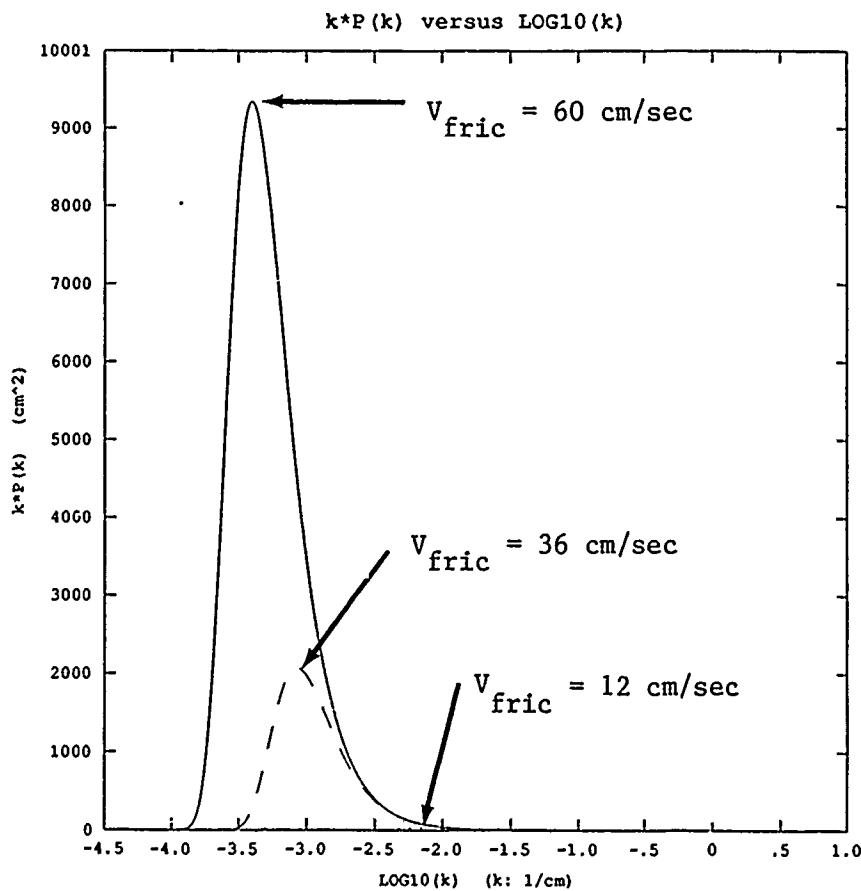


FIGURE 2.2:1. The Pierson-Stacy 1D elevation power spectrum  $k*P(k)$  versus  $\text{LOG10}(k)$ .

come directly from Pierson and Stacy [1973] with a small modification to allow rotation of the 2D spectra relative to the wind azimuth `wind_az`.

The directional spreading function is defined:

$$D(k, \phi, \text{wind\_az}) = (A * (1 - B)) + (B * C) \quad [2.2:13]$$

where

$$A = (8/(3\pi)) * (D^2), \quad [2.2:14]$$

$$B = \text{EXP}(-g^2/(2*k^2*v_{1950}^4)), \quad [2.2:15]$$

and

$$C = (A*D + 0.5)/\pi \quad [2.2:16]$$

such that

$$D = \text{COS}(\text{wind\_az}-\phi). \quad [2.2:17]$$

#### Creation of the Slope-Component Spectra

The third step is to derive the two slope-component spectra from the scalar power spectrum. The 2D spectrum must first be normalized by  $k$  over the range of wavenumber:

$$P_{2D}'(k, \phi) = P(k)*D(k, \phi)/k \quad [2.2:18]$$

since

$$P(k) = \int P_{2D}'(k, \phi)*k*d\phi = \int P_{2D}(k, \phi)*d\phi. \quad [2.2:19]$$

The two slope-component spectra are obtained by multiplying the scalar 2D elevation spectrum by the squares of the orthogonal components of the wave vector  $K$  that are aligned with the image coordinate system  $(x, y)$ :

The slope power spectrum is

$$S(k, \phi) = k^2 * P_{2D}'(k, \phi), \quad [2.2:20]$$

which leads to

$$S_x(k_x, k_y) = k_x^2 * P_{2D}'(k, \phi) = k_x^2 * P_{2D}'(k_x, k_y) \quad [2.2:21]$$

and

$$S_y(k_x, k_y) = k_y^2 * P_{2D}'(k, \phi) = k_y^2 * P_{2D}'(k_x, k_y), \quad [2.2:22]$$

since

$$|k| = \text{SQRT}(k_x^2 + k_y^2) = k \quad [2.2:23]$$

and

$$\begin{aligned} P_{2D}'(k) * dk &= P_{2D}(k_x, k_y) * k * dk \\ &= P_{2D}(k * \cos(\phi), k * \sin(\phi)) * k * dk \\ &= P_{2D}(k, \phi) * k * dk d\phi \\ &= P_{2D}'(k, \phi) * dk d\phi. \end{aligned} \quad [2.2:24]$$

The square roots of the slope-component power spectra yield the corresponding slope-component magnitude spectra, which are used as filters when generating the synthetic slope arrays:

$$M_x(k_x, k_y) = \text{SQRT}(S_x(k_x, k_y)) \quad [2.2:25]$$

and

$$M_y(k_x, k_y) = \text{SQRT}(S_y(k_x, k_y)). \quad [2.2:26]$$

#### Creation of a Frequency-Domain White Noise Image

The fourth step is to create a scalar matrix of spatial white noise, i.e. uniformly distributed, uncorrelated random

values with unit magnitudes ranging between +pi and -pi. The Fourier transform of this matrix yields a complex phase matrix with conjugate symmetry about its central ordinate:

$$\int_{-1}^{+1} [wgn(x,y)] = WGN(k_x, k_y) \quad [2.2:27]$$

#### Frequency-Domain Filtering and Back Transformation

WGN is filtered by  $M_x$  and  $M_y$  in the frequency domain to create the two complex Fourier transforms of the synthesized wave-slope components. The inverse Fourier transform yields the two slope-component arrays:

$$m_x(x,y) = \int_{-1}^{-1} [M_x(k_x, k_y) * WGN(k_x, k_y)] \quad [2.2:28]$$

and

$$m_y(x,y) = \int_{-1}^{-1} [M_y(k_x, k_y) * WGN(k_x, k_y)]. \quad [2.2:29]$$

Both  $m_x$  and  $m_y$  are real scalar arrays of slope components because of the conjugate symmetry.

#### Conversion to Slope Magnitude and Slope Azimuth

The fifth and final step is to convert the two slope-component arrays into a slope magnitude array  $\beta(x,y)$  and a slope azimuth array  $\alpha(x,y)$ , where  $\beta$  is the maximum

slope at the coordinate  $(x, y)$  and alpha is the azimuth angle to the direction of beta. The angles alpha and beta are related to  $m_x$  and  $m_y$  by the geometric relations:

$$\tan(m_x) = \tan(\beta) \cdot \cos(\alpha) \quad [2.2:30]$$

and

$$\tan(m_y) = \tan(\beta) \cdot \sin(\alpha). \quad [2.2:31]$$

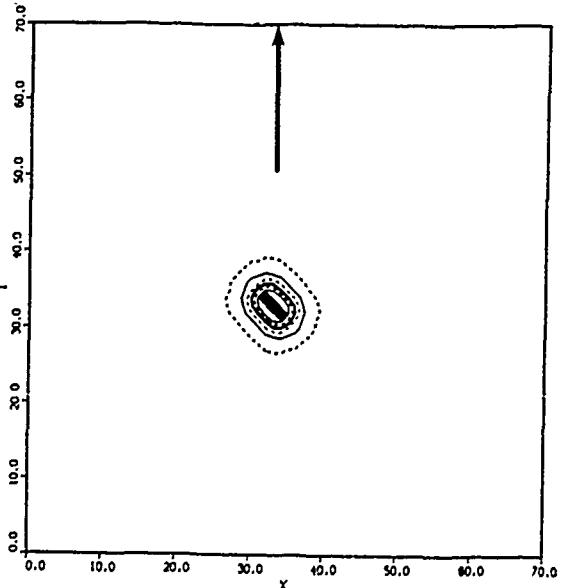
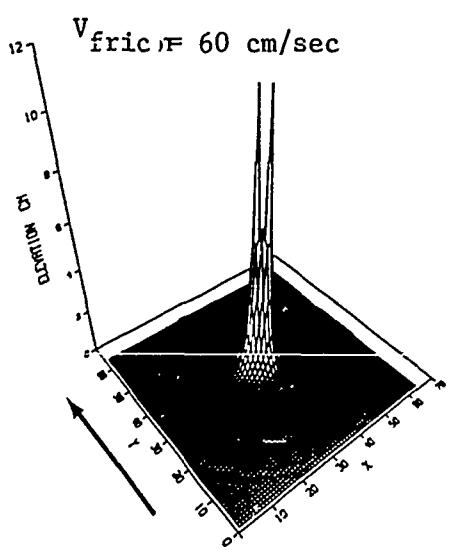
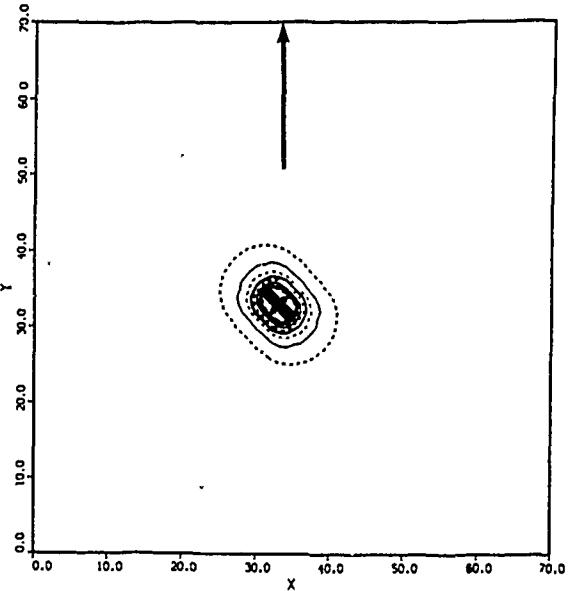
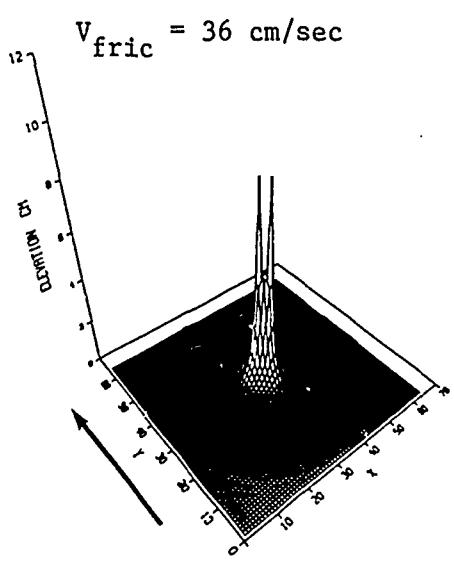
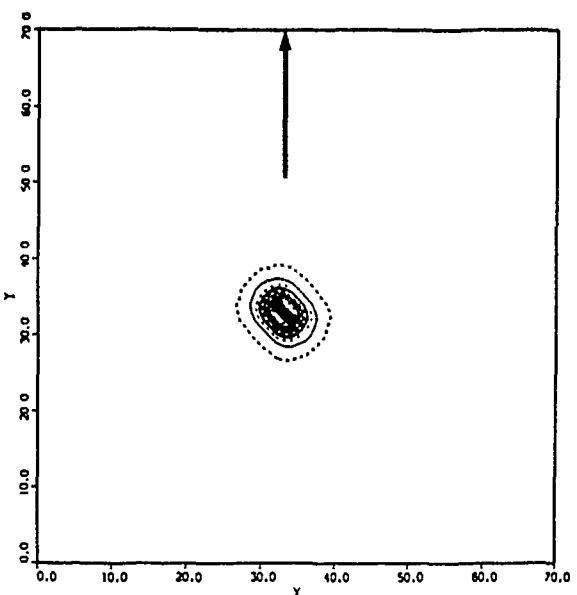
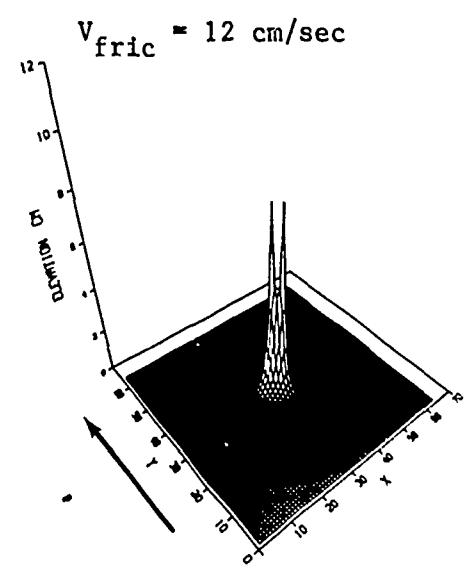
Therefore, the maximum slope magnitude of an array facet is

$$\beta = \text{ATAN} [\tan(m_x)^2 + \tan(m_y)^2] \quad [2.2:32]$$

and the azimuth angle in the direction of the maximum slope is

$$\alpha = \text{ATAN} [\tan(m_y)/\tan(m_x)]. \quad [2.2:33]$$

Figures 2.2:2 through 2.2:4 illustrate the 2D Pierson-Stacy elevation magnitude spectra for friction velocities of 12.0, 36.0, and 60.0 cm/sec, respectively. The spectra are oriented with wind azimuth = -45 deg (or +135 deg with ambiguity).

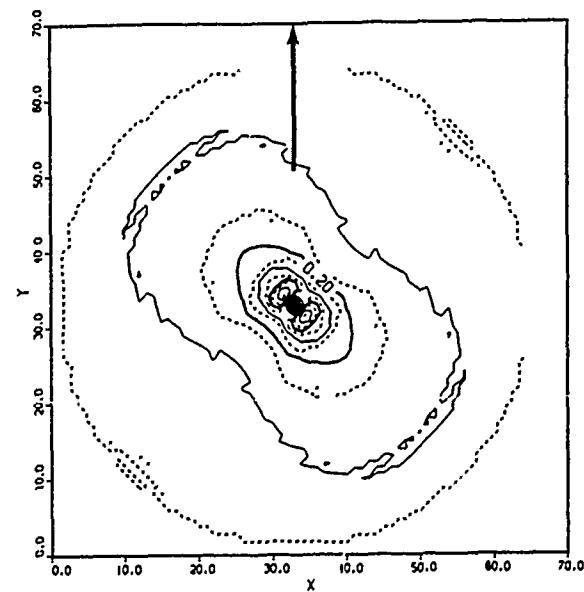
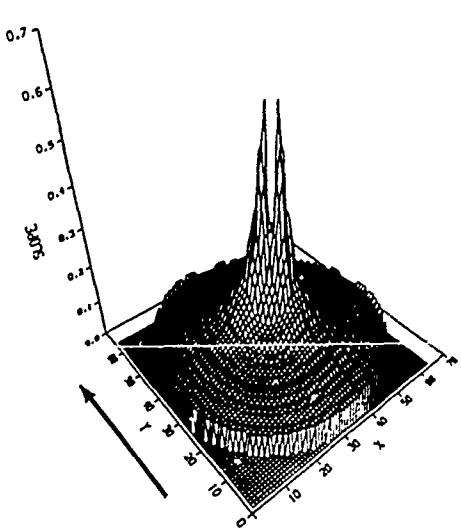
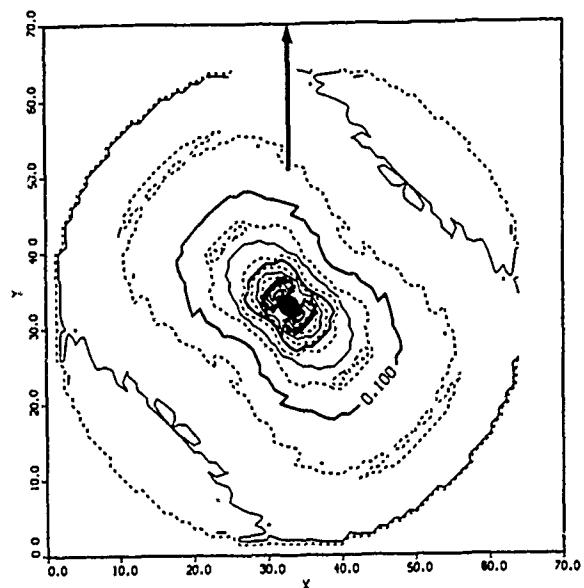
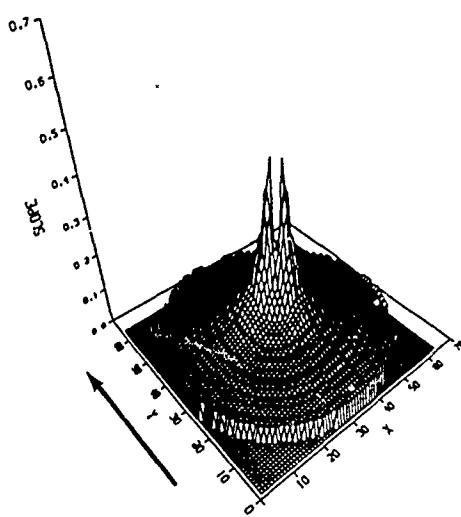
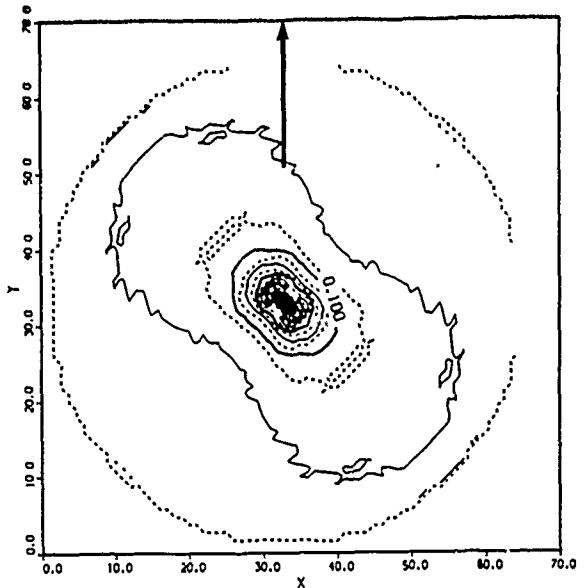
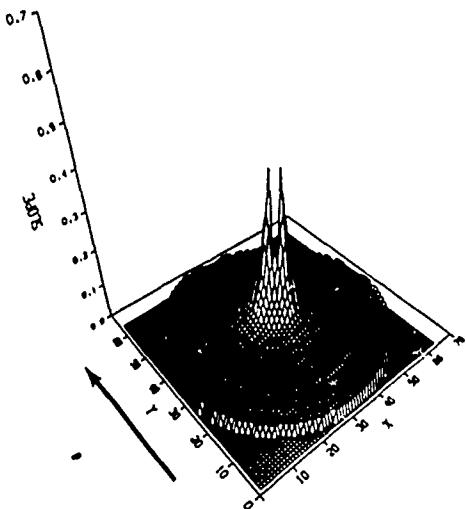


FIGURES 2.2:2-4. Elevation magnitude spectra,  $\sqrt{P(k_x, k_y)}$ .

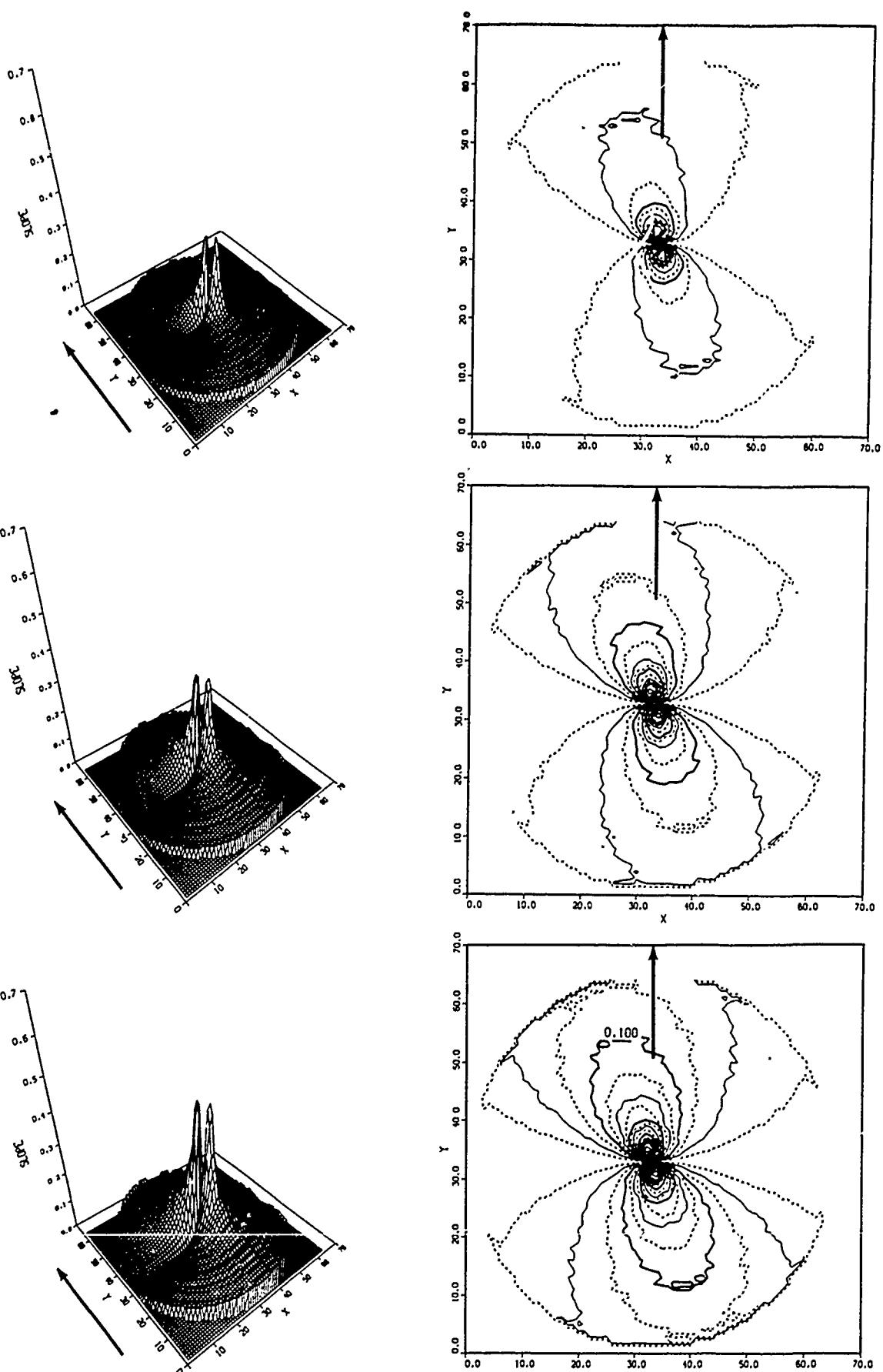
Figures 2.2:5 through 2.2:7 illustrate the 2D Pierson-Stacy slope magnitude spectra for friction velocities of 12.0, 36.0, and 60.0 cm/sec, respectively. The spectra are oriented with wind azimuth = -45 deg (or 135 deg with ambiguity).

Figures 2.2:8 through 2.2:10 illustrate the x-components of the above slope spectra. Figures 2.2:11 through 2.2:13 illustrate the y-component slope spectra. The spectra are oriented with wind azimuth = -45 deg (or 135 deg with ambiguity).

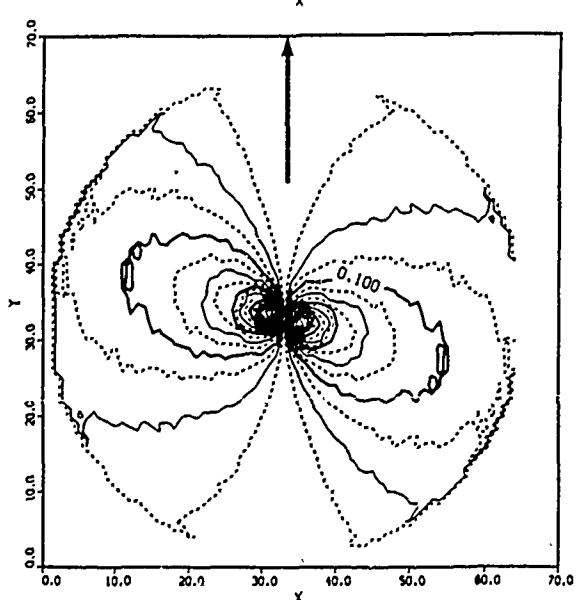
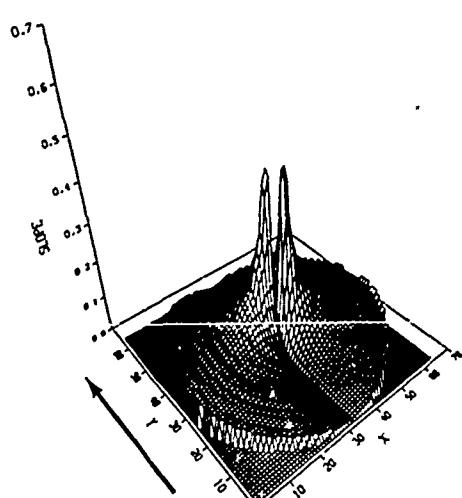
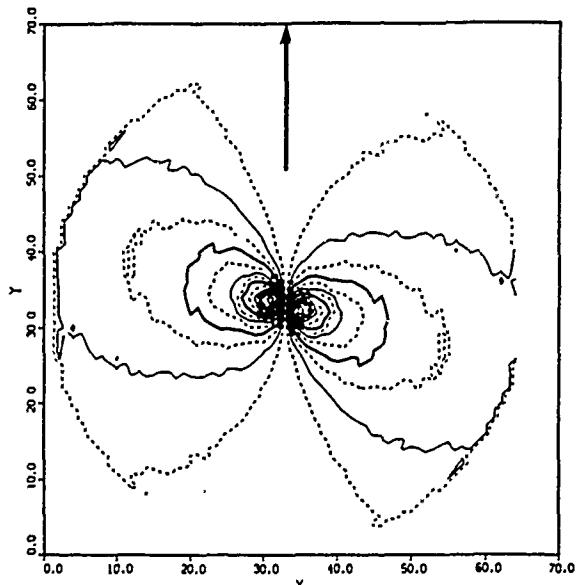
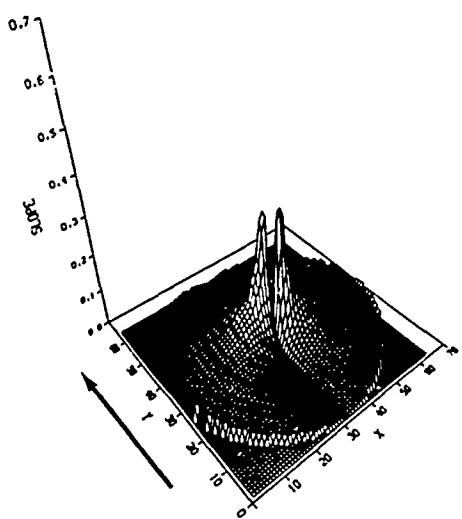
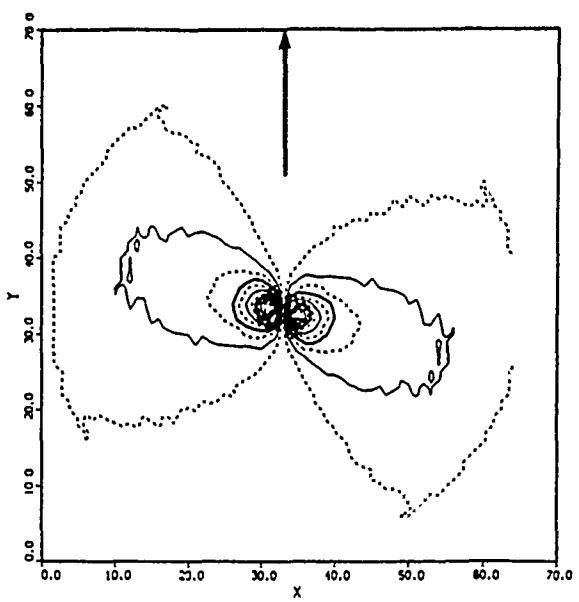
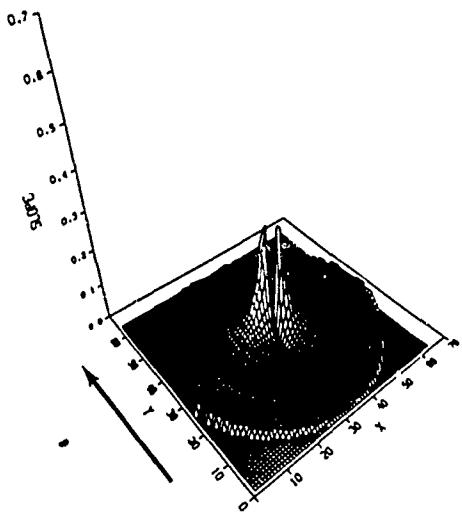
Figures 2.2:14 through 2.2:16 illustrate the sample slope distributions  $Pr(\beta, \alpha)$  for synthetic realizations using the above slope spectra. The spectra are oriented with wind azimuth = -45 deg (or 135 deg with ambiguity).



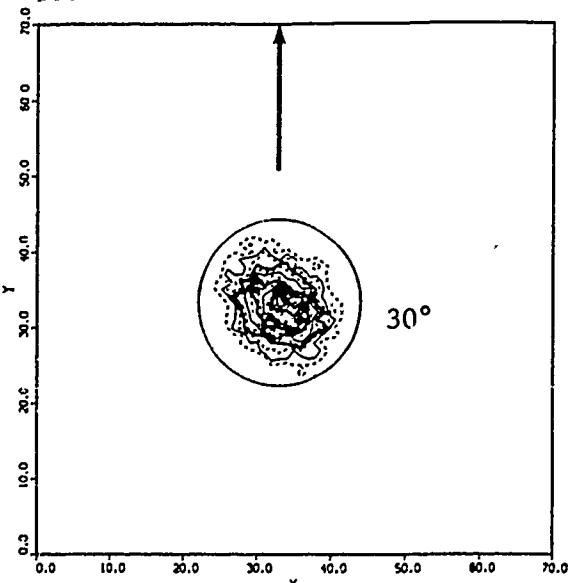
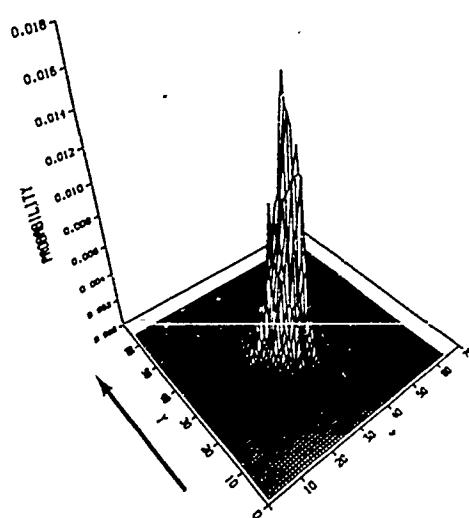
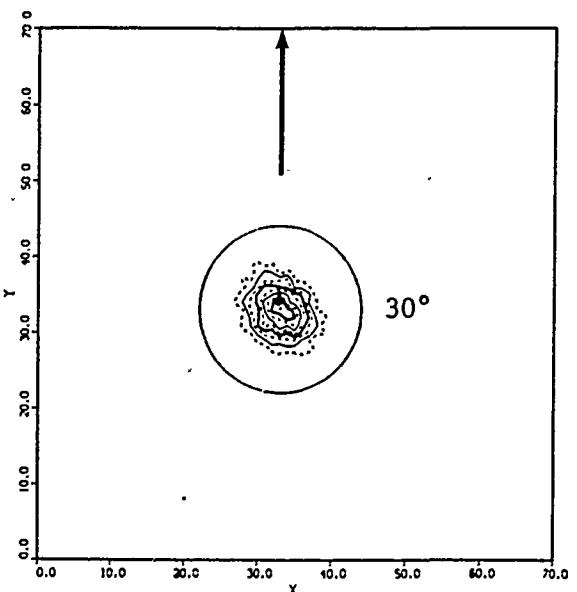
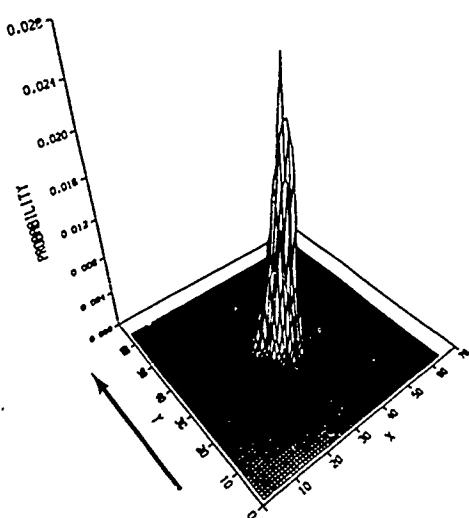
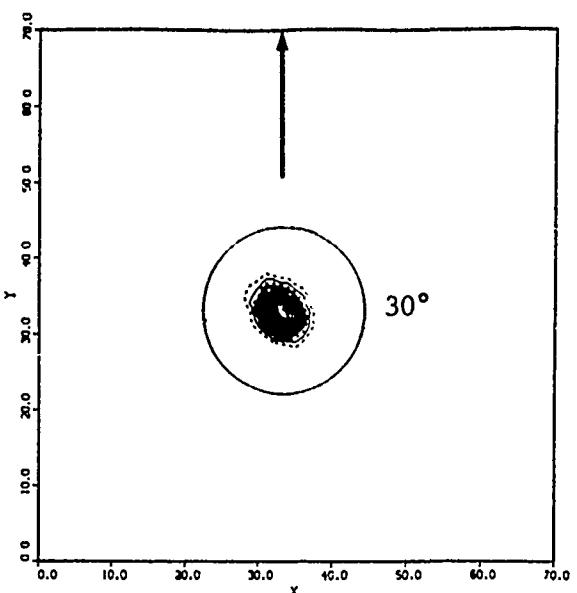
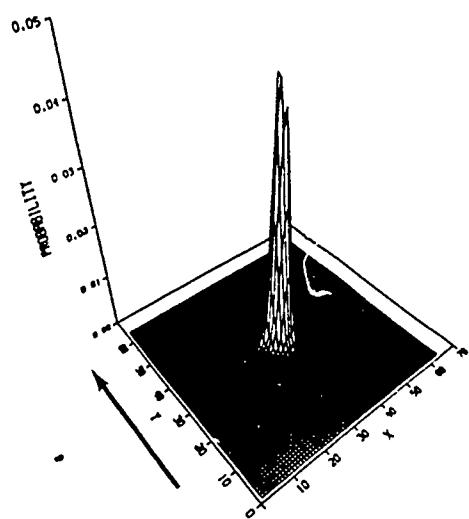
FIGURES 2.2:5-7. Slope magnitude spectra,  $M_0(k_x, k_y)$ .



FIGURES 2.2:8-10. X-components of slope magnitude spectra.



FIGURES 2.2:11-13. Y-components of slope magnitude spectra.



FIGURES 2.2:14-16. Sample slope probability distributions,  $\text{Pr}(\beta, \alpha)$ .

### Calculation of Total and Component Slope Variance

The slope variance is calculated from the 1D slope power spectrum for a specified range of wavenumber:

$$\text{VAR}(k_1, k_2) = \int_{k_1}^{k_2} S(k) * dk = \int_{k_1}^{k_2} k * S(k) * dLN(k) \quad [2.2:34]$$

The unresolved (i.e. both x- and y-components combined) slope variance is calculated from the 2D slope power spectrum for a specified range of wavenumber:

$$\text{VAR}(k_1, k_2) = \iint_{\vec{k}_1}^{\vec{k}_2} S(k, \phi) * dk d\phi = \iint_{\vec{k}_1}^{\vec{k}_2} S(k_x, k_y) * dk_x dk_y \quad [2.2:35]$$

Likewise, the x-component and y-component slope variance can be calculated by substituting  $S_x$  or  $S_y$  for  $S$  in Equation [2.2:35].

The 2D distribution of slopes in the downwind and crosswind directions correlate well with the results of Cox and Munk [1954a,b] and this is one measure of the adequacy of the Pierson-Stacy model. However, Pierson and Stacy [1973] suggest that the optical system used by Cox and Munk acted as a high-frequency spatial filter on the reflected

radiance from the sea surface. Pierson and Stacy concluded that the calculation of variance from the integration of their power spectrum over  $k$  and  $\phi$  yielded variances which fit reasonably well with the Cox and Munk model when they imposed an upper bound of  $k = 1.2$  radians per centimeter. This corresponds to a wavelength of 5.24 centimeters. To maintain consistency between the results of two previous efforts, this model assumes a maximum spatial frequency of  $k = 1.2$  radians per centimeter.

Figure 2.2:17 graphs  $k \cdot S(k)$  versus  $\log(k)$  for three friction velocities, 12.0, 36.0, and 60.0 cm/sec. This graph has the interesting property that the area under the curve equals the total variance and the area between two wavenumbers equals the fraction of the total variance between the two wavenumbers, as specified in Equation [2.2:34]. The wavenumber range sampled by the current study model is located between the two vertical lines.

Figure 2.2:18 graphs the fractional and total slope variance over the friction velocity range from 12.0 to 60.0 cm/sec. The lower line represents the integration of variance up to  $k_1 = 0.0375/\text{cm}$ .  $k_1$  is the fundamental spatial frequency for

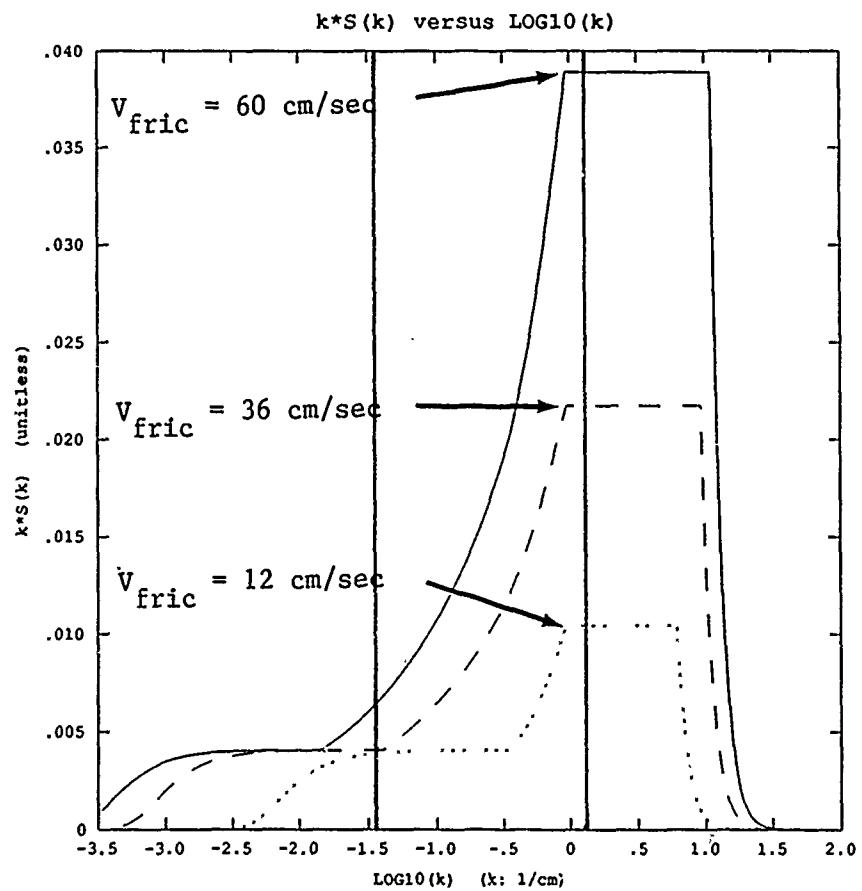


FIGURE 2.2:17. The 1D Pierson-Stacy slope power spectrum k\*S(k) versus LOG10(k).

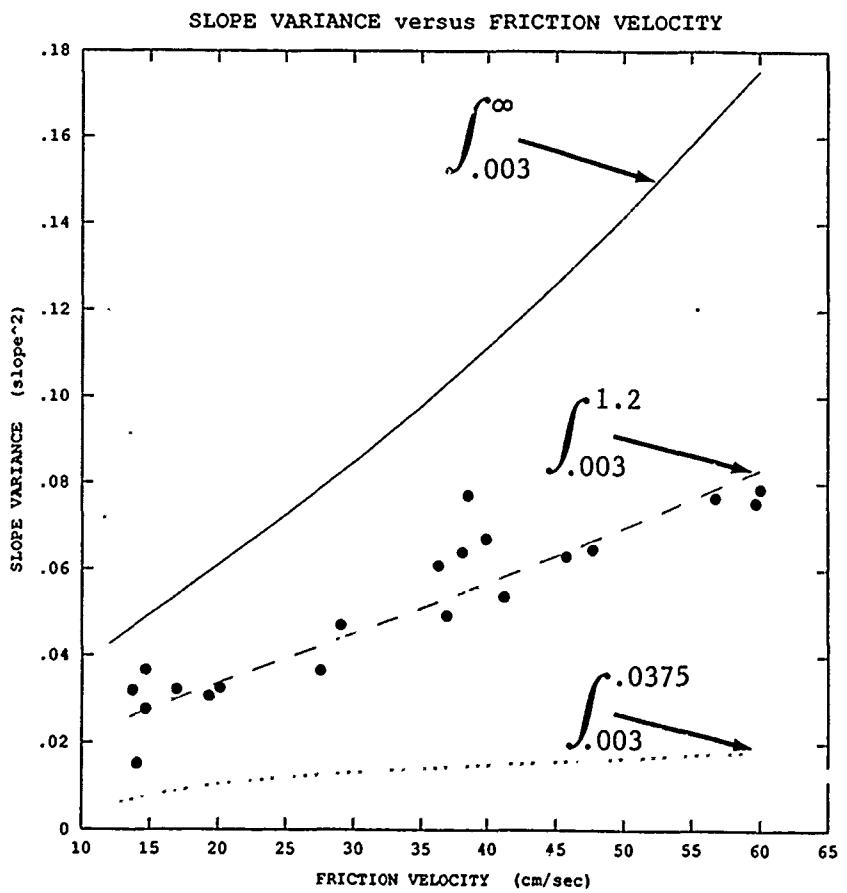
the current study model, that is

$$k_1 = k_{\max}/(n/2) = (1.2/\text{cm})/(64/2) = 0.0375/\text{cm} \quad [2.2:36]$$

The middle line represents the integration of variance up to  $k_2 = 1.2/\text{cm} = k_{\max}$ .  $k_{\max}$  is the maximum spatial frequency

for the current study model. The upper line represents the integration of variance over the entire range of wavenumber. The points represent the measurements of Cox and Munk [1954b].

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FIGURE 2.2:18. Slope Variance versus Friction Velocity.

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Figures 2.2:1, 2.2:17, and 2.2:18 are identical in magnitude to Figures 8.3, 8.4, and 9.1 of Pierson and Stacy [1973]. Note that there is a large fraction of the total slope

variance above  $k = 1.2/\text{cm}$ . For this and preceding studies, this fraction represents the sub-resolution slope variance; its effect on the spatial attenuation of the reflected and refracted radiance will be introduced into this study by a correlation filter model described in Chapter 2.5.

### 2.3 Synthesis of Skydome Downwelling Radiance

#### Chapman & Irani Model

Chapman and Irani [1981] incorporated a combination of the modified clear-sky luminance model of Hopkinson [1954] and the pure Rayleigh-scattering polarization model. The analytic Hopkinson model is defined:

$$L_{sky}(\mu, \nu) = L_{ref} * A * B \quad [2.3:1]$$

where

$L_{sky}(\mu, \nu)$  = the downwelling radiance at skydome coordinates  $(\mu, \nu)$ ,

$\mu$  = the zenith angle to the skydome location  $[0:+90 \text{ deg}]$ ,

$\nu$  = the azimuth angle to the skydome location  $[-180:+180 \text{ deg}]$ ,

$L_{ref}$  =  $L_{sky}(0, 0)$  = the reference radiance at zenith, [2.3:2]

$A = (1+\cos(\mu_0)^2)/(1-\cos(\mu_0)^2)$ , [2.3:3]

$B = 1-\exp(-0.32*\sec(\mu))$ , [2.3:4]

and

$$\begin{aligned} \cos(\mu_0) &= \cos(\theta_0) * \cos(\mu) \\ &\quad + \sin(\theta_0) * \sin(\mu) * \cos(\nu - \phi_0) \end{aligned} \quad [2.3:5]$$

such that

$\mu_0$  = the angle from the sun to a sky element,

$\theta_0$  = the zenith angle to the sun,

and

$\phi_{ij}$  = the azimuth angle to the sun.

Hopkinson's model was one of the earlier responses to a 1953 CIE proposal to establish a standard, analytic clear-sky radiance distribution to serve as a basis for daylight factor calculations. A broadband telephotometer was modified to measure the integrated radiance over a 5-degree Instantaneous Field Of View (IFOV). The actual measurements used in this model were made under the clear blue skies of Stockholm (approximately 60 degrees North latitude).

#### Current Study Model

This study uses the current CIE standard, analytic clear-sky radiance distribution [CIE,1973] for the sole reason that it is a current, accepted standard:

$$L_{sky}(\mu, \nu) = L_{ref} * A * B / C \quad [2.3:6]$$

where

$$A = 0.91 + 10 * EXP(-3 * \mu_0) + 0.45 * COS(\mu_0)^2, \quad [2.3:7]$$

$$B = 1 - EXP(-0.32 / COS(\mu)), \quad [2.3:8]$$

and

$$C = 0.274 * (0.91 + 10 * EXP(-3 * \theta_0) + 0.45 * COS(\theta_0)^2). \quad [2.3:9]$$

All other parameters are as previously defined. Although

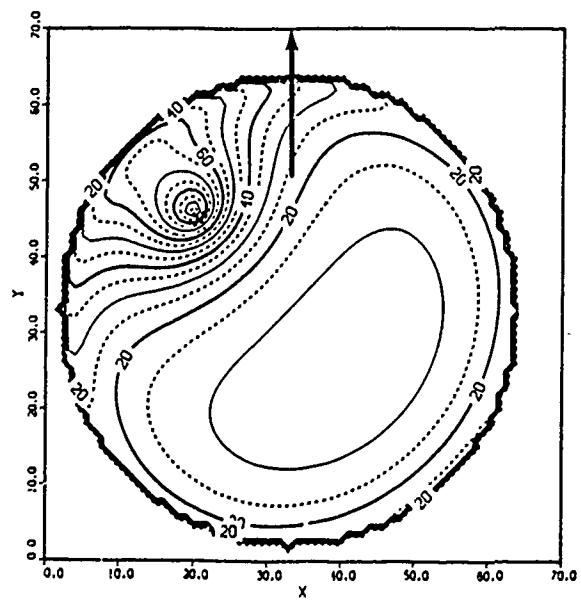
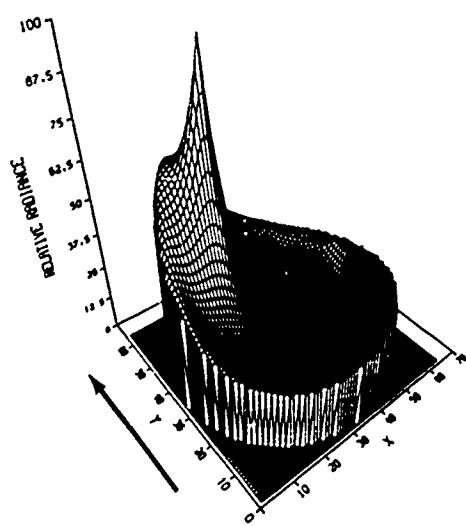
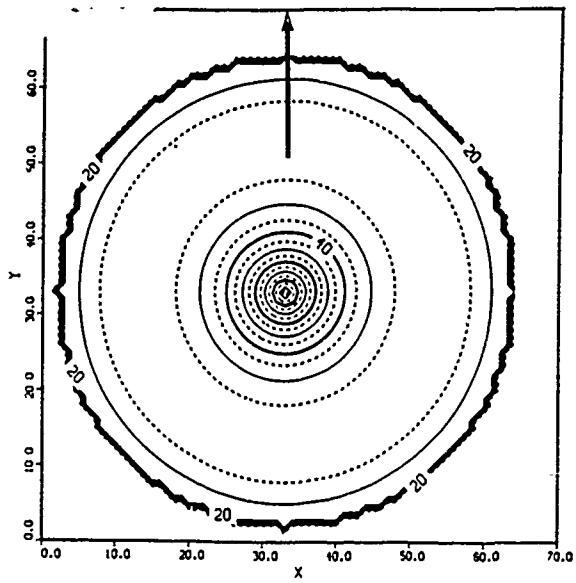
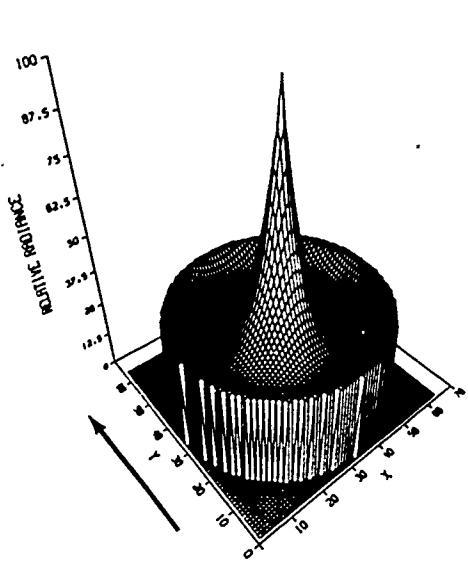
both the CIE model and the Hopkinson model provide broadband spectral radiance values, the results from the CIE model must be used with caution when applied to the narrowband spectral radiance in the vicinity of the peak solar spectral radiance, particularly in the vicinity of 460 nanometers. Chapman and Irani [1981] applied the Hopkinson model to their model requirement for a near-infrared narrowband distribution and noted that "measurements of sky radiance which we have made show [ $L_{ref}$ ] varies more than a factor of 3 for apparently clear skies."

This study requires only a reasonable relative spatial distribution of radiance and does not attempt to acquire absolute radiance values. Therefore, the convention used in this study sets the value of  $L_{ref}$  such that the radiance at the solar location is equal to 100.0 in arbitrary units of radiance. All other radiance values can be expressed as a relative percentage of the total radiance in the vicinity of the sun.

Figures 2.3:1 and 2.3:2 illustrate two of the 11 synthetic total skydome radiance distributions used in this study, for a solar azimuth of -45 degrees and solar zenith angles of 0.0 and 53.2 degrees, respectively. The arrows in these

(and following) figures point to zero azimuth.

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FIGURES 2.3:1-2. Synthetic skydome radiances distributions.

### Polarized Rayleigh Scattering Model

The fraction of the unpolarized solar radiance that is horizontally polarized by Rayleigh scattering in the atmosphere is defined:

$$H = 0.5*(1-\psi)+\psi*\cos(\nu_0)^2 \quad [2.3:10]$$

where

$$\psi = 0.94*(1-\cos(\mu_0)^2)/(1+\cos(\mu_0)^2), \quad [2.3:11]$$

and

$$\begin{aligned} \nu_0 = & \text{ACOS}[\text{SIN}(\theta)*\text{COS}(\theta_0) \\ & - \text{SIN}(\theta_0)*\text{COS}(\theta)*\text{COS}(\phi-\phi_0)/\text{SIN}(\mu_0)]. \end{aligned} \quad [2.3:12]$$

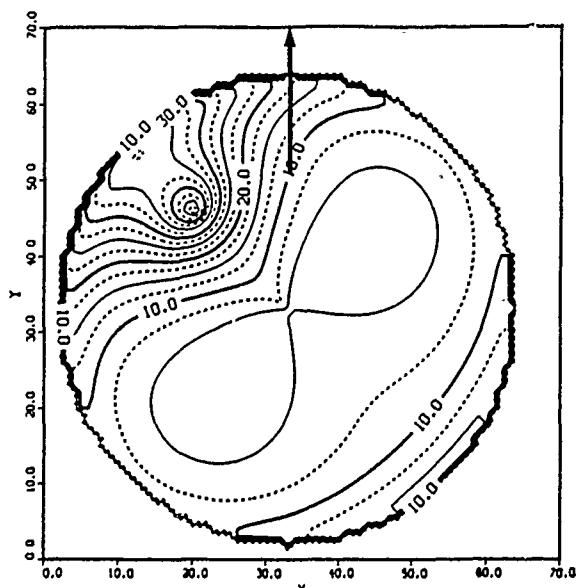
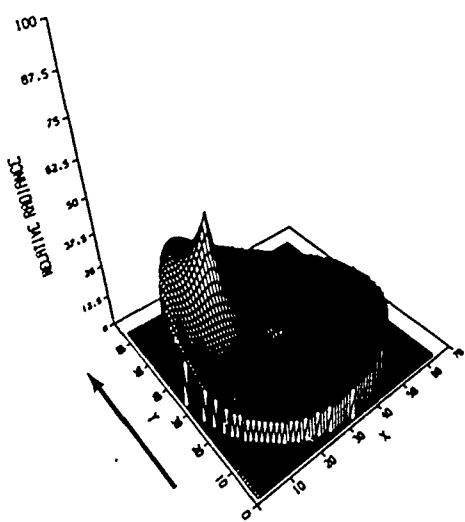
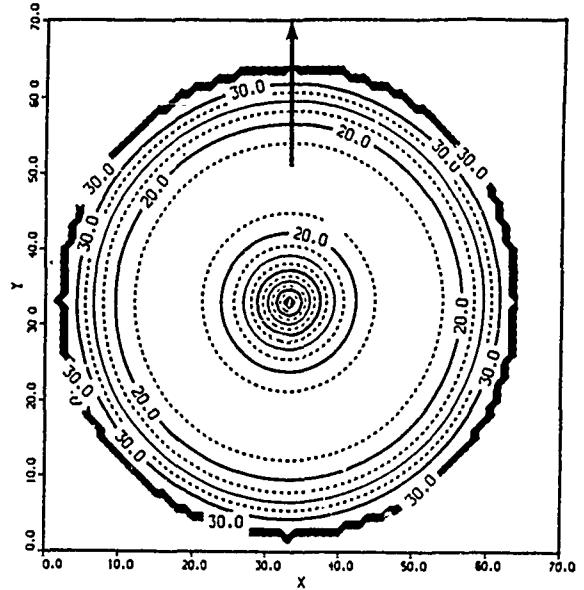
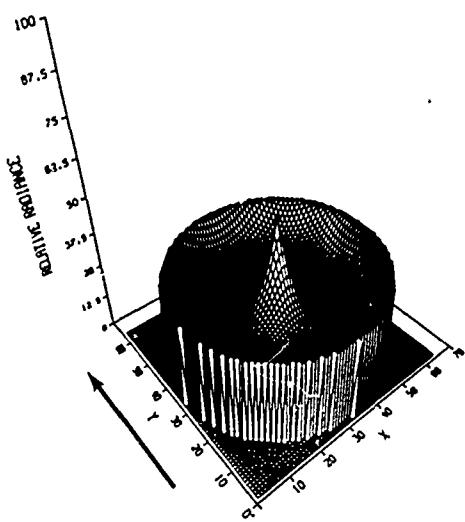
Likewise, the fraction of the unpolarized solar radiance that is vertically polarized by Rayleigh scattering in the atmosphere is defined:

$$V = 0.5*(1-\psi)+\psi*(1-\cos(\nu_0)^2). \quad [2.3:13]$$

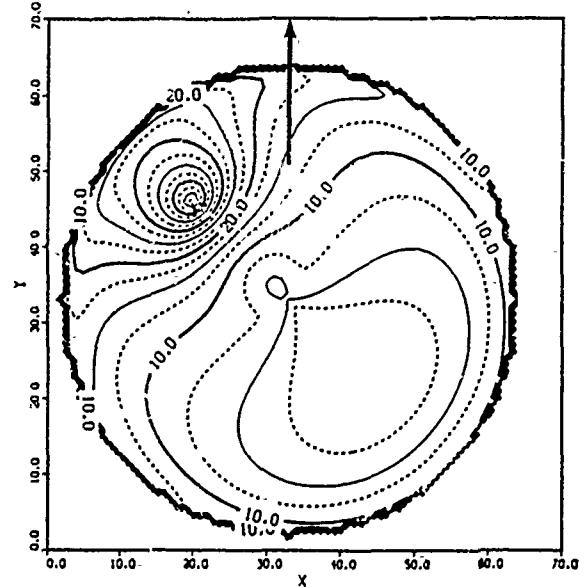
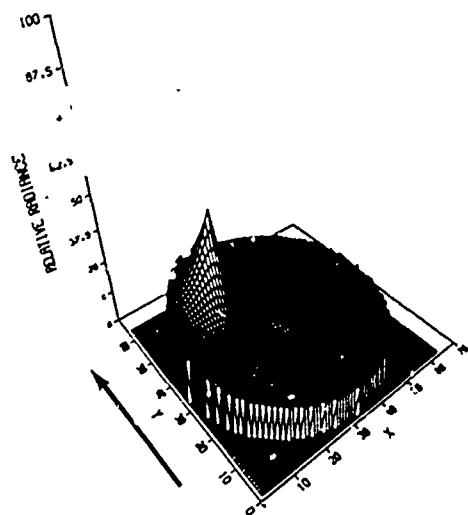
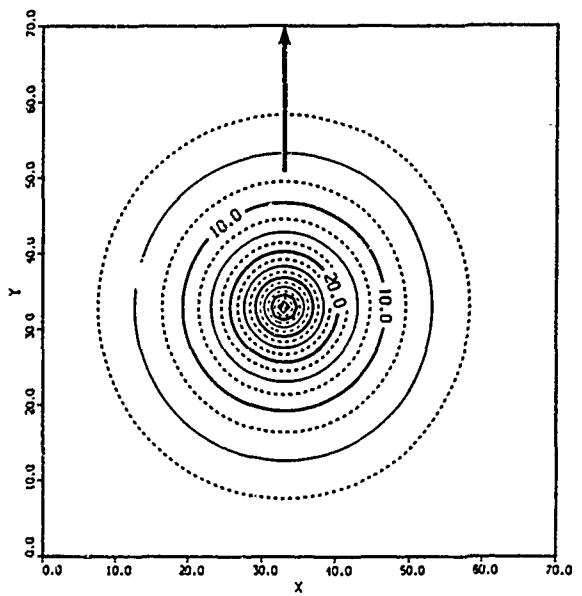
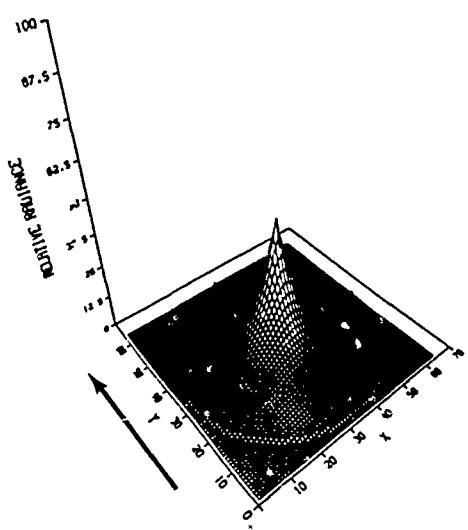
All other parameters are as previously defined:  $\nu_0$  is the angle between the vector parallel to the direction of polarization and the projection of that vector onto the surface coordinate plane.  $\psi$  includes the multiplicative depolarization term defined by Coulson [1975] to account for anisotropy of scattering.

Figures 2.3:3 and 2.3:4 illustrate the horizontally polarized component of the total radiance distributions illustrated in Figures 2.3:1 and 2.3:2.

Figures 2.3:5 and 2.3:6 illustrate the vertically polarized component of the total radiance distributions illustrated in Figures 2.3:1 and 2.3:2.



FIGURES 2.3:3-4. Horizontally polarized radiance components.



FIGURES 2.3:5-6. Vertically polarized radiance components.

## 2.4 Synthesis of Subsurface Upwelling Radiance

### Introduction

Existing linear or linearizable models for subsurface upwelling radiance employ one or more limiting assumptions:

- 1) the existence of a flat water surface,
- 2) an homogeneous subsurface profile,
- 3) a single-scattering approximation (i.e. clear or shallow water),
- 4) direct solar illumination only, and/or
- 5) diffuse sky illumination only [Aas, 1987; Gordon et al., 1975; Gordon & McCluney, 1975; Philpot, 1987; Zaneveld, 1982].

The minimum subsurface model for this synthesis must assume an homogeneous deepwater subsurface profile in order to simplify several radiometric and mechanical considerations:

- 1) the bottom is not visible and therefore not a radiometric factor,
- 2) surface waves do not mechanically refract (as they would in shallow water) and impart energy to the bottom,
- 3) the interaction of wave energy with subsurface waters is sufficient to make the inherent optical properties of the subsurface homogeneous at all radiometrically significant

depths, and

4) the subsurface is homogeneous to a sufficient optical depth in order to assume that the subsurface hemispheric radiance distribution (the 'subdome') is spatially invariant with respect to the image sample area of the water surface, i.e. the model assumes that each image sample area on the water surface 'sees' the same subsurface radiance distribution.

The calculation of upwelling subsurface radiance in the presence of a non-smooth water surface is complicated by several factors:

- 1) attenuation and refraction of the above-surface downwelling radiance by the rough water surface,
- 2) absorption and back-scattering (single or multiple) of the below-surface downwelling radiance due to the inherent optical properties of the water and its constituents,
- 3) attenuation and refraction of the below-surface upwelling radiance by the rough water surface, and
- 4) localized total internal reflection of the below-surface upwelling radiance where the angle between the surface normal vector and the upwelling radiance vector exceeds the critical angle (48.8 degrees at 460 nanometers).

A complete synthetic model for upwelling subsurface radiance would fully accommodate its parametric dependence on the four inherent optical properties of natural waters [Gordon et al., 1975]:

a = the absorption coefficient,

b = the scattering coefficient,

c = the attenuation coefficient, and

B(theta) = the volume-scattering coefficient,

with the following inter-relations:

c = b + a , and

$$b = 2\pi \int_0^{\pi} B(\theta) \sin(\theta) d\theta.$$

Two nonlinear integrations would be applied to accommodate attenuation of a non-flat water surface. The first integration would angularly distribute the refracted downwelling surface radiance based upon the slope distribution of the total surface area of the image. The second integration would angularly distribute the refracted upwelling subsurface radiance based upon the individual slope distribution of each image sample area.

In essence, at least three independent nonlinear operations would be required to compute the subsurface upwelling radiance; one intermediate operation for the single-scattering model. The only successful technique that can adequately model this level of complexity (and higher) is a Monte Carlo simulation [Plass & Kattawar, 1969], a method which is beyond the scope of this study.

#### Chapman & Irani Model

No subsurface model was defined by Chapman and Irani [1981] because they included the assumption of spectral sensing in the infrared where transmittance through water approaches zero.

#### Current Study Model

The approach for this study is to employ the results of previous Monte Carlo simulations as ad hoc analytic models. Results have been published by Plass, Kattawar and Guinn [1975, 1976] and will be used with the following assumptions:

- 1) the results are valid in a narrowband spectral region centered at 460 nanometers (which is near the wavelength of maximum transparency for clear water),
- 2) the optical depth tau of water is 10 at this wavelength,
- 3) the bottom layer has unit absorption, and

4) the results are valid over a broad band of wind velocities centered at 1030 cm/sec (measured at a height of 1950 cm). (This is equivalent to a friction velocity of 36 cm/sec for an homogeneous wind profile.)

Figures 2.4:1 through 2.4:5 illustrate the Monte Carlo results of Plass, Kattawar and Guinn [1976] for solar zenith angles of 0, 15, 32, 57, and 80 degrees, respectively.

The upwelling subsurface radiance observed just above the surface is given the following approximate analytic expression relative to  $L_{ref}$ , the reference sky radiance at zenith:

$$L_w(\beta, \alpha) = L_{ref} * R_0 * C * (1 + \cos(A * \beta))^2 / 2.0 \quad [2.4:1]$$

where

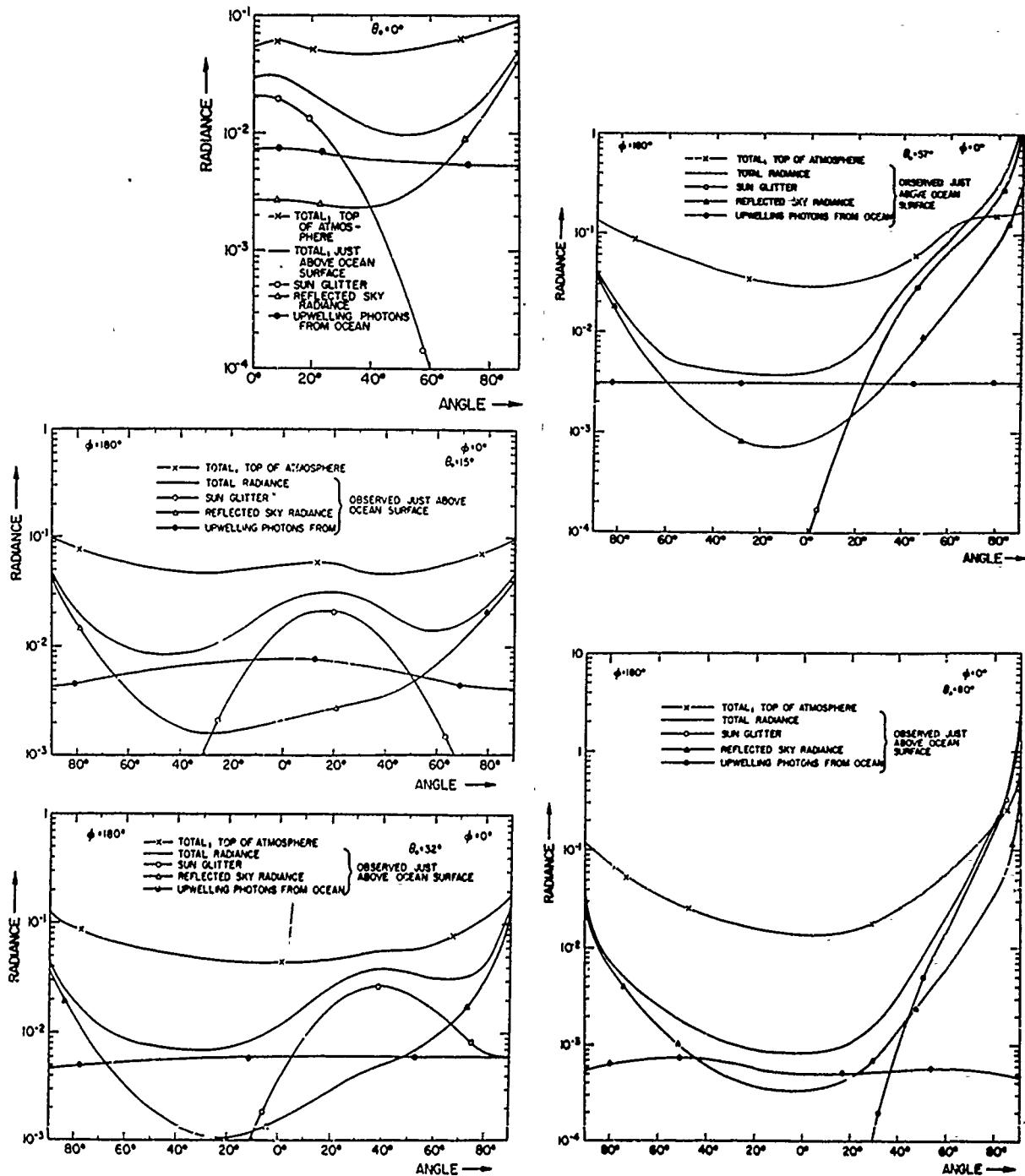
$L_w(\beta, \alpha)$  = the upwelling subsurface radiance observed at the wave-slope coordinates ( $\beta, \alpha$ ),

$\beta$  = the maximum wave-slope angle [0:+90 deg],

$\alpha$  = the azimuth angle to the wave-slope angle  $\beta$  [-180:+180 deg],

$L_{ref} = L_{sky}(0,0)$  = the reference radiance at zenith, [2.4:2]

$R_0 = 0.0208$  = the unpolarized Fresnel reflectivity at normal



FIGURES 2.4:1-5. Monte Carlo scattering results. (from Plass, Kattawar & Guinn, 1976)

incidence at 460 nanometers, [2.4:3]

$C = 0.32 \times 10^B$ , [2.4:4]

$A = \cos(\theta_0)^2$ , [2.4:5]

and

$B = \sin(1.1 \times \theta_0) \sqrt{3}$  [2.4:6]

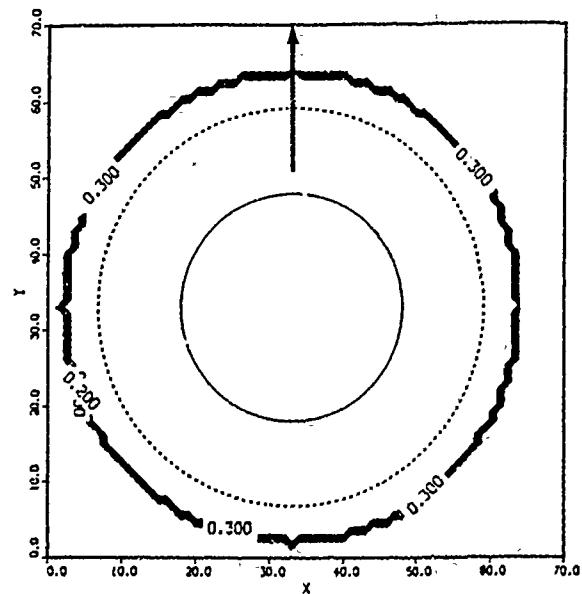
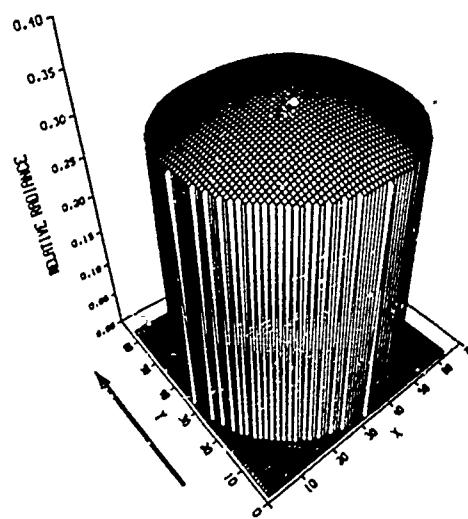
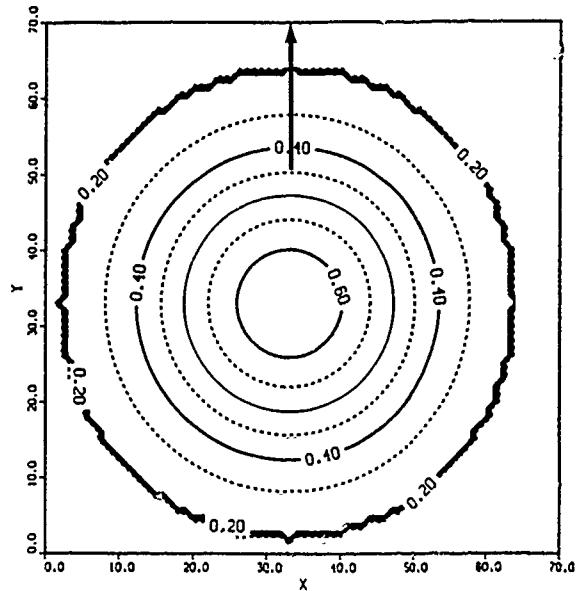
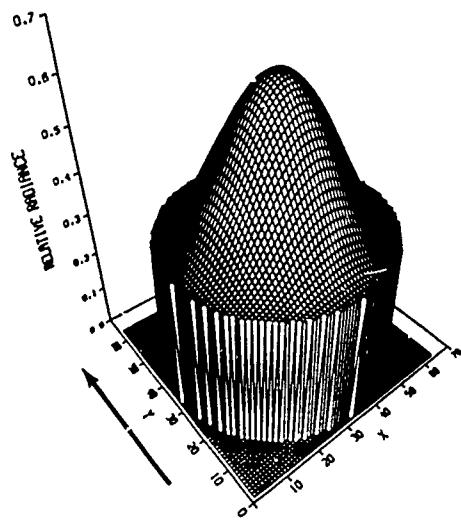
such that

$\theta_0$  = the zenith angle to the sun.

Note that this model has no dependence on  $\phi_0$ , the azimuth angle of the sun.

Again, this study requires only a reasonable relative spatial distribution of radiance and does not attempt to compute absolute radiance values. Therefore, the value of  $L_{ref}$  is set such that the radiance at the solar location is equal to 100.0 in arbitrary units of radiance. All other radiance values can then be expressed as a percentage of the total radiance in the vicinity of the sun.

Figures 2.4:6 and 2.4:7 illustrate the two synthetic total subdome radiance distributions for solar zenith angles of 0.0 and 53.2 degrees, respectively.



FIGURES 2.4:6-7. Synthetic subdome radiance distributions.

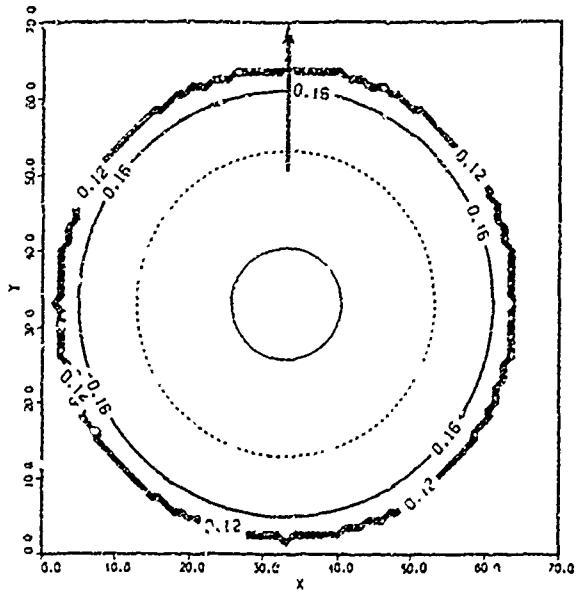
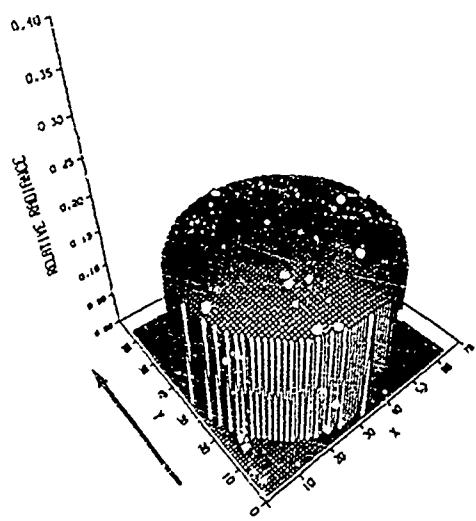
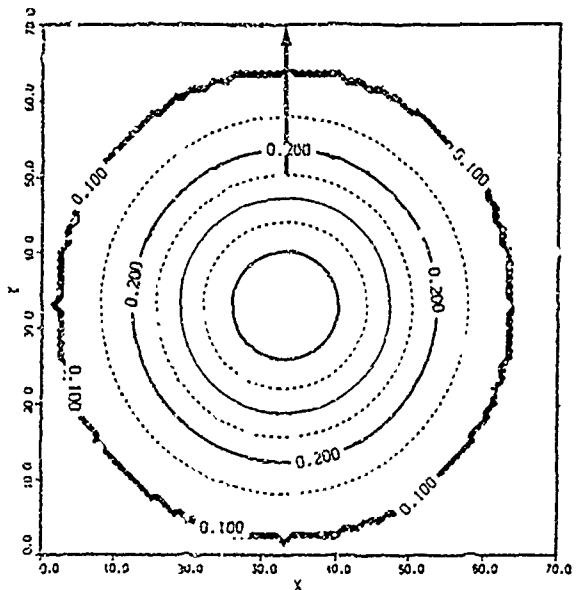
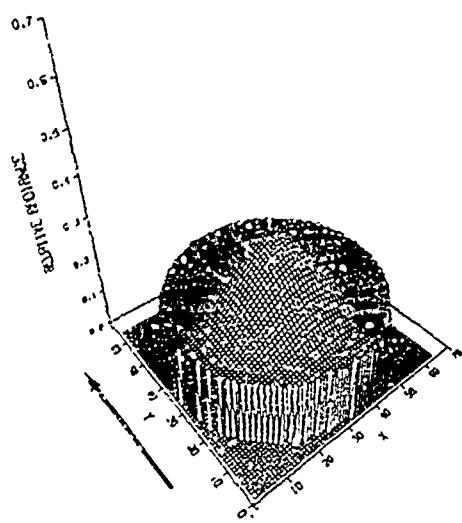
### Depolarized Random Scattering Model

Both Beckmann and Spizzichino [1963] and Konnen [1985] give evidence that the effect of rough water surfaces on the propagation of polarized downwelling radiance through the surface is a depolarization, either through multiple reflections or turbulent refraction. Therefore, this model will assume equal fractions of horizontally and vertically polarized radiance:

$$H = V = 0.5$$

[2.4:7]

Figures 2.4:8 and 2.4:9 illustrate the identical horizontally and vertically polarized components of the total upwelling subsurface (subdome) radiance distributions illustrated in Figures 2.4:6 and 2.4:7.



FIGURES 2.4;8-9. Horizontally and vertically polarized components of the synthetic subdome radiance distributions.

## 2.5 Synthesis of Reflected Skydome Radiance

### Chapman & Irani Model

Chapman and Irani [1981] incorporated a variant of the law of Malus into their calculation of reflected skydome radiance:

$$L_r = R * L_{sky}, \quad [5.2:1]$$

where

$R$  = the Fresnel reflectivity (for arbitrary polarization),  
and

$L_{sky}$  = the sky radiance (for arbitrary polarization).

Their paper does not indicate whether the attenuation due to surface-slope projection was incorporated in their model.

### Current Study Model

The radiance reflected from the water surface is calculated by the same method used by Cox and Munk [1954a] and Saunders [1967]:

Consider a small specular facet of the sea surface with area  $dA$  that is inclined from the zenith by the angle  $\beta$ . The projection of the facet onto a horizontal surface is

$$dA_h = dA \cdot \cos(\beta). \quad [2.5:2]$$

Likewise, the actual area of the sloping facet is

$$dA = dA_h \cdot \sec(\beta). \quad [2.5:3]$$

The projection of the actual facet area onto a plane normal to the incident radiance is

$$\begin{aligned} dA_i &= dA \cdot \cos(\omega) = dA \cdot \cos(\theta - \beta) \\ &= dA_h \cdot \sec(\beta) \cdot \cos(\omega), \end{aligned} \quad [2.5:4]$$

where

$\omega$  = the angle of incidence,

and

$\theta$  = the zenith angle of reflection.

Consider the probability distribution for surface slopes:

$$Pr(\beta, \alpha) = Pr(s_x, s_y) = Pr(dz/dx, dz/dy) \quad [2.5:5]$$

where

$\alpha$  = the azimuth angle to maximum slope angle  $\beta$ ,

and

$dz/dx, dz/dy$  = the slope components unique to  $\beta$  and  $\alpha$   
such that

$$dA_h = dx \cdot dy, \quad [2.5:6]$$

$$\text{TAN}(s_x) = \text{TAN}(dz/dx) = \text{TAN}(\beta)*\text{COS}(\alpha), \quad [2.5:7]$$

and

$$\text{TAN}(s_y) = \text{TAN}(dz/dy) = \text{TAN}(\beta)*\text{SIN}(\alpha). \quad [2.5:8]$$

The projection of the reflected radiance normal to the plane of reflection can then be defined in surface-slope coordinates ( $\beta, \alpha$ ) for any small, specular surface facet of slope =  $\beta(\alpha)$ :

$$\begin{aligned} L_r'(\beta, \alpha) &= L_r(\beta, \alpha)*\text{COS}(\theta) \\ &= R(\omega)*L_{\text{sky}}(\mu, \nu)*\text{COS}(\omega)*\text{SEC}(\beta), \end{aligned} \quad [2.5:9]$$

where

$R(\omega)$  = the Fresnel reflectivity of the water facet (for arbitrary polarization),

and

$L_{\text{sky}}(\mu, \nu)$  = the incident radiance of the skydome from the direction  $(\mu, \nu)$  that is reflected normal to the defined plane of reflection ( $\theta = 53.2$  deg,  $\phi = 180$  deg) for both polarizations. (The calculation of the incident vector coordinates from the reflection and surface normal vector coordinates is defined at the end of this chapter.)

The mean reflection radiance perceived by an observer at the reflection coordinates is similarly defined:

$$\begin{aligned}
 <L'> &= <L_r> * \cos(\theta) \\
 &= \int \int R(\omega) * L_{\text{sky}}(\beta, \alpha) * P_r(\beta, \alpha) \\
 &\quad * \cos(\omega) * \text{SEC}(\beta) * d\beta d\alpha \quad [2.5:10]
 \end{aligned}$$

over the full distribution of surface slopes. This value, combined with the small contribution due to refraction, is computed as the central ordinate of the radiance magnitude spectrum,  $M(0,0)$ , as an ensemble average of spectra. Because the reflection vector (i.e. line of sight) is fixed with respect to the other geometries,  $L_r$  can be linearly mapped to surface-slope coordinates  $(\beta, \alpha)$  once the skydome radiance distribution is uniquely determined.

Figure 2.5:1 illustrates the mapping of the function,  $\text{SEC}(\beta)$ , to the surface-slope coordinates  $(\beta, \alpha)$ . For this illustration, the function was truncated at  $\text{SEC}(60 \text{ deg})$  for values of  $\beta$  greater than 60 degrees. Note also that the function is not computed for values of  $\beta$  where  $(\omega - \mu) = (\theta - \omega) > 90$  degrees, i.e. the regions where surface slopes are not visible with respect to the line of sight.

Figure 2.5:2 illustrates the mapping of the function,  $\cos(\omega)$ , to the surface-slope coordinates  $(\beta, \alpha)$ .

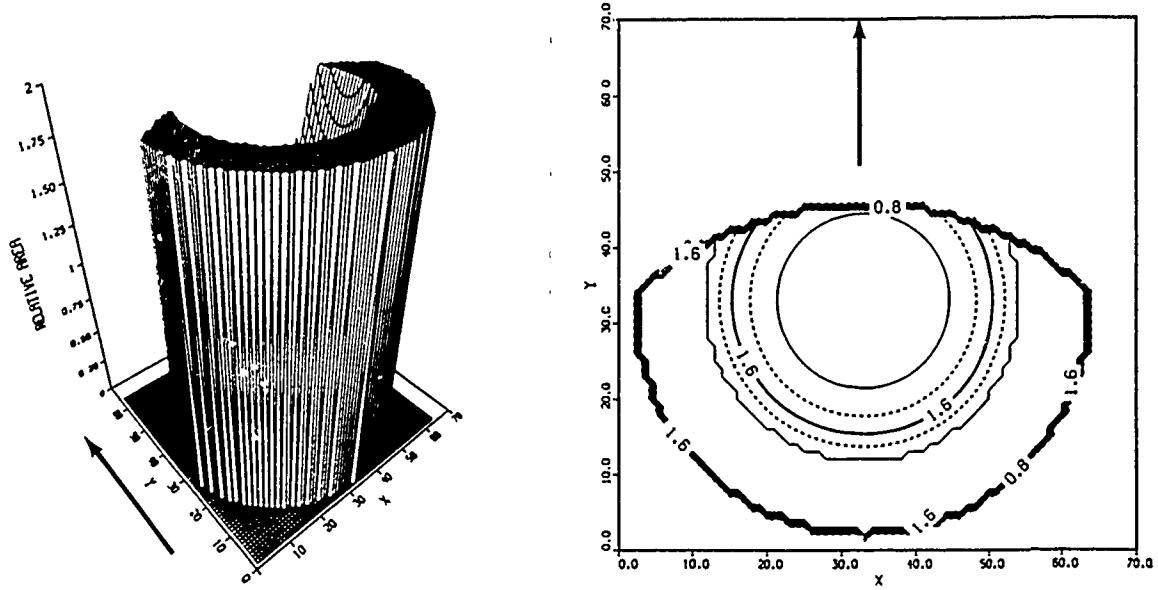


FIGURE 2.5:1. Map of  $\text{SEC}(\beta)$  to coordinates  $(\beta, \alpha)$ .

