

# Controlled-source electromagnetic survey design for geothermal applications

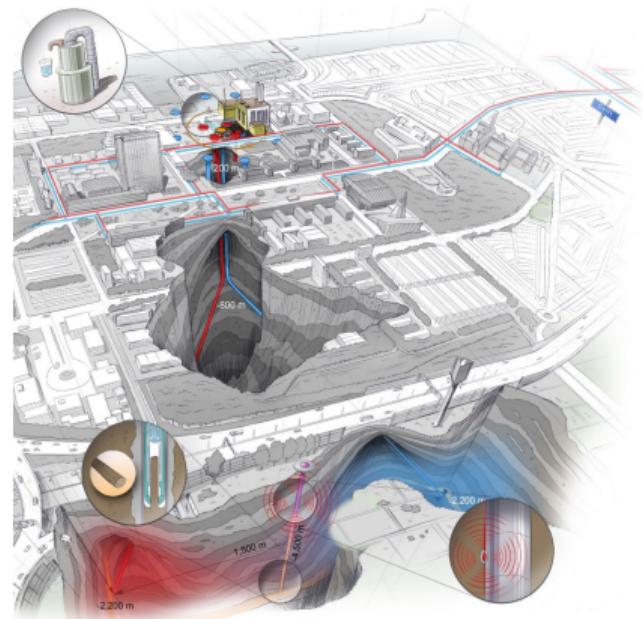
EasyGO Training: EM Modelling

Dr. Paula Rulff<sup>1</sup>

Dr. Dieter Werthmüller<sup>2</sup>

<sup>1</sup> TU Delft, NL

<sup>2</sup> ETH Zürich, CH



# Course info

<https://github.com/emsig/easygo-training-em>

**Dieter Werthmüller**

dieter.werthmuller@eaps.ethz.ch



**Paula Rulff**

p.rulff@tudelft.nl



**Suzanne van Noordt**

S.vanNoordt@tudelft.nl



# Learning Objectives

What would you like to learn?

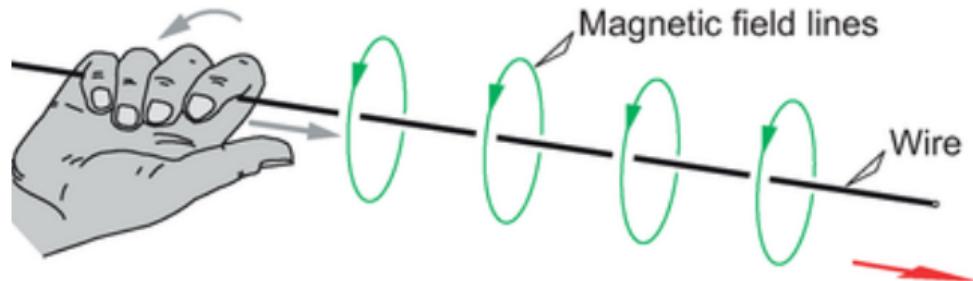
# Learning Objectives

What would you like to learn?

