

Estimating diffusion lengths of water isotopes from ice core data measured by continuous flow analysis

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Abstract.

We examine high-resolution water isotope data sets from continuous flow analysis (CFA) of the West Antarctic Ice Sheet (WAIS) Divide (WDC) and South Pole (SPC) ice cores. Spectral analysis of water isotope data reveals damping of high-frequency variations associated with diffusive smoothing of the isotopic profile in the firn layer of an ice sheet. This diffusion of water isotope ratios in ice cores can provide information about past firn conditions. The data spectra of windowed sections of CFA data from WDC and SPC show different characteristics as compared with similar discretely-sampled data sets, affecting the estimation of high-frequency damping due to the firn diffusion process. The cause of this spectral difference is not known, but we suggest possible explanations in the CFA measurement system and in natural firn processes. Without knowing the origin of this spectral difference, we can still parameterize a method to estimate the extent of diffusion in the spectra. In this study we use two modified techniques to estimate extent of diffusion in order to efficiently and accurately produce diffusion estimates for full ice core CFA data sets.

1. Introduction

To understand past and future climate change, water isotope data from ice cores have long been used as climate proxies, based on the temperature-dependent distillation of water isotope ratios (e.g. $\delta^{18}\text{O}$) in the atmosphere ([*Epstein et al.*, 1951; *Dansgaard*, 1954, 1964]). This indirect method for obtaining temperature records through past glacial and interglacial cycles relies on empirical correlations that only approximate the physical processes involved. A more direct approach uses the signal of isotope diffusion preserved in the ice, overcoming the many issues related to conventional isotope ratios [*Johnsen et al.*, 2000].

1.1. Stable Water Isotope Diffusion

Water isotope diffusion occurs primarily in the firn layer, snowfall in the upper tens of meters of an ice sheet that has yet to be fully compressed into ice. As firn is permeable, water molecules can diffuse through in the vapor phase, damping the seasonal variations and high-frequency noise of the original isotope signal. As the diffusion process depends on temperature, a temperature record can be obtained by measuring the extent of diffusion that has occurred [*Johnsen et al.*, 2000]. This method is independent of conventional assumptions about isotope fractionation before deposition, and thus improves on conventional ice core temperature methods. Beyond information about past temperature, diffusion estimates also provide constraints on past firn conditions and total ice thinning. With these motivations, estimations of diffusion have been made for ice cores in both Greenland and Antarctica [*Simonsen et al.*, 2011; *Gkinis et al.*, 2014; *van der Wel et al.*, 2015; *Jones et al.*, 2017a].

1.2. Continuous Flow Analysis

In the past, most ice core water isotope data have been measured discretely by melting vertical sections of the core. The isotope ratio of each discrete sample is measured on mass spectrometer or laser spectrometer. Recently, ice core water isotope measurements are often measured on continuous flow analysis (CFA) systems. These systems continuously feed the water stable isotopes of the melting core directly into a laser spectrometer, yielding high resolution data more easily than a discrete measurement process [*Gkinis et al.*, 2011; *Emanuelsson et al.*, 2015; *Jones et al.*, 2017b]. Depending on the amount of effort committed, discrete analyses can produce results ranging from one-meter for entire ice cores to 2.5-cm resolution for smaller sections of ice cores. Meanwhile, CFA analyses can easily produce results at half-cm resolution for an entire ice core record. This higher resolution greatly improves the ability to analyze diffusion lengths, which depend on information in the high frequencies of the data spectrum.

A number of studies have estimated diffusion lengths from discretely sampled water isotopes ([*Johnsen et al.*, 2000; *Simonsen et al.*, 2011; *Gkinis et al.*, 2014; *van der Wel et al.*, 2015]). While data spectra from discrete data match well with spectra from theoretically derived synthetic data [*Holme et al.*, 2017], spectra from CFA data have an additional characteristic in the mid-frequency range ($10 - 40$ cycles/m $\approx 2.5 - 10$ cm ice). Figure 1 demonstrates this difference by comparing the power spectra of discretely and continuously measured data. The CFA data spectra have a transition zone of sloping red noise from the damped frequencies into the system noise. The highest resolution discretely-sampled data sets do not show this transition zone in the spectrum.

Since estimating diffusion length depends on the shape of the spectrum, the presence of this transition can affect diffusion results. In this paper we describe how to estimate diffusion lengths on CFA data without the effect of this transition. We start from the approaches developed in *Johnsen et al.* [2000] and *Gkinis et al.* [2014] and propose two adjusted techniques to accommodate CFA water isotope data.

1.3. Water Isotope Data

We use data from the WAIS Divide and South Pole ice cores, both measured at the Institute of Arctic and Alpine Research (INSTAAR) at the University of Colorado. INSTAAR uses a CFA system to measure the water isotope ratios of $\delta^{18}\text{O}$, $\delta^{17}\text{O}$, and δD at a resolution of half-cm. For complete details on the INSTAAR CFA system, see *Jones et al.* [2017b]. Other published discrete and continuous water isotope data sets from Greenland and Antarctic cores presented in *Oerter et al.* [2004]; *Gkinis et al.* [2011]; *Steig et al.* [2013]; *Svensson et al.* [2015]; *Holme et al.* [2017] were used for comparison.

2. Diffusion Theory

The fundamental physics of advection and diffusion can describe mathematically the post-depositional effects on the profile of water isotopes in the ice core [*Johnsen, 1977*]. The majority of water isotope diffusion occurs in the firn layer where interconnected air pathways allow water vapor to diffuse vertically through the firn column. After firn densification has sealed off bubbles in the ice, vapor diffusion ceases and solid ice diffusion takes over. The process in solid ice has a diffusivity orders of magnitude smaller than that of vapor diffusion, and we do not consider it in this study.

The isotopic profile changes through the firn layer of an ice sheet due to the effects of diffusion across isotopic gradients and due to densification of the firn. As stated by *Johnsen* [1977] and subsequently used in several diffusion studies ([*Johnsen et al.*, 2000; *Simonsen et al.*, 2011; *Gkinis et al.*, 2014; *van der Wel et al.*, 2015; *Jones et al.*, 2017a; *Holme et al.*, 2017]), these changes to the isotopic profile can be described by Fick's second law, the basic advection-diffusion equation:

$$\frac{\partial \delta}{\partial t} = D \frac{\partial^2 \delta}{\partial z^2} - \dot{\epsilon} z \frac{\partial \delta}{\partial z} \quad (1)$$

where δ is the isotope ratio, D is the diffusivity coefficient, z is the vertical coordinate assuming an origin fixed on a sinking layer of firn, and $\dot{\epsilon}$ is the vertical strain rate. The coordinate system is fixed to a sinking layer of firn, so the term $\dot{\epsilon} z$ can be thought of as the velocity in this advection-diffusion framework. Considering a layer of firn with $z = 0$ at the vertical midpoint, $\dot{\epsilon} z$ is the speed at which a point distance z from $z = 0$ approaches the origin. A solution for the isotopic profile at time t and at depth z in the firn column is given by:

$$\delta(z, t) = S(t) \frac{1}{\sigma \sqrt{2\pi}} \int_{-\infty}^{\infty} \delta(z, 0) \exp\left(\frac{-(z - u)^2}{2\sigma^2}\right) du \quad (2)$$

where $S(t)$ is the total thinning the layer has experienced due to ice flow from $t = 0$ to $t = t'$:

$$S(t') = \exp\left(\int_0^{t'} \dot{\epsilon}(t) dt\right) \quad (3)$$

For analytical and statistical derivations of the solution in Equation 2, see Appendices A and B.

The amount of smoothing applied to the signal in Equation 2 can be quantified as a diffusion length, or the average vertical distance traveled by a water molecule before it

reaches the bottom of the firn. See Appendix C for a statistical derivation of diffusion length.

3. Estimating Diffusion Length from Data (methods)

With a mathematical understanding of water isotope diffusion, we can match diffusion lengths calculated from ice core data to modeled diffusion lengths to learn about past firn conditions, including temperature, densification, and ice thinning. To make this comparison, we require a method to estimate diffusion lengths from the ice core data. Equation 2 demonstrates that the ice core data, $\delta(z, t)$, is the convolution of the initial isotope signal $\delta(z, 0)$ with a Gaussian filter of standard deviation σ , the diffusion length:

$$\mathcal{G} = \frac{1}{\sigma\sqrt{2\pi}} \exp\left(\frac{-z^2}{2\sigma^2}\right) \quad (4)$$

Rather than solve this convolution to find the diffusion length σ , we take the Fourier transform to simplify from convolution to multiplication:

$$\delta(z, t) = \delta(z, 0) * \mathcal{G} \quad \Rightarrow \quad \hat{\delta}(z, t) = \hat{\delta}(z, 0)\hat{\mathcal{G}} \quad (5)$$

where $*$ represents the convolution and $\hat{}$ represents the Fourier transform. Furthermore, the Fourier transform of a Gaussian remains a Gaussian:

$$\mathfrak{F}(\mathcal{G}) = \hat{\mathcal{G}} = \exp\left(\frac{-k^2\sigma^2}{2}\right) \quad (6)$$

where k is wavenumber and σ is still the diffusion length.

With the assumption that the initial isotope signal $\delta(z, 0)$ can be approximated as white noise [Gkinis *et al.*, 2014], we can simply fit a Gaussian curve in frequency space to the data spectrum and solve for its standard deviation. The standard deviation σ_{std} of the

Gaussian in frequency space is related to the diffusion length σ by:

$$\sigma = \frac{1}{2\pi\sqrt{2}\sigma_{std}} \quad (7)$$

Appendix D derives this conversion factor. Repeating this method over consecutive windowed sections of data yields diffusion length estimates through the length of a core.