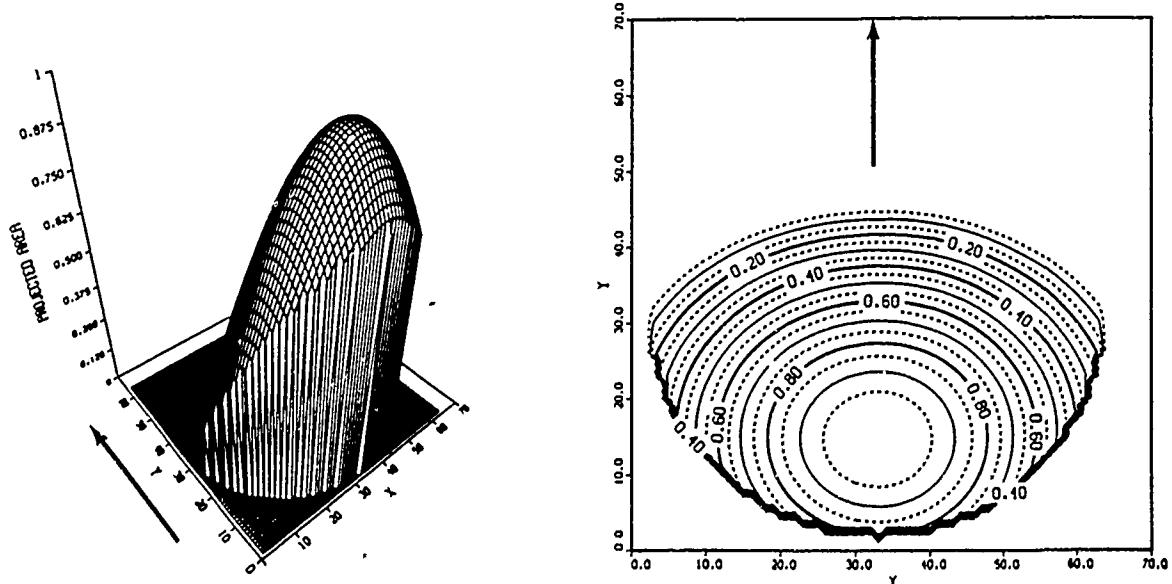


FIGURE 2.5:2. Map of  $\text{COS}(\omega)$  to coordinates  $(\beta, \alpha)$ .

Figures 2.5:3 and 2.5:4 illustrate the mapping of Fresnel reflectivity,  $R(\omega)$ , to surface-slope coordinates  $(\beta, \alpha)$  for horizontal and vertical polarizations, respectively.

Figures 2.5:5 and 2.5:6 illustrate the mapping of skydome radiance,  $L_{sky}(\mu, \nu)$ , to surface-slope coordinates  $(\beta, \alpha)$  for horizontal and vertical polarization, respectively. The solar zenith angle  $\theta_0$  in this example is 0 degrees. The original  $L_{sky}$  distributions in skydome coordinates  $(\mu, \nu)$  are illustrated in Figures 2.3:3 and 2.3:4.

Figures 2.5:7 and 2.5:8 illustrate the mapping of the reflected skydome radiance,  $L_r'$ , to the surface-slope coordinates  $(\beta, \alpha)$  for horizontal and vertical polarization, respectively. The solar zenith angle  $\theta_0$  in this example is 0 degrees.

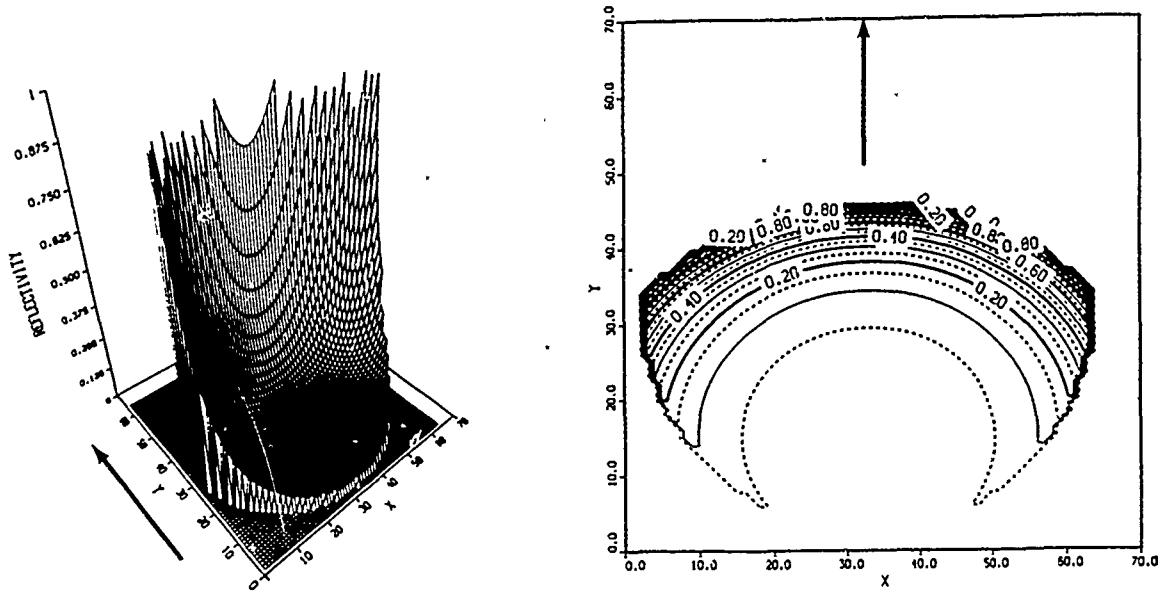


FIGURE 2.5:3. Map of Fresnel reflectivity,  $R(\omega)$ , to coordinates  $(\beta, \alpha)$  for horizontal polarization.

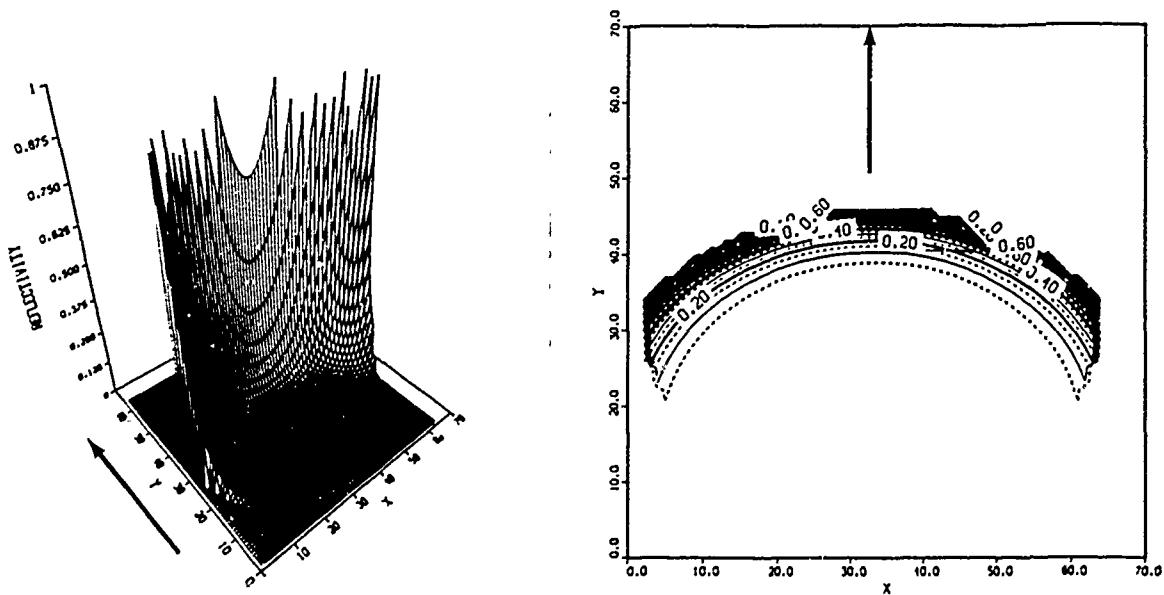


FIGURE 2.5:4. Map of Fresnel reflectivity,  $R(\omega)$ , to coordinates  $(\beta, \alpha)$  for vertical polarization.

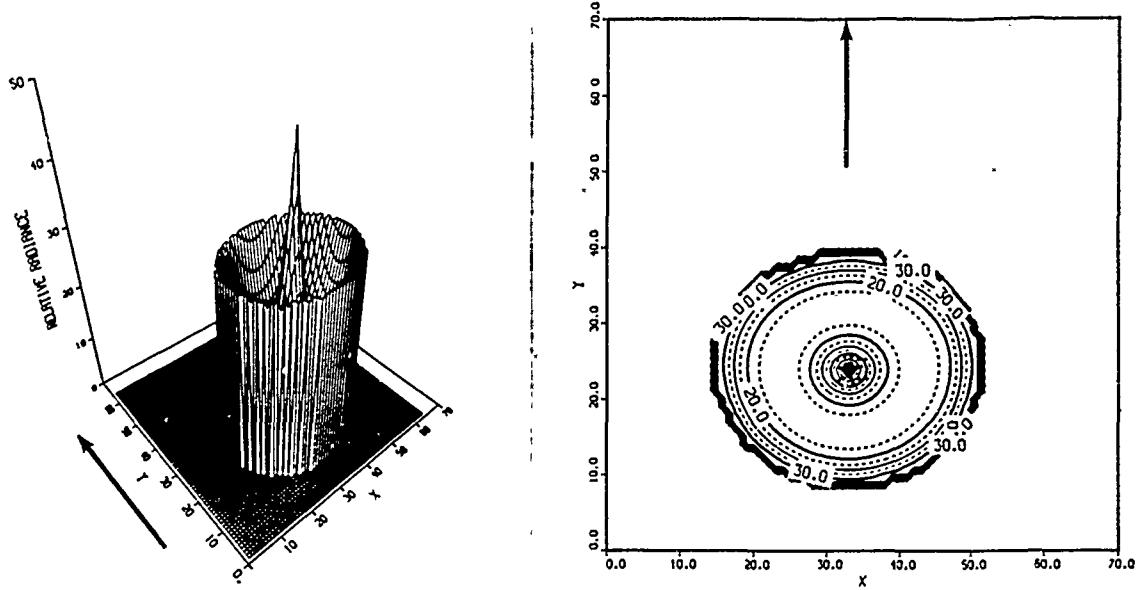


FIGURE 2.5:5. Map of skydome radiance,  $L_{sky}(\mu, \nu)$ , to coordinates (beta, alpha) for horizontal polarization.

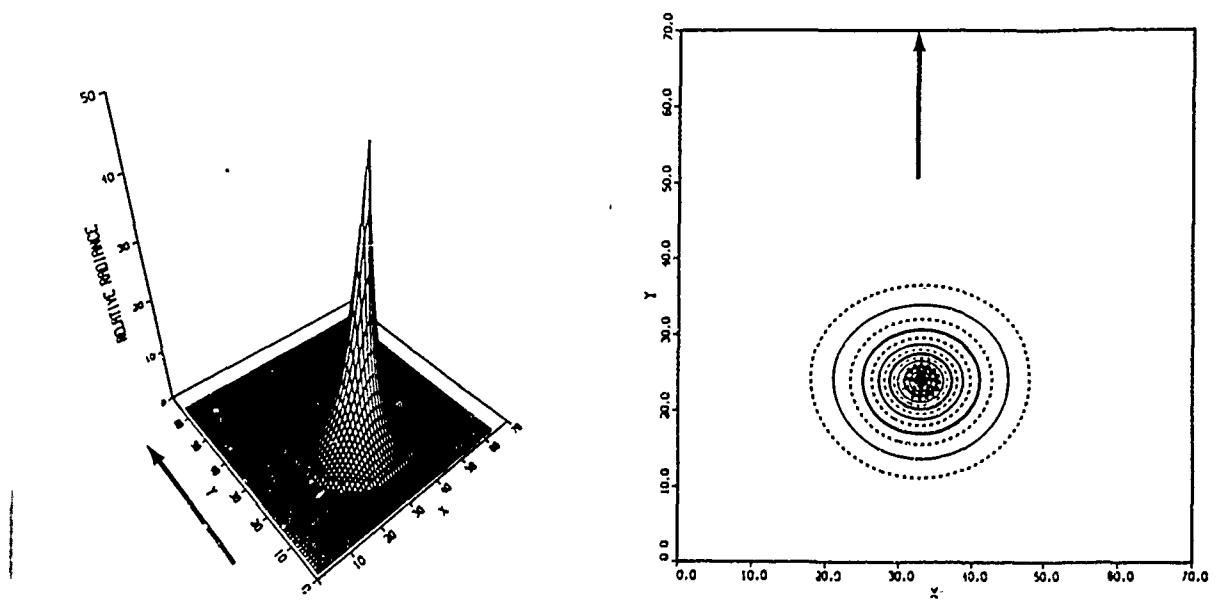


FIGURE 2.5:6. Map of skydome radiance,  $L_{sky}(\mu, \nu)$ , to coordinates (beta, alpha) for vertical polarization.

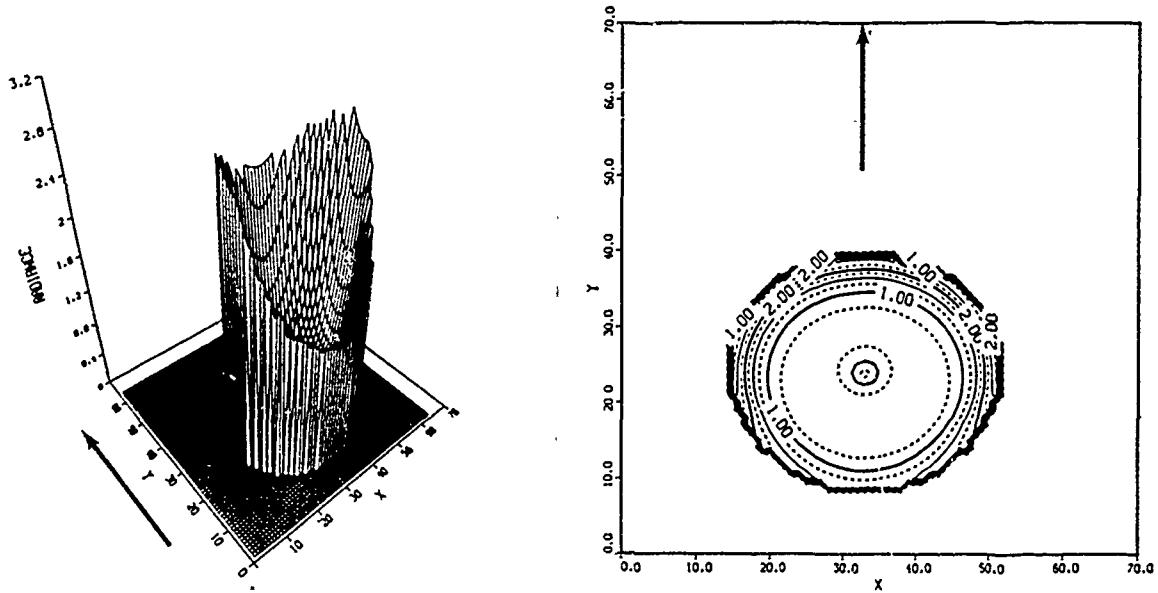


FIGURE 2.5:7. Map of reflected skydome radiance,  $L_r'$ , to coordinates (beta, alpha) for horizontal polarization.

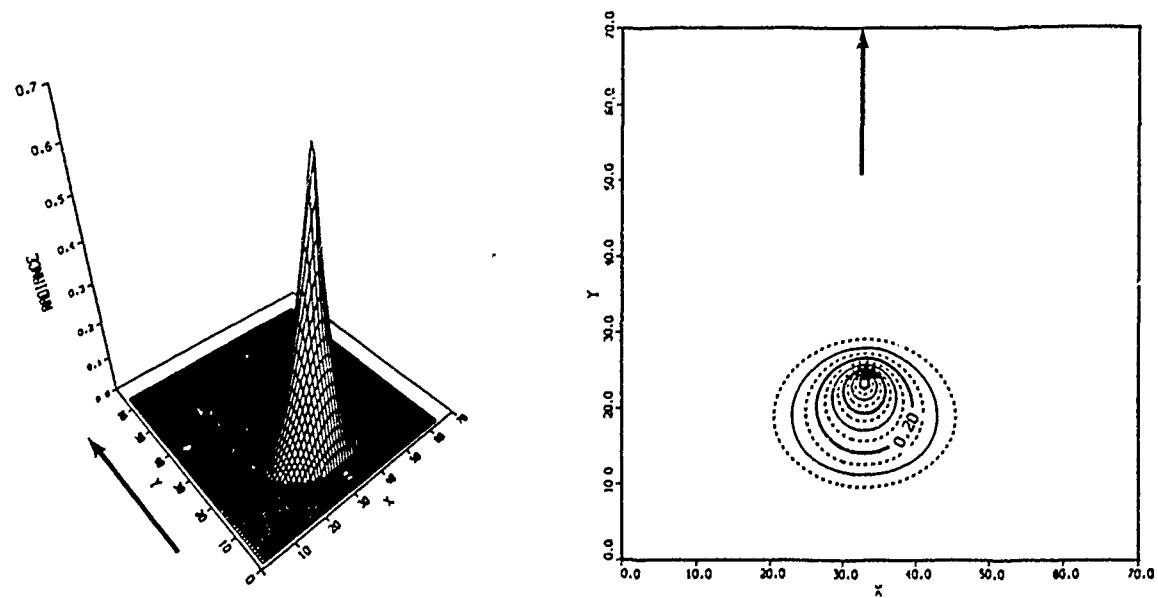


FIGURE 2.5:8. Map of reflected skydome radiance,  $L_r'$ , to coordinates (beta, alpha) for vertical polarization.

### Correlation of Sub-Resolution Wave-Slope Variance

The maximum wavenumber of the current study model is  $k_{max} = 1.2$  per centimeter to coincide with the results of other studies [Cox & Munk, 1954a,b; Pierson & Stacy, 1973]. This corresponds to a sample surface area of

$$A_h = 2.6 \text{ cm} \times 2.6 \text{ cm} = 6.76 \text{ cm}^2 \quad [2.5:11]$$

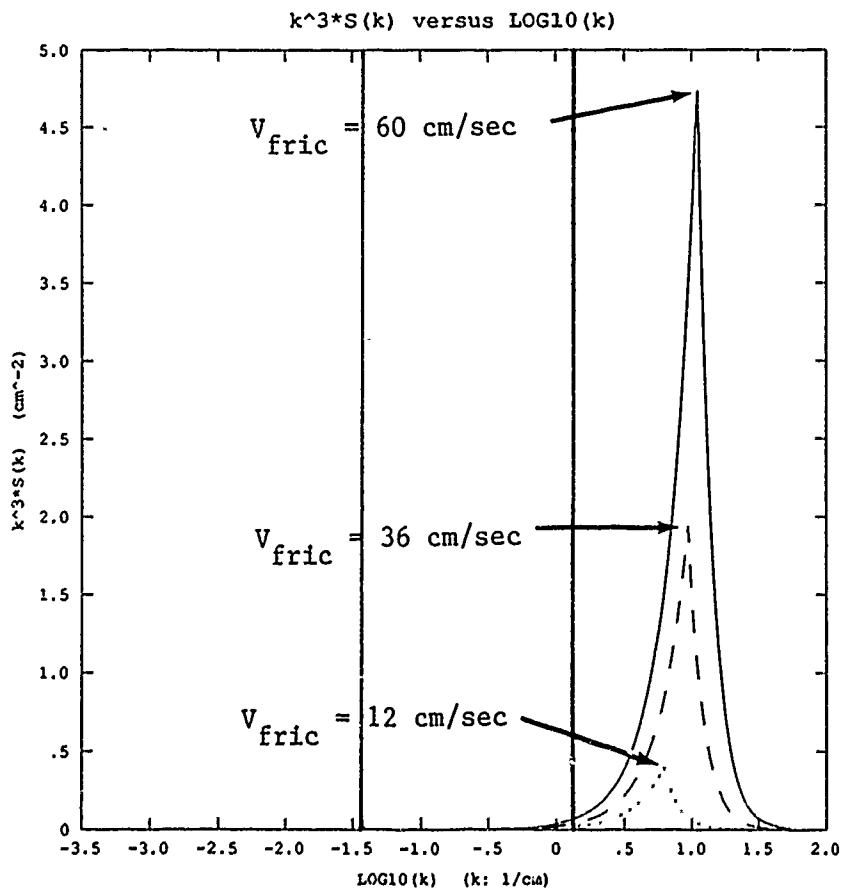
which is defined in the literature as the Cox-Munk window.

Because slope spectra (cite Figure 2.2:13) reveal significant slope energy between  $k = 1.2/\text{cm}$  and  $k = 30 \text{ cm}^{-1}$ , it is important to model the effect of this sub-resolution variance on the reconstruction of wave-slope spectra. However, it is unreasonable to synthesize large surfaces with sampling windows of the order  $A_h = 1.0 \text{ mm}^2$ , as required to consider  $k = 30$  per centimeter. Previous investigators [Chapman & Irani, 1981; Wilf & Manor, 1984; Schwartz & Hon, unpub.] all employed the Kirchoff approximation for reflection from rough surfaces [Kajiya, 1985; Beckmann & Spizzichino, 1963], which essentially replaces a rough surface facet by its mean tangent plane. This approximation is valid only if the rough sub-resolution surface is of low curvature. However, for wind-driven water-wave surfaces, it

is exactly the spectral region above  $k > 1.0$  per centimeter where all the spectral energy of curvature exists.

Figure 2.5:9 illustrates  $k^3 * S(k)$  versus  $k$ , the curvature power spectrum (after Pierson & Stacy, 1973). It is apparent that the Kirchoff approximation is not optimum for this particular circumstance.

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FIGURE 2.5:9. The 1D Pierson-Stacy curvature power spectrum  $k^3 * S(k)$  versus  $\text{LOG10}(k)$ .

### Current Study Model

To quote Stilwell and Pilon:

"The subresolution waves, especially those in the capillary regime, can have significant slopes. Since they are unresolved on the film, the intensity variations of the sky are partially smoothed. This effect can be observed by examining the reflection of a cloud in a ruffled water surface. As a result the small waves can smooth the sky luminance distribution and allow the use of photographic procedures on days in which the sky is far from a monotonic luminance function. Contingent only on the distribution of sub-resolution waves being homogeneous, the effective luminance distributions may extend the environmental conditions under which optical transforms are useful."

[Stilwell & Pilon, 1974]

An alternative method, then, for introducing the effect of sub-resolution wave slopes is to assume an homogeneous distribution and use it as a correlation filter to smooth the reflected skydome radiance. The correlation is nonlinear if executed in reflected skydome coordinates ( $\mu, \nu$ ). However, a linear correlation is possible if executed in surface slope coordinates ( $\beta, \alpha$ ).

The unresolved directional wave-slope variance for the wavenumber range  $k > 1.2/\text{cm}$  is calculated directly from the Pierson-Stacy model. A directional bivariate normal Gaussian distribution is calculated as a function of along-wind slope variance, cross-wind slope variance, and wind azimuth:

$$Pr'(\beta, \alpha) = \text{EXP}(-(L^2+M^2)/2)/(2\pi\text{DEV}_A\text{DEV}_C), \quad [2.5:12]$$

where

$$L = \text{TAN}(\beta)\text{COS}(\text{wind\_az}-\alpha)/\text{DEV}_A, \quad [2.5:13]$$

and

$$M = \text{TAN}(\beta)\text{SIN}(\text{wind\_az}-\alpha)/\text{DEV}_C, \quad [2.5:14]$$

such that

$$\text{DEV}_A = \text{SQRT}(\text{along-wind slope variance}),$$

$$\text{DEV}_C = \text{SQRT}(\text{cross-wind slope variance}),$$

and

$\text{wind\_az}$  - the wind azimuth (with 180-degree ambiguity).

The directional distribution of sub-resolution wave slopes,  $Pr'(\beta, \alpha)$ , is then correlated with the reflected sky-dome radiance,  $L_r'(\beta, \alpha)$ ,

$$L_r''(\beta, \alpha) = \int \int L_r'(b, a) * Pr'(b-\beta, a-\alpha) * da db \\ = L_r'(\beta, \alpha) \star \star Pr'(\beta, \alpha), \quad [2.5:15]$$

which follows directly from Equations [2.5:9] and [2.5:10].

(The  $\star \star$  denotes the double correlation operation.)

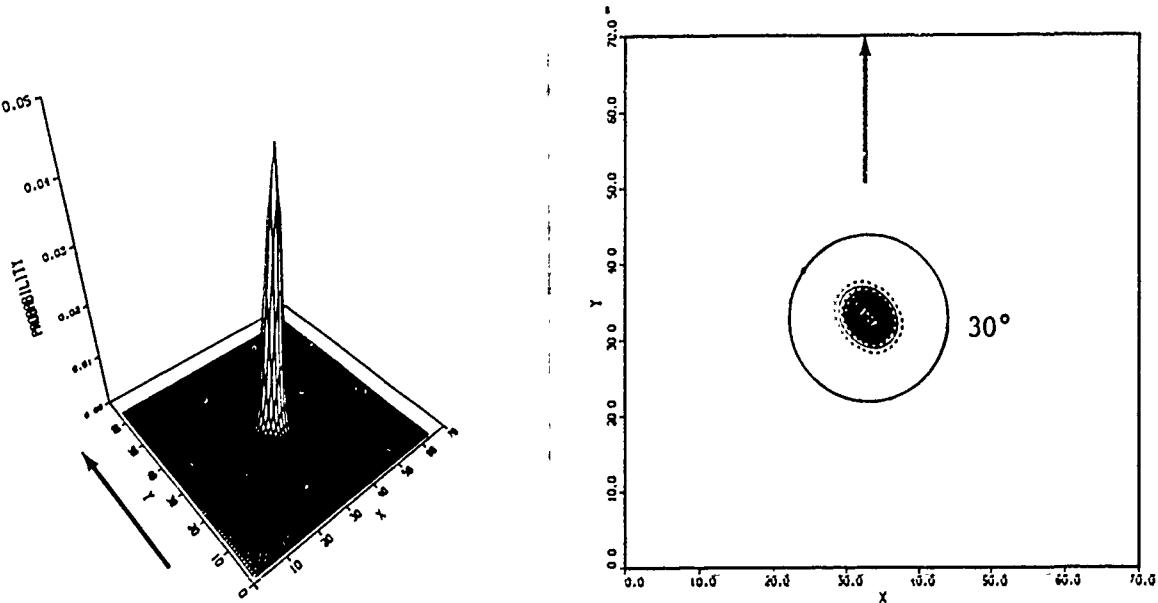
Note that because the sub-resolution wave-slope distribution is assumed to be both homogeneous and symmetric about its central ordinate, Equation [2.5:15] could also be implement-

ed as a double convolution operation:

$$\begin{aligned} L_r''(\beta, \alpha) &= L_r'(\beta, \alpha) \star \star P_r'(\beta, \alpha) \\ &- L_r'(\beta, \alpha) \bullet \bullet P_r'(\beta, \alpha). \end{aligned} \quad [2.5:16]$$

Figure 2.5:10 is the sub-resolution wave-slope distribution for a friction velocity of 12.0 cm/sec and a wind azimuth of -45 deg (or +135 deg with ambiguity).

Figures 2.5:11 and 2.5:12 illustrate the correlation of  $L_r'$ , the reflected skydome radiance (Figures 2.5:7 and 2.5:8), with  $P_r'$ , the sub-resolution wave-slope distribution.



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FIGURE 2.5:10. Sub-resolution wave-slope distribution  $Pr'$  for  $v_{fric} = 12.0$  cm/sec and  $wind\_az = -45$  deg.

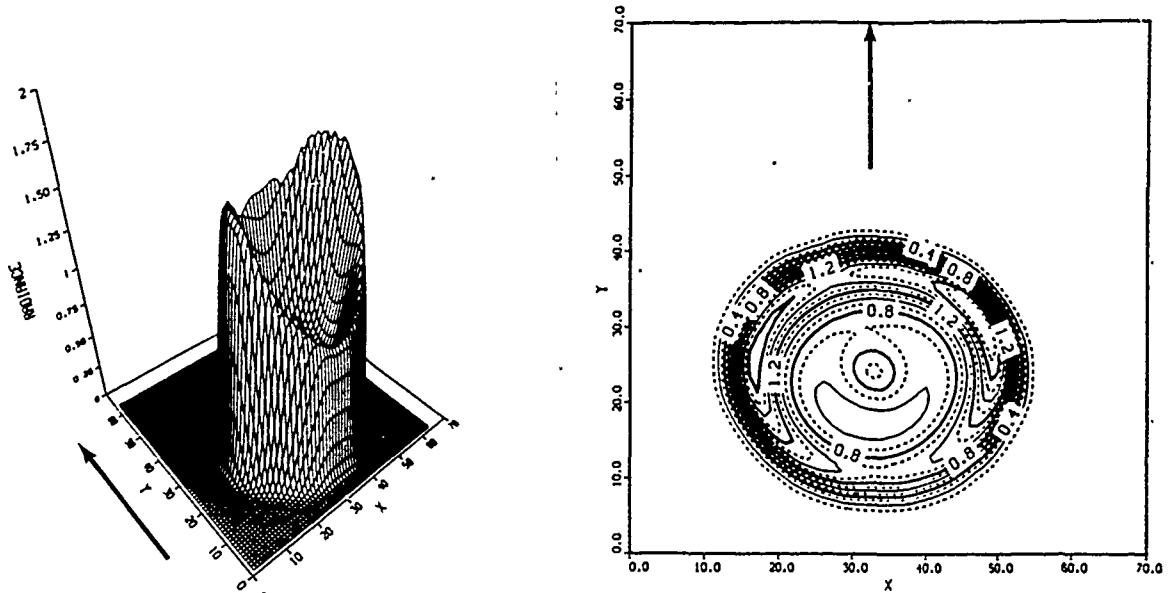


FIGURE 2.5:11. Effect of correlation of  $L_x'$  with sub-resolution wave-slope distribution  $Pr'$  (horizontal polarization).

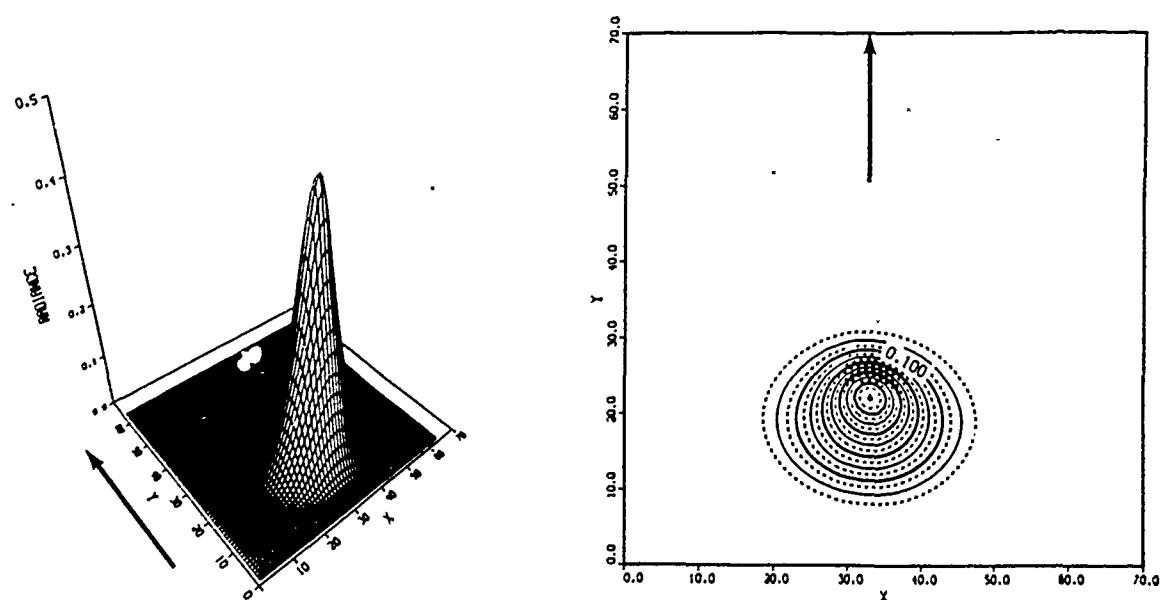


FIGURE 2.5:12. Effect of correlation of  $L_x'$  with sub-resolution wave-slope distribution  $Pr'$  (vertical polarization).

### Calculation of Reflection Angular Coordinates

The angular coordinates that describe the ray path due to reflection can be calculated if two of three vectors (the surface normal vector  $N$ , the incident vector  $I$ , and the reflection vector  $R$ ) are known.

Figure 2.5:13 illustrates the geometric relations and nomenclature for the current study model.

The reflection vector is fixed in this model. The observer is located at an azimuth angle  $\phi$  of 180 degrees and zenith angle  $\theta$  of 53.2 degrees (the Brewster angle) relative to the point of observation. The reflection vector  $R$  can be resolved into its unit vector coordinates in  $(x,y,z)$  space:

$$R_x = \text{SIN}(\theta) * \text{COS}(180 \text{ deg}), \quad [2.5:17]$$

$$R_y = \text{SIN}(\theta) * \text{SIN}(180 \text{ deg}), \quad [2.5:18]$$

and

$$R_z = \text{COS}(\theta). \quad [2.5:19]$$

The surface normal vector  $N$  and the incident vector  $I$  can be similarly resolved:

$$N_x = \text{SIN}(\beta) * \text{COS}(\alpha), \quad [2.5:20]$$

$$N_y = \sin(\beta) * \sin(\alpha), \quad [2.5:21]$$

and

$$N_z = \cos(\beta); \quad [2.5:22]$$

$$R_x = \sin(\mu) * \cos(\nu), \quad [2.5:23]$$

$$R_y = \sin(\mu) * \sin(\nu), \quad [2.5:24]$$

and

$$R_z = \cos(\mu). \quad [2.5:25]$$

Because the vector  $N$  must be in the plane containing  $I$  and  $R$  and because it must also bisect the angle between  $I$  and  $R$ , the following equations have unique solutions:

$$IN = RN = \cos(\omega), \quad [2.5:26]$$

$$IR = \cos(2\omega), \quad [2.5:27]$$

$$N = (R - I)/(2\cos(\omega)), \quad [2.5:28]$$

and

$$R = N*2*\cos(\omega) - I. \quad [2.5:29]$$

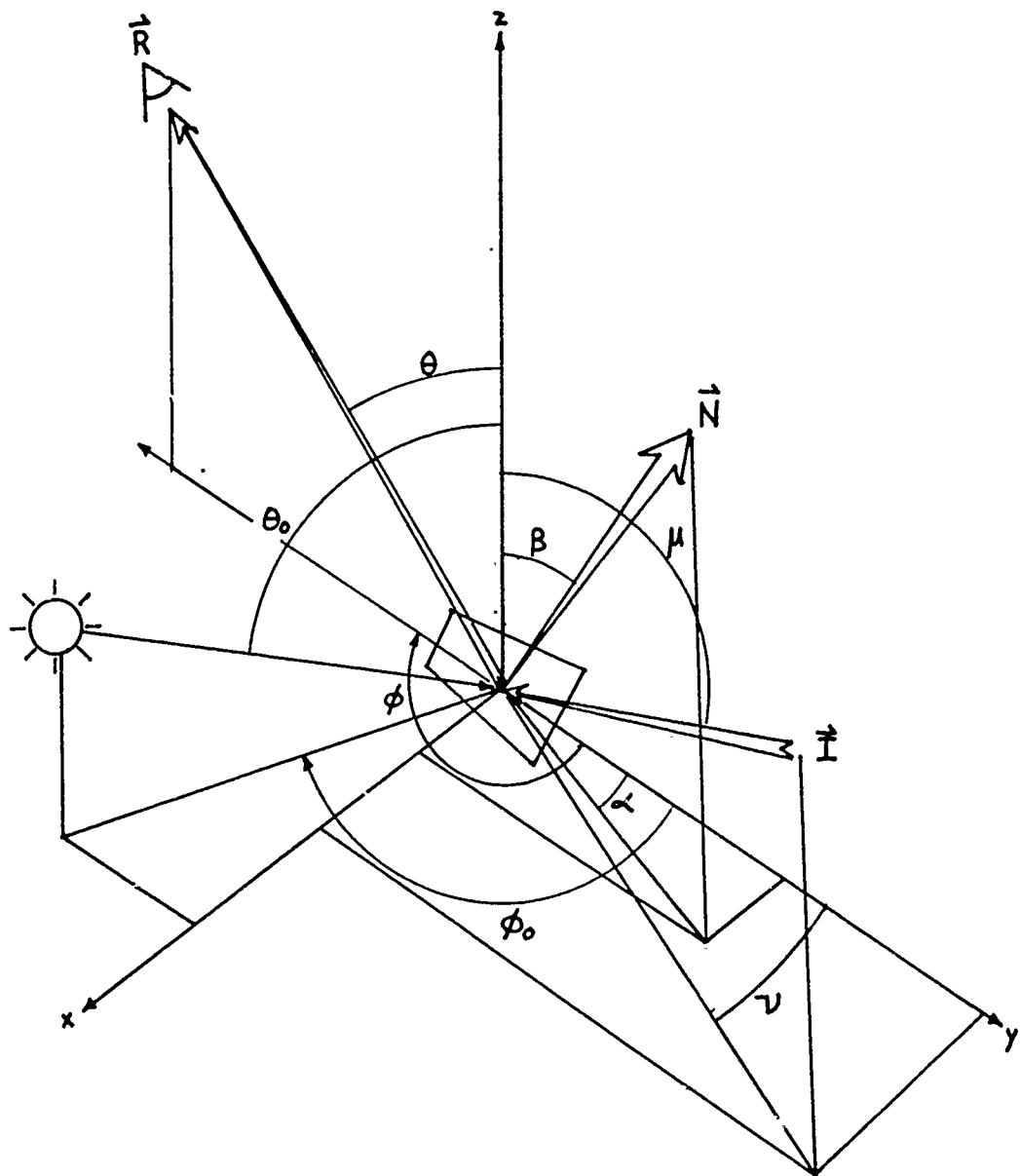


FIGURE 2.5:13. 3D geometric relations and nomenclature for the current study model.

## 2.6 Synthesis of Refracted Subsurface Radiance

### Chapman & Irani Model

No refracted subsurface radiance model was defined by Chapman and Irani [1981] because they included the assumption of spectral viewing in the infrared where the transmittance of water approaches zero.

### Current Study Model

As defined in Chapter 2.4, the upwelling subsurface radiance model of Plass et al. [1976] provides the radiance just above the ocean surface as a function of zenith angle of observation. Therefore, no angular refraction model is required. Only attenuation due to surface projection (as described in Chapter 2.5) needs to be considered:

The projection of the refracted radiance normal to the plane of observation can be defined, in surface-slope coordinates ( $\beta$ , $\alpha$ ) for any small, specular surface facet of slope:

$$L'_w(\beta, \alpha) = L_w(\beta, \alpha) * \cos(\omega) * \sec(\beta) \quad [2.6:1]$$

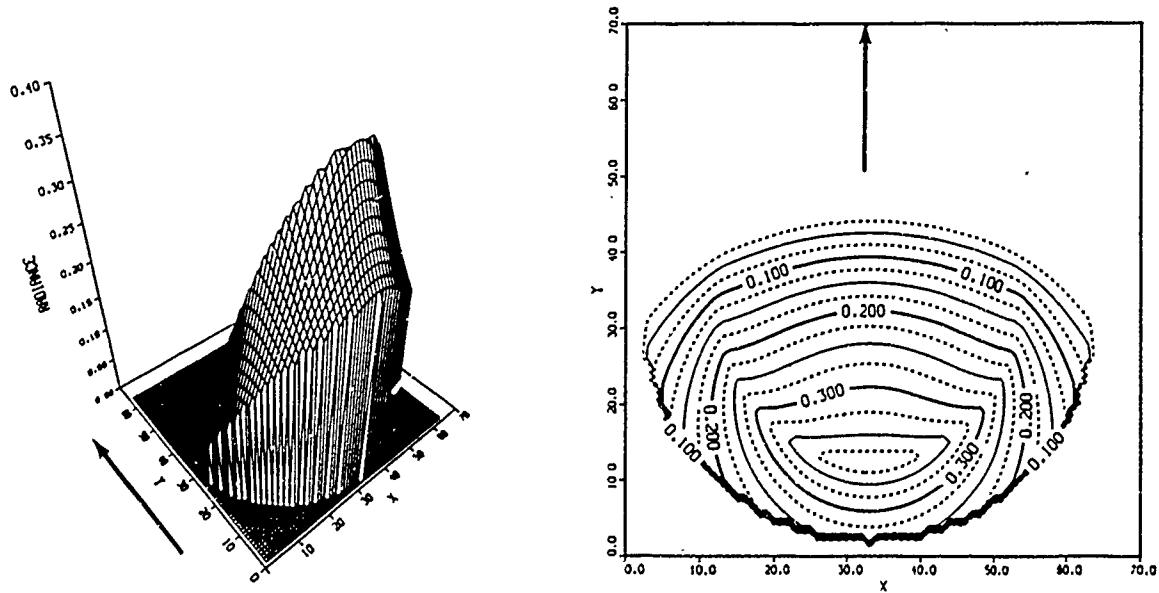
where

$L_w(\beta, \alpha)$  = the refracted subsurface radiance observed at wave-slope coordinates ( $\beta, \alpha$ ),  
and

$\omega$  is as previously defined.

Figure 2.6:1 illustrates the mapping of  $L_w'$ , the refracted subsurface radiance, to the surface-slope coordinates ( $\beta$ ,  $\alpha$ ) for identical horizontal and vertical polarizations, respectively. The solar zenith angle in this example is 0 degrees. The original unattenuated upwelling subsurface radiance distribution,  $L_w$ , is illustrated in Figure 2.4:8.. Note that the function, SEC( $\beta$ ), was truncated for values of  $\beta$  greater than 60 degrees (as in Figure 2.5:1), in order to maintain the scale of the illustration.

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FIGURE 2.6:1. Map of refracted subsurface radiance,  $L_w'$ , to coordinates ( $\beta, \alpha$ ) for both polarizations.

### Correlation of Sub-Resolution Wave-Slope Variance

No correlation of the refracted radiance is required for the current study model since the Monte Carlo simulation of Plass et al. [1976] has already incorporated the effect of wave slopes.

## 2.7 Generation of Synthetic Radiance Imagery

The integrated current study model can generate any combination of these synthetic radiance conditions:

- 1) Horizontal, Vertical polarization, or Both (unpolarized),
- 2) Reflected, Refracted radiance, or Both, and
- 3) Sub-resolution wave-slope attenuation (On or Off).

Because of Brewster-angle viewing relative to the mean sea surface, an initial assumption of the current study is that the effect of vertically polarized radiance is at a minimum and a comparison of horizontally polarized radiance spectra with unpolarized radiance spectra is sufficient to demonstrate the effect of vertical polarization. A second assumption is that the effect of upwelling subsurface refracted radiance is small relative to reflected radiance and a comparison of reflected radiance spectra with combined (reflected + refracted) radiance spectra is sufficient to demonstrate the effect of refracted radiance. A third assumption is that horizontally polarized, reflected radiance spectra synthesized without the sub-resolution wave-slope correlation filter will emulate the synthetic results of Chapman and Irani [1981].

The following radiance images (and their magnitude spectra) are synthesized in support of the current study:

$H1(x,y)$  = Horizontally polarized, reflected radiance image,

$L1(x,y)$  = Total (unpolarized), reflected radiance image,

$H0(x,y)$  = Horizontally polarized, reflected and refracted radiance image,

$L0(x,y)$  = Total (unpolarized), reflected and refracted radiance image, and

$H3(x,y)$  =  $H1(x,y)$  without the incorporation of the sub-resolution wave-slope correlation filter model.

The following pages illustrate individual realizations of synthetic images that were generated in support of the current study. All examples are  $L0$  images with contrasts normalized relative to a discrete 32-level gray scale: black corresponds to zero radiance and white corresponds to the maximum pixel radiance of the image. Only these examples are discrete: the actual syntheses are real scalar arrays.

Figures 2.7:1 through 2.7:3 illustrate synthetic images generated by filtering the same white noise array with slope spectra for  $v_{fric} = 12.0, 36.0$ , and  $60.0$  cm/sec, respectively.  $wind\_az = 135$  deg (-45 deg) and sun is at zenith.

FIGURE 2.7:1. L0 image.

vfric = 12.0 cm/sec  
Max pixel radiance L0 = 1.526  
Avg pixel radiance H1 = 0.578  
Avg pixel radiance H3 = 0.528  
Avg pixel radiance V1 = 0.032  
Avg pixel radiance L1 = 0.610  
Avg pixel radiance H2 = 0.106  
Avg pixel radiance V2 = 0.106  
Avg pixel radiance L2 = 0.212  
Avg pixel radiance H0 = 0.684  
Avg pixel radiance V0 = 0.138  
Avg pixel radiance L0 = 0.822

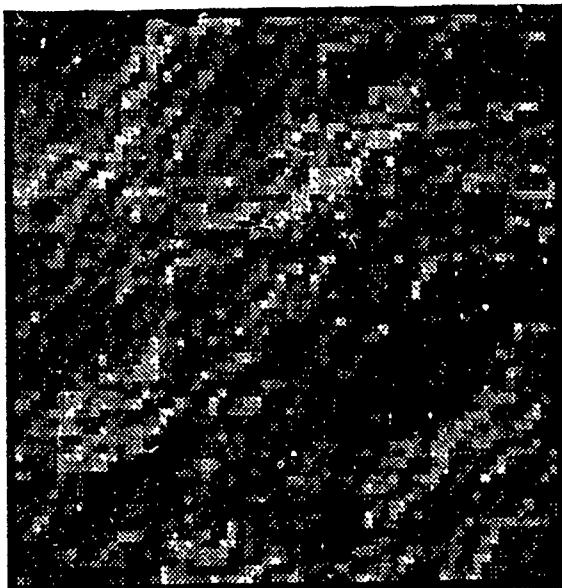


FIGURE 2.7:2. L0 image.

vfric = 36.0 cm/sec  
Max pixel radiance L0 = 1.894  
Avg pixel radiance H1 = 0.602  
Avg pixel radiance H3 = 0.564  
Avg pixel radiance V1 = 0.050  
Avg pixel radiance L1 = 0.652  
Avg pixel radiance H2 = 0.105  
Avg pixel radiance V2 = 0.105  
Avg pixel radiance L2 = 0.210  
Avg pixel radiance H0 = 0.707  
Avg pixel radiance V0 = 0.155  
Avg pixel radiance L0 = 0.862

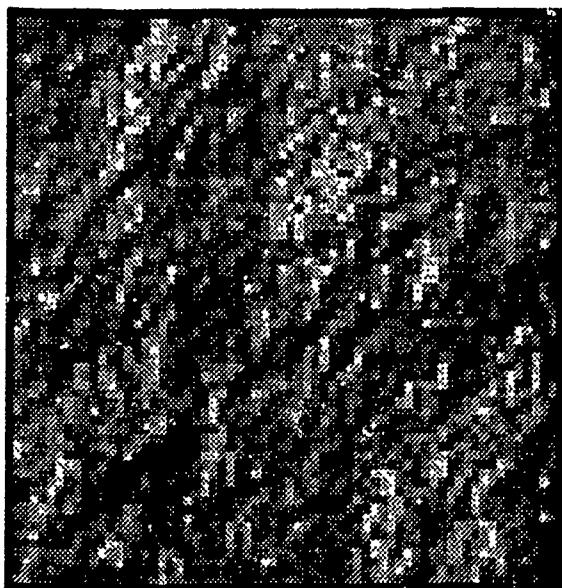
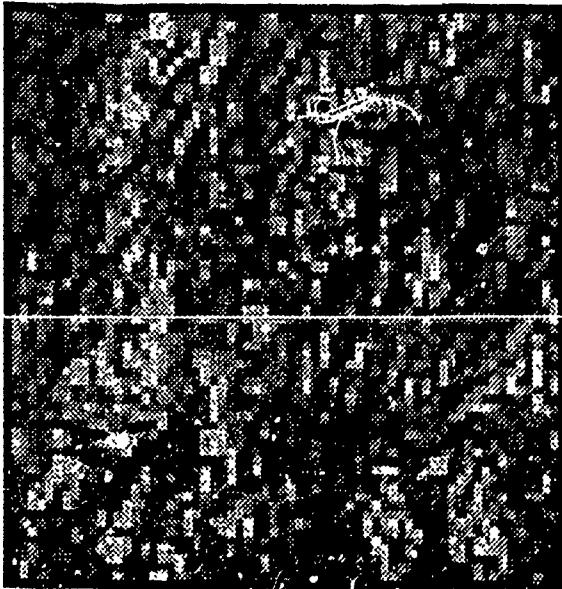


FIGURE 2.7:3. L0 image.

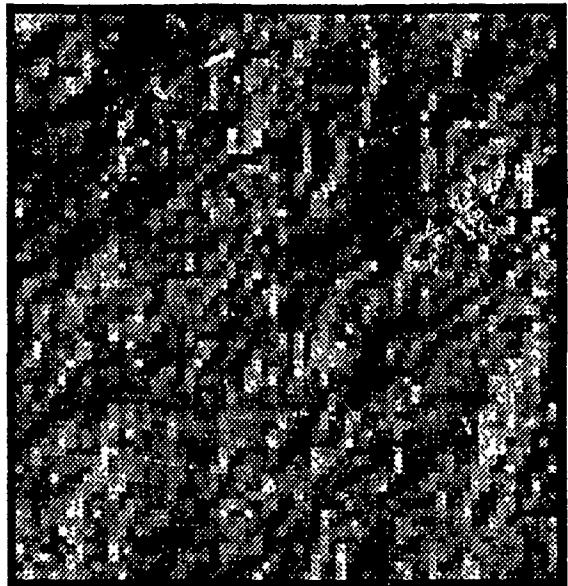
vfric = 60.0 cm/sec  
Max pixel radiance L0 = 1.737  
Avg pixel radiance H1 = 0.579  
Avg pixel radiance H3 = 0.574  
Avg pixel radiance V1 = 0.054  
Avg pixel radiance L1 = 0.633  
Avg pixel radiance H2 = 0.105  
Avg pixel radiance V2 = 0.105  
Avg pixel radiance L2 = 0.210  
Avg pixel radiance H0 = 0.684  
Avg pixel radiance V0 = 0.159  
Avg pixel radiance L0 = 0.843



The synthetic image in Figure 2.7:4 is a different synthetic realization (i.e. different white noise spectrum) using the same parameters as Figure 2.7:1. Note both the similarity of the image and its statistics to that of Figure 2.7:1.

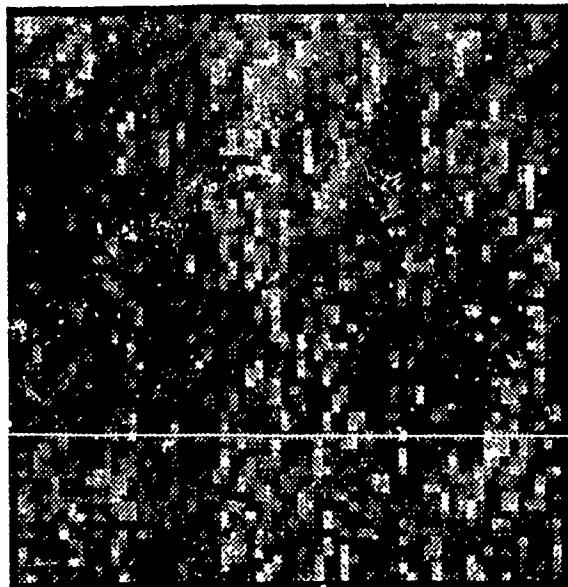
FIGURE 2.7:4. L0 image.

```
vfric = 12.0 cm/sec
Max pixel radiance L0 = 1.523
Avg pixel radiance H1 = 0.578
Avg pixel radiance H3 = 0.526
Avg pixel radiance V1 = 0.031
Avg pixel radiance L1 = 0.609
Avg pixel radiance H2 = 0.106
Avg pixel radiance V2 = 0.106
Avg pixel radiance L2 = 0.212
Avg pixel radiance H0 = 0.684
Avg pixel radiance V0 = 0.137
Avg pixel radiance L0 = 0.821
```



The final example in Figure 2.7:5 has  $v_{fric} = 60.0$  cm/sec and  $wind\_az$  is now 0 degrees.

FIGURE 2.7:5. L0 image.



## 2.8 Slope Spectra Analysis of Synthetic Images

Analysis of slope spectra involves several related figures of merit that are all based upon the difference in slope spectral response between the original slope magnitude spectra used in the synthesis and the final synthetic image magnitude spectra resulting from the nonlinear transformation of wave slope to radiance. These figures of merit can be correlated to either wave slope variance or the difference in wave slope variance.

The first figure of merit is the 2D difference (or error) spectrum  $D_2(k, \phi)$ . An ensemble average of four independent spectral realizations of the same synthetic geometry is calculated and the fundamental component  $N'(0,0)$  (that represents the mean image radiance) is then removed.

$$N'(k, \phi) = \sum_{i=1}^4 N_i(k, \phi) \quad [2.8:1]$$

$$N''(k, \phi) = N'(k, \phi) - N'(0,0) \quad [2.8:2]$$

where

$N$  = the magnitude spectrum of the synthetic radiance image,  
 $k$  = wavenumber,

and

phi = azimuth angle.

The modified image magnitude spectrum,  $N''(k, \phi)$ , is normalized so that the peak response matches the peak response of the original slope magnitude spectrum  $M_0(k, \phi)$ . The motivation is to provide a scaled difference (error) spectrum to illustrate the nonlinearity in slope spectral response of  $N_{\text{norm}}(k, \phi)$  relative to  $M_0(k, \phi)$  over the domains of wavenumber and azimuth.

$$N_{\text{norm}}(k, \phi) = N''(k, \phi) / \text{MAX}(M_0(k, \phi)) \quad [2.8:3]$$

$$D_2(k, \phi) = M_0(k, \phi) - N_{\text{norm}}(k, \phi) \quad [2.8:4]$$

Note that coordinates  $k$  and  $\phi$  are specified here for the consistency of this discussion only. The actual arrays used in this study are specified by rectangular spatial frequency (wavenumber) coordinates  $(k_x, k_y)$  such that

$$D_2(k_x, k_y) = D_2(k * \text{COS}(\phi), k * \text{SIN}(\phi)). \quad [2.8:5]$$

The second and third figures of merit are the 1D squared difference (error) spectra  $D_1(k)$  and  $D_1(\phi)$ . The motiva-

tion is to provide a scaled 1D variance spectrum to illustrate the nonlinearity in spectral response of  $N_{norm}(k)$  relative to  $M_0(k)$  over the domain of wavenumber only, and of  $N_{norm}(\phi)$  relative to  $M_0(\phi)$  over the domain of azimuth.

$$D_1(k) = \int |D_2(k, \phi)|^2 d\phi \quad [2.8:6]$$

$$D_1(\phi) = \int |D_2(k, \phi)|^2 dk \quad [2.8:7.]$$

The fourth figure of merit is the squared integrated value of the difference (error) in slope spectral response  $D_0$ . The 2D difference (error) spectrum  $D_2(k, \phi)$  is integrated with respect to both azimuth  $\phi$  and wavenumber  $k$ . The motivation is to provide a single number to describe the nonlinearity in slope spectral response in the calculation of the variance of  $N_{norm}$ , i.e. the estimate of the variance of  $M_0$ , relative to the actual slope variance of the original (input) spectrum.

$$D_0 = \int \int |D_2(k, \phi)|^2 dk d\phi. \quad [2.8:8]$$

The cross-wind and along-wind squared differences are also calculated:

$$D_{0c} = \int \int |D_2(k, \phi) * \sin(\text{wind\_az} - \phi)|^2 dk d\phi, \quad [2.8:9]$$

and

$$D_{0a} = \int \int |D_2(k, \phi) * \cos(\text{wind\_az} - \phi)|^2 dk d\phi. \quad [2.8:10]$$

These figures of merit are determined for each of four differentially polarized variations of the same synthetic image:

$H_1(x, y)$  = Horizontally polarized, reflected radiance image,

$L_1(x, y)$  = Total (unpolarized), reflected radiance image,

$H_0(x, y)$  = Horizontally polarized, reflected and refracted radiance image,

$L_0(x, y)$  = Total (unpolarized), reflected and refracted radiance image, and

$H_3(x, y)$  =  $H_1(x, y)$  without the incorporation of the sub-resolution wave-slope correlation filter model,

where

$$N_{ji}(k_x, k_y) = \left| \int_{-1}^{+1} IM_i(x, y) \right|$$

for  $IM = H_1, L_1, H_0, L_0, H_3$  and  $i = 1-4$ . [2.8:11]

Because of the large number of spectra to be processed, the Results will include only the  $D_1$  and  $D_0$  data.

Figure 2.8:1 illustrates a typical 2D difference spectrum  $D_2(k_x, k_y)$  for a single synthetic realization of image H1, where wind azimuth is 0 degrees, the sun is at zenith, and the wind friction velocity is 12.0 cm/sec. Note that this example is an 'alphascale' representation, where 'a - z' corresponds to increasing negative scaled values and 'A - Z' corresponds to increasing positive scaled values. Blanks correspond to near-zero deviations (errors) between spectra.

```

CASE01H1.A:HIGH,LOW,DELTA = 0.1485111 -4.0228695E-02 5.6041917E-03
bbbbabasabasabcbccdddee      a bB aadccdbchbbabbabbabbasab      31
aabbbbabbcrbcbbccacd A chaab bacAab A B dbchachbccaaabbasac      30
aabbbbabbcrbbhbhbbccRA      ab ab dc aa bB bbaAbbbchccabbbasab      29
abbchbaaabbbabbcod aa a cd a aabab bba Abcdcbchbbabbchabba      28
aababbbaaabcaasaca aa a A b abb aa Abb bb b casaaaabbabbbaa      27
aababbbaaabbbCABb AAC a a bb b abda C A Sbaccbbabbbaa      26
assabbbaaabbbCABAAC AAA a c assbbbab CAAAB Cchbbabbabbbaa      25
bbabbbaaabbbABCBAAB ABC bb b basaaa AA B 7Sbbabbabbab      24
assabbbaabbbABCBAAB bcsab a Ac aedda AAA Adccabbbab      23
acabbbchbcBABb A a aca a b a wccca aa A AAbabbabbab      22
abbasabbbBAA B A a aca bcb bba a B C A Abaaabb      21
abbasabbbCABA A A A Aaa AabbaA aabaa A A BAb A A AAbcaabba      20
assabbbaabbbAAA C aa bb a A a aaaaaa BAA Bcccaaaaa      19
abbaccBABAABCAc A b bab cadaab d A AABAA ABBAbcabbab      18
basabCABAABBCBAA abcccc a ABbbfec CA a ABABC AA Cdbsa      17
basasCBAABBCAABAA aaaaab ABBbbabbaabca b C CBB BA Caasa      16
basasCA AABABCABBS a ababab c a ABBAAACB BCA BBcaab      15
bbdbbaAABAABACTACA CB C aca a cdccb A B ABBCCCBBA ABBBab      14
baaCBBBCABCBCBDBBB AAB A AABbaAabcaAAA BAA BAAACB ABBBab      13
baaCBBCCCCCBBCACAAAB AA A Aaa Baaab ab ABCBA BBBBBBCCBAAABCDaa      12
aaBABBCCABCBCOCAA ABBCBDBaaabcc bcc a CEC BAAACCCACCCBCCDaa      11
baaaaa ABBA CRABBBBCBCCBCA A afa A a BBSBA BAACTCBDCCCCOAAAb      10
bdBCBCBDBBAAABBCBCCBCAACBCCR bbb AAAAAA BCCAAAB CDDCDCBCDCCBAAAb      9
aaBABBCCDAACODDDBBCCBCC B A Aaa a BBSBA CADDCCBCDCBCCBAAAb      8
aaCABBBBCBAAABDKECCBBDAAAB AA ccd BBDDBBCCBACB BCCCBCBCCBCC      7
ADCBBCBCBCDKECCCKEDDCBDBCC bccbcba CDBACB3BDCCCABDCCBDCBCC      6
bocCBBCBCDCCBDBCC BEEEDDCBA BEEEDFBCCFBCDGEB dEdDCBA BEEEDFBCCFBCDGEB      5
accCBCCCCCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      4
accCBCCCCCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      3
brcBAAABCDDCDCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      2
accBABBCCDDCDCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      1
BABBBCBCDDCDCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      0
abbBCCBCDDCDCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -1
bchBCCBCDDCDCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -2
abbBCCBCDDCDCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -3
abbBCCBCDDCDCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -4
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -5
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -6
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -7
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -8
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -9
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -10
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -11
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -12
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -13
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -14
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -15
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -16
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -17
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -18
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -19
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -20
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -21
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -22
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -23
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -24
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -25
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -26
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -27
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -28
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -29
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -30
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -31
accBACBCCACDCCBDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDCDDC      -32

```

FIGURE 2.8:1. 2D difference (error) spectrum,  $D_2(k_x, k_y)$ .

### 3.0 RESULTS AND DISCUSSION

The  $D_1$  results (slope variance versus wavenumber; slope variance versus azimuth) are listed in Appendix I. These results provide the most graphic illustration of the wave-slope error generated by the parametric variation of wind azimuth, solar position and wind velocity. The  $D_0$  results (partial and total integrated slope variance) are listed in Appendix II. The  $D_2$  difference (error) spectra (from which the  $D_1$  and  $D_0$  results are derived) are not included with this report to avoid cumbersome detail (i.e. 360 graphic examples of three-dimensional difference/error spectra).

### 3.1 Discussion of D<sub>0</sub> Results

Please refer to the D<sub>0</sub> results in Appendix II.

#### Parametric Variation of Mean Radiance

The mean radiance reflected/refracted in the direction of observation is the central ordinate of the forward Fourier transform of the radiance image and is calculated for each imaging geometry and for each of five polarized radiance conditions:

H<sub>1</sub> - reflected horizontal radiance,

L<sub>1</sub> - reflected total (horizontal + vertical) radiance,

H<sub>0</sub> - reflected & refracted horizontal radiance,

L<sub>0</sub> - reflected & refracted total radiance, and

H<sub>3</sub> - reflected horizontal radiance without the sub-resolution wave-slope model (essentially a reconstruction of the Chapman & Irani synthesis.)

H<sub>3</sub> will be compared with H<sub>1</sub> as an examination of the effect of sub-resolution wave slopes on model results.

As expected, the following conditions are observed:

- 1) The mean radiance is maximum when the imaging geometry favors specular reflection of the sun, i.e. the solar position is azimuth angle  $\phi_0 = 0.0$  deg and zenith angle  $\theta_0 = 53.2$  deg (denoted  $(\phi_0, \theta_0)$  for the remainder of this discussion). The second highest mean radiance is observed where the sun is at zenith (i.e (0,0)). Mean radiance decreases as sun azimuth varies from 0 degrees.
- 2) Under the conditions of maximum solar reflection and  $v_{fric} = 12.0$  cm/sec, the slightly disturbed water surface yields a mean radiance of approximately 1.98 for  $L_0$  (averaged over the four wind azimuths). This value compares favorably with the theoretical  $L_0$  value of 2.62 when the surface is a totally flat, specular surface. Mean radiance drops to 1.75 when  $v_{fric} = 36.0$  cm/sec and drops to 1.50 when  $v_{fric} = 60.0$  cm/sec. As the surface becomes rougher, the probability of reflecting facets having zero slope decreases and less of the direct solar radiance is reflected.
- 3) Under the conditions of maximum solar reflection and  $v_{fric} = 60.0$  cm/sec, mean radiance is highest (1.57 for  $L_0$ ) when the dominant wind azimuth  $wind\_az$  is orthogonal to the plane of reflection and lowest (1.45 for  $L_0$ ) when  $wind\_az$  is parallel to the plane of reflection. As the surface becomes

rougher due to wind, the slope distribution increases its spread along the azimuthal plane of the wind defined by wind\_az. When wind\_az is orthogonal to the plane of reflection, there is a greater probability that a facet will reflect radiance from the skydome near the plane of reflection than when wind\_az is parallel to the plane. With the sun located in this plane, the probability that a facet will reflect radiance from the vicinity of the sun is greater when wind\_az is orthogonal. This effect can be observed, to a lesser degree, for all sun orientations in the plane of reflection. The effect diminishes as vfric decreases because of the lower directional attenuation of the dominant low-frequency waves.

#### Demonstration of Statistical Model Stability

Comparisons of symmetric test cases with three different solar geometries provide an indication of the statistical stability of the complete model.

TABLE 3.1:1. Symmetric test cases and their parameters.

<u>Comparison 1</u>	<u>Comparison 2</u>	<u>Comparison 3</u>	vfric
12 and 24	07 and 19	11 and 23	12 cm/sec
60 and 72	55 and 67	59 and 71	36 cm/sec
36 and 48	31 and 43	35 and 47	60 cm/sec
Sun Coord	(0,0)	(180,53.2)	

The three sets of test cases have mirror symmetry with respect to observation: the sun is in the plane of reflection ( $\phi_{10} = 0$  or  $180$  deg) but the dominant wind azimuths `wind_az` for the comparison cases are directed  $45$  degrees away from the the plane of reflection and differ in sign only. The respective test cases should demonstrate similar statistics based on the ensemble average of the magnitude spectra of the four synthetic scenes within each test case.

TABLE 3.1:2. Comparison of mean radiances between symmetric geometries.

Comparison of symmetric test cases										
Condition	12	vs	24	60	vs	72	36	vs	48	
H1	1.7183		1.7183	1.4580		1.4629	1.2105		1.2145	
L1	1.7740		1.7739	1.5352		1.5404	1.2886		1.2929	
H0	1.8241		1.8241	1.5635		1.5682	1.3155		1.3196	
L0	1.9855		1.9854	1.7461		1.7511	1.4987		1.5030	
Avg Diff	$\pm 0.0027$	%		$\pm 0.3145$	%		$\pm 0.3147$	%		
Condition	07	vs	19	55	vs	67	31	vs	43	
H1	0.9664		0.9649	1.0081		1.0093	0.9621		0.9645	
L1	0.9760		0.9742	1.0300		1.0304	0.9966		0.9977	
H0	1.1519		1.1503	1.1918		1.1928	1.1436		1.1459	
L0	1.3469		1.3451	1.3975		1.3974	1.3596		1.3606	
Avg Diff	$\pm 0.1530$	%		$\pm 0.0622$	%		$\pm 0.1583$	%		
Condition	11	vs	23	59	vs	71	35	vs	47	
H1	0.4582		0.4577	0.4607		0.4619	0.4223		0.4237	
L1	0.4674		0.4668	0.4788		0.4800	0.4456		0.4470	
H0	0.5640		0.5635	0.5661		0.5672	0.5274		0.5288	
L0	0.6790		0.6784	0.6897		0.6907	0.6558		0.6571	
Avg Diff	$\pm 0.1036$	%		$\pm 0.2121$	%		$\pm 0.2766$	%		

Mean radiances (the ensemble DC components) are similar to within  $\pm 0.32\%$  among all comparisons. There does not appear

to be any correlation between average percent difference and solar geometry even though the mean radiance diminishes as the geometry moves away from maximum solar reflection.

The ensemble fundamental components differ more than expected and this is an indication that a larger ensemble may be one additional requirement for statistical stability.

TABLE 3.1:3. Comparison of fundamental components between symmetric geometries.

<u>Comparison of symmetric test cases</u>						
<u>Condition</u>	<u>12 vs 24</u>	<u>60 vs 72</u>	<u>36 vs 48</u>	<u>1</u>	<u> </u>	<u> </u>
H1	0.0359	0.0315	0.0243	0.0220	0.0183	0.0167
L1	0.0400	0.0374	0.0269	0.0245	0.0194	0.0177
H0	0.0347	0.0320	0.0229	0.0206	0.0170	0.0158
L0	0.0376	0.0350	0.0241	0.0218	0.0169	0.0159
Avg Diff	$\pm 8.3628\%$	$\pm 9.4935\%$	$\pm 7.6205\%$			
<u>Condition</u>	<u>07 vs 19</u>	<u>55 vs 67</u>	<u>31 vs 43</u>	<u>1</u>	<u> </u>	<u> </u>
H1	0.0216	0.0225	0.0120	0.0126	0.0060	0.0060
L1	0.0211	0.0223	0.0106	0.0114	0.0048	0.0048
H0	0.0194	0.0204	0.0094	0.0099	0.0058	0.0056
L0	0.0168	0.0180	0.0056	0.0061	0.0104	0.0086
Avg Diff	$\pm 5.2375\%$	$\pm 6.2567\%$	$\pm 5.1890\%$			
<u>Condition</u>	<u>11 vs 23</u>	<u>59 vs 71</u>	<u>35 vs 47</u>	<u>1</u>	<u> </u>	<u> </u>
H1	0.0124	0.0126	0.0083	0.0085	0.0051	0.0048
L1	0.0135	0.0138	0.0091	0.0094	0.0052	0.0049
H0	0.0112	0.0115	0.0069	0.0069	0.0040	0.0038
L0	0.0111	0.0114	0.0063	0.0063	0.0032	0.0032
Avg Diff	$\pm 2.2504\%$	$\pm 1.3861\%$	$\pm 4.1629\%$			

For these three geometries, the fundamental components are similar to within  $\pm 10\%$  among all comparisons. This may be a poor choice of imaging geometry for statistical validation

because the contrast ratios (fundamental/DC components) are not the highest among the full set of test cases. The main consideration that applies to these particular geometries is that a nearly saturated image (i.e a very high DC component relative to the remainder of the spatial spectrum) skews the confidence of spectral estimation in favor of the DC component, a parameter which has an expected value of zero in the slope spectral estimate, i.e. the water surface is expected to have a mean slope of zero. However, there does appear to be a correlation between a decrease in the mean radiance with a decrease in the average percent difference (and therefore increased statistical confidence) in the fundamental components. A fundamental consideration for this method is to determine an optimal geometry which removes spectral energy from the DC component and distributes it across the rest of the spatial spectrum with the intent of increasing the statistical stability of the non-DC portion of the spectrum.

#### Correlation of Contrast Ratio with Slope Variance

As explained in the Methods section, the DC component of the radiance image magnitude spectrum is removed and the remaining image spectrum is normalized with respect to the input slope magnitude spectrum  $M_0$  by a geometric scaling

factor between the peak slope components. The difference between the two spectra is then defined as the error, or difference, spectrum. Integration of the squared error over both wavenumber and azimuth yields the  $D_0$  results, the total slope variance due to errors/differences between the two spectra.

The contrast of the wave image was defined by Chapman and Irani as the ratio of the fundamental component to the DC component and was used as a figure of merit in their error analysis. A comparison of the contrast ratio of the image spectrum with the resulting integrated squared deviation (variance) from its difference spectrum demonstrates a good correlation between high contrast and low error. A simple matrix is presented, where the contrasts of the images for each geometry (with the exception of H3) are ranked from 1 to 4 (high contrast to low contrast) and the corresponding  $D_0$  results are likewise ranked from 1 to 4 (low to high error).

TABLE 3.1:4. Rank correlation matrices for contrast ratio vs. slope variance.

CONTRAST				
	1	2	3	4
V 1	22	1	1	0
A 2	1	18	3	2
R 3	1	3	14	6
4	0	2	6	16

CONTRAST				
	1	2	3	4
V 1	17	1	0	6
A 2	1	14	8	1
R 3	1	7	13	3
4	5	2	3	14

CONTRAST				
	1	2	3	4
V 1	9	4	5	6
A 2	2	10	6	6
R 3	4	9	6	5
4	9	1	7	7

The above results display a high correlation at  $v_{fric} = 60.0$  cm/sec and increasing loss of correlation as  $v_{fric}$  decreases. This can be explained as a result of decreasing statistical confidence and increasing ambiguity among contrast values and  $D_0$  values. For  $v_{fric} = 60.0$  cm/sec, the differences between contrast ratios and  $D_0$  values are generally on the order of  $\pm 10\%$ ; for  $v_{fric} = 36.0$  cm/sec, the differences are often less than  $\pm 5\%$ ; for  $v_{fric} = 12.0$  cm/sec, the dif-

ferences are often less than  $\pm 1\%$ . With the assumptions 1) that the fundamental component (the numerator of the contrast ratio) has an average model error tolerance of  $\pm 6\%$  (based on the analysis of Table 3.1:3) and 2) the corresponding  $D_0$  value may have a similar tolerance, it is reasonable to consider that the confidence of unambiguous rank-ordering should decrease for values at and below  $v_{fric} = 36.0$  cm/sec.

#### Determination of Optimal Imaging Geometry

Figure 3.1:1 provides a graphic comparison of the  $D_0$  values for each geometry. The intent of this comparison is to determine those imaging geometries which provide the lowest integrated squared errors (slope variance) for all sampled variations of dominant wind azimuth  $wind\_az$  and friction velocity  $v_{fric}$ .

Within this limited sample space, the unambiguous selection for optimal imaging of wind-wave slopes is a geometry which requires the sun to be directed 135 degrees away from the azimuth of reflection (i.e. observation) or, by complement, 45 degrees away from the solar azimuth of maximum specular reflection. The  $D_0$  results for this geometry are generally lowest for all polarized conditions.

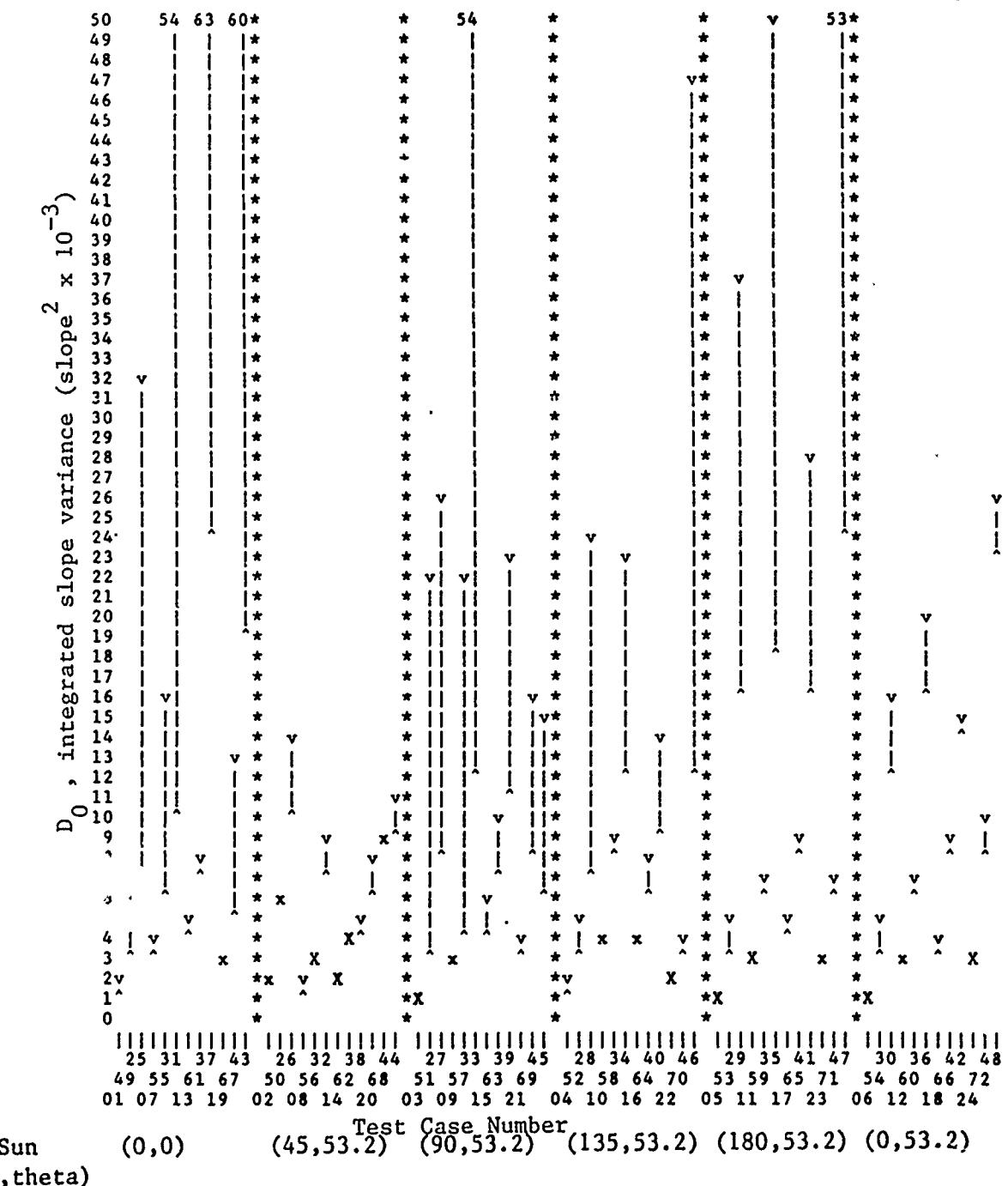


FIGURE 3.1:1. Range of slope variance ( $\text{error}^2$ ) versus test case number (slope variance expressed as  $\text{slope}^2 \times 10^{-3}$ ).

The second choice of imaging geometry places the sun at the position of maximum specular reflection but this is an ambiguous selection because other geometries provide lower D<sub>0</sub> results for particular values of wind\_az and vfric.

This result, from two-dimensional analysis of the current study model, is in full agreement with the results of the one-dimensional survey of Chapman and Irani [1981] and demonstrates a secondary validation of their approach.

#### Effect of Subsurface Refraction on Slope Variance

The following conditions are observed:

- 1) The effect of refraction on slope variance is minimal when contrast ratios are high relative to the set of all test cases. It was expected that the additional nonlinearity created by refraction in the transformation of slope to radiance would consistently generate higher slope variances in H<sub>0</sub> and L<sub>0</sub> relative to H<sub>1</sub> and L<sub>1</sub>. However, the stronger correlation between contrast and slope variance tends to swamp any effect created by subsurface refraction.
- 2) As expected, the effects of refraction are most noticeable when the sun is furthest away from point of maximum

specular reflection, i.e. when the proportion of refracted radiance relative to reflected radiance is largest. In these test cases, the contrast ratios approach a minimum among all test cases and the slope variances approach a maximum. The large differences between H1 and H0 and between L1 and L0 confirm that the decrease in contrast and increase in slope variance are both due to the effect of refraction on the minimum reflection cases.

#### Effect of Polarization on Slope Variance

A comparison of H1 with L1 indicates no apparent correlation between polarization and slope variance. Again, the stronger correlation between contrast and slope variance tends to swamp any effect created by differences in polarization. It was expected that the decrease in mean radiance from L1 to H1 would be offset by an increase in contrast and consequent decrease in slope variance. No such effect was discernible among this small sample. Because of Brewster-angle viewing, the effect due to the addition of vertically polarized radiance should be at a minimum among all observation geometries.

### 3.2 Discussion of D<sub>1</sub> Results

Please refer to the graphic D<sub>1</sub> results (slope variance versus wavenumber; slope variance versus azimuth) in Appendix I. Note that the slope variance for H1 is plotted as a solid line; L0 is plotted as a dashed line; H3 is plotted as a dotted line.

#### Slope Error as a Function of Wavenumber

The following conditions are observed:

- 1) For  $v_{fric} = 12.0$  cm/sec, the major source of slope error occurs in the vicinity of the dominant, low-spatial-frequency peak centered near  $k = 0.12$  rad/cm.
- 2) With the sole exception of the optimal imaging geometry (as determined in Chapter 3.1), the higher-spatial-frequency slope errors increase relative to the error peak near  $k = 0.12$  rad/cm as  $v_{fric}$  increases.
- 3) With the sole exception of the optimal imaging geometry, the higher spatial frequencies (above  $k = 0.12$  rad/cm) are the dominant source of slope error at  $v_{fric} = 60.0$  cm/sec.

4) At the optimal imaging geometry, the major source of slope error remains in the vicinity of the dominant, low-spatial-frequency peak centered near  $k = 0.12$  rad/cm for all values of  $v_{fric}$ .

5) At the optimal imaging geometry (and at the geometry of maximum solar reflection), the plots of H1 and L0 (as well as the unplotted L1 and H0 data) closely approximate each other within the set of test cases. The same plots within the other imaging geometries tend to approximate each other at  $v_{fric} = 12.0$  cm/sec and then deviate from each other at higher values of  $v_{fric}$ .

#### Variation of H3 with Respect to Wavenumber

H3 represents the results based on the model of Chapman and Irani where no sub-resolution wave-slope model is incorporated. H1 represents results based upon the inclusion of a radiance attenuation model to add the effect of sub-resolution wave slopes on scene radiance.

The following conditions are observed:

1) H3 does not demonstrate a good correlation with any of the model parameters as a function of wavenumber. Within the

limited scope of this experiment, H3 appears to be independent of the model parameters as a function of wavenumber.

- 2) As expected, H3 maps most closely with H1 when  $v_{fric} = 12.0$  cm/sec. The correlation filter that is used to attenuate radiance as a function of the sub-resolution wave-slope distribution provides a nearly impulse response at  $v_{fric} = 12.0$  cm/sec than at higher values of  $v_{fric}$  and the resulting difference between the H1 and H3 radiance models is slight.
- 3) Within the constraint of optimal imaging geometry, H3 maps closely to H1, L1, H0, and L0 when  $v_{fric} = 12.0$  cm/sec but deviates significantly at higher values of  $v_{fric}$ . The most general trend to note is that the slope error associated with H3 becomes much larger relative to the other four radiance conditions as wavenumber increases.
- 4) Within the constraint of maximum solar reflection geometry, H3 maps closely to H1, L1, H0 and L0 at higher values of  $v_{fric}$  but deviates when  $v_{fric} = 12.0$  cm/sec. This effect is opposite to that observed in 3) above.

The demonstrated independence of H3 as a function of wavenumber is also in agreement with the results of Chapman

and Irani. However, the significant result to note is that the other four radiance conditions do demonstrate a definite correlation with wavenumber at the optimal imaging geometry. This correlation is also exhibited in the geometry of maximum solar reflection, although to a lesser degree.

#### Slope Error as a Function of Azimuth

The following conditions are observed:

- 1) There is a definite azimuthal dependence between slope variance and the parameters of solar azimuth and dominant wind azimuth. This dependence is most readily observable in the two most optimal imaging geometries.
- 2) When  $v_{fric} = 12.0$  cm/sec and  $wind\_az$  is in the plane of reflection, a dominant error lobe is observable which has an orientation within 20 degrees of the azimuth orthogonal to the sun position. A minor error lobe is observable in the azimuthal plane of the sun, i.e. approximately orthogonal to the dominant lobe.
- 3) When  $v_{fric} = 12.0$  cm/sec and  $wind\_az$  is varied, the error lobes redistribute so that the dominant errors angularly skew toward the plane of the dominant wind azimuth.

4) As  $v_{fric}$  increases, the error lobes tend to angularly spread out so that the dominant errors redistribute between the solar azimuth and the dominant wind azimuth.

#### Variation of H3 with Respect to Azimuth

The following conditions are observed:

1) The H3 results follow the general trends outlined in Chapter 3.2 above. As discussed with respect to wavenumber, H3 maps closely to the results of the other four radiance conditions when  $v_{fric} = 12.0$  cm/sec.

2) Where H3 deviates significantly from the other four conditions, the additional slope error is observed to align with the plane of the solar azimuth. This result is most noticeable within the two most optimal geometries.

The demonstrated dependence of H3 as a function of azimuth is also in general agreement with the limited results of Chapman and Irani. However, the results for H1, L1, H0, and L0 demonstrate a more consistent correlation with the variation of model parameters.

#### 4.0 CONCLUSION

The results of the current study have established three main conclusions: 1) the limited results of Chapman and Irani [1981] have been generally verified, 2) the existence of an optimal imaging geometry for slope spectrum estimation is indicated, and 3) the enhanced model, incorporating the effects of sub-resolution wave slopes and subsurface refraction on observed radiance, demonstrates a significant effect on wave-slope spectra derived from imagery.

The most conclusive result is the analytic determination of the region containing an optimal geometry for the estimation of wave-slope spectra from imagery under the specified conditions of this study. Through the use of an integrated set of analytic models, the variation of a limited number of parameters has predicted the magnitude of slope variation as a function of azimuth and wavenumber. These predictions exceed the capability of a deterministic second-order theory. While this study has been directed toward a generalization of the Chapman and Irani study, the generalized model requires a more complete exploration of the parametric surface.

## Suggestions for Further Research

There are several potential directions for further research which will build upon this study. The types of research fall into two main categories: 1) comparison of the analytic model with empirical data derived from a physical experiment and 2) refinement and extension of the parametric surface exploration of the analytic model.

### Physical Experiment

It is obvious that an analytic model requires empirical results to verify its predictions. However, the complexity of this model and its limited utility presumes that the real world problem has an even higher level of complexity. The requirement still exists for the execution of a controlled experiment to calibrate the various techniques that attempt to measure the two-dimensional elevation/slope spectra of large water surfaces under natural conditions.

### Refinement of the Analytic Model

The current study verified the limited results of Chapman and Irani [1981] and amplified their method to execute an only slightly larger exploration of the paramet-

ric surface. There are many assumptions in this model which remain to be tested through a vigorous exploration of the currently unvaried parameters:

sample size and kmax

ensemble size

radiant wavelength and index of refraction

zenith angle of observation

Also, other analytic functions await comparison with the current analytic model and with existing empirical data:

input elevation spectrum

clear sky versus cloudy sky radiance

clear water versus turbid water refraction

sub-resolution wave-slope distributions

- Gram-Charlier versus Gaussian

A secondary (and more vigorous) approach to the simulation of wave shadowing could also be explored: the elevation power spectrum is available for the generation of synthetic height arrays which can be used to modify the radiance distributions at each surface element. A comparison of the current synthesis with this nonlinear masking technique and with real imagery could potentially verify the utility of the current synthesis at increasingly larger zenith angles of observation.

Additionally, a refined exploration of the current parameter set might establish an even more optimal imaging geometry. The results of the current study have placed the new point of departure at a solar azimuth angle 45 degrees from the point of maximum specular reflection. Small variations of solar azimuth and solar zenith angle from this point may isolate the ideal image geometry for slope spectrum estimation.

