**SUMMARY**

This document includes a summary of parameters and inputs for the tectonic subsidence model presented in Chapter 2.

**Tectonic subsidence model**

Representative stratigraphic thicknesses, lithological information, and inferred depositional environments were utilized to develop a tectonic subsidence curve for the entirety of the Trout Creek Sequence and McCoy Creek Group, as well as the overlying Paleozoic succession (through the Ely Limestone) within the model framework of Zhang et al. (2023a,b). All inputs for this portion of the model can be viewed in the first sheet (“Stratigraphy”) of Table SM4; a detailed discussion of the SubsidenceChron.jl model framework (and parameters therein) can be viewed within Zhang et al. (2023a,b).

Stratigraphic age constraints were integrated into a Bayesian age-height model framework (Schoene et al., 2019), encapsulated within the model framework of Zhang et al. 2023b), to generate model ages for all stratigraphic heights included within the tectonic subsidence curve model. Age controls for the Trout Creek Sequence and McCoy Creek Group, based on the age framework depicted in Fig. 6, are collated in the second sheet (“Age\_control”) of Table SM4. An estimated depositional age of strata below Trout Creek Sequence Unit 1, with large approximated uncertainty, was included to allow a continuous age-height curve to be modeled through the base of Unit 1. Paleozoic age controls include the base of the Pioche Shale, which can be correlated with the ca. 507 Ma base of the Bright Angel Shale (Karlstrom et al., 2020), and a minimum depositional age at the top of the Lower Pennsylvanian Ely Limestone (Long et al., 2022). The age-height model was then used to generate interpolated ages for all stratigraphic heights integrated into the tectonic subsidence curve, allowing tectonic subsidence to be plotted against time (blue curve, Fig. 8a).

The age-height model was also used to generate a model undecompacted sedimentation rate curve. Note that we include an age control point at the base of Marinoan glacial strata (diamictite sequence of Trout Creek Sequence Unit 7), with an estimated glacial onset age of ~640 Ma, in line with the conservative estimate used to calculate the linear Cryogenian sedimentation rates depicted with the dashed purple line in Fig. 8a. The incorporation of this approximate constraint does not appreciably change the structure of the tectonic subsidence curve, but allows for a better representation of sedimentation rate changes across the Cryogenian in the sedimentation rate model. All code used to generate the tectonic subsidence model and sedimentation rate curve depicted in Fig. 8a can be viewed within the Code directory in this repository.

A McKenzie-style idealized thermal subsidence curve was fit to the tectonic subsidence curve by generating posterior distributions of the timing of thermal subsidence onset (T0) and the crustal stretching factor (β) (Fig. 8c), which were then repeatedly sampled to generate a range of possible thermal subsidence curves (red curve, Fig. 8b). All calculated thermal subsidence curves utilize a lithospheric cooling constant, τ, of 65 Ma. The initial choice of a prior T0 was informed by geological evidence from the Cordilleran Laurentian margin, which suggests the occurrence of active rifting in the Cryogenian as evidenced by syn-extensional Cryogenian deposition in Death Valley (Macdonald et al., 2013; Nelson et al., 2020) and Cryogenian volcanism in Idaho (Isakson et al., 2022, Keeley et al., 2013). With a mid-Cryogenian (T0=651±5 Ma,1s) rift-drift transition prior set, beta-factor priors were then iteratively increased until β=5.9±.75 (1s), which resulted in the modeled tectonic subsidence curve being well-matched by idealized thermal subsidence curves generated from the resultant posterior β and T0 distributions (Fig. 8b,c). All code (modified from Zhang et al., 2023b) used to generate the tectonic subsidence curve and thermal subsidence curves and posterior parameter distributions in Fig. 8b,c can be viewed within the Code directory of this repository.

