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true true

05 enero, 2021

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1 Introduction

2 Data & Methods

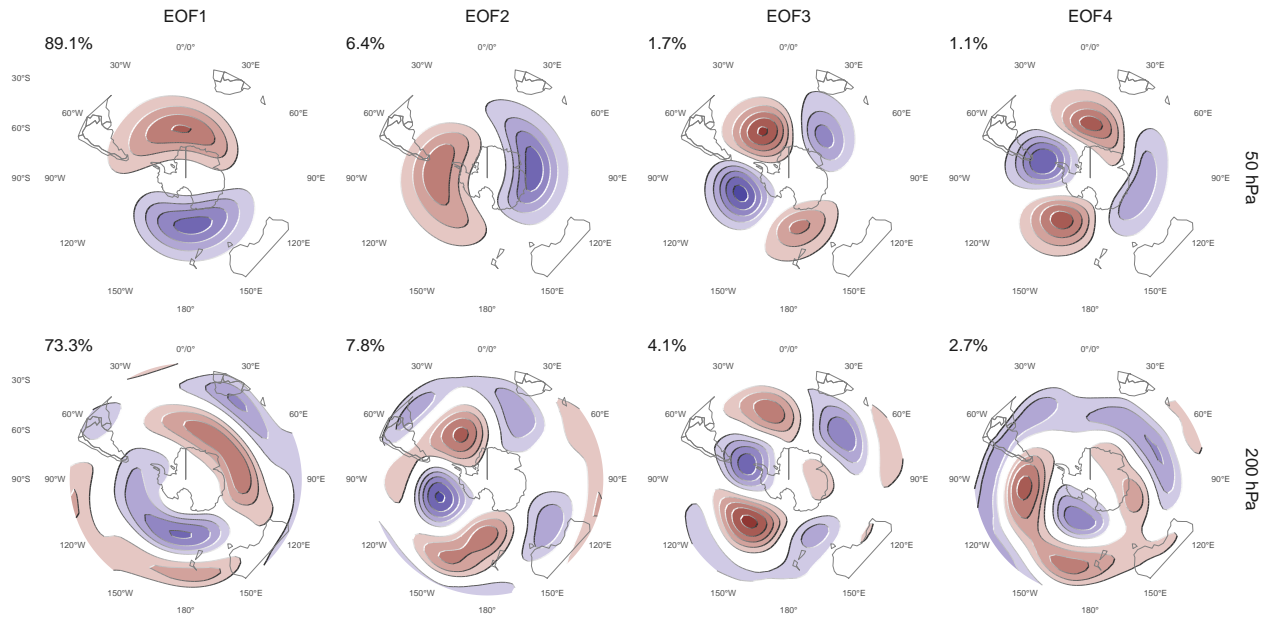
ERA5, ERA5 BE, ERA 20C

If

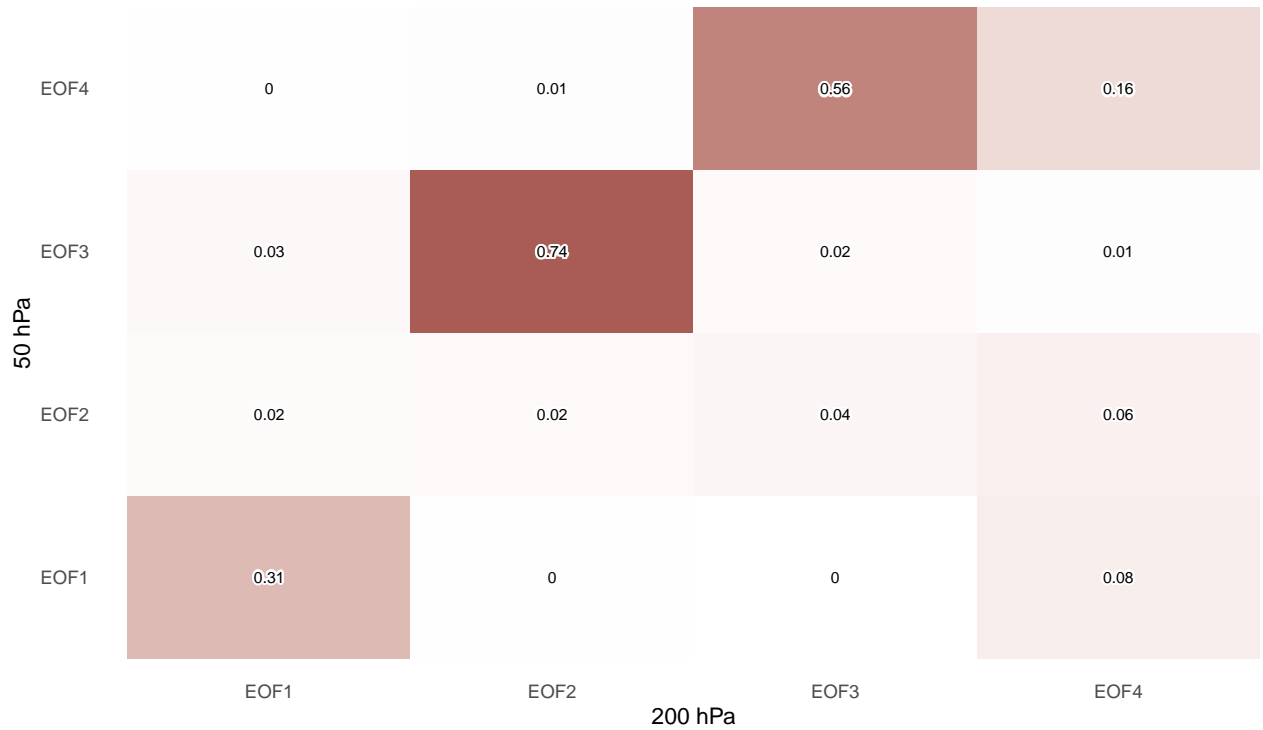
3 Results

Figure 1a show the spatial pattern of the 4 leading EOF of zonal anomalies of geopotential height for 50 hPa and 200 hPa computed separately. The percentages XXX . At 50 hPa, EOF1 and EOF2 on one hand, and EOF3 and EOF4 on the other are clearly pairs of zonal waves with wave numbers 1 and 2-3 respectively shifted by 1/4 wavelength. That is, each group of two EOFs represents the same zonal wave structure that changes in magnitude and location. Similarly, at 200 hPa EOF2 and EOF3 represent the same zonal wave structure with wave number ~3.

Furthermore, the wave 1 pattern represented by the first two EOFs at 50 hPa is similar to the wave 1 pattern shown as the leading EOF at 200 hPa, and the wave 2-3 pattern represented by the third and fourth EOF at 50 hPa is similar to the wave 3 pattern present in EOF2 and EOF3 at 200 hPa. These similarities suggest some level of joint variability across levels. Figure 1b shows the coefficient of determination between temporal series of each EOF at each level. EOF3 and EOF4 at 50 hPa are highly correlated with EOF2 and EOF3 at 200 hPa. Not only each pair of EOFs represent a single structure within each level, but they represent a coherent pattern of variability between levels. Similarly, the leading EOF of each level are relatively highly correlated, again suggesting an unique mode of joint variability.



(a) Spatial patterns (arbitrary units). The numbers at the top-left of each panel is the variance explained by each EOF.



(b) Coefficient of determination of the temporal index of each EOF between levels.

Figure 1: Leading 4 EOFs of zonal anomalies of geopotential height at 50 hPa and 200 hPa for the SON trimester and the period 1979 – 2009.