Fitting the data spectrum is complicated when working with CFA data (refer back to Figure 1). *Jones et al.* [2017a] avoids the effect of this transition zone noise by identifying the frequency at which the transition begins to affect the damping curvature of diffusion. The technique cuts off higher frequencies, and only includes lower frequencies in the Gaussian fit. Figure 2 shows examples of this cut-off technique for data from the WAIS Divide ice core (WDC) [*Jones et al.*, 2017b] and from the South Pole ice core (SPC). This technique relies on the fact that the damping signal left by natural diffusion is not wiped out by this transition zone noise.

In this study we follow the fitting approach of *Gkinis et al.* [2014], which uses a least-squares technique to fit the summation of two functions (P_s) to the spectra of ice core data (\mathbb{P}_s) by varying the four parameters P_0 , σ^2 , a_1 and σ_η^2 :

$$P_s = P_0 e^{-k^2 \sigma^2} + \frac{\sigma_\eta^2 \Delta}{|1 - a_1 \exp(-ik\Delta)|^2}, \quad (8)$$

where a_1 is the AR-1 coefficient, σ_η^2 is the variance of the noise signal and Δ is the sample size. The first function is a Gaussian, representing the high-frequency damping of firn diffusion, and the second function is an autoregressive noise signal of order 1 (AR-1), representing baseline noise introduced by measurement and post-depositional processes. Adding together these two functions avoids having to choose a cut-off frequency for the fit. Since it does not require any subjective choices, this multi-function technique can

be fully automated and produce results in less time and with less effort than the cut-off technique.

The multi-function technique has been used effectively to estimate diffusion lengths on many discretely sampled data sets [Gkinis *et al.*, 2014; Holme *et al.*, 2017]. It has been used on short sections of CFA data sets, but in some cases required extra boundary conditions in the parameterizations. However, when applying this technique to newer CFA data sets with a higher signal to noise ratio, the transition zone in the CFA spectra affects the technique's ability to effectively fit a Gaussian to the data. Figure 3 illustrates this issue with examples of CFA spectra from WDC and SPC. The presence of the transition zone force both the Gaussian and noise functions to attempt to accommodate the transition zone. The Gaussian function is pulled to a greater standard deviation, corresponding to a smaller diffusion length, and the noise level is raised slightly at lower frequencies. The poor fit of the data results in a poor representation of diffusion length.

4. Improving the fit for CFA data

In order to best fit the spectra of CFA data, we examined two different approaches. The first approach removes the transition zone by masking it with white noise, allowing the spectra to be fit effectively by two functions. The second approach improves the multi-function technique by including an extra function in the parameterization. Both techniques are introduced and tested on ice core data from WDC and SPC.

4.1. Technique 1: Adding white noise

The main difference between the new CFA data and the discrete data is the resolution and precision. This difference appears in the data spectra by lowering the measurement

white noise baseline for the high frequencies in the CFA data compared to the measurement baseline of the discrete data (refer to a figure). Previous studies of water isotope diffusion on both discrete and CFA data sets have not had the same precision and resolution as the measurements of WDC and SPC, but despite their lower precision, studies have still made sufficiently good fits [*Gkinis et al.*, 2014; *Holme et al.*, 2017]. One solution is to mask the transition zone in the CFA data. This shift is achieved by adding Gaussian-distributed noise in the time domain, which increases the white noise level in the frequency domain. This addition masks the transition zone and the power spectrum is now similar to that of older CFA and discrete measurements. Fig. 4 illustrates the effect of this technique.

We constructed a sensitivity test to quantify the best choice of noise level. The test uses a 500-year moving window. For each windowed section, 100 diffusion lengths are estimated by adding white noise with an increasing baseline to the data. A diffusion length estimate is available for each noise level as a function of age. By adding too much noise to the data we risk masking the climate signal. Hence, the optimal noise level is reached when the diffusion length estimates stop changing with increasing noise, when the gradient approaches 0. Fig 5 shows this gradient in diffusion length with respect to added noise. For WDC, adding a mean noise level of 0.4 ‰ for $\delta^{18}\text{O}$ and of 3.0 ‰ for δD seems to be the optimal solution.

4.2. Technique 2: Parameterizing a Multi-Function Fit

The second approach creates a parameterization of the total spectrum. In order to best fit the spectra of CFA data, we try three different parameterizations of a multi-function fit of all frequencies. In addition to the Gaussian that represents firn diffusion and the

autoregressive function that represents the baseline noise in Eq. 8, we add a third function to allow a better overall fit. We try three different functions and see which yields the best fits. First, we add a second AR-1 function, second, we add a second Gaussian curve, and third we add a folded normal distribution (FND) [*Tsagris et al.*, 2014]:

$$P_{FND} = P_{0_{FND}} \cdot e^{-(k \cdot \sigma_{FND})^2} \cdot |1 - \Phi(-ik\sigma_{FND})|^2, \quad (9)$$

where $\Phi(\sigma_{FND}) = 1/2 \cdot \text{erfc}(-\sigma_{FND}/2)$. Here P_0 and σ_{FND} are the two parameters that are being estimated from the fitting. The justification of fitting a folded normal distribution is that it reflects the uni-directional combined memory and diffusion induced by the one-way flow of the CFA system. Our goal here is not to explain physically what the transition zones represents, but simply to find the optimal parameterization to fit the data spectrum. The optimal parameterization will allow the first Gaussian to better fit the lower-frequency part of the spectrum, and thus to get a better estimate of diffusion length compared to Figure 3.

Figure 6 shows the parameterization that sums a Gaussian curve and two autoregressive noise functions to create a total fit for the data spectrum. With the inclusion of this third function in both WDC and SPC, the total fit is visually improved, as compared to the single-Gaussian fit in Figure 3. Figure 7 shows the parameterization that sums two Gaussian curves and one autoregressive noise function. Similarly, there is a visual improvement in the total fit as compared to the single-Gaussian fit. Finally, Figure 8 shows the parameterization that sums one Gaussian, one autoregressive noise function, and a folded normal distribution function.

5. Evaluating fitting techniques

5.1. Goodness of fit

The different fitting procedures in each technique are evaluated by calculating the adjusted goodness of fit (\bar{R}^2) between the spectra and parameterizations. The \bar{R}^2 is a goodness of fit estimation (R^2) that takes into account the number of fitted parameters (p):

$$\bar{R}^2 = 1 - (1 - R^2) \frac{n - 1}{n - p - 1}, \quad (10)$$

where n is the sample size. The \bar{R}^2 values enables a comparison between the parameterizations that use six fitting parameters with the parameterization that uses five parameters (the addition of white noise). The \bar{R}^2 values are plotted as a function of age in Figure 9. It is evident that all the parameterizations provide good fits to the data, but the best fits are obtained with the extra Gaussian and the FND. For simplicity and computational efficiency, we prefer the Gaussian curve over the FND.

5.2. Comparison of techniques

We use the WDC record to compare the diffusion length estimates between both techniques. In Figure 10, the diffusion length estimates are plotted with respect to age. Both methods reconstruct similar diffusion lengths, but the noise-adding method seems to be more stable in the glacial period. This stability is attributed to the fact that the measured water isotope signal seems to contain several noisy sections of half a meter or so in the glacial period, which will affect the shape of the spectra. By adding white noise to the data, the noisy data sections are masked, making the fit more stable.

Another way of validating the results is by comparing with the diffusion lengths estimated by the cut-off technique presented in *Jones et al.* [2017a]. While the cut-off

technique requires the choice of cut-off frequency, it avoids the effect of the transition zone because that part of the spectrum is not included in the fit. Since these results are not affected by the transition zone, we can use them as a benchmark with which to validate these new techniques. Figures 11 and 12 show these comparisons. The estimated diffusion lengths from this paper are very similar to the presented results in *Jones et al.* [2017a]. A few differences stand out in the comparison, such as the peak at an age of around 12,000 years before present. Likely, these differences were not found in *Jones et al.* [2017a] due to the lower resolution of that diffusion length record.

6. Application Example: South Pole Glacial-Interglacial Temperature

The water isotope data set from the South Pole ice core provides new information about climate in central Antarctica. From previous ice cores in the climatically distinct East Antarctica and West Antarctica, we have temperature estimates based on conventional assumptions about Rayleigh distillation. However, model estimates do not match with temperature estimates. The South Pole site spatially bridges the gap between those climatically distinct regions and diffusion-based temperature estimates will provide new insight into this model-data discrepancy.

We use both techniques described above to estimate the glacial-interglacial temperature change at the South Pole. A complete water isotope data set is not yet available from the South Pole ice core, but we can use two 20-meter windows of data, from the Holocene and Late Glacial, respectively, to make a diffusion-based estimate of the temperature change at this site. Using the adding-noise method, we calculate a change of about XX degrees, and using the multi-function parameterization method, we calculate a change of about 8 degrees C, as shown in Fig ???. These methods agree well with one another, and provide

the first insight the glacial-interglacial change at the South Pole. With the complete water isotope data set in the future, we will be able to carefully calculate a full temperature history, rather than only a temperature difference, to learn more about the climate in central Antarctica and how it does or does not agree with model simulations.

7. Discussion

Figure ?? and Table ?? demonstrate that for the CFA data sets for WDC and SPC, the noise-adding technique and the double-Gaussian parameterization of the multi-function parameterization technique work better than the single-Gaussian formulation used in previous studies. The cut-off fitting technique provides validation of the fits, but is overall less efficient than the multi-function technique and requires a subjective choice of cut-off value.