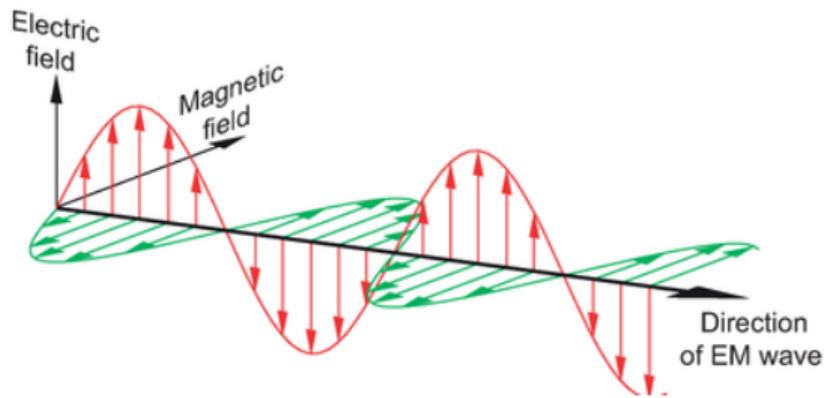
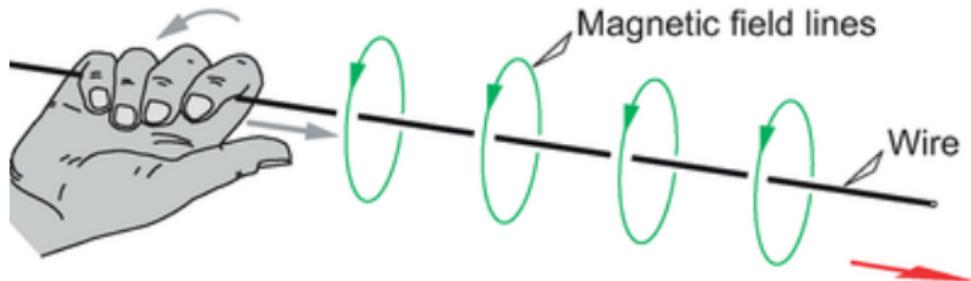
- **Relate** the electromagnetic geophysical method to the field setup and aim of investigation for geothermal monitoring
- **Use** *empymod* and *emg3d* to model electromagnetic data
- **Design** a field plan using controlled-source electromagnetics to monitor a near-surface geothermal project



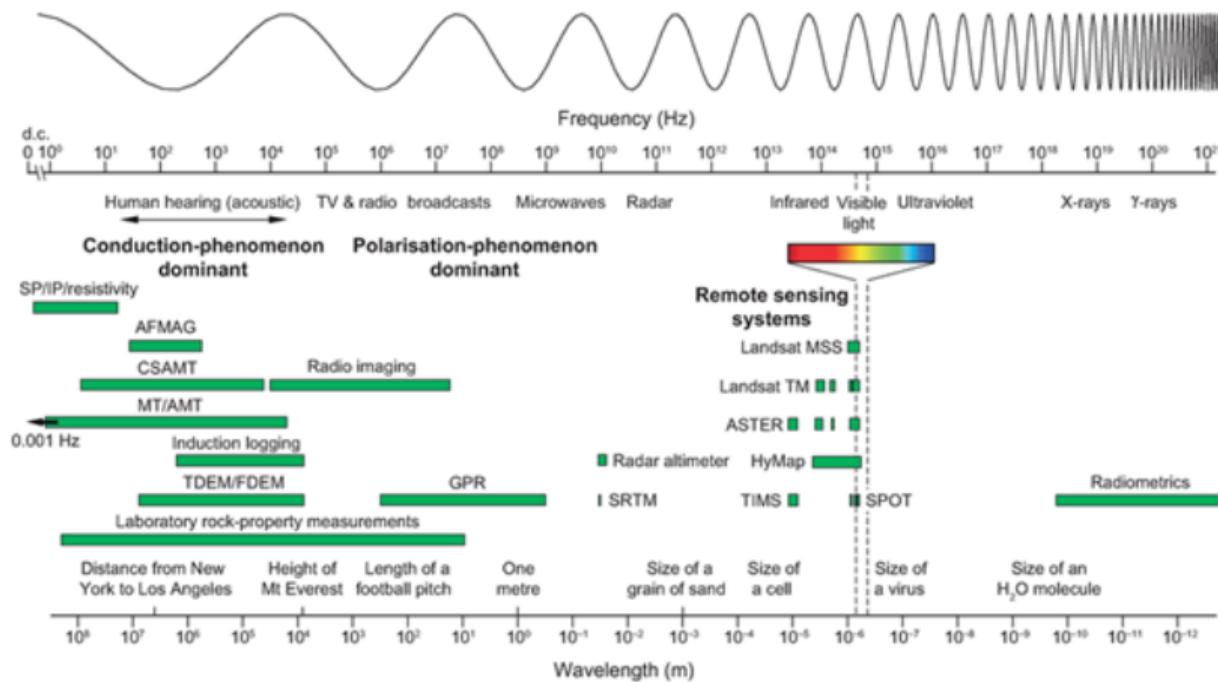
# Electromagnetism (EM)