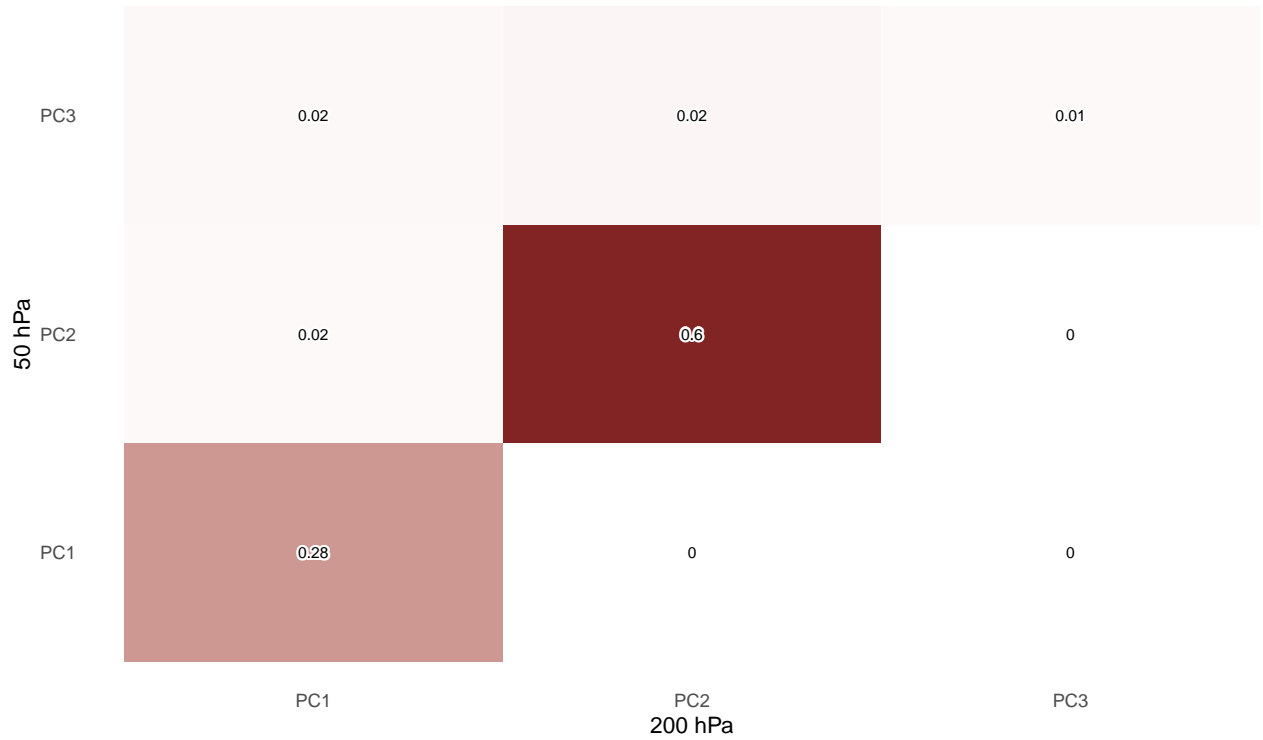
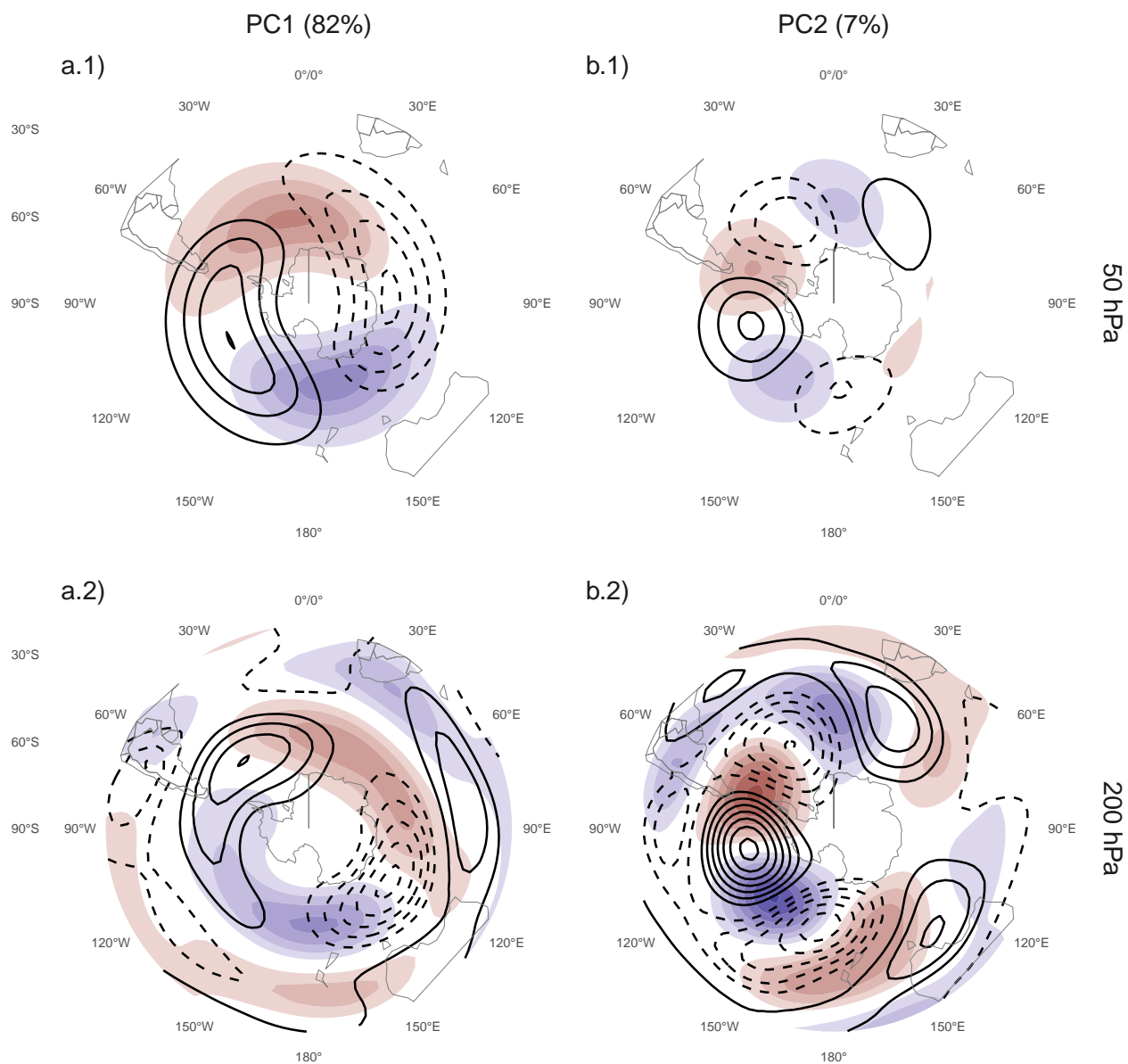
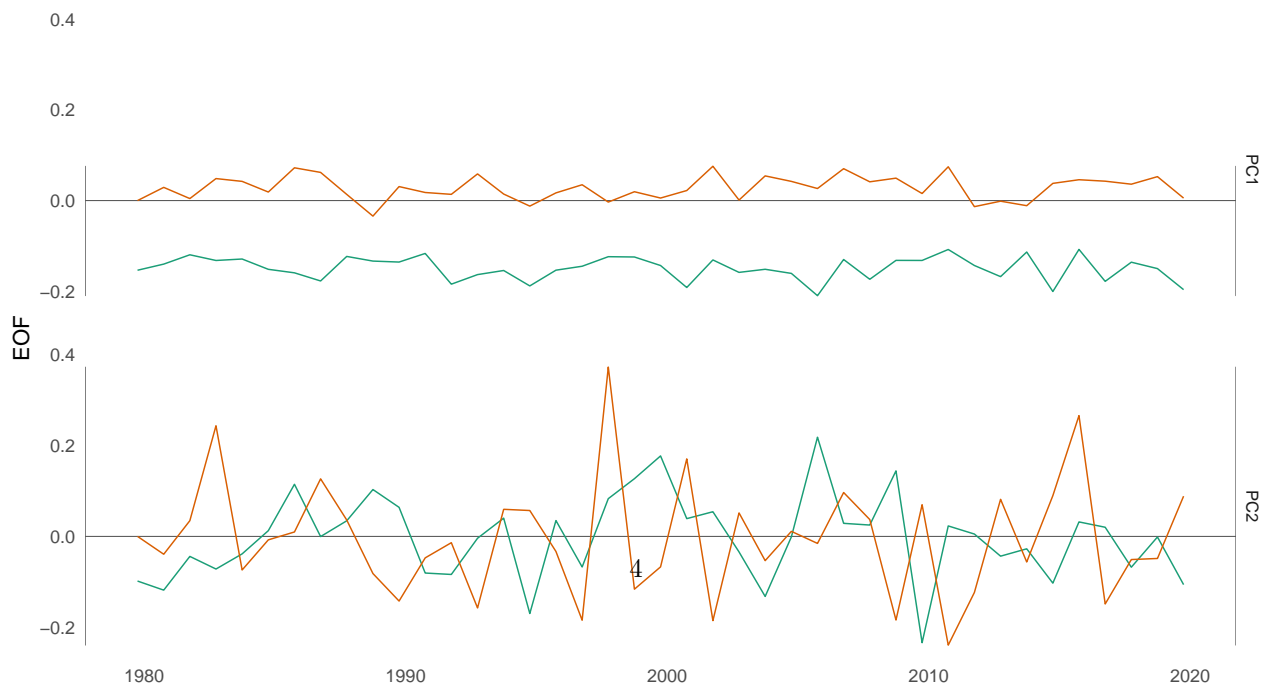


Figure 2: Coefficient of determination of the absolute magnitude of complex EOFs between 200 hPa and 50 hPa computing EOF separately for each level. The high correlation between levels for the two leading EOFs justifies treating the pattern as a mode of covariability between the stratosphere and the troposphere and, thus, computing the EOFs using both levels at the same time. (probably not to be shown.)



(a) Spatial patterns. Real part in shading, imaginary part in contours.