Finally, the method of analysis can be more robust. The current study deviates from the Chapman and Irani study with respect to the choice of figures of merit. However, the current study is consistent with the earlier study with respect to the implied assumption that the output radiance magnitude spectrum should be compared with the undecomposed slope magnitude spectrum. There is evidence [Stilwell, 1969; Kasevich, 1975] that the x-component slope magnitude spectrum alone or a linear combination of the two slope-component magnitude spectra may yield a more accurate synthetic representation of the radiance magnitude spectrum. A more detailed analysis, using results from the current study model, could verify this theory and provide a more accurate tool for the estimation of wave-slope spectra from imagery.

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## APPENDIX I - D<sub>1</sub> Results

The two-dimensional graphic results for the 72 combinations of solar position, dominant wind azimuth, and wind friction velocity are presented in Appendix I. Each page presents the graphic variation of wind friction velocity for a single combination of solar position and dominant wind azimuth:

v<sub>f</sub> = 12.0 cm/sec (Slope Variance versus Wavenumber)  
(Slope Variance versus Azimuth)

v<sub>f</sub> = 36.0 cm/sec (Slope Variance versus Wavenumber)  
(Slope Variance versus Azimuth)

v<sub>f</sub> = 60.0 cm/sec (Slope Variance versus Wavenumber)  
(Slope Variance versus Azimuth)

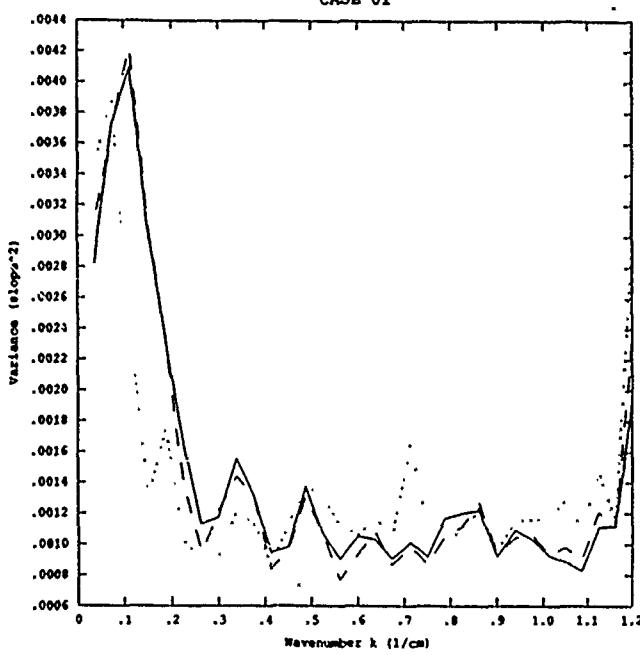
Each set of four pages presents the graphic variation of dominant wind velocity wind\_az for a single solar position:

First Page    wind\_az = 0 deg (or 180 deg with ambiguity)  
Second Page    wind\_az = 45 deg (or -135 deg with ambiguity)  
Third Page    wind\_az = 90 deg (or -90 deg with ambiguity)  
Fourth Page    wind\_Az = 135 deg (or -45 deg with ambiguity)

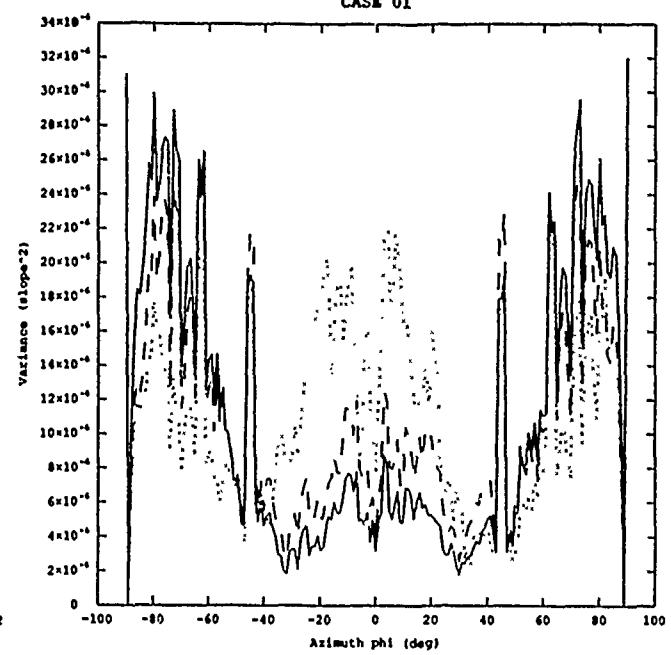
The graphic variation for the six solar positions is then presented in the following order:

First Set	phi <sub>0</sub>	=	0 deg	theta <sub>0</sub>	=	0.0 deg
Second Set	phi <sub>0</sub>	=	45 deg	theta <sub>0</sub>	=	53.2 deg
Third Set	phi <sub>0</sub>	=	90 deg	theta <sub>0</sub>	=	53.2 deg
Fourth Set	phi <sub>0</sub>	=	135 deg	theta <sub>0</sub>	=	53.2 deg
Fifth Set	phi <sub>0</sub>	=	180 deg	theta <sub>0</sub>	=	53.2 deg
Sixth Set	phi <sub>0</sub>	=	0 deg	theta <sub>0</sub>	=	53.2 deg

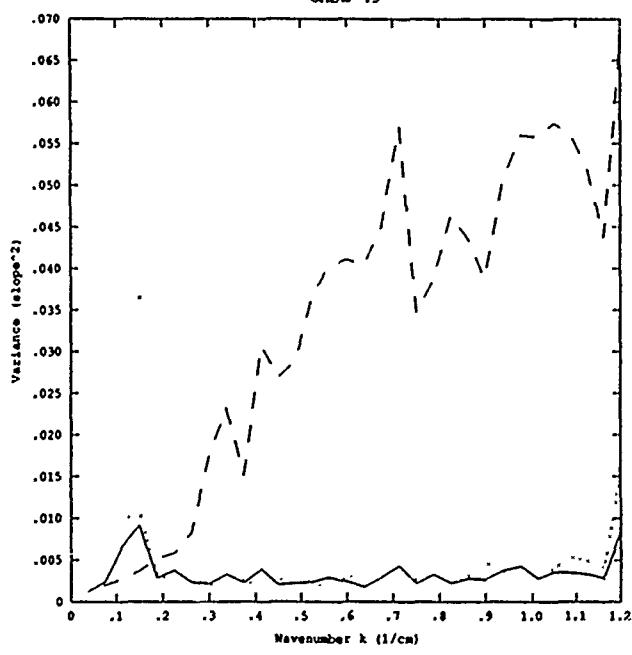
CASE 01



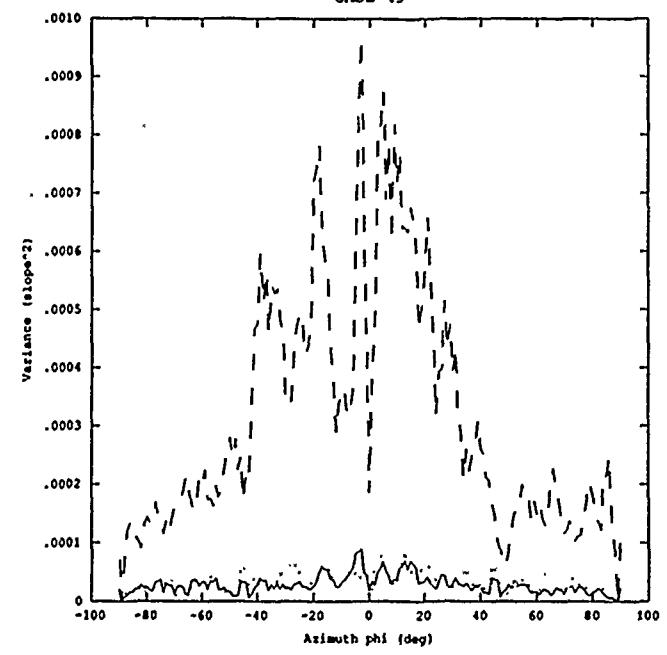
CASE 01



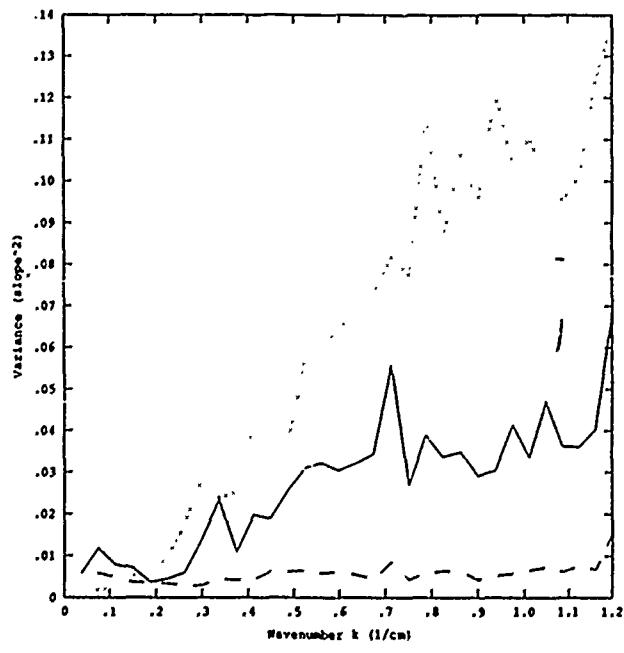
CASE 49



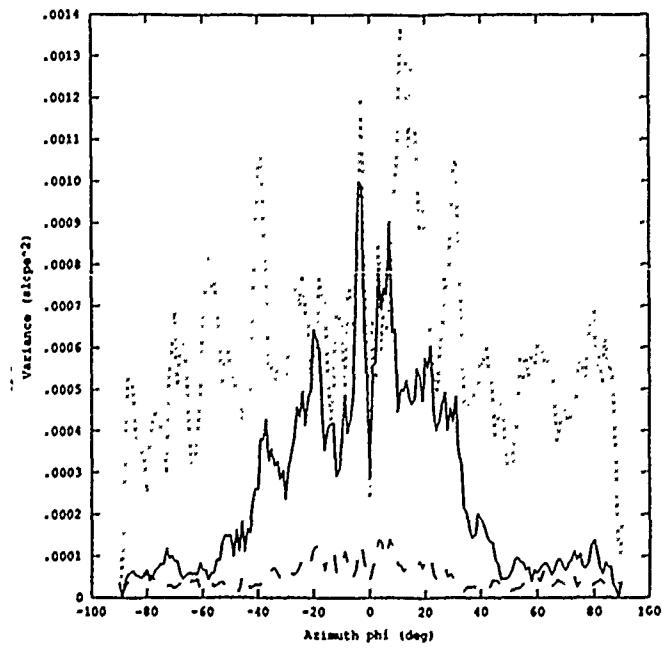
CASE 49

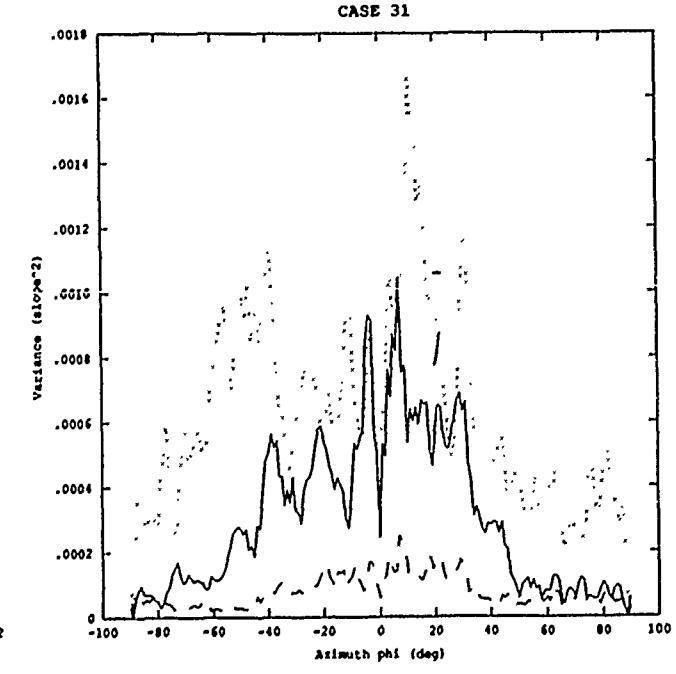
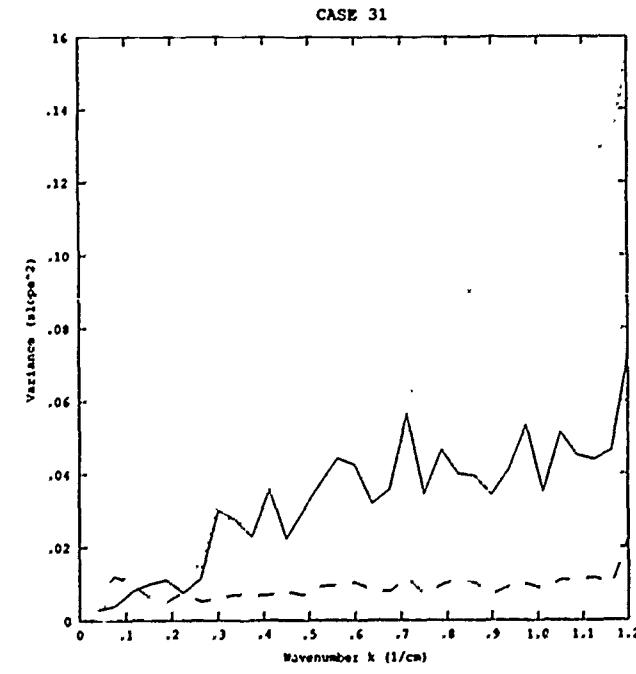
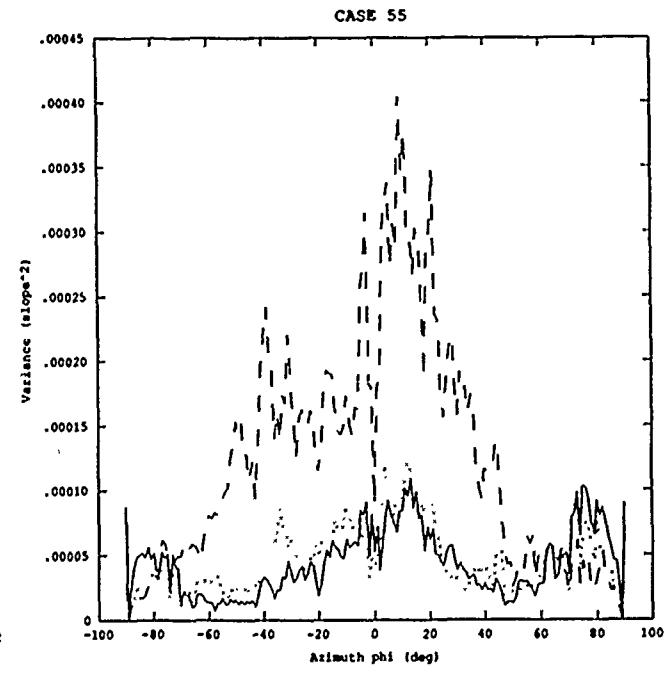
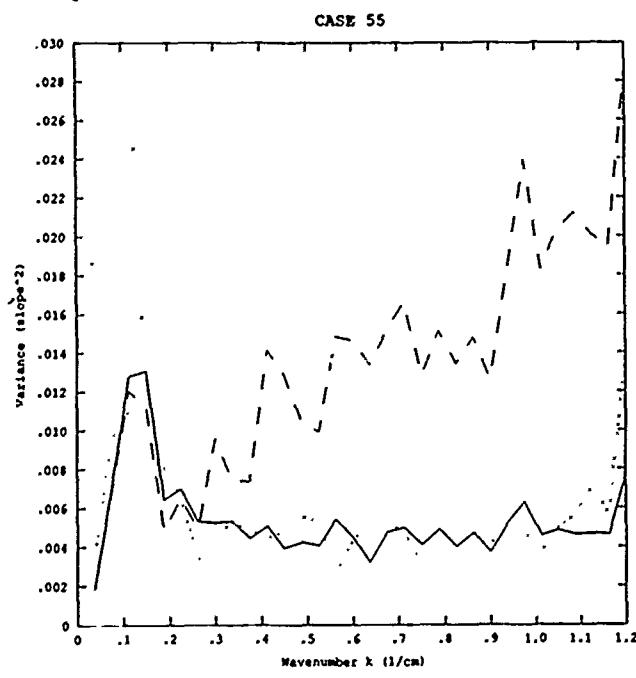
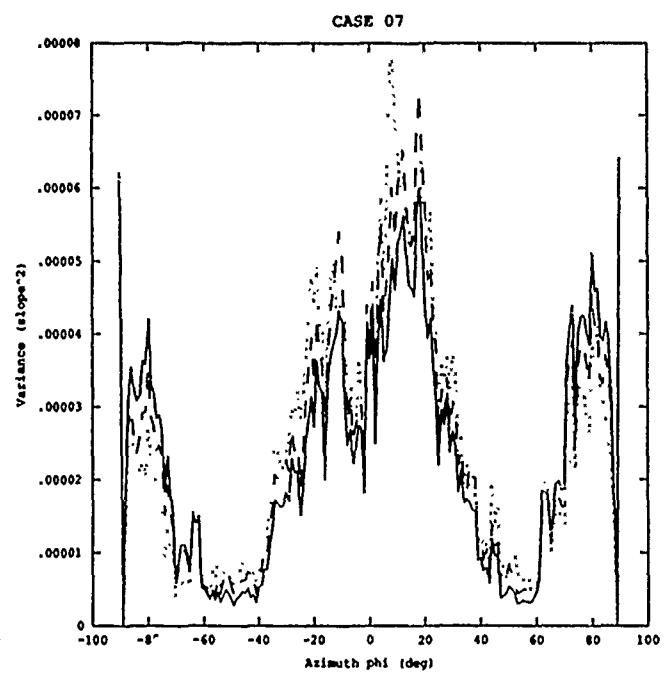
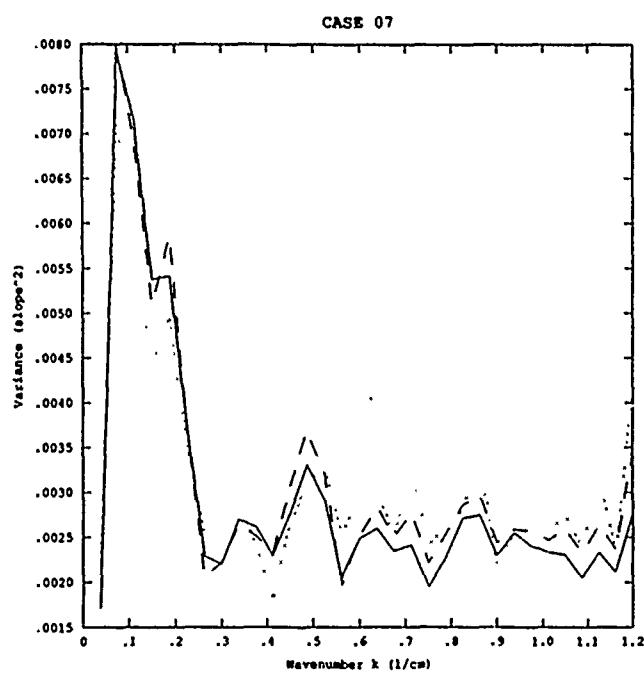


CASE 25

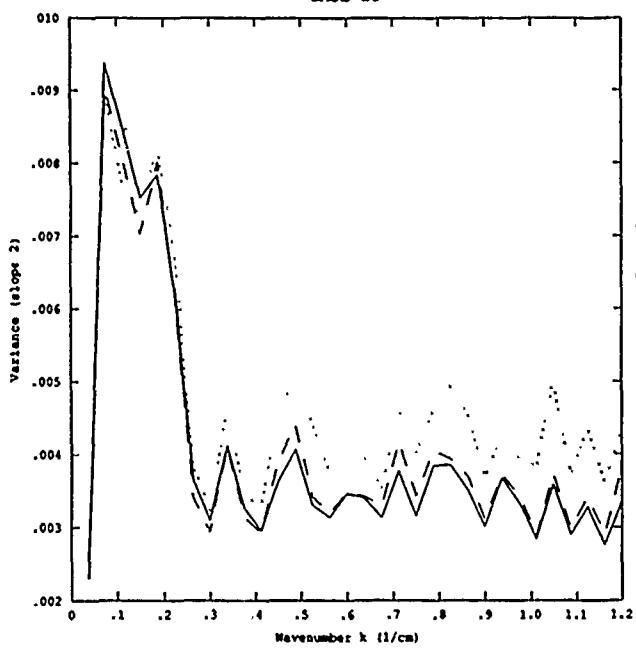


CASE 25

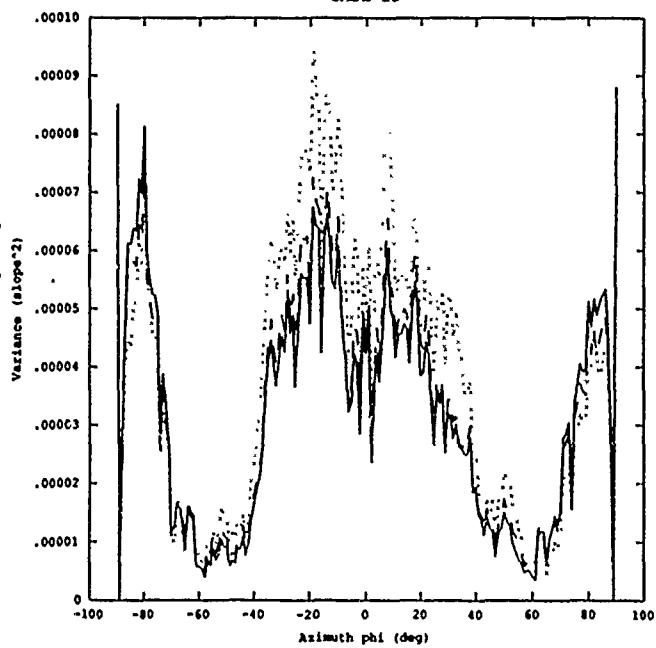




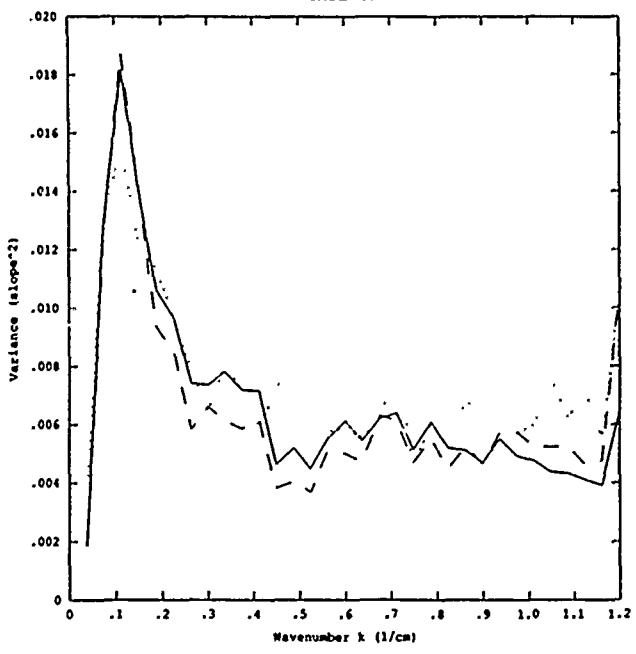
CASE 13



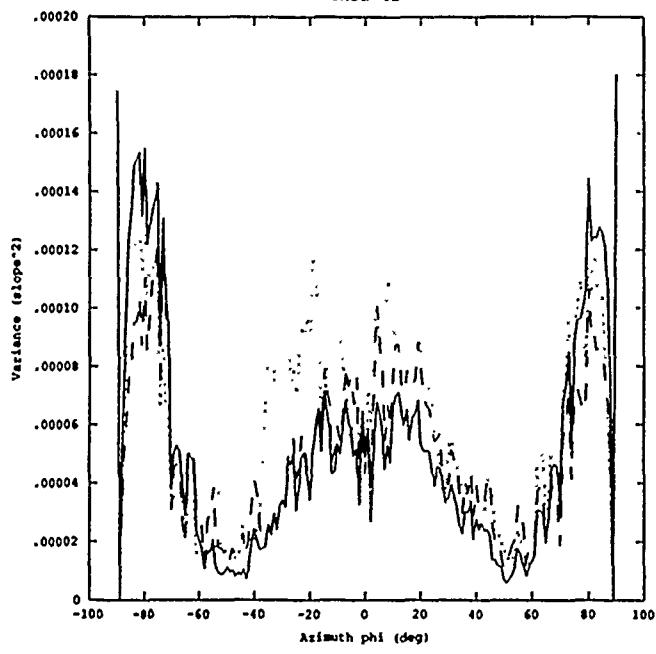
CASE 13



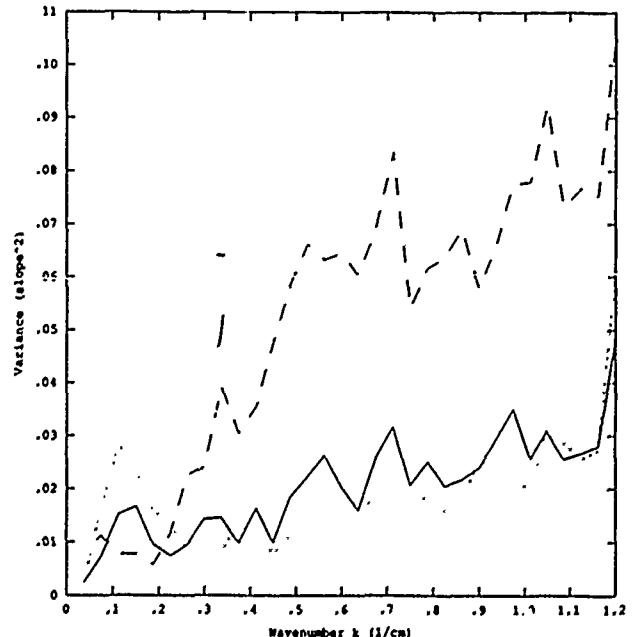
CASE 61



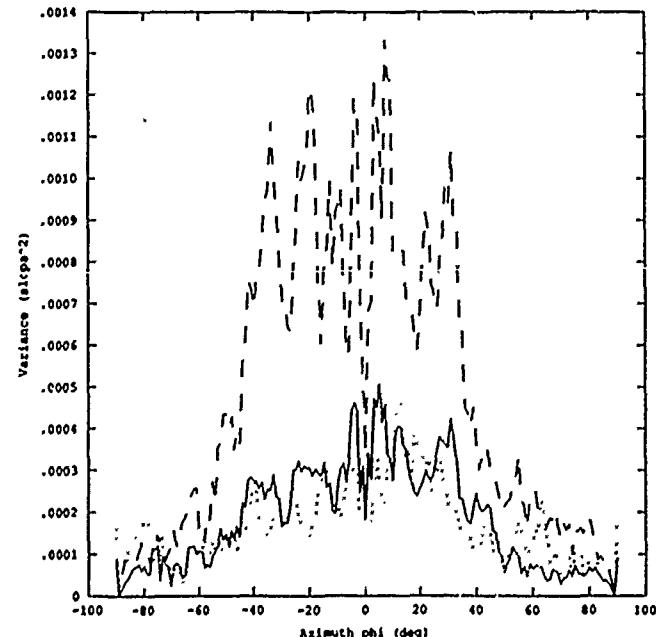
CASE 61



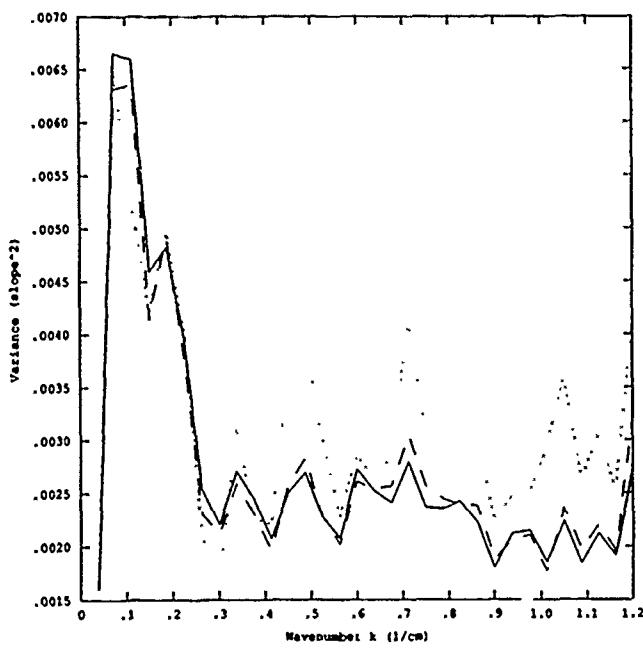
CASE 37



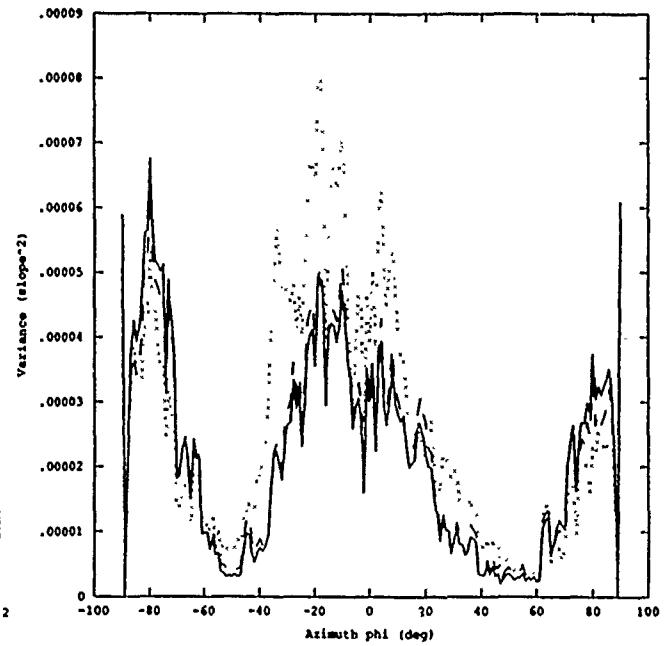
CASE 37



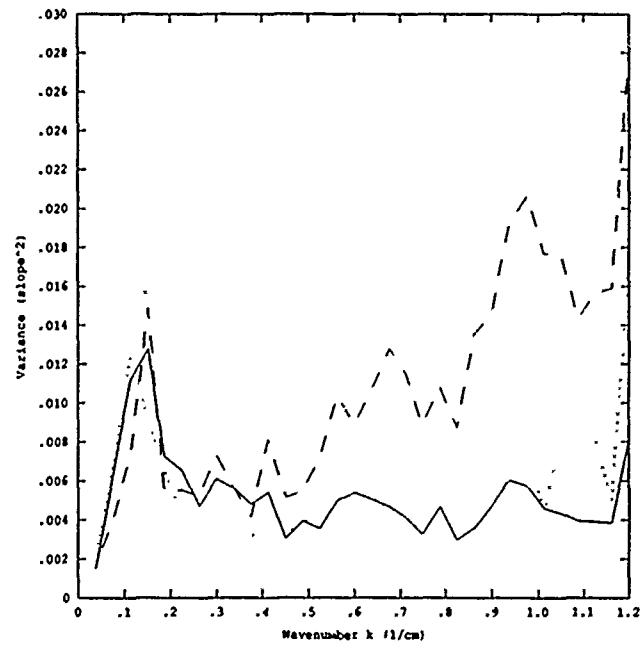
CASE 19



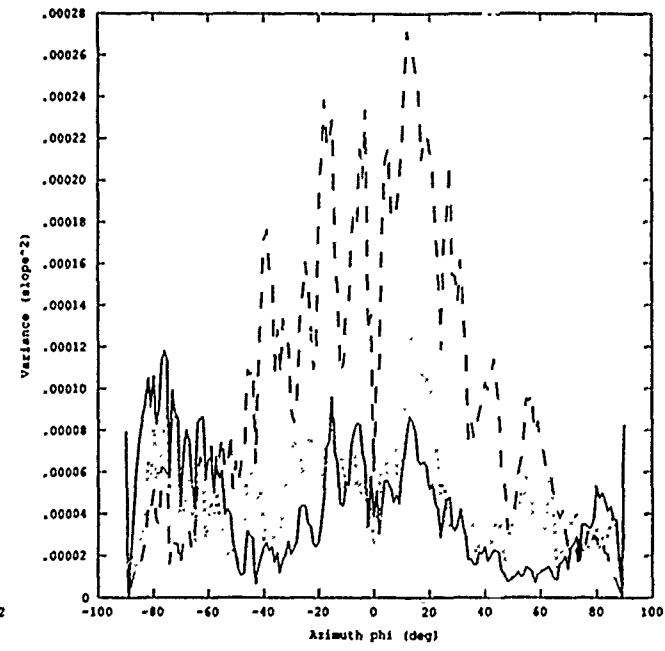
CASE 19



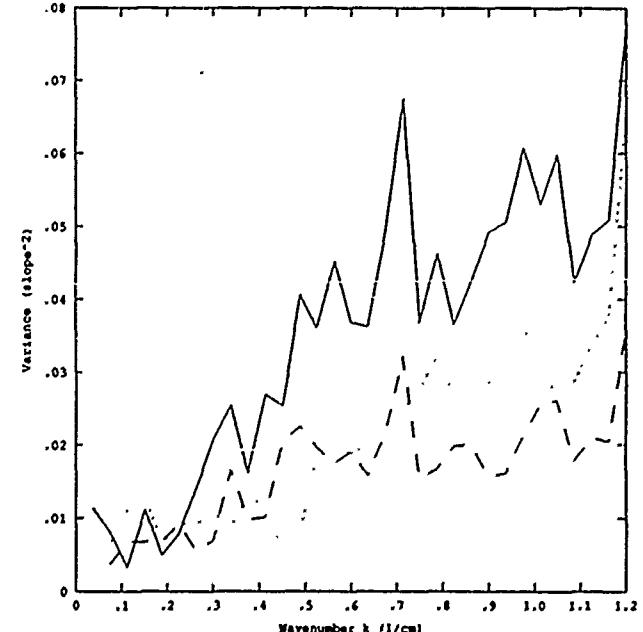
CASE 67



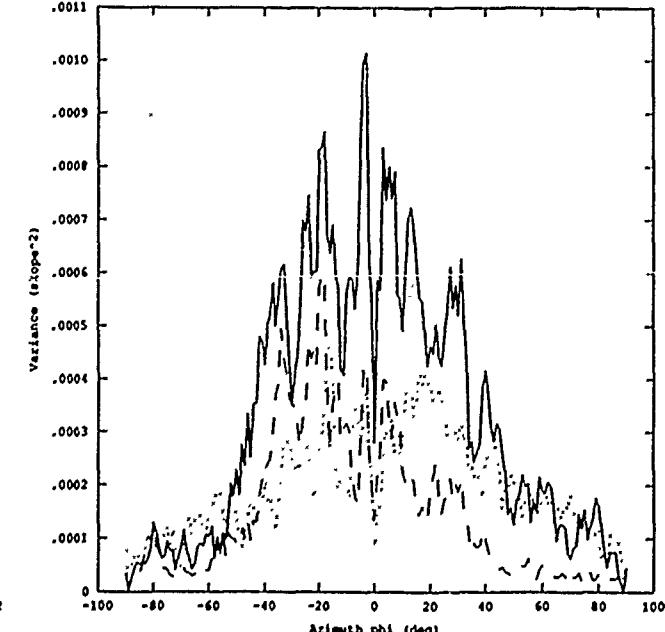
CASE 67

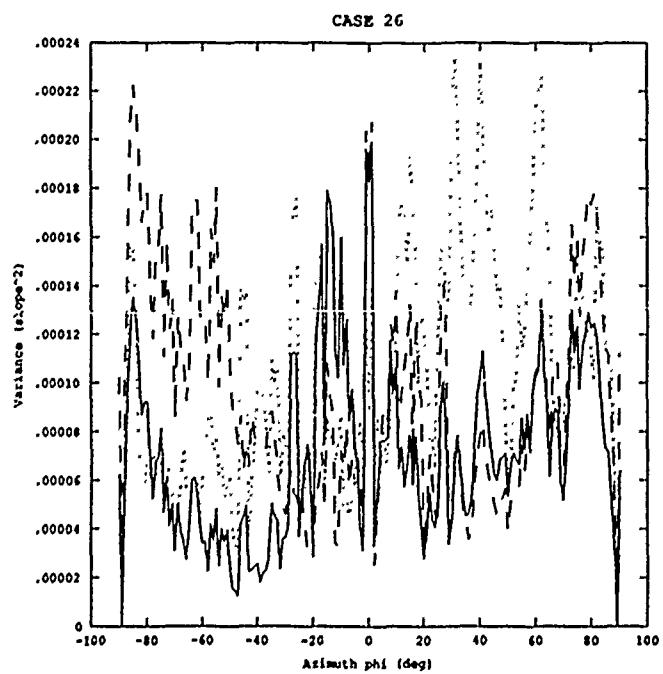
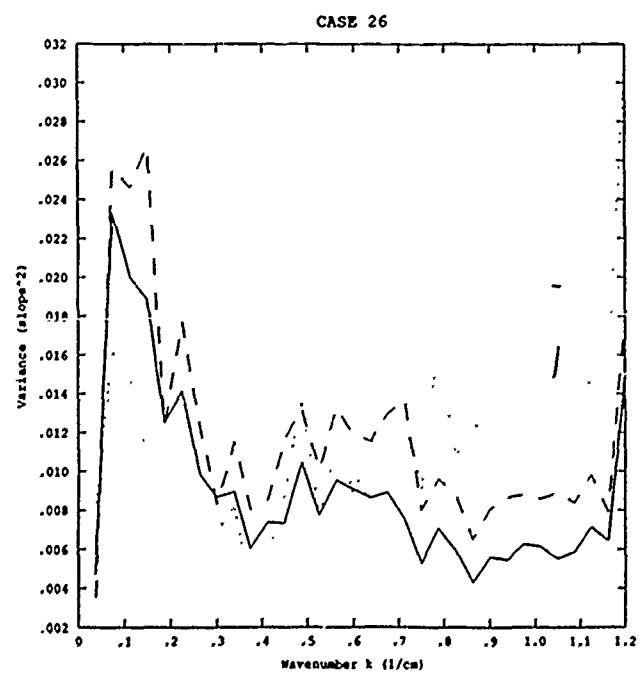
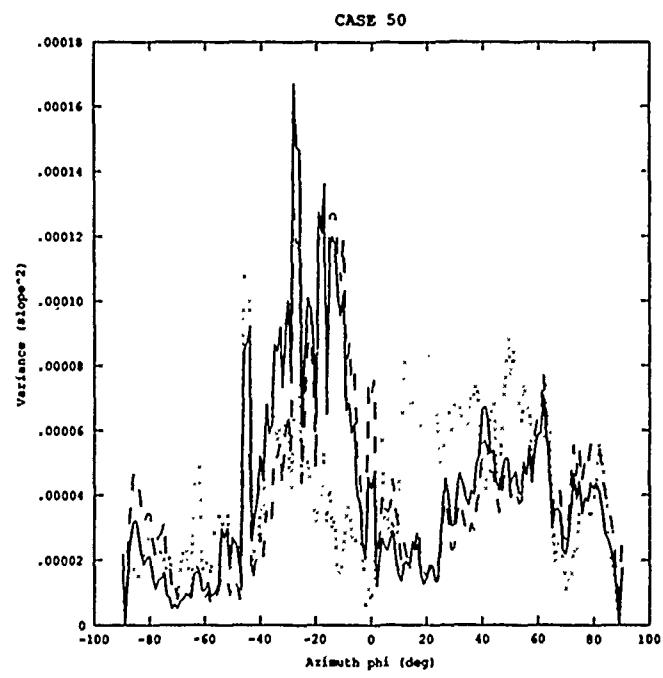
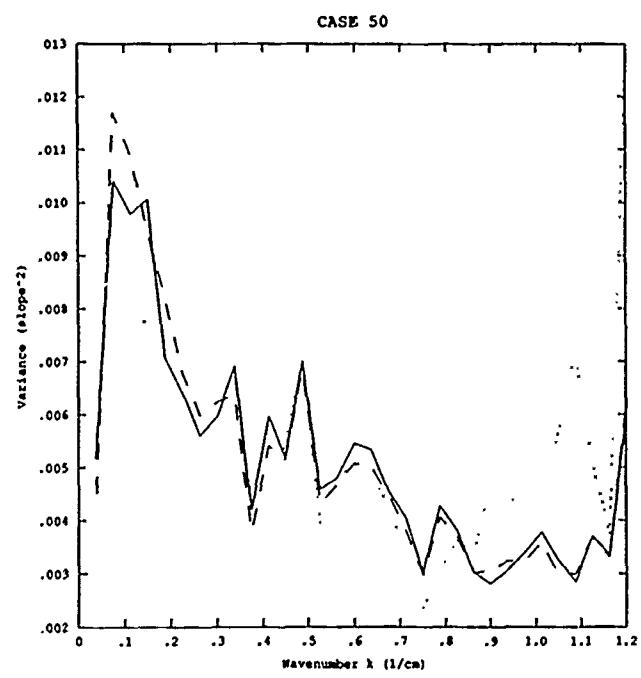
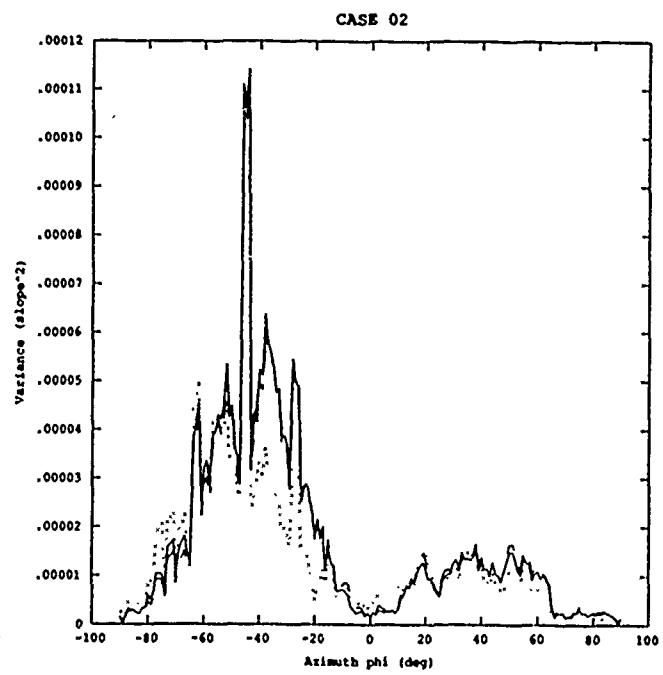
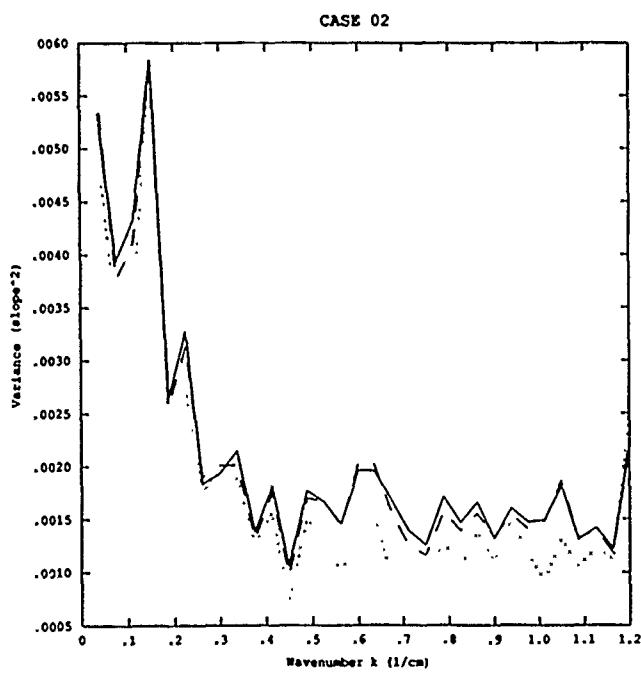


CASE 43

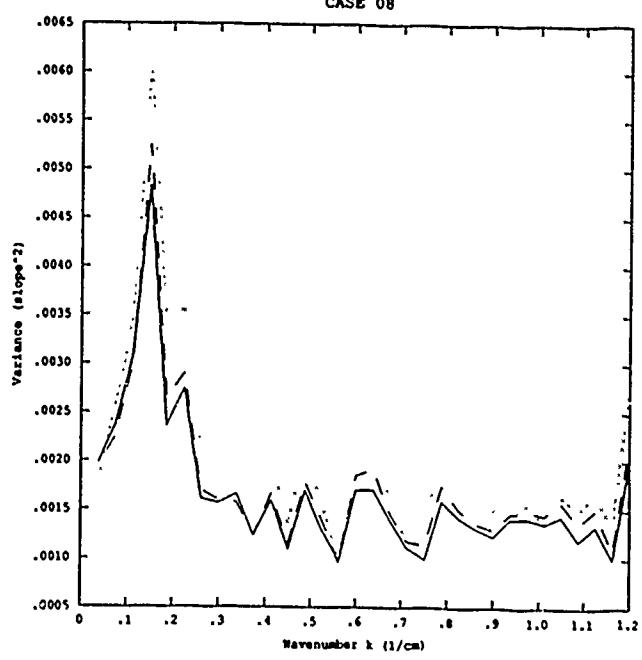


CASE 43

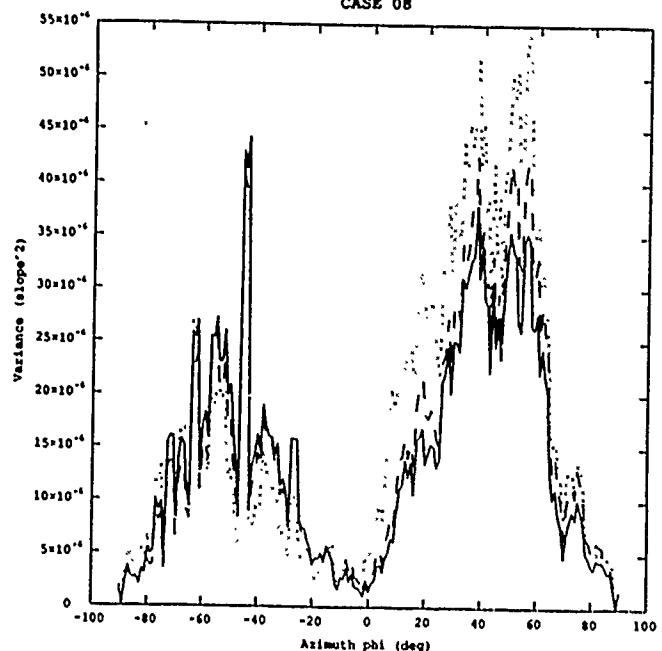




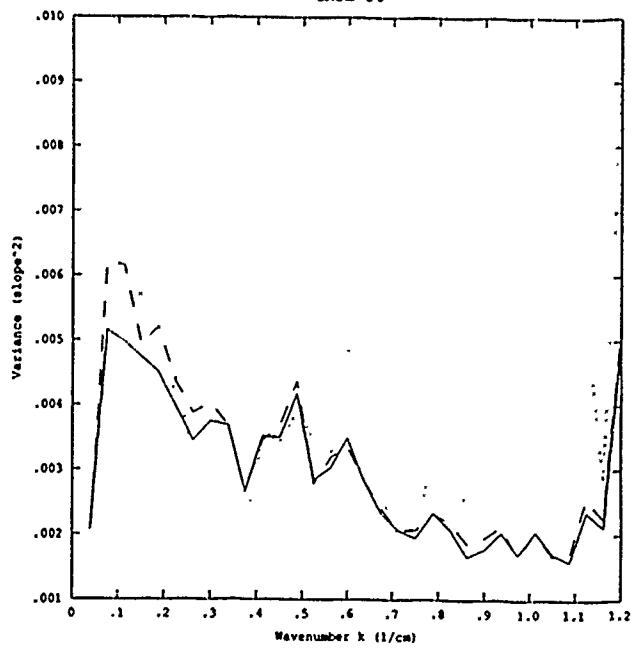
CASE 08



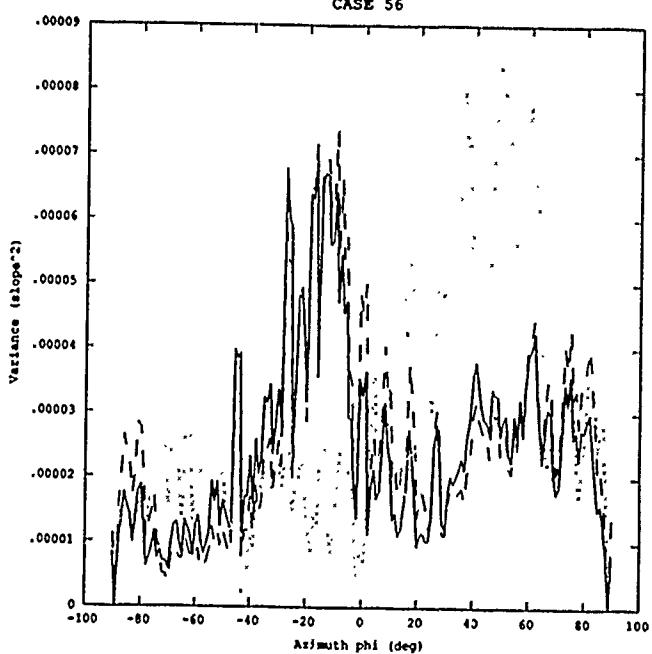
CASE 08



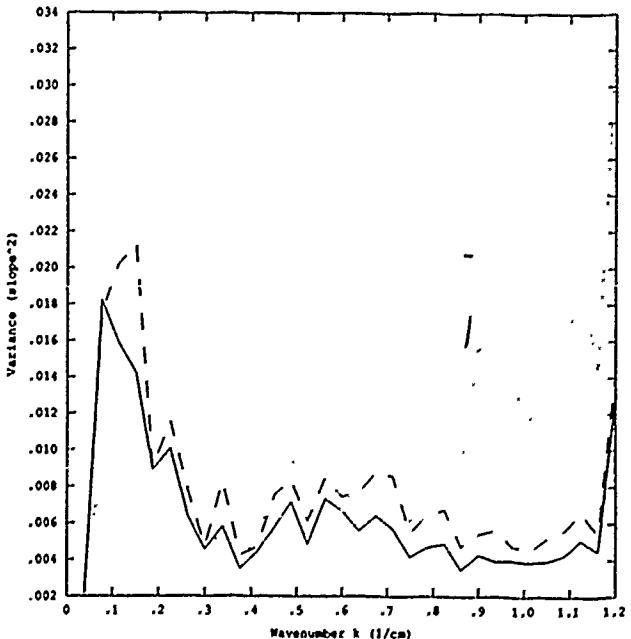
CASE 56



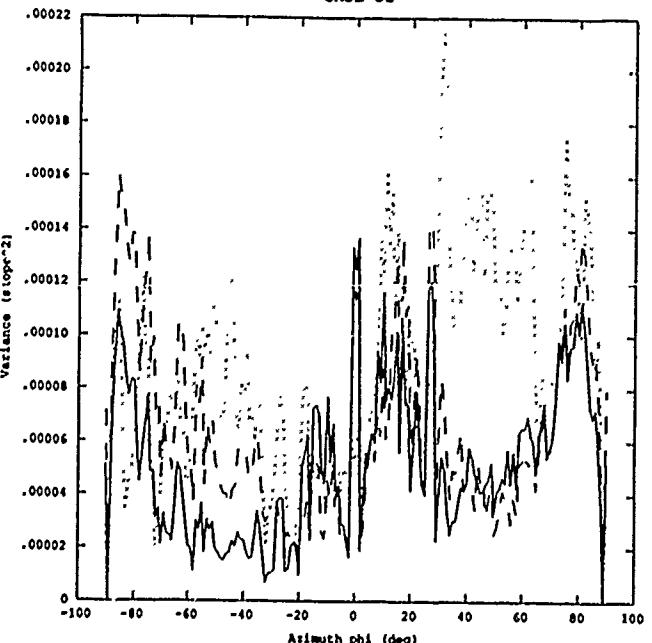
CASE 56



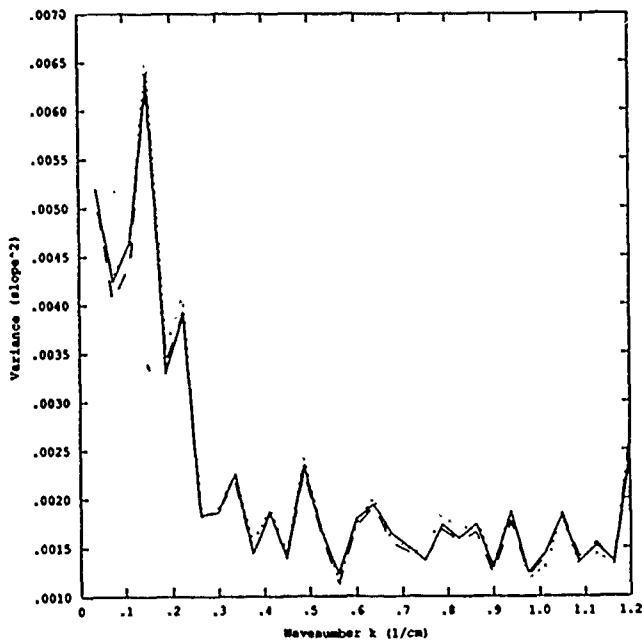
CASE 32



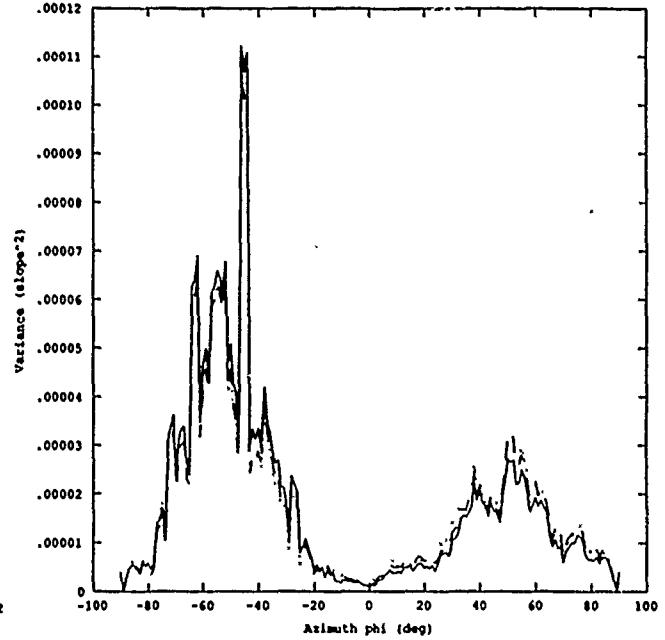
CASE 32



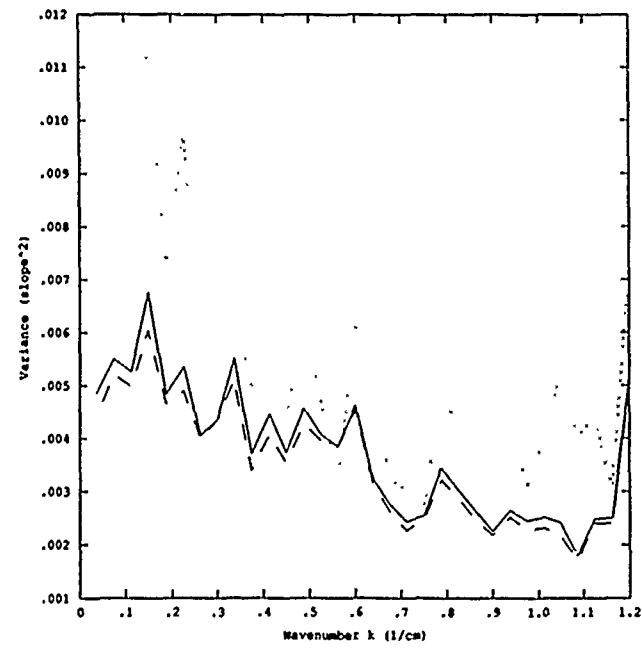
CASE 14



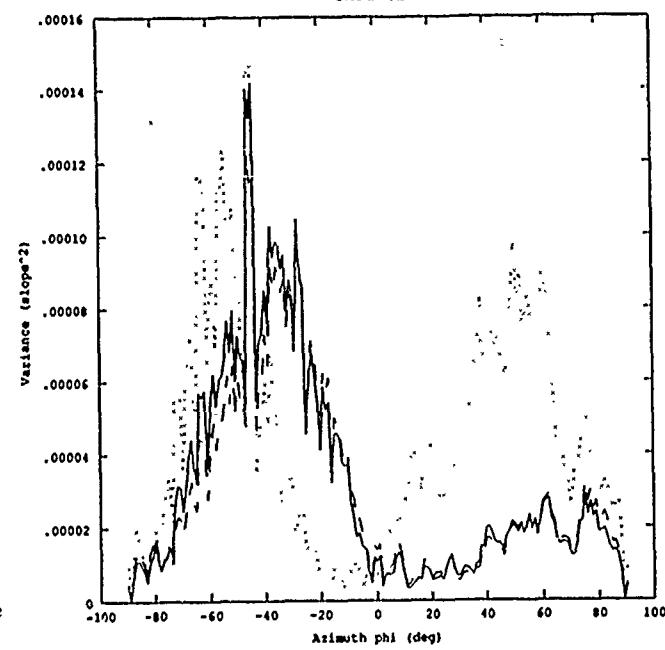
CASE 14



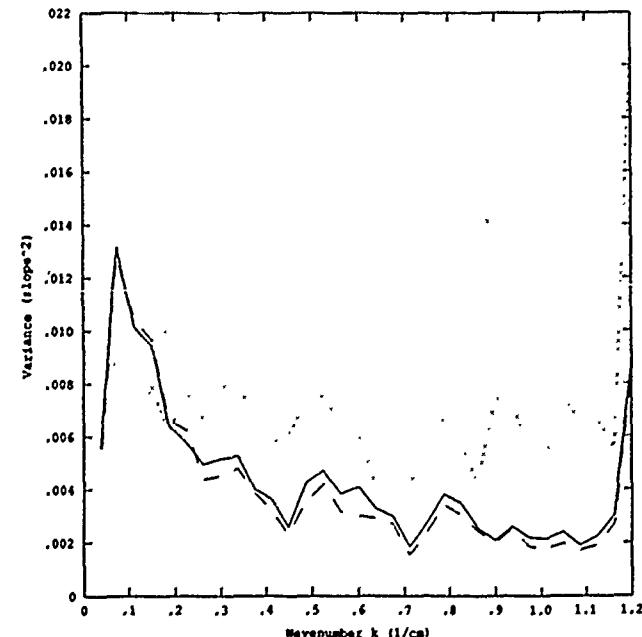
CASE 62



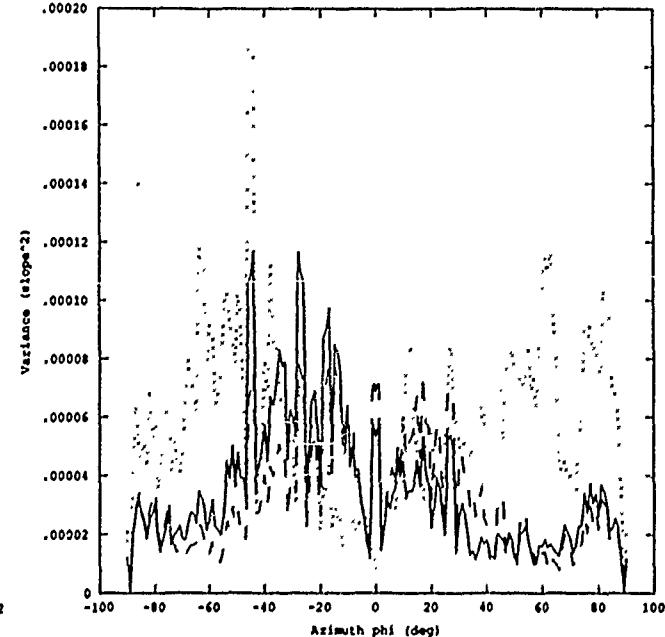
CASE 62

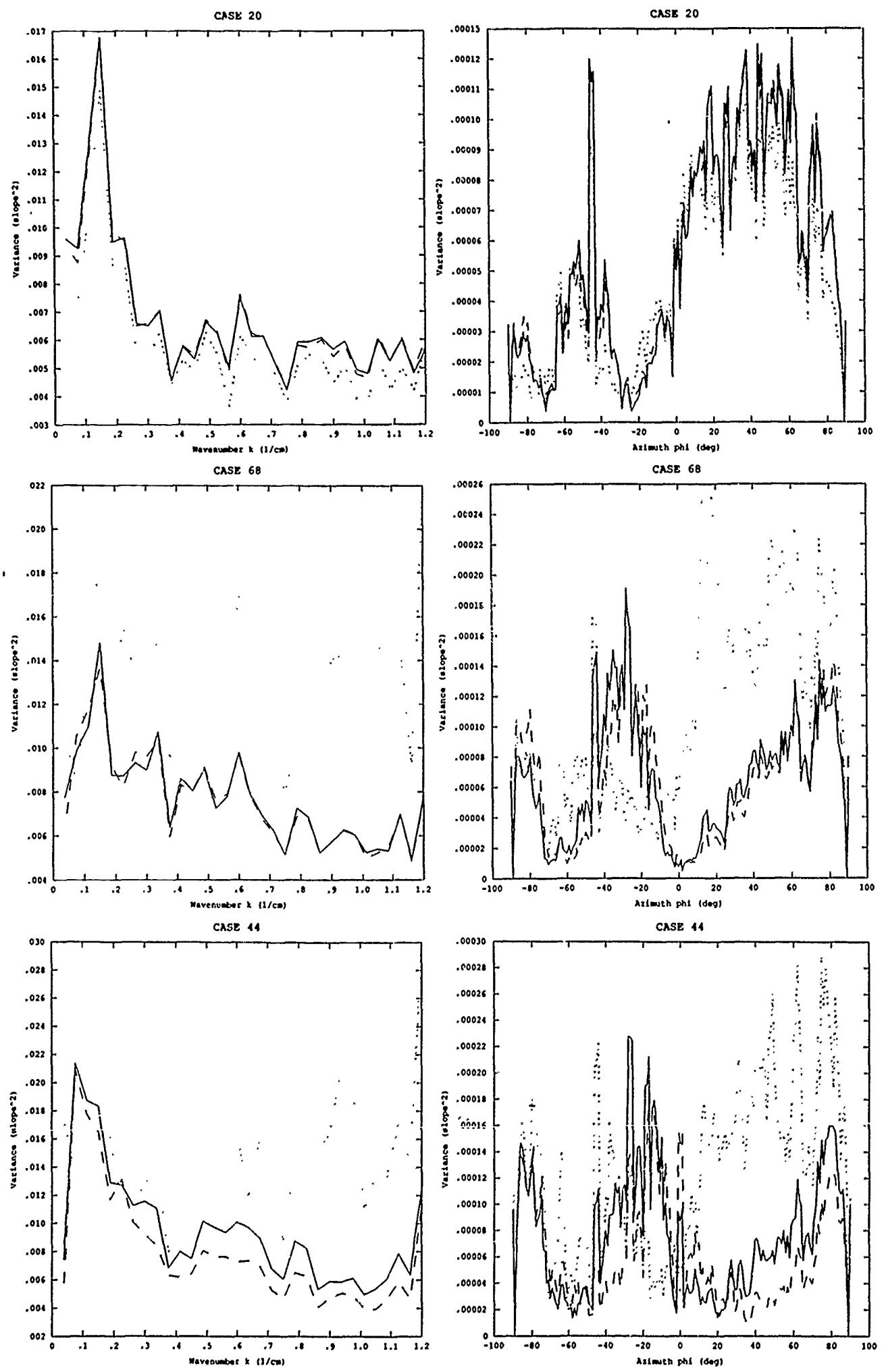


CASE 38

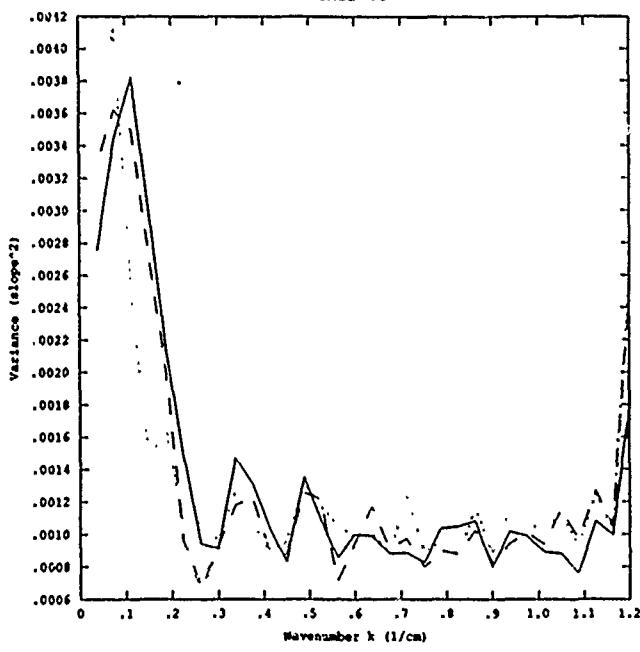


CASE 38

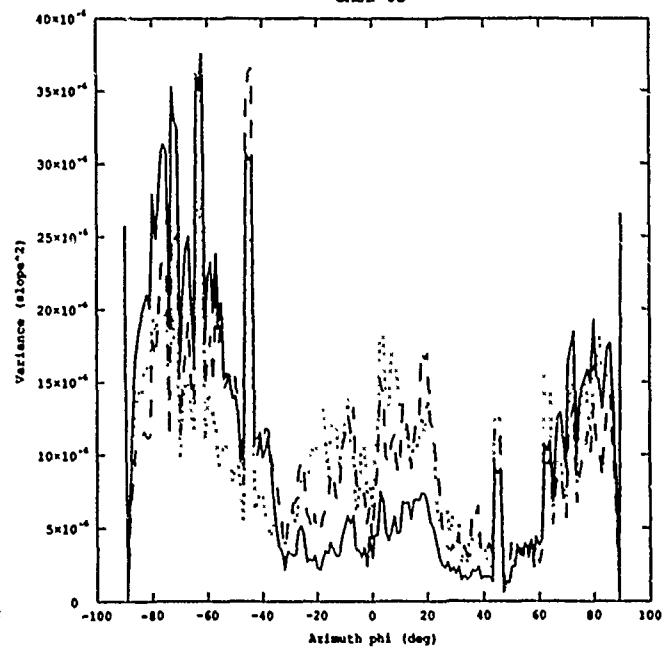




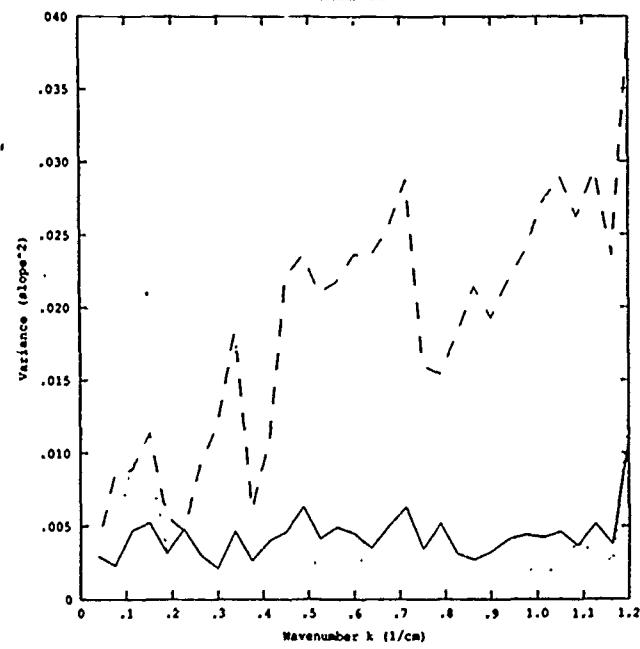
CASE 03



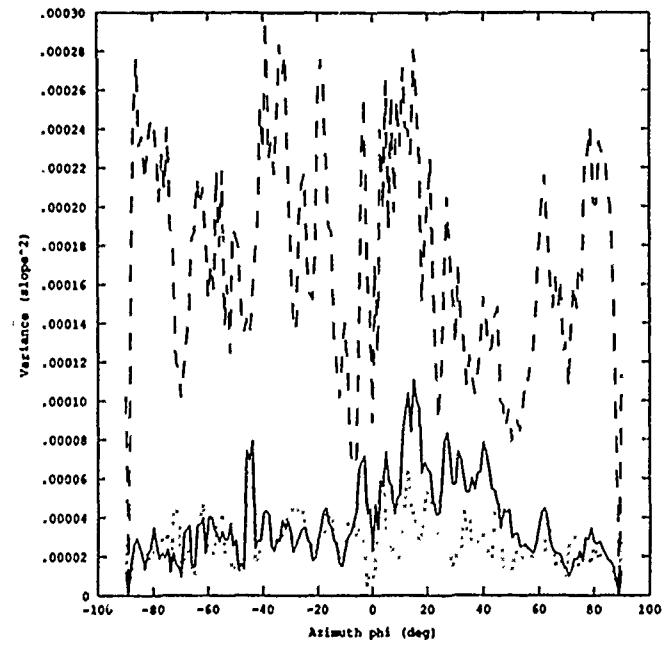
CASE 03



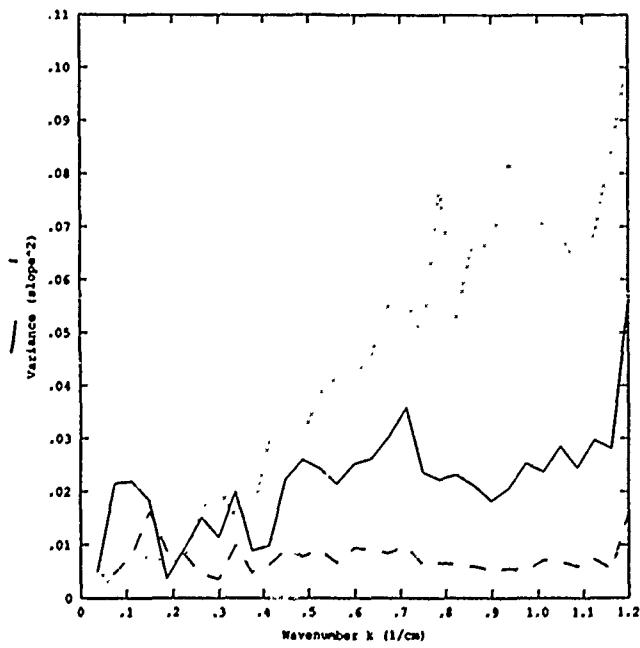
CASE 51



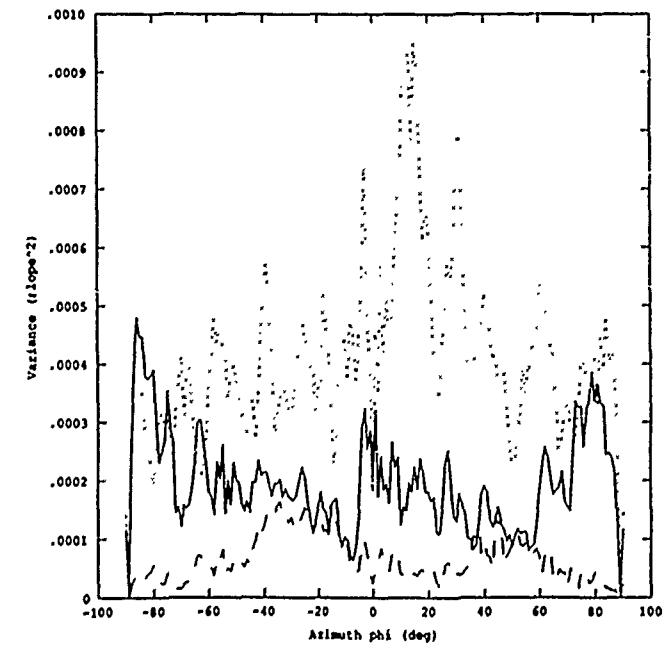
CASE 51

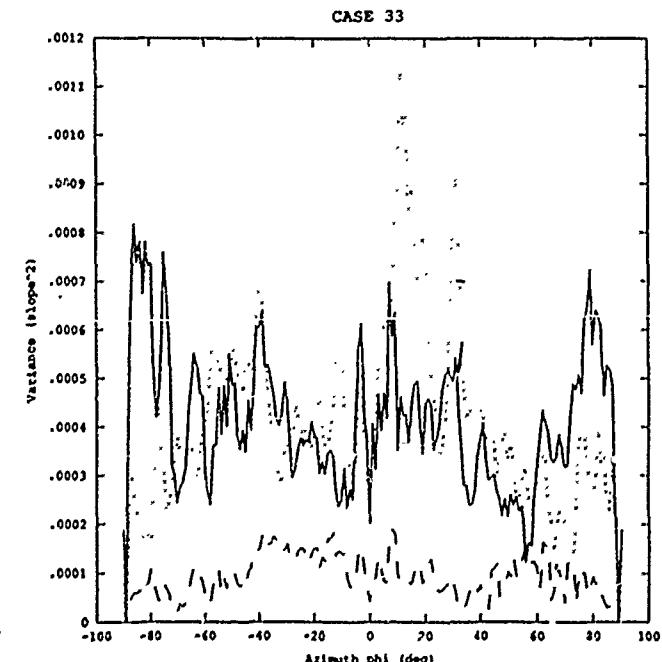
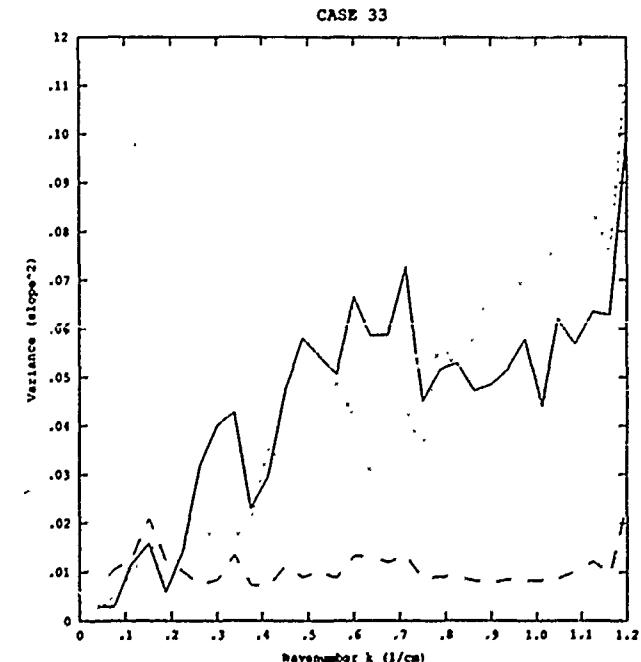
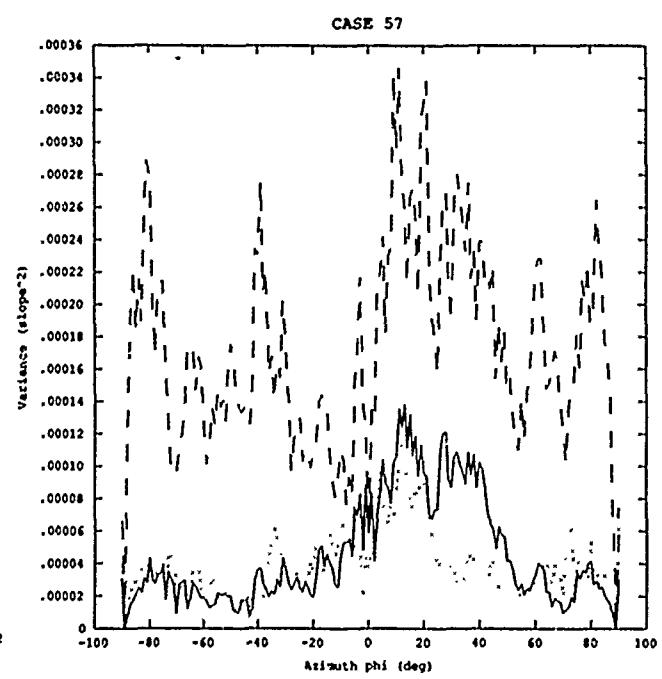
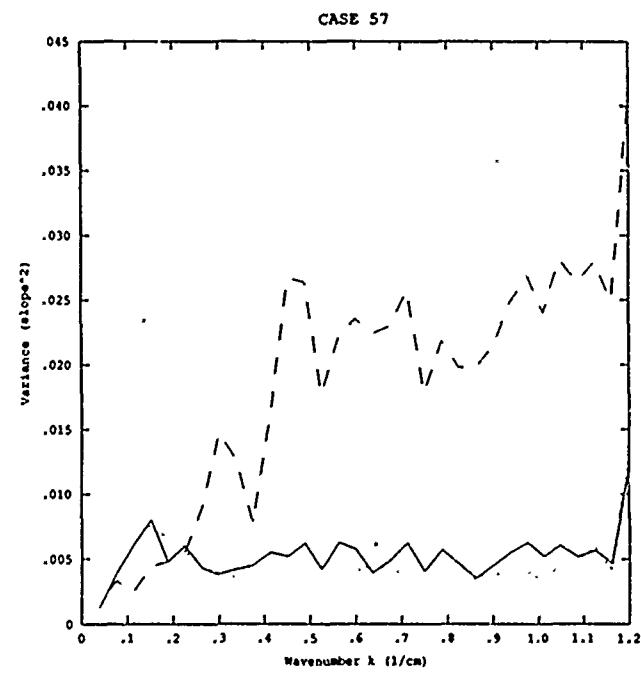
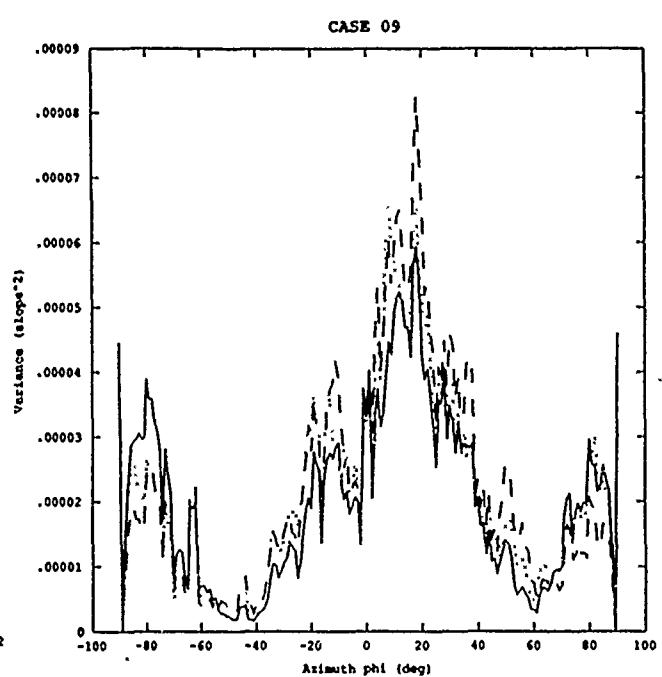
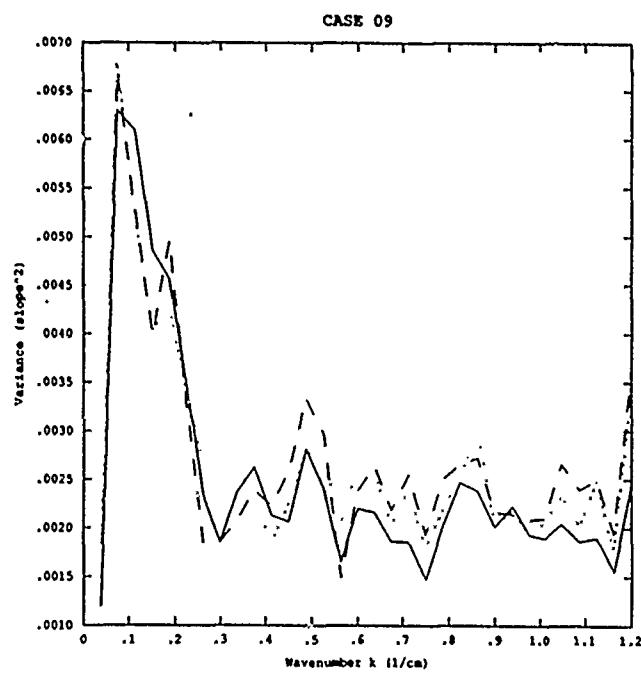