While including a second Gaussian in the multi-function technique improves fit, it remains to be explained how to physically interpret the red noise in the transition zone of the spectra. Possible origins of this transition include noise generated in the CFA measurement system, noise generated naturally in the firn, or some combination of both. In this section we explore each of these possibilities and how to gain insight into the origin of this noise.

7.1. Possible Noise Origin: CFA Measurement System

There are many possible sources of noise throughout the CFA measurement system that could contribute to the data spectrum between periods of 5 to 10 cm. Mixing and memory effects are known to occur throughout the system as sample water travels to the instrument through tubing and various reservoirs [*Gkinis et al.*, 2011]. For WDC, *Jones*

et al. [2017b] ran standards of ice through the CFA system and reported system-caused diffusion lengths of 0.7 cm and 0.8 cm for $\delta^{18}\text{O}$ and δD , respectively. However, system diffusion is unlikely to be the cause of the transition noise because, mathematically, it is expected to increase the total diffusion described by the main Gaussian, rather than add this sloping transition in the 5 to 10 cm period range.

A second possible source of this red noise in the CFA system is the Picarro instrument. *Gkinis* [2011] showed that the Picarro measurement can be affected by variations in water concentration within the instrument cavity. The INSTAAR CFA system has been carefully calibrated and set up to ensure water concentrations remain at a level that does not affect the measurement. However, small fluctuations in water concentration could still add noise to the data even if they do not change the isotope values significantly. While INSTAAR has carefully calibrated to address this issue, the process could be more complex than we understand and remain important.

Errors in depth registration within the CFA system could produce noise in this frequency range. We tested this possibility by adding random noise to a depth series and inspecting the resulting frequency spectra. This test showed that this source of noise does/does not remain a possibility. (Section still in progress)

(Get rid of following paragraph?) Figure ?? demonstrates the variability through the core of the diffusion length associated with the second Gaussian. The CFA system remains constant through the analysis of the entire core, and thus any effects solely due to the system would be expected to remain constant as well. This reasoning suggests that, according to Figure ??, there is a significant contribution from climate variation to the transition noise. However, inspection of individual spectra shows that there is an implicit

dependence of the diffusion length of the second Gaussian on that of the main Gaussian. Figure ?? shows that with a greater natural diffusion length, the transition zone is shifted toward lower frequencies and thus the diffusion length of the second Gaussian is greater as well. Conversely, with a smaller natural diffusion length, the transition zone is shifted toward higher frequencies and the second Gaussian diffusion length is also smaller. Therefore the transition noise may be affected directly by climate influence on the lower frequencies in the spectrum.

7.2. Possible Noise Origin: Natural Firn Diffusion Processes

Another possibility is that the red noise transition originates naturally in the firn rather than in the CFA system during the measurement. This effect could be explained by considering a more complex model of water isotope diffusion. In the *Johnsen et al.* [2000] model, all water molecules are treated identically, experiencing the same amount of time in the vapor phase relative to that spent in the solid ice phase. In reality, different molecules experience different amounts of time in the vapor phase relative to the solid phase. Some molecules may be trapped inside ice grains throughout their advection down the firn column, while others may remain on the surface of grains, allowing for many transitions back and forth to the vapor phase. Previous estimates (Johnsen and Whillans and Groote) claim that the lifetime of grains is short enough that no grains exist long enough to trap molecules and significantly affect the bulk diffusion length. However, if the lifetime of grains is longer than original estimates, this process could prevent all molecules from spending the same amount of time in the vapor phase, and thus could affect the diffusion length.

The *Johnsen et al.* [2000] model has been shown to work well in capturing a bulk diffusion length by fitting a Gaussian curve to the data spectrum, but the added complexity of allowing a range of individual diffusion lengths could contribute to the red noise transition in the higher frequencies. Mathematically, the *Johnsen et al.* [2000] model matches a diffusion length and corresponding Gaussian curve to the bulk water isotopes in the firn in a given window of data. A model that allows individual molecules to experience different amounts of time in the vapor phase, within some range, corresponds to a range of Gaussian curves. Each Gaussian curve is weighted by the number of molecules that experienced that amount of diffusion and the collection of weighted Gaussian curves sums to a non-Gaussian filter. In theory, this collection is infinite as it represents an infinite range of possible amounts of time spent in the vapor phase. In practice, this infinite collection can be approximated by a finite collection. As shown above in the multi-function parameterization, perhaps as few as two Gaussian curves could sufficiently represent the infinite summation.

To gain more insight into the origin of the transition zone noise, we could make new ice core measurements. To isolate CFA system effects, an approach is to make double measurements, continuous and discrete, on ice from SPC. Measuring the same meters of ice discretely at half-cm resolution as well as continuously on the CFA at the same resolution would provide a comparison to highlight where the noise originates. This would give us two data sets of the same resolution to compare the effects of the CFA directly to discretely sampled data.

8. Conclusions

In this study we examined the diffusion of water isotope data from the WAIS Divide and South Pole ice cores measured on the CFA system at INSTAAR. We observed that spectra from these CFA data, unlike spectra from comparable discretely sampled data, have a unique transition zone in the mid-frequencies. We found that the most effective ways to estimate diffusion lengths on these CFA data sets were the noise-adding technique and the double-Gaussian parameterization of the multi-function fitting technique. These methods are efficient in terms of time and effort required, and also effectively fit the data spectra with the transition zone.

The origin of the noise in the transition zone is unknown. It could be caused by effects in the CFA measurement system, by effects in natural diffusion in the firn, or by some combination of both. Further work is required to isolate the origin.

Appendix A: Analytical Derivation of Diffused Isotope Profile

Here is a complete derivation of the desired solution that I found in the slides of an MIT course “Mathematical Methods for Engineers” from 2006, which I found at <http://ocw.mit.edu/courses/mathematics/18-086-mathematical-methods-for-engineers-ii-spring-2006/readings/am54.pdf>. I’ve written up the steps as outlined by the MIT course, but added in many more detailed steps that are omitted on the course slides.

This derivation starts with the heat equation, with diffusivity D , and therefore is still not quite the solution I seek, which would also include an advection term. Equation (47)

starts off with the basic heat equation.

$$\frac{\partial u}{\partial t} = D \frac{\partial^2 u}{\partial x^2} \quad (\text{A1})$$

We can solve this PDE using the common technique of separation of variables. We assume that the solution $u(x, t)$ can be written as the product of two functions, one that depends only on t and one that depends only on x :

$$u(x, t) = G(t)E(x) \quad (\text{A2})$$

We can take derivatives of $u(x, t)$ this solution and rewrite equation (47) as

$$G'E = GE'' \quad (\text{A3})$$

$$\frac{G'(t)}{G(t)} = \frac{E''(x)}{E(x)} \quad (\text{A4})$$

Since these ratios are equal, they must be equal to some constant, and we can find a family of solutions that will satisfy equation (50):

$$E(x) = De^{ikx} \quad E''(x) = -Dk^2e^{ikx} \quad (\text{A5})$$

$$G(t) = e^{-k^2t} \quad G'(t) = -k^2e^{-k^2t}$$

check solutions:

$$\frac{G'(t)}{G(t)} = -k^2$$

$$\frac{E''(x)}{E(x)} = -k^2$$

Now we can write our solution as

$$u(x, t) = De^{ikx}e^{-k^2t} \quad (\text{A6})$$

Now account for all solutions from all linear combinations by integrating over all k :

$$u(x, t) = \frac{1}{2\pi D} \int_{-\infty}^{\infty} \hat{u}_0(k) e^{ikx} e^{-k^2t} dk \quad (\text{A7})$$

Here $\hat{u}_0(k)$ is included to satisfy the initial conditions $u(x, 0)$ at $t = 0$.

check:

$$u(x, 0) = \frac{1}{2\pi D} \int_{-\infty}^{\infty} \hat{u}_0(k) e^{ikx} dk \quad (\text{A8})$$

Using the definition of the inverse Fourier transform:

$$x(t) = \frac{1}{2\pi D} \int_{-\infty}^{\infty} \hat{x}(\omega) e^{i\omega t} d\omega \quad (\text{A9})$$

Comparing the above, we find the right hand side is exactly equal to our arbitrary initial condition:

$$u(x, 0) = u(x, 0) \quad (\text{A10})$$

So we have successfully derived the general solution as shown in equation (53).

To derive the fundamental solution for a specified initial condition, let us first consider an initial condition that is a delta function, or a point source: $u(x, 0) = \delta(x)$ at $t = 0$. This is a nice initial condition because its Fourier transform is $\hat{u}_0(k) = 1$.