Both observations motivate the decision of performing complex EOF jointly between levels. The computation of the EOFs was carried out using data from both levels at the same time, therefore, each complex EOF has a spatial part that depends on longitude, latitude and level, and a temporal part that depends only on time. Figures 3a and 3b show, respectively, the spatial and temporal parts of the first two leading complex EOFs of zonal anomalies of geopotential height at 50 hPa and 200 hPa. In the spatial patterns in Figure 3a, the real (in shading) and the imaginary parts (in contour) are in quadrature by construction, so that each EOF describes a single wave-like pattern whose amplitude and position (i.e. phase) is controlled by the magnitude and argument of the complex temporal EOF.

The wave patterns described by these complex EOFs match the patterns seen in the *naive* EOFs of Figure @ref(fig: cor-eof-naive-1): The first is a wave 1, while the second is a wave 3. Note that in Figure 3b, the real part of EOF1 has non zero mean. This is due to the fact that the geopotential fields that enter into the algorithm are anomalies with respect to the zonal mean, not the time mean. The real part of the first EOF is, therefore, capturing the mean zonally anomalous field *and the variability that projects onto it*.

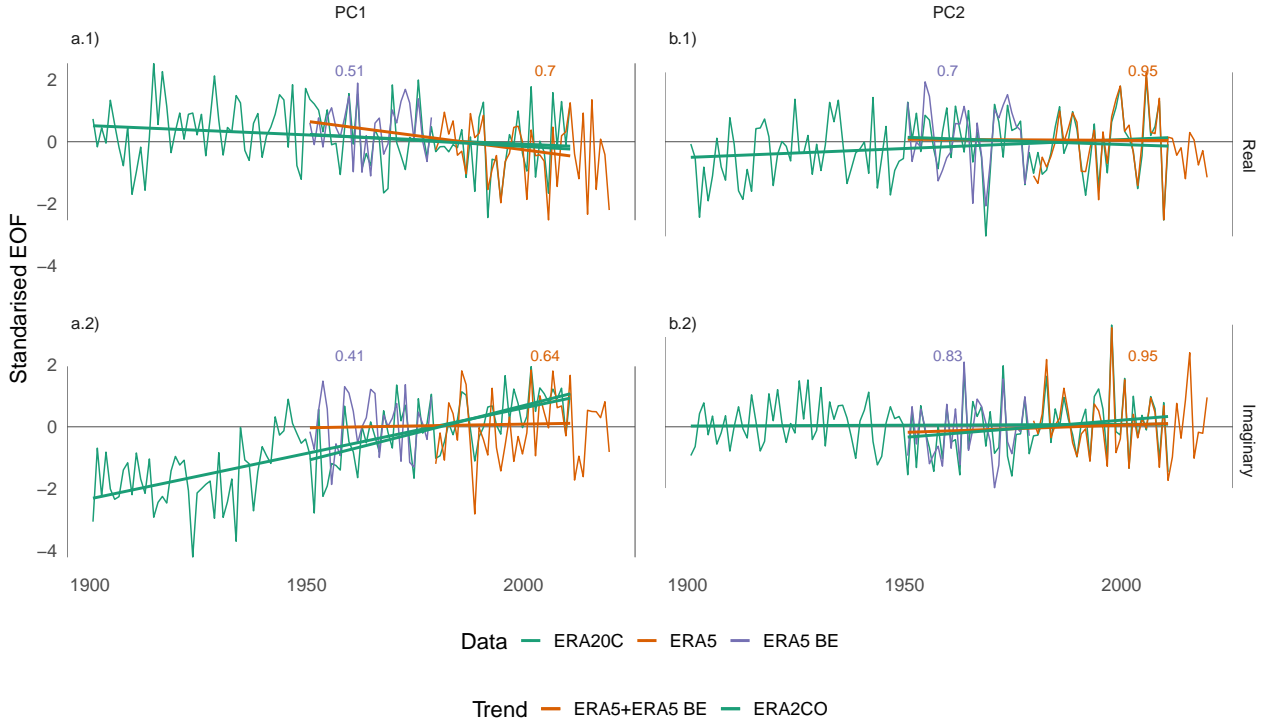


Figure 4: Temporal series extended using ERA5 Back extended preliminary edition (period 1950 – 1978), ERA5 (period 1979 – 2019) and ERA20C (period 1900 – 2010). Each series is computed by projecting monthly geopotential height zonal anomalies standardised by level south of 20°S onto the corresponding spatial pattern. In numbers, the correlation between the ERA20C-based series and the two ERA5-based series. The ERA5 + ERA5 BE and the ERA20C series are centred and scaled using the mean and standard deviation for their period of overlap (1950 – 2010).

Figure 4 shows temporal series of the two complex EOFs extended beyond the satellite era using the preliminary ERA5 back extension going back to 1950 and using ERA20C for the period 1900 – 2010. The hybrid ERA5 + ERA5 BE series and the ERA20C series agree relatively well during the period of overlap, although correlations are higher in the satellite period. The ERA20C time-series has a significant upward linear trend in the imaginary part of EOF1. However, there is no trend in the corresponding hybrid modern dataset during the period of overlap and the correlations between the old and modern series is much poorer. This suggests that the observed trends are not robust.

3.1 Regression

Interpretaciones:

- Parte real de EOF1: variabilidad que se proyecta en la dirección del campo medio
- Parte imaginaria de EOF1: variabilidad asociada a la parte simétrica del SAM
- EOF2: PSA1 y PSA2.
- Parte imaginaria: proyecta sobre la parte asimétrica del SAM (PSA2)
- Parte real: relativamente independiente del SAM

El primer EOF es esencialmente una medida de la variabilidad proyectada en la dirección el campo medio. La parte importante, de todas formas, es la Real.

3.2 Relationship with other indices

En la tropósfera, la parte imaginaria del PC2 captura la parte asimétrica del SAM. La parte real es muy parecida en estructura, pero como está desfasada, no coincide con el SAM y entonces no proyecta fuerte en él.

La Parte imaginaria está muy influenciada por ENSO, como se ve en la relación con el índice y en la regresión con la SST. La fase Real parece no tener influencia del océano tropical. Sin embargo, los flujos de acción de onda.

A su vez, la parte imaginaria del PC1 captura también parte del SAM, pero principalmente la parte simétrica.

3.2.1 Precipitation

4 Conclusions

5 References

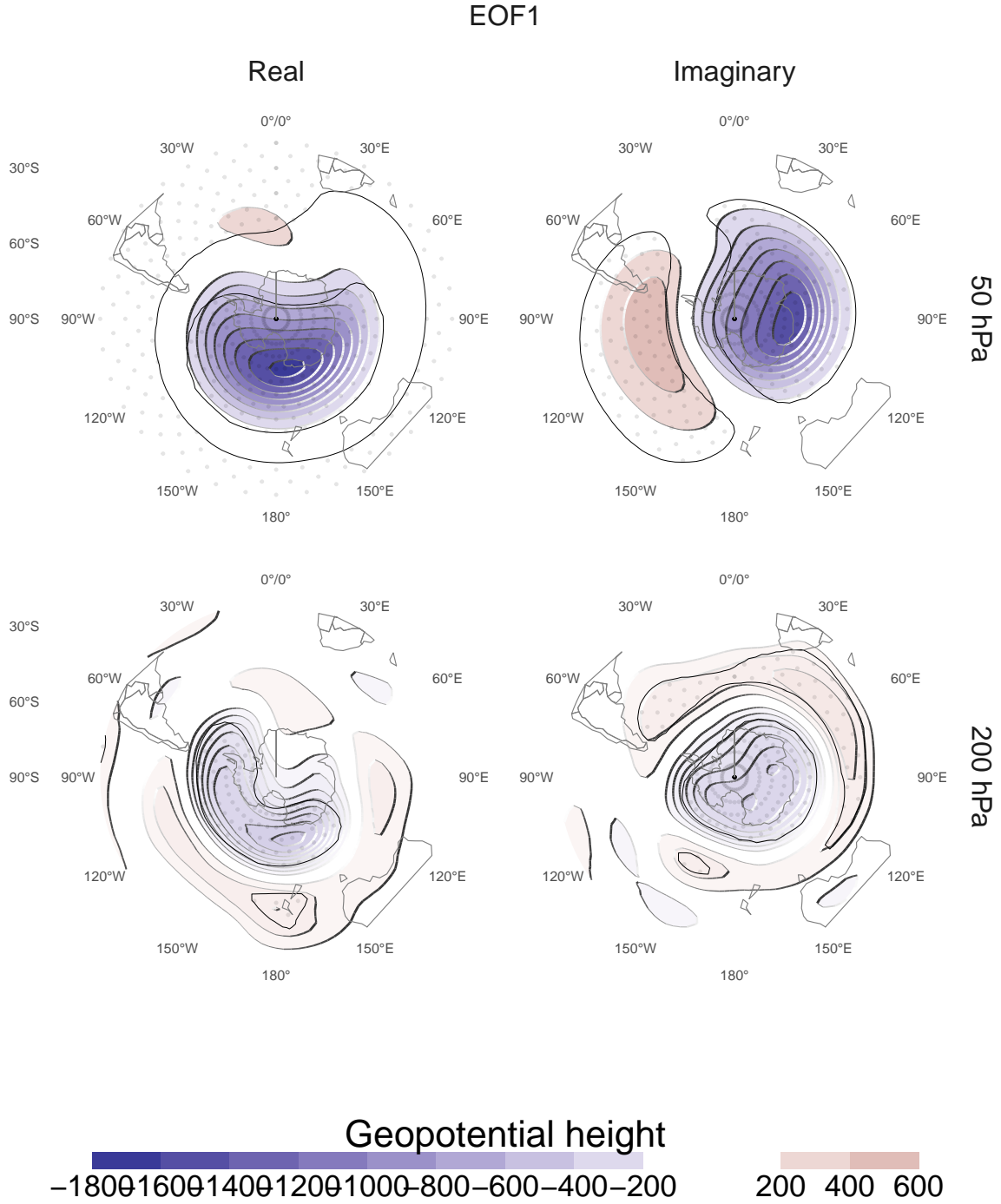


Figure 5: Regression coefficients of the real and imaginary part of the first complex EOF on SON geopotential height for the 1979 – 2019 period. These coefficients come from multiple linear regression involving the real and imaginary part of both EFO2.