CASE 27

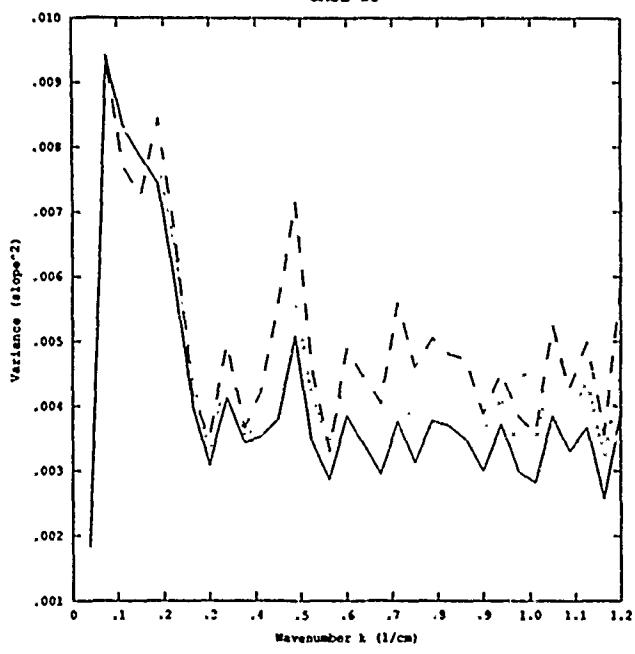


CASE 27

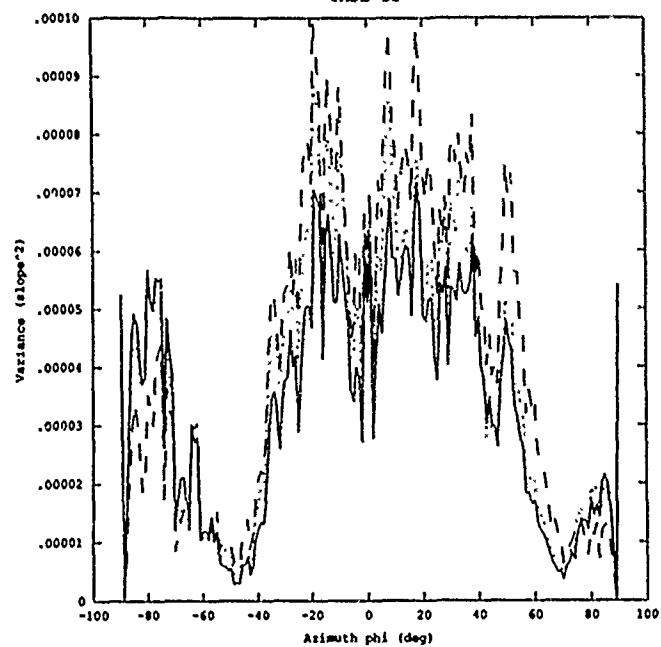




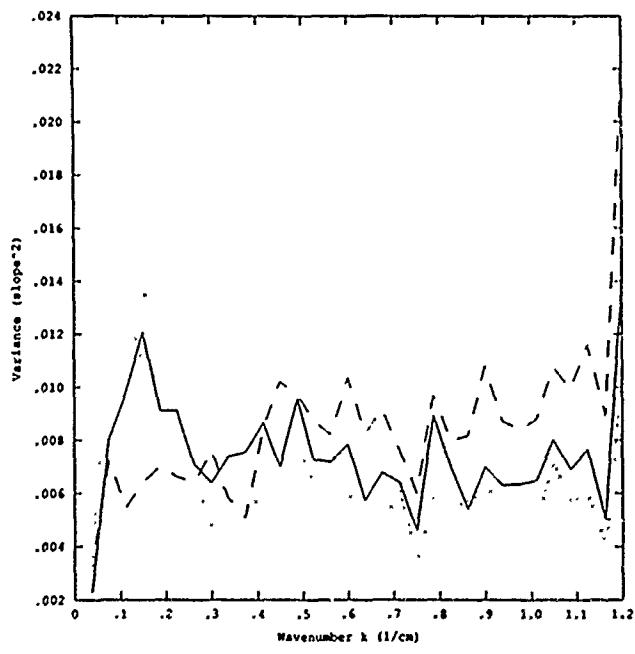
CASE 15



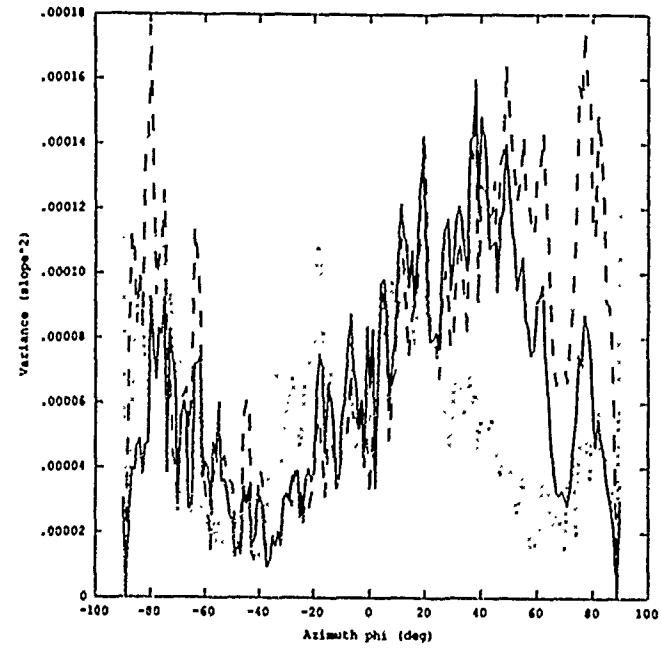
CASE 15



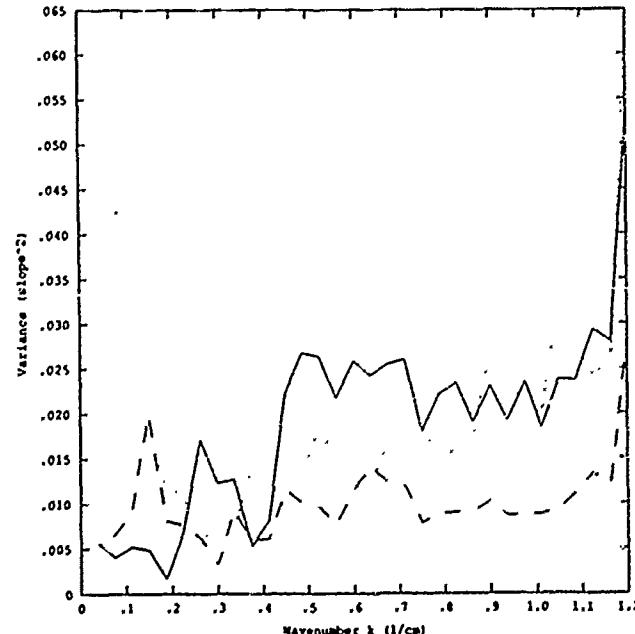
CASE 63



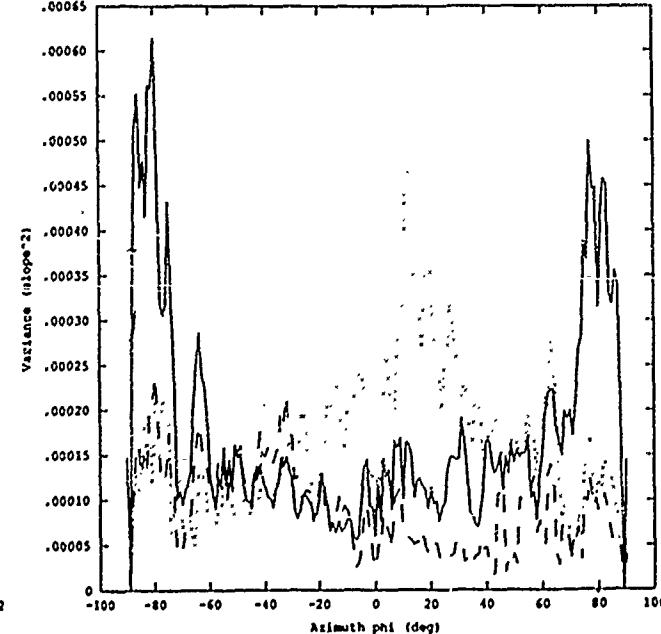
CASE 63



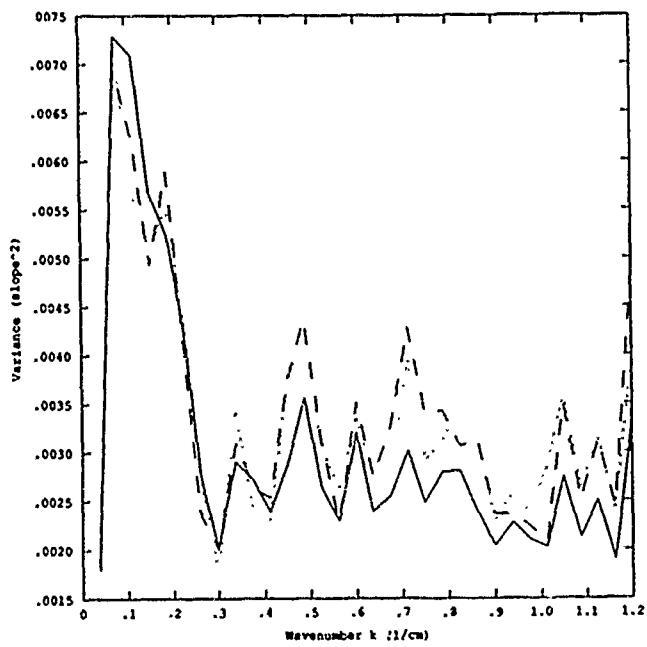
CASE 39



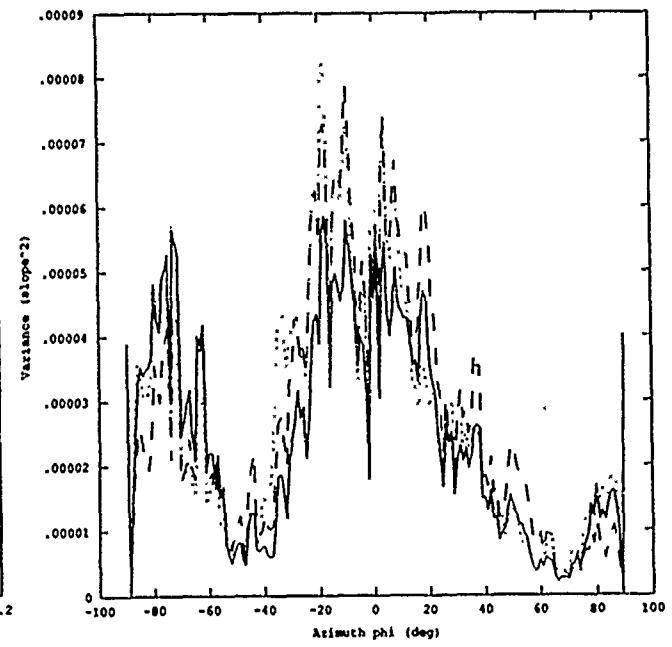
CASE 39



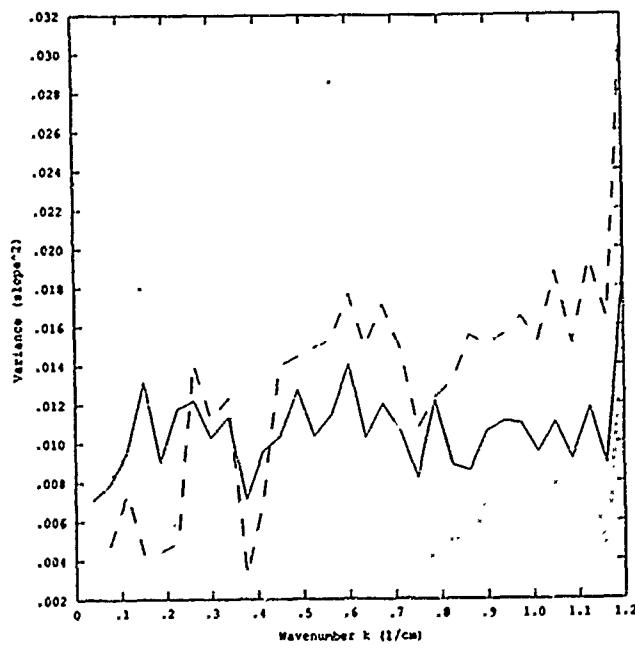
CASE 21



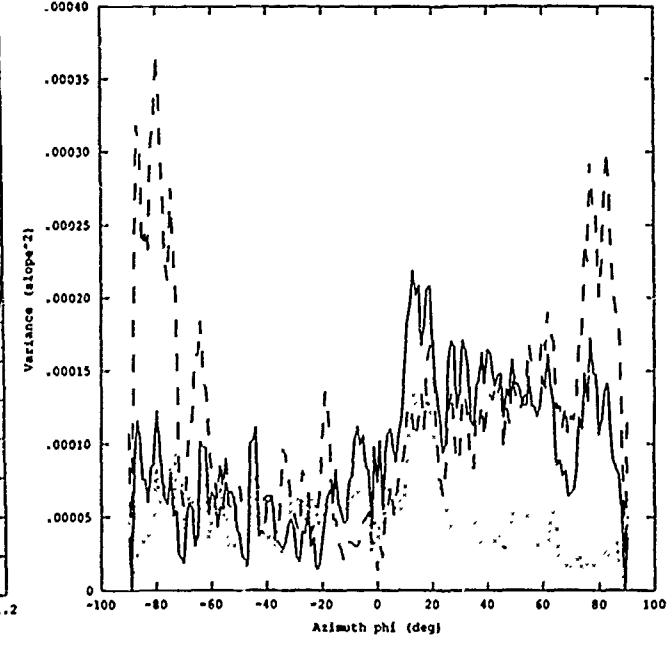
CASE 21



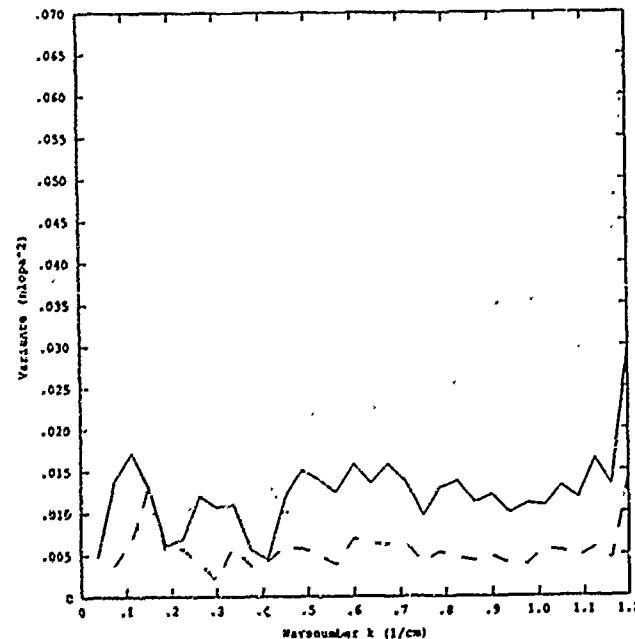
CASE 69



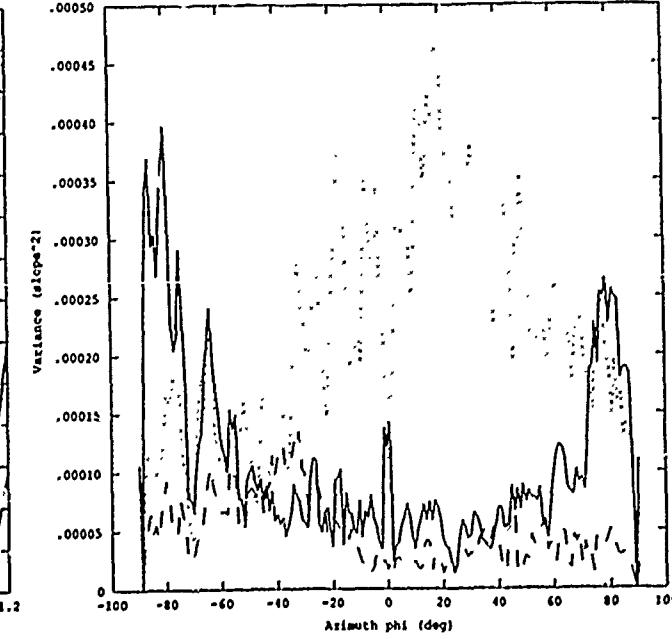
CASE 69

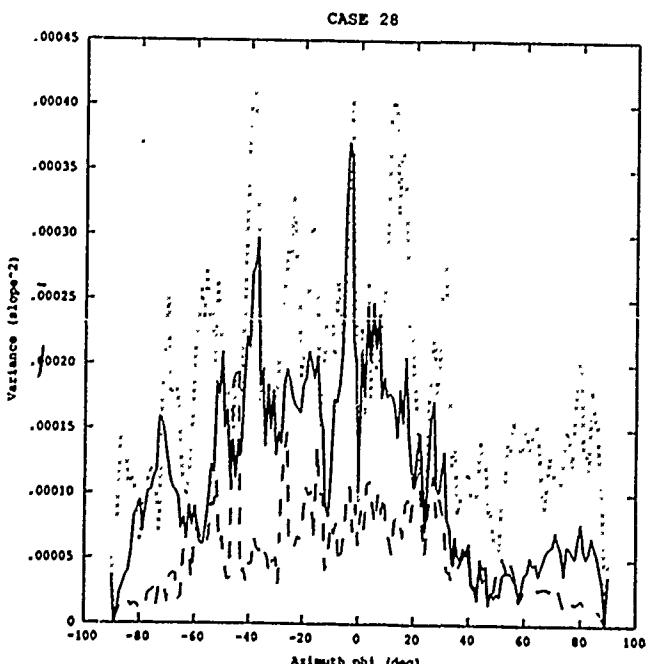
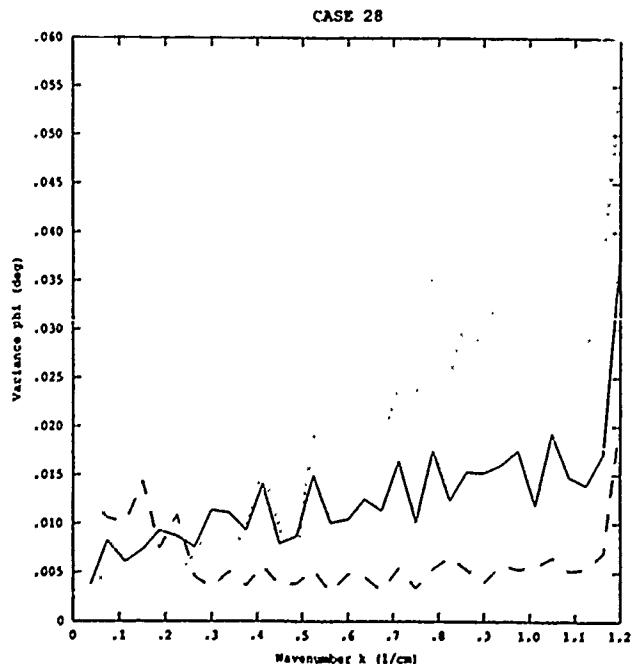
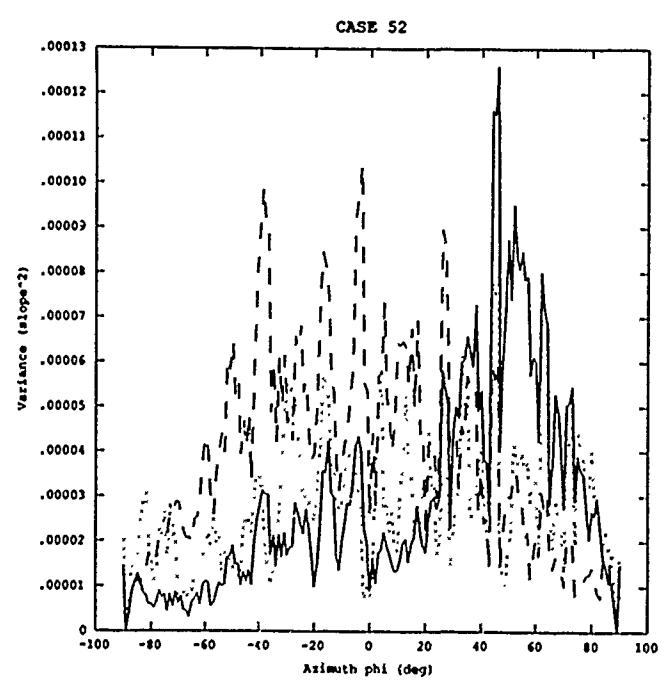
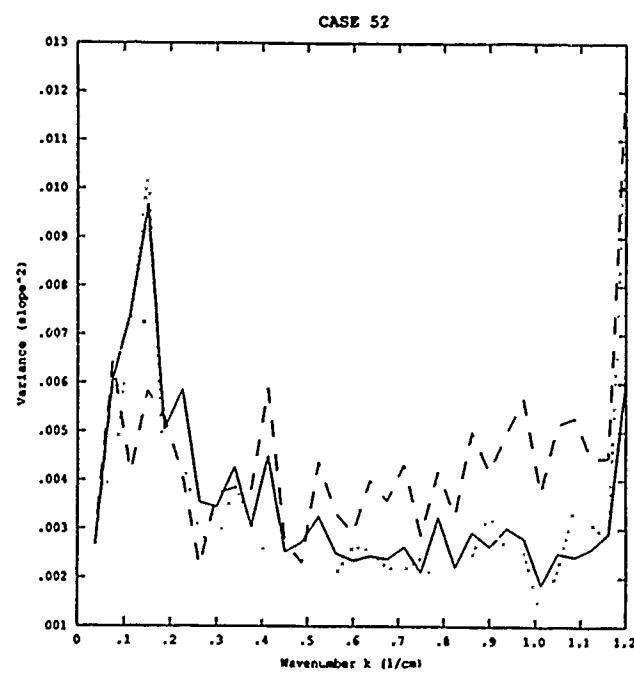
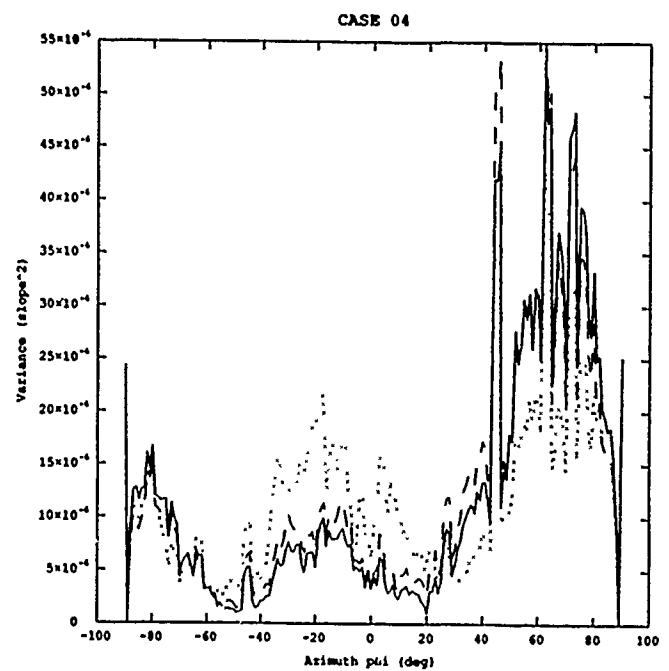
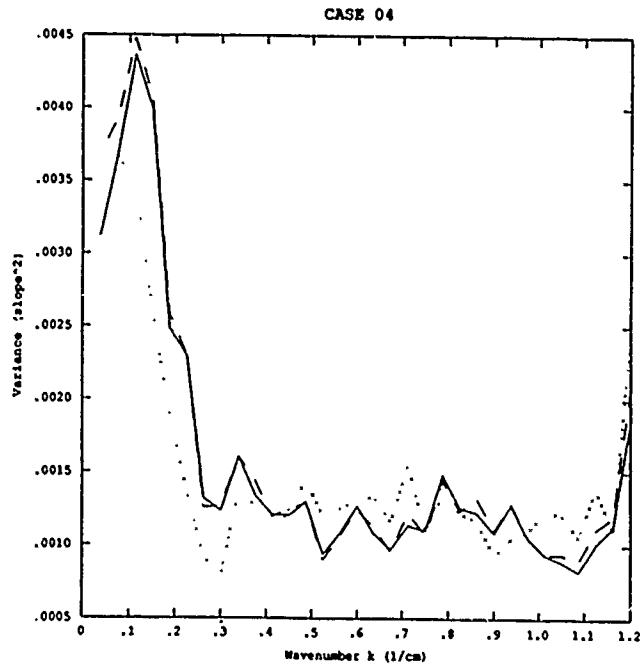


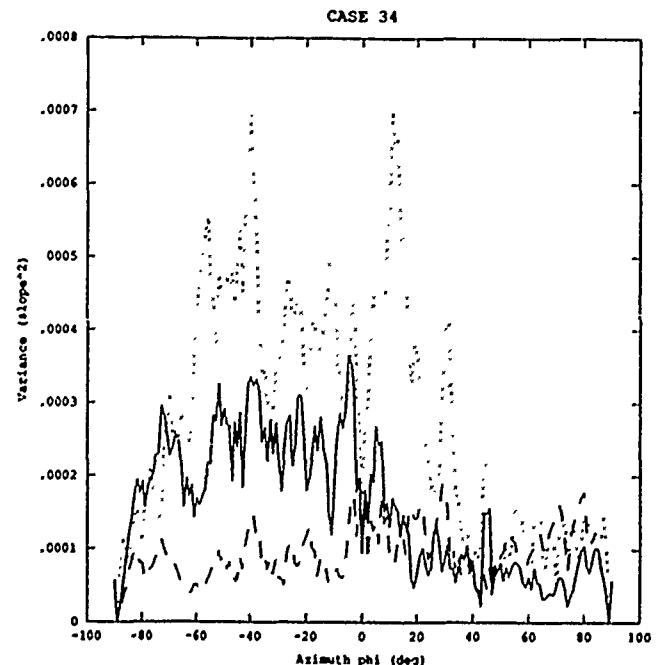
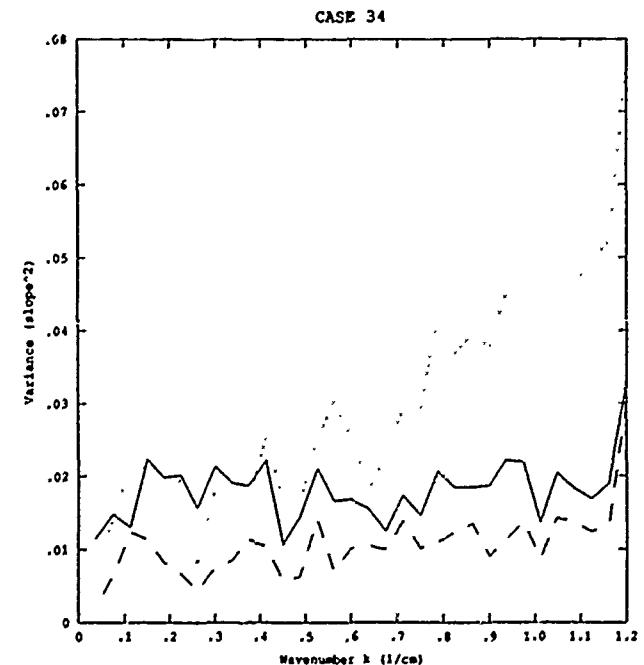
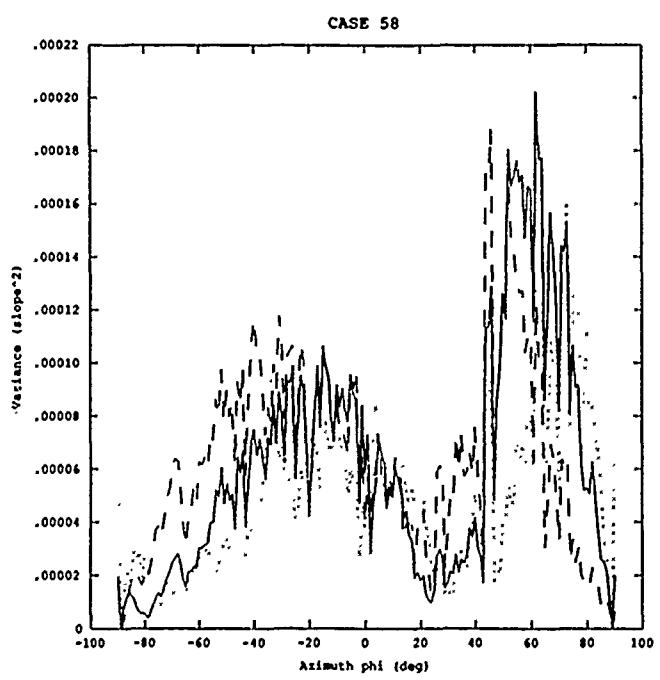
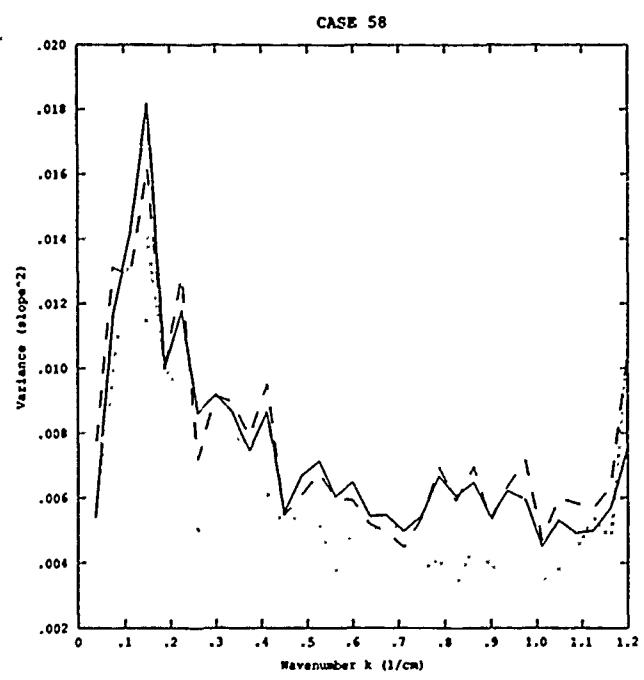
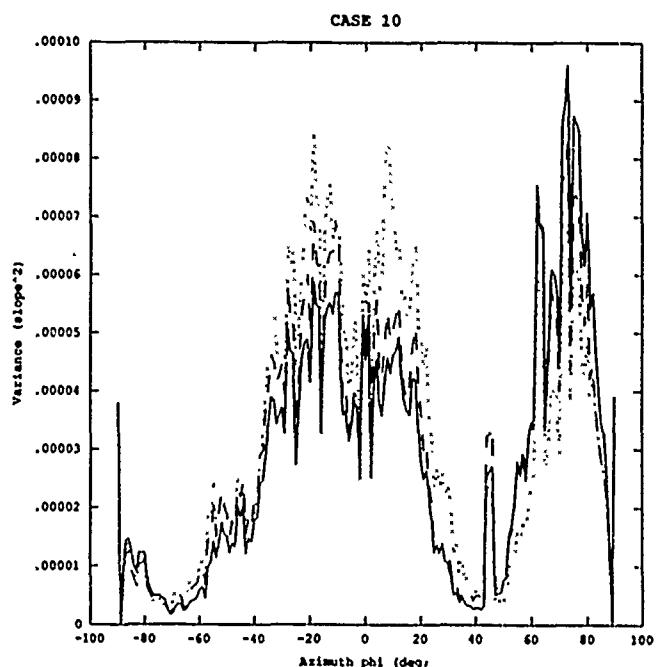
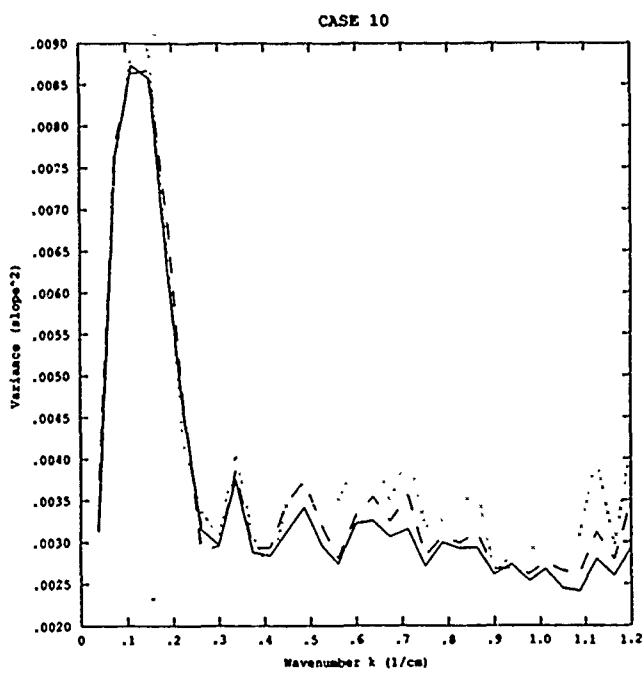
CASE 45

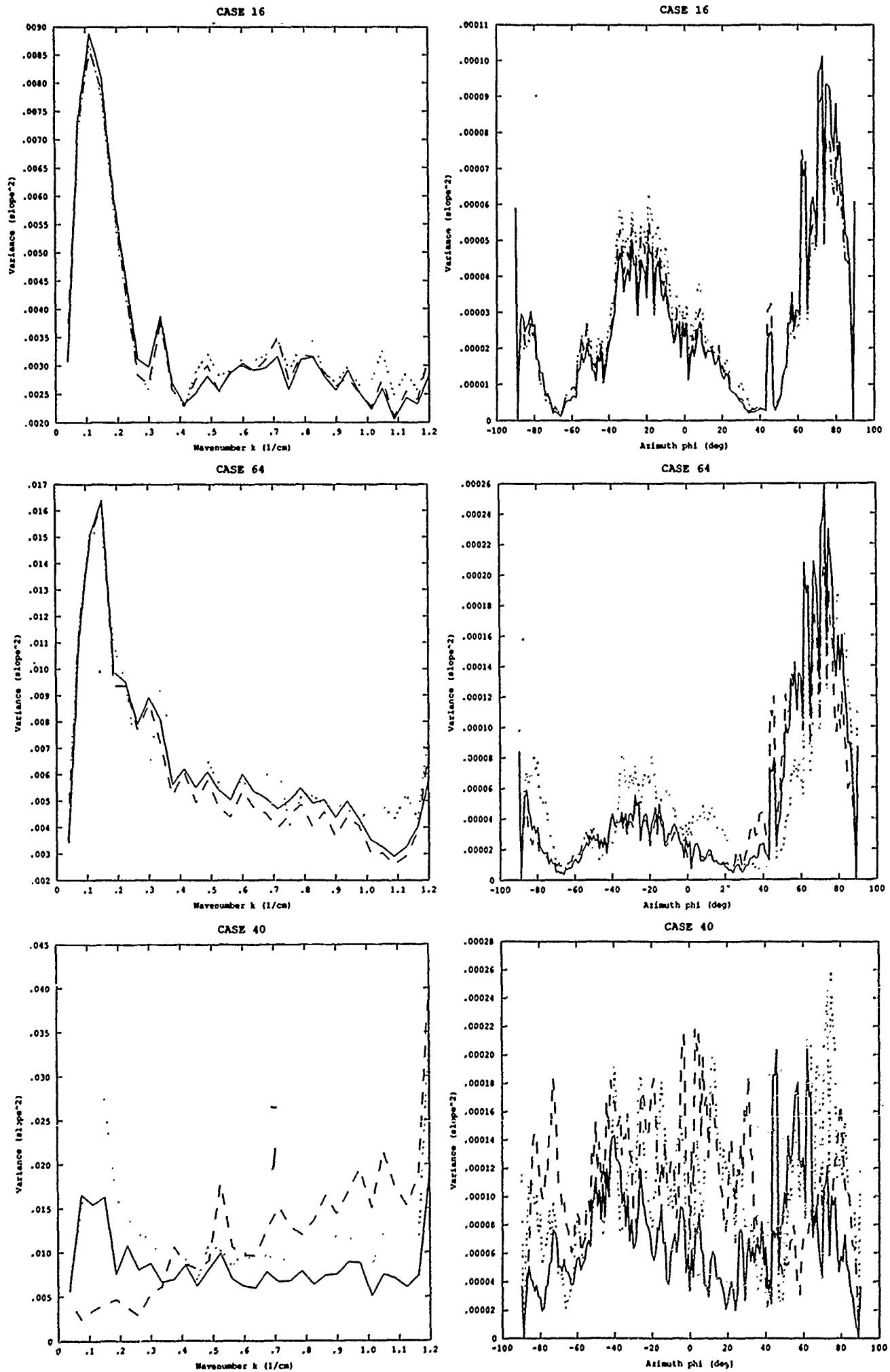


CASE 45

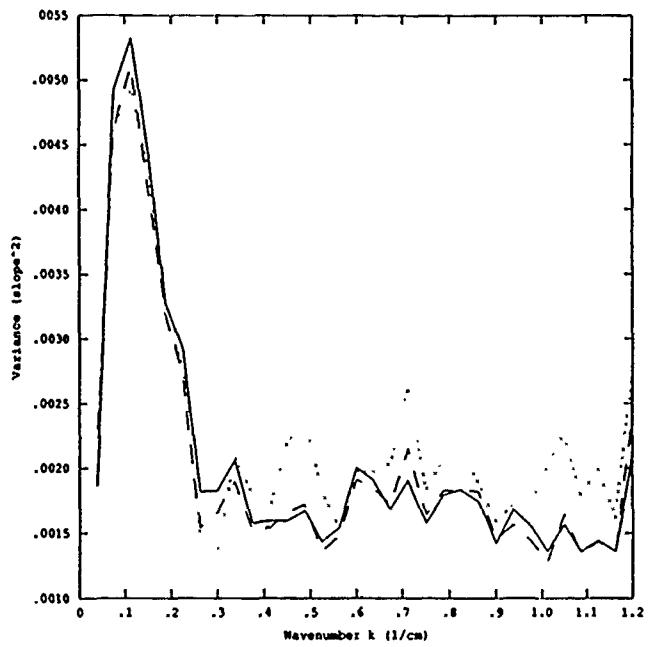




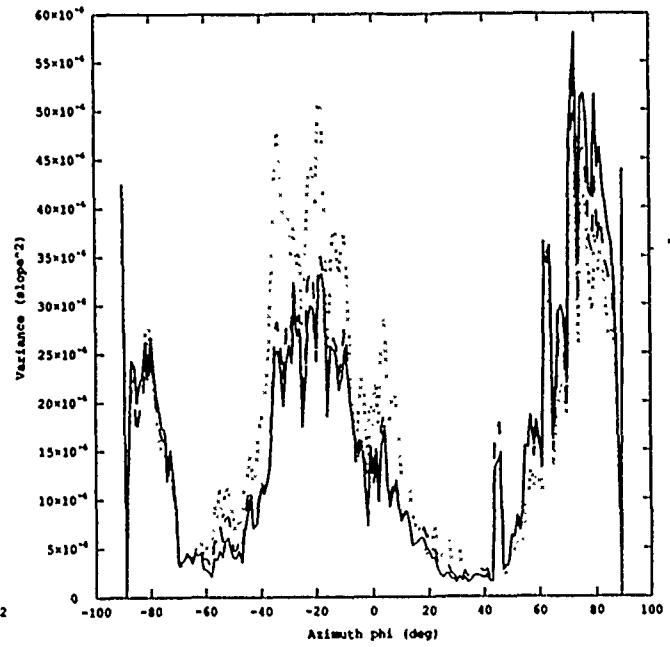




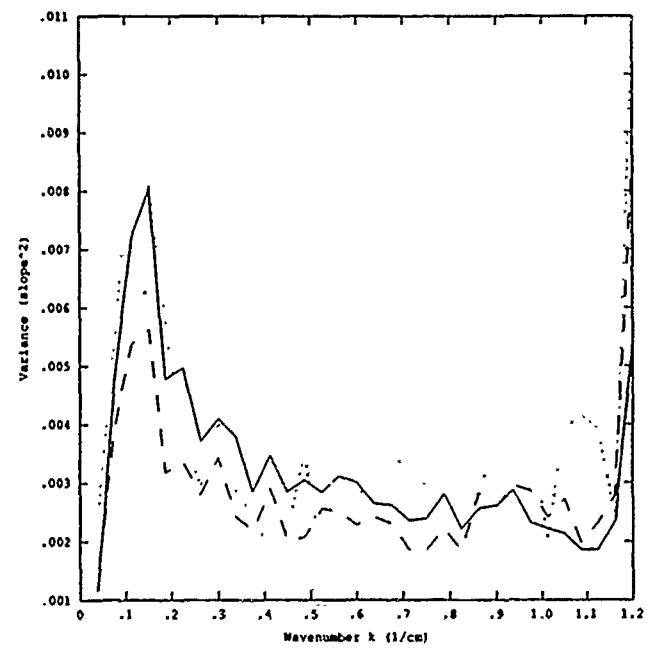
CASE 22



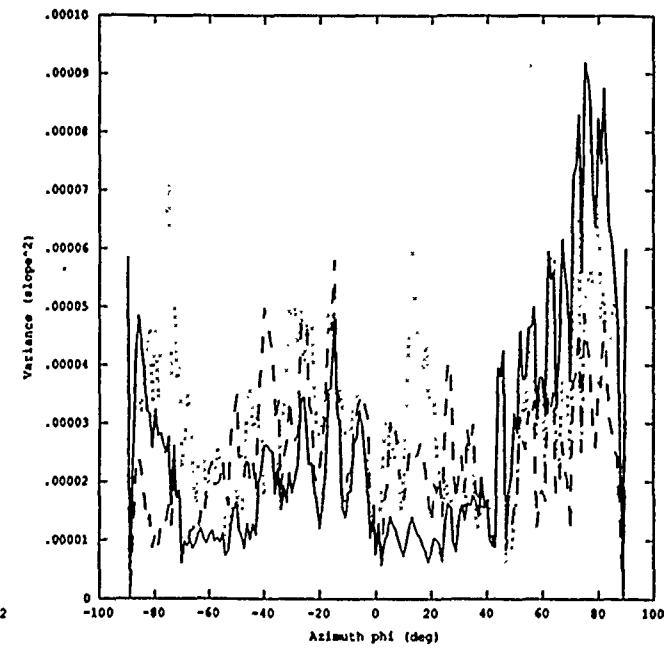
CASE 22



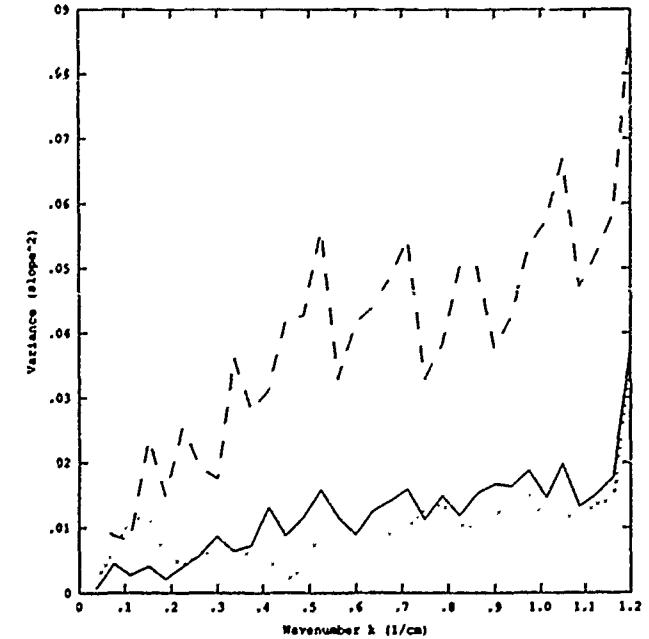
CASE 70



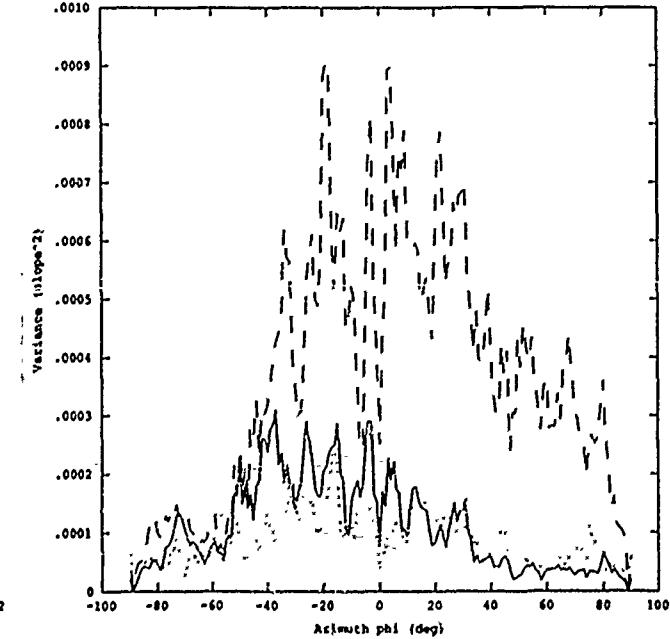
CASE 70



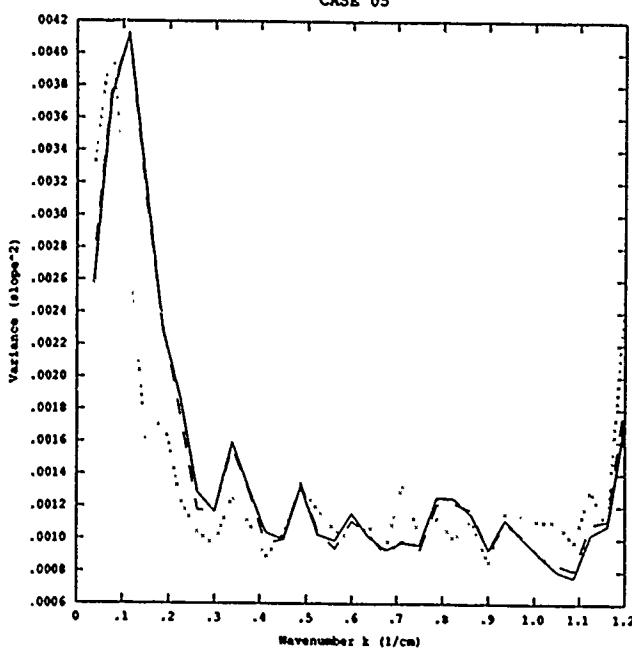
CASE 46



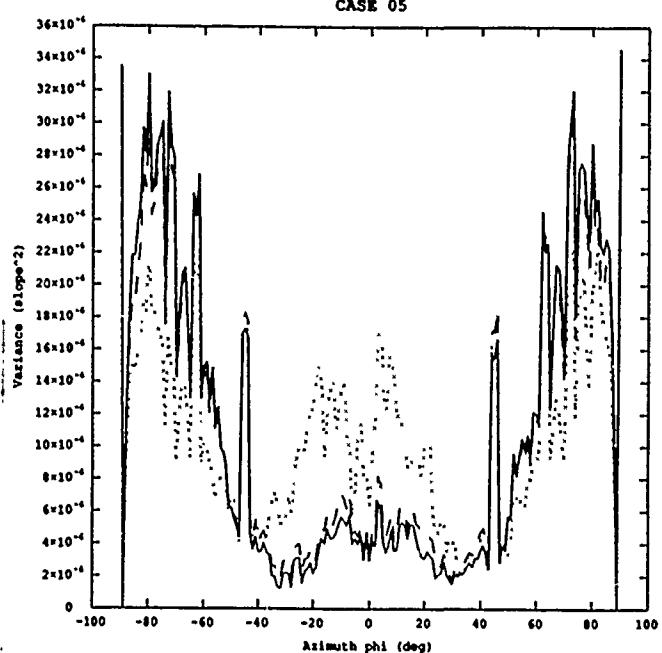
CASE 46



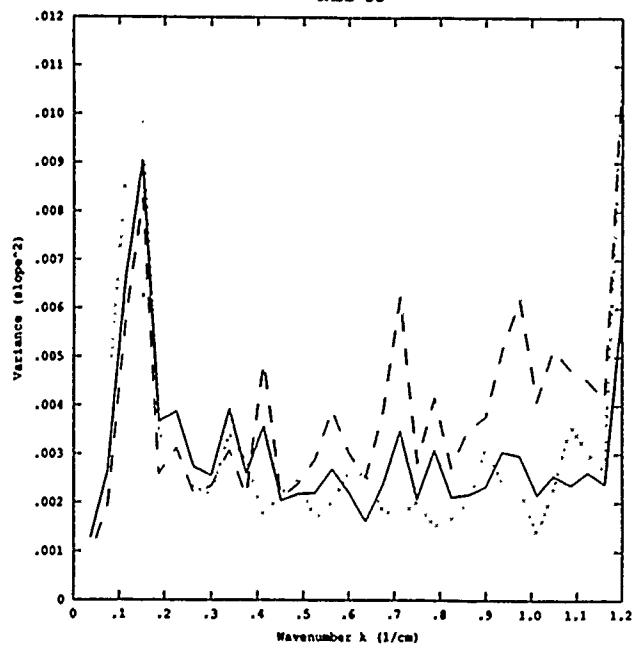
CASE 05



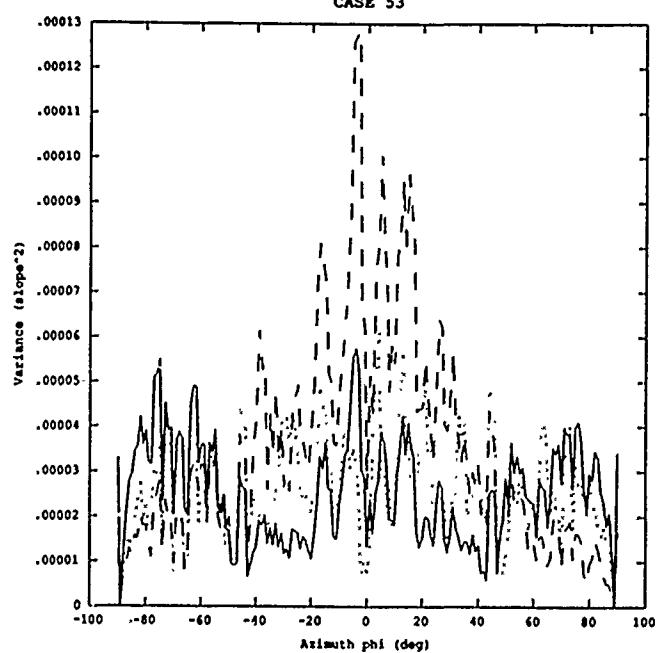
CASE 05



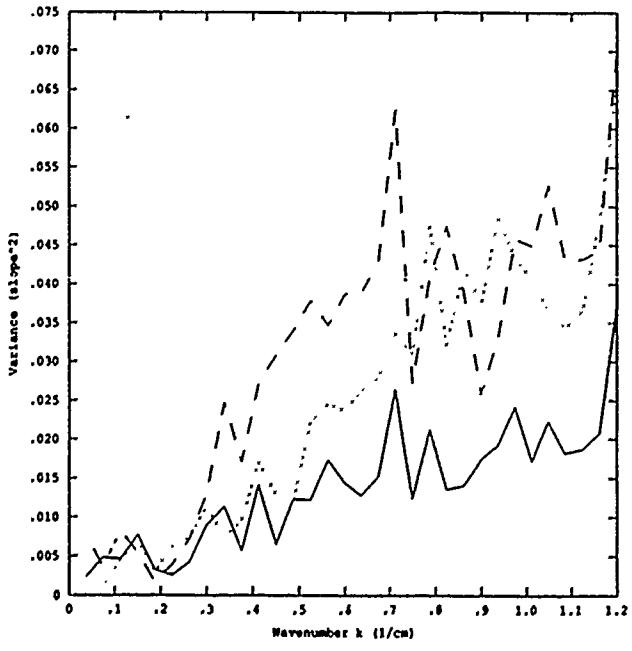
CASE 53



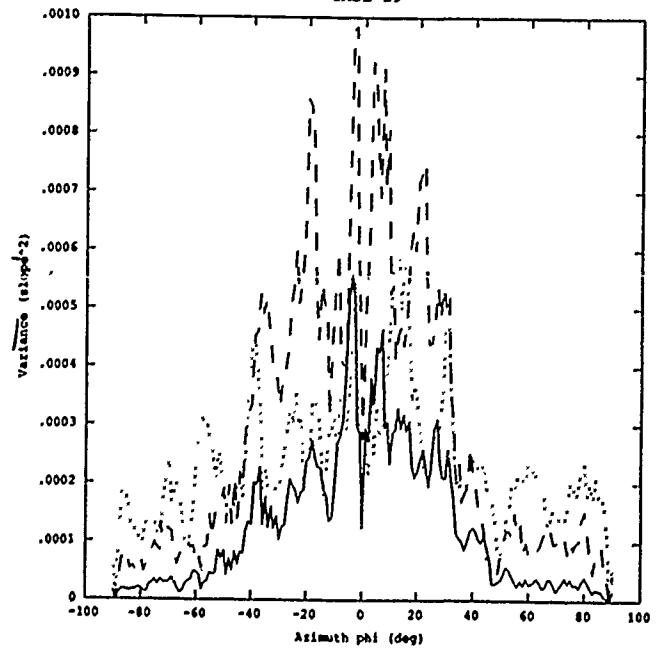
CASE 53

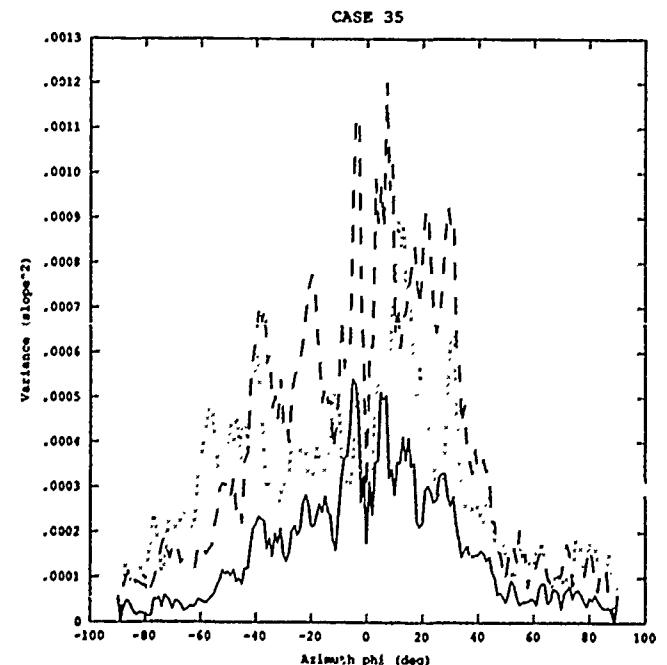
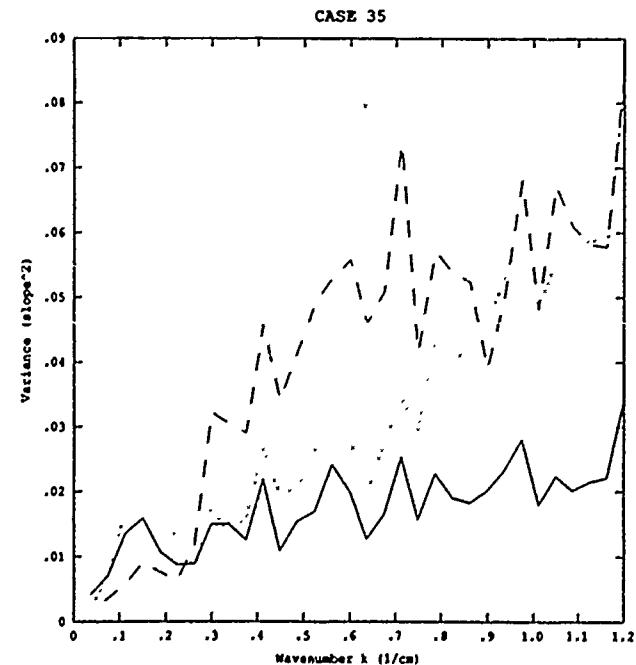
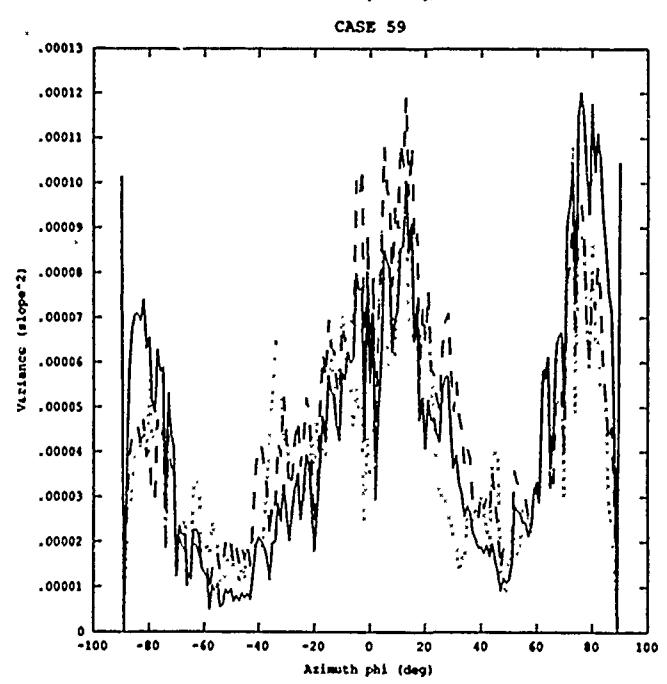
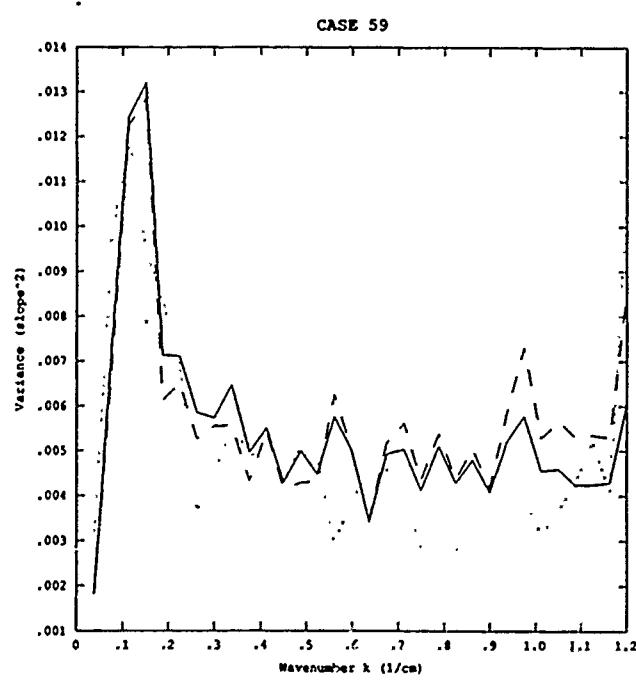
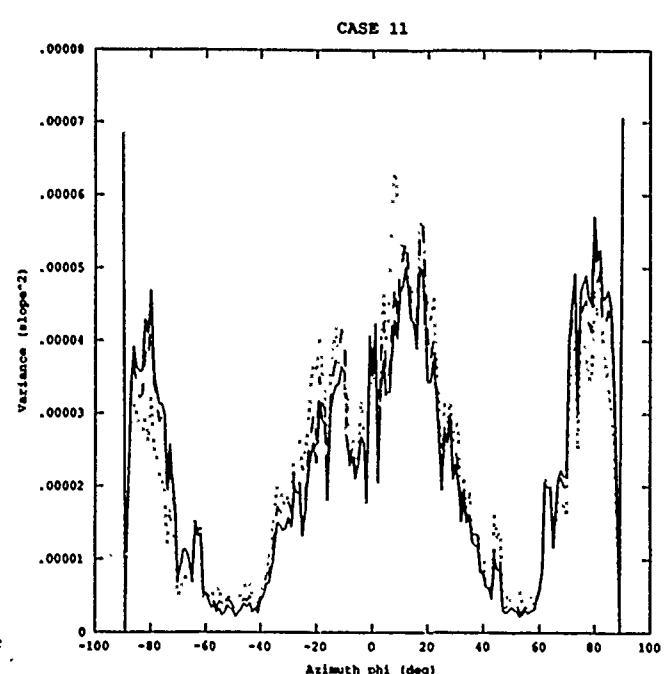
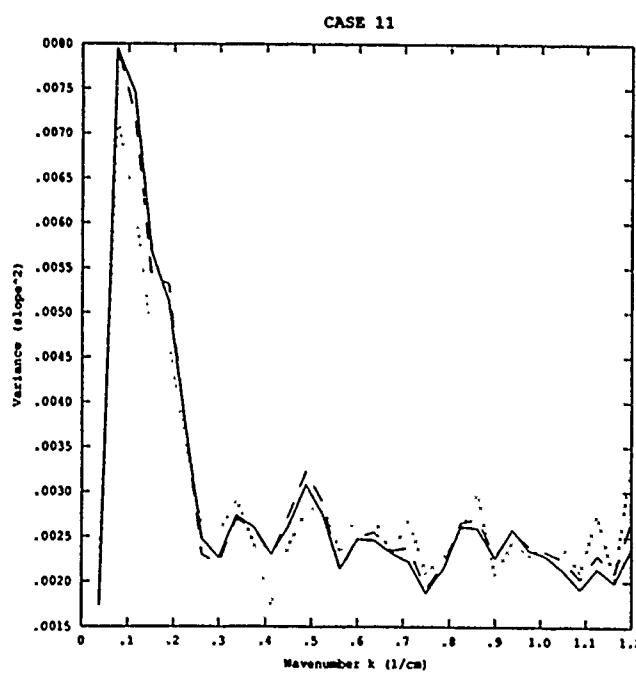


CASE 29

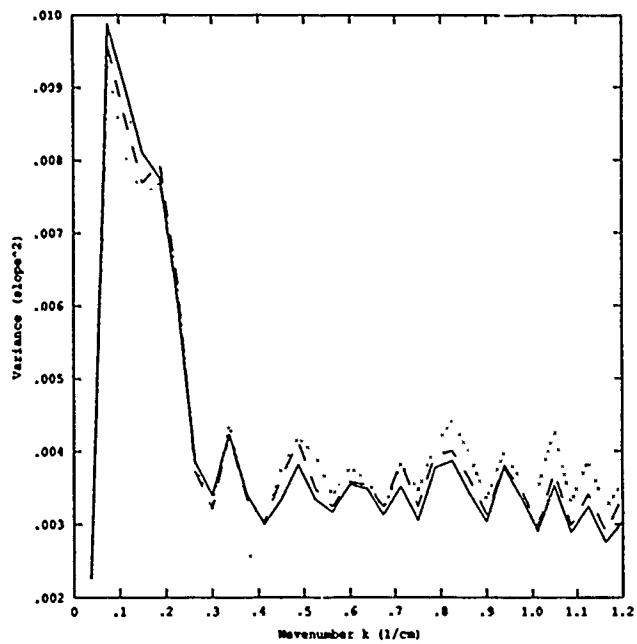


CASE 29

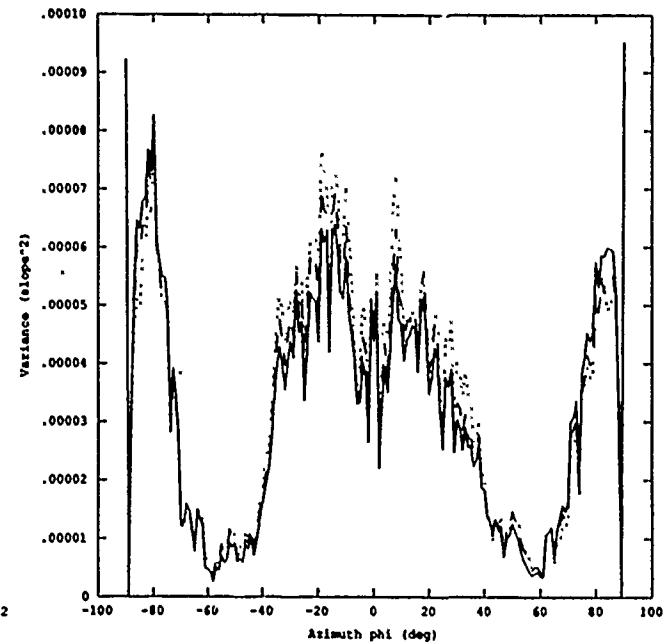




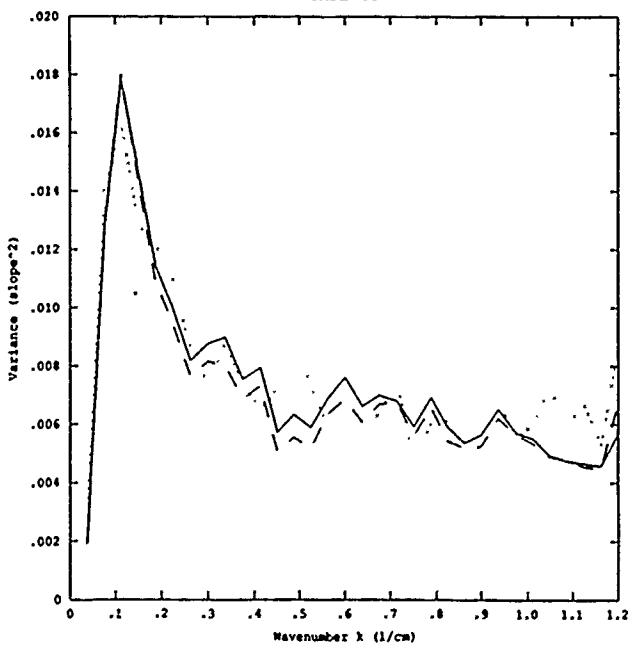
CASE 17



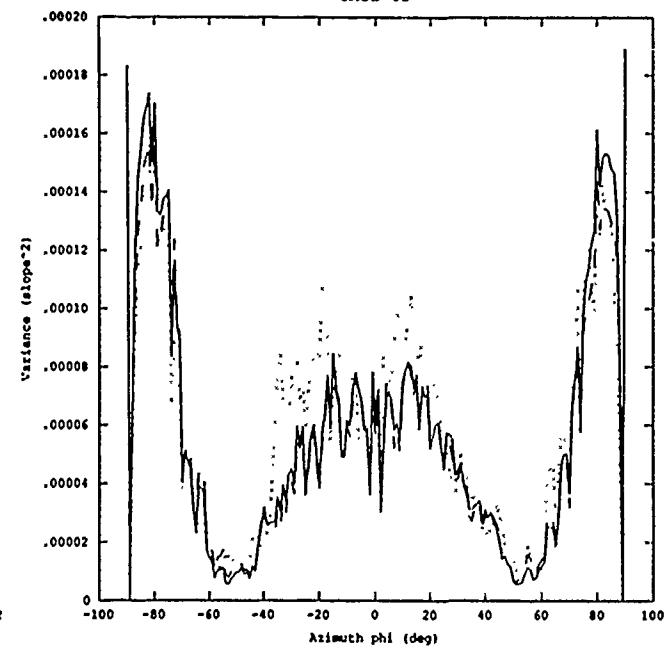
CASE 17



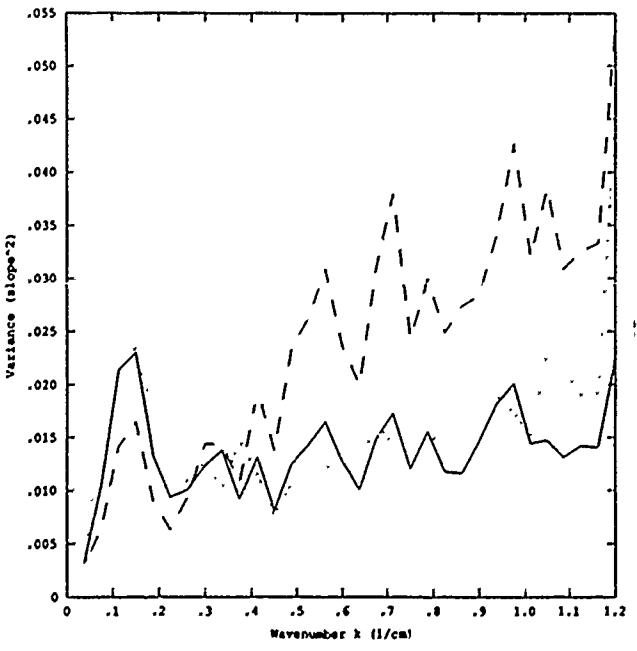
CASE 65



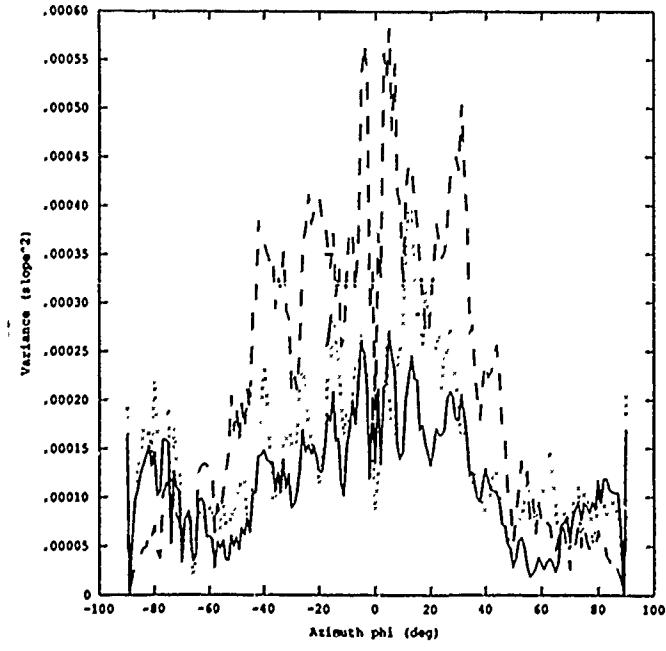
CASE 65

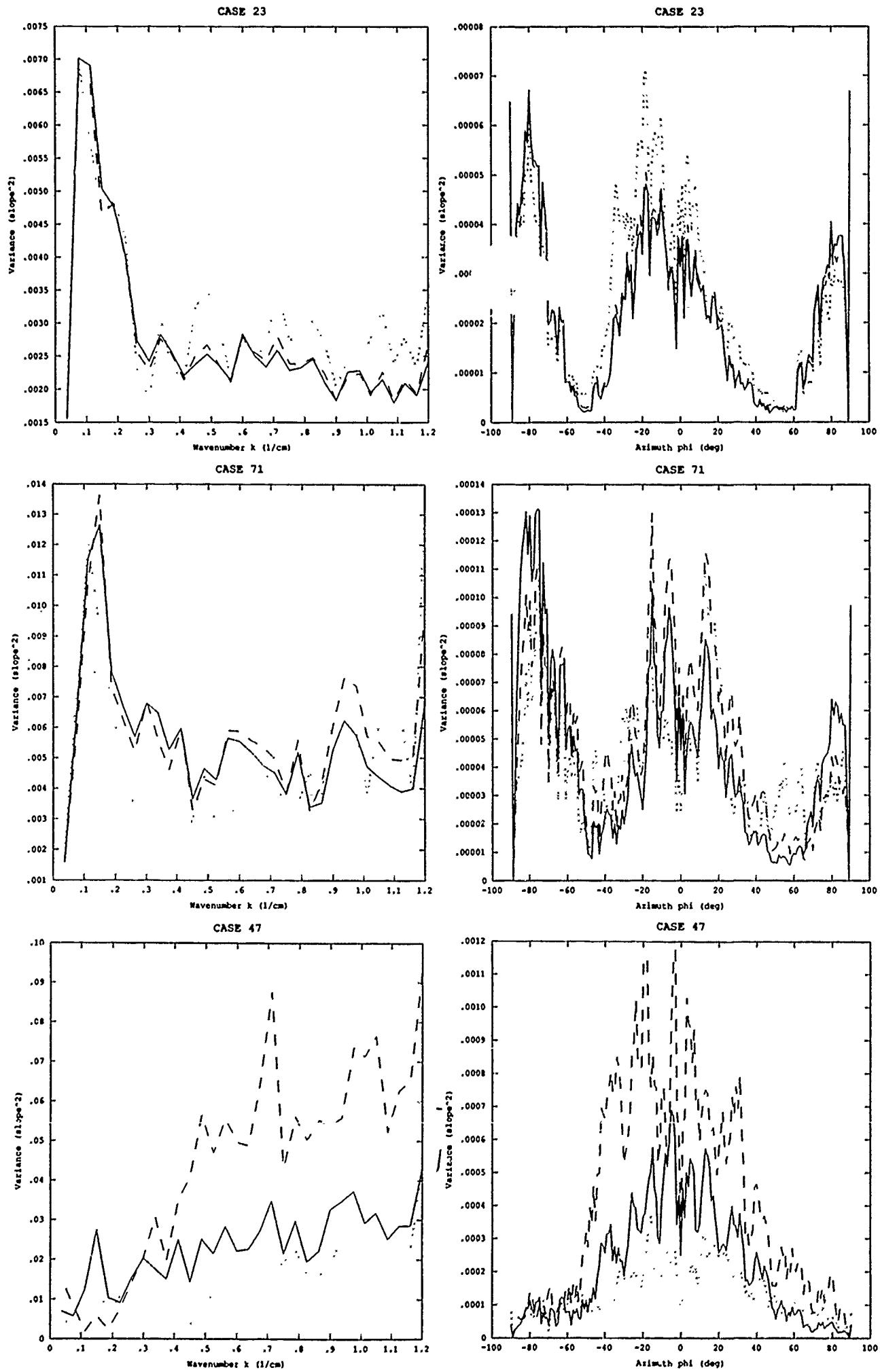


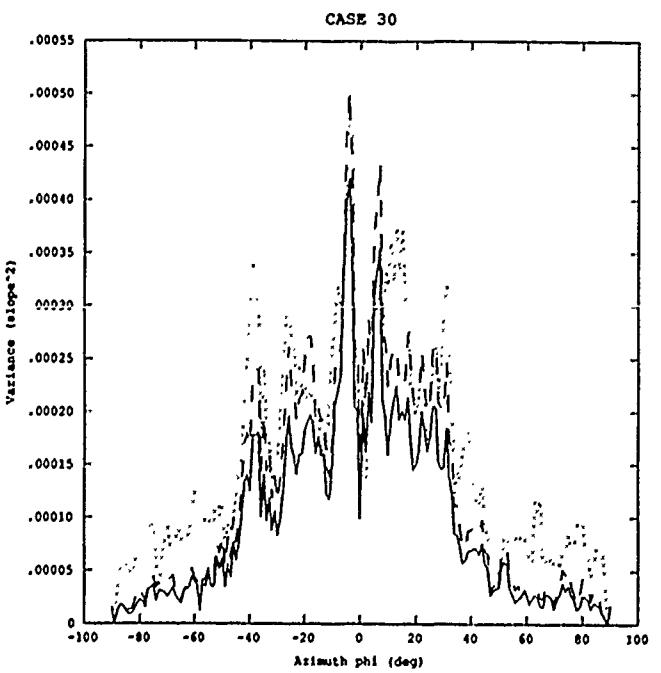
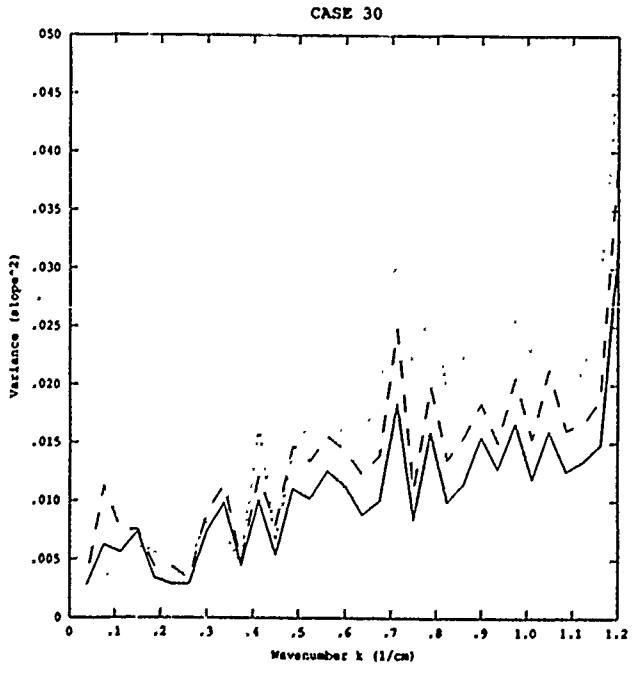
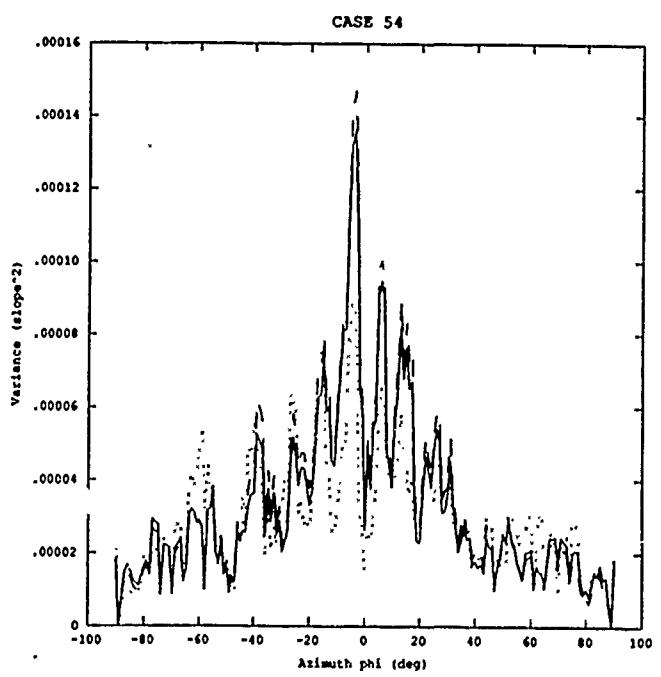
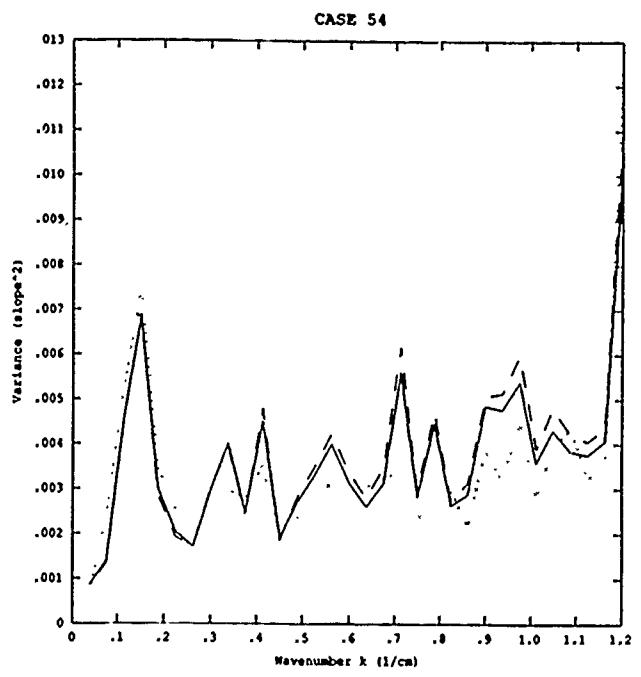
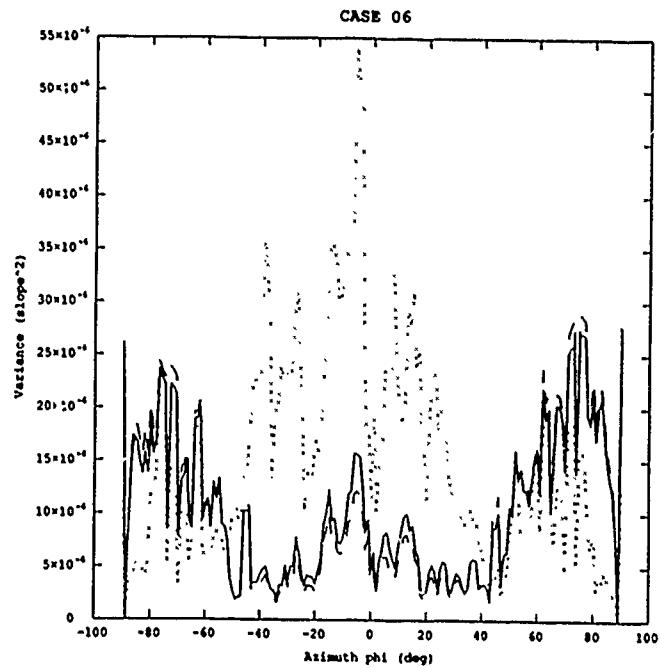
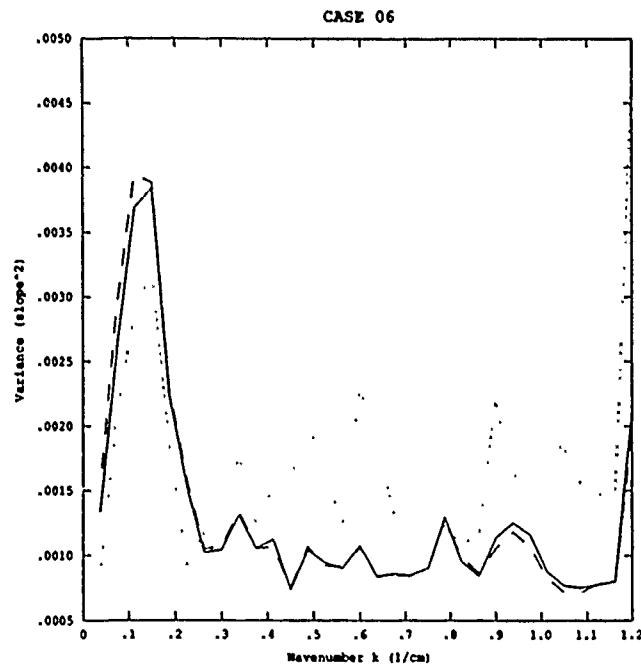
CASE 41

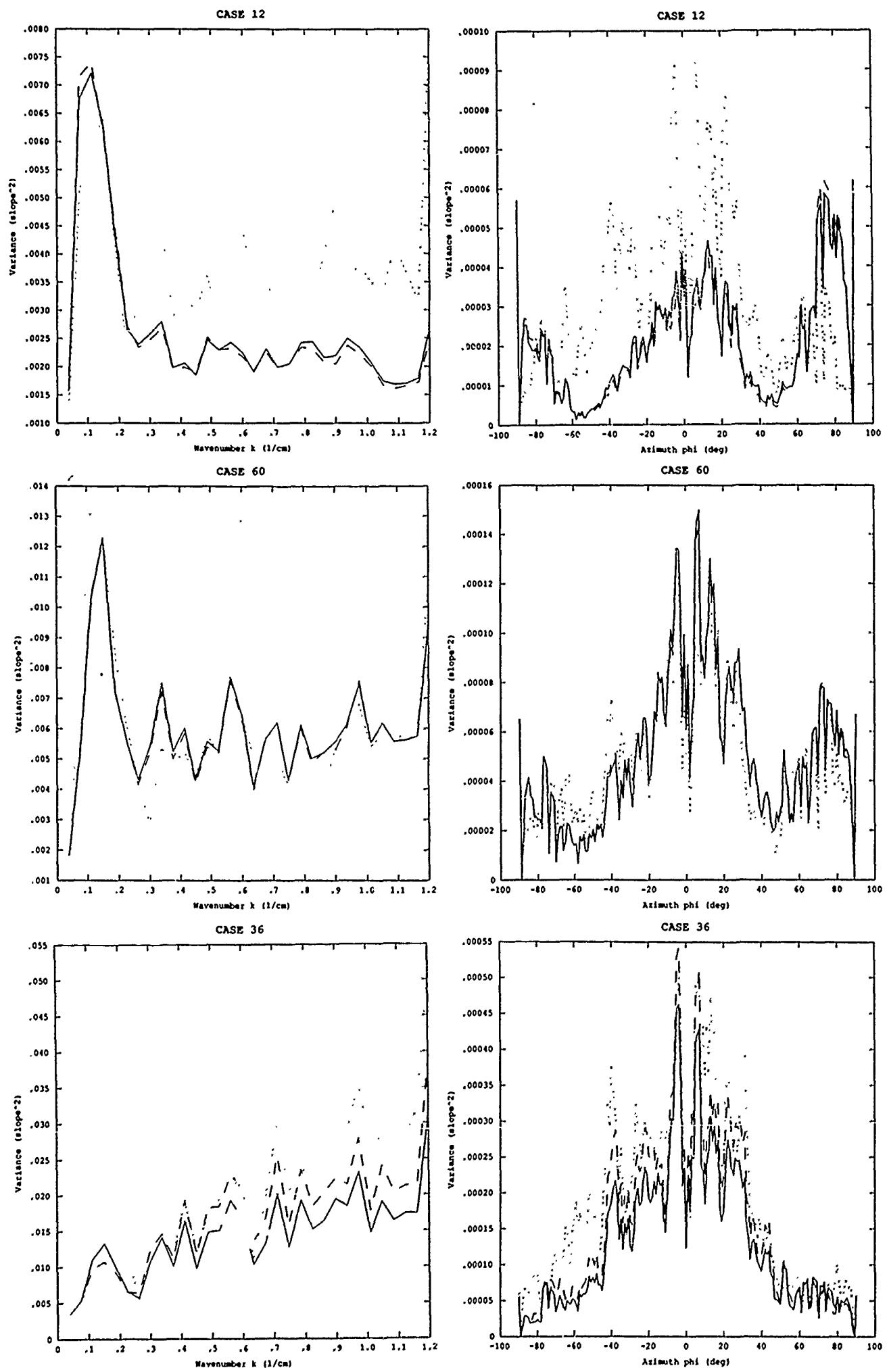


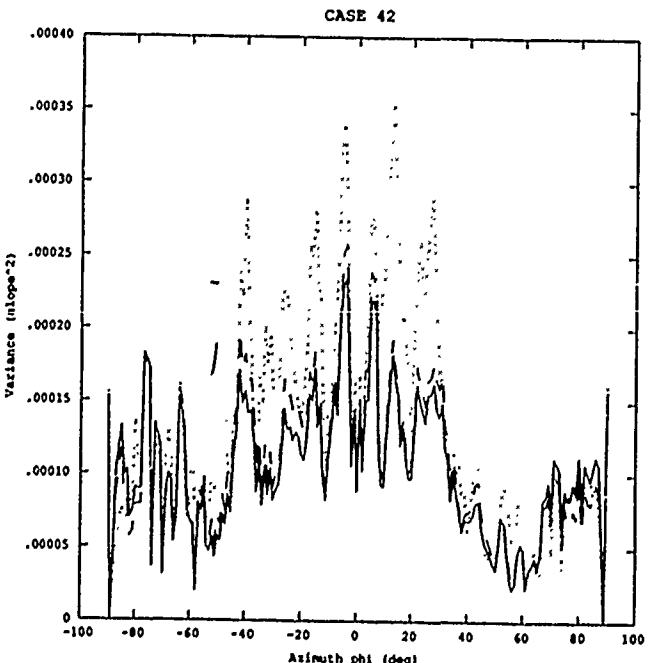
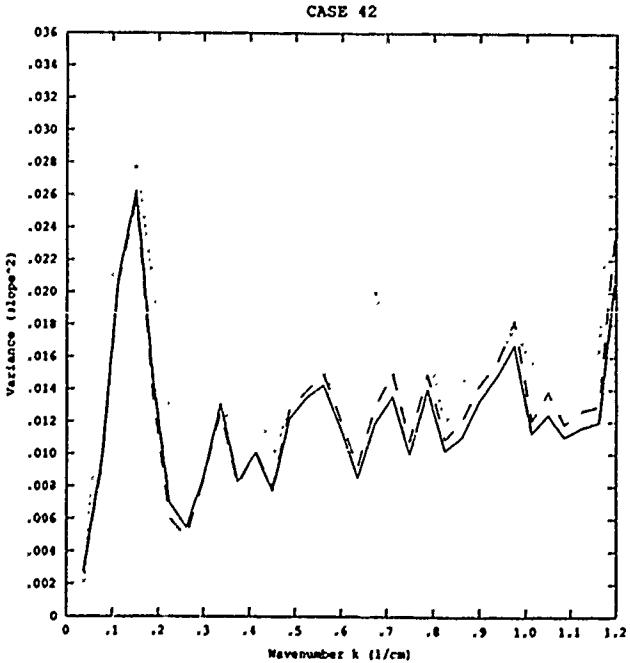
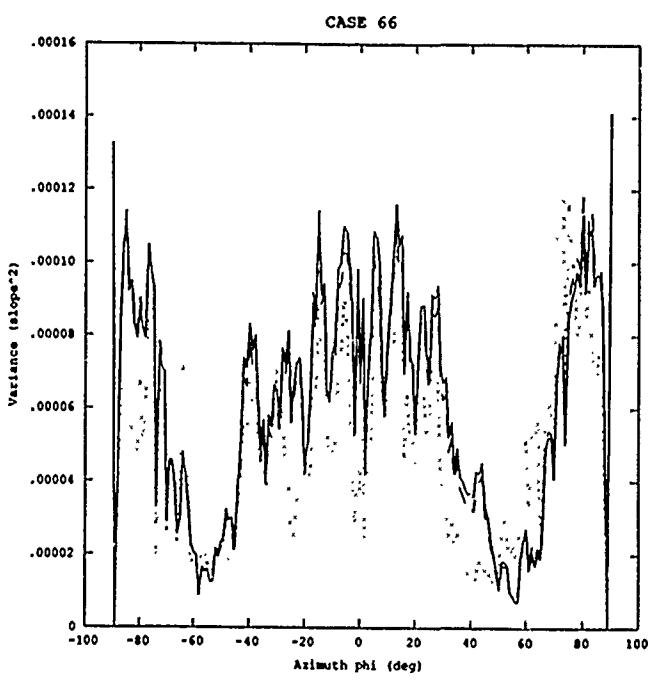
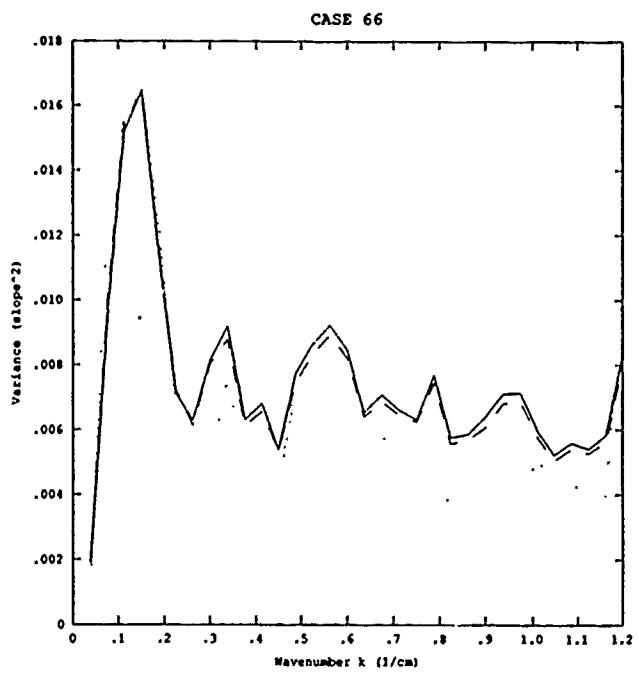
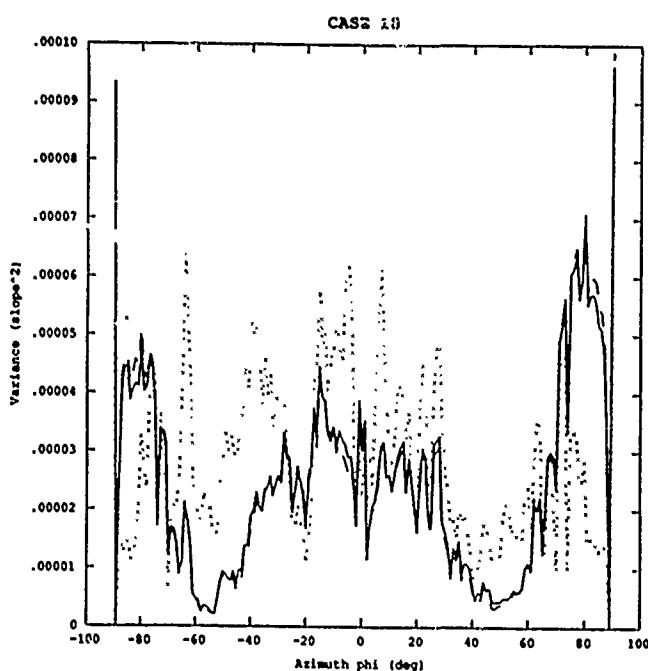
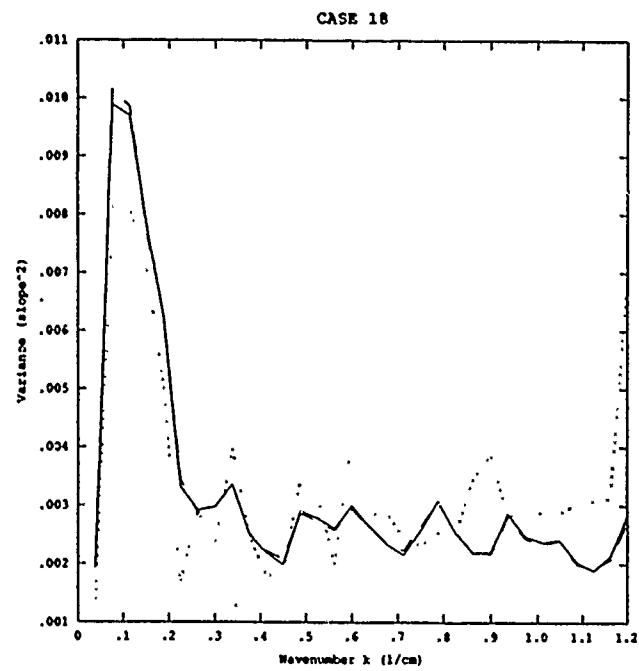
CASE 41