Plugging this initial condition into equation (53):

$$u(x, t) = \frac{1}{2\pi D} \int_{-\infty}^{\infty} e^{ikx} e^{-k^2 t} dk \quad (\text{A11})$$

Next take the partial derivative with respect to x of both sides:

$$\frac{\partial u}{\partial x} = \frac{1}{2\pi D} \int_{-\infty}^{\infty} ik e^{ikx} e^{-k^2 t} dk \quad (\text{A12})$$

Now rearrange to separate for integration by parts:

$$\frac{\partial u}{\partial x} = \frac{1}{2\pi D} \int_{-\infty}^{\infty} (ie^{ikx}) (ke^{-k^2 t}) dk \quad (\text{A13})$$

where the first grouping will be u and the second grouping will be dv . To compute the integration by parts we have:

$$\text{D R A F T} \quad u = ie^{ikx} \quad v = -\frac{1}{2t} e^{-k^2 t} \quad \text{May 17, 2017, 10:25am}$$

D R A F T

$$\frac{du}{dk} = -xe^{ikx} \frac{dv}{dk} = ke^{-k^2t}$$

From the equation for integration by parts,

$$\int_{-\infty}^{\infty} u dv = (uv)_{-\infty}^{\infty} - \int_{-\infty}^{\infty} v du \quad (\text{A14})$$

we get:

$$\frac{\partial u}{\partial x} = \frac{1}{2\pi D} \left[\left(-ie^{ikx} \frac{1}{2t} e^{-k^2t} \right)_{-\infty}^{\infty} - \int_{-\infty}^{\infty} \frac{1}{2t} e^{-k^2t} x e^{ikx} dk \right] \quad (\text{A15})$$

The term on the left inside the square brackets goes to zero when evaluated from $-\infty$ to ∞ , and we are left with:

$$\frac{\partial u}{\partial x} = -\frac{1}{4\pi t D} \int_{-\infty}^{\infty} e^{-k^2t} x e^{ikx} dk \quad (\text{A16})$$

Now compare equation (58) to the expression for $u(x, t)$ in equation (54) to see the simple result

$$\frac{\partial u}{\partial x} = -\frac{xu}{2t} \quad (\text{A17})$$

This is now a linear ODE that is solved by:

$$u = ce^{-x^2/4tD} \quad (\text{A18})$$

We can solve for c based on conservation of the initial condition through time:

$$\begin{aligned} \int_{-\infty}^{\infty} u(x, t) dx &= \int_{-\infty}^{\infty} u(x, 0) dx \\ &= \int_{-\infty}^{\infty} \delta(x) dx \\ &= 1 \end{aligned}$$

Apply this conservation criterion to solve for c :

$$\int_{-\infty}^{\infty} ce^{-x^2/4tD} dx = 1 \quad (\text{A19})$$

$$c = \frac{1}{\int_{-\infty}^{\infty} e^{-x^2/4tD} dx}$$

This is the integral of a Gaussian, which is known:

$$\int_{-\infty}^{\infty} e^{-ax^2} dx = \sqrt{\frac{\pi}{a}} \quad (\text{A21})$$

Here $a = \frac{1}{4tD}$, and thus we can solve for c as

$$c = \frac{1}{\sqrt{4tD\pi}} \quad (\text{A22})$$

So now our solution to equation (60) can be written

$$u(x, t) = \frac{1}{\sqrt{4\pi tD}} e^{-x^2/4tD} \quad (\text{A23})$$

This is the fundamental solution from a single point source (recall that our initial condition used in this solution was a single delta function).

Now let us consider the possibility that our initial condition delta function is instead located a different point $x = s$, such that

$$u(x, 0) = \delta(x - s) \quad \text{at} \quad t = 0 \quad (\text{A24})$$

Then the argument of the exponential in our solution shifts by s , i.e.

$$e^{-x^2/4tD} \quad \text{becomes} \quad e^{-(x-s)^2/4tD} \quad (\text{A25})$$

Because of linearity, any initial condition $u(x, 0)$ can be written as the combination of point sources:

$$u(x, 0) = \int_{\text{all } S} \delta(x - s) u(s, 0) ds \quad (\text{A26})$$

The solution for an initial condition extending over all points in x can be written as an integral of the responses to $\delta(x - s)$:

$$u(x, t) = \frac{1}{\sqrt{4\pi tD}} \int_{-\infty}^{\infty} u(s, 0) e^{-(x-s)^2/4tD} ds \quad (\text{A27})$$

And finally we have the solution we have been looking for. The one exception is that this solution does not include the $S(t)$ term that accounts for the thinning of the ice because this solution was derived from the pure diffusion equation, without accounting for any advection.

Appendix B: Statistical Derivation of Diffused Isotope Profile

Lasaga derives the same solution through a statistical framework, starting with the idea of a discrete random walk. A particle starts at $z = 0$ and at each time step can move a distance L either to the right with probability p or to the left with probability $q = 1 - p$. We would like to know what the probability is that after N steps the particle is at some position $z = mL$, where $-N \leq m \leq N$.

Let us define the number of steps the particle takes to the right, n_R , and the number of steps the particle takes to the left, n_L . With these definitions we can write:

$$N = n_R + n_L \qquad m = n_R - n_L \qquad (\text{B1})$$

and thus:

$$n_R = \frac{1}{2}(N + m) \qquad n_L = \frac{1}{2}(N - m) \qquad (\text{B2})$$

We can write the probability, $P_N(m)$ of arriving at position $z = mL$ after N steps as the product of the probability of taking a particular sequence of steps to that position, $p^{n_R}q^{n_L}$, times the number of different sequences of steps that will end at that position (because multiple sequences of steps will end at the same z position, i.e. RRRL, RRLR, RLRR, and LRRR all end at $m = 2$):

$$P_N(m) = p^{n_R}q^{n_L} \frac{N!}{n_R!n_L!} \qquad (\text{B3})$$

Plugging in equations (5) and (6),

$$P_N(m) = \frac{N!}{[\frac{1}{2}(N+m)]![\frac{1}{2}(N-m)]!} p^{n_R} q^{n_L} \quad (\text{B4})$$

We can now make a couple of assumptions to simplify equation (8). First, we can assume that $p = q = \frac{1}{2}$, that the particle is equally likely to step to the left or right. Second, we can assume that the particle has taken many steps ($N \gg 1$) and that the number of steps taken is much greater than the distance from the starting point ($N \gg m$). With these assumptions, we can use Stirling's Approximation:

$$\ln n! = n \ln n - n + \frac{1}{2} \ln n + \frac{1}{2} \ln 2\pi \quad (\text{B5})$$

Rewrite equation (8) with this approximation and set $p = q = \frac{1}{2}$ to get:

$$\begin{aligned} \ln P_N(m) = & N \ln N - \frac{1}{2}(N+m) \ln\left[\frac{1}{2}(N+m)\right] \\ & - \frac{1}{2}(N-m) \ln\left[\frac{1}{2}(N-m)\right] + \frac{1}{2} \ln N - \frac{1}{2} \ln\left[\frac{1}{2}(N+m)\right] \\ & - \frac{1}{2} \ln\left[\frac{1}{2}(N-m)\right] - \frac{1}{2} \ln(2\pi) + N \ln \frac{1}{2} \end{aligned} \quad (\text{B6})$$

Simplify with the following:

$$\begin{aligned} \ln(N+m) &= \ln \left[N \left(1 + \frac{m}{N} \right) \right] \\ &= \ln N + \ln \left(1 + \frac{m}{N} \right) \end{aligned} \quad (\text{B7})$$

For small x,

$$\ln(1+x) \sim x - \frac{1}{2}x^2 \quad (\text{B8})$$

and thus,

$$\ln \left(1 + \frac{m}{N} \right) = \frac{m}{N} - \frac{1}{2} \frac{m^2}{N^2} \quad (\text{B9})$$

So, equation (11) becomes:

$$\ln(N + m) = \ln N + \frac{m}{N} - \frac{1}{2} \frac{m^2}{N^2} \quad (\text{B10})$$

and, similarly,

$$\ln(N - m) = \ln N - \frac{m}{N} - \frac{1}{2} \frac{m^2}{N^2} \quad (\text{B11})$$

Plugging equations (13) and (14) into equation (10) and simplifying, we get:

$$\ln P_N(m) = -\frac{1}{2} \ln N + \frac{1}{2} \ln 2 - \frac{1}{2} \ln \pi - \frac{m^2}{2N} + \frac{m^2}{2N^2} \quad (\text{B12})$$

Neglecting the term of order $\frac{1}{N^2}$ and combining the remaining terms leaves us with:

$$P_N(m) = \left(\frac{1}{\pi N} \right)^{1/2} e^{-m^2/2N} \quad (\text{B13})$$

Equation (16) shows the Gaussian curve expected from diffusion from a point source, based only on the assumptions of a random walk and of many time steps.

This equation describes only the discrete probability of the particle taking steps of size L . If we want an expression for the continuous probability $W(z, t)$ that the particle will be at any point z at time t , we must generalize m to z and N to t .

Define the relations:

$$z = mL \quad N = vt \quad (\text{B14})$$

where v is the frequency of steps.

Because the particle can only reach an even numbered position at an even time step or an odd numbered position at an odd time step (i.e. N and m are both even or both odd), $P_N(m)$ is the discrete probability of the particle being anywhere between $z = mL$ and $z = (m + 2)L$. Thus the discrete probability can be related to the continuous probability

as:

$$P_N(m) = W(z, t)2L \quad (\text{B15})$$

Using this relation and plugging in known expressions:

$$\begin{aligned} W(z, t) &= \frac{P_N(m)}{2L} \\ &= \frac{1}{2L} \left(\frac{2}{\pi N} \right)^{1/2} e^{-(m^2/2N)} \\ &= \frac{1}{2L} \left(\frac{2}{\pi vt} \right)^{1/2} e^{-[(m^2/L^2)/2vt]} \\ &= \frac{1}{(2\pi L^2 vt)^{1/2}} e^{-(z^2/2vL^2 t)} \end{aligned} \quad (\text{B16})$$

And with the definition $D \equiv \frac{1}{2}vL^2$,

$$W(z, t) = \frac{1}{2\sqrt{\pi Dt}} e^{-(z^2/4Dt)} \quad (\text{B17})$$

We now have an expression for the continuous probability of particles diffusing to a particular point at a particular time that originate from a point source. However, we would like to know what this solution will look like for a continuous source with spatial extent. Because the diffusion equation is linear, it obeys the law of superposition, and we can generalize our point source solution by treating a continuous source as the sum of individual source slabs of arbitrarily small width dz .