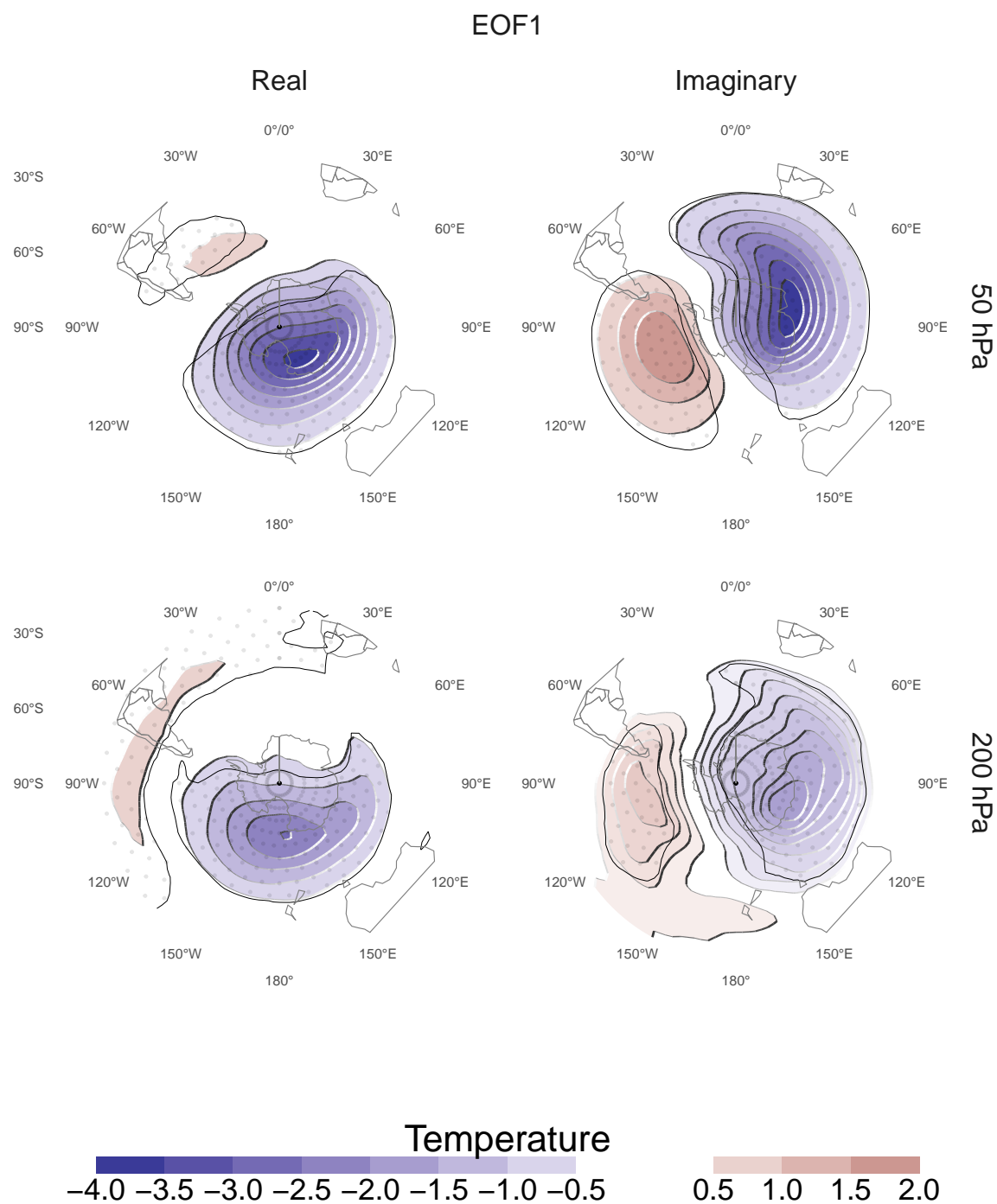


Figure 6: Same as Figure 5 but for air temperature.

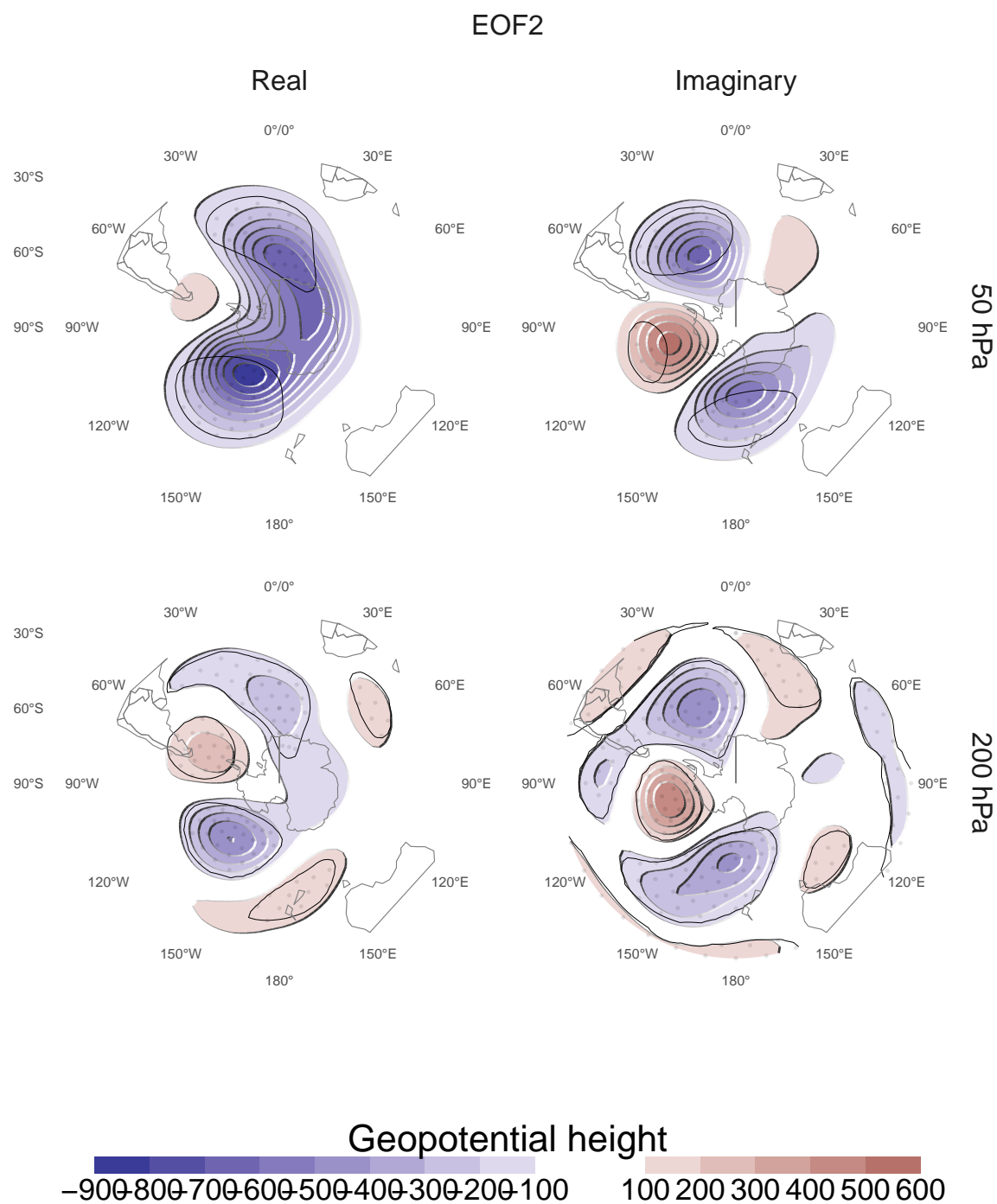


Figure 7: Same as Figure 5 but for the second EOF.