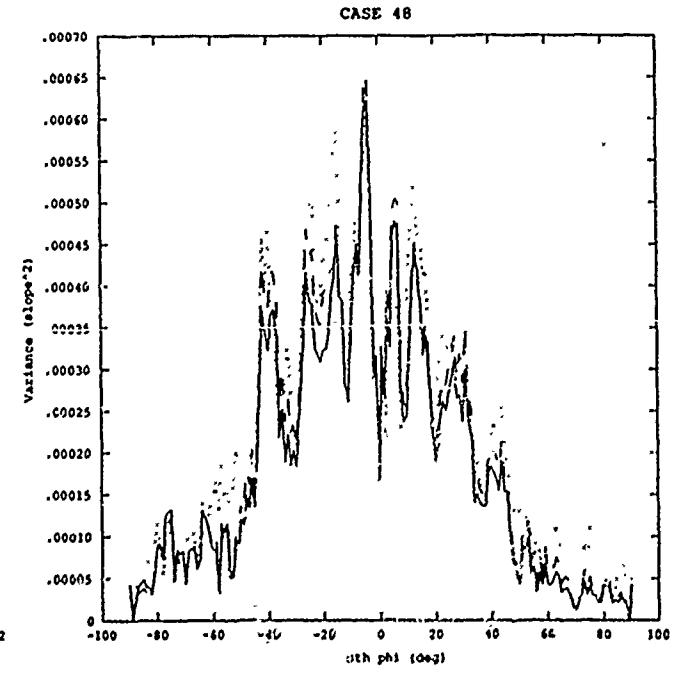
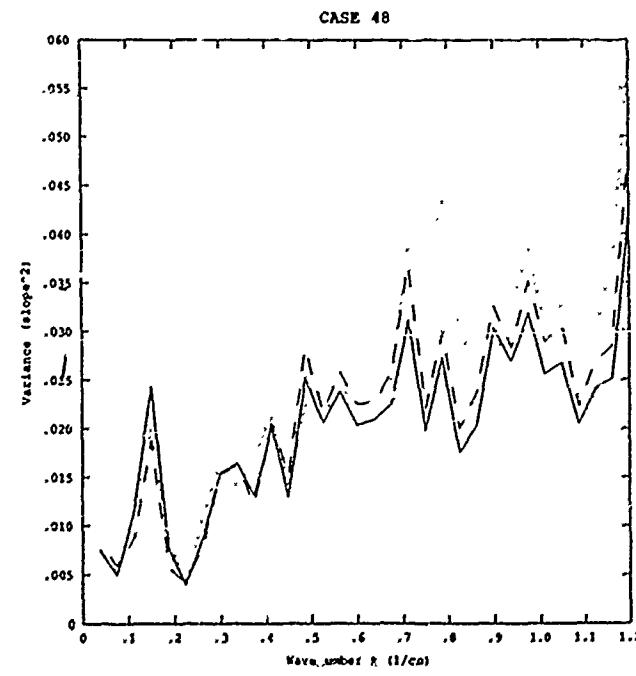
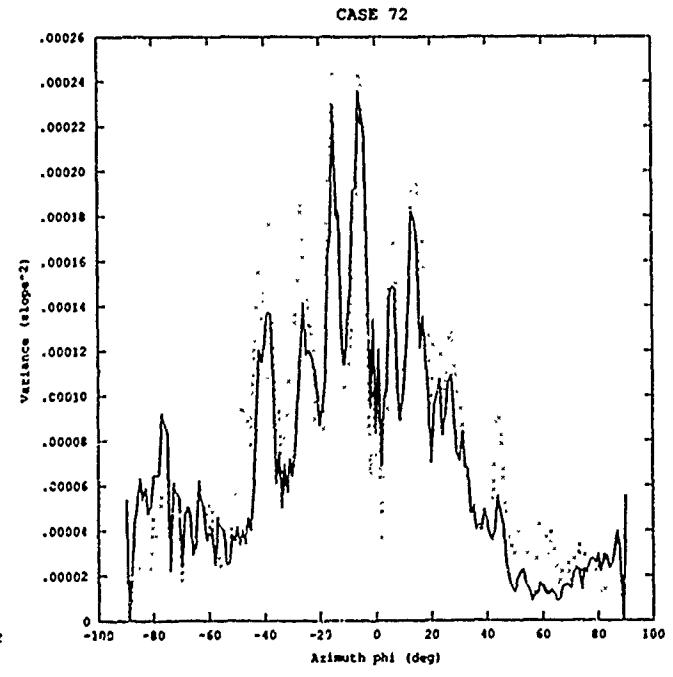
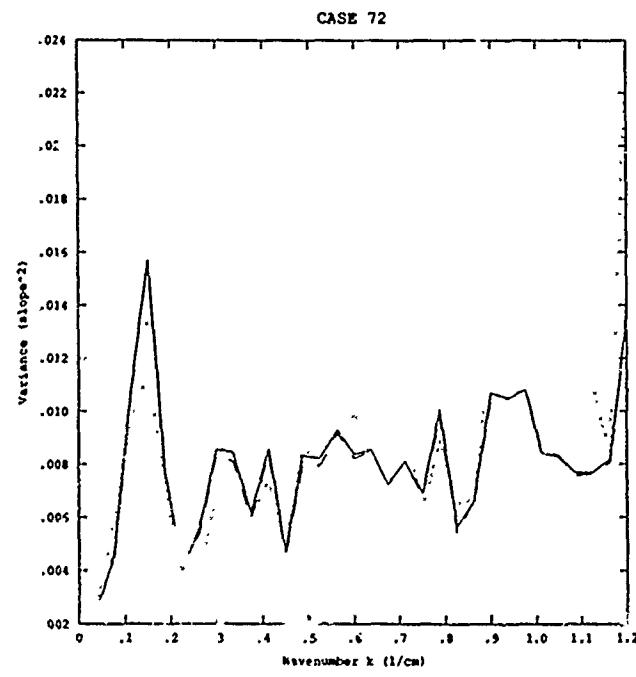
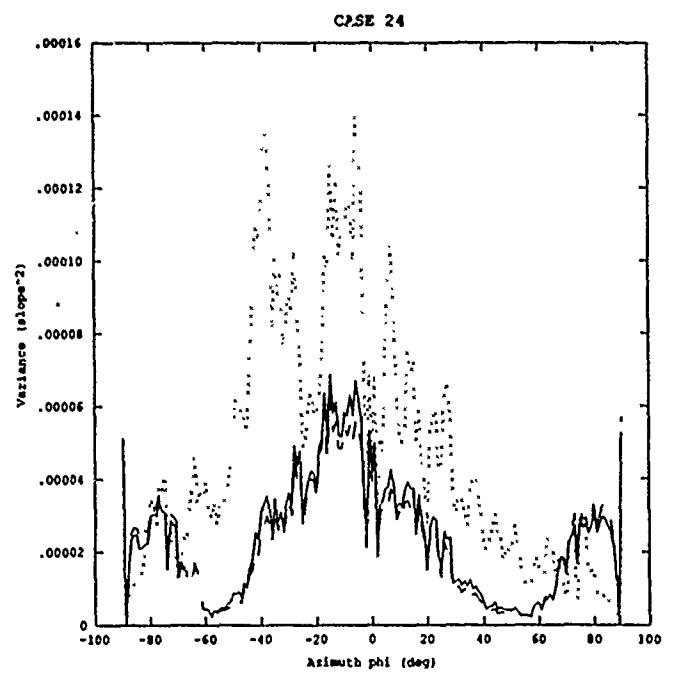
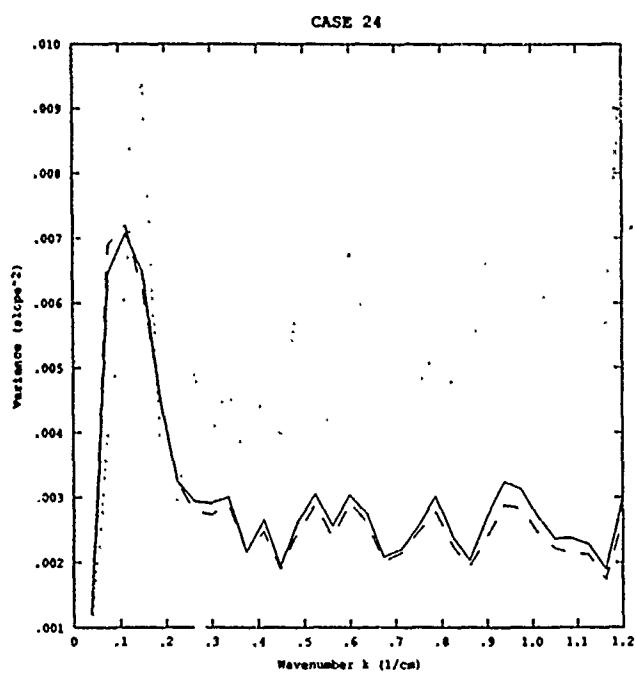












## APPENDIX II - D<sub>0</sub> Results

The one-dimensional graphic results for the 72 combinations of solar position, dominant wind azimuth, and wind friction velocity are presented in Appendix II.

The D<sub>1</sub> results for each test case are organized in the following format:

PARAMETERS - The format for the input file PARM.DAT is defined in Appendix III.

MAGNITUDE ARRAY SCALING FACTORS - For each of the five radiance conditions, the parameters are defined:

XX(0,0) = Mean Radiance  
High = Peak Radiance for k > 0  
Factor = Geometric Scaling Factor between Peak Radiance and Peak Slope Magnitude

INTEGRATED VARIANCE FROM POWER SPECTRA - The variance is calculated from S<sub>0</sub>, the input slope spectrum, and for each of the five normalized radiance spectra.

A,C,S - Along-Wind, Cross-Wind, & Summed Variance  
X,Y,S - Cross-Field, Along-Field, & Summed Variance

INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA - The variance is calculated from the error (difference) spectra between S<sub>0</sub>, the input slope spectrum, and each of the five normalized radiance spectra.

A,C,S - Along-Wind, Cross-Wind, & Summed Variance  
X,Y,S - Cross-Field, Along-Field, & Summed Variance

## CASE#1.DAT

## CASE#2.DAT

PARAMETERS	1.200000	3.750000E-02	64.000000	56789.00
12.000000	<b>0.000000E+00</b>	<b>0.000000E+00</b>	<b>64.000000</b>	<b>64.000000</b>
53.200000	<b>0.000000E+00</b>	<b>0.000000E+00</b>	<b>104.000000</b>	<b>9.885000</b>
1.337000	<b>1.020000E-02</b>	<b>99999.00</b>	<b>99999.00</b>	<b>99999.00</b>

## MAGNITUDE ARRAY SCALING FACTORS

H1( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>0.979596</b>	2.3192466E-02	15.01483
L1( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>0.9915391</b>	2.2538855E-02	16.27316
H0( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>1.163583</b>	<b>2.0704437E-02</b>	17.71529
L0( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>1.362177</b>	1.7567595E-02	28.87658
H3( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>0.8971584</b>	2.7659817E-02	13.26058

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S# A,C,S	<b>1.0605898E-02</b>	<b>6.3896347E-03</b>	<b>1.6995530E-02</b>
VAR H1 A,C,S	<b>1.0746664E-02</b>	<b>2.9954782E-03</b>	<b>1.3742088E-02</b>
VAR L1 A,C,S	<b>1.0789552E-02</b>	<b>3.0410733E-03</b>	<b>1.3830633E-02</b>
VAR H# A,C,S	<b>1.0751334E-02</b>	<b>3.0255551E-03</b>	<b>1.3776912E-02</b>
VAR L# A,C,S	<b>1.0901214E-02</b>	<b>3.1742391E-03</b>	<b>1.4075472E-02</b>
VAR H3 A,C,S	<b>1.2805673E-02</b>	<b>4.0093493E-03</b>	<b>1.6815042E-02</b>
VAR S# X,Y,S	<b>1.0605898E-02</b>	<b>6.3896337E-03</b>	<b>1.6995530E-02</b>
VAR H1 X,Y,S	<b>1.0746664E-02</b>	<b>2.9954782E-03</b>	<b>1.3742088E-02</b>
VAR L1 X,Y,S	<b>1.0789552E-02</b>	<b>3.0410733E-03</b>	<b>1.3830633E-02</b>
VAR H# X,Y,S	<b>1.0751334E-02</b>	<b>3.0255551E-03</b>	<b>1.3776912E-02</b>
VAR L# X,Y,S	<b>1.0901214E-02</b>	<b>3.1742391E-03</b>	<b>1.4075472E-02</b>
VAR H3 X,Y,S	<b>1.2805673E-02</b>	<b>4.0093493E-03</b>	<b>1.6815042E-02</b>

## MAGNITUDE ARRAY SCALING FACTORS

H1( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>0.6247715</b>	<b>1.7953391E-02</b>	<b>26.45263</b>
L1( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>0.6627284</b>	<b>2.0325290E-02</b>	<b>18.04575</b>
H0( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>0.7304618</b>	<b>1.6718665E-02</b>	<b>21.93866</b>
L0( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>0.8741091</b>	<b>1.7895389E-02</b>	<b>26.49618</b>
H3( $\theta$ , $\phi$ ), HIGH, FACTOR	<b>0.5820262</b>	<b>2.0975588E-02</b>	<b>17.49152</b>

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	<b>1.2719949E-03</b>	<b>1.7413879E-03</b>	<b>1.4928144E-03</b>
DIF L1 A,C,S	<b>1.2355722E-03</b>	<b>1.7443985E-03</b>	<b>1.4928144E-03</b>
DIF H# A,C,S	<b>1.2448757E-03</b>	<b>1.7367585E-03</b>	<b>1.4928144E-03</b>
DIF L# A,C,S	<b>1.1499249E-03</b>	<b>1.7373970E-03</b>	<b>1.4928144E-03</b>
DIF H3 A,C,S	<b>9.0592814E-04</b>	<b>1.7122239E-03</b>	<b>1.4928144E-03</b>
DIF H1 X,Y,S	<b>4.6939135E-04</b>	<b>1.2719949E-03</b>	<b>1.7413879E-03</b>
DIF L1 X,Y,S	<b>5.0949947E-04</b>	<b>1.2355722E-03</b>	<b>1.7443985E-03</b>
DIF H# X,Y,S	<b>4.9187889E-04</b>	<b>1.2448757E-03</b>	<b>1.7367585E-03</b>
DIF L# X,Y,S	<b>5.8747106E-04</b>	<b>1.1499249E-03</b>	<b>1.7373970E-03</b>
DIF H3 X,Y,S	<b>3.6629782E-04</b>	<b>9.0592814E-04</b>	<b>1.7122239E-03</b>

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S# A,C,S	<b>1.0605898E-02</b>	<b>6.3896337E-03</b>	<b>1.6995530E-02</b>
VAR H1 A,C,S	<b>9.7574415E-03</b>	<b>5.5585625E-03</b>	<b>1.5316016E-02</b>
VAR L1 A,C,S	<b>9.9829556E-03</b>	<b>5.0895498E-03</b>	<b>1.5072467E-02</b>
VAR H# A,C,S	<b>9.6167721E-03</b>	<b>5.9525055E-03</b>	<b>1.5569224E-02</b>
VAR L# A,C,S	<b>9.7021358E-03</b>	<b>5.6932503E-03</b>	<b>1.5395389E-02</b>
VAR H3 A,C,S	<b>1.0342992E-02</b>	<b>5.0727455E-03</b>	<b>1.5415745E-02</b>
VAR S# X,Y,S	<b>1.0605898E-02</b>	<b>6.3896347E-03</b>	<b>1.6995530E-02</b>
VAR H1 X,Y,S	<b>9.7574415E-03</b>	<b>5.5585625E-03</b>	<b>1.5316016E-02</b>
VAR L1 X,Y,S	<b>9.9829556E-03</b>	<b>5.0895498E-03</b>	<b>1.5072467E-02</b>
VAR H# X,Y,S	<b>9.6167721E-03</b>	<b>5.9525055E-03</b>	<b>1.5569224E-02</b>
VAR L# X,Y,S	<b>9.7021358E-03</b>	<b>5.6932503E-03</b>	<b>1.5395389E-02</b>
VAR H3 X,Y,S	<b>1.0342992E-02</b>	<b>5.0727455E-03</b>	<b>1.5415745E-02</b>

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	<b>1.3336245E-03</b>	<b>1.1877267E-03</b>	<b>2.5213528E-03</b>
DIF L1 X,Y,S	<b>1.1597133E-03</b>	<b>1.1613494E-03</b>	<b>2.3210510E-03</b>
DIF H# X,Y,S	<b>1.4386721E-03</b>	<b>1.1840228E-03</b>	<b>2.6226700E-03</b>
DIF L# X,Y,S	<b>1.3273243E-03</b>	<b>1.1429832E-03</b>	<b>2.4763971E-03</b>
DIF H3 X,Y,S	<b>1.0577217E-03</b>	<b>1.0952185E-03</b>	<b>2.1529442E-03</b>
DIF H1 X,Y,S	<b>1.3336245E-03</b>	<b>1.1877267E-03</b>	<b>2.5213528E-03</b>
DIF L1 X,Y,S	<b>1.1597133E-03</b>	<b>1.1613494E-03</b>	<b>2.3210510E-03</b>
DIF H# X,Y,S	<b>1.4386721E-03</b>	<b>1.1840228E-03</b>	<b>2.6226700E-03</b>
DIF L# X,Y,S	<b>1.3273243E-03</b>	<b>1.1429832E-03</b>	<b>2.4763971E-03</b>
DIF H3 X,Y,S	<b>1.0577217E-03</b>	<b>1.0952185E-03</b>	<b>2.1529442E-03</b>

## CASE#3.DAT

## CASE#4.DAT

PARAMETERS	1.200000E+00	3.750000E-02	64.000000	56789.66
L1(0,0),HIGH,FACTOR	12.000000E+00	8.000000E+00	64.000000	64.000000
H0(0,0),HIGH,FACTOR	53.200000	53.200000	9.885000	9.885000
L0(0,0),HIGH,FACTOR	1.337699	1.020000E-09	99999.66	99999.66

## MAGNITUDE ARRAY SCALING FACTORS

H1(0,0),HIGH,FACTOR	0.2317155	6.7422264E-03	54.40118
L1(0,0),HIGH,FACTOR	0.2561573	8.5639572E-03	42.82893
H0(0,0),HIGH,FACTOR	0.3374459	5.3447563E-03	68.62520
L0(0,0),HIGH,FACTOR	0.4675389	5.7723238E-03	63.54262
H3(0,0),HIGH,FACTOR	0.2079985	7.9102358E-03	46.36842

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	1.0605898E-02	6.3896347E-03	1.6995503E-02
VAR H1 A,C,S	1.0727595E-02	3.2693485E-03	1.3976925E-02
VAR L1 A,C,S	1.0926355E-02	3.2327457E-03	1.4159084E-02
VAR H2 A,C,S	1.0735123E-02	3.3982147E-03	1.4133333E-02
VAR L2 A,C,S	1.1239463E-02	3.6712477E-03	1.4910653E-02
VAR H3 A,C,S	1.2236980E-02	3.8402595E-03	1.6077235E-02

VAR S0 A,C,S	1.0605898E-02	6.3896347E-03	1.6995503E-02
VAR H1 A,C,S	1.0727595E-02	3.2693485E-03	1.3976925E-02
VAR L1 A,C,S	1.0926355E-02	3.2327457E-03	1.4159084E-02
VAR H2 A,C,S	1.0735123E-02	3.3982147E-03	1.4133333E-02
VAR L2 A,C,S	1.1239463E-02	3.6712477E-03	1.4910653E-02
VAR H3 A,C,S	1.2236980E-02	3.8402595E-03	1.6077235E-02

## INTEGRATED SQUARED DEVIATION FROM DIF J PECTRA

DIF H1 A,C,S	4.5836368E-04	1.1700063E-03	1.6283707E-03
DIF L1 A,C,S	5.0758856E-04	1.1616988E-03	1.66292928E-03
DIF H0 A,C,S	5.10066874E-04	1.0776187E-03	1.5876786E-03
DIF L0 A,C,S	6.5042389E-04	9.7214105E-04	1.6225645E-03
DIF H3 A,C,S	6.6028541E-04	9.5055831E-04	1.6108441E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	1.200000E-02	3.750000E-02	64.000000
DIF L1 X,Y,S	1.200000E+00	0.00000000E+00	64.000000
DIF H0 X,Y,S	53.200000	135.0000	53.200000
DIF L0 X,Y,S	1.337699	1.020000E-09	99999.66
DIF H3 X,Y,S	1.337699	1.020000E-09	99999.66

## INTEGRATED SQUARED DEVIATION FROM DIF J PECTRA

DIF H1 X,Y,S	4.5836368E-04	1.1700063E-03	1.6283707E-03
DIF L1 X,Y,S	5.0758856E-04	1.1616988E-03	1.66292928E-03
DIF H0 X,Y,S	5.10066874E-04	1.0776187E-03	1.5876786E-03
DIF L0 X,Y,S	6.5042389E-04	9.7214105E-04	1.6225645E-03
DIF H3 X,Y,S	6.6028541E-04	9.5055831E-04	1.6108441E-03

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	1.0605898E-02	6.3896347E-03	1.6995503E-02
VAR H1 A,C,S	1.0466638E-02	3.3429695E-03	1.4389523E-02
VAR L1 A,C,S	1.0444998E-02	3.2545668E-03	1.4299524E-02
VAR H2 A,C,S	1.0922271E-02	3.4321698E-03	1.4354471E-02
VAR L2 A,C,S	1.0991573E-02	3.4492635E-03	1.4449851E-02
VAR H3 A,C,S	1.23598813E-02	3.9367173E-03	1.6296523E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	5.8900731E-04	1.3127775E-03	1.9018683E-03
DIF L1 X,Y,S	5.7132472E-04	1.3075555E-03	1.8788876E-03
DIF H0 X,Y,S	6.5141202E-04	1.3073839E-03	1.9587951E-03
DIF L0 X,Y,S	6.9430331E-04	1.2740921E-03	1.9683971E-03
DIF H3 X,Y,S	7.7554921E-04	1.0569545E-03	1.8325031E-03

## CASE45.DAT

CASE06.DAT

## PARAMETERS

1.2E0000	3.7500001E-02	64.00000	56789.00
12.00000	9.0000000E+00	0.0000000E+00	64.00000
53.20000	180.00000	53.20000	9.0000000
1.337000	1.02899999E-09	99999.00	99999.00
			99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	0.4624957	1.3453400E-02	27.26337
L1(θ,θ), HIGH, FACTOR	0.473737886	1.4737497E-02	24.88788
H0(θ,θ), HIGH, FACTOR	0.5681861	1.2054356E-02	30.42769
L0(θ,θ), HIGH, FACTOR	0.6851693	1.1939812E-02	30.71951
H3(θ,θ), HIGH, FACTOR	0.4325745	1.5354494E-02	23.88788

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S# A,C,S	1.6605878E-02	6.3896347E-03	1.6995503E-02
VAR H1 A,C,S	1.9751618E-02	2.9393421E-03	1.3691191E-02
VAR L1 A,C,S	1.0789744E-02	2.9548895E-03	1.3744595E-02
VAR H0 A,C,S	1.0723243E-02	2.9455797E-03	1.3668792E-02
VAR L0 A,C,S	1.07566894E-02	2.9788330E-03	1.3735641E-02
VAR H3 A,C,S	1.2178957E-02	3.6580893E-03	1.5637045E-02
VAR S# X,Y,S	1.6605898E-02	6.3896347E-03	1.6995503E-02
VAR H1 X,Y,S	1.0751618E-02	2.9393421E-03	1.3691191E-02
VAR L1 X,Y,S	1.0789744E-02	2.9548895E-03	1.3744595E-02
VAR H0 X,Y,S	1.0723243E-02	2.9455797E-03	1.3668792E-02
VAR L0 X,Y,S	1.07566894E-02	2.9788330E-03	1.3735641E-02
VAR H3 X,Y,S	1.2178957E-02	3.6580893E-03	1.5637045E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S# A,C,S	1.0605398E-02	6.3896337E-03	1.6995503E-02
VAR H1 A,C,S	1.1827840E-02	3.5197808E-03	1.5347634E-02
VAR L1 A,C,S	1.1429348E-02	3.39404436E-03	1.4733412E-02
VAR H0 A,C,S	1.1892175E-02	3.5655941E-03	1.5457775E-02
VAR L0 A,C,S	1.1492500E-02	3.3585266E-03	1.4850751E-02
VAR H3 A,C,S	1.5504099E-02	6.0993349E-03	2.1603326E-02
VAR S# X,Y,S	1.0605398E-02	6.3896337E-03	1.6995503E-02
VAR H1 X,Y,S	1.1827840E-02	3.5197808E-03	1.5347634E-02
VAR L1 X,Y,S	1.1429348E-02	3.39404436E-03	1.4733412E-02
VAR H0 X,Y,S	1.1892175E-02	3.5655941E-03	1.5457775E-02
VAR L0 X,Y,S	1.1492500E-02	3.3585266E-03	1.4850751E-02
VAR H3 X,Y,S	1.5504099E-02	6.0993349E-03	2.1603326E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	4.0562998E-04	1.3370063E-03	1.7426370E-03
DIF L1 A,C,S	4.2029686E-04	1.3280771E-03	1.7483742E-03
DIF H0 A,C,S	4.1612997E-04	1.3227591E-03	1.7388887E-03
DIF L0 A,C,S	4.4931459E-04	1.2933519E-03	1.7413681E-03
DIF H3 A,C,S	6.4498034E-04	1.0170180E-03	1.6610981E-03
DIF H1 X,Y,S	4.0562998E-04	1.3370063E-03	1.7426370E-03
DIF L1 X,Y,S	4.2029686E-04	1.3280771E-03	1.7483742E-03
DIF H0 X,Y,S	4.1612997E-04	1.3227591E-03	1.7388887E-03
DIF L0 X,Y,S	4.4931459E-04	1.2933519E-03	1.7413681E-03
DIF H3 X,Y,S	6.4498034E-04	1.0170180E-03	1.6610981E-03

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S# A,C,S	1.0605398E-02	6.3896337E-03	1.6995503E-02
VAR H1 A,C,S	1.1827840E-02	3.5197808E-03	1.5347634E-02
VAR L1 A,C,S	1.1429348E-02	3.39404436E-03	1.4733412E-02
VAR H0 A,C,S	1.1892175E-02	3.5655941E-03	1.5457775E-02
VAR L0 A,C,S	1.1492500E-02	3.3585266E-03	1.4850751E-02
VAR H3 A,C,S	1.5504099E-02	6.0993349E-03	2.1603326E-02
VAR S# X,Y,S	1.0605398E-02	6.3896337E-03	1.6995503E-02
VAR H1 X,Y,S	1.1827840E-02	3.5197808E-03	1.5347634E-02
VAR L1 X,Y,S	1.1429348E-02	3.39404436E-03	1.4733412E-02
VAR H0 X,Y,S	1.1892175E-02	3.5655941E-03	1.5457775E-02
VAR L0 X,Y,S	1.1492500E-02	3.3585266E-03	1.4850751E-02
VAR H3 X,Y,S	1.5504099E-02	6.0993349E-03	2.1603326E-02

## CASE#7.DAT

## CASE#8.DAT

PARAMETERS	VALUES
H1(θ,θ), HIGH, FACTOR	0.963361
L1(θ,θ), HIGH, FACTOR	0.9759669
Hθ(θ,θ), HIGH, FACTOR	1.151857
Lθ(θ,θ), HIGH, FACTOR	1.346936
H3(θ,θ), HIGH, FACTOR	0.8803946
	1.0210000E-09
	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ)	2.1553256E-02	20.18031
L1(θ,θ)	2.11497065E-02	20.57411
Hθ(θ,θ)	1.9494136E-02	22.41539
Lθ(θ,θ)	1.6866012E-02	25.81924
H3(θ,θ)	2.3213871E-02	18.73676

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	1.0649075E-02	6.3449983E-03	1.6994014E-02
VAR H1 A,C,S	1.3324276E-02	8.3941165E-03	2.1718401E-02
VAR L1 A,C,S	1.3449105E-02	8.5145570E-03	2.196345E-02
VAR Hθ A,C,S	1.34508821E-02	8.5107880E-03	2.1961607E-02
VAR Lθ A,C,S	1.3808226E-02	8.8802287E-03	2.2728880E-02
VAR H3 A,C,S	1.41385735E-02	9.1494014E-03	2.3287963E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ X,Y,S	8.4970202E-03	8.4970109E-03	1.6994014E-02
VAR H1 X,Y,S	1.5874185E-02	5.9442070E-03	2.1718401E-02
VAR L1 X,Y,S	1.6688296E-02	5.9353625E-03	2.196345E-02
VAR Hθ X,Y,S	1.6626707E-02	5.9349309E-03	2.1961607E-02
VAR Lθ X,Y,S	1.6588242E-02	6.228271E-03	2.2728880E-02
VAR H3 X,Y,S	1.6673123E-02	6.6148997E-03	2.3287963E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	1.9298031E-03	1.6361463E-03	3.5659471E-03
DIF L1 A,C,S	1.9493713E-03	1.6535096E-03	3.6028787E-03
DIF Hθ A,C,S	1.9492735E-03	1.6531112E-03	3.6013813E-03
DIF Lθ A,C,S	2.0178822E-03	1.7137438E-03	3.7316259E-03
DIF H3 A,C,S	2.0287863E-03	1.6956879E-03	3.7244821E-03
DIF H1 X,Y,S	1.8270897E-03	1.7388582E-03	3.5659471E-03
DIF L1 X,Y,S	1.9004126E-03	1.7624733E-03	3.6028787E-03
DIF Hθ X,Y,S	1.8916044E-03	1.7897787E-03	3.6013813E-03
DIF Lθ X,Y,S	2.1048384E-03	1.6267921E-03	3.7316259E-03
DIF H3 X,Y,S	2.2399733E-03	1.4935071E-03	3.7244821E-03

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	1.0649075E-02	6.3449983E-03	1.6994014E-02
VAR H1 A,C,S	1.6529538E-02	4.8296824E-03	2.1559226E-02
VAR L1 A,C,S	1.7114335E-02	5.1686037E-03	2.2282979E-02
VAR Hθ A,C,S	1.6552282E-02	4.8371125E-03	2.1389436E-02
VAR Lθ A,C,S	1.7181354E-02	5.13899889E-03	2.2320338E-02
VAR H3 A,C,S	1.8383533E-02	5.6448882E-03	2.4628497E-02
VAR Sθ X,Y,S	8.4970202E-03	8.4970109E-03	1.6994014E-02
VAR H1 X,Y,S	1.0105016E-02	1.0543401E-02	2.1359226E-02
VAR L1 X,Y,S	1.2076455E-02	1.6206512E-02	2.2282979E-02
VAR Hθ X,Y,S	1.0657449E-02	1.0731947E-02	2.1389436E-02
VAR Lθ X,Y,S	1.1424644E-02	1.0895728E-02	2.2320338E-02
VAR H3 X,Y,S	1.2826443E-02	1.1201976E-02	2.4628497E-02
DIF H1 X,Y,S	9.7821641E-04	1.0335323E-03	2.0136009E-03
DIF L1 X,Y,S	1.0751315E-03	1.0785615E-03	2.1536872E-03
DIF Hθ X,Y,S	9.7513361E-04	1.0316728E-03	2.0068864E-03
DIF Lθ X,Y,S	1.0433987E-03	1.0785997E-03	2.1219798E-03
DIF H3 X,Y,S	1.2638383E-03	1.2224659E-03	2.4862455E-03

PARAMETERS		PARAMETERS	
1.20000E-02	3.7500001E-02	64.000000	56789.00
12.00000	0.0000000E+00	45.000000	64.000000
53.20000	99.000000	53.200000	9.8850000
1.3337000	1.0200000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1( $\theta, \theta$ ), HIGH, FACTOR 0.2319100  
 L1( $\theta, \theta$ ), HIGH, FACTOR 0.252389  
 H0( $\theta, \theta$ ), H, GH, FACTOR 0.3376779  
 L0( $\theta, \theta$ ), HIGH, FACTOR 0.4610146  
 H3( $\theta, \theta$ ), HIGH, FACTOR 0.2052276  
 VAR S0 A,C,S 1.0649075E-02 6.3449933E-03 1.6994014E-02  
 VAR H1 A,C,S 1.4888551E-02 7.7039288E-03 2.2592455E-02  
 VAR L1 A,C,S 1.44982882E-02 7.7737719E-03 2.2272666E-02  
 VAR H0 A,C,S 1.5721593E-02 7.9157355E-03 2.3637339E-02  
 VAR L0 A,C,S 1.6927495E-02 8.3681438E-03 2.4395257E-02  
 VAR H3 A,C,S 1.50988419E-02 8.1891386E-03 2.3287376E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 X,Y,S 3.4970202E-03 8.4970109E-03 1.6994014E-02  
 VAR H1 X,Y,S 1.5786350E-02 6.8061128E-03 2.2592455E-02  
 VAR L1 X,Y,S 1.5682261E-02 6.5897778E-03 2.2272666E-02  
 VAR H0 X,Y,S 1.6190013E-02 7.4472865E-03 2.3637339E-02  
 VAR L0 X,Y,S 1.6636493E-02 7.7588325E-03 2.4395257E-02  
 VAR H3 X,Y,S 1.6161703E-02 7.1258489E-03 2.3287376E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 1.7148135E-03 1.3221649E-03 3.0369780E-03  
 DIF L1 A,C,S 1.7441006E-03 1.3935184E-03 3.1376174E-03  
 DIF H0 A,C,S 1.7761995E-03 1.2635458E-03 3.0337465E-03  
 DIF L0 A,C,S 1.9153331E-03 1.3619228E-03 3.2773588E-03  
 DIF H3 A,C,S 1.8527806E-03 1.3862765E-03 3.2396505E-03  
 DIF H1 X,Y,S 1.6691079E-03 1.3678706E-03 3.0369780E-03  
 DIF L1 X,Y,S 1.7055691E-03 1.4288492E-03 3.1376174E-03  
 DIF H0 X,Y,S 1.7987142E-03 1.2359267E-03 3.0337465E-03  
 DIF L0 X,Y,S 2.0597964E-03 1.2175690E-03 3.2773588E-03  
 DIF H3 X,Y,S 1.9055881E-03 1.3288634E-03 3.2396505E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 2.3250512E-03 1.9767422E-03 1.6994014E-02  
 DIF L1 A,C,S 2.2959359E-03 1.9980911E-03 2.1792749E-02  
 DIF H0 A,C,S 2.3572170E-03 2.0493325E-03 2.2697223E-02  
 DIF L0 A,C,S 1.6436875E-02 1.697007E-02 1.1618906E-02  
 DIF H3 A,C,S 1.2265611E-02 1.2265611E-02 2.258692E-02  
 VAR S0 X,Y,S 8.4970202E-03 8.4970109E-03 1.6994014E-02  
 VAR H1 X,Y,S 1.5989481E-02 5.8033386E-03 2.1792749E-02  
 VAR L1 X,Y,S 1.6304408E-02 5.7926297E-03 2.2697223E-02  
 VAR H0 X,Y,S 1.5980539E-02 5.9968959E-03 2.1976655E-02  
 VAR L0 X,Y,S 1.6436875E-02 6.1560389E-03 2.258692E-02  
 VAR H3 X,Y,S 1.8946673E-02 6.8447292E-03 2.4911504E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

DIF H1 A,C,S 2.3250512E-03 1.9767422E-03 1.6994014E-02  
 DIF L1 A,C,S 2.2959359E-03 1.9980911E-03 2.1792749E-02  
 DIF H0 A,C,S 2.3572170E-03 2.0493325E-03 2.2697223E-02  
 DIF L0 A,C,S 1.6436875E-02 1.697007E-02 1.1618906E-02  
 DIF H3 A,C,S 1.2265611E-02 1.2265611E-02 2.258692E-02  
 VAR S0 X,Y,S 8.9936415E-03 8.9936415E-03 4.3617948E-03  
 VAR H1 X,Y,S 1.0155986E-02 1.0155986E-02 4.2855288E-03  
 VAR L1 X,Y,S 7.8516882E-03 7.8516882E-03 5.39599E-03  
 VAR H0 X,Y,S 7.8747338E-03 7.8747338E-03 5.5235339E-03  
 VAR L0 X,Y,S 9.0200000E-03 9.0200000E-03 4.4965528E-03  
 VAR H3 X,Y,S 9.99999.00 9.99999.00 4.8001758E-03

## CASE11.DAT

## CASE12.DAT

PARAMETERS		PARAMETERS	
1.2370000E-02	3.7500000E-02	1.2690000	3.7500000E-02
12.000000	0.0000000E+00	12.000000	0.0000000E+00
53.200000	100.00000	53.200000	9.8856000
1.3370000	1.0240000E-02	1.3370000	1.0200000E-02
			99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	0.4582272
L1(θ,θ), HIGH, FACTOR	0.4674655
H2(θ,θ), HIGH, FACTOR	0.56440151
L2(θ,θ), HIGH, FACTOR	0.6789812
H3(θ,θ), HIGH, FACTOR	0.4248831
	1.2372645E-02
	32.24118
	39.00066
	39.35696
	33.50769

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	1.0649975E-02
VAR H1 A,C,S	1.2865312E-02
VAR L1 A,C,S	1.2858900E-02
VAR H2 A,C,S	1.2942628E-02
VAR L2 A,C,S	1.3042567E-02
VAR H3 A,C,S	1.3312777E-02
	6.3449983E-03
	2.0991636E-03
	8.1248945E-03
	2.1621398E-03
	8.2899174E-03
	2.1861432E-03

VAR Sθ X,Y,S	8.4970202E-03
VAR H1 X,Y,S	1.5411285E-02
VAR L1 X,Y,S	1.5419249E-02
VAR H2 X,Y,S	1.5584986E-02
VAR L2 X,Y,S	1.5639862E-02
VAR H3 X,Y,S	1.5814118E-02
	8.4970109E-03
	2.09957269E-02
	8.0983795E-02
	2.1194205E-02
	8.1332467E-02
	2.1861432E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	1.8968448E-03
DIF L1 A,C,S	1.8962764E-03
DIF H2 A,C,S	1.9466531E-03
DIF L2 A,C,S	1.9159885E-03
DIF H3 A,C,S	1.9971788E-03
	1.6038719E-03
	1.6058472E-03
	1.61253349E-03
	1.6289004E-03
	1.6473396E-03
	1.6608448E-03
	1.6803905E-03
	1.7021203E-03
	1.7181376E-03
	1.7448852E-03
	1.7587053E-03
	1.7977506E-03
	1.805531E-03
	1.8388057E-03
	1.8518376E-03
	1.8799538E-03
	1.8891388E-03

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	1.0649983E-03
VAR H1 A,C,S	1.2168342E-02
VAR L1 A,C,S	1.1857739E-02
VAR H2 A,C,S	1.2192477E-02
VAR L2 A,C,S	1.1861636E-02
VAR H3 A,C,S	1.5325218E-02
	8.3549902E-03
	9.703888E-03
	1.0703888E-03
	1.1928142E-02
	2.0546551E-02
	3.4724467E-02
	4.995507E-02
	5.752196E-02
	6.3449983E-03
	7.5939805E-02
	12.19222
	16.86922
	12.52579
	11.57333
	10.03271

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	1.718274
L1(θ,θ), HIGH, FACTOR	1.773931
H2(θ,θ), HIGH, FACTOR	1.824061
L2(θ,θ), HIGH, FACTOR	1.985507
H3(θ,θ), HIGH, FACTOR	1.935853
	3.5939805E-02
	4.995507E-02
	5.752196E-02
	6.3449983E-03
	7.5939805E-02
	12.19222

## CASE13.DAT

## CASE14.DAT

PARAMETERS	3.7500001E-92	64.000000	56789.00
1.2000000	<b>6.0000000E+00</b>	<b>96.000000</b>	<b>64.000000</b>
12.000000	<b>5.0000000E+00</b>	<b>0.0000000E+00</b>	<b>160.000000</b>
53.200000	<b>1.0200000E-09</b>	<b>99999.00</b>	<b>99999.00</b>
1.3370000			

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ, φ), HIGH, FACTOR	<b>0.9463329</b>	1.6765045E-02	21.87794
L1(θ, φ), HIGH, FACTOR	<b>0.9533588</b>	1.6673784E-02	21.99745
H2(θ, φ), HIGH, FACTOR	<b>1.131199</b>	1.5293626E-02	23.98287
L2(θ, φ), HIGH, FACTOR	<b>1.324701</b>	1.3732953E-02	26.70839
H3(θ, φ), HIGH, FACTOR	<b>0.8646821</b>	1.5217487E-02	24.10287

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	<b>1.0605892E-02</b>	<b>6.3896294E-03</b>	<b>1.6995508E-02</b>
VCR H1 A,C,S	<b>7.5966557E-03</b>	<b>1.4392884E-02</b>	<b>2.1989554E-02</b>
VAR L1 A,C,S	<b>7.6394922E-03</b>	<b>1.4429245E-02</b>	<b>2.2059659E-02</b>
VAR H2 A,C,S	<b>7.6619784E-03</b>	<b>1.4468417E-02</b>	<b>2.2122462E-02</b>
VAR L2 A,C,S	<b>7.8185344E-03</b>	<b>1.4639320E-02</b>	<b>2.2457868E-02</b>
VAR H3 A,C,S	<b>8.6986593E-03</b>	<b>1.6235027E-02</b>	<b>2.4933942E-02</b>
VAR Sθ X,Y,S	<b>6.3896300E-03</b>	<b>1.0605880E-02</b>	<b>1.6995508E-02</b>
VAR H1 X,Y,S	<b>1.4392884E-02</b>	<b>7.5966552E-03</b>	<b>2.1989554E-02</b>
VAR L1 X,Y,S	<b>1.4429243E-02</b>	<b>7.6394922E-03</b>	<b>2.2059659E-02</b>
VAR H2 X,Y,S	<b>1.4466617E-02</b>	<b>7.6619759E-03</b>	<b>2.2122492E-02</b>
VAR L2 X,Y,S	<b>1.4639320E-02</b>	<b>7.8185434E-03</b>	<b>2.2457868E-02</b>
VAR H3 X,Y,S	<b>1.6235027E-02</b>	<b>8.6986593E-03</b>	<b>2.4933942E-02</b>

## INTEGRATED VARIANCE FROM POWER SPECTRA

H1(θ, φ), HIGH, FACTOR	<b>0.6396829</b>	1.8694512E-02	19.61998
L1(θ, φ), HIGH, FACTOR	<b>0.6658332</b>	1.9711750E-02	18.60744
H2(θ, φ), HIGH, FACTOR	<b>0.7455062</b>	1.8148819E-02	26.24986
L2(θ, φ), HIGH, FACTOR	<b>0.8776498</b>	1.8620441E-02	19.69838
H3(θ, φ), HIGH, FACTOR	<b>0.5800271</b>	1.8138792E-02	26.22163

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	<b>2.3762535E-03</b>	<b>2.4902624E-03</b>	<b>4.8664468E-03</b>
DIF L1 A,C,S	<b>2.35334810E-03</b>	<b>2.5223893E-03</b>	<b>4.8757835E-03</b>
DIF H2 A,C,S	<b>2.3462176E-03</b>	<b>2.53339912E-03</b>	<b>4.8892805E-03</b>
DIF L2 A,C,S	<b>2.28086337E-03</b>	<b>2.6498614E-03</b>	<b>4.9397286E-03</b>
DIF H3 A,C,S	<b>2.2555195E-03</b>	<b>3.153387E-03</b>	<b>5.5703185E-03</b>
DIF H1 X,Y,S	<b>2.4902824E-03</b>	<b>2.3762533E-03</b>	<b>4.8664468E-03</b>
DIF L1 X,Y,S	<b>2.5223891E-03</b>	<b>2.3534810E-03</b>	<b>4.8757835E-03</b>
DIF H2 X,Y,S	<b>2.5335920E-03</b>	<b>2.3462176E-03</b>	<b>4.8892805E-03</b>
DIF L2 X,Y,S	<b>2.6498614E-03</b>	<b>2.28086337E-03</b>	<b>4.9397286E-03</b>
DIF H3 X,Y,S	<b>3.3153387E-03</b>	<b>2.2555193E-03</b>	<b>5.5703185E-03</b>

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	<b>1.1008562E-03</b>	<b>1.5590338E-03</b>	<b>2.6679211E-03</b>
DIF L1 X,Y,S	<b>1.1370622E-03</b>	<b>1.6509298E-03</b>	<b>2.7879935E-03</b>
DIF H2 X,Y,S	<b>1.0981616E-03</b>	<b>1.4826785E-03</b>	<b>2.5808439E-03</b>
DIF L2 X,Y,S	<b>1.0966852E-03</b>	<b>1.5342711E-03</b>	<b>2.6339511E-03</b>
DIF H3 X,Y,S	<b>1.1225190E-03</b>	<b>1.5845113E-03</b>	<b>2.707024E-03</b>

## CASE15.DAT

## CASE16.DAT

PARAMETERS	
1.239999	3.750000E-02
12.98999	9.443333E+00
53.26000	9.090909
1.337000	1.323900E-09

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR SU A,C,S	1.0605882E-02
VAR HI A,C,S	6.3891296E-03
VAR LI A,C,S	1.5788292E-03
VAR L1 A,C,S	2.5523887E-02
VAR H0 A,C,S	1.6122219E-02
VAR L0 A,C,S	2.5790431E-02
VAR H3 A,C,S	1.6218759E-02
VAR L3 A,C,S	2.7143553E-02
VAR H2 A,C,S	1.9374779E-02
VAR L2 A,C,S	1.1877696E-02
VAR H4 A,C,S	1.7747095E-02
VAR L4 A,C,S	2.9824552E-02
VAR H5 A,C,S	1.6272576E-02
VAR L5 A,C,S	1.6703218E-02

## INTEGRATED SHAPED DEVIATION FROM DIF SPECTRA

DIF HI A,C,S	1.9768853E-03
DIF LI A,C,S	2.9677858E-03
DIF H0 A,C,S	3.2466248E-03
DIF L0 A,C,S	5.2519887E-03
DIF H3 A,C,S	5.1198294E-03
DIF L3 A,C,S	1.9168414E-03
DIF H4 A,C,S	3.2988921E-03
DIF L4 A,C,S	1.9924779E-02
DIF H5 A,C,S	1.7747005E-02
DIF L5 A,C,S	1.6703226E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 X,Y,S	5.3893300E-03
VAR HI X,Y,S	9.8175956E-03
VAR LI X,Y,S	1.6127215E-02
VAR H0 X,Y,S	1.6218753E-02
VAR L0 X,Y,S	1.7747095E-02
VAR H3 X,Y,S	1.6703218E-02
VAR L5 X,Y,S	1.60605882E-02
VAR H2 X,Y,S	9.8175956E-03
VAR L4 X,Y,S	2.5523887E-02
VAR H1 X,Y,S	2.5790431E-02
VAR L3 X,Y,S	2.7143553E-02
VAR H4 X,Y,S	1.6218759E-02
VAR L2 X,Y,S	1.8776968E-02
VAR H5 X,Y,S	1.6703226E-02

## INTEGRATED SHAPED DEVIATION FROM DIF SPECTRA

DIF HI X,Y,S	2.9677858E-03
DIF LI X,Y,S	3.2466248E-03
DIF H0 X,Y,S	5.2519887E-03
DIF L0 X,Y,S	5.1198294E-03
DIF H3 X,Y,S	3.2988921E-03
DIF L3 X,Y,S	1.9924779E-02
DIF H4 X,Y,S	1.7747005E-02
DIF L4 X,Y,S	1.6703218E-02
DIF H5 X,Y,S	1.60605882E-02
DIF L5 X,Y,S	9.8175956E-03
DIF H1 X,Y,S	2.5523887E-02
DIF L1 X,Y,S	2.7143553E-02
DIF H0 X,Y,S	1.6218759E-02
DIF L0 X,Y,S	1.8776968E-02
DIF H3 X,Y,S	1.6703226E-02

DIF HI X,Y,S	2.9677858E-03
DIF LI X,Y,S	3.2466248E-03
DIF H0 X,Y,S	5.2519887E-03
DIF L0 X,Y,S	5.1198294E-03
DIF H3 X,Y,S	3.2988921E-03
DIF L3 X,Y,S	1.9924779E-02
DIF H4 X,Y,S	1.7747005E-02
DIF L4 X,Y,S	1.6703218E-02
DIF H5 X,Y,S	1.60605882E-02
DIF L5 X,Y,S	9.8175956E-03
DIF H1 X,Y,S	2.5523887E-02
DIF L1 X,Y,S	2.7143553E-02
DIF H0 X,Y,S	1.6218759E-02
DIF L0 X,Y,S	1.8776968E-02
DIF H3 X,Y,S	1.6703226E-02

PARAMETERS	
1.200000	3.750000E-02
12.10000	9.000000E+00
53.20000	135.0000
1.337000	1.02550000E-09

## INTEGRATED VARIANCE FROM POWER SPECTRA

PARAMETERS	
1.200000	3.750000E-02
12.10000	9.000000E+00
53.20000	135.0000
1.337000	1.02550000E-09

PARAMETERS	
1.239999	3.750000E-02
12.98999	9.443333E+00
53.26000	9.090909
1.337000	1.323900E-09

## INTEGRATED VARIANCE FROM POWER SPECTRA

PARAMETERS	
1.200000	3.750000E-02
12.10000	9.000000E+00
53.20000	135.0000
1.337000	1.02550000E-09

## CASE17.DAT

## CASE18.DAT

PARAMETERS

1.209889	3.7529161E-02	64.00000	56789.49
12.999999	9.0990000E+00	98.00000	64.00000
53.20000	186.00000	53.20000	9.8854000
1.33799	1.0200000E-02	99999.98	99999.98

## MAGNITUDE ARRAY SCALING FACTORS

H5(θ, φ), HIGH, FACTOR	9.4501194	9.285694E-03	39.50002
L1(θ, φ), HIGH, FACTOR	9.4567370	9.9748746E-03	36.77000
H6(θ, φ), HIGH, FACTOR	8.5560228	8.46623397E-03	43.32331
L2(θ, φ), HIGH, FACTOR	9.6685336	8.2367899E-03	43.95006
H3(θ, φ), HIGH, FACTOR	8.4175338	8.6326599E-03	42.48008

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	1.0005882E-32	6.3896294E-03	1.6995548E-02
VAR H1 A,C,S	7.4455338E-03	1.4351268E-02	2.1796813E-02
VAR L1 A,C,S	7.5716837E-03	1.4558979E-02	2.2429759E-02
VAR H2 A,C,S	7.4706936E-03	1.4363205E-02	2.1133957E-02
VAR L2 A,C,S	7.6731676E-03	1.4659279E-02	2.2332411E-02
VAR H3 A,C,S	7.9204839E-03	1.5083088E-02	2.3663578E-02
VAR S0 X,Y,S	6.38953399E-03	1.8605882E-02	1.6995548E-02
VAR H1 X,Y,S	1.4351249E-02	7.4455533E-03	2.1796813E-02
VAR L1 X,Y,S	1.4558017E-02	7.5716833E-03	2.2429759E-02
VAR H2 X,Y,S	1.4363222E-02	7.4706981E-03	2.1133850E-02
VAR L2 X,Y,S	1.4659298E-02	7.6731676E-03	2.2332441E-02
VAR H3 X,Y,S	1.5083088E-02	7.9204858E-03	2.3693578E-02

## INTEGRATED SCATTERED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	2.4673914E-03	2.4226152E-03	4.9993203E-03
DIF L1 A,C,S	2.4541882E-03	2.4776152E-03	4.9993203E-03
DIF H2 A,C,S	2.4669615E-03	2.5192988E-03	4.9734833E-03
DIF L2 A,C,S	2.4494544E-03	2.4592316E-03	4.9861924E-03
DIF H3 A,C,S	2.4946224E-03	2.5900770E-03	4.9462244E-03
DIF H1 X,Y,S	2.4226152E-03	2.7789889E-03	5.1466175E-03
DIF L1 X,Y,S	2.5192988E-03	2.4541882E-03	4.9734833E-03
DIF H2 X,Y,S	2.4669615E-03	2.5192988E-03	4.9661924E-03
DIF L2 X,Y,S	2.4494544E-03	2.4669615E-03	4.9946224E-03
DIF H3 X,Y,S	2.4946224E-03	2.3677224E-03	5.1466175E-03

## INTEGRATED VARIANCE FROM DIF SPECTRA

VAR S0 A,C,S	1.0005882E-02	6.3896294E-03	1.6995548E-02
VAR H1 A,C,S	7.4455338E-03	1.4351268E-02	2.1796813E-02
VAR L1 A,C,S	7.5716833E-03	1.4558017E-02	2.2429759E-02
VAR H2 A,C,S	7.4706981E-03	1.4363222E-02	2.1133850E-02
VAR L2 A,C,S	7.6731676E-03	1.4659298E-02	2.2332441E-02
VAR H3 A,C,S	7.9204858E-03	1.5083088E-02	2.3693578E-02

## INTEGRATED SCATTERED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	2.3363772E-03	1.5902841E-03	3.9266543E-03
DIF L1 A,C,S	2.4134463E-03	2.4134463E-03	3.9866543E-03
DIF H2 A,C,S	2.4134463E-03	2.5752685E-03	3.9866543E-03
DIF L2 A,C,S	2.317305E-03	2.5896175E-03	3.8979221E-03
DIF H3 A,C,S	2.3966614E-03	2.5198431E-03	3.9359339E-03
DIF H1 X,Y,S	1.9596508E-03	2.1567922E-03	4.1164380E-03
DIF L1 X,Y,S	1.5732652E-03	2.3134463E-03	3.8979221E-03
DIF H2 X,Y,S	1.5806175E-03	2.3173617E-03	3.8979221E-03
DIF L2 X,Y,S	1.5498430E-03	2.3866994E-03	3.9359339E-03
DIF H3 X,Y,S	2.1567822E-03	2.5182228E-03	4.1164380E-03

## CASE19.DAT

## CASE20.DAT

PARAMETERS

1.29999E	3.7588691E-02	64.00000	56789.54
12.00000	6.0000000E+00	135.00000	64.00000
53.20000	8.0000000E+00	0.0000000E+00	100.00000
1.3370000	1.9296666E-09	99999.00	99999.00

MAGNITUDE ARRAY SIGNALING FACTORS

H1(θ, θ), HIGH, FACTOR	A.9648941	2.2538772E-02	19.29796
L1(θ, θ), HIGH, FACTOR	B.9744209	2.2283187E-02	19.51726
H2(θ, θ), HIGH, FACTOR	1.1583338	2.0418224E-02	21.39166
L2(θ, θ), HIGH, FACTOR	1.348093	1.8045209E-02	24.16543
H3(θ, θ), HIGH, FACTOR	0.8812735	2.3691868E-02	18.36643

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S <sub>0</sub> A,C,S	1.0649379E-02	6.3450006E-03	1.6944022E-02
VAR H <sub>1</sub> A,C,S	1.2474426E-02	7.693288E-03	2.0167669E-02
VAR L <sub>1</sub> A,C,S	1.2479666E-02	7.6789394E-03	2.0158570E-02
VAR H <sub>2</sub> A,C,S	1.2475561E-02	7.7030346E-03	2.0182539E-02
VAR L <sub>2</sub> A,C,S	1.2518639E-02	7.7522021E-03	2.0356998E-02
VAR H <sub>3</sub> A,C,S	1.4522227E-02	8.0869327E-03	2.3329284E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H <sub>1</sub> A,C,S	1.8731877E-03	1.4374188E-03	8.4970193E-03
DIF L <sub>1</sub> A,C,S	1.6732958E-03	1.4293238E-03	3.3025315E-03
DIF H <sub>2</sub> A,C,S	1.8729613E-03	1.4318322E-03	3.3447905E-03
DIF L <sub>2</sub> A,C,S	1.8812479E-03	1.4233522E-03	3.3645667E-03
DIF H <sub>3</sub> A,C,S	2.2386440E-03	1.5873476E-03	3.8259998E-03
DIF H <sub>1</sub> X,Y,S	1.5446621E-03	1.7659459E-03	3.3146858E-03
DIF L <sub>1</sub> X,Y,S	1.5611643E-03	1.7413466E-03	3.3025335E-03
DIF H <sub>2</sub> X,Y,S	1.5666699E-03	1.7441447E-03	3.3047985E-03
DIF L <sub>2</sub> X,Y,S	1.6255531E-03	1.6791188E-03	3.3045667E-03
DIF H <sub>3</sub> X,Y,S	2.2778877E-03	1.5482193E-03	3.8259998E-03

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S <sub>0</sub> A,C,S	1.0649379E-02	6.3450006E-03	1.6944022E-02
VAR H <sub>1</sub> A,C,S	1.2057226E-02	2.1421649E-02	3.3478897E-02
VAR L <sub>1</sub> A,C,S	1.1195125E-02	1.8774019E-02	2.9969225E-02
VAR H <sub>2</sub> A,C,S	1.2661862E-02	2.2619959E-02	3.5281815E-02
VAR L <sub>2</sub> A,C,S	1.247771E-02	2.1447405E-02	3.3925273E-02
VAR H <sub>3</sub> A,C,S	1.2326711E-02	1.9796136E-02	3.2032907E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H <sub>1</sub> A,C,S	2.8754934E-03	5.2791261E-03	8.1546092E-03
DIF L <sub>1</sub> A,C,S	2.6323320E-03	4.1007013E-03	6.7330245E-03
DIF H <sub>2</sub> A,C,S	3.0046252E-03	5.8335952E-03	8.8562272E-03
DIF L <sub>2</sub> A,C,S	2.8276551E-03	5.2983898E-03	8.1170527E-03
DIF H <sub>3</sub> A,C,S	2.4958027E-03	4.5361225E-03	7.0319348E-03
DIF H <sub>1</sub> X,Y,S	4.2035501E-03	3.9510699E-03	8.1546092E-03
DIF L <sub>1</sub> X,Y,S	3.6945355E-03	3.0384976E-03	6.7330245E-03
DIF H <sub>2</sub> X,Y,S	4.3325334E-03	4.5256889E-03	8.6582272E-03
DIF L <sub>2</sub> X,Y,S	4.1513611E-03	3.9656844E-03	8.1170527E-03
DIF H <sub>3</sub> X,Y,S	3.8882589E-03	3.1436733E-03	7.0319348E-03

## CASE21.DAT

## CASE22.DAT

PARAMETERS

1.200000	3.750000E-02	64.00000	56789.00
12.00000	11.900000E+00	135.00000	64.00000
53.20000	90.00000	53.20000	9.000000
1.337000	1.020000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	0.2329347	5.8767921E-03	74.01170
L1(θ,θ), HIGH, FACTOR	0.2533364	7.5843021E-03	57.96027
H2(θ,θ), HIGH, FACTOR	0.33886976	4.7378101E-03	91.80431
L2(θ,θ), HIGH, FACTOR	0.4648261	5.2006384E-03	83.63423
H3(θ,θ), HIGH, FACTOR	0.2058222	6.3799699E-03	68.17452

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	1.0049070E-02	6.3450000E-03	1.6994022E-02
VAR H1 A,C,S	1.2446433E-02	1.0162100E-02	2.2668571E-02
VAR L1 A,C,S	1.27966099E-02	9.8635970E-03	2.26559810E-02
VAR H2 A,C,S	1.24500500E-02	1.0752477E-02	2.3202548E-02
VAR L2 A,C,S	1.3410422E-02	1.1121614E-02	2.4539439E-02
VAR H3 A,C,S	1.40613366E-02	1.0235733E-02	2.4297021E-02
VAR S0 X,Y,S	8.49700063E-03	8.4970193E-03	1.6994022E-02
VAR H1 X,Y,S	1.5339355E-02	6.2691723E-03	2.2668571E-02
VAR L1 X,Y,S	1.6457189E-02	6.2626032E-03	2.26559810E-02
VAR H2 X,Y,S	1.5464541E-02	6.7379853E-03	2.3202548E-02
VAR L2 X,Y,S	1.7317633E-02	7.2217535E-03	2.4539439E-02
VAR H3 X,Y,S	1.7455615E-02	6.8414137E-03	2.4297021E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

H1(θ,θ), HIGH, FACTOR	0.3381905	1.1072428E-02	39.28239
L1(θ,θ), HIGH, FACTOR	0.35008317	1.2275637E-02	35.45208
H2(θ,θ), HIGH, FACTOR	0.4439534	9.8857451E-03	43.99784
L2(θ,θ), HIGH, FACTOR	0.5623575	9.0033313E-03	43.91976
H3(θ,θ), HIGH, FACTOR	0.3103087	1.1537333E-02	37.69938

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	2.0265912E-03	1.6520032E-03	3.6786739E-03
DIF L1 A,C,S	2.1226272E-03	1.6601700E-03	3.7828344E-03
DIF H2 A,C,S	1.9840101E-03	1.7092113E-03	3.7076463E-03
DIF L2 A,C,S	2.1955555E-03	1.8705910E-03	4.0060524E-03
DIF H3 A,C,S	2.2591699E-03	1.7671233E-03	4.0262924E-03
DIF H1 X,Y,S	2.1400000E-03	1.5345000E-03	3.6786739E-03
DIF L1 X,Y,S	2.2122643E-03	1.5705000E-03	3.7828344E-03
DIF H2 X,Y,S	2.2691835E-03	1.4384500E-03	3.7076463E-03
DIF L2 X,Y,S	2.4596657E-03	1.40660524E-03	4.0262924E-03
DIF H3 X,Y,S	2.5249876E-03	1.5013000E-03	4.0262924E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	1.2266146E-03	1.2517134E-03	2.4783262E-03
DIF L1 A,C,S	1.29222739E-02	5.7449564E-03	1.86633786E-02
DIF H2 A,C,S	1.1627851E-03	1.2330773E-02	5.9710772E-03
DIF L2 A,C,S	1.2279820E-03	1.2504510E-02	6.0011218E-03
DIF H3 A,C,S	1.5181506E-03	1.3805358E-02	6.4473688E-03
DIF H1 X,Y,S	9.3429571E-04	1.5440372E-03	2.4783262E-03
DIF L1 X,Y,S	1.0195084E-03	1.5655711E-03	2.5851532E-03
DIF H2 X,Y,S	9.0629433E-04	1.4910952E-03	2.3973873E-03
DIF L2 X,Y,S	9.8367140E-04	1.4524108E-03	2.4360835E-03
DIF H3 X,Y,S	1.2935828E-03	1.4174134E-03	2.7107988E-03

## CASE23.DAT

## CASE24.DAT

PARAMETERS	VALUES	PARAMETERS	VALUES
1.261000	3.750000E-02	64.00000	36789.00
12.00000	1.000000E+00	135.0000	64.00000
53.20000	180.0000	53.20000	9.885000
1.337000	1.020000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1 (0, 0), HIGH, FACTOR	0.4576920	1.2644354E-02	34.370886
L1 (0, 0), HIGH, FACTOR	0.4668326	1.3780513E-02	31.56279
H0 (0, 0), HIGH, FACTOR	0.5534549	1.1457671E-02	37.91159
L0 (0, 0), HIGH, FACTOR	0.6783584	1.1407901E-02	38.12721
H3 (0, 0), HIGH, FACTOR	0.4254104	1.3632638E-02	33.37491

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S 1.0619079E-02	6.345000E-03	1.6994022E-02
VAR H1 A,C,S 1.2577793E-02	7.7298668E-03	2.6298755E-02
VAR L1 A,C,S 1.2657733E-02	7.7422562E-03	2.0399585E-02
VAR H0 A,C,S 1.2533963E-02	7.7895065E-03	2.0251391E-02
VAR L0 A,C,S 1.2639223E-02	7.7414624E-03	2.0371718E-02
VAR H3 A,C,S 1.4649088E-02	8.4563708E-03	2.2505350E-02
VAR S0 X,Y,S 8.4970063E-03	8.4970193E-03	1.6994022E-02
VAR H1 X,Y,S 1.4935043E-02	5.3637377E-03	2.02298755E-02
VAR L1 X,Y,S 1.4994601E-02	5.4859204E-03	2.0399585E-02
VAR H0 X,Y,S 1.4890516E-02	5.3698851E-03	2.0251391E-02
VAR L0 X,Y,S 1.4748119E-02	5.4224979E-03	2.0371718E-02
VAR H3 X,Y,S 1.6330823E-02	6.1745331E-03	2.2505350E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 1.3788976E-03	1.4762600E-03	3.3551589E-03
DIF L1 A,C,S 1.8932822E-03	1.4735788E-03	3.3668624E-03
DIF H0 A,C,S 1.3746982E-03	1.4673986E-03	3.3420858E-03
DIF L0 A,C,S 1.3925200E-03	1.4586588E-03	3.3511859E-03
DIF H3 A,C,S 2.1329643E-03	1.5583459E-03	3.6113105E-03
DIF H1 X,Y,S 1.5232764E-03	1.8318817E-03	3.3551589E-03
DIF L1 X,Y,S 1.5574966E-03	1.8093779E-03	3.3668624E-03
DIF H0 X,Y,S 1.5223671E-03	1.8197218E-03	3.3420858E-03
DIF L0 X,Y,S 1.5740351E-03	1.7715720E-03	3.3511859E-03
DIF H3 X,Y,S 2.01438512E-03	1.6474569E-03	3.6913165E-03

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S 1.0619079E-02	6.3450000E-03	1.6994022E-02
VAR H1 A,C,S 1.5164242E-02	8.9362729E-03	2.4106666E-02
VAR L1 A,C,S 1.4539931E-02	8.4992275E-03	2.3930733E-02
VAR H0 A,C,S 1.5228570E-02	8.9961477E-03	2.4274705E-02
VAR L0 A,C,S 1.4670216E-02	8.5668219E-03	2.3230599E-02
VAR H3 A,C,S 2.2699258E-02	1.3838897E-02	3.5159152E-02
VAR S0 X,Y,S 8.4970063E-03	8.4970193E-03	1.6994022E-02
VAR H1 X,Y,S 1.7226971E-02	6.8735330E-03	2.4106666E-02
VAR L1 X,Y,S 1.6579717E-02	6.4539933E-03	2.3930733E-02
VAR H0 X,Y,S 1.7318441E-02	6.9528870E-03	2.4274705E-02
VAR L0 X,Y,S 1.6677121E-02	6.5535214E-03	2.3230599E-02
VAR H3 X,Y,S 2.2277975E-02	1.2872167E-02	3.5159152E-02
DIF H1 X,Y,S 2.0341277E-03	1.5854259E-03	3.6119528E-03
DIF L1 X,Y,S 1.9333727E-03	1.5377899E-03	3.4688619E-03
DIF H0 X,Y,S 2.0487618E-03	1.5877879E-03	3.6366410E-03
DIF L0 X,Y,S 1.9412488E-03	1.5298044E-03	3.4710569E-03
DIF H3 X,Y,S 3.8419475E-03	2.4770389E-03	6.3189911E-03
DIF H1 X,Y,S 2.1349054E-03	1.4846519E-03	3.6119528E-03
DIF L1 X,Y,S 1.5975699E-03	1.58666011E-03	3.4688619E-03
DIF H0 X,Y,S 2.1678414E-03	1.4687981E-03	3.6366410E-03
DIF L0 X,Y,S 1.9396804E-03	1.53313755E-03	3.4710569E-03
DIF H3 X,Y,S 4.2822739E-03	2.03367163E-03	6.3189911E-03

## CASE25.DAT

## CASE26.DAT

PARAMETERS	VALUES
1.200000	3.750000E-02
60.0000	64.00000
60.0000	0.0000000E+00
53.2000	0.4000000E+00
1.337000	1.0200000E-09

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ, φ), HIGH, FACTOR	0.9273533	5.862499E-03	94.74674
L1(θ, φ), HIGH, FACTOR	0.9724474	5.651150E-03	98.29025
H0(θ, φ), HIGH, FACTOR	1.108595	6.735320E-03	82.46668
L0(θ, φ), HIGH, FACTOR	1.334552	1.2614572E-02	44.03265
H3(θ, φ), HIGH, FACTOR	0.9492382	1.9311747E-02	28.76244

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 A,C,S	9.7741479E-02	4.5564171E-02	0.1433059
VAR L1 A,C,S	9.8582393E-02	4.0327933E-02	0.1389104
VAR H0 A,C,S	9.3708426E-02	4.0718209E-02	0.1344265
VAR L0 A,C,S	5.2038904E-02	1.7553733E-02	6.9392565E-02
VAR H3 A,C,S	0.1370472	9.4982139E-02	0.2311293
VAR S0 X,Y,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 X,Y,S	9.7741479E-02	4.5564171E-02	0.1433059
VAR L1 X,Y,S	9.8582393E-02	4.0327933E-02	0.1389104
VAR H0 X,Y,S	9.3708426E-02	4.0718209E-02	0.1344265
VAR L0 X,Y,S	5.2038904E-02	1.7553733E-02	6.9392565E-02
VAR H3 X,Y,S	0.1370472	9.4982139E-02	0.2311293

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 A,C,S	3.9124074E-02	4.5562354E-02	7.3686332E-02
VAR L1 A,C,S	3.2562204E-02	4.0115332E-02	7.2677612E-02
VAR H0 A,C,S	3.6109917E-02	4.5407683E-02	8.1597589E-02
VAR L0 A,C,S	4.0980197E-02	5.0730448E-02	9.1710791E-02
VAR H3 A,C,S	4.9804887E-02	4.1850913E-02	9.06555819E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	2.5031953E-02	7.6490908E-03	3.2681040E-02
DIF L1 A,C,S	1.9655135E-02	5.7003810E-03	2.5365555E-02
DIF H0 A,C,S	2.8931663E-02	5.8132333E-03	2.6744921E-02
DIF L0 A,C,S	4.1576944E-03	5.8132330E-03	2.6744921E-02
DIF H3 A,C,S	4.3148788E-02	3.2640668E-02	7.5789266E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF R1 X,Y,S	3.7500000E-02	64.00000	56789.00
DIF L1 X,Y,S	60.00000	0.0000000E+00	64.00000
DIF H0 X,Y,S	53.20000	45.00000	53.20000
DIF L0 X,Y,S	1.337000	1.0200000E-09	99999.00
DIF H3 X,Y,S	1.0200000E-09	99999.00	99999.00

## CASE27.DAT

## CASE28.DAT

PARAMETERS

1.266666	3.7500001E-02	64.00000	56789.00
68.00000	0.0000000E+00	0.0000000E+00	64.00000
53.20000	98.00000	53.20000	9.885000
1.337000	1.6260000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	0.2468761	2.5686375E-03	216.2442
L1(θ,θ), HIGH, FACTOR	0.2882440	2.9682522E-03	187.1313
Hθ(θ,θ), HIGH, FACTOR	0.351859	3.3398822E-03	166.3092
Lθ(θ,θ), HIGH, FACTOR	0.4981435	5.5835275E-03	99.48866
H3(θ,θ), HIGH, FACTOR	0.2282159	5.8940407E-03	94.23976

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	3.2197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 A,C,S	6.5942393E-02	6.1779581E-02	5.1268222
VAR L1 A,C,S	5.4496866E-02	5.6783097E-02	0.1112735
VAR Hθ A,C,S	6.6251434E-02	4.3985818E-02	0.1102372
VAR Lθ A,C,S	5.2103356E-02	2.5900764E-02	7.8804196E-02
VAR H3 A,C,S	0.1126230	7.6469347E-02	0.1890924
VAR Sθ X,Y,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 X,Y,S	6.5942393E-02	6.1779581E-02	5.1268222
VAR L1 X,Y,S	5.4496866E-02	5.6783097E-02	0.1112735
VAR Hθ X,Y,S	6.6251434E-02	4.3985818E-02	0.1102372
VAR Lθ X,Y,S	5.2103356E-02	2.5900764E-02	7.8804196E-02
VAR H3 X,Y,S	0.1126230	7.6469347E-02	0.1890924

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	1.18717564E-02	1.4425768E-02	2.6273319E-02
DIF L1 A,C,S	9.3972906E-03	1.2195912E-02	2.1593204E-02
DIF Hθ A,C,S	9.6537685E-03	7.7353340E-03	1.7389590E-02
DIF Lθ A,C,S	5.1600445E-03	3.7955195E-03	8.9556267E-03
DIF H3 A,C,S	2.9326528E-02	2.2957630E-02	5.2284252E-02
DIF H1 X,Y,S	1.18717564E-02	1.4425768E-02	2.6273319E-02
DIF L1 X,Y,S	9.3972906E-03	1.2195912E-02	2.1593204E-02
DIF Hθ X,Y,S	9.6537685E-03	7.7353340E-03	1.7389590E-02
DIF Lθ X,Y,S	5.1600445E-03	3.7955195E-03	8.9556267E-03
DIF H3 X,Y,S	2.9326528E-02	2.2957630E-02	5.2284252E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 A,C,S	6.4175986E-02	4.0676013E-02	6.1046520
VAR L1 A,C,S	7.1016520E-02	4.5425445E-02	6.1164419
VAR Hθ A,C,S	7.5869337E-02	5.2499799E-02	6.1283696
VAR Lθ A,C,S	4.2559907E-02	2.2385295E-02	6.494513E-02
VAR H3 A,C,S	7.6284966E-02	4.7607768E-02	6.1238929

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	9.8184776E-03	5.4872185E-03	1.5305679E-02
DIF L1 A,C,S	1.2118598E-02	6.9576045E-03	1.9076185E-02
DIF Hθ A,C,S	1.4724538E-02	9.1596805E-03	2.3889283E-02
DIF Lθ A,C,S	5.8698435E-03	2.7391686E-03	7.8000198E-03
DIF H3 A,C,S	1.3772319E-02	9.55384624E-03	2.3319764E-02

## CASE29.DAT

## CASE39.DAT

PARAMETERS	VALUES
H1(θ,θ), HIGH, FACTOR	0.4629414
L1(θ,θ), HIGH, FACTOR	0.4291545
H2(θ,θ), HIGH, FACTOR	0.5976812
L2(θ,θ), HIGH, FACTOR	0.6399342
H3(θ,θ), HIGH, FACTOR	0.4364834

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	0.4629414	4.5045316E-03	1.253398
L1(θ,θ), HIGH, FACTOR	0.4291545	4.3884119E-03	138.7103
H2(θ,θ), HIGH, FACTOR	0.5976812	3.6818879E-03	150.8609
L2(θ,θ), HIGH, FACTOR	0.6399342	3.4696939E-03	162.9442
H3(θ,θ), HIGH, FACTOR	0.4364834	1.3238719E-02	41.95670

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ X,Y,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 X,Y,S	7.5474535E-02	3.89262927E-02	0.1664375
VAR L1 X,Y,S	8.3656619E-02	3.5928752E-02	9.1195794
VAR H2 X,Y,S	8.7784559E-02	4.0736794E-02	0.1285213
VAR L2 X,Y,S	0.1072620	5.0374527E-02	0.1576564
VAR H3 X,Y,S	3.9198286E-02	5.2698194E-02	0.1417964
VAR Sθ X,Y,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 X,Y,S	7.5474535E-02	3.89262927E-02	0.1664375
VAR L1 X,Y,S	8.3656619E-02	3.5928752E-02	9.1195794
VAR H2 X,Y,S	8.7784559E-02	4.0736794E-02	0.1285213
VAR L2 X,Y,S	0.1072620	5.0374527E-02	0.1576564
VAR H3 X,Y,S	3.9198286E-02	5.2698194E-02	0.1417964

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	1.3958504E-02	3.5532389E-03	1.6611731E-02
DIF L1 X,Y,S	1.6288742E-02	4.7883856E-03	2.1063014E-02
DIF H2 X,Y,S	2.0065642E-02	6.2149307E-03	2.6824576E-02
DIF L2 X,Y,S	2.8159866E-02	9.2890101E-03	3.7439838E-02
DIF H3 X,Y,S	1.8331535E-02	1.1892287E-02	3.0133776E-02
DIF H1 X,Y,S	1.3958504E-02	3.5532389E-03	1.6611731E-02
DIF L1 X,Y,S	1.6288742E-02	4.7883856E-03	2.1063014E-02
DIF H2 X,Y,S	2.0065642E-02	6.2149307E-03	2.6824576E-02
DIF L2 X,Y,S	2.8159866E-02	9.2890101E-03	3.7439838E-02
DIF H3 X,Y,S	1.8331535E-02	1.1892287E-02	3.0133776E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ X,Y,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 X,C,S	6.5950297E-02	2.8206635E-02	9.4157174E-02
VAR L1 X,C,S	6.5326504E-02	2.7649645E-02	9.2976199E-02
VAR H2 X,C,S	6.8586290E-02	3.6690166E-02	9.5276266E-02
VAR L2 X,C,S	7.0767589E-02	3.2647718E-02	0.1034153
VAR H3 X,C,S	7.6551653E-02	3.9498415E-02	0.1166499
VAR Sθ X,Y,S	3.9197806E-02	2.3548655E-02	6.2746458E-02
VAR H1 X,Y,S	6.5950297E-02	2.8206635E-02	9.4157174E-02
VAR L1 X,Y,S	6.5326504E-02	2.7649645E-02	9.2976199E-02
VAR H2 X,Y,S	6.8586290E-02	3.6690166E-02	9.5276266E-02
VAR L2 X,Y,S	7.0767589E-02	3.2647718E-02	0.1034153
VAR H3 X,Y,S	7.6551653E-02	3.9498415E-02	0.1166499

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,C,S	9.7948662E-03	3.0024901E-03	1.2796566E-02
DIF L1 X,C,S	9.377541E-03	2.3989910E-03	1.2286661E-02
DIF H2 X,C,S	1.1351245E-02	3.4467283E-03	1.479798E-02
DIF L2 X,C,S	1.2535811E-02	3.8518833E-03	1.6387267E-02
DIF H3 X,C,S	1.3716412E-02	6.1443662E-03	1.9854786E-02
DIF H1 X,Y,S	9.7948662E-03	3.0024901E-03	1.2796566E-02
DIF L1 X,Y,S	9.377541E-03	2.3989910E-03	1.2286661E-02
DIF H2 X,Y,S	1.1351245E-02	3.4467283E-03	1.479798E-02
DIF L2 X,Y,S	1.2535811E-02	3.8518833E-03	1.6387267E-02
DIF H3 X,Y,S	1.3716412E-02	6.1443662E-03	1.9854786E-02

## CASE31.DAT

## CASE32.DAT

PARAMETERS	VALUES
H1( $\theta$ , $\varphi$ ), HIGH, FACTOR	3.7599991E-02
L1( $\theta$ , $\varphi$ ), HIGH, FACTOR	6.9999995E+00
H2( $\theta$ , $\varphi$ ), HIGH, FACTOR	11.0000002E+00
L2( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.0200000E-09
H3( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.3370000E-09

## MAGNITUDE ARRAY SCALING FACTORS

H1( $\theta$ , $\varphi$ ), HIGH, FACTOR	3.7621119	6.043743E-03	100.2137
L1( $\theta$ , $\varphi$ ), HIGH, FACTOR	4.7523417E-03	127.4598	
H2( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.143699	5.8417576E-03	103.6895
L2( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.359572	1.9445863E-02	57.98745
H3( $\theta$ , $\varphi$ ), HIGH, FACTOR	9.9714996	2.0316185E-02	29.81509

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9235316E-02	2.3511158E-02	6.2746525E-02
VAR H1 A,C,S	8.5413337E-02	6.9933858E-02	6.15533171
VAR L1 A,C,S	0.1944368	8.2555898E-02	0.1865927
VAR H2 A,C,S	9.62232435E-02	7.4296872E-02	0.1785142
VAR L2 A,C,S	4.4564535E-02	2.9915985E-02	7.4872285E-02
VAR H3 A,C,S	0.1223952	0.1082555	0.23386513
VAR S0 X,Y,S	3.1373270E-02	3.1373236E-02	6.2746525E-02
VAR H1 X,Y,S	9.9381964E-02	5.5935159E-02	9.15533171
VAR L1 X,Y,S	0.1167732	6.9019696E-02	0.1865927
VAR H2 X,Y,S	0.1800948	6.2419403E-02	0.1785142
VAR L2 X,Y,S	5.6221733E-02	2.4459579E-02	7.4872285E-02
VAR H3 X,Y,S	0.1341004	9.6550588E-02	0.23706513

## INTEGRATED VARIANCE FROM POWER SPECTRA

H1( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.2400000	3.7500001E-02	64.00000	56789.99
L1( $\theta$ , $\varphi$ ), HIGH, FACTOR	6.0000000	6.0000000E-09	45.00000	64.00000
H2( $\theta$ , $\varphi$ ), HIGH, FACTOR	53.20000	45.00000	53.20000	9.000000
L2( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.3370000	1.0200000E-09	99999.99	99999.99

PARAMETERS	VALUES		
H1( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.6330251E-02	37.00244	
L1( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.7139293E-02	35.34153	
H2( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.5358975E-02	39.50093	
L2( $\theta$ , $\varphi$ ), HIGH, FACTOR	1.57675855E-02	38.41609	
H3( $\theta$ , $\varphi$ ), HIGH, FACTOR	0.6380583	2.7716220E-02	21.81532

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	4.4948398E-03	3.2808841E-03	7.775179E-03
DIF L1 A,C,S	4.3185819E-03	3.1257775E-03	7.4423687E-03
DIF H2 A,C,S	5.1085956E-03	4.2404168E-03	9.349062E-03
DIF L2 A,C,S	5.0026742E-03	4.5356697E-03	9.5663415E-03
DIF H3 A,C,S	6.7753885E-03	4.8398757E-03	1.1615276E-02
DIF H1 X,Y,S	4.0122755E-03	3.7334587E-03	7.775179E-03
DIF L1 X,Y,S	3.9631591E-03	3.4792160E-03	7.4423687E-03
DIF H2 X,Y,S	4.2480249E-03	5.1009860E-03	9.349062E-03
DIF L2 X,Y,S	4.3309461E-03	5.2353991E-03	9.5663415E-03
DIF H3 X,Y,S	5.4955402E-03	6.1197085E-03	1.1615276E-02