If our point of interest is located at z' and one of the slabs making up the continuous source is located at z , we can write the contribution of this single slab to the concentration of particles at point z' as

$$c(z', t) = \frac{c_0}{2\sqrt{\pi Dt}} \exp\left(-\frac{(z' - z)^2}{4Dt}\right) dz \quad (\text{B18})$$

where c_0 is the initial concentration at point z .

We can sum the contributions from all the slabs making up the continuous source by integrating the above equation over the z values of the entire source. For a source of infinite extent we get:

$$c(z', t) = \int_{-\infty}^{\infty} \frac{c_0}{2\sqrt{\pi Dt}} \exp\left(-\frac{(z' - z)^2}{4Dt}\right) dz \quad (\text{B19})$$

Defining the diffusion length $\sigma \equiv \sqrt{2Dt}$, pulling the constants outside the integral, and replacing c with δ , z' with z , and z with u , we see that we have arrived at the solution presented by Johnsen in equation (2):

$$\delta(z, t) = \frac{1}{\sigma\sqrt{2\pi}} \int_{-\infty}^{\infty} \delta(z, 0) \exp\left(-\frac{(z - u)^2}{2\sigma^2}\right) du \quad (\text{B20})$$

The only difference between the above and equation (2) is the inclusion of the thinning function $S(t)$, which is included as a correction for the ice thinning, which has not yet been taken into account by this statistical derivation. Otherwise, following Lasaga's statistical approach, we have now derived the same diffusion solution as presented in Johnsen.

Appendix C: Statistical Derivation of Diffusion Length

The diffusion length, σ , which is the standard deviation of the Gaussian convolved with the original signal, can also be derived in a statistical manner. Lasaga starts by introducing the probability $W(Z, \tau)$ that an atom at position z at time t will be at position $z + Z$ at time $t + \tau$ through the effect of diffusion. If the concentration of profile at time t is known everywhere, then the concentration after diffusion acting over τ amount of time can be written as:

$$c(z, t + \tau) = \sum_{\text{all } Z} c(z - Z, t) W(Z, \tau) \quad (\text{C1})$$

This equation accounts for particles arriving at location z from all other locations $z - Z$.

Expanding the concentration terms as a Taylor series,

$$c(z, t + \tau) = c(z, t) + \tau \frac{\partial c}{\partial t} + \dots \quad (\text{C2})$$

and

$$c(z - Z, t) = c(z, t) - Z \frac{\partial c}{\partial z} + \frac{Z^2}{2} \frac{\partial^2 c}{\partial z^2} + \dots \quad (\text{C3})$$

Plugging these two concentration expressions into equation (68),

$$\begin{aligned} c(z, t) + \tau \frac{\partial c}{\partial t} + \dots = c(z, t) \sum_{all \ Z} W(Z, \tau) - \frac{\partial c}{\partial z} \sum_{all \ Z} ZW(Z, \tau) \\ + \frac{1}{2} \frac{\partial^2 c}{\partial z^2} \sum_{all \ Z} Z^2 W(Z, \tau) + \dots \end{aligned} \quad (\text{C4})$$

The particles must exist somewhere in space at time τ , so by the definition of probability:

$$\sum_{all \ Z} W(Z, \tau) = 1 \quad (\text{C5})$$

We can also use the definition of averages to write

$$\sum_{all \ Z} ZW(Z, \tau) = \langle Z \rangle \quad (\text{C6})$$

and

$$\sum_{all \ Z} Z^2 W(Z, \tau) = \langle Z^2 \rangle \quad (\text{C7})$$

Here $\langle \ \rangle$ represents a statistical average. We can now write the basic diffusion equation

(1) as:

$$\frac{\partial c}{\partial t} = \frac{\langle Z^2 \rangle}{2\tau} \frac{\partial^2 c}{\partial z^2} - \frac{\langle Z \rangle}{\tau} \frac{\partial c}{\partial z} \quad (\text{C8})$$

Comparing equations (1) and (73), we can write new statistical expressions for the diffusion coefficient and the velocity:

$$D = \frac{\langle Z^2 \rangle}{2\tau}$$

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(C9)

$$D = \frac{\langle Z^2 \rangle}{2\tau} \quad (\text{C10})$$

Equation (75) can be rearranged as

$$\begin{aligned} \langle Z^2 \rangle^{\frac{1}{2}} &= \sqrt{2D\tau} \\ &= \sigma \end{aligned} \quad (\text{C11})$$

This equation is equivalent to the diffusion length σ , which is used to describe the “extent” of diffusion. Equation (76) gives some insight into the statistical meaning of the diffusion length as the root mean square of the vertical position.

We can write equation (76) in its integrated form as:

$$\sigma^2 = \int_0^t 2D(\tau) d\tau \quad (\text{C12})$$

Taking the derivative $\frac{d}{dt}$ of both sides:

$$\frac{d\sigma^2}{dt} = 2D(t) \quad (\text{C13})$$

And we can now compare to equation (6) from Gkinis et al 2014, based directly off of Johnsen 1977:

$$\frac{d\sigma^2}{dt} - 2\dot{\epsilon}\sigma^2 = 2D(t) \quad (\text{C14})$$

Comparing equations (78) and (79), we have derived the same expression for sigma with the advection term neglected.

Appendix D: Derivation of Sigma Conversion Factor

This section derives the conversion factor between the standard deviation of the Gaussian in frequency space and the sigma diffusion length in the depth domain.

We start with the Gaussian (Green's Function) that describes diffusion:

$$G = \sqrt{\frac{1}{2\pi\sigma^2}} \exp\left(\frac{-x^2}{2\sigma^2}\right) \quad (D1)$$

This Gaussian is convolved with the original isotope signal in depth/time space to yield the diffused signal. Alternatively, this Gaussian can be Fourier transformed into frequency space to easily multiply the fft of the signal to, equivalently, yield the diffused signal.

The definition of transforming a Gaussian to frequency space is simple:

$$g(x) = \sqrt{\frac{\pi}{a}} \exp\left(\frac{-\pi^2 x^2}{a}\right) \Rightarrow \hat{g}(f) = \exp(-af^2) \quad (D2)$$

From Eq (1) above, we can solve for a by

$$\sqrt{\frac{1}{2\pi\sigma^2}} = \sqrt{\frac{\pi}{a}} \quad a = 2\pi^2\sigma^2 \quad (D3)$$

This expression for a yields the Gaussian we began with in (1). Now we can find the Fourier transform of (1) using the definition in (2).

$$\hat{G}(f) = \exp(-2\pi^2\sigma^2 f^2) \quad (D4)$$

We can use this to find, for some frequency, the amplitude A_σ of the diffused signal from the amplitude A_0 of the original signal as follows:

$$A_\sigma = A_0 \exp(-2\pi^2\sigma^2 f^2) \quad (D5)$$

We now have $\hat{G}(f)$ but we want $\hat{G}(k)$, a Gaussian in terms of k , where k is the wavenumber and $k = 2\pi f$, or $f = \frac{k}{2\pi}$. Plugging in this definition of k , we get:

$$\hat{G}(k) = \exp\left(-2\pi^2\sigma^2 \frac{k^2}{2^2\pi^2}\right) \quad \hat{G}(k) = \exp\left(\frac{-k^2\sigma^2}{2}\right) \quad (D6)$$

And again in terms of amplitude of a specific frequency,

$$A_\sigma = A_0 \exp\left(\frac{-k^2\sigma^2}{2}\right) \quad (D7)$$

This result agrees with Eqs (14) and (15) in Gkinis et al 2014.

As I've shown, the Gaussian filter yields the resulting amplitude of the diffused signal. In calculating diffusion length, we are fitting the power spectral density of the signal, and not the amplitude. Thus, we need to convert (6) and (8) to reflect the filter applied to the PSD rather than to the amplitude. Given that the power of a signal is defined as the square of the signal, we have

$$P_\sigma = P_0 \exp(-4\pi^2 \sigma^2 f^2) \quad (\text{D8})$$

and

$$P_\sigma = P_0 \exp(-k^2 \sigma^2) \quad (\text{D9})$$

where P_0 is the power spectral density of the original signal and P_σ is the power spectral density of the diffused signal.

Up to this point, σ has remained consistent and still refers to the standard deviation of the Gaussian in the depth before the Fourier transform, which is the diffusion length. When fitting a Gaussian to the isotope data, we do so in the frequency domain. Thus our fit gives us standard deviation of the Gaussian in frequency space, σ_{std} , which is not the same as σ diffusion length, the standard deviation of the Gaussian in the depth domain. We need to know how to convert from σ_{std} back to σ , which tells us the standard deviation of (1), and thus diffusion length.

To find the standard deviation in frequency space, we must convert Eq (11) to be in the standard Gaussian form $G(f) = \exp\left(\frac{-f^2}{2\sigma_{std}^2}\right)$.

$$\exp(-k^2 \sigma^2) = \exp\left(\frac{-f^2}{2\sigma_{std}^2}\right) \quad (\text{D10})$$

$$-k^2 \sigma^2 = \frac{-f^2}{2\sigma_{std}^2} \quad (\text{D11})$$

$$-2^2\pi^2 f^2 \sigma^2 = \frac{f^2}{2\sigma_{std}^2} \quad (D12)$$

$$\sigma^2 = \frac{1}{2^2\pi^2 2\sigma_{std}^2} \quad (D13)$$

$$\sigma = \frac{1}{2\pi\sqrt{2}\sigma_{std}} \quad (D14)$$

This is the conversion between the standard deviation σ_{std} of the Gaussian fit in the frequency domain, and the diffusion length σ , the standard deviation of the Gaussian in the depth domain.