EOF2

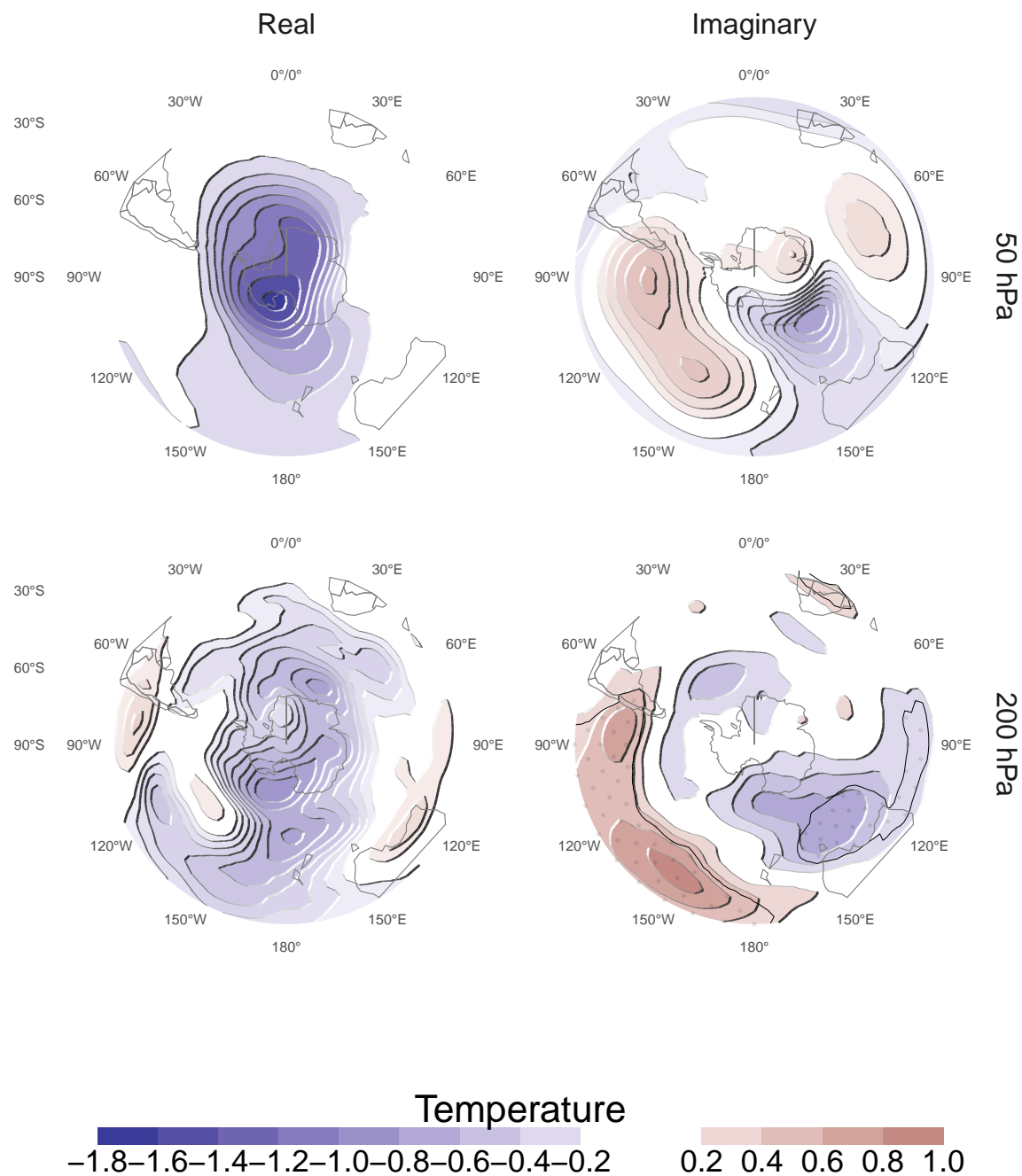


Figure 8: Same as Figure 5 but for the second EOF and temperature.

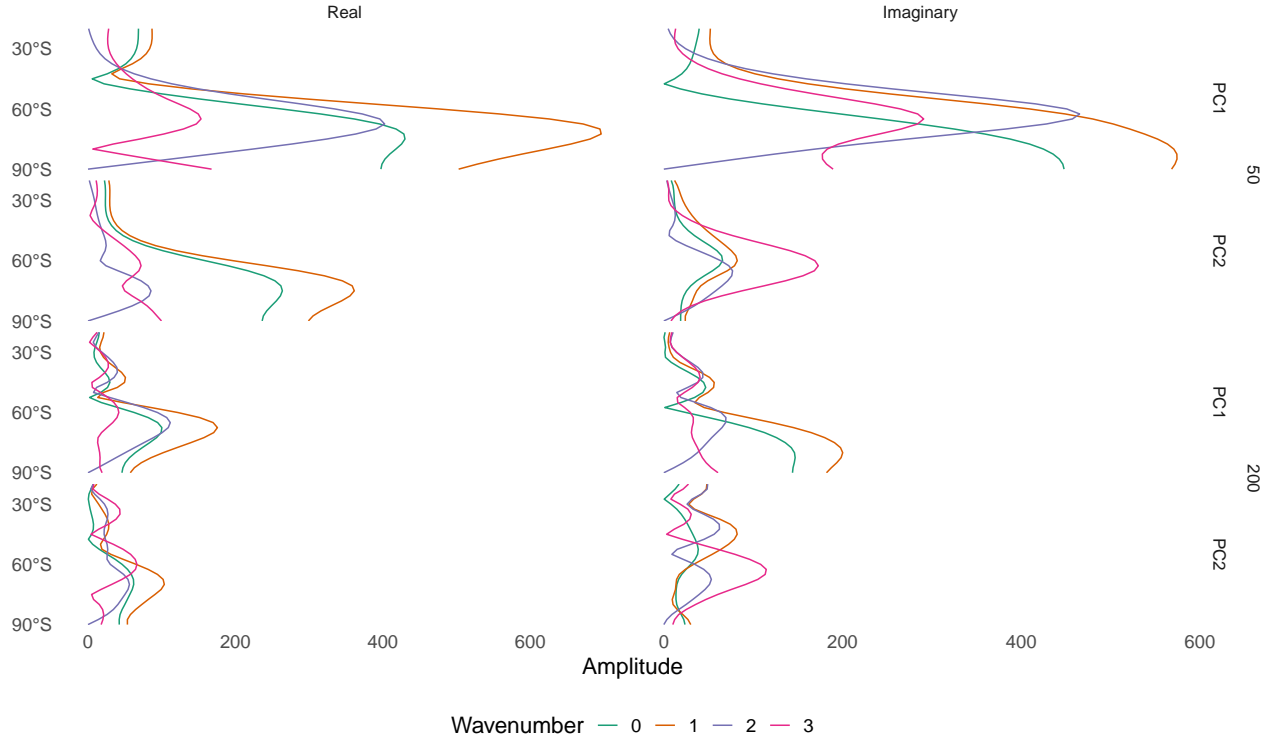


Figure 9: Amplitud de ondas zonales de la regresión con altura geopotencial (figuras anteriores)

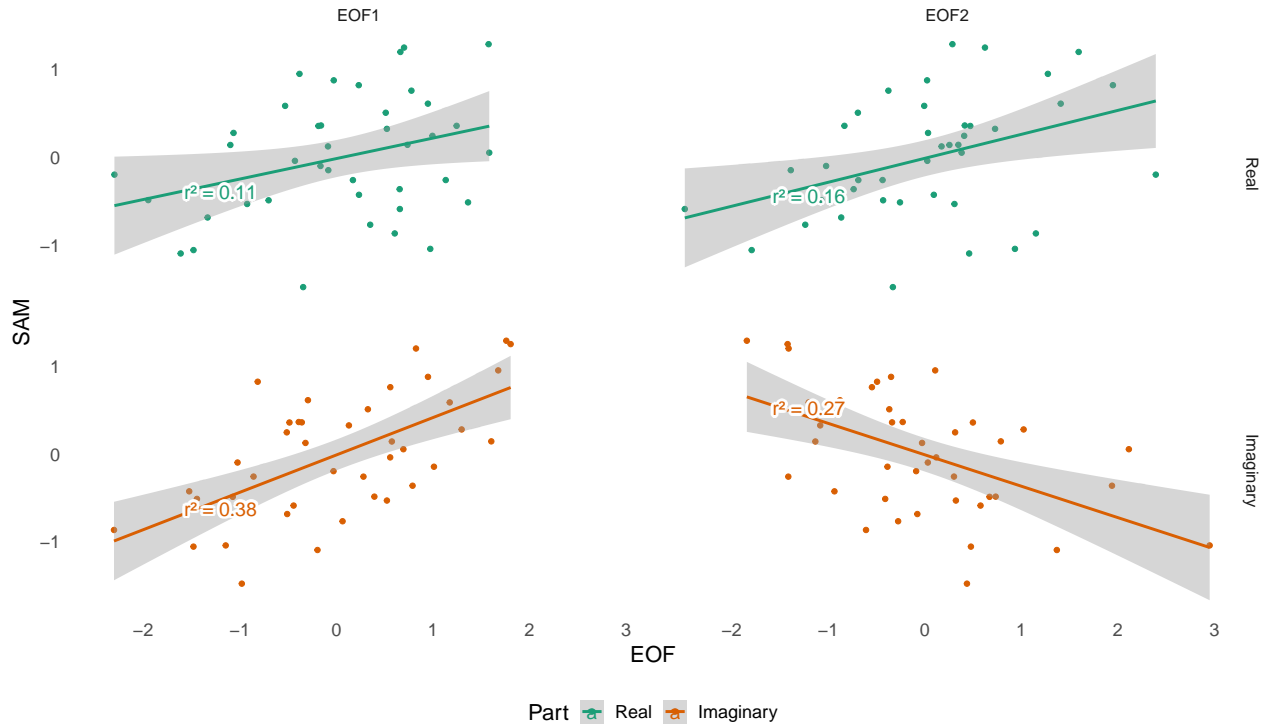


Figure 10: Standardised temporal EOFs and SAM index computed as the leading EOF of geopotential height anomalies at 700 hPa.