**ANALYTICAL RESULTS**

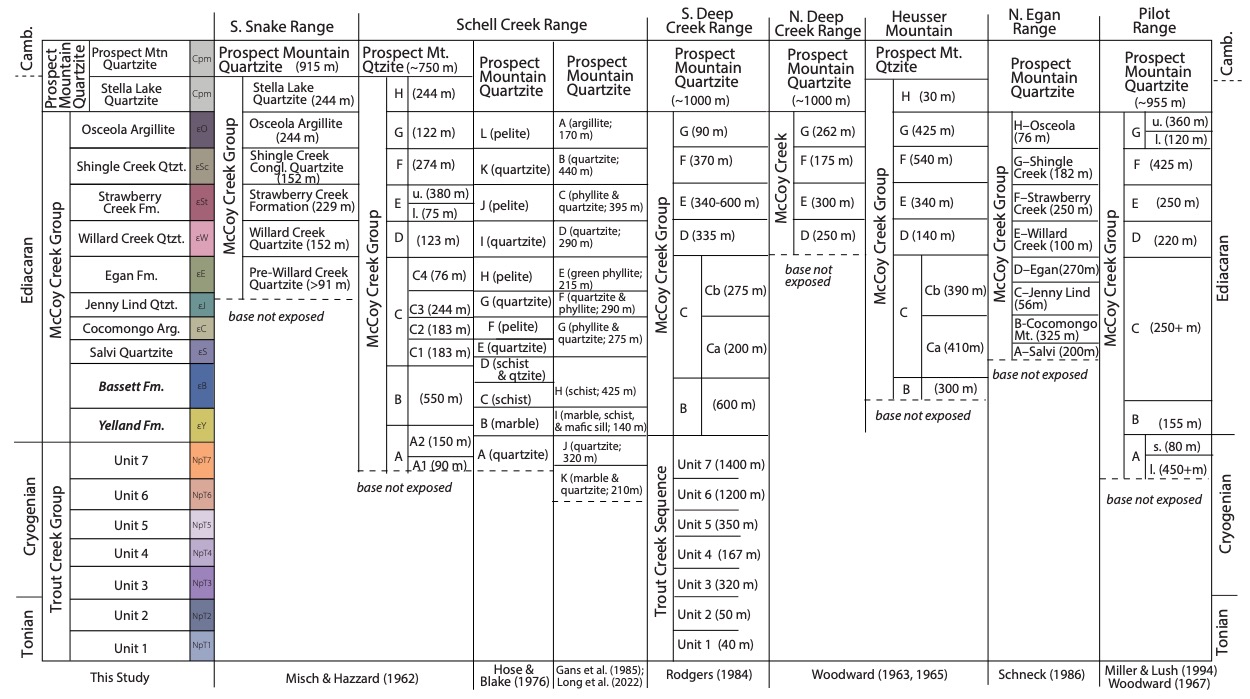
**U-Pb Zircon Geochronology**

Sandstone samples were collected through the Trout Creek Gp and McCoy Creek Gp in the Deep Creek, Schell Creek, and Egan Ranges to compliment the previous detrital zircon geochronological dataset of Yonkee et al. (2014) from the Deep Creek Range. Sample locations and new LA-ICPMS zircon data are collated in Table SM2. From TCU3, sample F2005-273.4/275 yielded an age spectrum with dominantly Stenian peaks, while sample TC4-1, from TCU4, features a spectrum with Stenian peaks, as well as a young peak that can be discretized into two distinct magmatic populations (688.5±2.44 Ma, n=6; 696.9±3.40 Ma, n=3; Fig. 1). From TCU7, sample F2003-3 yielded a dominant Stenian peak in the detrital zircon spectrum, with subsidiary 1300-1500Ma peaks. Sample F2002-204.9 from the Yelland Fm yielded a dominant peak around 1300 Ma, while samples F2002-273.4 and MC\_2 combined to yield dominant Stenian and 1300-1500Ma (syn-Picuris) peaks in the Bassett fm. The Egan Fm, represented by samples Egan 182.0 and Egan 233.0, features a dominant stenian peak and subsidiary syn-Picuris peak, while sample Egan 384.4 from the Willard Creek Quartzite and sample Egan 564.0 from the Strawberry Creek Fm display a similar pattern in the composite detrital zircon age spectrum for each respective unit. Sample M1901-10.25 from the Osceola Argillite yields a detrital zircon spectra with a dominant age peak between 1600-1800Ma, with a conspicuous lack of Stenian zircon. All detrital spectra, including those incorporating data from this study and those of Yonkee et al., 2014, are shown in Fig. 9 in the main manuscript.

**Carbonate carbon and oxygen isotope analyses**

We report 164 carbonate carbon (δ13C) and oxygen (δ18O) isotope analyses from the Trout Creek Sequence and the McCoy Creek Gp; all carbonate carbon and oxygen isotopic data, and sampling locations, are collated in Table SM3. δ13C values from Unit 2 of the Trout Creek Sequence vary from 0 to +4 ‰, and average +1.8 ‰. δ18O values vary from -5 to -16 ‰, average -9.3 ‰, and do not show significant covariance with δ13C. Three measurements from carbonates in the ~20 cm thick cap carbonate have δ13C values between -4.2 and -6.4 ‰, and δ18O values that average -16.5 ‰. Two carbonate intervals within the upper portions Unit 7 of the Trout Creek Sequence were sampled for carbonate chemostratigraphy, and host δ13C values between +5 and +9‰, with most values clustering between +7 and +8‰. The 8 m-thick cap dolostone of the basal Yelland Fm of the McCoy Creek Gp was sampled at 0.5 m resolution. These 16 samples have consistent δ13C values between -2.3 and -2.9 ‰, and δ18O values between -9.4 and -11.2 ‰ (Table SM3). Higher in unit B, δ13C values in limestones become more variable, but remain negative and do not covary with δ18O.