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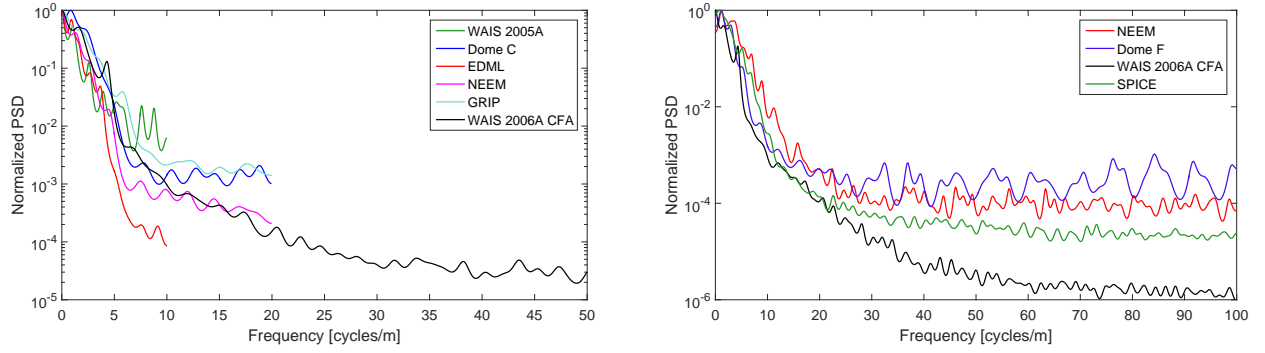


Figure 1: Left figure: The normalized PSD of five discretely measured $\delta^{18}\text{O}$ series plotted together with the PSD of a $\delta^{18}\text{O}$ WDC section. Right figure: The normalized PSD of four continuously measured δD series

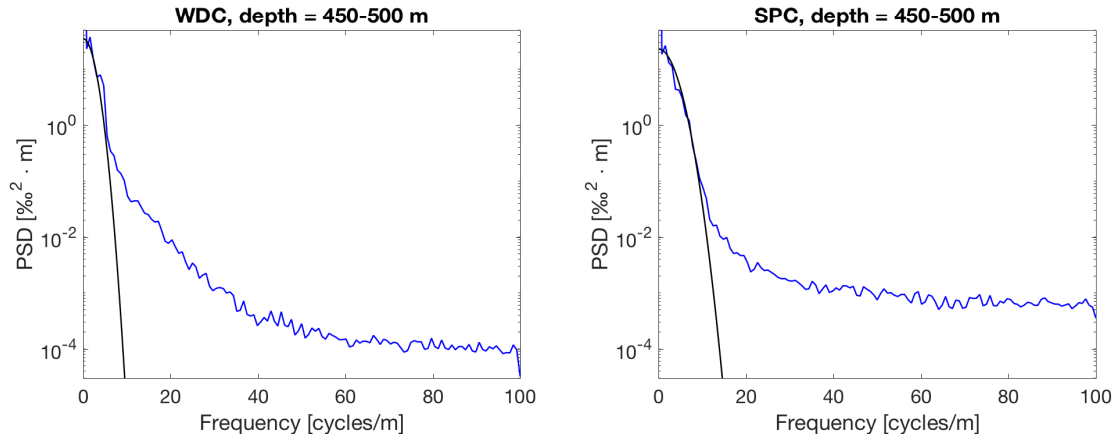


Figure 2: Cut-off technique on WDC and SPC at 450-500m depth

Table 1: Placeholder table

Treatments	Response 1	Response 2
Treatment 1	0.0003262	0.562
Treatment 2	0.0015681	0.910
Treatment 3	0.0009271	0.296

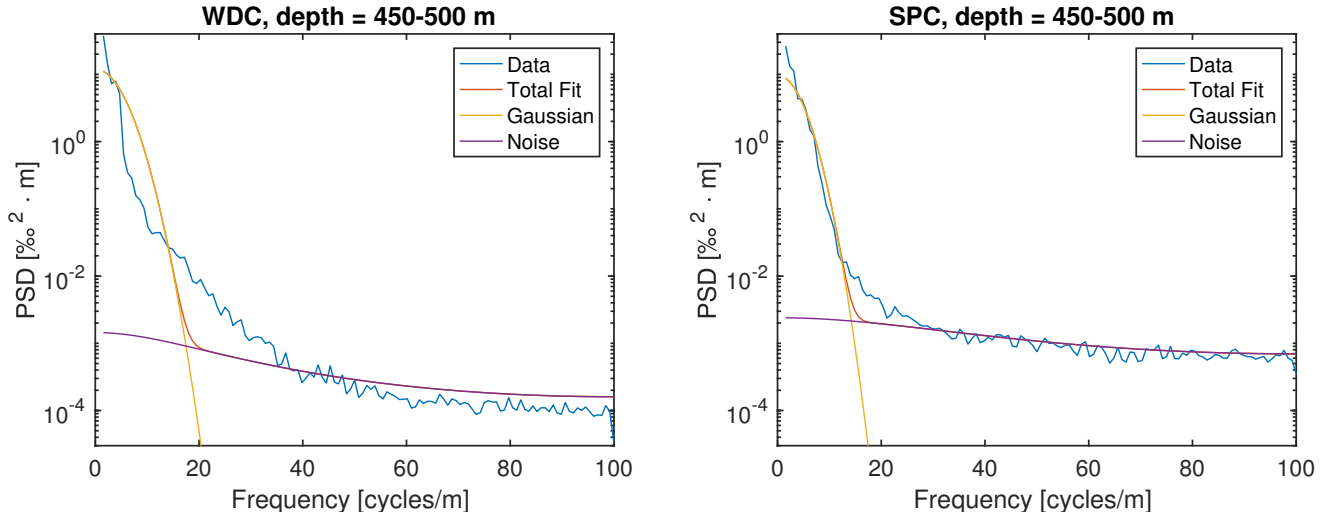


Figure 3: Single-Gaussian Multi-function fits for WDC and SPC at 450-500m depth

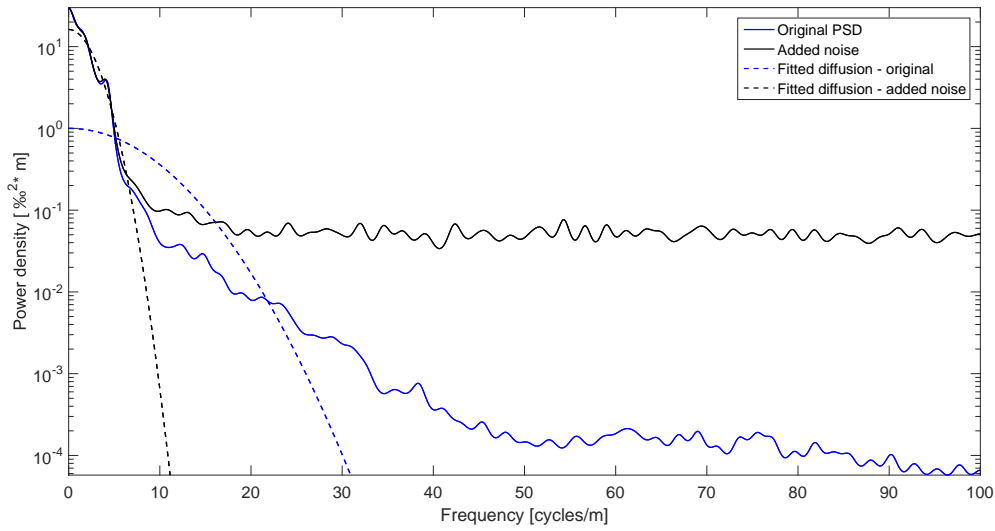


Figure 4: PSD of δD for a WDC section (solid curves). Blue curve represents the measured signal and the black curve represents the modified signal. The dashed lines represent the fitted Gaussian function which corresponds to the firm diffusion.

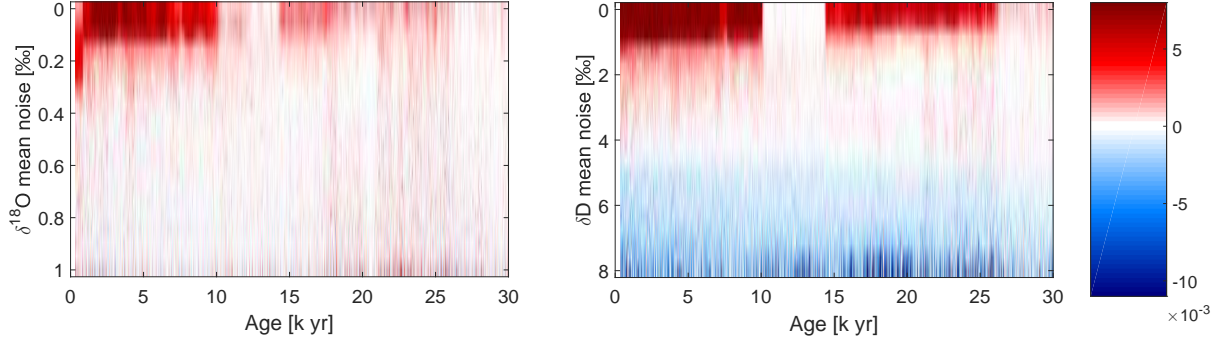


Figure 5: Gradient in estimated diffusion lengths with respect to noise level and age.