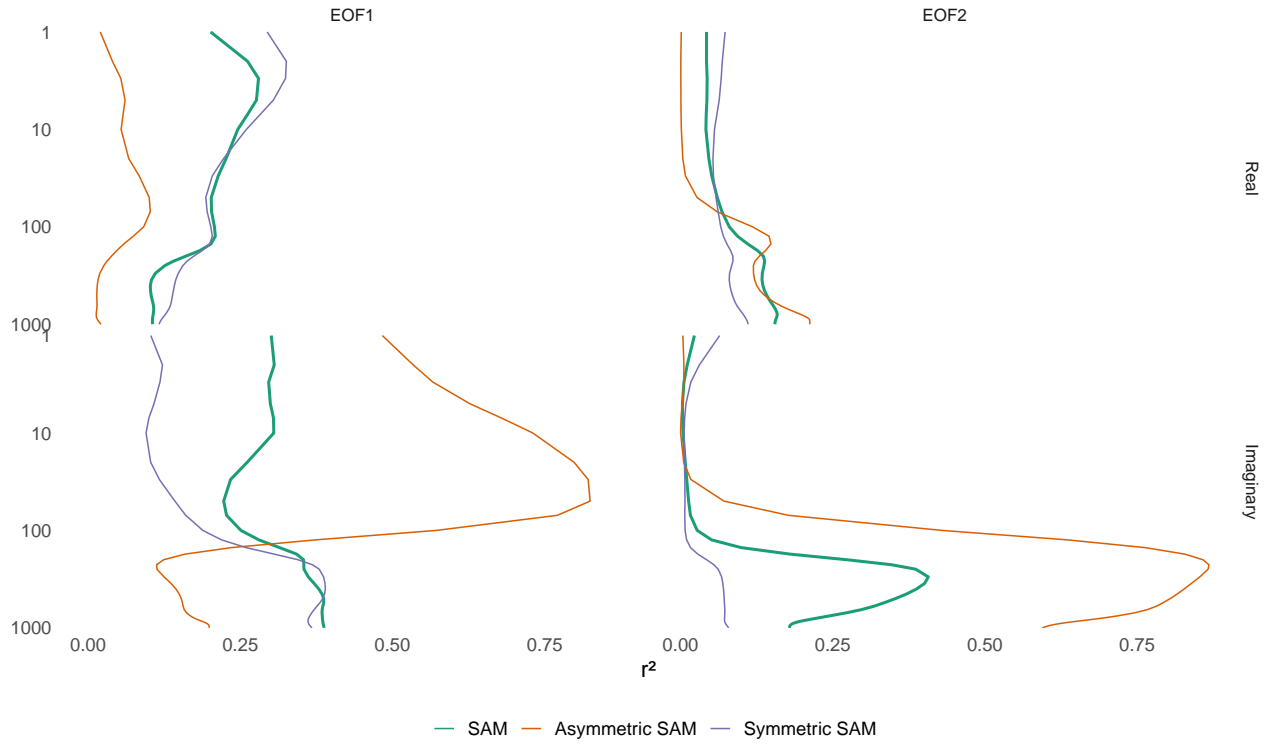


Figure 11: Coefficient of determination between the real and imaginary part of each EOF and the SAM, Asymmetric SAM and Symmetric SAM indices computed at each level according to Campitelli on press??.

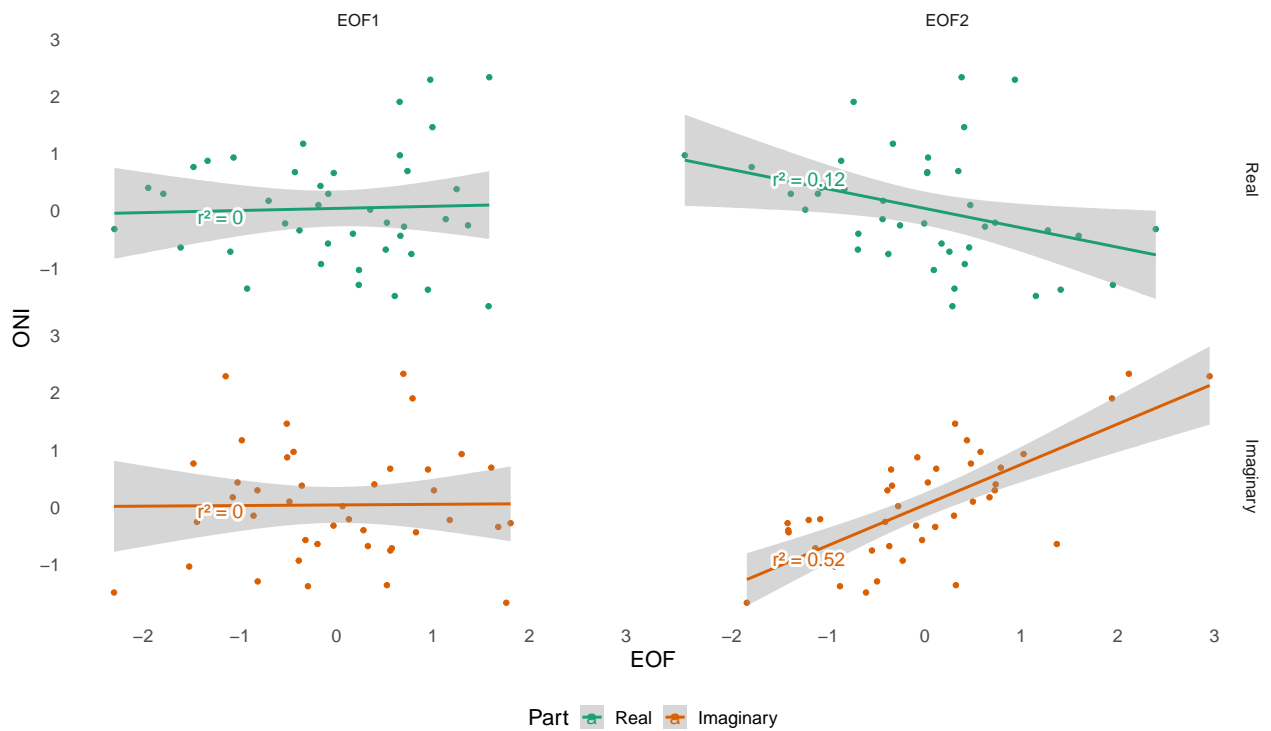


Figure 12: Same as Figure 10 but for the Oceanic Enso Index (REF??).