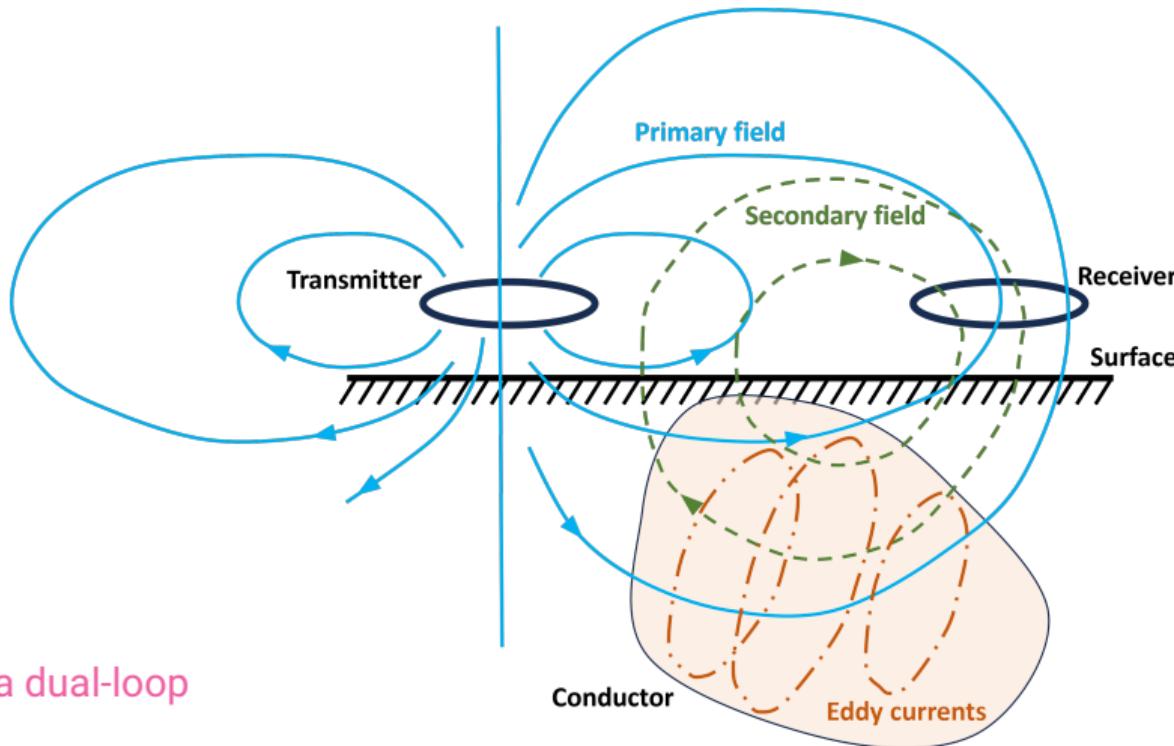
# Electromagnetism (EM)