## CASE33.DAT

## CASE34.DAT

PARAMETERS	VALUES	VALUES	VALUES	VALUES
H1(θ, φ), HIGH, FACTOR	3.7500001E-02	64.00000	56789.00	56789.00
L1(θ, φ), HIGH, FACTOR	0.0000000E+00	45.00000	64.00000	64.00000
H2(θ, φ), HIGH, FACTOR	53.20000	90.00000	53.20000	9.085000
L2(θ, φ), HIGH, FACTOR	1.337000	1.0200000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ, φ), HIGH, FACTOR	6.2558848
L1(θ, φ), HIGH, FACTOR	0.2974848
H2(θ, φ), HIGH, FACTOR	0.3619846
L2(θ, φ), HIGH, FACTOR	0.5816676E-03
H3(θ, φ), HIGH, FACTOR	0.2389007

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9233516E-02
VAR H1 A,C,S	0.183620
VAR L1 A,C,S	9.1453705E-02
VAR H2 A,C,S	5.4536007E-02
VAR L2 A,C,S	3.8610999E-02
VAR H3 A,C,S	0.1070286
VAR S0 X,Y,S	3.1373270E-02
VAR H1 X,Y,S	9.3411572E-02
VAR L1 X,Y,S	7.751569E-02
VAR H2 X,Y,S	6.021392E-02
VAR L2 X,Y,S	4.8622258E-02
VAR H3 X,Y,S	0.1075170
VAR S0 Y,Z,S	3.1373236E-02
VAR H1 Y,Z,S	9.9769168E-02
VAR L1 Y,Z,S	8.5771956E-02
VAR H2 Y,Z,S	5.0123371E-02
VAR L2 Y,Z,S	3.1035131E-02
VAR H3 Y,Z,S	0.1075170

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	2.5611673E-02
DIF L1 A,C,S	1.8818287E-02
DIF H2 A,C,S	8.8477894E-03
DIF L2 A,C,S	6.2458614E-03
DIF H3 A,C,S	2.5754157E-02
DIF H1 X,Y,S	2.5559429E-02
DIF L1 X,Y,S	1.7619775E-02
DIF H2 X,Y,S	1.4496628E-02
DIF L2 X,Y,S	6.9517533E-03
DIF H3 X,Y,S	3.9618825E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9235310E-02
VAR H1 A,C,S	4.9008753E-02
VAR L1 A,C,S	4.9714219E-02
VAR H2 A,C,S	6.5647051E-02
VAR L2 A,C,S	6.3585125E-02
VAR H3 A,C,S	6.8250697E-02
VAR S0 X,Y,S	3.1373270E-02
VAR H1 X,Y,S	6.0197884E-02
VAR L1 X,Y,S	6.3646488E-02
VAR H2 X,Y,S	5.8926665E-02
VAR L2 X,Y,S	4.8839333E-02
VAR H3 X,Y,S	8.6116895E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	7.6029864E-02
DIF L1 A,C,S	8.1197983E-03
DIF H2 A,C,S	8.6946338E-03
DIF L2 A,C,S	6.8893256E-03
DIF H3 A,C,S	1.4161595E-02
DIF H1 X,Y,S	1.2257195E-02
DIF L1 X,Y,S	1.3467896E-02
DIF H2 X,Y,S	1.0400696E-02
DIF L2 X,Y,S	6.8665138E-03
DIF H3 X,Y,S	2.2264501E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	4.36880662E-03
DIF L1 A,C,S	4.2145676E-03
DIF H2 A,C,S	3.79984403E-03
DIF L2 A,C,S	4.3222889E-03
DIF H3 A,C,S	1.02318335E-02
DIF H1 X,Y,S	144.8678
DIF L1 X,Y,S	143.7227
DIF H2 X,Y,S	159.4690
DIF L2 X,Y,S	148.1434
DIF H3 X,Y,S	59.26031

## CASE35.DAT

## CASE36.DAT

PARAMETERS	
1.200000	3.750000E-02
69.000000	6.000000E+00
53.200000	18.000000
1.333333	1.020000E-09

MAGNITUDE ARRAY SCALING FACTORS	
H1(θ, φ), HIGH, FACTOR	0.4223123
L1(θ, φ), HIGH, FACTOR	0.4456466
H0(θ, φ), HIGH, FACTOR	0.5274680
L0(θ, φ), HIGH, FACTOR	0.6553380
H3(θ, φ), HIGH, FACTOR	0.4511395

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S9 A,C,S	3.9235316E-02
VAR H1 A,C,S	6.1390684E-02
VAR L1 A,C,S	5.7166379E-02
VAR H0 A,C,S	7.4146517E-02
VAR L0 A,C,S	9.1796599E-02
VAR H3 A,C,S	8.5651226E-02

VAR S9 X,Y,S	3.1373238E-02
VAR H1 X,Y,S	7.4933864E-02
VAR L1 X,Y,S	7.0258787E-02
VAR H0 X,Y,S	8.7915689E-02
VAR L0 X,Y,S	8.1121249
VAR H3 X,Y,S	9.5835447E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	1.0843133E-02
DIF L1 A,C,S	9.8538867E-03
DIF H0 A,C,S	1.5493833E-02
DIF L0 A,C,S	2.5187775E-02
DIF H3 A,C,S	1.7953616E-02

DIF H1 X,Y,S	1.5581161E-02
DIF L1 X,Y,S	1.3594833E-02
DIF H0 X,Y,S	2.3886112E-03
DIF L0 X,Y,S	3.6516260E-02
DIF H3 X,Y,S	2.4870556E-02

PARAMETERS	
1.200000	3.750000E-02
69.000000	6.000000E+00
53.200000	18.000000
1.333333	1.020000E-09

## INTEGRATED VARIANCE FROM POWER SPECTRA

H1(θ, φ), HIGH, FACTOR	5.1119193E-03
L1(θ, φ), HIGH, FACTOR	5.2172765E-03
H0(θ, φ), HIGH, FACTOR	3.9918926E-03
L0(θ, φ), HIGH, FACTOR	3.2330176E-03
H3(θ, φ), HIGH, FACTOR	1.3186139E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S9 A,C,S	3.9235316E-02
VAR H1 A,C,S	5.6786552E-02
VAR L1 A,C,S	5.5940866E-02
VAR H0 A,C,S	5.9770368E-02
VAR L0 A,C,S	6.1936982E-02
VAR H3 A,C,S	6.7552032E-02
VAR S9 X,Y,S	3.1373238E-02
VAR H1 X,Y,S	6.7284595E-02
VAR L1 X,Y,S	6.6567779E-02
VAR H0 X,Y,S	6.9952718E-02
VAR L0 X,Y,S	7.2644828E-02
VAR H3 X,Y,S	7.7539324E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	8.9125957E-03
DIF L1 A,C,S	8.6913276E-03
DIF H0 A,C,S	9.6998829E-03
DIF L0 A,C,S	1.0287875E-02
DIF H3 A,C,S	1.1378960E-02
DIF H1 X,Y,S	1.2516121E-02
DIF L1 X,Y,S	1.2693929E-02
DIF H0 X,Y,S	1.3986604E-02
DIF L0 X,Y,S	1.4995794E-02
DIF H3 X,Y,S	1.6545562E-02

H1(θ, φ), HIGH, FACTOR	1.210453
L1(θ, φ), HIGH, FACTOR	1.288554
H0(θ, φ), HIGH, FACTOR	1.315598
L0(θ, φ), HIGH, FACTOR	1.49845
H3(θ, φ), HIGH, FACTOR	1.579957
H1(θ, φ), HIGH, FACTOR	1.8328948E-02
L1(θ, φ), HIGH, FACTOR	1.9428537E-02
H0(θ, φ), HIGH, FACTOR	1.7045917E-02
L0(θ, φ), HIGH, FACTOR	1.6871074E-02
H3(θ, φ), HIGH, FACTOR	4.4668247E-02

H1(θ, φ), HIGH, FACTOR	33.04766
L1(θ, φ), HIGH, FACTOR	31.17773
H0(θ, φ), HIGH, FACTOR	35.53513
L0(θ, φ), HIGH, FACTOR	35.90319
H3(θ, φ), HIGH, FACTOR	13.57886
H1(θ, φ), HIGH, FACTOR	6.2744525E-02
L1(θ, φ), HIGH, FACTOR	6.1068789
H0(θ, φ), HIGH, FACTOR	9.9442236E-02
L0(θ, φ), HIGH, FACTOR	9.1663128
H3(θ, φ), HIGH, FACTOR	0.1169983
H1(θ, φ), HIGH, FACTOR	6.2746525E-02
L1(θ, φ), HIGH, FACTOR	6.1068789
H0(θ, φ), HIGH, FACTOR	9.9442236E-02
L0(θ, φ), HIGH, FACTOR	9.1663128
H3(θ, φ), HIGH, FACTOR	0.1222876

## CASE37.DAT

## CASE38.DAT

PARAMETERS			
1.239989	3.7598991E-02	64.00000	56789.00
65.00000	0.0000000E+00	90.00000	64.00000
53.20000	6.0200000E+02	0.0000000E+00	100.0000
1.337600	1.6200000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	1.014915	6.9529314E-03	79.98769
L1(θ,θ), HIGH, FACTOR	1.040339	5.37154735E-03	163.4465
H2(θ,θ), HIGH, FACTOR	1.195437	5.5982117E-03	99.21972
L2(θ,θ), HIGH, FACTOR	1.403373	4.6612527E-03	119.1639
H3(θ,θ), HIGH, FACTOR	0.9827399	2.6173647E-02	21.22184

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9197754E-02	2.3508668E-02	6.2745025E-02
VAR H1 A,C,S	4.5189317E-02	6.5274857E-02	9.1104643
VAR L1 A,C,S	5.6384034E-02	7.8228109E-02	9.1346122
VAR H2 A,C,S	5.38818812E-02	8.3116300E-02	6.1429352
VAR L2 A,C,S	7.8886442E-02	6.1138923	6.1927782
VAR H3 A,C,S	4.2269852E-02	5.7646727E-02	9.9916525E-02
VAR S0 X,Y,S	2.3548666E-02	3.9197753E-02	6.27465025E-02
VAR H1 X,Y,S	6.5274857E-02	4.5189321E-02	6.1104640
VAR L1 X,Y,S	7.8228109E-02	5.6384038E-02	9.1346122
VAR H2 X,Y,S	8.3116300E-02	5.9818800E-02	6.1429352
VAR L2 X,Y,S	6.1138923	7.8886442E-02	9.1927782
VAR H3 X,Y,S	5.1646729E-02	4.2269852E-02	9.9916525E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9197750E-02	2.3548668E-02	6.2746502E-02
VAR H1 A,C,S	3.9737398E-02	1.4534242E-02	5.4271653E-02
VAR L1 A,C,S	4.0062048E-02	1.4783937E-02	5.4845951E-02
VAR H2 A,C,S	3.939721E-02	1.4033389E-02	5.3431772E-02
VAR L2 A,C,S	3.9414889E-02	1.3807974E-02	5.3222116E-02
VAR H3 A,C,S	4.5634720E-02	2.5410837E-02	7.1045533E-02
VAR S0 X,Y,S	2.3548666E-02	3.9197750E-02	6.2746502E-02
VAR H1 X,Y,S	1.4534242E-02	3.9737398E-02	5.4271653E-02
VAR L1 X,Y,S	1.4783937E-02	4.0062048E-02	5.4845951E-02
VAR H2 X,Y,S	1.4033389E-02	3.939721E-02	5.3431772E-02
VAR L2 X,Y,S	1.3807974E-02	3.9414889E-02	5.3222116E-02
VAR H3 X,Y,S	2.5410837E-02	4.5634720E-02	7.1045533E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	7.333877E-03	1.7199839E-02	2.4533655E-02
DIF L1 A,C,S	9.9001387E-03	2.4624147E-02	9.9001387E-03
DIF H2 A,C,S	1.1930142E-02	2.4624147E-02	3.4524292E-02
DIF L2 A,C,S	1.7047080E-02	4.5483440E-02	6.2331279E-02
DIF H3 A,C,S	8.0947044E-03	1.3472479E-02	2.2367224E-02
DIF H1 X,Y,S	1.7199839E-02	7.3338482E-03	2.4533655E-02
DIF L1 X,Y,S	2.4624147E-02	9.9001387E-03	3.4524292E-02
DIF H2 X,Y,S	2.11055646E-02	1.0930141E-02	3.8985766E-02
DIF L2 X,Y,S	4.5483444E-02	1.7647880E-02	6.2531278E-02
DIF H3 X,Y,S	1.3472479E-02	8.0947044E-03	2.2367224E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	1.9872761E-03	1.9872761E-03	5.2897511E-03
DIF L1 X,Y,S	3.3698229E-03	2.133954E-03	5.5037448E-03
DIF H2 X,Y,S	3.2262821E-03	1.7421397E-03	4.9684262E-03
DIF L2 X,Y,S	3.2466056E-03	1.6638766E-03	4.9104756E-03
DIF H3 X,Y,S	3.5617799E-03	4.7156513E-03	8.2774376E-03

CASE 53. DAT

CASE 49. DAT

PARAMETERS		ARRAY SCALING FACTORS
1.200000	3.750000E-02	64.000000
69.000000	0.000000E+00	90.000000
53.200000	94.000000	53.200000
1.337000	1.020000E-09	99999.00

MAGNITUDE ARRAY SCALING FACTORS

INTEGRATED VARIANCE FROM POWER SPECTRA					
VAR SW A,C,S	3.919775E-#2	2.3548668E-#2	6.2746592E-#2		
VAR H1 A,C,S	8.4266164E-#2	4.5169353E-#2	8.1293756		
VAR L1 A,C,S	7.3721884E-#2	3.8913845E-#2	6.1126357		
VAR H9 A,C,S	6.3756111E-#2	3.9179188E-#2	5.1692153		
VAR LS A,C,S	5.1156182E-#2	3.6385457E-#2	8.7461598E-#2		
VAR H3 A,C,S	5.2767665E-#2	5.6367524E-#2	9.1691351		

VAR S0	X,Y,S	2.3548665E-02	3.9197759E-02	6.2746582E-02
VAR H1	X,Y,S	4.5169353E-02	8.4206171E-02	8.1293756
VAR L1	X,Y,S	3.39138955E-02	7.372894E-02	8.1126557
VAR H0	X,Y,S	3.9179188E-02	6.3738111E-02	6.1629153
VAR L0	X,Y,S	3.5395457E-02	5.1156182E-02	8.7461598E-02
VAR H3	X,Y,S	5.5357524E-02	5.2767655E-02	8.1091351

INTEGRATED SQUARED DEVIATION FROM ONE CONSTRAINED

DIF H1 A,C,S	1. <sup>-38</sup> /1589E-02	7. <sup>-60</sup> /07724E-03	2. <sup>-27</sup> 62371E-02	DIF H1 A,C,S	5.255122E-03	5. <sup>-66</sup> 17922E-03	5.2545122E-03	DIF H1 X,Y,S	5. <sup>-66</sup> 17922E-03	5.2545122E-03	5.0617922E-03	DIF H1 A,C,S	1.0316323E-02
DIF L1 A,C,S	1. <sup>-125</sup> 16727E-02	5. <sup>-56</sup> 08908E-03	1. <sup>-68</sup> 11926E-02	DIF L1 A,C,S	5. <sup>-89</sup> 56979E-03	5. <sup>-47</sup> 9350E-03	5. <sup>-89</sup> 56979E-03	DIF L1 X,Y,S	5. <sup>-47</sup> 9350E-03	5. <sup>-89</sup> 56979E-03	5. <sup>-89</sup> 56979E-03	DIF L1 A,C,S	1.1282813E-02
DIF H9 A,C,S	9. <sup>-112</sup> 9271E-03	6. <sup>-177</sup> 9823E-03	1. <sup>-52</sup> 90914E-02	DIF H9 A,C,S	4. <sup>-38</sup> 00886E-03	5. <sup>-16</sup> 88110E-03	5. <sup>-38</sup> 00886E-03	DIF H9 X,Y,S	5. <sup>-16</sup> 88110E-03	9.4128232E-03	5. <sup>-16</sup> 88110E-03	DIF H9 A,C,S	9.4128232E-03
DIF L9 A,C,S	6. <sup>-68</sup> 13389E-03	5. <sup>-37</sup> 44632E-03	1. <sup>-28</sup> 55739E-02	DIF L9 A,C,S	6. <sup>-39</sup> 99051E-03	7. <sup>-97</sup> 45293E-03	6. <sup>-39</sup> 99051E-03	DIF L9 X,Y,S	7. <sup>-97</sup> 45293E-03	1.4371535E-02	7. <sup>-97</sup> 45293E-03	DIF L9 A,C,S	1.4371535E-02
DIF H3 A,C,S	9. <sup>-21</sup> 01041E-03	1. <sup>-23</sup> 06225E-02	2. <sup>-15</sup> 16293E-02	DIF H3 A,C,S	7.545995E-03	7. <sup>-63</sup> 41467E-03	7. <sup>-63</sup> 41467E-03	DIF H3 X,Y,S	7. <sup>-63</sup> 41467E-03	1.4580738E-02	1.4580738E-02	DIF H3 A,C,S	1.4580738E-02

INTEGRATED VARIANCE FROM POWER SP

VAR S9 A,C,S	3.919773E-#2	2.3548668E-#2	6.2746502E-#2
VAR H1 A,C,S	3.7705429E-#2	3.2959731E-#2	7.0665054E-#2
VAR L1 A,C,S	3.6225753E-#2	3.59667722E-#2	7.133671E-#2
VAR H# A,C,S	4.8155911E-#2	3.5730141E-#2	8.2886174E-#2
VAR L# A,C,S	6.1713938E-#2	4.3587826E-#2	8.1057019
VAR H3 A,C,S	3.2066310E-#2	4.6605191E-#2	7.2668247E-#2
VAR S# X,Y,S	2.3548668E-#2	3.9197735E-#2	6.2746502E-#2
VAR H1 X,Y,S	3.2959737E-#2	3.7705429E-#2	7.0665054E-#2
VAR L1 X,Y,S	3.59667722E-#2	3.62257539E-#2	7.133671E-#2
VAR H# X,Y,S	3.4730148E-#2	4.8155988E-#2	8.2886174E-#2
VAR L# X,Y,S	4.3987826E-#2	6.1713931E-#2	8.1057019
VAR H3 X,Y,S	4.6605191E-#2	7.2668247E-#2	7.2668247E-#2

INTEGRATED SQUARED DEVIATION FROM DIFFERENT SPECTRA

DIF H1	A, <sub>1</sub> C, <sub>5</sub>	5.2545122E-03	<b>5.6617922E-03</b>	1.0336323E-02
DIF L1	A, <sub>1</sub> C, <sub>5</sub>	5.8925697E-03	5.4794358E-03	1.1282913E-02
DIF H9	A, <sub>1</sub> C, <sub>5</sub>	4.3940886E-03	5.1688111E-03	9.4128232E-03
DIF L9	A, <sub>1</sub> C, <sub>5</sub>	6.3969551E-03	7.9746293E-03	1.4377535E-02
DIF H3	A, <sub>1</sub> C, <sub>5</sub>	7.5465995E-03	<b>7.6341467E-03</b>	1.4599738E-02
DIF H1	X, <sub>1</sub> Y, <sub>5</sub>	5.6617922E-03	5.2545122E-03	1.0336323E-02
DIF L1	X, <sub>1</sub> Y, <sub>5</sub>	5.4794358E-03	5.8925697E-03	1.1282913E-02
DIF H9	X, <sub>1</sub> Y, <sub>5</sub>	5.1688111E-03	4.3940886E-03	9.4128232E-03
DIF L9	X, <sub>1</sub> Y, <sub>5</sub>	7.9746293E-03	6.3969551E-03	1.4377535E-02
DIF H3	X, <sub>1</sub> Y, <sub>5</sub>	<b>7.6341467E-03</b>	7.5465995E-03	1.4599738E-02

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MAGNITUDE ARRAY SCALING FACTORS			
H1(θ, φ), HIGH, FACTOR	0.3517864	5.3896423E-03	104.61
L1(θ, φ), HIGH, FACTOR	0.3779574	5.3783792E-03	103.27
H2(θ, φ), HIGH, FACTOR	0.4569333	4.3628574E-03	127.31
L2(θ, φ), HIGH, FACTOR	0.5882912	3.6361918E-03	152.75
H3(θ, φ), HIGH, FACTOR	0.3433594	1.3940362E-02	42.594

INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S9 A,C,S	3.9197739E-02	2.3586668E-02	6.2746562E-02
VAR H1 A,C,S	3.7765429E-02	3.29597737E-02	7.0665054E-02
VAR L1 A,C,S	3.6275738E-02	3.58687722E-02	7.1336717E-02
VAR H0 A,C,S	4.8155911E-02	3.47301414E-02	8.2886174E-02
VAR L0 A,C,S	6.1713938E-02	4.35878265E-02	8.1557019
VAR H3 A,C,S	3.20663100E-02	4.6651915E-02	7.2668247E-02

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VAR S9 X,Y,S 2.3548666E-02 3.9197755E-02 6.2746592E-02
VAR H1 X,Y,S 3.2959737E-02 3.7705429E-02 7.0665854E-02
VAR L1 X,Y,S 3.5060722E-02 3.6275730E-02 7.1336471E-02
VAR H0 X,Y,S 3.47330148E-02 4.81559886E-02 8.2888174E-02
VAR L0 X,Y,S 4.3987826E-02 6.1713934E-02 8.1057819E-02
VAR H3 X,Y,S 4.6605591E-02 3.2663100E-02 7.268247E-02

```

## CASE41.DAT

## CASE42.DAT

PARAMETERS				
1.269999	3.7599991E-02	64.00000	56789.00	
69.00000	11.0000000E+00	99.00000	64.00000	56789.00
53.20000	187.00000	53.20000	9.00000	64.00000
1.337000	1.0230000E-09	99999.00	99999.00	9.00000

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH,FACTOR	6.45000024	5.75392625E-03	96.88653
L1(θ,θ), HIGH,FACTOR	9.4721721	5.9106438E-03	93.97504
H2(θ,θ), HIGH,FACTOR	6.5559694	4.5894321E-03	121.2665
L2(θ,θ), HIGH,FACTOR	6.6825038	3.6556339E-03	151.9444
H3(θ,θ), HIGH,FACTOR	9.4597161	1.4531134E-02	38.22503

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9197750E-02	2.3548668E-02	6.2746502E-02
VAR H1 A,C,S	3.8998570E-02	5.1483624E-02	8.2352395E-02
VAR L1 A,C,S	3.8995654E-02	5.2157864E-02	8.3153382E-02
VAR H2 A,C,S	3.6297856E-02	5.6593390E-02	9.2881508E-02
VAR L2 A,C,S	4.7612212E-02	7.0089296E-02	0.1179016
VAR H3 A,C,S	3.6605757E-02	5.4495001E-02	9.1101691E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9197750E-02	2.3548668E-02	6.2746502E-02
VAR H1 A,C,S	3.0496454E-02	4.5963850E-02	7.6462389E-02
VAR L1 A,C,S	2.9871946E-02	4.5693219E-02	7.5565264E-02
VAR H2 A,C,S	3.1964581E-02	4.6871119E-02	7.8835532E-02
VAR L2 A,C,S	3.2619074E-02	4.7265410E-02	7.9884544E-02
VAR H3 A,C,S	3.6881376E-02	5.5162832E-02	9.2044286E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	6.38117943E-03	9.6851252E-03	1.6566923E-02
DIF L1 A,C,S	6.9883633E-03	9.7055520E-03	1.6693849E-02
DIF H2 A,C,S	6.7129722E-03	1.28507733E-02	1.9563788E-02
DIF L2 A,C,S	8.2504469E-03	2.9348772E-02	2.8599132E-02
DIF H3 A,C,S	8.2379738E-03	1.1345849E-02	1.9583818E-02
DIF H1 X,Y,S	9.5851252E-03	6.8817951E-03	1.6566923E-02
DIF L1 X,Y,S	9.1055528E-03	6.9883633E-03	1.6693849E-02
DIF H2 X,Y,S	1.28507733E-02	6.7129722E-03	1.9563788E-02
DIF L2 X,Y,S	2.0348772E-02	8.2504469E-03	2.8599132E-02
DIF H3 X,Y,S	1.1345849E-02	8.2379738E-03	1.9583818E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	6.7044087E-03	6.7044087E-03	6.7044087E-03
DIF L1 X,Y,S	7.616080E-03	6.7819682E-03	1.439794E-02
DIF H2 X,Y,S	6.7819682E-03	6.616080E-03	1.439794E-02
DIF L2 X,Y,S	8.4724575E-03	6.5988544E-03	8.4724575E-03
DIF H3 X,Y,S	8.6222962E-03	6.5366919E-03	1.5063323E-02
DIF H1 X,Y,S	7.9095867E-03	1.1387363E-02	1.8847823E-02
DIF L1 X,Y,S	7.616080E-03	6.7819682E-03	1.439794E-02
DIF H2 X,Y,S	6.7819682E-03	6.5988544E-03	1.5063323E-02
DIF L2 X,Y,S	8.6222962E-03	6.5366919E-03	1.5188391E-02
DIF H3 X,Y,S	8.2379738E-03	1.1387363E-02	1.8847823E-02

## CASE43.DAT

## CASE44.DAT

PARAMETERS			
1.239999	3.750000E-02	64.00000	56789.00
60.00000	6.000000E+00	135.00000	64.00000
53.28999	6.000000E+00	6.000000E+00	106.00000
1.337699	1.020000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1( $\theta$ , $\theta$ ), HIGH, FACTOR	0.9644694	6.0477504E-03	100.1577
L1( $\theta$ , $\theta$ ), HIGH, FACTOR	0.9977114	4.7524729E-03	127.4555
H0( $\theta$ , $\theta$ ), HIGH, FACTOR	1.145888	5.6885116E-03	108.0917
L0( $\theta$ , $\theta$ ), HIGH, FACTOR	1.3686579	8.6866880E-03	70.37886
H3( $\theta$ , $\theta$ ), HIGH, FACTOR	0.2634795	2.7588138E-02	21.95614

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 X,C,S	3.12515324E-02	2.3511153E-02	6.2746540E-02
VAR H1 X,C,S	0.0000000E+00	7.2344963E-02	0.1604165
VAR L1 X,C,S	0.1120533	8.4735818E-02	0.1967889
VAR H0 X,C,S	0.1165317	8.1798628E-02	0.1923299
VAR L0 X,C,S	6.999692E-02	4.2369232E-02	6.1122187
VAR H3 X,C,S	6.0336849E-02	5.6128416E-02	6.1224654

VAR S0 X,Y,S	3.1373221E-02	3.1373285E-02	6.2746540E-02
VAR H1 X,Y,S	0.163749	5.8041781E-02	0.1604165
VAR L1 X,Y,S	0.122555	7.3533140E-02	0.1967889
VAR H0 X,Y,S	0.1214180	7.0912361E-02	0.1923299
VAR L0 X,Y,S	7.4581482E-02	3.7637267E-02	6.1122187
VAR H3 X,Y,S	7.2553653E-02	4.9881533E-02	6.1224654

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,C,S	2.2659284E-02	2.07775738E-02	4.2834956E-02
DIF L1 X,C,S	3.2225952E-02	2.7256173E-02	5.9526172E-02
DIF H0 X,C,S	3.1803602E-02	2.55556639E-02	5.6562312E-02
DIF L0 X,C,S	1.1974258E-02	7.7757440E-03	1.9749997E-02
DIF H3 X,C,S	1.1888423E-02	1.3477480E-02	2.5335893E-02

DIF H1 X,Y,S	3.1510934E-02	1.1291650E-02	4.2834956E-02
DIF L1 X,Y,S	4.285111E-02	1.6651921E-02	5.9526172E-02
DIF H0 X,Y,S	4.1397352E-02	1.5164952E-02	5.6562312E-02
DIF L0 X,Y,S	1.46666886E-02	5.1033535E-03	1.9749997E-02
DIF H3 X,Y,S	1.5911007E-02	9.4248891E-03	2.5335893E-02

PARAMETERS			
1.239999	3.75000001E-02	64.00000	56789.00
60.00000	6.00000000E+00	135.00000	64.00000
53.28999	6.00000000E+00	6.000000E+00	106.00000
1.337699	1.02000000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1( $\theta$ , $\theta$ ), HIGH, FACTOR	0.9644694	6.0477504E-03	100.1577
L1( $\theta$ , $\theta$ ), HIGH, FACTOR	0.9977114	4.7524729E-03	127.4555
H0( $\theta$ , $\theta$ ), HIGH, FACTOR	1.145888	5.6885116E-03	108.0917
L0( $\theta$ , $\theta$ ), HIGH, FACTOR	1.3686579	8.6866880E-03	70.37886
H3( $\theta$ , $\theta$ ), HIGH, FACTOR	0.2634795	2.7588138E-02	21.95614

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	3.9235324E-02	2.3511153E-02	6.2746540E-02
VAR H1 A,C,S	3.6602188E-02	3.7495667E-02	7.4987166E-02
VAR L1 A,C,S	3.6888333E-02	3.7912898E-02	7.4881223E-02
VAR H0 A,C,S	3.9446899E-02	3.49944735E-02	7.3851191E-02
VAR L0 A,C,S	4.5594045E-02	3.1682167E-02	7.1676321E-02
VAR H3 A,C,S	4.2731281E-02	5.4614733E-02	9.7345948E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	3.1373221E-02	3.1373285E-02	6.2746540E-02
DIF L1 X,Y,S	2.0854776E-02	5.0332262E-02	7.4987166E-02
DIF H0 X,Y,S	2.3907822E-02	5.08993438E-02	7.4981223E-02
DIF L0 X,Y,S	2.3373680E-02	5.04778180E-02	7.3851191E-02
DIF H3 X,Y,S	2.2349650E-02	4.9336262E-02	7.1676321E-02
DIF H1 X,Y,S	4.1691132E-02	5.56554798E-02	9.7345948E-02
DIF L1 X,Y,S	5.6165154E-03	5.9375442E-03	1.1311817E-02
DIF H0 X,Y,S	5.827542E-03	5.2488730E-03	1.0276408E-02
DIF L0 X,Y,S	4.9301089E-03	4.5550293E-03	9.4486189E-03
DIF H3 X,Y,S	7.5716464E-03	9.8272707E-03	1.7398888E-02

## CASE45.DAT

## CASE46.DAT

PARAMETERS

1.200000	3.750000E-02	64.00000	56789.98
60.00000	6.000000E+00	135.00000	64.00000
53.20000	95.00000	53.20000	9.885000
1.337000	1.020000E-09	99999.98	99999.98

MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH,FACTOR	<b>4.2764742</b>	4.3552535E-03	139.7217
L1(θ,θ), HIGH,FACTOR	<b>8.3181950</b>	4.8266877E-03	125.4958
H0(θ,θ), HIGH,FACTOR	<b>0.3815361</b>	5.2463133E-03	115.4588
L0(θ,θ), HIGH,FACTOR	<b>0.5282398</b>	7.1965095E-03	84.16946
H3(θ,θ), HIGH,FACTOR	<b>0.2408874</b>	7.3930799E-03	81.93199

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	3.9235524E-02	2.3511153E-02	6.2746540E-02
VAR H1 A,C,S	5.7363757E-02	3.9168654E-02	9.5531928E-02
VAR L1 A,C,S	5.5166858E-02	3.6478655E-02	9.1644727E-02
VAR H0 A,C,S	5.9837051E-02	2.7602131E-02	8.7439194E-02
VAR L0 A,C,S	5.5034626E-02	2.0541701E-02	7.5576233E-02
VAR H3 A,C,S	6.7575462E-02	6.4217843E-02	6.1317935
VAR Sθ X,Y,S	3.1373221E-02	3.1373285E-02	6.2746540E-02
VAR H1 X,Y,S	3.4720883E-02	6.181004E-02	9.6531920E-02
VAR L1 X,Y,S	3.1966712E-02	5.9678055E-02	9.1644727E-02
VAR H0 X,Y,S	3.7687242E-02	4.9752016E-02	8.7439194E-02
VAR L0 X,Y,S	3.7388057E-02	3.8168241E-02	7.5576253E-02
VAR H3 X,Y,S	7.3791683E-02	5.8901865E-02	6.1317935

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	7.9855518E-03	6.7463833E-03	1.4731930E-02
DIF L1 A,C,S	7.6957452E-03	6.1803544E-03	1.3876104E-02
DIF H0 A,C,S	5.8496566E-03	3.3727728E-03	9.2224190E-03
DIF L0 A,C,S	3.9558657E-03	2.5750398E-03	6.5509075E-03
DIF H3 A,C,S	1.1963857E-02	1.5883598E-02	2.7849454E-02
DIF H1 X,Y,S	5.0826883E-03	9.6492562E-03	1.4731930E-02
DIF L1 X,Y,S	5.0724349E-03	8.8035633E-03	1.3876104E-02
DIF H0 X,Y,S	3.7676422E-03	5.4546194E-03	9.2224190E-03
DIF L0 X,Y,S	3.0880465E-03	3.4590012E-03	6.5309075E-03
DIF H3 X,Y,S	1.6311446E-02	1.1506323E-02	2.7849454E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	3.9235324E-02	2.3511153E-02	6.2746540E-02
VAR H1 A,C,S	6.9906557E-02	3.7103798E-02	6.1070945
VAR L1 A,C,S	6.3975126E-02	3.5636675E-02	9.9611056E-02
VAR H0 A,C,S	9.9926673E-02	7.4242972E-02	6.1732691
VAR L0 A,C,S	8.7486044E-02	8.6918227E-02	6.1740442
VAR H3 A,C,S	5.2485534E-02	3.3552527E-02	8.6038247E-02
VAR Sθ X,Y,S	3.1373221E-02	3.1373285E-02	6.2746540E-02
VAR H1 X,Y,S	6.021552E-02	4.6876889E-02	6.1070945
VAR L1 X,Y,S	5.6941785E-02	4.2670138E-02	9.2611056E-02
VAR H0 X,Y,S	9.3571775E-02	7.9697780E-02	6.1732691
VAR L0 X,Y,S	8.1031918	7.1212366E-02	6.1740442
VAR H3 X,Y,S	5.1600997E-02	3.5029218E-02	8.6038247E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	9.0039792E-03	5.2763522E-03	1.42800319E-02
DIF L1 A,C,S	7.6870466E-03	5.1715402E-03	1.2858892E-02
DIF H0 A,C,S	2.2270523E-02	1.9917550E-02	4.2188037E-02
DIF L0 A,C,S	2.0860326E-02	2.6431640E-02	4.7311906E-02
DIF H3 A,C,S	6.2509872E-03	5.4180198E-03	1.166927E-02
DIF H1 X,Y,S	9.4788177E-03	4.8015028E-03	1.42800319E-02
DIF L1 X,Y,S	8.5954555E-03	4.2331291E-03	1.2858892E-02
DIF H0 X,Y,S	2.5858266E-02	1.6329806E-02	4.2188037E-02
DIF L0 X,Y,S	3.1387044E-02	1.5924517E-02	4.7311906E-02
DIF H3 X,Y,S	6.8672239E-03	4.8018075E-03	1.166927E-02

## CASE47.DAT

## CASE48.DAT

PARAMETERS				
1.200000	3.750000E-02	64.00000	56789.00	56789.00
60.00000	0.0000000E+00	135.00000	64.00000	64.00000
53.20000	160.00000	53.20000	9.885000	9.885000
1.337000	1.0260000E-09	99999.00	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR 0.4237172 4.7737663E-03 126.8870  
 L1(θ,θ), HIGH, FACTOR 0.4469372 4.9122730E-03 123.3893  
 H2(θ,θ), HIGH, FACTOR 0.5287082 3.8294455E-03 158.1767  
 L2(θ,θ), HIGH, FACTOR 0.6570732 3.2215109E-03 188.0264  
 H3(θ,θ), HIGH, FACTOR 0.4478621 1.6117226E-02 37.58270

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S 3.9235321E-02 2.3511153E-02 6.2746540E-02  
 VAR H1 A,C,S 6.9921263E-02 5.7392698E-02 0.1272242  
 VAR L1 A,C,S 6.4745525E-02 5.3593340E-02 0.1188492  
 VAR H2 A,C,S 8.1152737E-02 6.6467784E-02 0.1476208  
 VAR L2 A,C,S 0.1029618 7.9122963E-02 0.1826849  
 VAR H3 A,C,S 5.6721538E-02 4.5701035E-02 0.1924226

VAR S0 X,Y,S 3.1373221E-02 3.1373285E-02 6.2746540E-02  
 VAR H1 X,Y,S 8.7471075E-02 3.9753021E-02 0.1272242  
 VAR L1 X,Y,S 8.1157431E-02 3.6891993E-02 0.1188492  
 VAR H2 X,Y,S 9.7956933E-02 5.6536402E-02 0.1476208  
 VAR L2 X,Y,S 0.1156321 6.7653065E-02 0.1826849  
 VAR H3 X,Y,S 6.3618645E-02 3.8893838E-02 0.1924226

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 1.4375166E-02 1.3534699E-02 2.7989793E-02  
 DIF L1 A,C,S 1.2663497E-02 1.1846535E-02 2.4510011E-02  
 DIF H2 A,C,S 1.9292468E-02 1.7999308E-02 3.7291750E-02  
 DIF L2 A,C,S 2.8421668E-02 2.4890551E-02 5.3312205E-02  
 DIF H3 A,C,S 8.7638423E-03 8.9753375E-03 1.7739195E-02

DIF H1 X,Y,S 2.1028444E-02 6.8813562E-03 2.7989793E-02  
 DIF L1 X,Y,S 1.8103335E-02 6.4897035E-03 2.4510011E-02  
 DIF H2 X,Y,S 2.8014641E-02 9.2771063E-03 3.7291750E-02  
 DIF L2 X,Y,S 3.9093109E-02 1.4309298E-02 5.3312205E-02  
 DIF H3 X,Y,S 1.1329573E-02 6.40961157E-03 1.7739195E-02

PARAMETERS				
1.200000	3.7500000E-02	64.00000	56789.00	56789.00
60.00000	0.0000000E+00	135.00000	64.00000	64.00000
53.20000	160.00000	53.20000	9.885000	9.885000
1.337000	1.0260000E-09	99999.00	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR 1.200000 3.7500000E-02 64.00000  
 L1(θ,θ), HIGH, FACTOR 60.00000 0.0000000E+00 135.00000  
 H2(θ,θ), HIGH, FACTOR 53.20000 0.0000000E+00 53.20000  
 L2(θ,θ), HIGH, FACTOR 1.337000 1.0260000E-09 99999.00

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S 3.9235324E-02 2.3511153E-02 6.2746540E-02  
 VAR H1 A,C,S 6.8277337E-02 5.1553868E-02 0.1188492  
 VAR L1 A,C,S 6.7759126E-02 5.1629056E-02 0.1187883  
 VAR H2 A,C,S 6.9843724E-02 5.2926088E-02 0.1227797  
 VAR L2 A,C,S 7.0252812E-02 5.3775244E-02 0.1246961  
 VAR H3 A,C,S 7.9254314E-02 5.7385482E-02 0.1366400

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 1.3345358E-02 1.045781E-02 2.4361162E-02  
 DIF L1 A,C,S 1.3844545E-02 1.0758821E-02 2.3823369E-02  
 DIF H2 A,C,S 1.4004633E-02 1.1593582E-02 2.55598235E-02  
 DIF L2 A,C,S 1.4382414E-02 1.1879852E-02 2.6262311E-02  
 DIF H3 A,C,S 1.5857097E-02 1.3181894E-02 2.9038986E-02

PARAMETERS	1.200000E-02	64.00000	56789.00
36.00000	6.0000000E+00	6.0000000E+00	64.00000
53.20000	9.0000000E+00	9.0000000E+00	100.00000
1.337000	1.0240000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	0.9919936	1.02226829E-02	42.07218
L1(θ,θ), HIGH, FACTOR	1.021375	8.9383568E-03	53.52649
H2(θ,θ), HIGH, FACTOR	1.1755#2	7.2849345E-03	59.71810
L2(θ,θ), HIGH, FACTOR	1.388892	3.7794111E-03	113.8445
H3(θ,θ), HIGH, FACTOR	0.9699911	3.2236971E-02	13.34694

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	2.2875596E-02	1.3741832E-02	3.6617365E-02
VAR H1 X,Y,S	2.7738085E-02	9.8730875E-03	3.7611037E-02
VAR L1 X,Y,S	3.0132117E-02	1.2010077E-02	4.2241268E-02
VAR Hθ A,C,S	3.7174772E-02	1.5638402E-02	5.2813258E-02
VAR Lθ A,C,S	8.1552714E-02	4.4886183E-02	9.1264988
VAR H3 A,C,S	2.6457888E-02	1.2447299E-02	3.8905188E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	1.2000000	3.7500000E-02	64.00000
36.00000	6.0000000E+00	6.0000000E+00	64.00000
53.20000	9.0000000E+00	9.0000000E+00	53.20000
1.337000	1.0200000E-09	99999.00	99999.00

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	2.3373240E-03	1.6385566E-03	3.97588803E-03
DIF L1 A,C,S	3.2513610E-03	1.6287981E-03	4.8891624E-03
DIF Hθ A,C,S	5.2680099E-03	1.8442333E-03	7.1123256E-03
DIF Lθ A,C,S	2.7525896E-02	1.2715299E-02	4.0241159E-02
DIF H3 A,C,S	2.5818859E-03	2.0210675E-03	4.6628825E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	2.3373240E-03	1.6385566E-03	3.7715081E-03	2.4070079E-03	6.1785025E-03
DIF L1 X,Y,S	3.2513610E-03	1.6287981E-03	3.7621958E-03	2.3792775E-03	6.1414777E-03
DIF Hθ X,Y,S	5.2680099E-03	1.8442333E-03	3.7118474E-03	2.4410449E-03	6.1528822E-03
DIF Lθ X,Y,S	2.7525896E-02	1.2715299E-02	4.0221474E-03	6.1665196E-03	6.1665196E-03
DIF H3 X,Y,S	2.5818859E-03	2.0210675E-03	4.6628825E-03	5.7223723E-03	5.7223723E-03

## CASE51.DAT

## CASE52.DAT

PARAMETERS	3.750000E-02	64.00000E-02	56789.00
H1(0,0), HIGH, FACTOR	<b>9.245829E-02</b>	<b>3.0857385E-03</b>	<b>139.4366</b>
L1(0,0), HIGH, FACTOR	<b>9.2828577</b>	<b>4.0909674E-03</b>	<b>105.1991</b>
H0(0,0), HIGH, FACTOR	<b>9.3511569</b>	<b>1.8517275E-03</b>	<b>232.3587</b>
L0(0,0), HIGH, FACTOR	<b>9.4935394</b>	<b>1.8395105E-03</b>	<b>233.9319</b>
H3(0,0), HIGH, FACTOR	<b>9.2264158</b>	<b>9.3421191E-03</b>	<b>46.05617</b>

## MAGNITUDE ARRAY SCALING FACTORS

H1(0,0), HIGH, FACTOR	9.245829E-02	3.0857385E-03	139.4366
L1(0,0), HIGH, FACTOR	9.2828577	4.0909674E-03	105.1991
H0(0,0), HIGH, FACTOR	9.3511569	1.8517275E-03	232.3587
L0(0,0), HIGH, FACTOR	9.4935394	1.8395105E-03	233.9319
H3(0,0), HIGH, FACTOR	9.2264158	9.3421191E-03	46.05617

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S 2.2875576E-02	1.3741837E-02	3.6617365E-02
VAR H1 X,Y,S 3.0781627E-02	1.6182279E-02	4.6563383E-02
VAR L1 A,C,S 2.6155148E-02	1.2529409E-02	3.8684562E-02
VAR H0 A,C,S 4.0117637E-02	1.4438586E-02	7.8556113E-02
VAR L0 A,C,S 4.0999041E-02	4.2181652E-02	9.0880809E-02
VAR H3 A,C,S 2.1318743E-02	1.1072985E-02	3.5391659E-02

## INTEGRATED VARIANCE FROM DIF SPECTRA

VAR S0 X,Y,S 2.2875596E-02	1.3741632E-02	3.6617365E-02
VAR H1 X,Y,S 3.0781627E-02	1.6182279E-02	4.6563383E-02
VAR L1 X,Y,S 2.6155148E-02	1.2529409E-02	3.8684562E-02
VAR H0 X,Y,S 4.0117637E-02	1.4438586E-02	7.8556113E-02
VAR L0 X,Y,S 4.0999041E-02	4.2181652E-02	9.0880809E-02
VAR H3 X,Y,S 2.1318743E-02	1.1072985E-02	3.5391659E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 3.10330346E-03	2.0351906E-03	5.1382259E-03
DIF L1 A,C,S 2.1644537E-03	1.7119035E-03	3.8763594E-03
DIF H0 A,C,S 8.4576504E-03	7.8865669E-03	1.6258262E-02
DIF L0 A,C,S 1.1178816E-02	1.1397973E-02	2.2496756E-02
DIF H3 A,C,S 2.0141555E-03	1.8592439E-03	3.8734048E-03

DIF H1 X,Y,S 3.10330346E-03	2.0351906E-03	5.1382259E-03
DIF L1 X,Y,S 2.1644537E-03	1.7119035E-03	3.8763594E-03
DIF H0 X,Y,S 8.4576504E-03	7.8865669E-03	1.6258262E-02
DIF L0 X,Y,S 1.1178816E-02	1.1397973E-02	2.2496756E-02
DIF H3 X,Y,S 2.0141555E-03	1.8592439E-03	3.8734048E-03

## INTEGRATED VARIANCE FROM DIF SPECTRA

VAR S0 A,C,S 2.2875596E-02	1.3741632E-02	3.6617365E-02
VAR H1 X,Y,S 2.3228347E-02	9.65622339E-03	3.2878597E-02
VAR L1 A,C,S 2.2699920E-02	8.6279567E-03	3.1318933E-02
VAR H0 A,C,S 2.473116E-02	1.2923197E-02	3.7653337E-02
VAR L0 A,C,S 2.7655602E-02	1.6622925E-02	4.4278629E-02
VAR H3 A,C,S 2.495991E-02	1.1669224E-02	3.6029186E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 2.1472595E-03	2.1110419E-03	4.258283E-03
DIF L1 A,C,S 1.8792632E-03	2.2672899E-03	4.1465331E-03
DIF H0 A,C,S 2.7653303E-03	1.7531270E-03	4.5184731E-03
DIF L0 A,C,S 3.4181611E-03	1.5189931E-03	4.5189931E-03
DIF H3 A,C,S 2.0962798E-03	1.9215163E-03	4.0177908E-03

## CASE53.DAT

## CASE54.DAT

PARAMETERS	5.7500001E-02	64.00000	56789.98
36.00000	1.9000000E+00	9.0000000E+00	64.00000
53.20000	189.0000	53.20000	9.885600
1.337000	1.0200000E-09	99999.99	99999.99

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ, θ), HIGH,FACTOR	0.4468286	8.1063444E-03	53.07756
L1(θ, θ), HIGH,FACTOR	0.4687488	8.8254223E-03	48.75291
H0(θ, θ), HIGH,FACTOR	0.55361650	6.3570538E-03	67.65368
L0(θ, θ), HIGH,FACTOR	0.6794215	5.3272257E-03	80.76718
H3(θ, θ), HIGH,FACTOR	0.4541388	1.8322222E-02	23.43323

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2875596E-02	1.3741832E-02	3.6617365E-02
VAR H1 A,C,S	2.5626143E-02	7.9436274E-03	3.3569779E-02
VAR L1 A,C,S	2.4475290E-02	7.4378434E-03	3.1912349E-02
VAR H0 A,C,S	2.8331993E-02	9.6641919E-03	3.7996173E-02
VAR L0 A,C,S	3.0994585E-02	1.1691522E-02	4.2686116E-02
VAR H3 A,C,S	2.4376977E-02	1.0333366E-02	3.4710389E-02
VAR S0 X,Y,S	2.2875596E-02	1.3741832E-02	3.6617365E-02
VAR H1 X,Y,S	2.5626143E-02	7.9436274E-03	3.3569779E-02
VAR L1 X,Y,S	2.4475290E-02	7.4378434E-03	3.1912349E-02
VAR H0 X,Y,S	2.8331993E-02	9.6641919E-03	3.7996173E-02
VAR L0 X,Y,S	3.0994585E-02	1.1691522E-02	4.2686116E-02
VAR H3 X,Y,S	2.4376977E-02	1.0333366E-02	3.4710389E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

MAGNITUDE ARRAY SCALING FACTORS			
H1(θ, θ), HIGH,FACTOR	1.419238	2.3482593E-02	18.32277
L1(θ, θ), HIGH,FACTOR	1.5000022	2.6038338E-02	16.52429
H0(θ, θ), HIGH,FACTOR	1.524574	2.1785045E-02	19.75948
L0(θ, θ), HIGH,FACTOR	1.710694	2.2621470E-02	19.02621
H3(θ, θ), HIGH,FACTOR	1.747829	4.9159654E-02	8.752462
VAR S0 A,C,S	2.2875596E-02	1.3741832E-02	3.6617365E-02
VAR H1 A,C,S	3.1561270E-02	1.1317615E-02	4.2878933E-02
VAR L1 A,C,S	2.9407065E-02	1.0278562E-02	3.9885592E-02
VAR H0 A,C,S	3.3254158E-02	1.2308263E-02	4.5562429E-02
VAR L0 A,C,S	3.2318294E-02	1.1925312E-02	4.423664E-02
VAR H3 A,C,S	2.9667298E-02	1.2695429E-02	4.2332742E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	1.5849394E-03	2.0581731E-03	3.6431155E-03
DIF L1 A,C,S	1.3911274E-03	2.2197051E-03	3.6168340E-03
DIF H0 A,C,S	2.3139163E-03	1.6631188E-03	3.9770352E-03
DIF L0 A,C,S	3.1526222E-03	1.5316515E-03	4.7227278E-03
DIF H3 A,C,S	1.9861875E-03	1.8810238E-03	3.7872111E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	2.9267534E-03	1.5024336E-03	4.4291830E-03
DIF L1 X,Y,S	2.4162212E-03	1.5926182E-03	4.0000422E-03
DIF H0 X,Y,S	3.4228893E-03	1.5926182E-03	4.0000422E-03
DIF L0 X,Y,S	3.1800981E-03	1.4710914E-03	4.8939098E-03
DIF H3 X,Y,S	2.4560019E-03	1.4761277E-03	4.6571150E-03
DIF H1 X,Y,S	2.4560019E-03	1.6565275E-03	4.1125366E-03
DIF L1 X,Y,S	2.3197051E-03	1.5926182E-03	4.0000422E-03
DIF H0 X,Y,S	3.1800981E-03	1.4761277E-03	4.6571150E-03
DIF L0 X,Y,S	3.1800981E-03	1.4761277E-03	4.6571150E-03
DIF H3 X,Y,S	2.4560019E-03	1.6565275E-03	4.1125366E-03

## CASE55.DAT

## CASE56.DAT

PARAMETERS	VALUES		
1.200000	3.750000E-02	64.00000	56789.00
36.00000	6.000000E-02	45.00000	64.00000
53.20000	1.000000E-02	8.000000E+00	100.00000
1.337000	1.0260000E-02	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ, φ), HIGH, FACTOR	1.000000
L1(θ, φ), HIGH, FACTOR	1.030000
H2(θ, φ), HIGH, FACTOR	1.191794
L2(θ, φ), HIGH, FACTOR	1.397467
H3(θ, φ), HIGH, FACTOR	0.9544371

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2876349E-02
VAR H1 A,C,S	2.4932304E-02
VAR L1 A,C,S	2.5128221E-02
VAR H2 A,C,S	2.7334739E-02
VAR L2 A,C,S	4.0377144E-02
VAR H3 A,C,S	2.5035574E-02
VAR S0 X,Y,S	1.8308684E-02
VAR H1 X,Y,S	3.0777288E-02
VAR L1 X,Y,S	3.0866221E-02
VAR H2 X,Y,S	3.2971494E-02
VAR L2 X,Y,S	4.7355749E-02
VAR H3 X,Y,S	2.9058346E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	3.5822173E-03
DIF L1 A,C,S	3.6251219E-03
DIF H2 A,C,S	3.9830925E-03
DIF L2 A,C,S	8.1229135E-03
DIF H3 A,C,S	3.5124288E-03
DIF H1 X,Y,S	3.2788040E-03
DIF L1 X,Y,S	3.5215805E-03
DIF H2 X,Y,S	4.4697393E-03
DIF L2 X,Y,S	1.1597508E-02
DIF H3 X,Y,S	3.5349955E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	2.0839428E-03
DIF L1 X,Y,S	2.9847021E-03
DIF H2 X,Y,S	3.3533848E-03
DIF L2 X,Y,S	8.1356466E-03
DIF H3 X,Y,S	3.1106551E-03
DIF H1 A,C,S	6.4661605E-03
DIF L1 A,C,S	6.6898273E-03
DIF H2 A,C,S	7.3364791E-03
DIF L2 A,C,S	1.4598572E-02
DIF H3 A,C,S	1.4472049E-02
DIF H1 X,Y,S	6.4661605E-03
DIF L1 X,Y,S	6.6898273E-03
DIF H2 X,Y,S	7.3364791E-03
DIF L2 X,Y,S	1.4598572E-02
DIF H3 X,Y,S	1.4472049E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2896349E-02
VAR H1 A,C,S	2.8118761E-02
VAR L1 A,C,S	2.88805107E-02
VAR H2 A,C,S	2.7468899E-02
VAR L2 A,C,S	2.7455218E-02
VAR H3 A,C,S	3.5680031E-02
VAR S0 X,Y,S	1.8308684E-02
VAR H1 X,Y,S	1.5327693E-02
VAR L1 X,Y,S	1.6160009E-02
VAR H2 X,Y,S	1.4515343E-02
VAR L2 X,Y,S	1.4462239E-02
VAR H3 X,Y,S	2.5007652E-02
DIF H1 A,C,S	1.7003248E-03
DIF L1 A,C,S	1.6836234E-03
DIF H2 A,C,S	1.8888653E-03
DIF L2 A,C,S	1.9113834E-03
DIF H3 A,C,S	1.5002813E-03
DIF H1 X,Y,S	2.1444261E-03
DIF L1 X,Y,S	2.0392630E-03
DIF H2 X,Y,S	2.2675071E-03
DIF L2 X,Y,S	2.2735209E-03
DIF H3 X,Y,S	2.1332924E-03

## CASE57.DAT

## CASE58.DAT

PARAMETERS	VALUES
H1(θ,θ),HIGH,FACTOR	3.7500000E-02
L1(θ,θ),HIGH,FACTOR	45.00000E-06
H0(θ,θ),HIGH,FACTOR	90.00000E-06
L0(θ,θ),HIGH,FACTOR	1.0200000E-09
H3(θ,θ),HIGH,FACTOR	1.3370000E-09
	56789.00
	64.00000
	9.885000
	99999.00
	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

	VALUES
H1(θ,θ)	3.4278177E-03
L1(θ,θ)	4.5686848E-03
H0(θ,θ)	2.2072336E-03
L0(θ,θ)	1.9812812E-03
H3(θ,θ)	8.4968824E-03
	136.8854
	102.7031
	212.5821
	236.4677
	55.22241

## INTEGRATED VARIANCE FROM POWER SPECTRA

	VALUES
VAR Sθ A,C,S	2.2893349E-02
VAR H1 A,C,S	3.6698870E-02
VAR L1 A,C,S	3.2737046E-02
VAR H0 A,C,S	4.7983419E-02
VAR L0 A,C,S	5.8331008E-02
VAR H3 A,C,S	2.7338993E-02
	1.3720096E-02
	5.6166839E-02
	4.9382728E-02
	2.9956756E-02
	7.7940151E-02
	4.0132224E-02
	9.8443732E-02
	1.7793855E-02
	4.5132779E-02
	3.6617398E-02
	2.5070749E-02
	2.2533653E-02
	2.8147370E-02
	1.8644666E-02
	2.7983220E-02
	2.0256875E-02
	3.9513933E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

	VALUES
DIF H1 A,C,S	3.4431513E-02
DIF L1 A,C,S	3.1539153E-02
DIF H0 A,C,S	4.1750338E-02
DIF L0 A,C,S	4.9793373E-02
DIF H3 A,C,S	2.9584415E-02
	2.1735352E-02
	1.7843559E-02
	3.6186155E-02
	7.7940151E-02
	4.8724598E-02
	9.8443732E-02
	1.5620386E-02
	4.5132779E-02
	3.8300684E-02
	5.6166839E-02
	4.9382728E-02
	2.9956756E-02
	7.7940151E-02
	4.0132224E-02
	9.8443732E-02
	1.7793855E-02
	4.5132779E-02

	VALUES
DIF H1 X,Y,S	4.38835063E-03
DIF L1 X,Y,S	3.0821103E-03
DIF H0 X,Y,S	3.3888897E-03
DIF L0 X,Y,S	1.2222515E-02
DIF H3 X,Y,S	3.0416066E-03
	2.1938167E-03
	6.2781228E-03
	1.9556158E-03
	5.037259E-03
	2.9356376E-03
	5.0377259E-03
	5.5826241E-03
	1.33972263E-02
	1.0074758E-02
	2.2296293E-02
	2.6883495E-03
	5.6419498E-03
	6.2781228E-03
	1.9556158E-03
	5.037259E-03
	2.9356376E-03
	5.0377259E-03
	5.5826241E-03
	1.33972263E-02
	1.0074758E-02
	2.2296293E-02
	2.6883495E-03
	5.6419498E-03
	6.2781228E-03
	1.9556158E-03
	5.037259E-03
	2.9356376E-03
	5.0377259E-03
	5.5826241E-03
	1.33972263E-02
	1.0074758E-02
	2.2296293E-02
	2.6883495E-03
	5.6419498E-03
	6.2781228E-03
	1.9556158E-03
	5.037259E-03
	2.9356376E-03
	5.0377259E-03
	5.5826241E-03
	1.33972263E-02
	1.0074758E-02
	2.2296293E-02
	2.6883495E-03
	5.6419498E-03
	6.2781228E-03
	1.9556158E-03
	5.037259E-03
	2.9356376E-03
	5.0377259E-03
	5.5826241E-03
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	5.0377259E-03
	5.5826241E-03
	1.33972263E-02
	1.0074758E-02
	2.229

PARAMETERS		PARAMETERS	
1.269999	3.759999E-02	64.00000	56789.00
36.00000	6.000000E+00	45.00000	64.00000
53.26200	18.00000	53.26000	9.000000
1.337000	1.4200000E-02	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR 0.4606020 8.3272867E-03 56.34700  
 L1(θ,θ), HIGH, FACTOR 0.4787847 9.1470275E-03 51.29736  
 H2(θ,θ), HIGH, FACTOR 0.5661303 6.8956339E-03 68.04571  
 L2(θ,θ), HIGH, FACTOR 0.6896612 6.2586539E-03 74.97113  
 H3(θ,θ), HIGH, FACTOR 0.4534822 1.6790370E-02 27.94568

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S 2.2896349E-02 1.3720999E-02 3.6617398E-02  
 VAR H1 A,C,S 2.4670066E-02 1.6607841E-02 4.1277926E-02  
 VAR L1 A,C,S 2.3779966E-02 1.6805403E-02 3.9825354E-02  
 VAR H2 A,C,S 2.5589671E-02 1.7084536E-02 4.3174207E-02  
 VAR L2 A,C,S 2.5517885E-02 1.8420999E-02 4.3936838E-02  
 VAR H3 A,C,S 2.3509864E-02 1.7276773E-02 4.0786620E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 3.7012997E-03 2.9785683E-03 6.6218679E-03  
 DIF L1 A,C,S 3.6467665E-03 2.8734591E-03 6.58922735E-03  
 DIF H2 A,C,S 3.7645189E-03 2.9906456E-03 6.7551611E-03  
 DIF L2 A,C,S 3.7919371E-03 3.0525867E-03 6.8445322E-03  
 DIF H3 A,C,S 3.2746440E-03 2.7746820E-03 6.0493373E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S 2.9996426E-03 3.6222239E-03 6.6218679E-03  
 DIF L1 X,Y,S 2.7162652E-03 3.8039565E-03 6.52022735E-03  
 DIF H2 X,Y,S 3.4732311E-03 3.2819160E-03 6.7551611E-03  
 DIF L2 X,Y,S 3.7213966E-03 3.1231460E-03 6.8445322E-03  
 DIF H3 X,Y,S 2.9117456E-03 3.1375880E-03 6.0493373E-03

DIF H1 X,Y,S 4.3482827E-03 3.6222239E-03 6.6218679E-03  
 DIF L1 X,Y,S 3.8179136E-03 3.0788708E-03 6.8967813E-03  
 DIF H2 X,Y,S 4.6653212E-03 2.8339245E-03 7.4992460E-03  
 DIF L2 X,Y,S 4.2864555E-03 2.8968912E-03 7.18333365E-03  
 DIF H3 X,Y,S 4.0759719E-03 2.9332972E-03 7.9892641E-03

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S 2.2896349E-02 1.3720999E-02 3.6617398E-02  
 VAR H1 A,C,S 2.7537160E-02 1.9721597E-02 4.7558761E-02  
 VAR L1 A,C,S 2.6192579E-02 1.8081724E-02 4.4874277E-02  
 VAR H2 A,C,S 2.8228475E-02 2.04097911E-02 4.8636351E-02  
 VAR L2 A,C,S 2.7191071E-02 1.9745568E-02 4.6973586E-02  
 VAR H3 A,C,S 2.7756875E-02 2.0940965E-02 4.8697911E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S 3.9982699E-03 3.2676957E-03 7.2659547E-03  
 DIF L1 A,C,S 3.8213280E-03 3.0754611E-03 6.8967873E-03  
 DIF H2 A,C,S 4.1017981E-03 3.3975362E-03 7.4992460E-03  
 DIF L2 A,C,S 3.939324E-03 3.2440133E-03 7.18333365E-03  
 DIF H3 A,C,S 3.6100477E-03 3.3992219E-03 7.9892641E-03

## CASE61.DAT

## CASE62.DAT

PARAMETERS	
1.286999	3.7500001E-02
36.00000	6.0000000E+00
53.20000	9.0000000E+00
1.337000	1.0240000E-09

PARAMETERS	
1.200000	1.2000000E-02
36.00000	6.0000000E+00
53.20000	9.0000000E+00
1.337000	1.0240000E-09

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	1.033541	1.3824857E-02	31.12261
L1(θ,θ), HIGH, FACTOR	1.049619	1.3159366E-02	32.69648
Hθ(θ,θ), HIGH, FACTOR	1.217480	1.1826988E-02	36.38272
Lθ(θ,θ), HIGH, FACTOR	1.416898	9.1755698E-03	46.89246
H3(θ,θ), HIGH, FACTOR	0.9452016	2.5605777E-02	16.88343

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR Sθ A,C,S	2.28756699E-62	1.3741827E-02	3.6617476E-02
VAR H1 A,C,S	1.5198626E-62	2.3858388E-02	3.7956923E-02
VAR L1 A,C,S	1.2982123E-62	2.3298597E-02	3.6280677E-02
VAR Hθ A,C,S	1.3624724E-62	2.3938668E-02	3.7633376E-02
VAR Lθ A,C,S	1.1547533E-62	2.3945739E-02	3.8893045E-02
VAR H3 A,C,S	1.5130291E-62	2.6785214E-02	4.1915528E-02
VAR Sθ X,Y,S	1.3741827E-02	2.28756699E-02	3.6617476E-02
VAR H1 X,Y,S	2.3858388E-02	1.3198625E-02	3.76569235E-02
VAR L1 X,Y,S	2.3298597E-02	1.2982123E-02	3.6280677E-02
VAR Hθ X,Y,S	2.3938668E-02	1.3624724E-02	3.7633376E-02
VAR Lθ X,Y,S	2.3945739E-02	1.4547331E-02	3.8893045E-02
VAR H3 X,Y,S	2.6785214E-02	1.5130291E-02	4.1915528E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

H1(θ,θ), HIGH, FACTOR	0.7009528	2.0538900E-02	28.97942
L1(θ,θ), HIGH, FACTOR	0.7583873	2.1795165E-02	19.74131
Hθ(θ,θ), HIGH, FACTOR	0.8153951	1.9791866E-02	21.73949
Lθ(θ,θ), HIGH, FACTOR	0.9692719	2.0358259E-02	21.13466
H3(θ,θ), HIGH, FACTOR	0.6416683	2.2322357E-02	19.27587

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	5.0728116E-03	2.958076E-03	7.9786275E-03
DIF L1 A,C,S	5.01615873E-03	2.8238779E-03	7.8854589E-03
DIF Hθ A,C,S	4.7870567E-03	3.0402613E-03	7.8273862E-03
DIF Lθ A,C,S	4.3403627E-03	3.3905051E-03	7.7398728E-03
DIF H3 A,C,S	4.044138E-03	4.2536971E-03	9.8581011E-03
DIF H1 X,Y,S	2.0538900E-03	5.0728116E-03	7.9786275E-03
DIF L1 X,Y,S	2.0238779E-03	5.9615873E-03	7.8854589E-03
DIF Hθ X,Y,S	2.0402613E-03	4.7870567E-03	7.8273862E-03
DIF Lθ X,Y,S	2.0390505E-03	4.3403627E-03	7.7398728E-03
DIF H3 X,Y,S	2.044138E-03	4.2536971E-03	9.8581011E-03
DIF H1 X,Y,S	2.3741212E-03	2.1158669E-03	4.4899923E-03
DIF L1 X,Y,S	2.3937355E-03	2.3741212E-03	4.7989669E-03
DIF Hθ X,Y,S	2.3122445E-03	1.90681132E-03	4.2139612E-03
DIF Lθ X,Y,S	2.3107405E-03	1.9418765E-03	4.2526186E-03
DIF H3 X,Y,S	2.6743594E-03	3.6627159E-03	6.2770671E-03

## CASE63.DAT

## CASE64.DAT

PARAMETERS	3.7500001E-02	64.00000	56789.00
H1(0,0), HIGH, FACTOR	0.2716419	3.5562778E-03	126.9875
L1(0,0), HIGH, FACTOR	0.3030993	4.6642888E-03	92.24665
H2(0,0), HIGH, FACTOR	0.3770842	2.8510308E-03	150.9153
L2(0,0), HIGH, FACTOR	0.5139449	2.9183644E-03	147.4566
H3(0,0), HIGH, FACTOR	0.2396687	7.0414296E-03	11.8478

## MAGNITUDE ARRAY SCALING FACTORS

H1(0,0), HIGH, FACTOR	2.2875609E-02	1.3741827E-02	3.6617476E-02
VAR H1 A,C,S	2.9145464E-02	2.9051362E-02	5.8196776E-02
VAR L1 A,C,S	2.3997868E-02	2.7719352E-02	5.1517271E-02
VAR H2 A,C,S	3.5384227E-02	2.6496310E-02	6.1790522E-02
VAR L2 A,C,S	3.8975518E-02	2.7065914E-02	6.5111156E-02
VAR H3 A,C,S	1.8681361E-02	2.6974686E-02	4.5111144E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2875609E-02	1.3741827E-02	3.6617476E-02
VAR H1 X,Y,S	2.9145464E-02	2.9051362E-02	5.8196776E-02
VAR L1 X,Y,S	2.7719352E-02	5.1517271E-02	5.1517271E-02
VAR H2 X,Y,S	2.6496310E-02	6.1790522E-02	6.1790522E-02
VAR L2 X,Y,S	2.7065914E-02	6.5111156E-02	6.5111156E-02
VAR H3 X,Y,S	2.6974686E-02	4.5111144E-02	4.5111144E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	4.8567243E-03	4.8957509E-03	8.9524612E-03
DIF L1 A,C,S	3.4942393E-03	4.3109371E-03	7.8951668E-03
DIF H2 A,C,S	5.2267285E-03	4.2165019E-03	9.4432374E-03
DIF L2 A,C,S	5.7947575E-03	4.4651241E-03	1.0257389E-02
DIF H3 A,C,S	3.7915939E-03	4.1655595E-03	7.8971507E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	1.3741827E-02	3.6617476E-02	3.6617476E-02
DIF H1 A,C,S	5.2692886E-03	2.3126544E-03	7.5819413E-03
DIF L1 A,C,S	5.5762194E-03	2.1723052E-03	7.7485209E-03
DIF H2 A,C,S	4.8892768E-03	2.4411457E-03	7.2414191E-03
DIF L2 A,C,S	4.7918153E-03	2.3999628E-03	7.1827699E-03
DIF H3 A,C,S	5.2472735E-03	2.7843381E-03	8.0316206E-03
DIF H1 X,Y,S	2.3126542E-03	5.2692886E-03	7.5819413E-03
DIF L1 X,Y,S	2.1723052E-03	5.5762194E-03	7.7485209E-03
DIF H2 X,Y,S	2.4411457E-03	4.8892768E-03	7.2414191E-03
DIF L2 X,Y,S	2.3999628E-03	4.7918153E-03	7.1827699E-03
DIF H3 X,Y,S	2.7843381E-03	5.2472735E-03	8.0316206E-03
DIF H1 X,Y,S	4.3957509E-03	4.8567243E-03	8.9524612E-03
DIF L1 X,Y,S	4.3109371E-03	3.4942393E-03	7.8951668E-03
DIF H2 X,Y,S	4.2165019E-03	5.2267285E-03	9.4432374E-03
DIF L2 X,Y,S	4.4651241E-03	5.7947575E-03	1.0257389E-02
DIF H3 X,Y,S	4.1655595E-03	3.7915939E-03	7.8971507E-02

## CASE65.DAT

## CASE66.DAT

PARAMETERS	PARAMETERS	PARAMETERS	PARAMETERS
1.246999	3.7500000E-02	64.000000	56789.00
36.000000	0.0000000E+00	98.000000	64.000000
53.24999	180.00000	53.230000	9.8850000
1.337000	1.0200000E-09	99999.00	99999.00
<b>MAGNITUDE ARRAY SCALING FACTORS</b>			
H1(θ,θ), HIGH, FACTOR	0.4768368	8.3096810E-03	51.77876
L1(θ,θ), HIGH, FACTOR	0.4378539	9.2170909E-03	46.68122
H2(θ,θ), HIGH, FACTOR	0.3822732	7.1989958E-03	59.70820
L2(θ,θ), HIGH, FACTOR	0.7839387	6.99352589E-03	61.50809
H3(θ,θ), HIGH, FACTOR	0.4598535	1.3799264E-02	31.16029
<b>INTEGRATED VARIANCE FROM POWER SPECTRA</b>			
VAR Sθ A,C,S	2.2875569E-02	1.3741827E-02	3.6617476E-02
VAR H1 A,C,S	1.3674027E-02	2.59366678E-02	3.9588887E-02
VAR L1 A,C,S	1.32336524E-02	2.52800034E-02	3.8516447E-02
VAR H2 A,C,S	1.3916882E-02	2.5928233E-02	3.9844947E-02
VAR L2 A,C,S	1.3597312E-02	2.5103339E-02	3.8709443E-02
VAR H3 A,C,S	1.4384513E-02	2.6577163E-02	4.0961765E-02
<b>INTEGRATED VARIANCE FROM POWER SPECTRA</b>			
VAR Sθ X,Y,S	1.3741827E-02	2.2875569E-02	3.6617476E-02
VAR H1 X,Y,S	2.59066682E-02	1.3674027E-02	3.9588887E-02
VAR L1 X,Y,S	2.52800034E-02	1.32336524E-02	3.8516447E-02
VAR .θ X,Y,S	2.5978123E-02	1.3916882E-02	3.9844947E-02
VAR Lθ X,Y,S	2.5133139E-02	1.3597312E-02	3.8709443E-02
VAR H3 X,Y,S	2.6577163E-02	1.4384513E-02	4.0961765E-02
<b>INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA</b>			
DIF H1 A,C,S	5.4172128E-03	3.3877317E-03	8.0049350E-03
DIF L1 A,C,S	5.5912733E-03	3.1814300E-03	8.0049350E-03
DIF H2 A,C,S	5.177919E-03	3.4612788E-03	8.7722048E-03
DIF L2 A,C,S	5.1417882E-03	3.273394E-03	8.4150197E-03
DIF H3 A,C,S	5.2094539E-03	3.98973074E-03	9.1989497E-03
DIF H1 X,Y,S	3.3877320E-03	5.4172128E-03	8.0049350E-03
DIF L1 X,Y,S	3.1814300E-03	5.5912733E-03	8.7722048E-03
DIF H2 X,Y,S	3.4682788E-03	5.177919E-03	8.6465339E-03
DIF L2 X,Y,S	3.2732943E-03	5.1417882E-03	8.4150197E-03
DIF H3 X,Y,S	3.9895074E-03	5.2094539E-03	9.1989497E-03
<b>INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA</b>			
DIF H1 A,C,S	4.4486602E-03	4.4486602E-03	4.4486602E-03
DIF L1 A,C,S	4.6931454E-03	4.6931454E-03	4.6931454E-03
DIF H2 A,C,S	4.3531740E-03	4.7344123E-03	4.7344123E-03
DIF L2 A,C,S	4.5077540E-03	4.3388284E-03	4.3388284E-03
DIF H3 A,C,S	4.1983472E-03	3.3415579E-03	7.5397133E-03
DIF H1 X,Y,S	4.5673372E-03	4.4486602E-03	4.4486602E-03
DIF L1 X,Y,S	4.1230894E-03	4.6931454E-03	4.6931454E-03
DIF H2 X,Y,S	4.7344123E-03	4.3531740E-03	4.3531740E-03
DIF L2 X,Y,S	4.3388284E-03	4.5877540E-03	4.5877540E-03
DIF H3 X,Y,S	3.3415579E-03	4.1989497E-03	4.1989497E-03

## CASE67.DAT

## CASE68.DAT

PARAMETERS	VALUES	PARAMETERS	VALUES
1.288888	5.7588881E-02	64.000000	54789.00
36.89888	1.0000000E+00	135.000000	64.000000
53.24000	8.0000000E+00	9.0000000E+00	100.000000
1.337902	1.0200000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(0,3),HIGH,FACTOR	1.0599322	1.2579173E-02	37.30120
L1(0,0),HIGH,FACTOR	1.038421	1.1376479E-02	41.24460
H0(0,0),HIGH,FACTOR	1.192787	9.8751177E-03	47.51377
L0(0,0),HIGH,FACTOR	1.397352	6.1078724E-03	76.82189
:3(0,0),HIGH,FACTOR	0.9466138	3.0066632E-02	15.69906

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2896349E-02	1.3720991E-02	3.6617409E-02
VAR H1 A,C,S	2.2270119E-02	1.6937997E-02	3.9208118E-02
VAR L1 A,C,S	2.1884004E-02	1.6655752E-02	3.8539786E-02
VAR H0 A,C,S	2.4900779E-02	1.9618975E-02	4.4519123E-02
VAR L0 A,C,S	3.5643976E-02	2.9418312E-02	6.5062165E-02
VAR H3 A,C,S	2.4357142E-02	1.8293021E-02	4.3130141E-02
VAR S0 X,Y,S	1.8308669E-02	1.83086718E-02	3.6617409E-02
VAR H1 X,Y,S	2.7911861E-02	1.1216183E-02	3.9208118E-02
VAR L1 X,Y,S	2.72120336E-02	1.1326938E-02	3.8539786E-02
VAR H0 X,Y,S	3.4953338E-02	1.5555579E-02	4.4519123E-02
VAR L0 X,Y,S	4.197153E-02	2.39063678E-02	6.5062165E-02
VAR H3 X,Y,S	2.8492155E-02	1.46338031E-02	4.3130141E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	3.5608774E-03	2.6737752E-03	6.2346621E-03
DIF L1 A,C,S	3.4625535E-03	2.6554539E-03	6.1179668E-03
DIF H0 A,C,S	3.8525991E-03	3.1431833E-03	6.9957866E-03
DIF L0 A,C,S	6.5320776E-03	6.4665019E-03	1.2978511E-02
DIF H3 A,C,S	3.5736639E-03	3.1849579E-03	6.7585637E-03
DIF H1 X,Y,S	2.88338950E-03	3.3515764E-03	6.2346621E-03
DIF L1 X,Y,S	2.8458915E-03	3.2720229E-03	6.1179668E-03
DIF H0 X,Y,S	4.0688887E-03	2.9869893E-03	6.9957866E-03
DIF L0 X,Y,S	9.0789441E-03	3.9195726E-03	1.2998511E-02
DIF H3 X,Y,S	3.5822843E-03	3.1762756E-03	6.7585637E-03

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2896348E-02	1.3720991E-02	3.6617409E-02
VAR H1 A,C,S	1.9237043E-02	2.8393969E-02	4.7509985E-02
VAR L1 A,C,S	1.8541198E-02	2.9075243E-02	4.7616467E-02
VAR H0 A,C,S	2.0772345E-02	2.7617762E-02	4.8399131E-02
VAR L0 A,C,S	2.1250617E-02	2.7648635E-02	4.8899274E-02
VAR H3 A,C,S	2.7202634E-02	4.1428769E-02	6.8631262E-02
VAR S0 X,Y,S	1.8308669E-02	1.8308718E-02	3.6617409E-02
VAR H1 X,Y,S	1.7395154E-02	3.045865E-02	4.7548985E-02
VAR L1 X,Y,S	1.8382244E-02	2.9234193E-02	4.7616467E-02
VAR H0 X,Y,S	1.6419856E-02	3.1970184E-02	4.8399131E-02
VAR L0 X,Y,S	1.6219653E-02	3.2679614E-02	4.8899274E-02
VAR H3 X,Y,S	3.4012966E-02	3.4618289E-02	6.8631262E-02

DIF H1 A,C,S	4.7903918E-03	4.4349264E-03	9.2235994E-03
DIF L1 A,C,S	4.9364921E-03	4.545715E-03	9.4853559E-03
DIF H0 A,C,S	4.6613873E-03	4.4657434E-03	9.8671265E-03
DIF L0 A,C,S	4.7856172E-03	4.4680686E-03	9.1656856E-03
DIF H3 A,C,S	5.4739285E-03	9.6529689E-03	1.5126849E-02
DIF H1 X,Y,S	4.1626957E-03	5.0626225E-03	9.2235994E-03
DIF L1 X,Y,S	4.3743184E-03	5.1553268E-03	9.4853559E-03
DIF H0 X,Y,S	3.9418144E-03	5.1233154E-03	9.9671265E-03
DIF L0 X,Y,S	3.9191355E-03	5.2465596E-03	9.1656856E-03
DIF H3 X,Y,S	7.3268284E-03	7.799984E-03	1.5126849E-02

PARAMETERS

1.200000E+00	3.750000E-02	64.00000	56709.00
36.00000	11.00000E+00	135.0000	64.00000
53.20000	99.00000	53.20000	9.885000
1.337000	1.020000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ, φ), HIGH,FACTOR	0.2618286	3.0775403E-03	152.4654
L1(θ, φ), HIGH,FACTOR	0.296994	4.2020525E-03	111.6641
H2(θ, φ), HIGH,FACTOR	0.3671474	2.3882277E-03	196.4713
L2(θ, φ), HIGH,FACTOR	0.5076298	2.5744168E-03	182.2628
H3(θ, φ), HIGH,FACTOR	0.2279700	8.0634207E-03	58.19698

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2896318E-02	1.3720991E-02	3.6617409E-02
VAR H1 A,C,S	2.8317876E-02	3.7324071E-02	6.5655912E-02
VAR L1 A,C,S	2.2365561E-02	2.8782798E-02	5.1168374E-02
VAR H2 A,C,S	3.6413241E-02	4.0560913E-02	7.6914117E-02
VAR L2 A,C,S	3.9827824E-02	3.6340915E-02	7.6168917E-02
VAR H3 A,C,S	2.38355922E-02	2.0562879E-02	4.4398893E-02

## INTEGRATED VARIANCE FROM DIF SPECTRA

DIF H1 A,C,S	4.8439694E-03	7.8908669E-02	1.8398710E-02
DIF L1 A,C,S	3.5474587E-02	3.6221360E-02	6.5655912E-02
DIF H2 A,C,S	2.9836617E-02	2.1311663E-02	5.1148374E-02
DIF L2 A,C,S	3.2718193E-02	4.4196033E-02	7.6914117E-02
DIF H3 A,C,S	2.8946398E-02	1.5458345E-02	4.4398893E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2896338E-02	1.3720991E-02	3.6617409E-02
VAR H1 A,C,S	2.4100667E-02	9.5773051E-03	3.3677969E-02
VAR L1 A,C,S	2.2515753E-02	9.4384765E-03	3.1956017E-02
VAR H2 A,C,S	2.6060198E-02	1.0337497E-02	3.6397662E-02
VAR L2 A,C,S	2.6063361E-02	1.1178694E-02	3.7245091E-02
VAR H3 A,C,S	2.3449168E-02	1.2349859E-02	3.5799056E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	1.8308669E-02	1.8308669E-02	1.8308669E-02
DIF L1 X,Y,S	2.1522341E-02	1.2155637E-02	3.3677969E-02
DIF H2 X,Y,S	2.1087847E-02	1.0868939E-02	3.1956017E-02
DIF L2 X,Y,S	2.1889100E-02	1.4508486E-02	3.6397662E-02
DIF H3 X,Y,S	2.1840112E-02	1.5455066E-02	3.7245091E-02
VAR H3 X,Y,S	2.2961654E-02	1.2837982E-02	3.5799056E-02
DIF H1 X,Y,S	4.8439694E-03	7.8908669E-02	1.2742856E-02
DIF L1 X,Y,S	3.8871660E-03	4.5577637E-03	8.5162157E-03
DIF H2 X,Y,S	6.3451394E-03	9.9539766E-03	1.6299114E-02
DIF L2 X,Y,S	6.9748438E-03	8.6326208E-03	1.5307472E-02
DIF H3 X,Y,S	3.5549233E-03	3.2773714E-03	6.9322861E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	1.3030567E-03	2.2827483E-03	3.9332751E-03
DIF L1 X,Y,S	1.1757435E-03	3.0554512E-03	4.2311889E-03
DIF H2 X,Y,S	1.560321E-03	1.9771134E-03	3.5375110E-03
DIF L2 X,Y,S	1.7011950E-03	1.770751E-03	3.4721713E-03
DIF H3 X,Y,S	1.8865734E-03	2.6803578E-03	4.5669228E-03

## CASE71.DAT

## CASE72.DAT

PARAMETERS		PARAMETERS	
1.200000	3.750000E-02	64.00000	56789.00
36.00000	0.00000000E+00	135.00000	64.00000
53.20000	188.00000	53.20000	9.885000
1.337000	1.02000000E-09	99999.00	99999.00

## MAGNITUDE ARRAY SCALING FACTORS

H1(θ,θ), HIGH, FACTOR	0.4618788	8.4568635E-03	55.48373
L1(θ,θ), HIGH, FACTOR	0.48000421	9.3807466E-03	50.01929
H2(θ,θ), HIGH, FACTOR	0.5671896	6.9197142E-03	67.98891
L2(θ,θ), HIGH, FACTOR	0.6906798	6.3975949E-03	74.38942
H3(θ,θ), HIGH, FACTOR	0.4499777	1.6611675E-02	28.24639

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 A,C,S	2.2896348E-02	1.3720991E-02	3.6617409E-02
VAR H1 A,C,S	2.3192881E-02	1.6888340E-02	4.0061269E-02
VAR L1 A,C,S	2.2891735E-02	1.5953334E-02	3.8045823E-02
VAR H2 A,C,S	2.4590764E-02	1.8404999E-02	4.2995792E-02
VAR L2 A,C,S	2.4889313E-02	1.8936923E-02	4.3622497E-02
VAR H3 A,C,S	2.3630153E-02	1.6822167E-02	4.0452335E-02

## INTEGRATED VARIANCE FROM POWER SPECTRA

VAR S0 X,Y,S	1.83088669E-02	1.8308718E-02	3.6617409E-02
VAR H1 X,Y,S	2.9178008E-02	1.0083181E-02	4.0061269E-02
VAR L1 X,Y,S	2.7625229E-02	1.0219757E-02	3.8045823E-02
VAR H2 X,Y,S	3.0888277E-02	1.2897485E-02	4.2995792E-02
VAR L2 X,Y,S	3.0666077E-02	1.2669165E-02	4.3622497E-02
VAR H3 X,Y,S	2.73822283E-02	1.3870044E-02	4.0452335E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	3.7582247E-03	2.8456589E-03	6.6038906E-03
DIF L1 A,C,S	3.6714688E-03	2.7342158E-03	6.4856213E-03
DIF H2 A,C,S	3.9164853E-03	3.0353272E-03	6.9518303E-03
DIF L2 A,C,S	2.9198939E-03	3.0880728E-03	7.0039663E-03
DIF H3 A,C,S	3.3268879E-03	2.8100666E-03	6.1359473E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	2.8968644E-03	3.7670148E-03	6.6038906E-03
DIF L1 X,Y,S	2.5124999E-03	3.8931221E-03	6.4956213E-03
DIF H2 X,Y,S	3.5682137E-03	3.3836088E-03	6.9518303E-03
DIF L2 X,Y,S	3.7646889E-03	3.2392820E-03	7.0039663E-03
DIF H3 X,Y,S	2.9566681E-03	3.1882794E-03	6.1359473E-03

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 A,C,S	5.442944E-03	4.3215174E-03	9.7644553E-03
DIF L1 A,C,S	4.898834E-03	3.7930399E-03	5.3902365E-02
DIF H2 A,C,S	5.8019441E-03	4.1672667E-02	1.8507935E-02
DIF L2 A,C,S	3.9966891E-02	1.7533956E-02	5.7500160E-02
DIF H3 A,C,S	5.7735159E-03	4.6678586E-03	1.0441364E-02

## INTEGRATED SQUARED DEVIATION FROM DIF SPECTRA

DIF H1 X,Y,S	6.6231498E-03	3.1413157E-03	9.7644553E-03
DIF L1 X,Y,S	5.6312272E-03	3.1484521E-03	8.7797344E-03
DIF H2 X,Y,S	7.2277886E-03	3.1777681E-03	1.0405492E-02
DIF L2 X,Y,S	6.5191011E-03	3.1376509E-03	9.6567500E-03
DIF H3 X,Y,S	7.2094752E-03	3.2318889E-03	1.0441364E-02

### 6.3 Appendix III - FORTRAN 77 Code

The input data files, FORTRAN programs, and FORTRAN subroutines that were created in support of this study are listed as follows:

#### Input Data Files

PARM.DAT  
V\_DATA.DAT

#### FORTRAN Programs

SURF.FOR  
LIGHTBLUE.FOR  
LIGHT.FOR  
ANALYS.FOR  
PSVARY.FOR

#### FORTRAN Subroutines

ALPHA.FOR  
ALPHAS.FOR  
CHECKERBOARD.FOR  
DARKNESS.FOR  
DEEPBLUE.FOR  
DIST\_GEN.FOR  
FFT2D.FOR  
PLT2DFILE.FOR  
PLT2DOUTPUT.FOR  
PLT3DFILE.FOR  
PLT3DINPUT.FOR  
PLT3DOUTPUT.FOR  
PSVARY1.FOR  
PSVARY2.FOR  
RALPH.FOR  
RIMP.FOR  
SEA\_SPEC.FOR  
SKYBLUE.FOR  
SKYMAP.FOR  
SPREADVARY.FOR  
WGN\_SPEC.FOR  
WGN2D.FOR  
WIND.FOR

Input Data Files

TYPE PARM.DAT

4	4		
1.20	0.0375	64.0	56789.0
60.0	0.0	135.0	64.0
53.2	0.0	53.2	9.885
1.337	1.02E-9	99999.0	99999.0

\*

*****	PARM.DAT	PARAMETERS	*****
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*	PARM_ROWS	PARM_COLS	
*	K_NYQ	DELTA_K	N1(SURF_ARRAY) ISEED(RANDOMIZER)
*	WIND_V(Z)	Z (CM)	WIND_AZ N(DOME_ARRAY)
*	THETA(SENSOR)	PHIO(SUN)	THETAO(SUN) LREF(LSKY(0,0))
*	INDEX_REAL	INDEX_IMAG	UNUSED UNUSED

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*****	*****	*****	*****
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*	NOTES: LREF = 100.0 FOR THETAO = 0.0 DEG		
*	LREF = 32.65 FOR THETAO = 26.6 DEG		
*	LREF = 9.885 FOR THETAO = 53.2 DEG		

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*****	*****	*****	*****
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## TYPE V\_DATA.DAT

9.3570001E-02	5.000000	98.63132	116.0000	118.7493 124.0000
7.1240805E-02	6.000000	122.4473	143.0000	146.5888 153.0000
2.8380003E-02	10.00000	227.0883	262.0000	267.3242 278.0000
1.8863197E-02	12.00000	284.7601	326.0000	333.0432 346.0000
1.0929998E-02	15.00000	376.4139	428.0000	436.7679 453.0000
7.5672008E-03	18.00000	468.2427	530.0000	540.6674 561.0000
7.0200004E-03	20.00000	524.0226	593.0000	604.4945 627.0000
8.8527985E-03	24.00000	614.9089	698.0000	711.4752 738.0000
9.8099969E-03	25.00000	634.1133	721.0000	734.7032 763.0000
1.7019998E-02	30.00000	719.6121	824.0000	840.3198 874.0000
2.7672853E-02	35.00000	797.0169	918.0000	937.8427 977.0000
3.0168798E-02	36.00000	812.0167	937.0000	956.8661 997.0000
4.1279998E-02	40.00000	870.8839	1010.000	1031.828 1075.000
4.7484916E-02	42.00000	899.7244	1045.000	1068.715 1115.000
5.7569999E-02	45.00000	942.3242	1098.000	1123.386 1173.000
6.8561204E-02	48.00000	984.1787	1151.000	1177.311 1231.000
7.6380000E-02	50.00000	1011.687	1185.000	1212.867 1269.000
9.3171470E-02	54.00000	1065.795	1253.000	1283.069 1343.000
9.7606353E-02	55.00000	1079.138	1270.000	1300.435 1362.000
0.1211800	60.00000	1144.791	1353.000	1386.207 1453.000
0.1470531	65.00000	1208.744	1434.000	1470.277 1543.000
0.1525004	66.00000	1221.338	1450.000	1486.895 1562.000
0.1751914	70.00000	1271.084	1514.000	1552.736 1631.000
0.1870752	72.00000	1295.587	1545.000	1585.286 1665.000
0.2055700	75.00000	1331.893	1592.000	1633.663 1717.000
0.2248644	78.00000	1367.675	1638.000	1681.516 1768.000
0.2381700	80.00000	1391.246	1669.000	1713.134 1802.000
0.2658397	84.00000	1437.728	1729.000	1775.710 1869.000
0.2729771	85.00000	1449.214	1744.000	1791.219 1886.000
0.3099800	90.00000	1505.859	1818.000	1867.983 1968.000
0.3491700	95.00000	1561.244	1891.000	1943.485 2049.000
0.3572698	96.00000	1572.174	1905.000	1958.439 2065.000
0.3905400	100.0000	1615.421	1962.000	2017.781 2129.000
0.4076971	102.0000	1636.766	1990.000	2047.173 2161.006
0.4340843	105.0000	1668.444	2032.000	2090.922 2208.000
0.4612525	108.0000	1699.723	2074.000	2134.271 2254.000
0.4797982	110.0000	1720.359	2101.000	2162.954 2285.000
0.5179287	114.0000	1761.123	2156.000	2219.813 2347.000
0.5276778	115.0000	1771.210	2170.000	2233.923 2362.000
0.5777199	120.0000	1821.038	2237.000	2303.869 2437.000
0.6299220	125.0000	1869.881	2303.000	2372.831 2512.000
0.6406213	126.0000	1879.535	2315.000	2386.508 2527.000
0.6842815	130.0000	1917.775	2368.000	2440.843 2585.000
0.7066289	132.0000	1936.674	2354.000	2467.789 2615.000
0.7407966	135.0000	1964.753	2433.000	2507.938 2658.000
0.7757397	139.0000	1992.513	2471.000	2547.769 2701.000

FORTRAN Programs

PROGRAM SURF