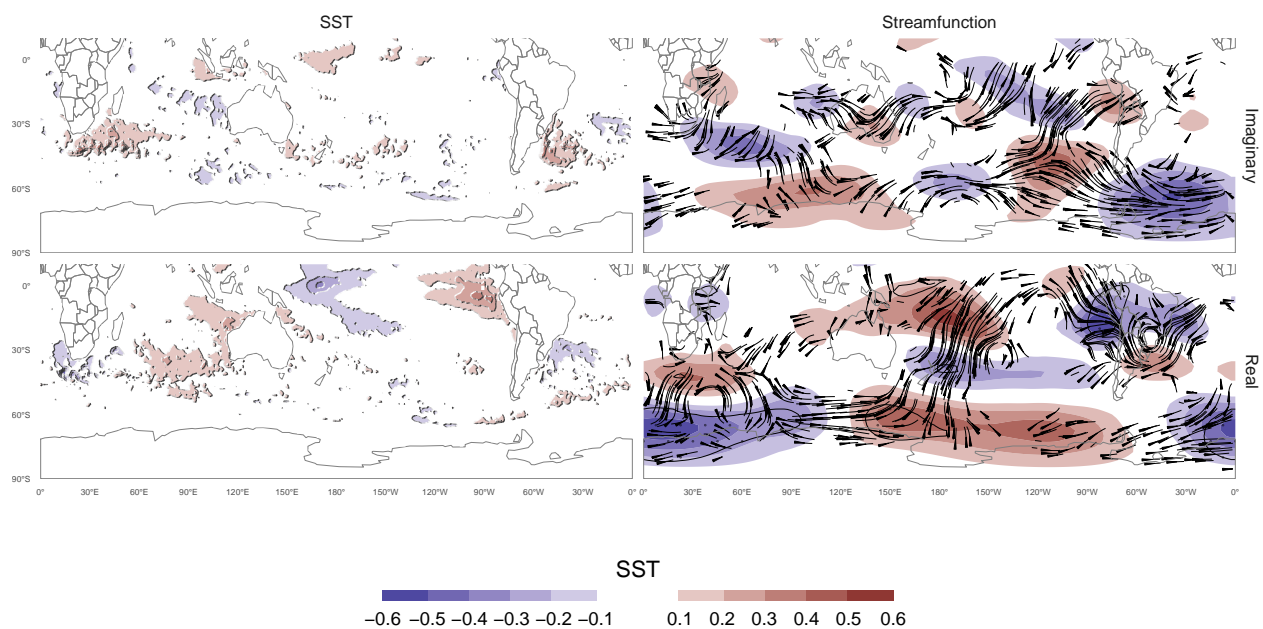


Figure 13: Regresión con PSI en 200 hPa y sst, PC1

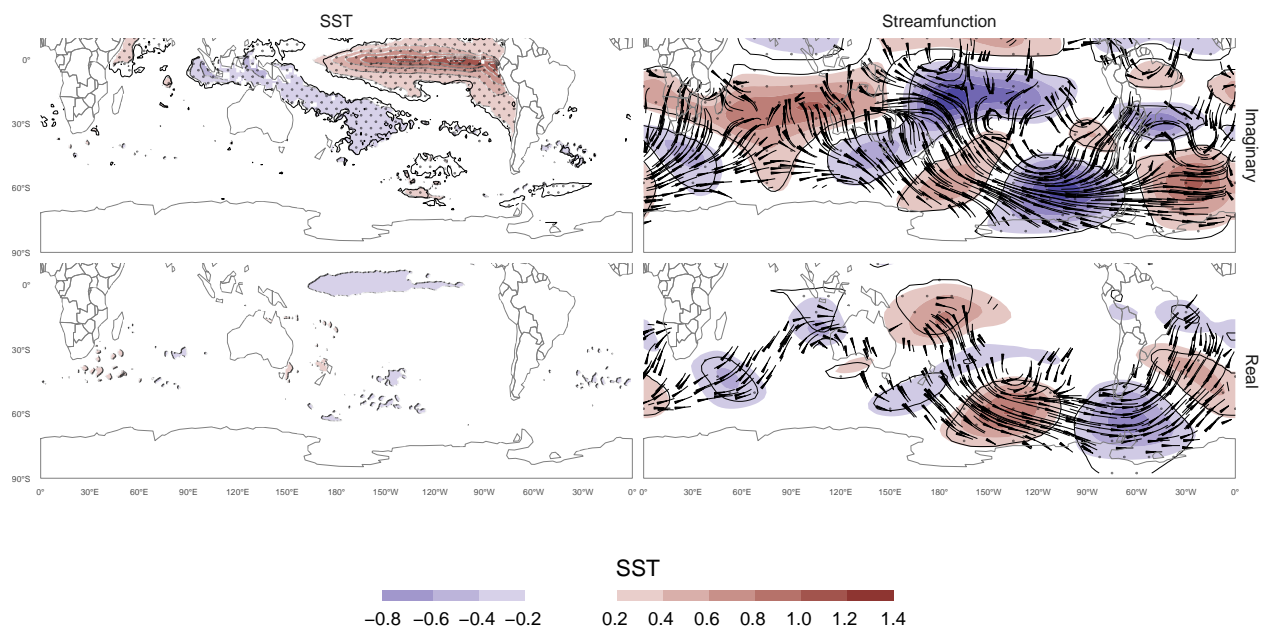


Figure 14: Regresión con PSI en 200 hPa y sst, PC1

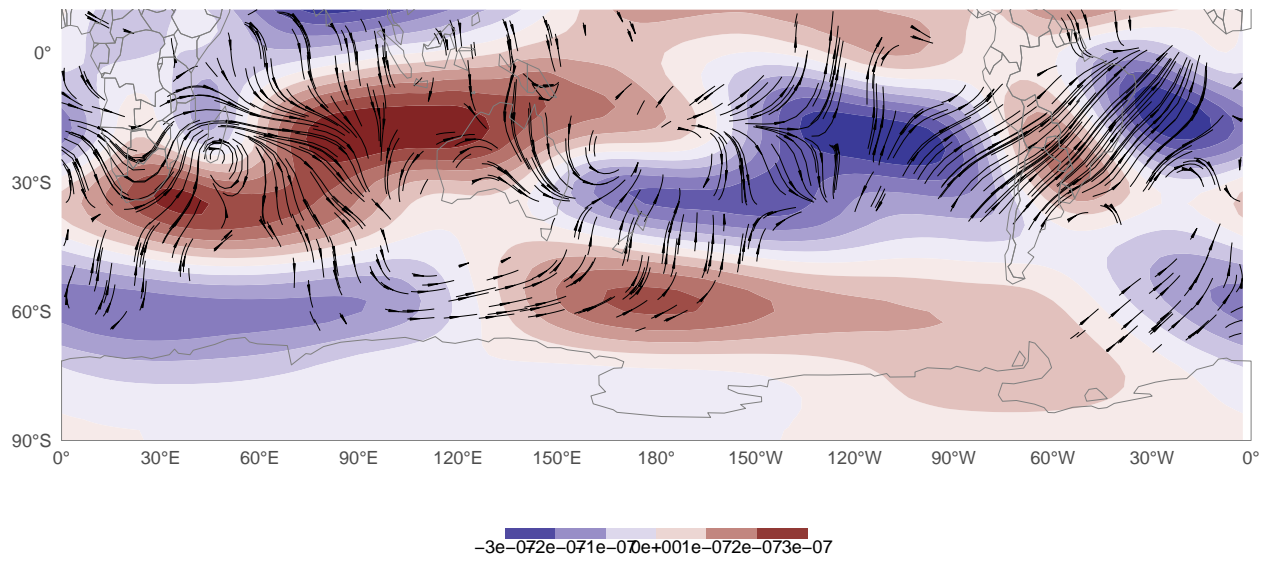


Figure 15: mean streamfunction