# EM Geophysics

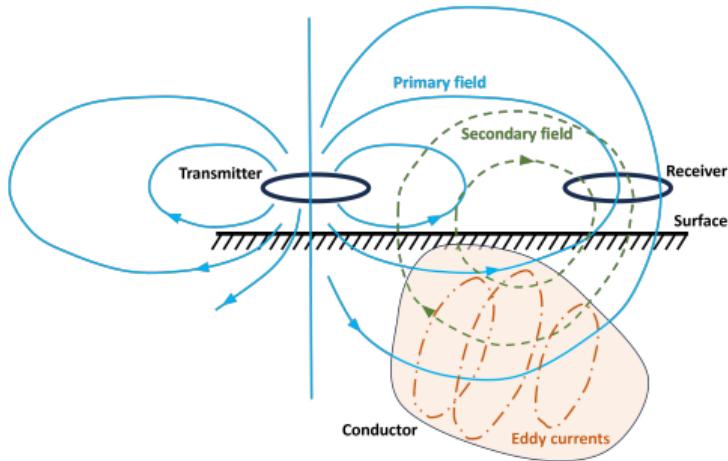


# EM Geophysics



Play with a dual-loop

# EM Geophysics



Play with a dual-loop

Maxwell's equations with a time dependency of  $e^{i\omega t}$  in frequency-domain:

$$\nabla \times E = -i\omega \mu H \quad (\text{Faraday's law})$$

$$\nabla \times H = \left( \frac{1}{\rho} + i\omega \epsilon \right) E + J_p \quad (\text{Ampère's law})$$

$$\nabla \cdot (\epsilon E) = q \quad (\text{Gauss' law})$$

$$\nabla \cdot \mu H = 0,$$

$\omega$ : angular frequency [rad/s]

$E$ : electric field [V/m]

$H$ : magnetic field [A/m]

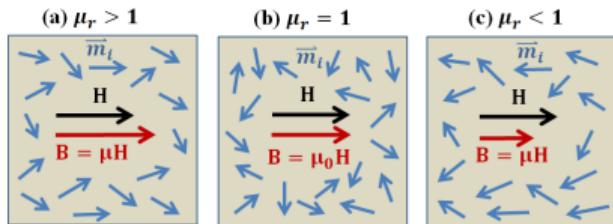
$J_p$ : impressed current density of the source

$q$ : volume electric charge density [ $C/m^3$ ]

$i$ : imaginary unit

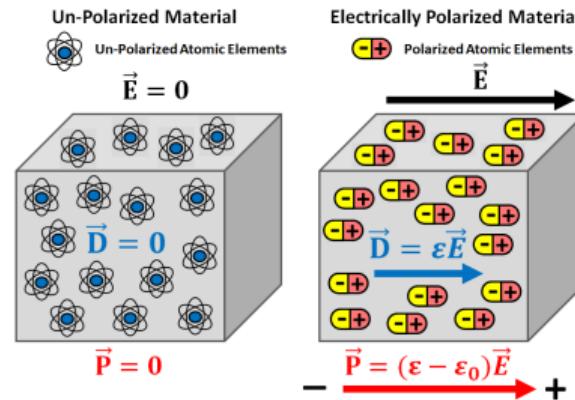
# Physical properties

## Magnetic permeability $\mu$



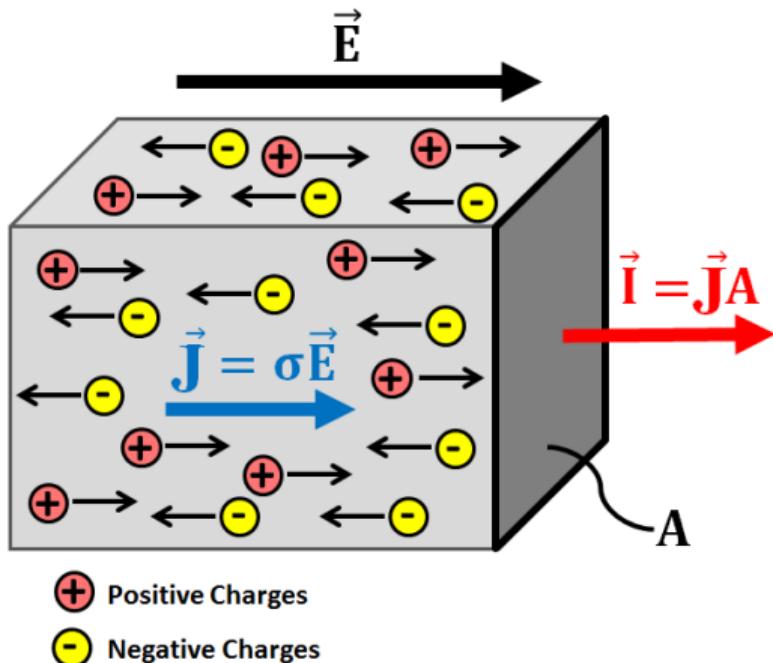
Characterises the degree of induced magnetism a material experiences under the influence of an external magnetic field.