```

*
* CREATE SYNTHETIC WIND-DRIVEN WATERWAVE SURFACE OF TWO SLOPE-COMPONENT
* (X & Y) ARRAYS (+ OPTIONAL ELEVATION ARRAY)
* CONVERT SLOPE-COMPONENT ARRAYS INTO MAXIMUM-SLOPE & SLOPE-AZIMUTH ARRAY
*
* INPUTS:
* WIND_V = WIND_VELOCITY [CM/SEC] { > 0 }
* Z = HEIGHT OF WIND_V MEASUREMENT [CM] { > 0 }
* WIND_AZ = WIND_AZIMUTH [DEG] [-180..+180]
* K_NYQ = MAX_WAVENUMBER [CY/CM] { > 0 }
* DELTA_K = DELTA_WAVENUMBER [CY/CM] { > 0 }
*
IMPLICIT NONE
INTEGER I,I1,I2,J,J1,J2,N,N1,N2,N3,P,P1,P2,Q,Q1,Q2
PARAMETER(N=64,P=-N/2,Q=N/2-1)
PARAMETER(N1=64,P1=-N1/2,Q1=N1/2-1)
INTEGER*4 ISEED
INTEGER KCOR(P:Q,P:Q)
REAL ANGO,ANG1,ANG2,ANG3,BIAS1,BIAS2,DELTA_K,K_NYQ,PI,SCALE1,SCALE2
REAL SLOPE,SUM1,SUM2,TEMP,VAR_UD,VAR_US,VAR_CD,VAR_CS,VAR_SD,VAR_SS
REAL WGHT,WIND_AZ,WIND_V,X1,Y1,X2,Y2,Z
REAL F1(P:Q,P:Q),F2(P:Q,P:Q),F3(P:Q,P:Q),F4(P:Q,P:Q),W1(P:Q,P:Q)
REAL DIST1(P1:Q1,P1:Q1),DIST2(P1:Q1,P1:Q1),FX1(P:Q,P:Q),FY1(P:Q,P:Q)
REAL XDIST1(P1:Q1,2),XDIST2(P1:Q1,2),YDIST1(P1:Q1,2),YDIST2(P1:Q1,2)
REAL PARM(-2:1,-2:1),VARY(0:4,0:4)
REAL RY(P1:Q1,P1:Q1),IY(P1:Q1,P1:Q1),MY(P1:Q1,P1:Q1),PY(P1:Q1,P1:Q1)
*** COMPLEX G1(P:Q,P:Q),G2(P:Q,P:Q)
*** COMPLEX G3(P:Q,P:Q),G4(P:Q,P:Q)
COMPLEX GX1(P:Q,P:Q),GY1(P:Q,P:Q),CWN(P:Q,P:Q)
CHARACTER*1 OPT
CHARACTER*18 FNAME
*
PI=3.14159
*
WRITE(*,*) ''
WRITE(*,*) 'ENTER FILENAME OF PARAMETERS = PARM'
*** READ(*,300) FNAME
FNAME='PARM.DAT'
*
WRITE(*,*) ''
WRITE(*,*) 'READING PARM.....'
CALL PLT3DINPUT(FNAME,PARM,N3,-2,1)
*
WRITE(*,*) ''
DO I=-2,1
WRITE(*,*) (PARM(I,J),J=-2,1)
END DO
*
K_NYQ=PARM(-2,-2)

```

```

DELTAK=PARM(-2,-1)
N2=INT(PARM(-2,0))
N3=INT(PARM(-1,1))
ISEED=JINT(PARM(-2,1))
WIND_V=PARM(-1,-2)
Z=PARM(-1,-1)
WIND_AZ=PARM(-1,0)
*
IF ((N.NE.N2).OR.((K_NYQ/DELTAK).NE.FLOAT(N/2))) THEN
    WRITE(*,*) 'INCORRECT PARAMETER FILE [N(SURF)]'
    GOTO 5
ENDIF
*
IF (N1.NE.N3) THEN
    WRITE(*,*) 'INCORRECT PARAMETER FILE [N1(DOME)]'
    GOTO 5
ENDIF
*
* CREATE N*N COMPLEX MATRIX GWN(L,M) OF NORMALIZED, FREQUENCY-DOMAIN
* WHITE GAUSSIAN NOISE
*
CALL WGN_SPEC(W1,GWN,ISEED,N,P,Q)
*
WIND_AZ=WIND_AZ*(PI/180.0)
*
CALL SEA_SPEC(WIND_V,Z,WIND_AZ,DELTAK,F1,F2,F3,F4,FX1,FY1,KCOR,N,P,Q)
*
WRITE(*,*) ''
WRITE(*,*) 'TSLOPE = TOTAL (X+Y) SLOPE SPECTRUM=SQRT[(K^2)*S(L,M)]'
IF (N.EQ.64) CALL ALPHA(F1,64)
FNAME='TSLOPE.RAE'
CALL PLT3DOUTPUT(FNAME,F1,N,P,Q)
*** CALL PLT3DFILE(F1,N)
*
WRITE(*,*) ''
WRITE(*,*) 'HEIGHT = ELEVATION SPECTRUM=SQRT[S(L,M)]'
IF (N.EQ.64) CALL ALPHA(F2,64)
FNAME='HEIGHT.RAE'
CALL PLT3DOUTPUT(FNAME,F2,N,P,Q)
*** CALL PLT3DFILE(F2,N)
*
WRITE(*,*) ''
WRITE(*,*) 'XSLOPE = X-COMPONENT SLOPE SPECTRUM=SQRT[(L^2)*S(L,M)]'
IF (N.EQ.64) CALL ALPHA(F3,64)
FNAME='XSLOPE.RAE'
CALL PLT3DOUTPUT(FNAME,F3,N,P,Q)
*** CALL PLT3DFILE(F3,N2)
*
WRITE(*,*) ''
WRITE(*,*) 'YSLOPE = Y-COMPONENT SLOPE SPECTRUM=SQRT[(M^2)*S(L,M)]'
IF (N.EQ.64) CALL ALPHA(F4,64)
FNAME='YSLOPE.RAE'

```

```

        CALL PLT3DOUTPUT(FNAME,F4,N,P,Q)
***      CALL PLT3DFILE(F4,N)
*
* FILTER MATRIX GWN WITH SLOPE-COMPONENT SPECTRUMS FX1 & FY1
* OPTIONAL FILTERING OF TOTAL_SLOPE SPECTRUM G1 & ELEVATION SPECTRUM G2
*
*      DO I=P,Q
*          DO J=P,Q
*
***          G1(I,J)=CMPLX(0.0,F1(I,J))*GWN(I,J)
***          G2(I,J)=CMPLX(0.0,F2(I,J))*GWN(I,J)
          GX1(I,J)=CMPLX(0.0,FX1(I,J))*GWN(I,J)
          GY1(I,J)=CMPLX(0.0,FY1(I,J))*GWN(I,J)
*
*      END DO
*      END DO
*
* BACK-FOURIER-TRANSFORM THE FILTERED SLOPE-COMPONENT SPECTRUMS GX1 & GY1
* OPTIONAL BACK-FOURIER-TRANSFORM OF G1 & G2
*
***      CALL FFT2D(G1,N,+1)
***      CALL FFT2D(G2,N,+1)
          CALL FFT2D(GX1,N,+1)
          CALL FFT2D(GY1,N,+1)
*
* OPTIONAL COMPLEX DISPLAY (REAL,IMAG,MAG,PHS) OF G1,G2,GX1 & GX2
*
***      IF (N.EQ.64) THEN
***          CALL RALPH(G1,64,RY,IY,MY,PY)
***          CALL RALPH(G2,64,RY,IY,MY,PY)
***          CALL RALPH(GX1,64,RY,IY,MY,PY)
***          CALL RALPH(GY1,64,RY,IY,MY,PY)
***      ENDIF
*
          SCALE1=FLOAT(Q1)
          SCALE2=FLOAT(Q1)/90.0
          WIND_AZ=WIND_AZ*(180.0/PI)
          WGHT=1.0/FLOAT(N*N)
*
* RESCALE THE TRANSFORMED SLOPE-COMPONENT ARRAYS GX1 & GY1
* CONVERT SLOPE-COMPONENTS INTO MAXIMUM_SLOPE ARRAY=FY1(X,Y) [DEG] [0..+90]
* AND SLOPE AZIMUTH ARRAY=FX1(X,Y) [DEG] [-180..+180]
* CALCULATE ALONG-WIND, CROSS-WIND, & COMBINED VARIANCES
* CREATE SLOPE DISTRIBUTION ARRAYS=DIST(SLOPE,AZIMUTH) [PROBABILITY] [0..1]
*
*      DO I=P,Q
*          DO J=P,Q
*
          X1=TAN(REAL(GX1(I,J))*DELTAK)
          Y1=TAN(REAL(GY1(I,J))*DELTAK)
*
*      IF (X1.NE.0.0) THEN

```

```

        FX1(I,J)=ATAN2D(Y1,X1)
    ELSE
        FX1(I,J)=0.0
    ENDIF
*
        SLOPE=SQRT((X1**2)+(Y1**2))
        FY1(I,J)=ATAND(SLOPE)
*
        ANG0=COSD(FX1(I,J))
        ANG1=SIND(FX1(I,J))
*
        X1=ANG0*SLOPE
        Y1=ANG1*SLOPE
*
        X2=ANG0*FY1(I,J)
        Y2=ANG1*FY1(I,J)
*
        ANG2=COSD(WIND_AZ-FX1(I,J))
        ANG3=SIND(WIND_AZ-FX1(I,J))
*
        VAR_UD=VAR_UD+WGHT*(ANG2*FY1(I,J))**2
        VAR_US=VAR_US+WGHT*(ANG2*SLOPE)**2
*
        VAR_CD=VAR_CD+WGHT*(ANG3*FY1(I,J))**2
        VAR_CS=VAR_CS+WGHT*(ANG3*SLOPE)**2
*
        VAR_SD=VAR_SD+WGHT*(FY1(I,J)**2)
        VAR_SS=VAR_SS+WGHT*(SLOPE**2)
*
        I1=INT(X1*SCALE1+SIGN(0.5,ANG0))
        J1=INT(Y1*SCALE1+SIGN(0.5,ANG1))
*
        I2=INT(X2*SCALE2+SIGN(0.5,ANG0))
        J2=INT(Y2*SCALE2+SIGN(0.5,ANG1))
*
        IF (ABS(I1*J1).LE.(Q1**2)) DIST1(I1,J1)=DIST1(I1,J1)+WGHT
        DIST2(I2,J2)=DIST2(I2,J2)+WGHT
*
    END DO
    END DO
*
    WRITE(*,*) ' '
    WRITE(*,*) 'WANT CROSS-SECTION DISTRIBUTIONS FOR PLT2D? [N]'
    READ(*,300) OPT
*
    IF ((OPT.EQ.'Y').OR.(OPT.EQ.'y')) THEN
*
        DO I=P1,Q1
        DO J=P1,Q1
*
            XDIST1(I,1)=FLOAT(I)/SCALE1
            XDIST1(I,2)=DIST1(I,0)

```

```

*
      YDIST1(J,1)=FLOAT(J)/SCALE1
      YDIST1(J,2)=DIST1(0,J)
*
      XDIST2(I,1)=FLOAT(I)/SCALE2
      XDIST2(I,2)=DIST2(I,0)
*
      YDIST2(J,1)=FLOAT(J)/SCALE2
      YDIST2(J,2)=DIST2(0,J)
*
      END DO
      END DO
*
      WRITE(*,*) ''
      WRITE(*,*) 'CROSS-SECTION DISTRIBUTIONS FOR PLT2D'
      WRITE(*,*) ''
      WRITE(*,*) 'XDIST1'
      CALL PLT2DFILE(XDIST1,N1)
      WRITE(*,*) ''
      WRITE(*,*) 'YDIST1'
      CALL PLT2DFILE(YDIST1,N1)
      WRITE(*,*) ''
      WRITE(*,*) 'XDIST2'
      CALL PLT2DFILE(XDIST2,N1)
      WRITE(*,*) ''
      WRITE(*,*) 'YDIST2'
      CALL PLT2DFILE(YDIST2,N1)
*
      ENDIF

* BIAS ALL ZERO VALUES IN F3 FOR ALPHASCALE DISPLAY
*
***      BIAS1=0.02/-25.0
***      BIAS2=0.02/-25.0
*
***      DO I=P1,Q1
***      DO J=P1,Q1
***          IF (DIST1(I,J).EQ.0.0) DIST1(I,J)=BIAS1
***          IF (DIST2(I,J).EQ.0.0) DIST2(I,J)=BIAS2
***      END DO
***      END DO
*
      WRITE(*,*) ''
      WRITE(*,*) 'SLOPE DISTRIBUTION ARRAY=DIST1 [PROB] (SLOPE COORDINATES)'
      IF (N1.EQ.64) CALL ALPHA(DIST1,64)
      FNAME='DIST1.RAE'
      CALL PLT3DOOUTPUT(FNAME,DIST1,N1,P1,Q1)
***      CALL PLT3DFILE(DIST1,N1)
*
      WRITE(*,*) ''
      WRITE(*,*) 'CROSS-WIND SLOPE VAR [SLOPE^2] & DEV= ',VAR_CS,SQRT(VAR_CS)
      WRITE(*,*) 'ALONG-WIND SLOPE VAR [SLOPE^2] & DEV= ',VAR_US,SQRT(VAR_US)

```

```

      WRITE(*,*) 'TOTAL-WIND SLOPE VAR [SLOPE^2] & DEV= ',VAR_SS,SQRT(VAR_SS)
*
      WRITE(*,*) ''
      WRITE(*,*) 'SLOPE DISTRIBUTION ARRAY=DIST2 [PROB] (FILENAME=DIST.RAY)'
      IF (N1.EQ.64) CALL ALPHA(DIST2,64)
      FNAME='DIST2.RAE'
      CALL PLT3DOUTPUT(FNAME,DIST2,N1,P1,Q1)
***   CALL PLT3DFILE(DIST2,N1)
*
      WRITE(*,*) ''
      WRITE(*,*) 'CROSS-WIND SLOPE VAR [DEG^2] & DEV= ',VAR_CD,SQRT(VAR_CD)
      WRITE(*,*) 'ALONG-WIND SLOPE VAR [DEG^2] & DEV= ',VAR_UD,SQRT(VAR_UD)
      WRITE(*,*) 'TOTAL-WIND SLOPE VAR [DEG^2] & DEV= ',VAR_SD,SQRT(VAR_SD)
*
      K_NYQ=FLOAT(N/2)*DELTA_K
      CALL PSVARY1(K_NYQ,N,WIND_V,Z,VARY)
*
      WRITE(*,*) ''
      WRITE(*,*) 'MAXIMUM-SLCOPE ARRAY=BETA [DEG] (FILENAME=BETA.RAY)'
      IF (N.EQ.64) CALL ALPHA(FY1,64)
      FNAME='BETA.RAE'
      CALL PLT3DOUTPUT(FNAME,FY1,N,P,Q)
***   CALL PLT3DFILE(FY1,N)
*
      WRITE(*,*) ''
      WRITE(*,*) 'SLOPE-AZIMUTH ARRAY=ALPH [DEG] (FILENAME=ALPH.RAY)'
      IF (N.EQ.64) CALL ALPHA(FX1,64)
      FNAME='ALPH.RAE'
      CALL PLT3DOUTPUT(FNAME,FX1,N,P,Q)
***   CALL PLT3DFILE(FX1,N)
*
      300   FORMAT (A)
*
      5     END

```

```

PROGRAM LIGHTBLUE
*
* CREATE UNPOLARIZED SKYDOME HEMISPHERICAL RADIANCE DISTRIBUTION
* CREATE HORIZONTALLY_ & VERTICALLY_POLARIZED RADIANCE COMPONENTS
*
* MAP SKYDOME RADIANCE DISTRIBUTIONS TO REFLECTED SURFACE COORDINATES
*
* CREATE UNPOLARIZED SUBDOME HEMISPHERICAL RADIANCE DISTRIBUTION
* CREATE HORIZONTALLY_ & VERTICALLY_POLARIZED RADIANCE COMPONENTS
*
* LSUB IS ALREADY LINEARLY MAPPED TO REFRACTED SURFACE COORDINATES
*
* INPUT:
* LREF = REFERENCE RADIANCE MEASURED AT THE ZENITH POINT, LSKY(0,0) [UNIT]
* PHIO = SUN AZIMUTH [DEG] [-180..+180]
* THETA = SENSOR ZENITH ANGLE [DEG] [0..+90]
* ZO = SOLAR ZENITH ANGLE [DEG] [0..+90]
*
* OUTPUT:
* LSKY = UNPOLARIZED SKYDOME SPATIAL RADIANCE DISTRIBUTION
* HSKY = HORIZONTALLY_POLARIZED LSKY
* VSKY = VERTICALLY_POLARIZED LSKY
*
* LSRF = LSKY MAPPED TO REFLECTED SURFACE COORDINATES (ALPH,BETA)
* HSRF = HSKY MAPPED TO REFLECTED SURFACE COORDINATES (ALPH,BETA)
* VSRF = VSKY MAPPED TO REFLECTED SURFACE COORDINATES (ALPH,BETA)
*
* LSUB = UNPOLARIZED SUBDOME SPATIAL RADIANCE DISTRIBUTION
* HSUB = HORIZONTALLY_POLARIZED LSUB
* VSUB = VERTICALLY_POLARIZED LSUB
*
IMPLICIT NONE
INTEGER I,J,N,N1,P,Q
PARAMETER (N=64,P=-N/2,Q=N/2-1)
REAL LSKY(P:Q,P:Q),HSKY(P:Q,P:Q),VSKY(P:Q,P:Q)
REAL LSRF(P:Q,P:Q),HSRF(P:Q,P:Q),VSRF(P:Q,P:Q)
REAL LSUB(P:Q,P:Q),HSUB(P:Q,P:Q),VSUB(P:Q,P:Q)
REAL LREF,PHIO,PI,THETA,ZO
REAL PARM(-2:1,-2:1)
CHARACTER*18 FNAME
*
PI=3.14159
*
WRITE(*,*) 'ENTER FILENAME OF PARAMETERS = PARM'
FNAME='PARM.DAT'
*
WRITE(*,*) 'READING PARM.....'
CALL PLT3DINPUT(FNAME,PARM,N1,-2,1)
*
DO I=-2,1
WRITE(*,*) (PARM(I,J),J=-2,1)
END DO

```

```
*  
N1=INT(PARM(-1,1))  
THETA=PARM(0,-2)  
PHIO=PARM(0,-1)  
Z0=PARM(0,0)  
LREF=PARM(0,1)  
  
*  
IF (N.NE.N1) THEN  
    WRITE(*,*) 'INCORRECT PARAMETER FILE'  
    GOTO 1  
ENDIF  
  
***  
WRITE(6,*) 'INPUT LREF [RADIANCE],PHIO [DEG],Z0 [DEG],THETA [DEG] ='  
READ(*,*) LREF,PHIO,Z0,THETA  
  
***  
WRITE(*,*) 'LREF,PHIO,Z0,THETA=',LREF,PHIO,Z0,THETA  
  
*  
PHIO=PHIO*(PI/180.0)  
Z0=Z0*(PI/180.0)  
  
***  
CALL SKYBLUE(LREF,PHIO,Z0,N,P,Q,LSKY,HSKY,VSKY)  
  
*  
CALL SKYMAP(LREF,PHIO,Z0,THETA,N,P,Q,LSRF,HSRF,VSRF)  
  
*  
CALL DEEPBLUE(LREF,Z0,N,P,Q,LSUB,HSUB,VSUB)  
  
*  
1    END
```

PROGRAM LIGHT

```

*
*   * CREATES OBJECT SPATIAL RADIANCE DISTRIBUTION IN THE DIRECTION OF SENSOR
*
*   * CALCULATES FORWARD FAST FOURIER TRANSFORMS OF SYNTHETIC IMAGES
*
*   IMPLICIT NONE
INTEGER I,I1,I2,INV,J,J1,J2,N,N1,N2,NTMP,P,P1,P2,Q,Q1,Q2
*   PARAMETER(N=64,P=-N/2,Q=N/2-1)
*   PARAMETER(N1=64,P1=-N1/2,Q1=N1/2-1)
*   PARAMETER(N2=64,P2=-N2/2,Q2=N2/2-1)
REAL IV1,IV2,IV3,NV1,NV2,NV3,RV1,RV2,RV3
REAL D1(P:Q,P:Q),D2(P:Q,P:Q),D3(P:Q,P:Q),D4(P:Q,P:Q)
REAL DARK1(P:Q,P:Q),DARK2(P:Q,P:Q),DELTA
REAL DST1(P:Q,P:Q),DST2(P:Q,P:Q),DST3(P:Q,P:Q),DST4(P:Q,P:Q)
REAL GAMH(P:Q,P:Q),GAMS(P:Q,P:Q),GAMV(P:Q,P:Q)
REAL HRAD(P1:Q1,P1:Q1),LRAD(P1:Q1,P1:Q1),VRAD(P1:Q1,P1:Q1)
REAL HSKY(P:Q,P:Q),LSKY(P:Q,P:Q),VSKY(P:Q,P:Q)
REAL HSUB(P:Q,P:Q),LSUB(P:Q,P:Q),VSUB(P:Q,P:Q)
REAL PARM(-2:1,-2:1),V1250,VARY(0:4,0:4)
REAL ALPH(P1:Q1,P1:Q1),COSALPHA,SINALPHA
*   REAL BETA(P1:Q1,P1:Q1),COSBETA,SINBETA
REAL DEVA,DEVC,DEVT
REAL GAMMAHOP,GAMMASUM,GAMMAVER
REAL MU,COSMU,SINMU
REAL NMBR(P:Q,P:Q),REF1,REF2,SWITCH
REAL NU,COSNU,SINNU,TANNU1,TANNU2
REAL OMEGA,COSOMEGA,SINOMEGA
*   REAL INDEXI,INDEXR,RHO,RHO1
REAL THETA,COSTHETA,SINTHETA
REAL DELTA_K,K_NYQ,PHI,PHI1,PHI2,RADIUS,RANGE,TEMP,WIND_V,WIND_AZ,Z
*** REAL RY(P:Q,P:Q),IY(P:Q,P:Q),MY(P:Q,P:Q),PY(P:Q,P:Q)
REAL ATH1(P:Q,P:Q),ATS1(P:Q,P:Q),ATV1(P:Q,P:Q)
REAL ATH2(P:Q,P:Q),ATS2(P:Q,P:Q),ATV2(P:Q,P:Q)
REAL HMAG0(P1:Q1,P1:Q1),LMAG0(P1:Q1,P1:Q1),VMAG0(P1:Q1,P1:Q1)
REAL HMAG1(P1:Q1,P1:Q1),LMAG1(P1:Q1,P1:Q1),VMAG1(P1:Q1,P1:Q1)
REAL HMAG2(P1:Q1,P1:Q1),LMAG2(P1:Q1,P1:Q1),VMAG2(P1:Q1,P1:Q1)
REAL HMAG3(P1:Q1,P1:Q1)
*   REAL RDH1(P2:Q2,P2:Q2),RDS1(P2:Q2,P2:Q2),RDV1(P2:Q2,P2:Q2)
REAL RDH2(P:Q,P:Q),RDS2(P:Q,P:Q),RDV2(P:Q,P:Q)
REAL RFLH(P1:Q1,P1:Q1),RFLS(P1:Q1,P1:Q1),RFLV(P1:Q1,P1:Q1)
REAL RFLX(P1:Q1,P1:Q1)
REAL RFRH(P1:Q1,P1:Q1),RFRS(P1:Q1,P1:Q1),RFRV(P1:Q1,P1:Q1)
REAL HAVG,LAVG,VAVG,HAVG1,LAVG1,VAVG1,HAVG2,LAVG2,VAVG2,XAVG
COMPLEX GD1(P2:Q2,P2:Q2),GH1(P2:Q2,P2:Q2),GV1(P2:Q2,P2:Q2)
*   COMPLEX GD2(P2:Q2,P2:Q2),GH2(P2:Q2,P2:Q2),GV2(P2:Q2,P2:Q2)
COMPLEX SH0(P1:Q1,P1:Q1),SL0(P1:Q1,P1:Q1),SV0(P1:Q1,P1:Q1)
*   COMPLEX SH1(P1:Q1,P1:Q1),SL1(P1:Q1,P1:Q1),SV1(P1:Q1,P1:Q1)
COMPLEX SH2(P1:Q1,P1:Q1),SL2(P1:Q1,P1:Q1),SV2(P1:Q1,P1:Q1)
COMPLEX SH3(P1:Q1,P1:Q1)
CHARACTER*18 FNAME0,FNAME1,FNAME2,FNAME3,FNAME4,FNAME5,FNAME6
CHARACTER*18 FNAME7,FNAME8,FNAME

```

```

*
***      WRITE(*,*) 'ENTER FILENAME OF PARAMETERS = PARM'
***      READ(*,300) FNAME0
      FNAME0='PARM.DAT'

*
***      WRITE(*,*) 'ENTER FILENAME OF MAXIMUM_SLOPE ARRAY = BETA'
***      READ(*,300) FNAME1
      FNAME1='BETA.RAE'

*
***      WRITE(*,*) 'ENTER FILENAME OF SLOPE_AZIMUTH ARRAY = ALPH'
***      READ(*,300) FNAME2
      FNAME2='ALPH.RAE'

*
***      WRITE(*,*) 'ENTER FILENAME OF SKYDOME_RADIANCE ARRAY = LSRF'
***      READ(*,300) FNAME3
      FNAME3='LSRF.RAY'

*
***      WRITE(*,*) 'ENTER FILENAME OF HORZ-POLARIZED LSRF = HSRF'
***      READ(*,300) FNAME4
      FNAME4='HSRF.RAY'

*
***      WRITE(*,*) 'ENTER FILENAME OF VERT-POLARIZED LSRF = VSRF'
***      READ(*,300) FNAME5
      FNAME5='VSRF.RAY'

*
***      WRITE(*,*) 'ENTER FILENAME OF SUBDOME_RADIANCE ARRAY = LSUB'
***      READ(*,300) FNAME6
      FNAME6='LSUB.RAY'

*
***      WRITE(*,*) 'ENTER FILENAME OF HORZ-POLARIZED LSUB = HSUB'
***      READ(*,300) FNAME7
      FNAME7='HSUB.RAY'

*
***      WRITE(*,*) 'ENTER FILENAME OF VERT-POLARIZED LSUB = VSUB'
***      READ(*,300) FNAME8
      FNAME8='VSUB.RAY'

*
      WRITE(*,*) 'READING PARM.....'
      CALL PLT3DINPUT(FNAME0,PARM,NTMP,-2,1)

*
      DO I=-2,1
         WRITE(*,*) (PARM(I,J),J=-2,1)
      END DO

*
      K_NYQ=PARM(-2,-2)
      DELTA_K=PARM(-2,-1)
      NTMP=INT(PARM(-2,0))

*
      IF ((N1.NE.NTMP).OR.((K_NYQ/DELTA_K).NE.FLOAT(NTMP/2))) THEN
         WRITE(*,*) 'INCORRECT PARAMETER FILE [N(SURF)]'
         GOTO 5
      ENDIF

```

```

*
      WIND_V=PARM(-1,-2)
      Z=PARM(-1,-1)
      WIND_AZ=PARM(-1,0)
      NTMP=INT(PARM(-1,1))
*
      IF (N.NE.NTMP) THEN
          WRITE(*,*) 'INCORRECT PARAMETER FILE [N1(DOME)]'
          GOTO 5
      ENDIF
*
      THETA=PARM(0,-2)
      RANGE=PARM(0,-1)
      INDEXR=PARM(1,-2)
      INDEXI=PARM(1,-1)
*
      WRITE(*,*) 'READING ARRAYS.....'
*
*** WRITE(*,*) 'READING BETA.....'
CALL PLT3DINPUT(FNAME1,BETA,NTMP,P,Q)
*
      IF (N1.NE.NTMP) THEN
          WRITE(*,*) 'BETA & DEFAULT ARRAY SIZES DO NOT MATCH'
          GOTO 5
      ENDIF
*
*** WRITE(*,*) 'READING ALPH.....'
CALL PLT3DINPUT(FNAME2,ALPH,NTMP,P,Q)
*
      IF (N1.NE.NTMP) THEN
          WRITE(*,*) 'ALPH & DEFAULT ARRAY SIZES DO NOT MATCH'
          GOTO 5
      ENDIF
*
*** WRITE(*,*) 'READING LSKY.....'
CALL PLT3DINPUT(FNAME3,LSKY,NTMP,P,Q)
*
*** IF (N.NE.NTMP) THEN
***     WRITE(*,*) 'LSKY & DEFAULT ARRAY SIZES DO NOT MATCH'
***     GOTO 5
*** ENDIF
*
*** WRITE(*,*) 'READING HSKY.....'
CALL PLT3DINPUT(FNAME4,HSKY,NTMP,P,Q)
*
      IF (N.NE.NTMP) THEN
          WRITE(*,*) 'HSKY & DEFAULT ARRAY SIZES DO NOT MATCH'
          GOTO 5
      ENDIF
*
*** WRITE(*,*) 'READING VSKY.....'
CALL PLT3DINPUT(FNAME5,VSKY,NTMP,P,Q)

```

```

*
    IF (N.NE.NTMP) THEN
        WRITE(*,*) 'VSKY & DEFAULT ARRAY SIZES DO NOT MATCH'
        GOTO 5
    ENDIF
*
***   WRITE(*,*) 'READING LSUB.....'
***   CALL PLT3DINPUT(FNAME6,LSUB,NTMP,P,Q)
*
***   IF (N.NE.NTMP) THEN
***       WRITE(*,*) 'LSUB & DEFAULT ARRAY SIZES DO NOT MATCH'
***       GOTO 5
***   ENDIF
*
***   WRITE(*,*) 'READING HSUB.....'
CALL PLT3DINPUT(FNAME7,HSUB,NTMP,P,Q)
*
IF (N.NE.NTMP) THEN
    WRITE(*,*) 'HSUB & DEFAULT ARRAY SIZES DO NOT MATCH'
    GOTO 5
ENDIF
*
***   WRITE(*,*) 'READING VSUB.....'
CALL PLT3DINPUT(FNAME8,VSUB,NTMP,P,Q)
*
IF (N.NE.NTMP) THEN
    WRITE(*,*) 'VSUB & DEFAULT ARRAY SIZES DO NOT MATCH'
    GOTO 5
ENDIF
*
***   WRITE(*,*) 'BETA'
***   IF (N1.EQ.64) CALL ALPHA(BETA,64)
*
***   WRITE(*,*) 'ALPH'
***   IF (N1.EQ.64) CALL ALPHA(ALPH,64)
*
***   WRITE(*,*) 'LSKY'
***   IF (N.EQ.64) CALL ALPHA(LSKY,64)
*
***   WRITE(*,*) 'HSKY'
***   IF (N.EQ.64) CALL ALPHA(HSKY,64)
*
***   WRITE(*,*) 'VSKY'
***   IF (N.EQ.64) CALL ALPHA(VSKY,64)
*
***   WRITE(*,*) 'LSUB'
***   IF (N.EQ.64) CALL ALPHA(LSUB,64)
*
***   WRITE(*,*) 'HSUB'
***   IF (N.EQ.64) CALL ALPHA(HSUB,64)
*
***   WRITE(*,*) 'VSUB'

```

```

***      IF (N.EQ.64) CALL ALPHA(VSUB,64)
*
* NOTE: NO SLOPE VARIANCE CALCULATIONS REQUIRED WHEN SURFACE IS FLAT
* IE: IF WIND_V = 0.0 THEN MEAN SLOPE = 0.0 AND SLOPE VARIANCE = 0.0
* SLOPE DISTRIBUTION = DELTA FUNCTION
*
*      IF (WIND_V.EQ.0.0) GOTO 1
*
* CALCULATE ALONG-WIND, CROSS-WIND & TOTAL-WIND SLOPE VARIANCES ABOVE, BELOW
* & WITHIN THE SPATIAL (WAVENUMBER) FREQUENCY RANGE OF THE MODEL SURFACE
*
*      CALL PSVARY2(K_NYQ,N,WIND_V,Z,V1250,VARY)
*
* CREATE HIGH_K_SLOPE_DISTRIBUTION ARRAY = DST1
* (K_NYQ < K < INF)
* FOR CORRELATION OF SURFACE_REFLECTED_SKYDOME RADIANCE AND
* FOR CORRELATION OF SURFACE_REFRACTED_SUBDOME RADIANCE
*
*      WRITE(*,*) ''
*      WRITE(*,*) 'CALCULATING HIGH_K_SLOPE_DISTRIBUTIONS'
*
*      DEVA=SQRT(VARY(4,3)-VARY(3,3))
*      DEV_C=SQRT(VARY(4,2)-VARY(3,2))
*      WRITE(*,*) 'VAR_C,DEV_C      = ',DEV_C**2,DEV_C
*      WRITE(*,*) 'VAR_A,DEV_A      = ',DEVA**2,DEVA
*      DEVT=SQRT((DEVA**2)+(DEV_C**2))
*      WRITE(*,*) 'VAR_TOT,DEV_TOT = ',DEVT**2,DEVT
*
*      CALL DIST_GEN(1,N,P,Q,WIND_AZ,V1250,DEVA,DEV_C,D1,D2,D3,D4)
*
*      DO I=P,Q
*      DO J=P,Q
*          IF (D3(I,J).GT.0.0) DST1(I,J)=D3(I,J)
*          D1(I,J)=0.0
*          D2(I,J)=0.0
*          D3(I,J)=0.0
*          D4(I,J)=0.0
*      END DO
*      END DO
*
* CREATE MID_K_SLOPE_DISTRIBUTION ARRAY = DST2
* (DELTA_K < K < INF)
*
***      WRITE(*,*) ''
***      WRITE(*,*) 'CALCULATING MID_K_SLOPE_DISTRIBUTIONS'
*
***      DEVA=SQRT(VARY(4,3)-VARY(2,3))
***      DEV_C=SQRT(VARY(4,2)-VARY(2,2))
***      WRITE(*,*) 'VAR_C,DEV_C      = ',DEV_C**2,DEV_C
***      WRITE(*,*) 'VAR_A,DEV_A      = ',DEVA**2,DEVA
***      DEVT=SQRT((DEVA**2)+(DEV_C**2))
***      WRITE(*,*) 'VAR_TOT,DEV_TOT = ',DEVT**2,DEVT

```

```

*
*** CALL DIST_GEN(1,N,P,Q,WIND_AZ,V1250,DEVA,DEVC,D1,D2,D3,D4)
*
*** DO I=P,Q
*** DO J=P,Q
***     IF (D3(I,J).GT.0.0) DST2(I,J)=D3(I,J)
***     D1(I,J)=0.0
***     D2(I,J)=0.0
***     D3(I,J)=0.0
***     D4(I,J)=0.0
*** END DO
*** END DO
*
* CREATE FULL_K_SLOPE_DISTRIBUTION ARRAY = DST3
* (0 < K < INF)
*
*** WRITE(*,*) ''
*** WRITE(*,*) 'CALCULATING FULL_K_SLOPE_DISTRIBUTIONS'
*
*** DEVA=SQRT(VARY(4,3))
*** DEVVC=SQRT(VARY(4,2))
*** WRITE(*,*) 'VAR_C,DEV_C      = ',DEVVC**2,DEVVC
*** WRITE(*,*) 'VAR_A,DEV_A      = ',DEVA**2,DEVA
*** DEVT=SQRT((DEVA**2)+(DEVVC**2))
*** WRITE(*,*) 'VAR_TOT,DEV_TO    = ',DEVT**2,DEVT
*
*** CALL DIST_GEN(1,N,P,Q,WIND_AZ,V1250,DEVA,DEVC,D1,D2,D3,D4)
*
*** DO I=P,Q
*** DO J=P,Q
***     IF (D3(I,J).GT.0.0) DST3(I,J)=D3(I,J)
***     D1(I,J)=0.0
***     D2(I,J)=0.0
***     D3(I,J)=0.0
***     D4(I,J)=0.0
*** END DO
*** END DO
*
* CREATE SAMPLE_K_SLOPE_DISTRIBUTION ARRAY = DST4
* (DELTA_K < K < K_NYQ)
*
*** WRITE(*,*) ''
*** WRITE(*,*) 'CALCULATING SAMPLE_K_SLOPE_DISTRIBUTIONS'
*
*** DEVA=SQRT(VARY(3,3)-VARY(2,3))
*** DEVVC=SQRT(VARY(3,2)-VARY(2,2))
*** WRITE(*,*) 'VAR_C,DEV_C      = ',DEVVC**2,DEVVC
*** WRITE(*,*) 'VAR_A,DEV_A      = ',DEVA**2,DEVA
*** DEVT=SQRT((DEVA**2)+(DEVVC**2))
*** WRITE(*,*) 'VAR_TOT,DEV_TCT   = ',DEVT**2,DEVT
*
*** CALL DIST_GEN(1,N,P,Q,WIND_AZ,V1250,DEVA,DEVC,D1,D2,D3,D4)

```

```

*
***      DO I=P,Q
***      DO J=P,Q
***          IF (D3(I,J).GT.0.0) DST4(I,J)=D3(I,J)
***          D1(I,J)=0.0
***          D2(I,J)=0.0
***          D3(I,J)=0.0
***          D4(I,J)=0.0
***      END DO
***      END DO
*
* CALCULATE ANGULAR SURFACE_REFLECTED_SKYDOME_RADIANCE DISTRIBUTIONS:
*
* 1)CREATE ANGULAR POLARIZED (HORZ & VERT) FRESNEL_COEFFICIENT DISTRIBUTION
*   RELATIVE TO SENSOR AND SURFACE COORDINATES
*
*   CREATE ANGULAR SURFACE_PROJECTION_ATTENUATION_COEFFICIENT DISTRIBUTION
*   RELATIVE TO SENSOR AND SURFACE COORDINATES
*
1      CALL DARKNESS(P,Q,THETA,INDEXR,INDEXI,GAMH,GAMV,DARK1)
*
* 2)CREATE ATTENUATED_SKYDOME_RADIANCE DISTRIBUTIONS
*
      DO I=P,Q
      DO J=P,Q
*
      ATH1(I,J)=GAMH(I,J)*DARK1(I,J)*HSKY(I,J)
      ATV1(I,J)=GAMV(I,J)*DARK1(I,J)*VSKY(I,J)
      ATS1(I,J)=ATH1(I,J)+ATV1(I,J)
*
      IF (WIND_V.EQ.0.0) THEN
          RDH1(I,J)=ATH1(I,J)
          RDV1(I,J)=ATV1(I,J)
          RDS1(I,J)=ATS1(I,J)
      ENDIF
*
      END DO
      END DO
*
      IF (WIND_V.EQ.0.0) GOTO 2
*
      WRITE(*,*) 'ATH1 = ATTENUATED HORIZONTAL_SKYDOME_RADIANCE'
***      IF (N.EQ.64) CALL ALPHA(ATH1,64)
***      CALL PLT3DFILE(ATH1,N)
*
      WRITE(*,*) 'ATV1 = ATTENUATED VERTICAL_SKYDOME_RADIANCE'
***      IF (N.EQ.64) CALL ALPHA(ATV1,64)
***      CALL PLT3DFILE(ATV1,N)
*
      WRITE(*,*) 'ATS1 = ATTENUATED TOTAL_SKYDOME_RADIANCE'
***      IF (N.EQ.64) CALL ALPHA(ATS1,64)
***      CALL PLT3DFILE(ATS1,N)

```

```

*
* 3) CORRELATE SKYDOME BY HIGH_K_WATER_SURFACE_FILTER=DST1 (K > K_NYQ)
*
      WRITE(*,*) 'DST1 = HIGH_K SLOPE PROBABILITY DISTRIBUTION'
*** IF (N.EQ.64) CALL ALPHA(DST1,64)
*** CALL PLT3DFILE(DST1,N)
*
      DO I=P,Q
      DO J=P,Q
          GD1(I,J)=CMPLX(DST1(-I,-J),0.0)
          GH1(I,J)=CMPLX(ATH1(I,J),0.0)
          GV1(I,J)=CMPLX(ATV1(I,J),0.0)
      END DO
      END DO
*
      CALL FFT2D(GD1,N2,-1)
      CALL FFT2D(GH1,N2,-1)
      CALL FFT2D(GV1,N2,-1)
*
      DO I=P2,Q2
      DO J=P2,Q2
          GH1(I,J)=GH1(I,J)*GD1(I,J)
          GV1(I,J)=GV1(I,J)*GD1(I,J)
      END DO
      END DO
*
      CALL FFT2D(GH1,N2,+1)
      CALL FFT2D(GV1,N2,+1)
*
      TEMP=FLOAT(N2*N2)
*
      DO I=P2,Q2
      DO J=P2,Q2
*
          RDH1(I,J)=CABS(GH1(I,J))/TEMP
          RDV1(I,J)=CABS(GV1(I,J))/TEMP
          RDS1(I,J)=RDH1(I,J)+RDV1(I,J)
*
      END DO
      END DO
*
      2      WRITE(*,*) 'RDH1 = FILTERED ATH1'
*** IF (N2.EQ.64) CALL ALPHA(RDH1,64)
*** CALL PLT3DFILE(RDH1,N2)
*
      WRITE(*,*) 'RDV1 = FILTERED ATVi'
*** IF (N2.EQ.64) CALL ALPHA(RDV1,64)
*** CALL PLT3DFILE(RDV1,N2)
*
      WRITE(*,*) 'RDS1 = FILTERED ATSl'
*** IF (N2.EQ.64) CALL ALPHA(RDS1,64)
*** CALL PLT3DFILE(RDS1,N2)

```

```

*
* CALCULATE ANGULAR SUBSURFACE_REFRACTED_SUBDOME_RADIANCE DISTRIBUTIONS:
*
* 1)CREATE ANGULAR POLARIZED (HORZ & VERT) UPWELLING_RADIANCE DISTRIBUTION
*   RELATIVE TO SENSOR COORDINATES
*
*   CREATE ANGULAR SUBSURFACE_PROJECTION_ATTENUATION_COEFFICIENT DISTRIBUTION
*   RELATIVE TO SENSOR COORDINATES
*
* 2)CREATE ATTENUATED_SUBDOME_RADIANCE DISTRIBUTIONS
*
      DO I=P,Q
      DO J=P,Q
*
      ATH2(I,J)=DARK1(I,J)*HSUB(I,J)
      ATV2(I,J)=DARK1(I,J)*VSUB(I,J)
      ATS2(I,J)=ATH2(I,J)+ATV2(I,J)
*
***      IF (WIND_V.EQ.0.0) THEN
            RDH2(I,J)=ATH2(I,J)
            RDV2(I,J)=ATV2(I,J)
            RDS2(I,J)=ATS2(I,J)
***      ENDIF
*
      END DO
      END DO
*
      SWITCH = 1
      IF (SWITCH.EQ.1) GOTO 3
*
***      IF (WIND_V.EQ.0.0) GOTO 3
*
      WRITE(*,*) 'ATH2 = ATTENUATED HORZ_POLARIZED_SUBDOME_RADIANCE'
      IF (N.EQ.64) CALL ALPHA(ATH2,64)
      CALL PLT3DFILE(ATH2,N)
*
      WRITE(*,*) 'ATV2 = ATTENUATED VERT_POLARIZED_SUBDOME_RADIANCE'
      IF (N.EQ.64) CALL ALPHA(ATV2,64)
      CALL PLT3DFILE(ATV2,N)
*
      WRITE(*,*) 'ATS2 = ATTENUATED TOTAL_SUBDOME_RADIANCE'
      IF (N.EQ.64) CALL ALPHA(ATS2,64)
      CALL PLT3DFILE(ATS2,N)
*
* 3)CORRELATE SUBDOME BY HIGH_K_WATER_SURFACE_FILTER=DST1 (K > K_NYQ)
* NOTE:DST1 WAS PREVIOUSLY FORWARD_TRANSFORMED BY REFLECTION ALGORITHM
*
      DO I=P,Q
      DO J=P,Q
          GD2(I,J)=CMPLX(DST1(-I,-J),0.0)
          GH2(I,J)=CMPLX(ATH2(I,J),0.0)
          GV2(I,J)=CMPLX(ATV2(I,J),0.0)

```

```

        END DO
        END DO

    *
        CALL FFT2D(GD2,N2,-1)
        CALL FFT2D(GH2,N2,-1)
        CALL FFT2D(GV2,N2,-1)

    *
        DO I=P2,Q2
        DO J=P2,Q2
            GH2(I,J)=GH2(I,J)*GD2(I,J)
            GV2(I,J)=GV2(I,J)*GD2(I,J)
        END DO
        END DO

    *
        CALL FFT2D(GH2,N2,+1)
        CALL FFT2D(GV2,N2,+1)

    *
        TEMP=FLOAT(N2*N2)

    *
        DO I=P2,Q2
        DO J=P2,Q2
    *
        RDH2(I,J)=CABS(GH2(I,J))/TEMP
        RDV2(I,J)=CABS(GV2(I,J))/TEMP
        RDS2(I,J)=RDH2(I,J)+RDV2(I,J)

    *
        END DO
        END DO

    *
3       WRITE(*,*) 'RDH2 = FILTERED ATH2'
***      IF (N2.EQ.64) CALL ALPHA(RDH2,64)
***      CALL PLT3DFILE(RDH2,N2)

    *
        WRITE(*,*) 'RDV2 = FILTERED ATV2'
***      IF (N2.EQ.64) CALL ALPHA(RDV2,64)
***      CALL PLT3DFILE(RDV2,N2)

    *
        WRITE(*,*) 'RDS2 = FILTERED ATS2'
***      IF (N2.EQ.64) CALL ALPHA(RDS2,64)
***      CALL PLT3DFILE(RDS2,N2)

    *
* CALCULATE REFLECTED & REFRACTED RADIANCE DISTRIBUTIONS IN DIRECTION OF SENSOR
    *
        DELTA=90.0/FLOAT(Q)
        COSTHETA=COSD(THETA)
        SINTHETA=SIND(THETA)

    *
* IV1,IV2,IV3 ARE THE RESOLVED COMPONENTS OF THE SENSOR COORDINATE VECTOR
    *
        IV1=SINTHETA*SIND(180.0)
        IV2=SINTHETA*COSD(180.0)
        IV3=COSTHETA

```

```

*
DO I=P1,Q1
DO J=P1,Q1
*
COSBETA=COSD(BETA(I,J))
SINBETA=SIND(BETA(I,J))
COSALPHA=COSD(ALPH(I,J))
SINALPHA=SIND(ALPH(I,J))
*
* NV1,NV2,NV3 ARE THE RESOLVED COMPONENTS OF THE SURFACE NORMAL VECTOR
*
NV1=SINBETA*SINALPHA
NV2=SINBETA*COSALPHA
NV3=COSBETA
*
* DOT PRODUCT OF IV & NV = |IV||NV|(COSOMEGA) = COSOMEGA
*
COSOMEGA=(IV1*NV1)+(IV2*NV2)+(IV3*NV3)
OMEGA=ACOSD(COSOMEGA)
SINOMEGA=SIND(OMEGA)
*
* RV1,RV2,RV3 ARE THE RESOLVED COMPONENTS OF THE REFLECTED SKYDOME VECTOR
*
RV1=(NV1*2.0*COSOMEGA)-IV1
RV2=(NV2*2.0*COSOMEGA)-IV2
RV3=(NV3*2.0*COSOMEGA)-IV3
*
COSMU=RV3
MU=ACOSD(COSMU)
SINMU=SIND(MU)
*
IF (MU.GT.0.0) THEN
    NU=ATAN2D(RV1,RV2)
ELSE
    NU=0.0
ENDIF
*
COSNU=COSD(NU)
SINNU=SIND(NU)
*
RADIUS=MU/DELTA
I2=INT(COSNU*(RADIUS+0.5))
J2=INT(SINNU*(RADIUS+0.5))
*
I1=INT(COSALPHA*(BETA(I,J)/DELTA)+0.5)
J1=INT(SINALPHA*(BETA(I,J)/DELTA)+0.5)
*
RFLH(I,J)=RDH1(I1,J1)
RFLV(I,J)=RDV1(I1,J1)
RFLS(I,J)=RFLH(I,J)+RFLV(I,J)
*
RFRH(I,J)=RDV2(I1,J1)

```

```

RFRV(I,J)=RDV2(I1,J1)
RFRS(I,J)=RFRH(I,J)+RFRV(I,J)
*
HRAD(I,J)=RFLH(I,J)+RFRH(I,J)
VRAD(I,J)=RFLV(I,J)+RFRV(I,J)
LRAD(I,J)=HRAD(I,J)+VRAD(I,J)
*
IF (MU.LE.90.0) THEN
    RFLX(I,J)=ATH1(I1,J1)
    NMBR(I2,J2)=NMBR(I2,J2)+1.0
ELSEIF (RDS1(I1,J1).GT.0.0) THEN
    REF1=REF1+1.0
    REF2=REF2+1.0
ELSE
    REF2=REF2+1.0
ENDIF
*
4    END DO
END DO
*
TEMP=FLOAT(N1*N1)
*
DO I=P1,Q1
DO J=P1,Q1
*
SH1(I,J)=CMPLX(RFLH(I,J),0.0)
*** SV1(I,J)=CMPLX(RFLV(I,J),0.0)
SL1(I,J)=CMPLX(RFLS(I,J),0.0)
*
HAVG1=HAVG1+RFLH(I,J)/TEMP
VAVG1=VAVG1+RFLV(I,J)/TEMP
LAVG1=LAVG1+RFLS(I,J)/TEMP
*
*** SH2(I,J)=CMPLX(RFRH(I,J),0.0)
*** SV2(I,J)=CMPLX(RFRV(I,J),0.0)
*** SL2(I,J)=CMPLX(RFRS(I,J),0.0)
*
HAVG2=HAVG2+RFRH(I,J)/TEMP
VAVG2=VAVG2+RFRV(I,J)/TEMP
LAVG2=LAVG2+RFRS(I,J)/TEMP
*
SH0(I,J)=CMPLX(HRAD(I,J),0.0)
*** SV0(I,J)=CMPLX(VRAD(I,J),0.0)
SL0(I,J)=CMPLX(LRAD(I,J),0.0)
*
HAVG=HAVG+HRAD(I,J)/TEMP
VAVG=VAVG+VRAD(I,J)/TEMP
LAVG=LAVG+LRAD(I,J)/TEMP
*
SH3(I,J)=CMPLX(RFLX(I,J),0.0)
*
XAVG=XAVG+RFLX(I,J)/TEMP

```

```

*
      END DO
      END DO

*
      WRITE(*,*) ''
      WRITE(*,*) 'SYNTHETIC RADIANCE IMAGES'

*
      WRITE(*,*) ''
      WRITE(*,*) 'NMBR=PIXEL COUNTS REFLECTING FROM SKYDOME LOCATION'
***   IF (N.EQ.64) CALL ALPHA(NMBR,64)
***   CALL PLT3DFILE(NMBR,N)

*
      WRITE(*,*) ''
      WRITE(*,*) 'NMBR OF 2NDARY REFLECTORS & RATIO = ',REF2,REF2/TEMP

      WRITE(*,*) '(IE PIXELS WITH MU > 90.0 DEGREES)'

*
      WRITE(*,*) ''
      WRITE(*,*) 'NMBR THAT REFLECT SOME SKY & RATIO = ',REF1,REF1/TEMP

      WRITE(*,*) '(IE PIXELS WITH 90.0 DEGREES < MU < MU_MAX)'

*
      WRITE(*,*) ''
      WRITE(*,*) 'RFLH=REFLECTED HORIZ_POLARIZED RADIANCE'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',HAVG1
***   IF (N1.EQ.64) CALL ALPHA(RFLH,64)
***   CALL PLT3DFILE(RFLH,N1)

*
      WRITE(*,*) ''
      WRITE(*,*) 'RFLX=RFLH WITHOUT SUB-RESOLUTION FILTERING'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',XAVG
***   IF (N1.EQ.64) CALL ALPHA(RFLX,64)
***   CALL PLT3DFILE(RFLX,N1)

*
      WRITE(*,*) ''
      WRITE(*,*) 'RFLV=REFLECTED VERT_POLARIZED RADIANCE'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',VAVG1
***   IF (N1.EQ.64) CALL ALPHA(RFLV,64)
***   CALL PLT3DFILE(RFLV,N1)

*
      WRITE(*,*) ''
      WRITE(*,*) 'RFLS=REFLECTED TOTAL RADIANCE'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',LAVG1
***   IF (N1.EQ.64) CALL ALPHA(RFLS,64)
***   CALL PLT3DFILE(RFLS,N1)

*
      WRITE(*,*) ''
      WRITE(*,*) 'RFRH=REFRACTED HORIZ_POLARIZED RADIANCE'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',HAVG2
***   IF (N1.EQ.64) CALL ALPHA(RFRH,64)
***   CALL PLT3DFILE(RFRH,N1)

```

```

        WRITE(*,*) ''
        WRITE(*,*) 'RFRV=REFRACTED VERT_POLARIZED RADIANCE'
        WRITE(*,*) 'MEAN PIXEL RADIANCE = ',VAVG2
*** IF (N1.EQ.64) CALL ALPHA(RFRV,64)
*** CALL PLT3DFILE(RFRV,N1)
*
        WRITE(*,*) ''
        WRITE(*,*) 'RFRS=REFRACTED TOTAL RADIANCE'
        WRITE(*,*) 'MEAN PIXEL RADIANCE = ',LAVG2
*** IF (N1.EQ.64) CALL ALPHA(RFRS,64)
*** CALL PLT3DFILE(RFRS,N1)
*
        WRITE(*,*) ''
        WRITE(*,*) 'HRAD=REFLECTED & REFRACTED HORIZ_POLARIZED RADIANCE'
        WRITE(*,*) 'MEAN PIXEL RADIANCE = ',HAVG
*** IF (N1.EQ.64) CALL ALPHA(HRAD,64)
*** CALL PLT3DFILE(HRAD,N1)
*
        WRITE(*,*) ''
        WRITE(*,*) 'VRAD=REFLECTED & REFRACTED VERT_POLARIZED RADIANCE'
        WRITE(*,*) 'MEAN PIXEL RADIANCE = ',VAVG
*** IF (N1.EQ.64) CALL ALPHA(VRAD,64)
*** CALL PLT3DFILE(VRAD,N1)
*
        WRITE(*,*) ''
        WRITE(*,*) 'LRAD=REFLECTED & REFRACTED TOTAL RADIANCE'
        WRITE(*,*) 'MEAN PIXEL RADIANCE = ',LAVG
*** IF (N1.EQ.64) CALL ALPHA(LRAD,64)
*** CALL PLT3DFILE(LRAD,N1)
*
        CALL FFT2D(SL1,N1,-1)
        CALL FFT2D(SH1,N1,-1)
*** CALL FFT2D(SV1,N1,-1)
*
*** CALL FFT2D(SL2,N1,-1)
*** CALL FFT2D(SH2,N1,-1)
*** CALL FFT2D(SV2,N1,-1)
*
        CALL FFT2D(SL0,N1,-1)
        CALL FFT2D(SH0,N1,-1)
*** CALL FFT2D(SV0,N1,-1)
*
        CALL FFT2D(SH3,N1,-1)
*
        DO I=P1,Q1
        DO J=P1,Q1
*
        HMAG1(I,J)=CABS(SH1(I,J))/TEMP
*** VMAG1(I,J)=CABS(SV1(I,J))/TEMP
        LMAG1(I,J)=CABS(SL1(I,J))/TEMP
*
*** HMAG2(I,J)=CABS(SH2(I,J))/TEMP

```

```

***      VMAG2(I,J)=CABS(SV2(I,J))/TEMP
***      LMAG2(I,J)=CABS(SL2(I,J))/TEMP
*
      HMAGO(I,J)=CABS(SH0(I,J))/TEMP
***      VMAGO(I,J)=CABS(SV0(I,J))/TEMP
      LMAGO(I,J)=CABS(SL0(I,J))/TEMP
*
      HMAG3(I,J)=CABS(SH3(I,J))/TEMP
*
      END DO
      END DO
*
      WRITE(*,*) ' '
      WRITE(*,*) 'FORWARD TRANSFORMS OF RADIANCE IMAGES'
*
      WRITE(*,*) 'HMAG1= HORZ/REFLECTED RAD MAG SPECTRUM'
      TEMP=HMAG1(0,0)
      WRITE(*,*) 'HMAG1(0,0) = ',HMAG1(0,0)
***      HMAG1(0,0)=0.0
***      IF (N1.EQ.64) CALL ALPHA(HMAG1,64)
***      HMAG1(0,0)=TEMP
      FNAME='HMAG1.RAD'
      CALL PLT3DOUTPUT(FNAME,HMAG1,N1,P1,Q1)
      CALL PLT3DFILE(HMAG1,N1)
*
      WRITE(*,*) 'HMAG3= HMAG1 WITHOUT SUB-RESOLUTION FILTER'
      TEMP=HMAG3(0,0)
      WRITE(*,*) 'HMAG3(0,0) = ',HMAG3(0,0)
***      HMAG3(0,0)=0.0
***      IF (N1.EQ.64) CALL ALPHA(HMAG3,64)
***      HMAG3(0,0)=TEMP
      FNAME='HMAG3.RAD'
      CALL PLT3DOUTPUT(FNAME,HMAG3,N1,P1,Q1)
      CALL PLT3DFILE(HMAG3,N1)
*
      WRITE(*,*) 'LMAG1= TOTAL/REFLECTED RAD MAG SPECTRUM'
      TEMP=LMAG1(0,0)
      WRITE(*,*) 'LMAG1(0,0) = ',LMAG1(0,0)
***      LMAG1(0,0)=0.0
***      IF (N1.EQ.64) CALL ALPHA(LMAG1,64)
***      LMAG1(0,0)=TEMP
      FNAME='LMAG1.RAD'
      CALL PLT3DOUTPUT(FNAME,LMAG1,N1,P1,Q1)
      CALL PLT3DFILE(LMAG1,N1)
*
      WRITE(*,*) 'HMAGO= HORZ/REFL & REFR RAD MAG SPECTRUM'
      TEMP=HMAGO(0,0)
      WRITE(*,*) 'HMAGO(0,0) = ',HMAGO(0,0)
***      HMAGO(0,0)=0.0
***      IF (N1.EQ.64) CALL ALPHA(HMAGO,64)
***      HMAGO(0,0)=TEMP
      FNAME='HMAGO.RAD'

```

```
      CALL PLT3DOUTPUT(FNAME,HMAGO,N1,P1,Q1)
***      CALL PLT3DFILE(HMAGO,N1)
*
      WRITE(*,*) 'LMAGO= TOTAL/REFL & REFR RAD MAG SPECTRUM'
      TEMP=LMAGO(0,0)
      WRITE(*,*) 'LMAGO(0,0) = ',LMAGO(0,0)
***      LMAGO(0,0)=0.0
***      IF (N1.EQ.64) CALL ALPHA(LMAGO,64)
***      LMAGO(0,0)=TEMP
      FNAME='LMAGO.RAD'
      CALL PLT3DOUTPUT(FNAME,LMAGO,N1,P1,Q1)
***      CALL PLT3DFILE(LMAGO,N1)
*
      GOTO 1000
5      WRITE(*,*) 'BAD FILE OPENING'
300    FORMAT (A)
1000   END
```

PROGRAM ANALYS

```
*  
* ANALYSIS PROGRAM FOR THE GENERATION AND ANALYSIS OF 2D DIFFERENCE  
* SPECTRA  
*  
IMPLICIT NONE  
INTEGER I,I1,I2,INV,J,J1,J2,KINT,N,N1,N2,NTMP,P,PHINT,P1,P2,Q,Q1,Q2,R  
PARAMETER(N=64,P=-N/2,Q=N/2-1,R=Q+1)  
PARAMETER(N1=64,P1=-N1/2,Q1=N1/2-1)  
PARAMETER(N2=64,P2=-N2/2,Q2=N2/2-1)  
REAL ANG0,ANG1,COSA1,COSA2,K,L,M,KMAG,DELTA_K,DELTA_PHI,PI,SINA1,SINA2  
REAL DAT1,DAT2,DAT3,DAT4,DAT5  
REAL DATH1,DATH1,DATH0,DATH0,DATH3,DATH0  
REAL DST1(P:Q,P:Q)  
REAL HIGH,LOW,FACTOR  
REAL FACTOR1,FACTOR2,FACTOR3,FACTOR4,FACTOR5  
REAL PARM(-2:1,-2:1),V1250  
REAL K_NYQ,PHI,PHI1,PHI2,RADIUS,RANGE,TEMP,TEMP1,WIND_V,WIND_AZ,Z  
REAL HMAG0(P1:Q1,P1:Q1),LMAG0(P1:Q1,P1:Q1),VMAG0(P1:Q1,P1:Q1)  
REAL HMAG1(P1:Q1,P1:Q1),LMAG1(P1:Q1,P1:Q1),VMAG1(P1:Q1,P1:Q1)  
REAL SLOPE(P1:Q1,P1:Q1),LMAG2(P1:Q1,P1:Q1),VMAG2(P1:Q1,P1:Q1)  
REAL HMAG3(P1:Q1,P1:Q1)  
REAL D1KH1(1:R,2),D1KL1(1:R,2)  
REAL D1KH0(1:R,2),D1KL0(1:R,2)  
REAL D1KH3(1:R,2)  
*  
REAL SLP0(-90:90,2)  
*  
REAL D1PH1(-90:+90,2),D1PL1(-90:+90,2)  
REAL D1PH0(-90:+90,2),D1PL0(-90:+90,2)  
REAL D1PH3(-90:+90,2)  
REAL D1FH1(-90:+90,2),D1FL1(-90:+90,2)  
REAL D1FH0(-90:+90,2),D1FL0(-90:+90,2)  
REAL D1FH3(-90:+90,2)  
REAL D2H1(P:Q,P:Q),D2L1(P:Q,P:Q)  
REAL D2H0(P:Q,P:Q),D2L0(P:Q,P:Q)  
REAL D2H3(P:Q,P:Q)  
REAL DIFA1,DIFC1,DIFS1  
REAL DIFA2,DIFC2,DIFS2  
REAL DIFA3,DIFC3,DIFS3  
REAL DIFA4,DIFC4,DIFS4  
REAL DIFAS,DIFC5,DIFS5  
REAL DIFX1,DIFY1  
REAL DIFX2,DIFY2  
REAL DIFX3,DIFY3  
REAL DIFX4,DIFY4  
REAL DIFX5,DIFY5  
REAL VARA0,VARC0,VARS0  
REAL VARA1,VARC1,VARS1  
REAL VARA2,VARC2,VARS2  
REAL VARA3,VARC3,VARS3  
REAL VARA4,VARC4,VARS4
```

```

REAL VARA5,VARC5,VARS5
REAL VARX0,VARY0
REAL VARX1,VARY1
REAL VARX2,VARY2
REAL VARX3,VARY3
REAL VARX4,VARY4
REAL VARX5,VARY5
*
REAL VARXX
*
REAL DVRA1,DVRC1,DVRS1
REAL DVRA2,DVRC2,DVRS2
REAL DVRA3,DVRC3,DVRS3
REAL DVRA4,DVRC4,DVRS4
REAL DVRA5,DVRC5,DVRS5
COMPLEX GD1(P2:Q2,P2:Q2),GH1(P2:Q2,P2:Q2),GV1(P2:Q2,P2:Q2)
COMPLEX GD2(P2:Q2,P2:Q2),CH2(P2:Q2,P2:Q2),GV2(P2:Q2,P2:Q2)
COMPLEX SH0(P1:Q1,P1:Q1),SL0(P1:Q1,P1:Q1),SV0(P1:Q1,P1:Q1)
COMPLEX SH1(P1:Q1,P1:Q1),SL1(P1:Q1,P1:Q1),SV1(P1:Q1,P1:Q1)
COMPLEX SH2(P1:Q1,P1:Q1),SL2(P1:Q1,P1:Q1),SV2(P1:Q1,P1:Q1)
COMPLEX SH3(P1:Q1,P1:Q1)
CHARACTER*18 FNAME0,FNAME1,FNAME2,FNAME3,FNAME4,FNAME5,FNAME6
CHARACTER*18 FNAME7,FNAME8,FNAME
*
PI=3.141593
*
FNAME0='PARM.DAT'
FNAME1='HMAG1.RAD'
FNAME2='LMAG1.RAD'
FNAME3='HMAGO.RAD'
FNAME4='LMAGO.RAD'
FNAME5='HMAG3.RAD'
FNAME6='TSLOPE.RAE'
*
CALL PLT3DINPUT(FNAME0,PARM,NTMP,-2,1)
*
DO I=-2,1
    WRITE(*,*) (PARM(I,J),J=-2,1)
END DO
*
K_NYQ=PARM(-2,-2)
DELTA_K=PARM(-2,-1)
NTMP=INT(PARM(-2,0))
*
IF ((N1.NE.NTMP)..OR.((K_NYQ/DELTA_K).NE.FLOAT(NTMP/2))) THEN
    WRITE(*,*) 'INCORRECT PARAMETER FILE [N(SURF)]'
    GOTO 5
ENDIF
*
WIND_AZ=PARM(-1,0)
NTMP=INT(PARM(-1,1))
*

```

```

        IF (N.NE.NTMP) THEN
            WRITE(*,*) 'INCORRECT PARAMETER FILE (N1(DOME))'
            GOTO 5
        ENDIF
        *
        WRITE(*,*) 'READING ARRAYS.....'
        *
        CALL PLT3DINPUT(FNAME1, HMAG1, NTMP, P, Q)
        CALL PLT3DINPUT(FNAME2, LMAG1, NTMP, P, Q)
        CALL PLT3DINPUT(FNAME3, HMAGO, NTMP, P, Q)
        CALL PLT3DINPUT(FNAME4, LMAGO, NTMP, P, Q)
        CALL PLT3DINPUT(FNAME5, HMAG3, NTMP, P, Q)
        CALL PLT3DINPUT(FNAME6, SLOPE, NTMP, P, Q)
        *
        DAT1=HMAG1(0,0)
        DAT2=LMAG1(0,0)
        DAT3=HMAGO(0,0)
        DAT4=LMAGO(0,0)
        DAT5=HMAG3(1,0)
        *
        EMAG1(0,0)=0.0
        LMAG1(0,0)=0.0
        HMAGO(0,0)=0.0
        LMAGO(0,0)=0.0
        HMAG3(0,0)=0.0
        *
        *** WRITE(*,*) 'HMAG1'
        *** IF (N1.EQ.64) CALL ALPHA(HMAG1,64)
        *
        *** WRITE(*,*) 'LMAG1'
        *** IF (N1.EQ.64) CALL ALPHA(LMAG1,64)
        *
        *** WRITE(*,*) 'HMAGO'
        *** IF (N.EQ.64) CALL ALPHA(HMAGO,64)
        *
        *** WRITE(*,*) 'LMAGO'
        *** IF (N.EQ.64) CALL ALPHA(LMAGO,64)
        *
        *** WRITE(*,*) 'HMAG3'
        *** IF (N.EQ.64) CALL ALPHA(HMAG3,64)
        *
        *** WRITE(*,*) 'SLOPE'
        *** IF (N.EQ.64) CALL ALPHA(SLOPE,64)
        *
        WRITE(*,*) 'HMAG1'
        IF (N1.EQ.64) CALL ALPHAS(HMAG1,64,DATH1,LOW,FACTOR)
        *
        WRITE(*,*) 'LMAG1'
        IF (N1.EQ.64) CALL ALPHAS(LMAG1,64,DATL1,LOW,FACTOR)
        *
        WRITE(*,*) 'HMAGO'
        IF (N.EQ.64) CALL ALPHAS(HMAGO,64,DATH0,LOW,FACTOR)

```

```

*
      WRITE(*,*) 'LMAG0'
      IF (N.EQ.64) CALL ALPHAS(LMAG0,64,DATL0,LOW,FACTOR)
*
      WRITE(*,*) 'HMAG3'
      IF (N.EQ.64) CALL ALPHAS(HMAG3,64,DATH3,LOW,FACTOR)
*
      WRITE(*,*) 'SLOPE'
      IF (N.EQ.64) CALL ALPHAS(SLOPE,64,DATSO,LOW,FACTOR)
*
      FACTOR1=DATSO/DATH1
      FACTOR2=DATSO/DATL1
      FACTOR3=DATSO/DATH0
      FACTOR4=DATSO/DATL0
      FACTOR5=DATSO/DATH3
*
      DO I=P,Q
      DO J=P,Q
      HMAG1(I,J)=HMAG1(I,J)*FACTOR1
      LMAG1(I,J)=LMAG1(I,J)*FACTOR2
      HMAG0(I,J)=HMAG0(I,J)*FACTOR3
      LMAG0(I,J)=LMAG0(I,J)*FACTOR4
      HMAG3(I,J)=HMAG3(I,J)*FACTOR5
*
      D2H1(I,J)=SLOPE(I,J)-HMAG1(I,J)
      D2L1(I,J)=SLOPE(I,J)-LMAG1(I,J)
      D2H0(I,J)=SLOPE(I,J)-HMAG0(I,J)
      D2L0(I,J)=SLOPE(I,J)-LMAG0(I,J)
      D2H3(I,J)=SLOPE(I,J)-HMAG3(I,J)
      END DO
      END DO
*
      WRITE(*,*) '2D DIFFERENCE SPECTRA'
*
      WRITE(*,*) ' '
      WRITE(*,*) 'D2H1=SLOPE-HMAG1'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',DAT1
      WRITE(*,*) 'FACTOR1           = ',FACTOR1
***  IF (N.EQ.64) CALL ALPHA(D2H1,64)
***  CALL PLT3DFILE(D2H1,N)
*
      WRITE(*,*) ' '
      WRITE(*,*) 'D2L1=SLOPE-LMAG1'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',DAT2
      WRITE(*,*) 'FACTOR2           = ',FACTOR2
***  IF (N.EQ.64) CALL ALPHA(D2L1,64)
***  CALL PLT3DFILE(D2L1,N)
*
      WRITE(*,*) ' '
      WRITE(*,*) 'D2H0=SLOPE-HMAG0'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',DAT3
      WRITE(*,*) 'FACTOR3           = ',FACTOR3

```

```

***      IF (N.EQ.64) CALL ALPHA(D2H0,64)
***      CALL PLT3DFILE(D2H0,N)
*
      WRITE(*,*) ''
      WRITE(*,*) 'D2L0=SLOPE-LMAG0'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',DATA4
      WRITE(*,*) 'FACTOR4           = ',FACTOR4
      IF (N.EQ.64) CALL ALPHA(D2L0,64)
***      CALL PLT3DFILE(D2L0,N)
*
      WRITE(*,*) ''
      WRITE(*,*) 'D2H3=SLOPE-HMAG3'
      WRITE(*,*) 'MEAN PIXEL RADIANCE = ',DAT5
      WRITE(*,*) 'FACTORS            = ',FACTORS5
***      IF (N.EQ.64) CALL ALPHA(D2H3,64)
***      CALL PLT3DFILE(D2H3,N)
*
50      DO I=P,Q
         DO J=P,Q
*
         L=FLOAT(I)
         M=FLOAT(J)
         KMAG=SQRT((L**2)+(M**2))
         KINT=INT(KMAG+0.5)
         K=KMAG*DELTA_K
*
         IF ((K.EQ.0.0).OR.(KINT.GT.R)) GOTO 10
*
         ANG0=ATAN2D(M,L)
         ANG1=WIND_AZ-ANG0
*
         IF (ABS(ANG0).LE.90.0) THEN
             PHINT=INT(ANG0+SIGN(0.5,ANG0))
         ELSE
             PHINT=INT((ANG0-SIGN(180.0,ANG0))-SIGN(0.5,ANG0))
         ENDIF
*
         IF (ABS(PHINT).GT.90.0) WRITE(*,*) 'PHINT.GT.+/- 90.0'
*
         COSA1=COSD(ANG0)**2
         SINA1=SIND(ANG0)**2
         COSA2=COSD(ANG1)**2
         SINA2=SIND(ANG1)**2
*
         TEMP=(SLOPE(I,J)**2)*(DELTA_K**2)
         VARSO=VARSO+TEMP
         VARAO=VARAO+TEMP*COSA2
         VARC0=VARC0+TEMP*SINA2
         VARK0=VARK0+TEMP*COSA1
         VARY0=VARY0+TEMP*SINA1
*
         TEMP=(HMAG1(I,J)**2)*(DELTA_K**2)

```

```

TEMP1=(D2H1(I,J)**2)*(DELTA_K**2)
VARS1=VARS1+TEMP
VARA1=VARA1+TEMP*COSA2
VARC1=VARC1+TEMP*SINA2
VARX1=VARX1+TEMP*COSA1
VARY1=VARY1+TEMP*SINA1
DIFX1=DIFX1+TEMP1*COSA1
DIFY1=DIFY1+TEMP1*SINA1
DIFS1=DIFS1+TEMP1
DIFA1=DIFA1+TEMP1*COSA2
DIFC1=DIFC1+TEMP1*SINA2
D1KH1(KINT,2)=D1KH1(KINT,2)+(TEMP1/DELTA_K)
D1PH1(PHINT,2)=D1PH1(PHINT,2)+(TEMP1/DELTA_K)
D1PH1(PHINT,1)=D1PH1(PHINT,1)+1.0
*
TEMP=(LMAG1(I,J)**2)*(DELTA_K**2)
TEMP1=(D2L1(I,J)**2)*(DELTA_K**2)
VARS2=VARS2+TEMP
VARA2=VARA2+TEMP*COSA2
VARC2=VARC2+TEMP*SINA2
VARX2=VARX2+TEMP*COSA1
VARY2=VARY2+TEMP*SINA1
DIFX2=DIFX2+TEMP1*COSA1
DIFY2=DIFY2+TEMP1*SINA1
DIFS2=DIFS2+TEMP1
DIFA2=DIFA2+TEMP1*COSA2
DIFC2=DIFC2+TEMP1*SINA2
D1KL1(KINT,2)=D1KL1(KINT,2)+(TEMP1/DELTA_K)
D1PL1(PHINT,2)=D1PL1(PHINT,2)+(TEMP1/DELTA_K)
D1PL1(PHINT,1)=D1PL1(PHINT,1)+1.0
*
TEMP=(HMAGO(I,J)**2)*(DELTA_K**2)
TEMP1=(D2H0(I,J)**2)*(DELTA_K**2)
VARS3=VARS3+TEMP
VARA3=VARA3+TEMP*COSA2
VARC3=VARC3+TEMP*SINA2
VARX3=VARX3+TEMP*COSA1
VARY3=VARY3+TEMP*SINA1
DIFX3=DIFX3+TEMP1*COSA1
DIFY3=DIFY3+TEMP1*SINA1
DIFS3=DIFS3+TEMP1
DIFA3=DIFA3+TEMP1*COSA2
DIFC3=DIFC3+TEMP1*SINA2
D1KH0(KINT,2)=D1KH0(KINT,2)+(TEMP1/DELTA_K)
D1PH0(PHINT,2)=D1PH0(PHINT,2)+(TEMP1/DELTA_K)
D1PH0(PHINT,1)=D1PH0(PHINT,1)+1.0
*
TEMP=(LMAGO(I,J)**2)*(DELTA_K**2)
TEMP1=(D2L0(I,J)**2)*(DELTA_K**2)
VARS4=VARS4+TEMP
VARA4=VARA4+TEMP*COSA2
VARC4=VARC4+TEMP*SINA2

```

```

VARX4=VARX4+TEMP*COSA1
VARY4=VARY4+TEMP*SINA1
DIFX4=DIFX4+TEMP1*COSA1
DIFY4=DIFY4+TEMP1*SINA1
DIFS4=DIFS4+TEMP1
DIFA4=DIFA4+TEMP1*COSA2
DIFC4=DIFC4+TEMP1*SINA2
D1KL0(KINT,2)=D1KL0(KINT,2)+(TEMP1/DELTA_K)
D1PL0(PHINT,2)=D1PL0(PHINT,2)+(TEMP1/DELTA_K)
D1PL0(PHINT,1)=D1PL0(PHINT,1)+1.0
*
TEMP=(HMAG3(I,J)**2)*(DELTA_K**2)
TEMP1=(D2H3(I,J)**2)*(DELTA_K**2)
VARSS=VARSS+TEMP
VARA5=VARA5+TEMP*COSA2
VARC5=VARC5+TEMP*SINA2
VARX5=VARX5+TEMP*COSA1
VARY5=VARY5+TEMP*SINA1
DIFX5=DIFX5+TEMP1*COSA1
DIFY5=DIFY5+TEMP1*SINA1
DIFS5=DIFS5+TEMP1
DIFA5=DIFA5+TEMP1*COSA2
DIFC5=DIFC5+TEMP1*SINA2
D1KH3(KINT,2)=D1KH3(KINT,2)+(TEMP1/DELTA_K)
D1PH3(PHINT,2)=D1PH3(PHINT,2)+(TEMP1/DELTA_K)
D1PH3(PHINT,1)=D1PH3(PHINT,1)+1.0
*
10    END DO
      END DO
*
DO I=1,R
      D1KH1(I,1)=FLOAT(I)*DELTA_K
      D1KL1(I,1)=FLOAT(I)*DELTA_K
      D1KH0(I,1)=FLOAT(I)*DELTA_K
      D1KL0(I,1)=FLOAT(I)*DELTA_K
      D1KH3(I,1)=FLOAT(I)*DELTA_K
END DO
*
DO I=-90,+90
      D1FH1(I,1)=FLOAT(I)
      D1FL1(I,1)=FLOAT(I)
      D1FH0(I,1)=FLOAT(I)
      D1FL0(I,1)=FLOAT(I)
      D1FH3(I,1)=FLOAT(I)
*
      IF (D1PH1(I,1).GT.0.0) D1PH1(I,2)=D1PH1(I,1)/D1PH1(I,1)
      IF (D1PL1(I,1).GT.0.0) D1PL1(I,2)=D1PL1(I,1)/D1PL1(I,1)
      IF (D1PH0(I,1).GT.0.0) D1PH0(I,2)=D1PH0(I,1)/D1PH0(I,1)
      IF (D1PL0(I,1).GT.0.0) D1PL0(I,2)=D1PL0(I,1)/D1PL0(I,1)
      IF (D1PH3(I,1).GT.0.0) D1PH3(I,2)=D1PH3(I,1)/D1PH3(I,1)
*
      END DO

```