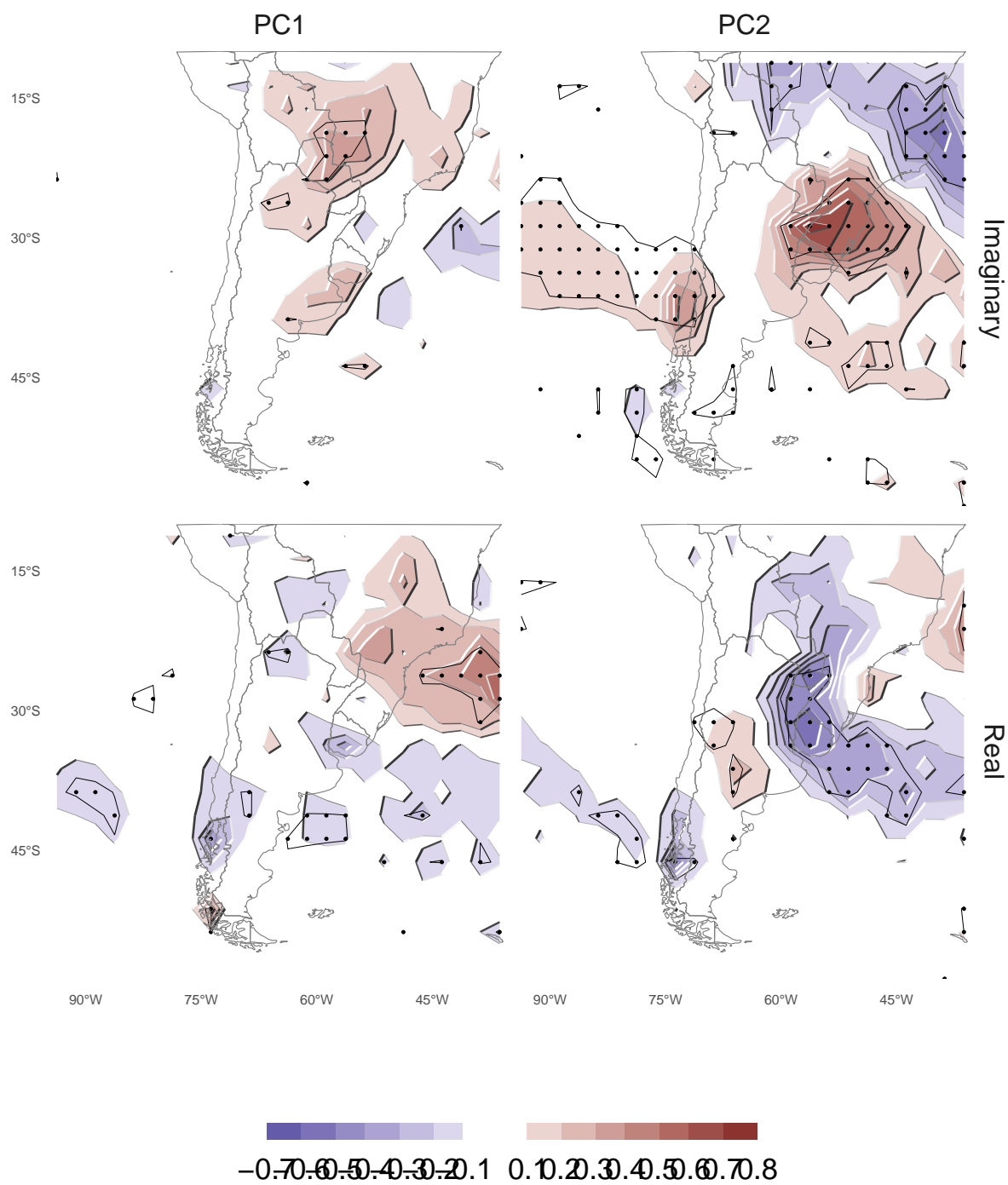


Figure 16: Regresión de cada PC con la precipitación

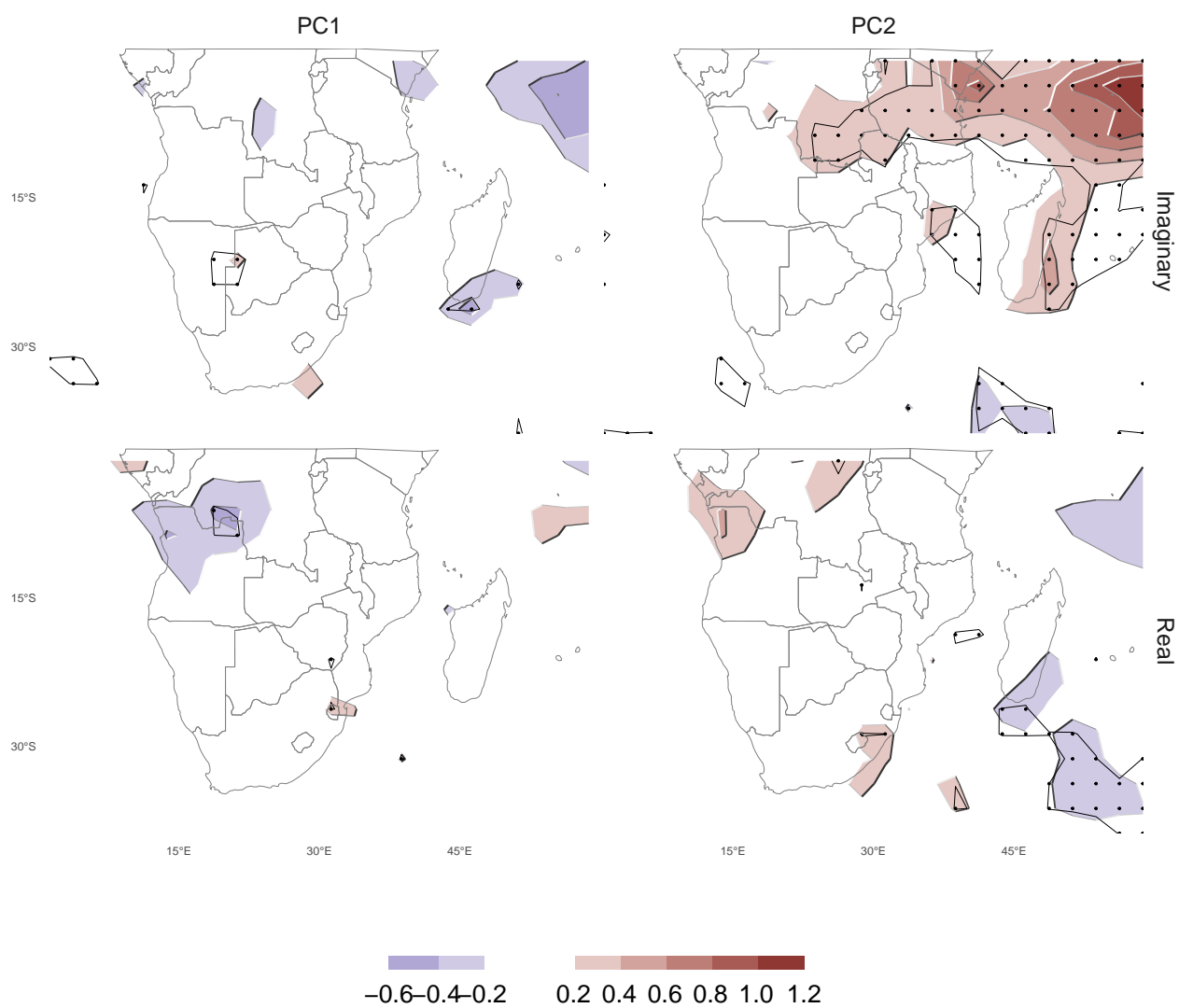


Figure 17: Regresión de cada PC con la precipitación

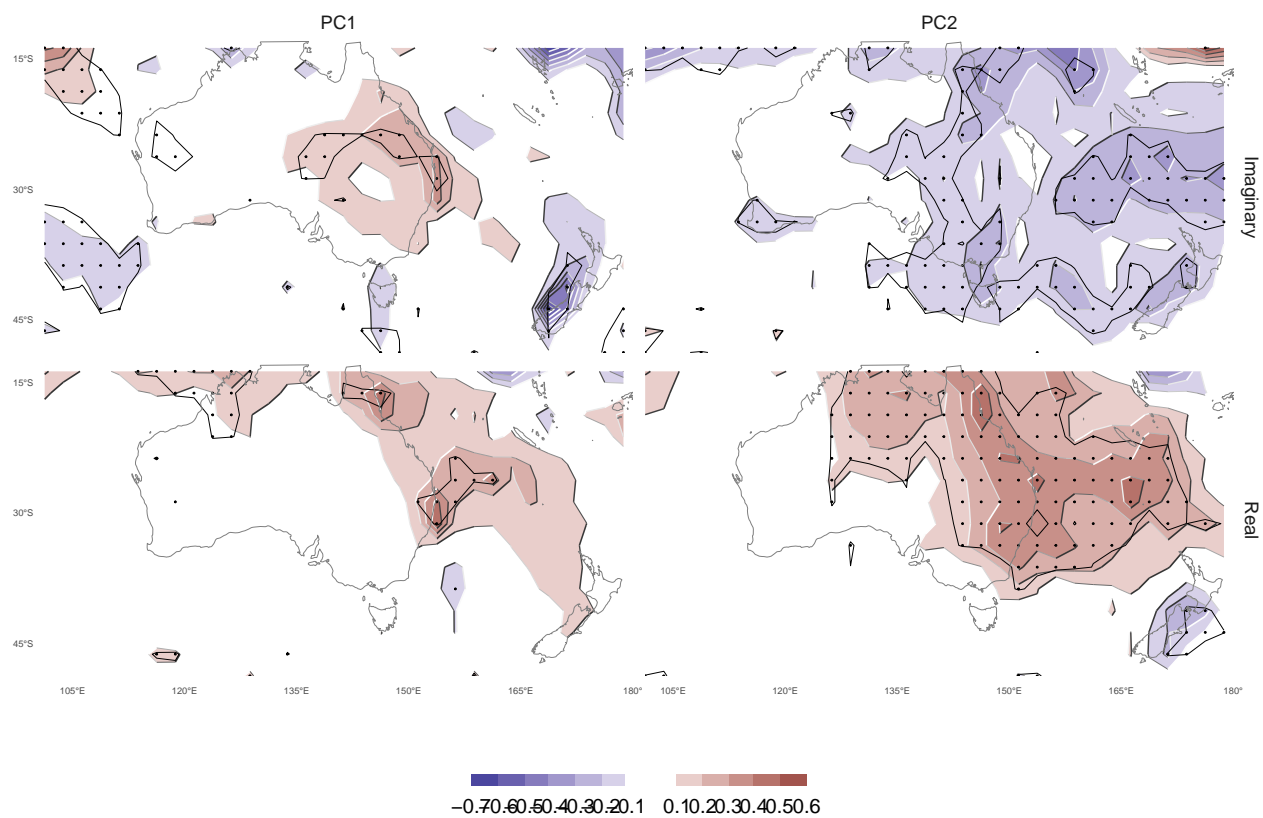


Figure 18: Regresión de cada PC con la precipitación