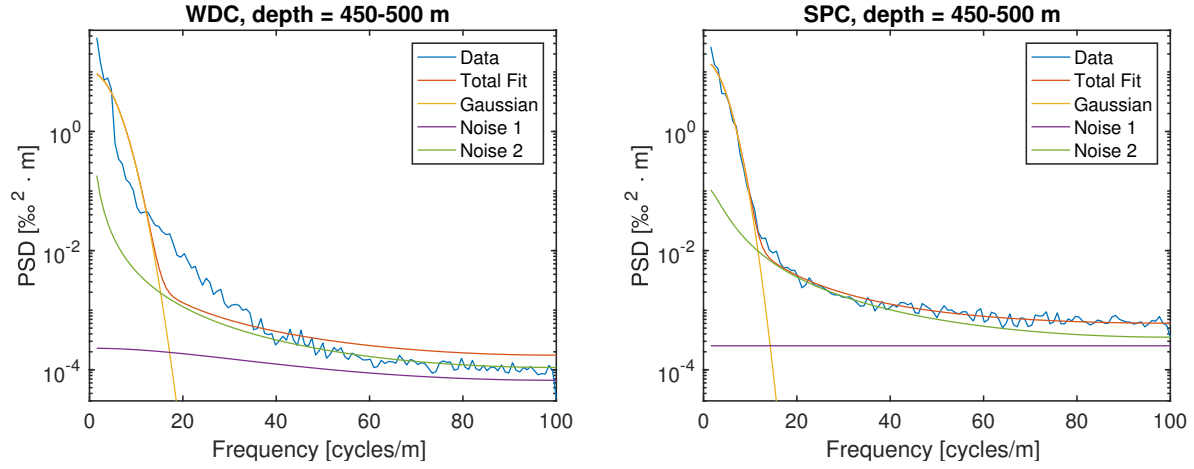


Figure 6: WDC and SPC spectra fits with a single Gaussian and two autoregressive noise functions

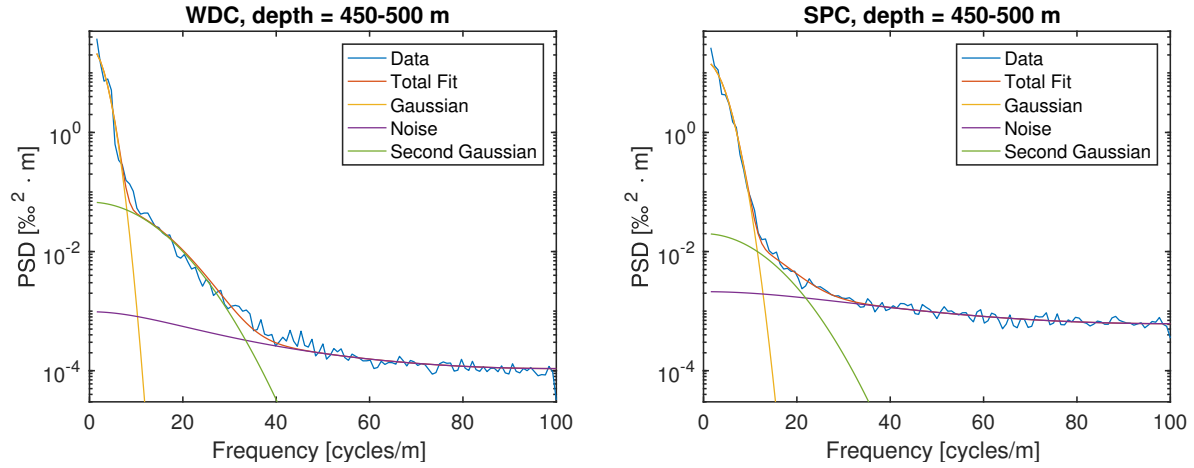


Figure 7: WDC and SPC spectra fits with double-Gaussian multi-function technique

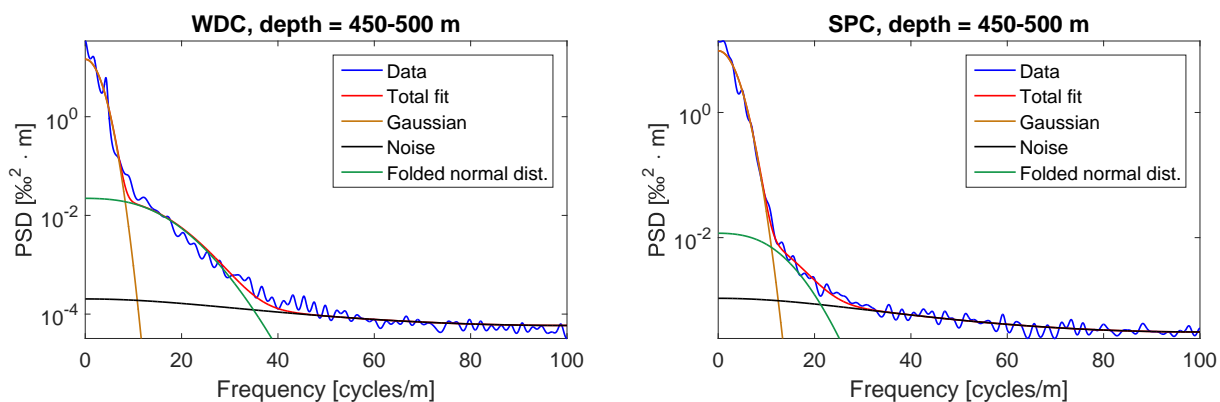


Figure 8: WDC and SPC spectra fits with double-Gaussian multi-function technique

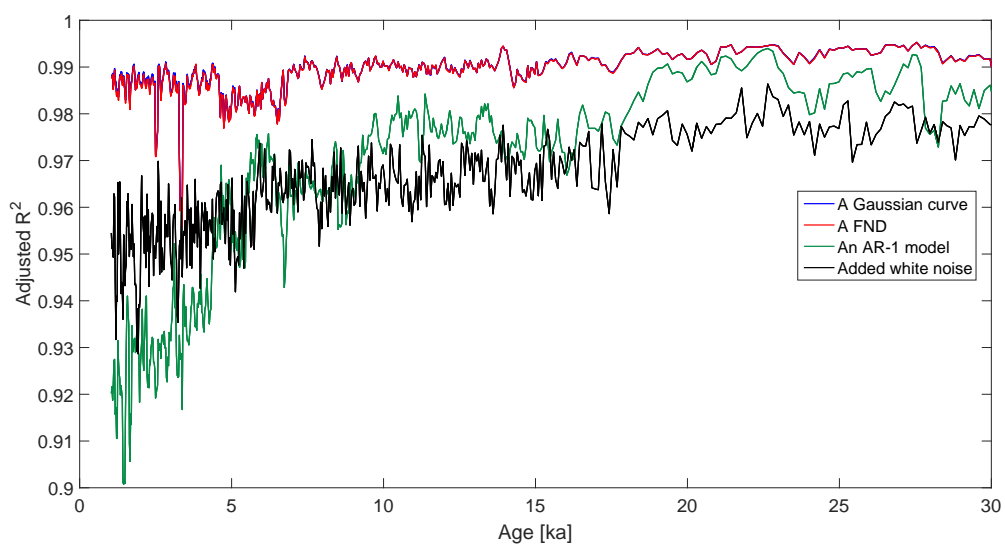


Figure 9: The adjusted goodness of fit calculations through age.

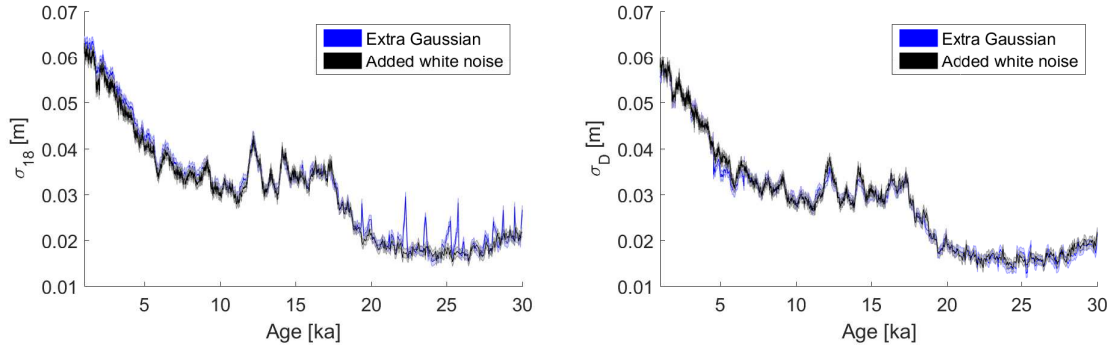


Figure 10: WDC diffusion lengths of $\delta^{18}\text{O}$ (left) and δD (right). The method that has an extra Gaussian is represented by the blue curve and the addition of white noise is represented by the black curve.