```

*
      DO I=-90,+90
*
      IF ((I.GT.-89).AND.(I.LT.89)) THEN
          D1FH1(I,2)=(D1PH1(I-1,2)+D1PH1(I,2)+D1PH1(I+1,2))/3.0
***       D1FL1(I,2)=(D1PL1(I-1,2)+D1PL1(I,2)+D1PL1(I+1,2))/3.0
***       D1FH0(I,2)=(D1PH0(I-1,2)+D1PH0(I,2)+D1PH0(I+1,2))/3.0
***       D1FL0(I,2)=(D1PL0(I-1,2)+D1PL0(I,2)+D1PL0(I+1,2))/3.0
***       D1FH3(I,2)=(D1PH3(I-1,2)+D1PH3(I,2)+D1PH3(I+1,2))/3.0
      ELSEIF (I.EQ.-90) THEN
          D1FH1(I,2)=(D1PH1(+89,2)+D1PH1(I,2)+D1PH1(I+1,2))/3.0
***       D1FL1(I,2)=(D1PL1(+89,2)+D1PL1(I,2)+D1PL1(I+1,2))/3.0
***       D1FH0(I,2)=(D1PH0(+89,2)+D1PH0(I,2)+D1PH0(I+1,2))/3.0
***       D1FL0(I,2)=(D1PL0(+89,2)+D1PL0(I,2)+D1PL0(I+1,2))/3.0
***       D1FH3(I,2)=(D1PH3(+89,2)+D1PH3(I,2)+D1PH3(I+1,2))/3.0
      ELSEIF (I.EQ.+90) THEN
          D1FH1(I,2)=(D1PH1(I-1,2)+D1PH1(I,2)+D1PH1(-89,2))/3.0
***       D1FL1(I,2)=(D1PL1(I-1,2)+D1PL1(I,2)+D1PL1(-89,2))/3.0
***       D1FH0(I,2)=(D1PH0(I-1,2)+D1PH0(I,2)+D1PH0(-89,2))/3.0
***       D1FL0(I,2)=(D1PL0(I-1,2)+D1PL0(I,2)+D1PL0(-89,2))/3.0
***       D1FH3(I,2)=(D1PH3(I-1,2)+D1PH3(I,2)+D1PH3(-89,2))/3.0
      ENDIF
*
      END DO
*
      WRITE(*,*) '           PARAMETERS'
      WRITE(*,*) ',PARM(-2,-2),PARM(-2,-1),PARM(-2,0),PARM(-2,1)
*
      WRITE(*,*) ',PARM(-1,-2),PARM(-1,-1),PARM(-1,0),PARM(-1,1)
*
      WRITE(*,*) ',PARM(+0,-2),PARM(+0,-1),PARM(+0,0),PARM(+0,1)
*
      WRITE(*,*) ',PARM(+1,-2),PARM(+1,-1),PARM(+1,0),PARM(+1,1)
      WRITE(*,*) '
*
      WRITE(*,*) '           MAGNITUDE ARRAY SCALING FACTORS'
      WRITE(*,*) 'H1(0,0),HIGH,FACTOR',DAT1,DATH1,FACTOR1
      WRITE(*,*) 'L1(0,0),HIGH,FACTOR',DAT2,DATL1,FACTOR2
      WRITE(*,*) 'H0(0,0),HIGH,FACTOR',DAT3,DATH0,FACTOR3
      WRITE(*,*) 'L0(0,0),HIGH,FACTOR',DAT4,DATL0,FACTOR4
      WRITE(*,*) 'H3(0,0),HIGH,FACTOR',DAT5,DATH3,FACTOR5
      WRITE(*,*) '
*
      WRITE(*,*) '           INTEGRATED VARIANCE FROM POWER SPECTRA'
      WRITE(*,*) 'VAR S0 A,C,S',VARA0,VARC0,VARS0
      WRITE(*,*) 'VAR H1 A,C,S',VARA1,VARC1,VARS1
      WRITE(*,*) 'VAR L1 A,C,S',VARA2,VARC2,VARS2
      WRITE(*,*) 'VAR H0 A,C,S',VARA3,VARC3,VARS3
      WRITE(*,*) 'VAR L0 A,C,S',VARA4,VARC4,VARS4
      WRITE(*,*) 'VAR H3 A,C,S',VARA5,VARC5,VARS5
      WRITE(*,*) '

```



```

        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        *
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        *
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        *
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        WRITE(1,*)
        *
        CLOSE(1)
        *
        FNAME1='K1.DT'
        FNAME2='K2.DT'
        FNAME3='K3.DT'
        FNAME4='K4.DT'
        FNAME5='K5.DT'
        *
        * CALL PLT2DOUTPUT(FNAME1,D1KH1,R)
        *** CALL PLT2DOUTPUT(FNAME2,D1KL1,R)
        *** CALL PLT2DOUTPUT(FNAME3,D1KHO,R)
        * CALL PLT2DOUTPUT(FNAME4,D1KLO,R)
        * CALL PLT2DOUTPUT(FNAME5,D1KH3,R)
        *
        FNAME1='P1.DT'
        FNAME2='P2.DT'
        FNAME3='P3.DT'
        FNAME4='P4.DT'
        FNAME5='P5.DT'
        *
        * CALL PLT2DOUTPUT(FNAME1,D1FH1,181)
        * CALL PLT2DOUTPUT(FNAME2,D1FL1,181)
        * CALL PLT2DOUTPUT(FNAME3,D1FH0,181)

```

```
CALL PLT2DOUTPUT(FNAME4,D1FL0,181)
CALL PLT2DOUTPUT(FNAME5,D1FH3,181)
*
GOTO 1000
5    WRITE(*,*) 'BAD FILE OPENING'
300    FORMAT (A)
1000   END
```

```

PROGRAM PSVARY
IMPLICIT NONE
INTEGER I,J,N,P,Q
REAL A,ALPH,BETA,D,DELT,DELTA_K,ENUG,G,K,K_MAX,KNU,K1,K2,K3
REAL L,M,P2,PI,SPEC
REAL VAR,VAR1,VAR2,VAR3,VAR4,VARS,VAR6,VAR7,VAR8,VAR9,VAR10,VAR11
REAL VAR12,VAR13,VAR14,VAR15,VAR16
REAL V_FRIC,V_MIN,V250,V1000,V1250,V1950,WIND_V,Z
REAL F1(-800:400,2),F2(-800:400,2),F3(-800:400,2),F4(-800:400,2)
REAL F5(-800:400,2)

*
A=1.0
ALPH=0.0081
* ALPH=PHILLIPS CONSTANT [ ]
BETA=0.74
* BETA= [UNITLESS]
ENUG=1.473E-4
* ENUG=E/(NU*G) [UNITLESS]
G=980.0
* G=ACCELERATION OF GRAVITY [CM/SEC^2]
K2=0.359
* [1/CM]
K3=0.942
* [1/CM]
K_MAX=3.63
* [1/CM]
PI=3.141593
*
V_MIN=12.0
* V_MIN=MINIMUM FRICTION VELOCITY [CM/SEC]
*
      WRITE(*,*) 'ENTER WIND_V [CM/SEC^2] & Z [CM] '
      READ(*,*) WIND_V,Z
*
* WIND_V=WIND VELOCITY AT HEIGHT Z [CM/SEC^2]
* Z=HEIGHT OF WIND_V MEASUREMENT ABOVE WATER SURFACE [CM]
*
      CALL WIND(WIND_V,Z,V_FRIC,V250,V1000,V1250,V1950)
*
D=(1.274+(0.0268*V_FRIC)+(6.03E-5*(V_FRIC**2)))**2
K1=(K2*(V_MIN**2))/(V_FRIC**2)
KNU=(0.5756*SQRT(V_FRIC)*K_MAX)/(D**0.16667)
P2=LOG10(D/(V_FRIC/V_MIN))/LOG10(K3/K2)
*
DELT=0.010
VAR=0.0
*
DO I=-800,400
L=FLOAT(I-200)*DELT
K=EXP(L)
F1(I,1)=K
F2(I,1)=LOG10(K)

```

```

F3(I,1)=LOG10(K)
F4(I,1)=LOG10(K)
F5(I,1)=K
*
IF ((0.0.LT.K).AND.(K.LE.K1)) THEN
SPEC=(ALPH/(2.0*(K**3)))*EXP((-1.0*BETA*(G**2))/((V1950**4)*(K**2)))
GOTO 1
ENDIF
*
IF ((K1.LT.K).AND.(K.LE.K2)) THEN
SPEC=ALPH/(2.0*SQRT(K1)*(K**2.5))
GOTO 1
ENDIF
*
IF ((K2.LT.K).AND.(K.LE.K3)) THEN
SPEC=(ALPH*D)/(2.0*(K3**P2)*(K***(3.0-P2)))
GOTO 1
ENDIF
*
IF ((K3.LT.K).AND.(K.LE.KNU)) THEN
SPEC=(ALPH*D)/(2.0*(K**3))
GOTO 1
ENDIF
*
IF (KNU.LT.K) THEN
SPEC=(ENUG*(V_FRIC**3)*(K_MAX**6))/(K**9)
ENDIF
*
1 F1(I,2)=(K**2)*SPEC
F5(I,2)=SQRT((K**2)*SPEC)
F2(I,2)=(K**3)*SPEC
F3(I,2)=(K*SPEC)
F4(I,2)=(K**5)*SPEC
VAR=VAR+(F2(I,2)*DELT)
*
*** WRITE(*,*) L,X,VAR
*
IF (EXP(L-DELT).LE.0.0030000) VAR1=VAR
IF (EXP(L-DELT).LE.0.0046875) VAR2=VAR
IF (EXP(L-DELT).LE.0.0093750) VAR3=VAR
IF (EXP(L-DELT).LE.0.0187500) VAR4=VAR
IF (EXP(L-DELT).LE.0.0375000) VAR5=VAR
IF (EXP(L-DELT).LE.0.0750000) VAR6=VAR
IF (EXP(L-DELT).LE.0.1500000) VAR7=VAR
IF (EXP(L-DELT).LE.0.3000000) VAR8=VAR
IF (EXP(L-DELT).LE.0.6000000) VAR9=VAR
IF (EXP(L-DELT).LE.1.2000000) VAR10=VAR
IF (EXP(L-DELT).LE.2.4000000) VAR11=VAR
IF (EXP(L-DELT).LE.4.8000000) VAR12=VAR
IF (EXP(L-DELT).LE.9.6000000) VAR13=VAR
IF (EXP(L-DELT).LE.19.200000) VAR14=VAR
IF (EXP(L-DELT).LE.38.200000) VAR15=VAR

```

```

IF (EXP(L-DELT).LE.300.00000) VAR16=VAR
*
END DO
*
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.0030000/CM = ',VAR1
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.0046875/CM = ',VAR2
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.0093750/CM = ',VAR3
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.0187500/CM = ',VAR4
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.0375000/CM = ',VAR5
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.0750000/CM = ',VAR6
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.1500000/CM = ',VAR7,VAR7-VAR2
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.3000000/CM = ',VAR8,VAR8-VAR3
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 0.6000000/CM = ',VAR9,VAR9-VAR4
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 1.2000000/CM = ',VAR10,VAR10-VAR5
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 2.4000000/CM = ',VAR11,VAR11-VAR6
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 4.8000000/CM = ',VAR12,VAR12-VAR7
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 9.6000000/CM = ',VAR13,VAR13-VAR8
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 19.200000/CM = ',VAR14,VAR14-VAR9
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 38.400000/CM = ',VAR15,VAR15-VAR10
WRITE(*,*) 'CUM SLOPE VARIANCE UP TO 300.00000/CM = ',VAR16
WRITE(*,*) 'TOTAL SLOPE VARIANCE = ',VAR
*
WRITE(*,*) 'K^2*SPEC VS K'
CALL PLT2DFILE(F1,1201)
*
WRITE(*,*) 'K^2*SPEC VS K'
CALL PLT2DFILE(F5,1201)
*
WRITE(*,*) 'K^3*SPEC VS LOG10(K)'
CALL PLT2DFILE(F2,1201)
*
WRITE(*,*) 'K^1*SPEC VS LOG10(K)'
CALL PLT2DFILE(F3,1201)
*
WRITE(*,*) 'K^5*SPEC VS LOG10(K)'
CALL PLT2DFILE(F4,1201)
*
END

```

FORTRAN Subroutines

```
SUBROUTINE ALPHA(MAT1,N)
IMPLICIT NONE
INTEGER I,J,K,L,M,N
REAL MAT1(N,64),LOW,HIGH,FACTOR,Q
CHARACTER*1 CHOICE,LTR
CHARACTER*64 A,B,C,D,TEMP,PICS(208)

*
WRITE(*,*) 'WANT AN ALPHASCALE IMAGE? [N]'
READ(*,300) CHOICE
IF (CHOICE.NE.'Y') GOTO 2
*
LOW=MAT1(1,1)
HIGH=MAT1(1,1)
DO I=1,N
    DO J=1,64
        IF (MAT1(I,J).LT.LOW) LOW=MAT1(I,J)
        IF (MAT1(I,J).GT.HIGH) HIGH=MAT1(I,J)
    END DO
END DO
*
IF (HIGH.LT.0.0) THEN
HIGH=0.0
FACTOR=LOW/-26.5
GOTO 1
ENDIF
*
IF (LOW.GT.0.0) THEN
LOW=0.0
FACTOR=HIGH/26.5
GOTO 1
ENDIF
*
IF (HIGH.GE.-1.0*LOW) THEN
FACTOR=HIGH/26.5
ELSE
FACTOR=LOW/-26.5
ENDIF
*
IF (FACTOR.EQ.0.0) FACTOR=1.0
*
1 WRITE(*,*) 'ALPHASCALE:HIGH,LOW,FACTOR ..',HIGH,LOW,FACTOR
*
DO I=N,1,-1
M=I-N/2-1
TEMP=' '
    DO J=1,64
K=INT((MAT1(I,J))/FACTOR)
        IF (K.GT.0) THEN
            L=K+64
            LTR=CHAR(L)
```

```

        ELSEIF (K.LT.0) THEN
            L=-1*K+96
            LTR=CHAR(L)
        ELSE
            LTR=CHAR(32)
        ENDIF
        TEMP=TEMP(1:J-1)//LTR
        END DO
        WRITE(*,*) TEMP,' ',M
        PICS(N-I+1)=TEMP
        END DO
*
A=' '
B='----- ++++++++'
C='33322222222211111111000000000000000001111111122222222233'
D='2109876543210987654321098765432101234567890123456789012345678901'

        WRITE(*,*) A
        WRITE(*,*) B
        WRITE(*,*) C
        WRITE(*,*) D
*
PICS(N+1)=A
PICS(N+2)=B
PICS(N+3)=C
PICS(N+4)=D
*
CALL PICFILE(PICS,N,HIGH,LOW,FACTOR)
*
2      RETURN
300    FORMAT (A)
      END

SUBROUTINE ALPHAS(MAT1,N,HIGH,LOW,FACTOR)
IMPLICIT NONE
INTEGER I,J,K,L,M,N
REAL MAT1(N,64),LOW,HIGH,FACTOR,Q
CHARACTER*1 CHOICE,LTR
CHARACTER*64 A,B,C,D,TEMP,PICS(208)
*
LOW=MAT1(1,1)
HIGH=MAT1(1,1)
DO I=1,N
    DO J=1,64
        IF (MAT1(I,J).LT.LOW) LOW=MAT1(I,J)
        IF (MAT1(I,J).GT.HIGH) HIGH=MAT1(I,J)
    END DO
END DO
*
IF (HIGH.LT.0.0) THEN
    HIGH=0.0

```

```

FACTOR=LOW/-26.5
GOTO 1
ENDIF
*
IF (LOW.GT.0.0) THEN
LOW=0.0
FACTOR=HIGH/26.5
GOTO 1
ENDIF
*
IF (HIGH.GE.-1.0*LOW) THEN
FACTOR=HIGH/26.5
ELSE
FACTOR=LOW/-26.5
ENDIF
*
IF (FACTOR.EQ.0.0) FACTOR=1.0
*
*
1 WRITE(*,*) 'WANT AN ALPHASCALE IMAGE? [N]'
READ(*,300) CHOICE
IF (CHOICE.NE.'Y') GOTO 2
WRITE(*,*) 'ALPHASCALE:HIGH,LOW,FACTOR =',HIGH,LOW,FACTOR
*
DO I=N,1,-1
M=I-N/2-1
TEMP=' '
DO J=1,64
K=INT((MAT1(I,J))/FACTOR)
IF (K.GT.0) THEN
L=K+64
LTR=CHAR(L)
ELSEIF (K.LT.0) THEN
L=-1*K+96
LTR=CHAR(L)
ELSE
LTR=CHAR(32)
ENDIF
TEMP=TEMP(1:J-1)||LTR
END DO
WRITE(*,*) TEMP,' ',M
PIGS(N-I+1)=TEMP
END DO
*
A=
B='----- ++++++++
C='333222222222111111111000000000000001111111112222222233'
D='210987654321098765432109876543210123456789012345678901'

WRITE(*,*) A
WRITE(*,*) B
WRITE(*,*) C

```

```

        WRITE(*,*) D
*
        PICS(N+1)=A
        PICS(N+2)=B
        PICS(N+3)=C
        PICS(N+4)=D
*
        CALL PICFILE(PICS,N,HIGH,LOW,FACTOR)
*
2      RETURN
300    FORMAT (A)
      END

      SUBROUTINE CHECKERBOARD(X,N)
      IMPLICIT NONE
      INTEGER I,J,N
      COMPLEX X(1:N,1:N)
      DO I=1,N
        DO J=1,N
          X(I,J)=X(I,J)*(-1.0)**((I-1)+(J-1))
        END DO
      END DO
      RETURN
      END

      SUBROUTINE DARKNESS(P,Q,THETA,INDEXR,INDEXI,GAMH,GAMV,DARK)
*
*   * CREATES POLARIZED (HORZ & VERT) FRESNEL_COEFFICIENT DISTRIBUTIONS
*   * RELATIVE TO SENSOR AND SURFACE COORDINATES
*
*   * CREATES SURFACE_PROJECTION_ATTENUATION_COEFFICIENT DISTRIBUTION
*   * RELATIVE TO SENSOR AND SURFACE COORDINATES
*
      IMPLICIT NONE
      INTEGER I,J,N,P,Q
      REAL I1,I2,I3,N1,N2,N3,R1,R2,R3
      REAL ALPH,COSALPHA,SINALPHA
      REAL BETA,COSBETA,SINBETA,SECBETA
      REAL DELTA,GAMMAHOR,GAMMAPAR,GAMMAPER,GAMMASUM,GAMMAVER
      REAL DARK(P:Q,P:Q),GAMH(P:Q,P:Q),GAMV(P:Q,P:Q)
      REAL DRK1(-32:31,-32:31),DRK2(-32:31,-32:31),GAMS(-32:31,-32:31)
      REAL L,M,MU,COSMU,SINMU
      REAL NU,COSNU,SINNU
      REAL OMEGA,COSOMEGA,COSOMEGA2,SINOMEGA
      REAL INDEXI,INDEXR,RHO,RHO1
      REAL PI,PROJECT,RO,THETA
      REAL PHI,PHI1,PHI2,RADIUS
*
      DELTA=90.0/FLOAT(Q)
      N=Q-P+1

```

```

PI=3.141593
R0=(INDEXR-1.0)**2/(INDEXR+1.0)**2
*
WRITE(*,*)
WRITE(*,*) 'THETA, INDEXR, INDEXI, R0= ', THETA, INDEXR, INDEXI, R0
*
* I1,I2,I3 ARE THE RESOLVED COMPONENTS OF THE SENSOR COORDINATE VECTOR
*
I1=SIND(THETA)*COSD(180.0)
I2=SIND(THETA)*SIND(180.0)
I3=COSD(THETA)*(+1.0)
*
DO I=P,Q
DO J=P,Q
*
L=INT(I)*DELTA
M=INT(J)*DELTA
MU=SQRT((L**2)+(M**2))
*
IF (MU.GT.FLOAT(Q)*DELTA) GOTO 1
*
IF (MU.GT.0.0) THEN
    NU=ATAN2D(M,L)
ELSE
    NU=0.0
ENDIF
*
* N1,N2,N3 ARE THE RESOLVED COMPONENTS OF THE SURFACE NORMAL VECTOR
*
N1=SIND(MU)*COSD(NU)
N2=SIND(MU)*SIND(NU)
N3=COSD(MU)*(+1.0)
*
* DOT PRODUCT OF I & N = |I||N|(COSOMEGA) = COSOMEGA
*
COSOMEGA=(I1*N1)+(I2*N2)+(I3*N3)
OMEGA=ACOSD(COSOMEGA)
SINOMEGA=SIND(OMEGA)
*
IF (COSOMEGA.LE.0.0) GOTO 1
*
BETA=MU
*
IF (BETA.GT.60.0) BETA=60.0
*** IF (BETA.LT.90.0) THEN
    SEC_BETA=1.0/(COSD(BETA))
*** ELSE
***     SEC_BETA=0.0
*** ENDIF
*
PROJECT=COSOMEGA*SEC_BETA
DARK(I,J)=PROJECT

```

```

*
      IF (SINOMEGA.EQ.0.0) THEN
          GAMMAHOR=R0
          GAMMAVER=R0
          GOTO 3
      ENDIF
*
      RHO=ASIND(SINOMEGA/INDEXR)
      GAMMAHOR=(SIND(OMEGA-RHO)**2)/(SIND(OMEGA+RHO)**2)
      GAMMAVER=(TAND(OMEGA-RHO)**2)/(TAND(OMEGA+RHO)**2)
*
      3  GAMMASUM=0.5*(GAMMAHOR+GAMMAVER)
*
      GAMH(I,J)=GAMMAHOR
      GAMV(I,J)=GAMMAVER
*
      IF (N.EQ.64) THEN
          GAMS(I,J)=GAMMASUM
          DRK1(I,J)=SECBETA
          DRK2(I,J)=COSOMEGA
      ENDIF
*
      1  END DO
      END DO
*
      IF (N.EQ.64) THEN
          WRITE(*,*) 'GAMH= HORZ_POLARIZED FRESNEL COEFF DISTRIBUTION'
          *** CALL ALPHA(GAMH,64)
          *** CALL PLT3DFILE(GAMH,64)
*
          WRITE(*,*) 'GAMV= VERT_POLARIZED FRESNEL COEFF DISTRIBUTION'
          *** CALL ALPHA(GAMV,64)
          *** CALL PLT3DFILE(GAMV,64)
*
          WRITE(*,*) 'GAMS= UNPOLARIZED FRESNEL COEFF DISTRIBUTION'
          *** CALL ALPHA(GAMS,64)
          *** CALL PLT3DFILE(GAMS,64)
*
          WRITE(*,*) 'DRK1= ACTUAL AREA OF SLOPING SURFACE'
          *** CALL ALPHA(DRK1,64)
          *** CALL PLT3DFILE(DRK1,64)
*
          WRITE(*,*) 'DRK2= PROJECTION OF AREA NORMAL TO INCIDENCE'
          *** CALL ALPHA(DRK2,64)
          *** CALL PLT3DFILE(DRK2,64)
*
          WRITE(*,*) 'DARK= DRK1*DRK2= SECBETA*COSOMEGA'
          *** CALL ALPHA(DARK,64)
          *** CALL PLT3DFILE(DARK,64)
      ENDIF
*
      RETURN

```

```

5      WRITE(*,*) 'BAD FILE OPENING'
300      FORMAT (A)
      END

      SUBROUTINE DEEPBLUE(LREF,Z0,N,P,Q,LSUB,HSUB,VSUB)

      * CREATE UNPOLARIZED SUBDOME HEMISPHERICAL RADIANCE DISTRIBUTION
      * RESOLVE HORIZONTALLY & VERTICALLY POLARIZED RADIANCE COMPONENTS
      * BASED UPON THE MONTE CARLO RESULTS OF PLASS, KATTAWAR & GUINN [1976]
      *
      * INPUT:
      * LREF = REFERENCE RADIANCE MEASURED AT THE ZENITH POINT, LSKY(0,0) [UNIT]
      * Z0 = SOLAR ZENITH ANGLE
      *
      * OUTPUT:
      * LSUB = UNPOLARIZED SUBDOME HEMISPHERICAL RADIANCE DISTRIBUTION
      * HSUB = HORIZONTALLY_POLARIZED LSUB
      * VSUB = VERTICALLY_POLARIZED LSUB
      *
      IMPLICIT NONE
      INTEGER I,J,N,P,Q
      REAL LSUB(P:Q,P:Q),HSUB(P:Q,P:Q),VSUB(P:Q,P:Q)
      REAL A,B,C,COSMU,COSMU2,COSNU,L,LREF,M,MU,PHI,PI,PSI,R0,THETA,Z0
      CHARACTER*18 FNAME
      *
      PI=3.14159
      *
      DO I=P+1,Q
      DO J=P+1,Q
      *
      L=FLOAT(I)
      M=FLOAT(J)
      THETA=(PI/2.0)*(SQRT((L**2)+(M**2))/FLOAT(Q))
      IF (THETA.GT.PI/2.0) THEN
          LSUB(I,J)=0.0
          HSUB(I,J)=0.0
          VSUB(I,J)=0.0
          GOTO 4
      ENDIF
      *
      IF (THETA.GT.0.0) THEN
          PHI=ATAN2D(M,L)
      ELSE
          PHI=0.0
      ENDIF
      *
      A=COS(Z0)**2
      B=(SIN(1.1*Z0))**SQRT(3.0)
      C=0.32*(10.0**B)
      R0=0.02
      LSUB(I,J)=LREF*R0*C*((1.0+(COS(A*THETA)**2))/2.0)

```

```

*
* CALCULATE HORIZONTALLY_ & VERTICALLY_POLARIZED FRACTIONS
* VIA RANDOM SCATTERING MODEL
*
      HSUB(I,J)=LSUB(I,J)*0.5
      VSUB(I,J)=HSUB(I,J)
*
4      END DO
      END DO
*
***      WRITE(*,*) 'LSUB = UNPOLARIZED SUBDOME RADIANCE DISTRIBUTION'
***      IF (N.EQ.64) CALL ALPHA(LSUB,64)
      FNAME='LSUB.RAY'
***      CALL PLT3DOUTPUT(FNAME,LSUB,N,P,Q)
***      CALL PLT3DFILE(LSUB,N)
*
      WRITE(*,*) 'HSUB = HORIZONTALLY_POLARIZED LSUB'
      IF (N.EQ.64) CALL ALPHA(HSUB,64)
      FNAME='HSUB.RAY'
      CALL PLT3DOUTPUT(FNAME,HSUB,N,P,Q)
***      CALL PLT3DFILE(HSUB,N)
*
      WRITE(*,*) 'VSUB = VERTICALLY_POLARIZED LSUB'
      IF (N.EQ.64) CALL ALPHA(VSUB,64)
      FNAME='VSUB.RAY'
      CALL PLT3DOUTPUT(FNAME,VSUB,N,P,Q)
***      CALL PLT3DFILE(VSUB,N)
*
      RETURN
END

```

SUBROUTINE DIST\_GEN(OPT,N1,P,Q,WIND\_AZ,V1250,DEVA,DEVc,D1,D2,D3,D4)

```

*
* CREATES SURFACE SLOPE PROBABILITY DISTRIBUTIONS
*
* OUTPUTS:
* D1      = UNNORMALIZED GAUSSIAN DISTRIBUTION      (MIN PROB < 1/N1^2)
* D2      = UNNORMALIZED GRAM-CHARLIER DISTRIBUTION (MIN PROB < 1/N1^2)
* D3      = NORMALIZED GAUSSIAN DISTRIBUTION       (MIN PROB = 1/N1^2)
* D4      = NORMALIZED GRAM-CHARLIER DISTRIBUTION (MIN PROB = 1/N1^2)
*
* INPUTS:
* OPT     = SELECT 1 OF 2 DISTRIBUTION TYPES:
*          1 = SLOPE ANGLE [DEG]
*          2 = SLOPE ANGLE [TAND(DEG)]
* N1      = SPATIAL ARRAY SIZE
* P & Q   = DISTRIBUTION ARRAY DIMENSIONS
* WIND_AZ = AZIMUTH OF WIND DIRECTION
* V1250   = WIND VELOCITY AT 1250 CM ABOVE SURFACE
* DEVA    = ALONG-WIND SLOPE DEVIATION [TAND(DEG)]
* DEVc    = CROSS-WIND SLOPE DEVIATION [TAND(DEG)]

```

```

*
      IMPLICIT NONE
      INTEGER I,J,N1,OPT,P,Q
      REAL ANG,ANG1,DELTA,DEVA,DEVC,C03,C04,C21,C22,C40,GC1,GC2,GC3,GC4 .
      REAL GRAMCHAR,L,L1,M,M1,PI,RAD,SUM1,SUM2,SUM3,SUM4,VAR1,VAR2,VAR3
      REAL VAR4,V1250,WGHT0,WGHT1,WGHT2,WGHT3,WGHT4,WIND_AZ
      REAL D1(P:Q,P:Q),D2(P:Q,P:Q),D3(P:Q,P:Q),D4(P:Q,P:Q)
*
* GRAM-CHARLIER SERIES COEFFICIENTS FROM COX-MUNK STUDY
*
      C03=0.04-0.00033*V1250
      C04=0.23
      C21=0.61-0.000086*V1250
      C22=0.12
      C40=0.40
*
      PI=3.141593
      WIND_AZ=WIND_AZ*(PI/180.0)
*
      IF (OPT.EQ.1) DELTA=90.0/FLOAT(Q)
      IF (OPT.EQ.2) DELTA=1.0/FLOAT(Q)
*
      DO I=P+1,Q
      DO J=P+1,Q
      L=FLOAT(I)*DELTA
      M=FLOAT(J)*DELTA
*
      RAD=SQRT((L**2)+(M**2))
*
      IF (RAD.GE.(FLOAT(Q)*DELTA)) GOTO 1
*
      IF ((M.EQ.0.0).AND.(L.EQ.0.0)) THEN
          L1=0.0
          M1=0.0
          GOTO 2
      ENDIF
*
      ANG=ATAN2(M,L)
      ANG1=WIND_AZ-ANG
*
      IF (OPT.EQ.1) THEN
          L1=(TAND(RAD)/DEVA)*COS(ANG1)
          M1=(TAND(RAD)/DEVC)*SIN(ANG1)
      ENDIF
*
      IF (OPT.EQ.2) THEN
          L1=(RAD/DEVA)*COS(ANG1)
          M1=(RAD/DEVC)*SIN(ANG1)
      ENDIF
*
* CALCULATE GRAM-CHARLIER EXPANSION
*

```

```

2      GC1=1.0-(0.5*C21*((M1**2)-1.0)*L1)-(C03*((L1**3)-3.0*L1)/6.0)
GC2=C40*((L1**4)-6.0*(L1**2)+3.0)/24.0
GC3=C22*((L1**2)-1.0)*((M1**2)-1.0)/4.0
GC4=C04*((M1**4)-6.0*(M1**2)+3.0)/24.0
GRAMCHAR=GC1+GC2+GC3+GC4
*
D1(I,J)=EXP(-0.5*((L1**2)+(M1**2)))/(2.0*PI*DEVA*DEVC)
D2(I,J)=GRAMCHAR*D1(I,J)
SUM1=SUM1+D1(I,J)
SUM2=SUM2+D2(I,J)
*
1      END DO
END DO
*
* BIAS ALL ZERO VALUES IN DIST1 & DIST2 FOR ALPHASCALE DISPLAY
* CALCULATE VARIANCES OF DIST1 & DIST2
* NORMALIZE DIST1 & DIST2
*
WGHT0=1.0/(FLOAT(N1)**2)
WGHT1=0.0
WGHT2=0.0
*** WGHT1=-0.005
*** WGHT2=-0.005
*
DO I=P+1,Q
DO J=P+1,Q
*
L=FLOAT(I)*DELTA
M=FLOAT(J)*DELTA
RAD=SQRT((L**2)+(M**2))
*
IF (RAD.GE.(FLOAT(Q)*DELTA)) GOTO 3
*
D1(I,J)=D1(I,J)/SUM1
D2(I,J)=D2(I,J)/SUM2
*
VAR1=VAR1+D1(I,J)*(RAD**2)
VAR2=VAR2+D2(I,J)*(RAD**2)
*
IF (D1(I,J).EQ.0.0) D1(I,J)=WGHT1
*
IF (D1(I,J).LT.WGHT0) THEN
    D3(I,J)=WGHT1
ELSE
    D3(I,J)=D1(I,J)
    SUM3=SUM3+D3(I,J)
ENDIF
*
IF (D2(I,J).EQ.0.0) D2(I,J)=WGHT2
*
IF (D2(I,J).LT.WGHT0) THEN
    D4(I,J)=WGHT2

```

```

        ELSE
            D4(I,J)=D2(I,J)
            SUM4=SUM4+D4(I,J)
        ENDIF
    *
    3    END DO
    END DO
    *
    * BIAS ALL ZERO VALUES IN DIST3 & DIST4 FOR ALPHASCALE DISPLAY
    * CALCULATE VARIANCES OF DIST3 & DIST4
    * NORMALIZE DIST3 & DIST4
    *
    WGHT3=0.0
    WGHT4=0.0
    *** WGHT3=-0.005
    *** WGHT4=-0.005
    *
    DO I=P+1,Q
    DO J=P+1,Q
    *
    L=FLOAT(I)*DELTA
    M=FLOAT(J)*DELTA
    RAD=SQRT((L**2)+(M**2))
    *
    IF (RAD.GE.(FLOAT(Q)*DELTA)) GOTO 4
    *
    IF (D3(I,J).GE.WGHT0) THEN
        D3(I,J)=D3(I,J)/SUM3
        VAR3=VAR3+D3(I,J)*(RAD**2)
    ELSE
        D3(I,J)=WGHT3
    ENDIF
    *
    IF (D4(I,J).GE.WGHT0) THEN
        D4(I,J)=D4(I,J)/SUM4
        VAR4=VAR4+D4(I,J)*(RAD**2)
    ELSE
        D4(I,J)=WGHT4
    ENDIF
    *
    4    END DO
    END DO
    *
    WRITE(*,*) ''
    WRITE(*,*) 'SUM1,SUM2= ',SUM1,SUM2
    *
    IF (OPT.EQ.1) THEN
        WRITE(*,*) 'DEV1,DEV2= '.TAND(SQRT(VAR1)),TAND(SQRT(VAR2))
    ELSE
        WRITE(*,*) 'DEV1,DEV2= ',SQRT(VAR1),SQRT(VAR2)
    ENDIF
    *

```

```

      WRITE(*,*) 'SUM3,SUM4= ',SUM3,SUM4
      *
      IF (OPT.EQ.1) THEN
          WRITE(*,*) 'DEV3,DEV4= ',TAND(SQRT(VAR3)),TAND(SQRT(VAR4))
      ELSE
          WRITE(*,*) 'DEV3,DEV4= ',SQRT(VAR3),SQRT(VAR4)
      ENDIF
      *
      WRITE(*,*) ''
      *
      *** WRITE(*,*) 'UNNORMALIZED GAUSSIAN SLOPE DISTRIBUTION ARRAY'
      *** IF (N1.EQ.64) CALL ALPHA(D1,64)
      *** CALL PLT3DFILE(D1,N1)
      *
      *** WRITE(*,*) 'NORMALIZED GAUSSIAN SLOPE DISTRIBUTION ARRAY'
      *** IF (N1.EQ.64) CALL ALPHA(D3,64)
      *** CALL PLT3DFILE(D3,N1)
      *
      *** WRITE(*,*) 'UNNORMALIZED GRAM-CHARLIER SLOPE DISTRIBUTION ARRAY'
      *** IF (N1.EQ.64) CALL ALPHA(D2,64)
      *** CALL PLT3DFILE(D2,N1)
      *
      *** WRITE(*,*) 'NORMALIZED GRAM-CHARLIER SLOPE DISTRIBUTION ARRAY'
      *** IF (N1.EQ.64) CALL ALPHA(D4,64)
      *** CALL PLT3DFILE(D4,N1)
      *
      RETURN
END

SUBROUTINE FFT2D(X,N,INV)
*
* CALCULATES FORWARD & BACKWARD FAST_FOURIER_TRANSFORM
*
* OUTPUT:
* X = TRANSFORMED COMPLEX SQUARE ARRAY
*
* INPUT:
* X = UNTRANSFORMED SQUARE ARRAY
* N = DIMENSION OF X
* INV = DIRECTION OF TRANSFORMATION:
*       -1 = FORWARD TRANSFORM
*       +1 = BACKWARD TRANSFORM
*
* IMPLICIT NONE
REAL ANGLE,PI,RTEM,WPWR
INTEGER I,INV,IREM,IT,ITER,J,J1,J2,K,L,M,MXP,M1,N,NXP,NXP2,N1,N2
COMPLEX X(1:N,1:N),W,T
*
PI=3.141593
ITER=0
IREM=N

```

```

*
10      IREM=INT(IREM/2)
*
*          IF (IREM.EQ.0) GOTO 20
*
*          ITER=ITER+1
GOTO 10
*
20      CONTINUE
*
CALL CHECKERBOARD(X,N)
*
DO 80 M1=1,2
DO 70 L=1,N
*
NXP2=N
*
DO 50 IT=1,ITER
*
NXP=NXP2
NXP2=INT(NXP/2)
*
IF (NXP2.EQ.0) GOTO 50
*
WPWR=PI/FLOAT(NXP2)
*
DO 40 M=1,NXP2
*
ANGLE=FLOAT(M-1)*WPWR
W=CMPLX(COS(ANGLE),FLOAT(INV)*SIN(ANGLE))
*
DO 40 MXP=NXP,N,NXP
*
J1=MXP-NXP+M
J2=J1+NXP2
*
IF (M1.EQ.1) THEN
    T=X(J1,L)-X(J2,L)
    X(J1,L)=X(J1,L)+X(J2,L)
    X(J2,L)=T*W
ELSE
    T=X(L,J1)-X(L,J2)
    X(L,J1)=X(L,J1)+X(L,J2)
    X(L,J2)=T*W
ENDIF
*
40      CONTINUE
50      CONTINUE
*
N2=N/2
N1=N-1
J=1

```

```

*
      DO I=1,N1
*
      IF (I.GE.J) GOTO 55
*
      IF (M1.EQ.1) THEN
          T=X(J,L)
          X(J,L)=X(I,L)
          X(I,L)=T
      ELSE
          T=X(L,J)
          X(L,J)=X(L,I)
          X(L,I)=T
      ENDIF
*
55      K=N2
*
60      IF (K.GE.J) GOTO 65
*
65      J=J-K
      K=K/2
      GOTO 60
      J=J+K
*
      END DO
70      CONTINUE
80      CONTINUE
*
      CALL CHECKERBOARD(X,N)
*
***      IF (INV.EQ.-1) GOTO 90
*
***      RTEM=FLOAT(N*N)
*
***      DO 90 J=1,N
***      DO 85 I=1,N
*
***          X(I,J)=X(I,J)/RTEM
*
***85      CONTINUE
***90      CONTINUE
*
      RETURN
      END

```

```

SUBROUTINE PICFILE(DATA,N,A,B,C)
IMPLICIT NONE
INTEGER I,J,N
REAL A,B,C
CHARACTER*64 DATA(204)
CHARACTER*1 OPT

```

```

CHARACTER*10 FNAME
*
      WRITE(*,*) 'WANT AN ALPHASCALE PIC FILE? [N] '
      READ(*,300) OPT
      IF (OPT.NE.'Y') GOTO 5
*
      WRITE(*,*) 'ENTER PIC FILENAME -> '
      READ(*,300) FNAME
      OPEN(3,FILE=FNAME,STATUS='UNKNOWN',ERR=10)
*
      WRITE(3,*) FNAME,:HIGH,LOW,DELTA =',A,B,C
      DO I=1,N
      WRITE(3,*) DATA(I),N/2-I
      END DO
*
      *** DO I=N+1,N+4
      *** WRITE(3,*) DATA(I)
      *** END DO
*
      CLOSE (3)
      WRITE(*,*) 'FILENAME = ',FNAME,' IS WRITTEN'
5      RETURN
10     WRITE(*,*) 'BAD FILE OPENING'
300    FORMAT (A)
      END

```

```

SUBROUTINE PLT2DFILE(X,N)
IMPLICIT NONE
INTEGER I,J,N
REAL X(N,2),Y,Z
CHARACTER*1 CHOICE
CHARACTER*18 FILENAME
WRITE(*,*) 'WANT A PLT2D OUTPUT FILE? [N] '
READ(*,300) CHOICE
IF (CHOICE.NE.'Y') GOTO 1
WRITE(*,*) 'ENTER PLT2D OUTPUT FILENAME ==>'
READ(*,300) FILENAME
OPEN(1,FILE=FILENAME,STATUS='NEW',ERR=5)
DO I=1,N
      WRITE(1,*)(X(I,J),J=1,2)
END DO
CLOSE(1)
WRITE(*,*) FILENAME,' HAS BEEN WRITTEN AS AN OUTPUT FILE'
1      RETURN
5      WRITE(*,*) 'BAD FILE OPENING'
300    FORMAT (A)
      END

```

```
SUBROUTINE PLT2DOUTPUT(FILENAME,X,N)
```

```

IMPLICIT NONE
INTEGER I,J,N
REAL X(N,2),Y,Z
CHARACTER*1 CHOICE
CHARACTER*18 FILENAME
*
WRITE(*,*) 'WANT A PLT2D OUTPUT FILE? [N] FILE =',FILENAME
*
READ(*,300) CHOICE
IF (CHOICE.NE.'Y') GOTO 1
*** WRITE(*,*) 'ENTER PLT2D OUTPUT FILENAME ==>'
**** READ(*,300) FILENAME
OPEN(1,FILE=FILENAME,STATUS='NEW',ERR=5)
DO I=1,N
WRITE(1,*)(X(I,J),J=1,2)
END DO
CLOSE(1)
WRITE(*,*) FILENAME,'HAS BEEN WRITTEN AS AN OUTPUT FILE'
1 RETURN
5 WRITE(*,*) 'BAD FILE OPENING'
300 FORMAT (A)
END

```

```

SUBROUTINE PLT3DFILE(X,N)
IMPLICIT NONE
INTEGER I,J,N
REAL X(N,N),Y,Z
CHARACTER*1 CHOICE
CHARACTER*18 FILENAME
WRITE(*,*) 'WANT A PLT3D OUTPUT FILE? [N] '
READ(*,300) CHOICE
IF (CHOICE.NE.'Y') GOTO 1
WRITE(*,*) 'ENTER PLT3D OUTPUT FILENAME ==>'
READ(*,300) FILENAME
OPEN(1,FILE=FILENAME,STATUS='NEW',ERR=5)
WRITE(1,*) N,N
DO I=1,N
WRITE(1,*)(X(I,J),J=1,N)
END DO
CLOSE(1)
WRITE(*,*) FILENAME,'HAS BEEN WRITTEN AS AN OUTPUT FILE'
1 RETURN
5 WRITE(*,*) 'BAD FILE OPENING'
300 FORMAT (A)
END

```

```

SUBROUTINE PLT3DINPUT(FILENAME,X,N,P,Q)
IMPLICIT NONE
INTEGER T,F,N,P,P1,Q,Q1
REAL X(F:P:Q)

```

```

CHARACTER*18 FILENAME
*
*** WRITE(*,*) 'ENTER PLT3D INPUT FILENAME ==>'
*** READ(*,300) FILENAME
*
OPEN(1,FILE=FILENAME,STATUS='OLD',ERR=5)
READ(1,*) N,N
P1=-N/2
Q1=N/2-1
*
DO I=P1,Q1
READ(1,*)(X(I,J),J=P1,Q1)
END DO
*
CLOSE(1)
WRITE(*,*) FILENAME,'HAS BEEN INPUT'
1 RETURN
5 WRITE(*,*) 'BAD FILE OPENING'
300 FORMAT (A)
END

SUBROUTINE PLT3DOUTPUT(FILENAME,X,N,P,Q)
IMPLICIT NONE
INTEGER I,J,N,P,P1,Q,Q1
REAL X(P:Q,P:Q)
CHARACTER*1 CHOICE
CHARACTER*18 FILENAME
*
WRITE(*,*) 'WANT A PLT3D OUTPUT FILE? [N] FILE= ',FILENAME
*
READ(*,300) CHOICE
IF (CHOICE.NE.'Y') GOTO 1
*
*** WRITE(*,*) 'ENTER PLT3D OUTPUT FILENAME ==>'
*** READ(*,300) FILENAME
*
OPEN(1,FILE=FILENAME,STATUS='NEW',ERR=5)
WRITE(1,*) N,N
P1=-N/2
Q1=N/2-1
DO I=P1,Q1
WRITE(1,*)(X(I,J),J=P1,Q1)
END DO
CLOSE(1)
WRITE(*,*) FILENAME,'HAS BEEN OUTPUT'
1 RETURN
5 WRITE(*,*) 'BAD FILE OPENING'
300 FORMAT (A)
END

```

```

SUBROUTINE PSVARY1(KNYQ,N,WIND_V,Z,VARY)
*
* CALCULATES ALONG-WIND, CROSS-WIND AND TOTAL VARIANCE OF SLOPE DISTRIBUTION
* FOR THREE SPECTRAL REGIONS: 1) BELOW LOWEST SAMPLE FREQUENCY, 2) WITHIN
* SAMPLED SPECTRAL REGION, & 3) ABOVE HIGHEST SAMPLE FREQUENCY
*
IMPLICIT NONE
INTEGER I,J,N,P,Q
REAL A,ALPH,ANG,BETA,D,DELT,DELTA,ENUG,G,K,K_MAX,KNU,KNYQ
REAL K1,K2,K3,L,M,P2,PI,SPEC,SPREAD_TOT,TEMP
REAL VAR_S,VAR_X,VAR_Y,VAR1_S,VAR1_X,VAR1_Y,VAR2_S,VAR2_X,VAR2_Y
REAL V_FRIC,V_MIN,V250,V1000,V1250,V1950,WIND_V,Z
REAL LNSPEC(-800:400,2),SPREAD(0:179),VARY(0:4,0:4)
CHARACTER*1 OPT
*
** WIND_V=WIND VELOCITY AT HEIGHT Z [CM/SEC^2]
** Z=HEIGHT OF WIND_V MEASUREMENT ABOVE WATER SURFACE [CM]
*
A=1.0
ALPH=0.0081
** ALPH=PHILLIPS CONSTANT []
BETA=0.74
** BETA= [UNITLESS]
ENUG=1.473E-4
** ENUG=E/(NU*G) [UNITLESS]
G=980.0
** G=ACCELERATION OF GRAVITY [CM/SEC^2]
K2=0.359
** [1/CM]
K3=0.942
** [1/CM]
K_MAX=3.63
** '' 'M'
J=3.141593
*
V_MIN=1.0
** V_MIN=MINIMUM FRICTION VELOCITY [CM/SEC]
*
WRITE(*,*) 'WANT FULL-SPECTRUM STATISTICS? (N) '
READ(*,300) OPT
IF (OPT.NE.'Y') GOTO 2
*
WRITE(*,*) 'CALCULATING FULL-SPECTRUM STATISTICS.....'
*
CALL WIND(WIND_V,Z,V_FRIC,V250,V1000,V1250,V1950)
*
D=(1.274+(0.0268*V_FRIC)+(6.03E-5*(V_FRIC**2)))**2
K1=(K2*(V_MIN**2))/(V_FRIC**2)
KNU=(0.5756*SQRT(V_FRIC)*K_MAX)/(D**0.16667)
P2=LOG10(D/(V_FRIC/V_MIN))/LOG10(K3/K2)
*
DELT=0.01

```

```

DELTA=PI/180.0
VAR_S=0.0
VAR_X=0.0
VAR_Y=0.0
VARY(0,2)=0.0
VARY(0,3)=0.0
VARY(0,4)=0.0
*
DO I=-800,400
L=FLOAT(I)*DELT
K=EXP(L)
LNSPEC(I,1)=LOG10(K)
*
CALL SPREADVARY(K,V1950,SPREAD,SPREAD_TOT)
*
IF ((0.0.LT.K).AND.(K.LE.K1)) THEN
SPEC=(ALPH/(2.0*(K**3)))*EXP((-1.0*BETA*(G**2))/((V1950**4)*(K**2)))
GOTO 1
ENDIF
*
IF ((K1.LT.K).AND.(K.LE.K2)) THEN
SPEC=ALPH/(2.0*SQRT(K1)*(K**2.5))
GOTO 1
ENDIF
*
IF ((K2.LT.K).AND.(K.LE.K3)) THEN
SPEC=(ALPH*D)/(2.0*(K3**P2)*(K***(3.0-P2)))
GOTO 1
ENDIF
*
IF ((K3.LT.K).AND.(K.LE.KNU)) THEN
SPEC=(ALPH*D)/(2.0*(K**3))
GOTO 1
ENDIF
*
IF (KNU.LT.K) THEN
SPEC=(ENUG*(V_FRIC**3)*(K_MAX**6))/(K**9)
ENDIF
*
1 LNSPEC(I,2)=(K**3)*SPEC
*
DO J=0,179
ANG=FLOAT(J)*DELT
TEMP=(SPREAD(J)*LNSPEC(I,2)*DELT*DELT)
VAR_S=VAR_S+TEMP
VAR_X=VAR_X+((COS(ANG)**2)*TEMP)
VAR_Y=VAR_Y+((SIN(ANG)**2)*TEMP)
END DO
*
*** WRITE(*,*) L,K,VAR
*
IF (EXP(L-DELT).LE.0.003) THEN

```

```

        VARY(1,1)=0.003
        VARY(1,2)=VAR_X
        VARY(1,3)=VAR_Y
        VARY(1,4)=VAR_S
    ENDIF
*
    IF (EXP(L-DELT).LE.(KNYQ/FLOAT(N/2))) THEN
        VARY(2,1)=KNYQ/FLOAT(N/2)
        VARY(2,2)=VAR_Y
        VARY(2,3)=VAR_X
        VARY(2,4)=VAR_S
    ENDIF
*
    IF (EXP(L-DELT).LE.KNYQ) THEN
        VARY(3,1)=KNYQ
        VARY(3,2)=VAR_Y
        VARY(3,3)=VAR_X
        VARY(3,4)=VAR_S
    ENDIF
*
    IF (EXP(L-DELT).LE.300.0) THEN
        VARY(4,1)=300.0
        VARY(4,2)=VAR_Y
        VARY(4,3)=VAR_X
        VARY(4,4)=VAR_S
    ENDIF
*
END DO
*
WRITE(*,*) 'RESULTS OF PSVARY1'
*** WRITE(*,*) 'SPREAD_TOT = ',SPREAD_TOT
WRITE(*,*) ''
*
DO I=1,4
WRITE(*,*) 'WAVE SLOPE VARIANCE UP TO ',VARY(I,1),' CY/CM & DIFFER'
WRITE(*,*) 'CUM CROSS-WIND VARIANCE = ',VARY(I,2),VARY(I,2)-VARY(I-1,2)

WRITE(*,*) 'CUM ALONG-WIND VARIANCE = ',VARY(I,3),VARY(I,3)-VARY(I-1,3)
WRITE(*,*) 'CUM COMBINED VARIANCE = ',VARY(I,4),VARY(I,4)-VARY(I-1,4)
WRITE(*,*) ''
END DO
*
2      RETURN
300    FORMAT (A)
END

```

SUBROUTINE PSVARY2(KNYQ,N,WIND\_V,Z,V1250,VARY)

```

*
* CALCULATES ALONG-WIND, CROSS-WIND AND TOTAL VARIANCE OF SLOPE DISTRIBUTION
* FOR THREE SPECTRAL REGIONS: 1) BELOW LOWEST SAMPLE FREQUENCY, 2) WITHIN
* SAMPLED SPECTRAL REGION, & 3) ABOVE HIGHEST SAMPLE FREQUENCY

```

```

*
      IMPLICIT NONE
      INTEGER I,J,N,P,Q
      REAL A,ALPH,ANG,BETA,D,DELT,DELTA,ENUG,G,K,K_MAX,KNU,KNYQ
      REAL K1,K2,K3,L,M,P2,PI,SPEC,SPREAD_TOT,TEMP
      REAL VAR_S,VAR_X,VAR_Y,VAR1_S,VAR1_X,VAR1_Y,VAR2_S,VAR2_X,VAR2_Y
      REAL V_FRIC,V_MIN,V250,V1000,V1250,V1950,WIND_V,Z
      REAL LNSPEC(-800:400,2),SPREAD(0:179),VARY(0:4,0:4)
      CHARACTER*1 OPT
*
** WIND_V=WIND VELOCITY AT HEIGHT Z [CM/SEC^2]
** Z=HEIGHT OF WIND_V MEASUREMENT ABOVE WATER SURFACE [CM]
*
      A=1.0
      ALPH=0.0081
** ALPH=PHILLIPS CONSTANT []
      BETA=0.74
** BETA= [UNITLESS]
      ENUG=1.473E-4
** ENUG=E/(NU*G) [UNITLESS]
      G=980.0
** G=ACCELERATION OF GRAVITY [CM/SEC^2]
      K2=0.359
** [1/CM]
      K3=0.942
** [1/CM]
      K_MAX=3.63
** [1/CM]
      PI=3.141593
**
      V_MIN=12.0
** V_MIN=MINIMUM FRICTION VELOCITY [CM/SEC]
*
*** WRITE(*,*) 'WANT FULL-SPECTRUM STATISTICS? [N] '
*** READ(*,300) OPT
*** IF ((OPT.NE.'Y').OR.(OPT.NE.'y')) GOTO 2
*
      WRITE(*,*) 'CALCULATING FULL-SPECTRUM STATISTICS.....'
*
      CALL WIND(WIND_V,Z,V_FRIC,V250,V1000,V1250,V1950)
*
      D=(1.274+(0.0268*V_FRIC)/(6.03E-5*(V_FRIC**2)))**2
      K1=(K2*(V_MIN**2))/(V_FRIC**2)
      KNU=(0.5756*SQRT(V_FRIC)*K_MAX)/(D**0.16667)
      P2=LOG10(D/(V_FRIC/V_MIN))/LOG10(K3/K2)
*
      DELT=0.01
      DELTA=PI/180.0
      VAR_S=0.0
      VAR_X=0.0
      VAR_Y=0.0
      VARY(0,2)=0.0

```

```

VARY(0,3)=0.0
VARY(0,4)=0.0
*
DO I=-800,400
L=FLOAT(I)*DELT
K=EXP(L)
LNSPEC(I,1)=LOG10(K)
*
CALL SPREADVARY(K,V1950,SPREAD,SPREAD_TOT)
*
IF ((0.0.LT.K).AND.(K.LE.K1)) THEN
SPEC=(ALPH/(2.0*(K**3)))*EXP((-1.0*BETA*(G**2))/((V1950**4)*(K**2)))
GOTO 1
ENDIF
*
IF ((K1.LT.K).AND.(K.LE.K2)) THEN
SPEC=ALPH/(2.0*SQRT(K1)*(K**2.5))
GOTO 1
ENDIF
*
IF ((K2.LT.K).AND.(K.LE.K3)) THEN
SPEC=(ALPH*D)/(2.0*(K3**P2)*(K***(3.0-P2)))
GOTO 1
ENDIF
*
IF ((K3.LT.K).AND.(K.LE.KNU)) THEN
SPEC=(ALPH*D)/(2.0*(K**3))
GOTO 1
ENDIF
*
IF (KNU.LT.K) THEN
SPEC=(ENUG*(V_FRIC**3)*(K_MAX**6))/(K**9)
ENDIF
*
1 LNSPEC(I,2)=(K**3)*SPEC
*
DO J=0,179
ANG=FLOAT(J)*DELT
TEMP=(SPREAD(J)*LNSPEC(I,2)*DELT*DELT)
VAR_S=VAR_S+TEMP
VAR_X=VAR_X+((COS(ANG)**2)*TEMP);
VAR_Y=VAR_Y+((SIN(ANG)**2)*TEMP)
END DO
*
*** WRITE(*,*) L,K,VAR
*
IF (EXP(L-DELT).LE.0.003) THEN
VARY(I,1)=0.003
VARY(I,2)=VAR_X
VARY(I,3)=VAR_Y
VARY(I,4)=VAR_S
ENDIF

```

```

*
    IF (EXP(L-DELT).LE.(KNYQ/FLOAT(N/2))) THEN
        VARY(2,1)=KNYQ/FLOAT(N/2)
        VARY(2,2)=VAR_Y
        VARY(2,3)=VAR_X
        VARY(2,4)=VAR_S
    ENDIF
*
    IF (EXP(L-DELT).LE.KNYQ) THEN
        VARY(3,1)=KNYQ
        VARY(3,2)=VAR_Y
        VARY(3,3)=VAR_X
        VARY(3,4)=VAR_S
    ENDIF
*
    IF (EXP(L-DELT).LE.300.0) THEN
        VARY(4,1)=300.0
        VARY(4,2)=VAR_Y
        VARY(4,3)=VAR_X
        VARY(4,4)=VAR_S
    ENDIF
*
    END DO
*
    WRITE(*,*) 'RESULTS OF PSVARY1'
***   WRITE(*,*) 'SPREAD_TOT = ',SPREAD_TOT
    WRITE(*,*) ''
*
    DO I=1,4
    WRITE(*,*) 'WAVE SLOPE VARIANCE UP TO ',VARY(I,1),' CY/CM & DIFFER'
    WRITE(*,*) 'CUM CROSS-WIND VARIANCE = ',VARY(I,2),VARY(I,2)-VARY(I-1,2)

    WRITE(*,*) 'CUM ALONG-WIND VARIANCE = ',VARY(I,3),VARY(I,3)-VARY(I-1,3)
    WRITE(*,*) 'CUM COMBINED VARIANCE = ',VARY(I,4),VARY(I,4)-VARY(I-1,4)
    WRITE(*,*) ''
    END DO
*
2      RETURN
300  FORMAT (A)
      END

```

SUBROUTINE RALPH(X,N,RY,IY,MY,PY)

```

*
* COMPOSITE SUBROUTINE TO DISPLAY COMPLEX ARRAY AS REAL, IMAGINARY,
* MAGNITUDE, AND PHASE COMPONENTS
*
* INPUT:
* COMPLEX ARRAY X(N,N)
*
```

```

* OUTPUTS:
* REAL ARRAY      RY(N,N)
* IMAGINARY ARRAY IY(N,N)
* MAGNITUDE ARRAY MY(N,N)
* PHASE ARRAY     PY(N,N)
*
IMPLICIT NONE
INTEGER N
COMPLEX X(N,N)
REAL RY(N,N),IY(N,N),MY(N,N),PY(N,N)
*
CALL RIMP(X,N,RY,IY,MY,PY)
*
WRITE(*,*) 'REAL COMPONENT OF COMPLEX ARRAY'
CALL ALPHA(RY,N)
*
WRITE(*,*) 'IMAGINARY COMPONENT OF COMPLEX ARRAY'
CALL ALPHA(IY,N)
*
WRITE(*,*) 'MAGNITUDE COMPONENT OF COMPLEX ARRAY'
CALL ALPHA(MY,N)
*
WRITE(*,*) 'PHASE COMPONENT OF COMPLEX ARRAY'
CALL ALPHA(PY,N)
*
RETURN
END

SUBROUTINE RIMP(X,N,RY,IY,MY,PY)
IMPLICIT NONE
INTEGER I,J,N
COMPLEX X(N,N)
REAL RY(N,N),IY(N,N),MY(N,N),PY(N,N),HALFPI
*
HALFPI=3.141593/2.0
*
DO I=1,N
    DO J=1,N
        RY(I,J)=REAL(X(I,J))
        IY(I,J)=AIMAG(X(I,J))
        MY(I,J)=CABS(X(I,J))
        IF ((RY(I,J).EQ.0.0).AND.(IY(I,J).EQ.0.0)) THEN
            PY(I,J)=0.0
        ELSEIF (IY(I,J).EQ.0.0) THEN
            PY(I,J)=SIGN(HALFPI,RY(I,J))
        ELSE
            PY(I,J)=ATAN2(IY(I,J),RY(I,J))
        ENDIF
    END DO
END DO
RETURN

```

END

```
SUBROUTINE SEA_SPEC(W_V,Z,W_AZ,DELTA_K,F1,F2,F3,F4,FX1,FY1,KCOR,N,P,Q)
*
* GENERATES PIERSON-STACY 2D SLOPE & ELEVATION SPECTRA IN RECT COORDINATES [L,M]
*
* CALCULATES VARIANCE OF SPECTRA FROM 1*DELTA_K TO (N/2-1)*DELTA_K
*
* INPUTS:
* W_V      = WIND_VELOCITY MEASUREMENT AT HEIGHT Z ABOVE WATER SURFACE [CM/SEC]
* Z        = HEIGHT OF WIND_VELOCITY MEASUREMENT [CM]
* W_AZ     = AZIMUTHAL DIRECTION OF WIND RELATIVE TO SENSOR [RAD]
* DELTA_K  = DELTA SPATIAL FREQUENCY [CY/CM]
*
* OUTPUTS:
* F1      = ABS( TOTAL (X+Y) SLOPE SPECTRUM) [ ]
* F2      = ABS(ELEVATION SPECTRUM) [CM]
* F3      = ABS(X-COMPONENT SLOPE SPECTRUM) [ ]
* F4      = ABS(Y-COMPONENT SLOPE SPECTRUM) [ ]
* FX1     = X-COMPONENT SLOPE SPECTRUM [ ]
* FY1     = Y-COMPONENT SLOPE SPECTRUM [ ]
* KCOR    = INTEGER(K+OFF) POINTER ARRAY TO POLAR-RECTANGULAR CORRECTION FACTOR
*
IMPLICIT NONE
INTEGER I,J,N,P,Q
INTEGER KCOR(P:Q,P:Q)
REAL A,ALPH,ANG0,ANG1,BETA,D,DELTA_K,ENUG,G,K,KINT,KMAG,K_MAX,KNU
REAL K1,K2,K3,L,M,P2,PI,SPANG,SPEC,SPREAD,SPRED0,SPRED1,SPRED2,SPRED3
REAL TEMP,VAR_X,VAR_Y,VAR_ELV,VAR_TOT,V_FRIC,V_MIN,V250,V1000,V1250
REAL V1950,W_AZ,W_V,X2,Y2,Z
REAL F1(P:Q,P:Q),F2(P:Q,P:Q),F3(P:Q,P:Q),F4(P:Q,P:Q)
REAL FX1(P:Q,P:Q),FY1(P:Q,P:Q),SPRDCOR(0:4096)
REAL CSPAN(0:512),SPANGCOR(0:512),SPECCOR(0:512,1:2)
*
A=1.0
ALPH=0.0081
* ALPH=PHILLIPS CONSTANT [UNITLESS]
BETA=0.74
* BETA= [ ]
ENUG=1.473E-4
* ENUG=E/(NU*G) [ ]
G=980.0
* G=ACCELERATION OF GRAVITY [CM/SEC^2]
K2=0.359
* [CY/CM]
K3=0.942
* [CY/CM]
K_MAX=3.63
* [CY/CM]
PI=3.141593
*
```

```

V_MIN=12.0
* V_MIN=MINIMUM FRICTION VELOCITY (CM/SEC)
*
CALL WIND(W_V,Z,V_FRIC,V250,V1000,V1250,V1950)
*
D=(1.274+(0.0268*V_FRIC)+(6.03E-5*(V_FRIC**2)))**2
K1=(K2*(V_MIN**2))/(V_FRIC**2)
KNU=((0.5756*SQRT(V_FRIC)*K_MAX)/(D**(0.16667)))
P2=LOG10(D/(V_FRIC/V_MIN))/LOG10(K3/K2)
*
DO I=P,Q
DO J=P,Q
L=FLOAT(I)
M=FLOAT(J)
KMAG=SQRT((L**2)+(M**2))
K=KMAG*DELTA_K
IF((K.EQ.0.0).OR.(K.GE.(FLOAT(-P)*DELTA_K))) THEN
SPANG=0.0
GOTO 2
ENDIF
ANG0=ATAN2(M,L)
ANG1=W_AZ-ANG0
*
* CALCULATE 2D DIRECTIONAL SPREADING FUNCTION
*
SPREDO=COS(ANG1)**2
SPRED1=(8.0/(3.0*PI))*(SPREDO**2)
SPRED2=EXP((-1.0*(G**2))/(2.0*(K**2)*(V1950**4)))
SPRED3=(1.0-(A/2.0)+(A*SPREDO))/PI
SPREAD=((SPRED1*(1.0-SPRED2))+(SPRED3*SPRED2))
*
* CALCULATE POLAR-TO-RECTANGULAR CORRECTION FACTORS
*
KINT=4*INT(KMAG-0.25)
SPRDCOR(KINT)=SPRDCOR(KINT)+SPREAD
KCOR(I,J)=KINT
*
* CALCULATE 1D ELEVATION POWER SPECTRUM
*
IF ((0.0.LT.K).AND.(K.LE.K1)) THEN
SPEC=(ALPH/(2.0*(K**3)))*EXP((-1.0*BETA*(G**2))/((V1950**4)*(K**2)))
GOTO 1
ENDIF
*
IF ((K1.LT.K).AND.(K.LE.K2)) THEN
SPEC=ALPH/(2.0*SQRT(K1)*(K**2.5))
GOTO 1
ENDIF
*
IF ((K2.LT.K).AND.(K.LE.K3)) THEN
SPEC=(ALPH*D)/(2.0*(K3**P2)*(K***(3.0-P2)))
GOTO 1

```

```

        ENDIF
*
        IF ((K3.LT.K).AND.(K.LE.KNU)) THEN
          SPEC=(ALPH*D)/(2.0*(K**3))
          GOTO 1
        ENDIF
*
        IF (KNU.LT.K) THEN
          SPEC=(ENUG*(V_FRIC**3)*(K_MAX**6))/(K**9)
        ENDIF
*
* CALCULATE UNCORRECTED 2D SLOPE & ELEVATION SPECTRA
*
1       SPANG=SQRT(SPREAD*SPEC)
        X2=(L*DELTA_K)
        Y2=(M*DELTA_K)
*
2       F1(I,J)=K*SPANG
        F2(I,J)=SPANG
        FX1(I,J)=X2*SPANG
        FY1(I,J)=Y2*SPANG
        F3(I,J)=ABS(X2*SPANG)
        F4(I,J)=ABS(Y2*SPANG)
*
* OPTION: CALCULATE POLAR-RECTANGULAR CORRECTION DATA
*
***      SPECCOR(KINT,1)=SPECCOR(KINT,1)+((K**2)*SPEC)
***      SPECCOR(KINT,2)=SPECCOR(KINT,2)+1.0
***      SPANGCOR(KINT)=SPANGCOR(KINT)+F1(I,J)**2
*
        END DO
        END DO
*
* CORRECT SPECTRA FOR POLAR-TO-RECTANGULAR CONVERSION
* CALCULATE CUMULATIVE VARIANCES
*
        DO I=P,Q
        DO J=P,Q
*
        KMAG=SQRT(FLOAT(I)**2+FLOAT(J)**2)
*
        KINT=KCOR(I,J)
        IF (KINT.EQ.0) GOTO 3
*
        TEMP=SQRT(SPRDCOR(KINT)*DELTA_K)
*
        F1(I,J)=F1(I,J)/TEMP
        VAR_TOT=VAR_TOT+(F1(I,J)**2)*(DELTA_K**2)
        CSPAN(KINT)=CSPAN(KINT)+F1(I,J)**2
        F2(I,J)=F2(I,J)/TEMP
        VAR_ELV=VAR_ELV+(F2(I,J)**2)*(DELTA_K**2)
        FX1(I,J)=FX1(I,J)/TEMP

```

```

        VAR_X=VAR_X+(FX1(I,J)*(.2)**(DELTA_K**2))
        FY1(I,J)=FY1(I,J)/TEMP
        VAR_Y=VAR_Y+(FY1(I,J)**2)*(DELTA_K**2)
        F3(I,J)=F3(I,J)/TEMP
        F4(I,J)=F4(I,J)/TEMP
*
* ADDITIONAL VARIANCE FOR DOUBLY-SAMPLED NYQUIST FREQUENCY
*
        IF (KMAG.GT.FLOAT(Q)) THEN
            VAR_TOT=VAR_TOT+(F1(I,J)**2)*(DELTA_K**2)
            VAR_ELV=VAR_ELV+(F2(I,J)**2)*(DELTA_K**2)
            VAR_X=VAR_X+(FX1(I,J)**2)*(DELTA_K**2)
            VAR_Y=VAR_Y+(FY1(I,J)**2)*(DELTA_K**2)
        ENDIF
*
3      END DO
        END DO
*
* OPTION: WRITE OUT POLAR-RECTANGULAR CORRECTION PARAMETERS
*
***      WRITE(*,*) 'K,SPANGCOR,SPRDGCR,CSPAN'
***      DO I=1,Q+1
***      K=FLOAT(I)*DELTA_K
***      WRITE(*,*) K,SPANGCOR(I),SPRDGCR(I),CSPAN(I)
***      END DO
*
        WRITE(*,*) 'TOTAL SLOPE VARIANCE OF THE Y-COMPONENTS = ',VAR_Y
        WRITE(*,*) 'TOTAL SLOPE VARIANCE OF THE X-COMPONENTS = ',VAR_X
        WRITE(*,*) 'TOTAL SLOPE COMBINED VARIANCE           = ',VAR_TOT
        WRITE(*,*) 'TOTAL ELEVATION VARIANCE     [CM^2]       = ',VAR_ELV
*
        RETURN
    END

SUBROUTINE SKYBLUE(LREF,PHIO,ZO,N,P,Q,LSKY,HSKY,VSKY)
*
* CREATE UNPOLARIZED SKYDOME HEMISPHERICAL RADIANCE DISTRIBUTION
* RESOLVE HORIZONTALLY_ & VERTICALLY_POLARIZED RADIANCE COMPONENTS
*
* INPUT:
* LREF = REFERENCE RADIANCE MEASURED AT THE ZENITH POINT, LSKY(0,0) [UNIT]
* PHIO = SUN AZIMUTH [DEG] {-180..+180}
* ZO   = SUN ECLINATION
*
* OUTPUT:
* LSKY = UNPOLARIZED SKYDOME ANGULAR RADIANCE DISTRIBUTION
* HSKY = HORIZONTAL_POLARIZED LSKY
* VSKY = VERTICALLY_POLARIZED LSKY
*
IMPLICIT NONE
INTEGER I,J,N,P,Q

```

```

REAL LSKY(P:Q,P:Q),HSKY(P:Q,P:Q),VSKY(P:Q,P:Q)
REAL A,B,C,COSMU,COSMU2,COSNU,H,L,LREF,M,MU,PHI,PHIO,PI,PSI,THETA,Z0
CHARACTER*18 FNAME
*
PI=3.14159
*
DO I=P+1,Q
DO J=P+1,Q
*
L=FLOAT(I)
M=FLOAT(J)
THETA=(PI/2.0)*(SQRT((L**2)+(M**2))/FLOAT(Q))
*
IF (THETA.GT.PI/2.0) THEN
    LSKY(I,J)=0.0
    HSKY(I,J)=0.0
    VSKY(I,J)=0.0
    GOTO 4
ENDIF
*
IF (THETA.EQ.0.0) THEN
    PHI=0.0
    GOTO 1
ENDIF
*
PHI=ATAN2(M,L)
*
1 COSMU=COS(Z0)*COS(THETA)+SIN(Z0)*SIN(THETA)*COS(PHI-PHIO)
MU=ACOS(COSMU)
A=0.91+10.0*EXP(-3.0*MU)+0.45*(COSMU**2.0)
IF (THETA.GE.1.57) THEN
    B=1.0
    GOTO 2
ENDIF
B=1.0-EXP(-0.32/COS(THETA))
2 C=0.274*(0.91+10.0*EXP(-3.0*Z0)+0.45*((COS(Z0))**2.0))
LSKY(I,J)=LREF*A*B/C
*
* CALCULATE HORIZONTALLY & VERTICALLY POLARIZED FRACTIONS
* VIA PURE RAYLEIGH SCATTERING MODEL
*
IF (MU.EQ.0.0) THEN
    PSI=0.0
    GOTO 3
ENDIF
*
COSMU2=COSMU**2
COSNU=(SIN(THETA)*COS(Z0)-SIN(Z0)*COS(THETA)*COS(PHI-PHIO))/SIN(MU)
PSI=0.94*((1.0-COSMU2)/(1.0+COSMU2))
3 H=.5*(1.0-PSI)+(PSI*(COSNU**2))
HSKY(I,J)=LSKY(I,J)*H
VSKY(I,J)=LSKY(I,J)*(1.0-H)

```

```

*
4      END DO
END DO
*
      WRITE(*,*) 'LSKY = UNPOLARIZED SKYDOME ANGULAR RADIANCE DISTRIBUTION'
      IF (N.EQ.64) CALL ALPHA(LSKY,64)
      FNAME='LSKY.RAY'
      CALL PLT3DOUTPUT(FNAME,LSKY,N,P,Q)
      CALL PLT3DFILE(LSKY,N)
*****
      WRITE(*,*) 'HSKY = HORIZONTALLY_POLARIZED LSKY'
      IF (N.EQ.64) CALL ALPHA(HSKY,64)
      FNAME='HSKY.RAY'
      CALL PLT3DOUTPUT(FNAME,HSKY,N,P,Q)
      CALL PLT3DFILE(HSKY,N)
*****
      WRITE(*,*) 'VSKY = VERTICALLY_POLARIZED LSKY'
      IF (N.EQ.64) CALL ALPHA(VSKY,64)
      FNAME='VSKY.RAY'
      CALL PLT3DOUTPUT(FNAME,VSKY,N,P,Q)
      CALL PLT3DFILE(VSKY,N)
*****
      RETURN
END

SUBROUTINE SKYMAP(LREF,PHIO,Z0,THETA,N,P,Q,LSRF,HSRF,VSRF)
*
* CREATE UNPOLARIZED SKYDOME HEMISPHERICAL RADIANCE DISTRIBUTION
* IN SURFACE SLOPE COORDINATES (BETA,ALPHA) RELATIVE TO SENSOR
* RESOLVE HORIZONTALLY_ & VERTICALLY_POLARIZED RADIANCE COMPONENTS
*
* INPUT:
* LREF = REFERENCE RADIANCE MEASURED AT THE ZENITH POINT, LSKY(0,0) [UNIT]
* PHIO = SUN AZIMUTH [DEG] [-180..+180]
* THETA = SENSOR AZIMUTH ANGLE [DEG] [0..+90]
* Z0   = SUN AZIMUTH ANGLE [DEG] [0..+90]
*
* OUTPUT:
* LSRF = LINEAR MAP OF LSKY TO SURFACE COORDINATES (ALPH,BETA)
* HSRF = LINEAR MAP OF HSKY TO SURFACE COORDINATES (ALPH,BETA)
* VSRF = LINEAR MAP OF VSKY TO SURFACE COORDINATES (ALPH,BETA)
*
IMPLICIT NONE
INTEGER I,J,N,P,Q
REAL ALPH,BETA,COSOMEGA,DELTA,IV1,IV2,IV3,NV1,NV2,NV3,RV1,RV2,RV3,THET
REAL LSRF(P:Q,P:Q),HSRF(P:Q,P:Q),VSRF(P:Q,P:Q)
REAL A,B,C,COSMU,COSNU2,COSNU,H,L,LREF,M,MU,PHI,PHIO,PI,PSI,THETA,Z0
CHARACTER*18 FNAME
*
PI=3.14159
DELTA=90.0/FLOAT(Q)

```

```

*
* IV1,IV2,IV3 ARE THE RESOLVED COMPONENTS OF THE SENSOR COORDINATE VECTOR
*
        IV1=SIND(THETA)*SIND(180.0)
        IV2=SIND(THETA)*COSD(180.0)
        IV3=COSD(THETA)

*
        DO I=P+1,Q
        DO J=P+1,Q

*
        L=FLOAT(I)*DELTA
        M=FLOAT(J)*DELTA

*
        BETA=SQRT((L**2)+(M**2))

*
        IF (BETA.GE.90.0) GOTO 4

*
        IF (BETA.GT.0.0) THEN
            ALPH=ATAN2D(M,L)
        ELSE
            ALPH=0.0
        ENDIF

*
* NV1,NV2,NV3 ARE THE RESOLVED COMPONENTS OF THE SURFACE NORMAL VECTOR
*
        NV1=SIND(BETA)*SIND(ALPH)
        NV2=SIND(BETA)*COSD(ALPH)
        NV3=COSD(BETA)

*
* DOT PRODUCT OF IV & NV = |IV||NV|(-OSOMEGA) = COSOMEGA
*
        COSOMEGA=(IV1*NV1)+(IV2*NV2)+(IV3*NV3)

*
* RV1,RV2,RV3 ARE THE RESOLVED COMPONENTS OF THE REFLECTED SKYDOME VECTOR
*
        RV1=(NV1*2.0*COSOMEGA)-IV1
        RV2=(NV2*2.0*COSOMEGA)-IV2
        RV3=(NV3*2.0*COSOMEGA)-IV3

*
        THET=ACOS(RV3)

*
        IF (THET.GT.0.0) THEN
            PHI=ATAN2(RV1,RV2)
        ELSE
            PHI=0.0
        ENDIF

*
        IF (THET.GT.PI/2.0) THEN
            LSRF(I,J)=0.0
            HSRF(I,J)=0.0
            VSRF(I,J)=0.0
            GOTO 4

```

```

        ENDIF
*
1      COSMU=COS(Z0)*COS(THET)+SIN(Z0)*SIN(THET)*COS(PHI-PHIO)
      MU=ACOS(COSMU)
      A=0.91+10.0*EXP(-3.0*MU)+0.45*(COSMU**2.0)
      IF (THET.GE.1.57) THEN
          B=1.0
          GOTO 2
      ENDIF
      B=1.0-EXP(-0.32/COS(THET))
2      C=0.274*(0.91+10.0*EXP(-3.0*Z0)+0.45*((COS(Z0))**2.0))
      LSRF(I,J)=LREF*A*B/C
*
* CALCULATE HORIZONTALLY & VERTICALLY POLARIZED FRACTIONS
* VIA PURE RAYLEIGH SCATTERING MODEL
*
      IF (MU.EQ.0.0) THEN
          PSI=0.0
          GOTO 3
      ENDIF
*
      COSMU2=COSMU**2
      COSNU=(SIN(THET)*COS(Z0)-SIN(Z0)*COS(THET)*COS(PHI-PHIO))/SIN(MU)
      PSI=0.94*((1.0-COSMU2)/(1.0+COSMU2))
3      H=(0.5*(1.0-PSI))+(PSI*(COSNU**2))
      HSRF(I,J)=LSRF(I,J)*H
      VSRF(I,J)=LSRF(I,J)*(1.0-H)
*
4      END DO
      END DO
*
*** WRITE(*,*) 'LSRF = LSKY MAPPED TO SURFACE COORDINATES'
*** IF (N.EQ.64) CALL ALPHA(LSRF,64)
*** FNAME='LSRF.RAY'
*** CALL PLT3DOUTPUT(FNAME,LSRF,N,P,Q)
*** CALL PLT3DFILE(LSRF,N)
*
*** WRITE(*,*) 'HSRF = HSKY MAPPED TO SURFACE COORDINATES'
*** IF (N.EQ.64) CALL ALPHA(HSRF,64)
*** FNAME='HSRF.RAY'
*** CALL PLT3DOUTPUT(FNAME,HSRF,N,P,Q)
*** CALL PLT3DFILE(HSRF,N)
*
*** WRITE(*,*) 'VSRF = VSKY MAPPED TO SURFACE COORDINATES'
*** IF (N.EQ.64) CALL ALPHA(VSRF,64)
*** FNAME='VSRF.RAY'
*** CALL PLT3DOUTPUT(FNAME,VSRF,N,P,Q)
*** CALL PLT3DFILE(VSRF,N)
*
      RETURN
      END

```

```

SUBROUTINE SPREADVARY(K,V1950,F3,SPREAD_TOT)
IMPLICIT NONE
INTEGER I,J,N,P,Q
REAL A,ANG1,DELTA,G,K
REAL PI,SPREAD,SPREDO,SPRED1,SPRED2,SPRED3
REAL SPREAD_TOT,V_FRIC,V_MIN,V1950,X2,Y2,Z
REAL F3(0:179)

*
A=1.0
G=980.0
PI=3.141593
DELTA=PI/180.0
SPREAD_TOT=0.0

*
DO I=0,179
    ANG1=FLOAT(I)*DELTA
    SPREDO=COS(ANG1)**2
    SPRED1=(8.0/(3.0*PI))*(SPREDO**2)
    SPRED2=EXP((-1.0*(G**2))/(2.0*(K**2)*(V1950**4)))
    SPRED3=(1.0~(A/2.0)+(A*(SPREDO)))/PI
    SPREAD=(SPRED1*(1.0-SPRED2))+(SPRED3*SPRED2)
    SPREAD_TOT=SPREAD_TOT+SPREAD*DELTA
    F3(I)=SPREAD
END DO
*
RETURN
END

```

```

SUBROUTINE WGN_SPEC(W1,GWN,ISEED,N,P,Q)
*
* CREATE FREQUENCY DOMAIN REPRESENTATION OF WHITE GAUSSIAN NOISE
* OUTPUT:
* N*N COMPLEX MATRIX GWN(P:Q,P:Q) WHERE:
*      P=-N/2 AND Q=N/2-1
*      REAL(WGN(I,J)) VARIES FROM -1.0 TO +1.0
*      IMAG(WGN(I,J)) VARIES FROM -1.0 TO +1.0
*      MAG (WGN(I,J)) EQUALS +1.0 EVERYWHERE
*      PHS (WGN(I,J)) VARIES FROM -PI TO +PI
*
IMPLICIT NONE
INTEGER I,J,N,P,Q
INTEGER*4 ISEED
REAL VAR,X1,X2,Y1,Y2
REAL W1(P:Q,P:Q)
COMPLEX GWN(P:Q,P:Q)
*
VAR=FLOAT(N*N)
*
* CREATE N*N REAL MATRIX W1(P:Q,P:Q) OF PSEUDORANDOM VALUES
*
```

```

      CALL WGN2D(N,VAR,ISEED,W1)
*
* INSERT W1 VALUES INTO REAL COMPONENTS OF N*N COMPLEX MATRIX GWN(P:Q,P:Q)
*
      DO I=P,Q
        DO J=P,Q
          GWN(I,J)=CMPLX(W1(I,J),0.0)
        END DO
      END DO
*
* FORWARD-FFT THE COMPLEX MATRIX GWN(P:Q,P:Q) INTO THE FREQUENCY DOMAIN
*
      CALL FFT2D(GWN,N,-1)
*
* NORMALIZE THE MAGNITUDE OF GWN(P:Q,P:Q)
*
      DO I=P,Q
        DO J=P,Q
          GWN(I,J)=GWN(I,J)/CABS(GWN(I,J))
        END DO
      END DO
*
      RETURN
    END

SUBROUTINE WGN2D(N,VAR,ISEED,WAY)
REAL WAY(N,N),VAR,SUM,TEMP,TVAR
INTEGER I,J,K,N
INTEGER*4 ISEED
*
TVAR=SQRT(VAR)*1.414159
*
DO I=1,N
  DO J=1,N
    SUM=0.0
    DO K=1,6
      TEMP = RAN (ISEED)
      SUM = SUM + TEMP
    END DO
    WAY(I,J)=TVAR*(SUM-3.0)
  END DO
END DO
RETURN
END

SUBROUTINE WIND(WIND_V,Z,V_FRIC,V250,V1000,V1250,V1950)
*
* OUTPUTS WIND VELOCITIES AT HEIGHTS 0 CM, 250 CM, 1250 CM, & 1950 CM
* ABOVE WATER SURFACE BASED ON INPUT WIND_VELOCITY MEASUREMENT AND MEASUREMENT
* HEIGHT ABOVE WATER SURFACE. AN HOMOGENEOUS WIND PROFILE IS ASSUMED.

```

```

*
* INPUTS:
* WIND_V = INPUT WIND_VELOCITY MEASUREMENT [CM/SEC]
* Z      = MEASUREMENT HEIGHT ABOVE WATER SURFACE [CM]
* OUTPUTS:
* V_FRIC = WIND VELOCITY AT    0 CM          [CM/SEC]
* V250   = WIND VELOCITY AT  250 CM          [CM/SEC]
* V1000  = WIND VELOCITY AT 1000 CM          [CM/SEC]
* V1250  = WIND VELOCITY AT 1250 CM          [CM/SEC]
* V1950  = WIND VELOCITY AT 1950 CM          [CM/SEC]
*
IMPLICIT NONE
INTEGER I,J,N
REAL RATIO,WIND_V,V_FRIC,V250,V1000,V1250,V1950,X(46,6),Y,Z,Z0
*
OPEN(1,FILE='V_DATA.DAT',STATUS='OLD',ERR=5)
DO I=1,46
READ(1,*) (X(I,J),J=1,6)
END DO
CLOSE(1)
*
IF (Z.EQ.1950.0) J=6
IF (Z.EQ.1250.0) J=5
IF (Z.EQ.1000.0) J=4
IF (Z.EQ.250.0) J=3
IF (Z.EQ.0.0) J=2
*
DO I=1,46
IF (X(I,J).GT.WIND_V) THEN
  RATIO=(WIND_V-X(I-1,J))/(X(I,J)-X(I-1,J))
  V_FRIC=X(I-1,2)+(RATIO*(X(I,2)-X(I-1,2)))
  GOTO 1
ENDIF
END DO
*
1 Z0=(0.684/V_FRIC)+(4.28E-5*(V_FRIC**2))-0.0443
V250=(V_FRIC/0.4)*LOG(250.0/Z0)
V1000=(V_FRIC/0.4)*LOG(1000.0/Z0)
V1250=(V_FRIC/0.4)*LOG(1250.0/Z0)
V1950=(V_FRIC/0.4)*LOG(1950.0/Z0)
*
WRITE(*,*) 'V_FRIC = ',V_FRIC
WRITE(*,*) 'V250   = ',V250
WRITE(*,*) 'V1000  = ',V1000
WRITE(*,*) 'V1250  = ',V1250
WRITE(*,*) 'V1950  = ',V1950
*
RETURN
5 WRITE(*,*) 'BAD FILE OPENING FOR V_DATA.DAT'
END

```

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