The Osceola Fm of the McCoy Creek Gp was sampled for carbonate chemostratigraphy in the S. Snake Range and the Pilot Range. In the S. Snake Range, four samples were analyzed from the ~20 cm thick limestone horizon. The lower horizon yielded δ13C values of -11.37 and -11.24 ‰, and the upper horizon yielded δ13C values of -9.19 and -9.39. In the Pilot Range, δ13C values increase up-section from ~-9.5 to -8.5 ‰ over a 70m interval.

**Fig. SM1**

**Figure SM1:** Unit nomenclature for the Trout Creek Sequence and McCoy Creek Group, including previous naming schemes (right) and the full stratigraphic nomenclature (including map unit colors and symbols) used in this study (left). Note that the new nomenclature is an amalgamation of the naming schemes of Misch and Hazzard (1962), Rodgers (1984), and Schneck (1986); two new formalized units are highlighted with bold italicized text. Formalization tables for the Yelland Fm and Bassett Fm can be found in Table SM1.

**Fig. SM2**

A map of a geological study

Description automatically generated

**Figure SM2:** Geologic map of the Heusser Mountain study area (see Fig. 1b). In addition to measured section F2262, a schematic section through the Osceola Fm (Fig. 5 inset) at this locality was approximated from map relationships and geometries.

**Fig. SM3**

A chart of different types of soil

Description automatically generated

**Figure SM3: Version of Figure 6 from main manuscript, with references for all geochronological constraints. A)** Generalized lithostratigraphy of the Trout Creek Sequence and McCoy Creek Group. **B)** Carbonate δ13C chemostratigraphy of the Trout Creek Sequence and McCoy Creek Group, correlated with **C)** a compiled global Tonian-Ediacaran δ13C chemostratigraphy, amalgamated from the compilations of Yang et al. (2021) and Park et al. (2019). All chemostratigraphic data from this study, including sampling location coordinates, are compiled in Table SM3.

**Fig. SM4**



**Figure SM4.** Version of Figure 10A, including references to graphical components of the figure. Proterozoic sedimentary cover of Laurentia, schematically depicting the inferred extent of pre-Grenville sedimentary cover (transparent blue polygon) and overlying Stenian (syn-Grenville) sediments (transparent green polygon). Remnant pre-Grenville quartzites, including modern outcrop and well-characterized subsurface extents, are depicted in dark blue (modified from Carlson, 1970; Southwick et al., 1986; Jones III and Thrane, 2012, Medaris et al., 2021, Mahatma et al., 2022, and Brennan et al., 2021). We infer the extent of pre-Stenian sedimentary cover by broadly outlining the extent of these remnant outcrops and subsurface occurrences. The surface outcrop and well-characterized subsurface extent of Stenian syn-Grenville quartzites, including the Middle Run Fm (Clay et al., 2021) in the subsurface of Indiana Ohio, and Kentucky, the Jacobsville Fm (Hodgin et al., 2022) of the Midcontinent, the Hazel and Lanoria Formations (Spencer et al., 2014) of Texas, the Unkar Group (Timmons et al., 2012) of Arizona, the Crystal Springs Formation (Mahon et al., 2014) of California, and distal basins in the Canadian Arctic (Rainbird et al., 2012) are depicted with dark green polygons. We infer the putative extent of syn-Grenville sedimentary cover by outlining the extent of Stenian quartzite outcrops and subsurface occurrences, and extend this polygon to cover a consistent distance inboard of the Grenville front; this distance is a conservative estimate for pro-Grenville sediment extent, given evidence for long-traveled pro-Grenville sedimentation in northwest Canada (Rainbird et al., 2012). The extent of the Laurentian craton (dark gray) and rift positions/extents are modified from Macdonald et al., 2022 and Poole et al., 2005; ages and extents of cratonic basement provinces (italicized labels) are compiled from Clay et al., 2021(a) and Brennan et al., 2021(b).

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