## Dielectric permittivity $\epsilon$



Defines how strongly a material becomes electrically polarised under the influence of an electric field.

# Electrical resistivity $\rho$



$\rho [\Omega\text{m}]$

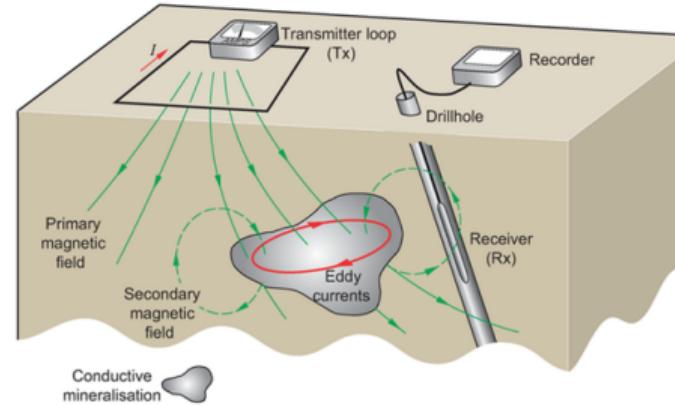
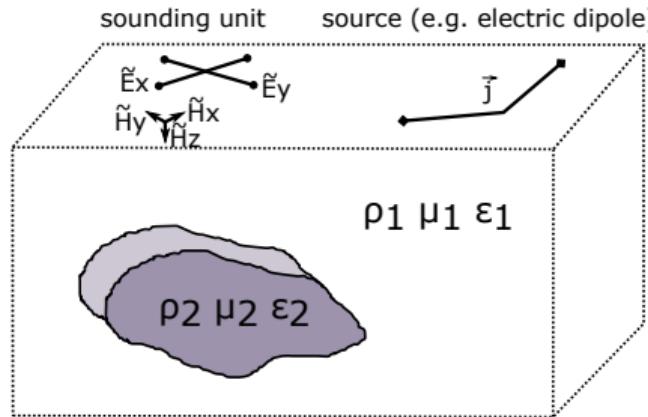
- Quantifies how easily electrical charges move through a given material when subjected to an applied electric field.
- Is the inverse of conductivity  $\sigma$ .
- Determines - in combination with the source current frequency and measurement geometry - the depth of investigation.
- $\rho$  of a subsurface material depends on its mineralogy and pore-water properties.

# Controlled-source Electromagnetics (CSEM)

Investigating the electric and magnetic properties of the Earth by measuring electric and magnetic fields generated by a transmitter and the conductive subsurface.

# Controlled-source Electromagnetics (CSEM)

Investigating the electric and magnetic properties of the Earth by measuring electric and magnetic fields generated by a transmitter and the conductive subsurface.



$\tilde{E}$ : electric field component  
 $\tilde{H}$ : magnetic field component

# Modelling

Total-field formulation for the electric field  $E$  in frequency domain with time dependence  $e^{i\omega t}$ :

$$\nabla \times \frac{1}{\mu} \nabla \times E + i\omega \frac{1}{\rho} E - \omega^2 \epsilon E = -i\omega J_p \quad \text{in } \Omega,$$
$$\hat{n} \times E = 0 \quad \text{on } \partial\Omega$$

$J_p$ : source ( $\approx 10^{-1} - 10^4$  Hz)