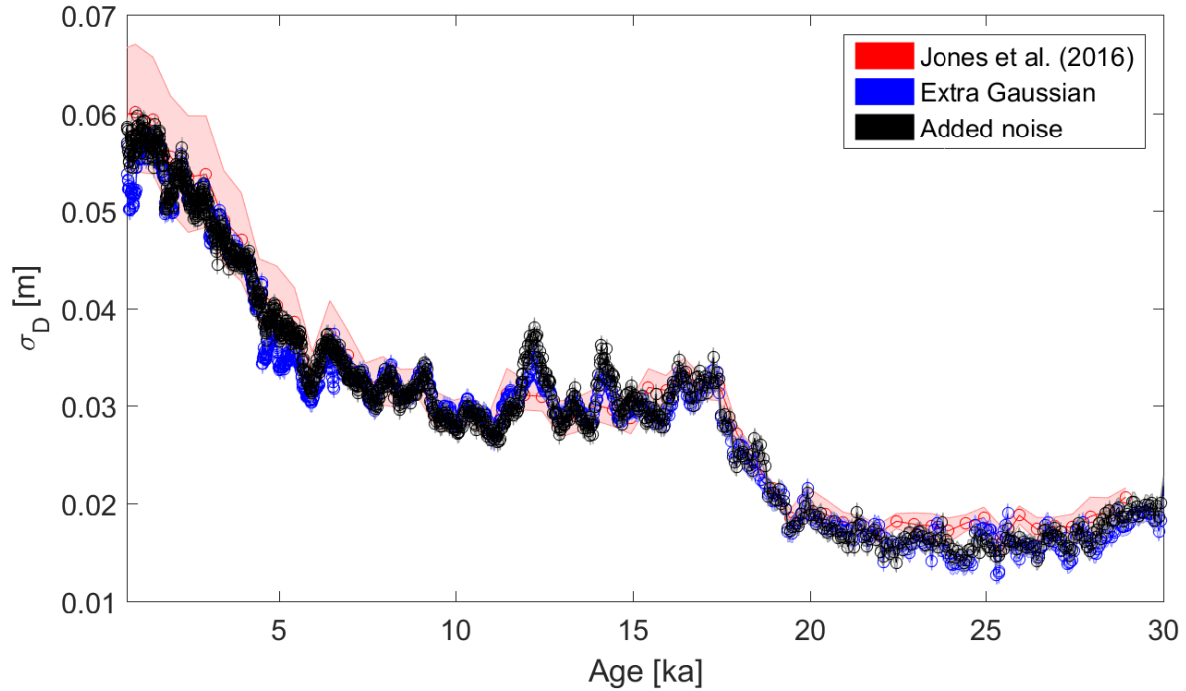


Figure 11: Estimated WDC diffusion lengths of δD compared with ?

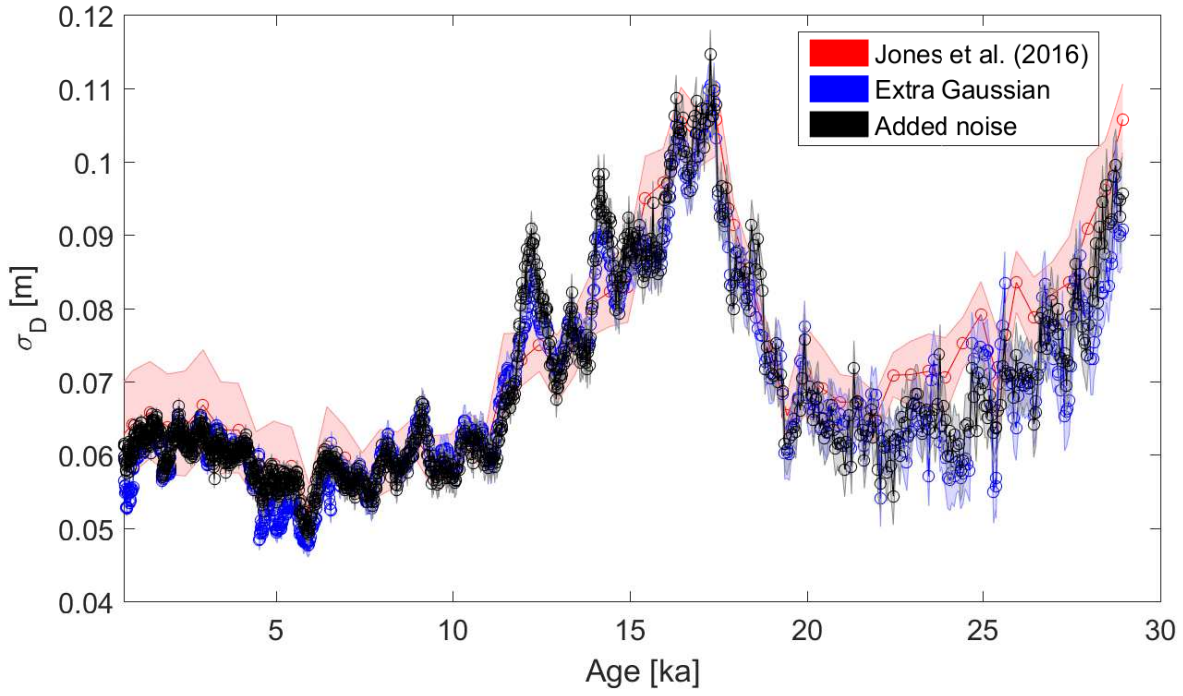


Figure 12: Thinning corrected WDC diffusion lengths compared of δD with ?

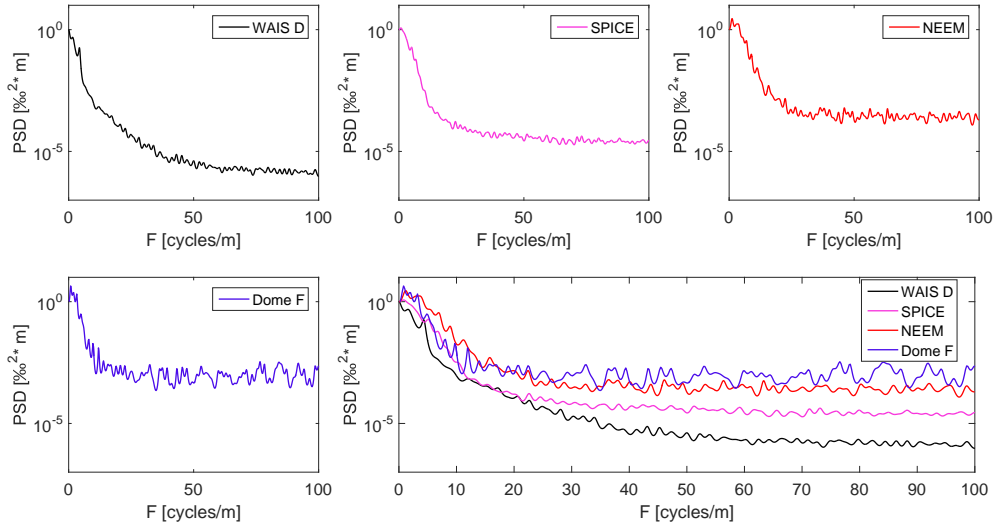


Figure 13: Normalized power spectral densities of δD records measured with different CFA systems.

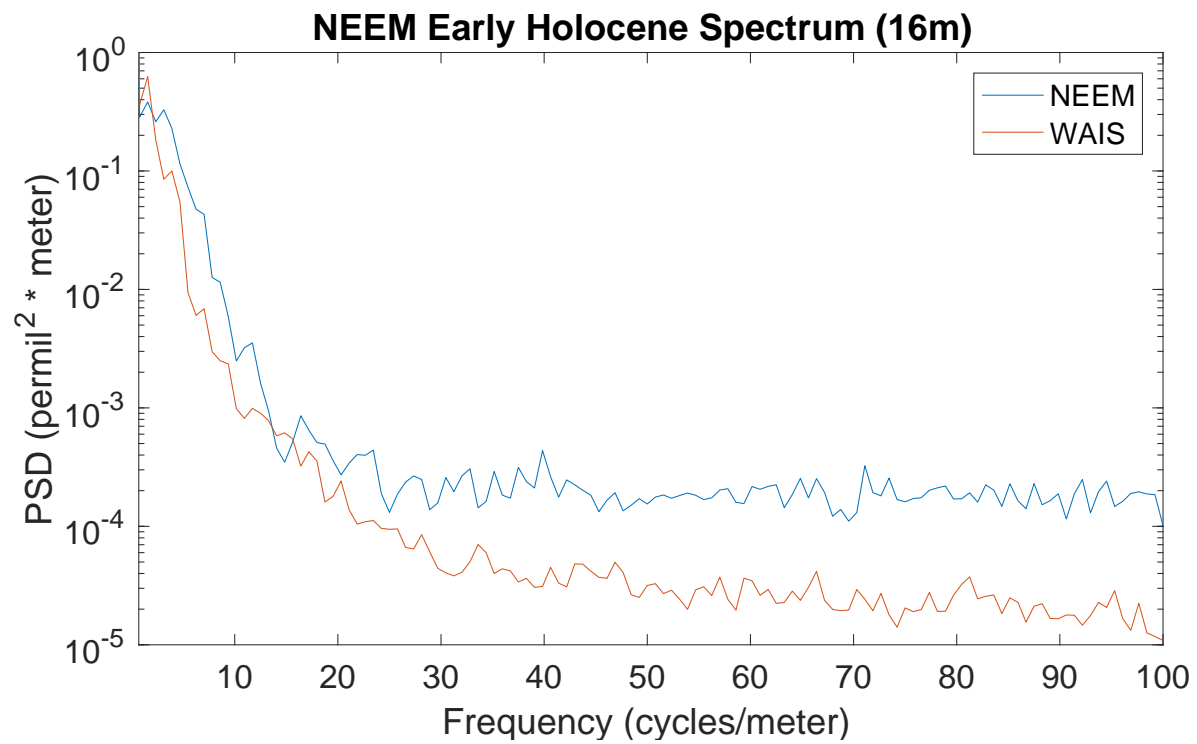


Figure 14: Spectra from NEEM 1382-1398 meters, CFA data measured on a different system