$\omega$ : angular frequency

$\rho$ : electrical resistivity

$\epsilon$ : dielectric permittivity

$\mu$ : magnetic permeability

# Modelling

Total-field formulation for the electric field  $E$  in frequency domain with time dependence  $e^{i\omega t}$ :

$$\nabla \times \frac{1}{\mu} \nabla \times E + i\omega \frac{1}{\rho} E - \omega^2 \epsilon E = -i\omega J_p \quad \text{in } \Omega,$$
$$\hat{n} \times E = 0 \quad \text{on } \partial\Omega$$

$J_p$ : source ( $\approx 10^{-1} - 10^4$  Hz)

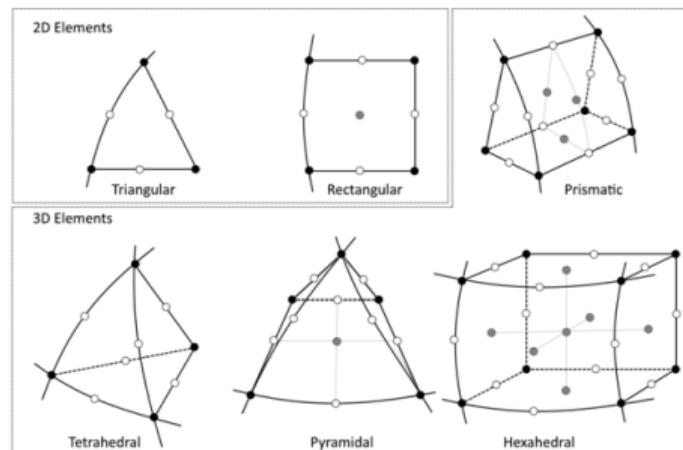
$\omega$ : angular frequency

$\rho$ : electrical resistivity

$\epsilon$ : dielectric permittivity

$\mu$ : magnetic permeability

## Discretisation



<https://www.comsol.com>

# Modelling

Layered-Earth modelling with empymod using `empymod.ipynb`



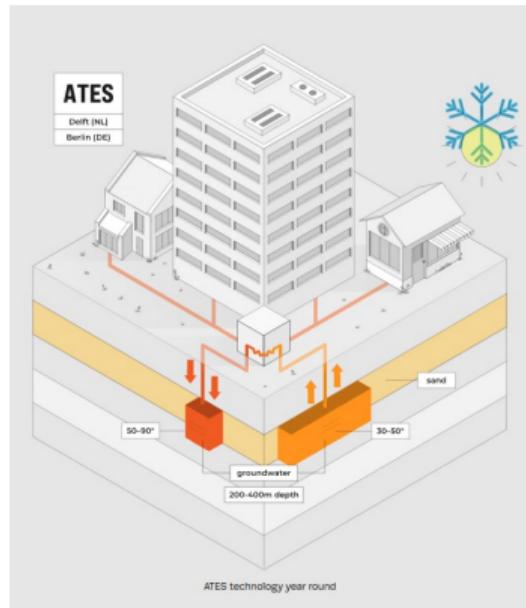
# Aquifer Thermal Energy Storage (ATES) in Delft

To account for seasonal supply and demand fluctuations, seasonal shallow heat storage will be connected to continuous deep geothermal exploration.

# Aquifer Thermal Energy Storage (ATES) in Delft

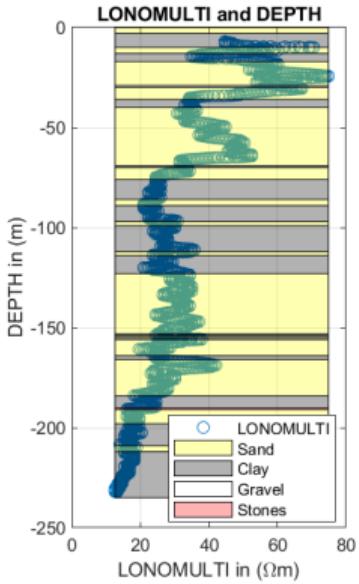
To account for seasonal supply and demand fluctuations, seasonal shallow heat storage will be connected to continuous deep geothermal exploration.

- Store heat surplus in summer and supply this heat in winter, thereby reducing the need for fossil fuel energy to meet peak heat demand
- Monitoring changes in subsurface temperatures due to heat injection and extraction operations

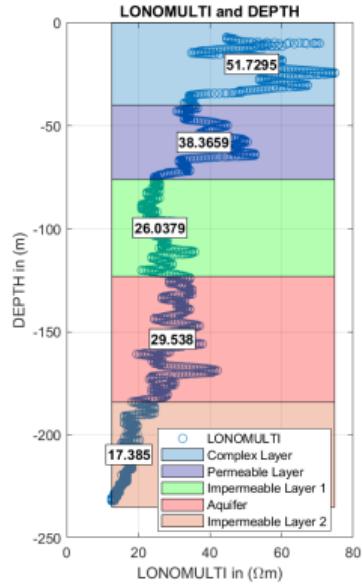


[https://www.push-it-thermalstorage.eu/  
technologies/#ates](https://www.push-it-thermalstorage.eu/technologies/#ates)

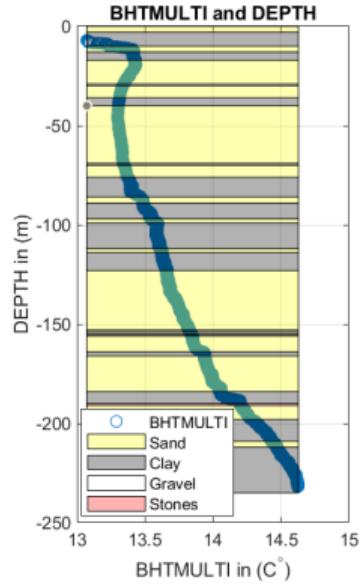
# ATES site: Subsurface Structure



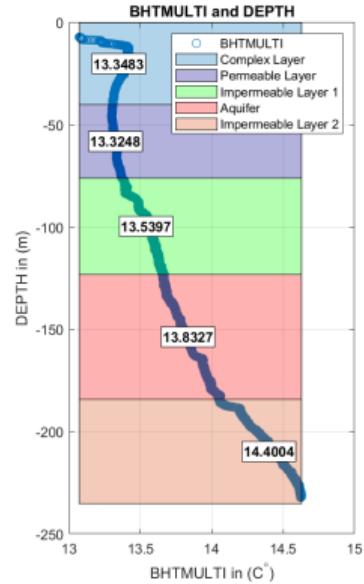
Stratigraphy with long-normal resistivity log [3]



Averaged resistivities [3]

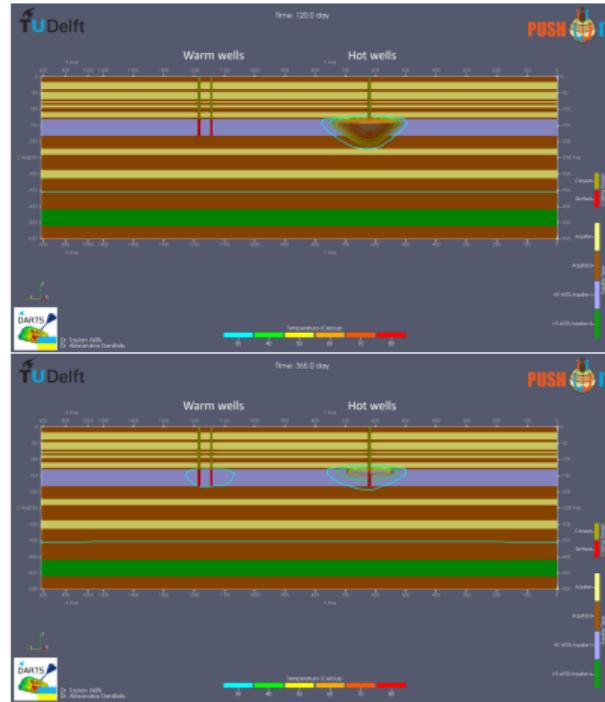
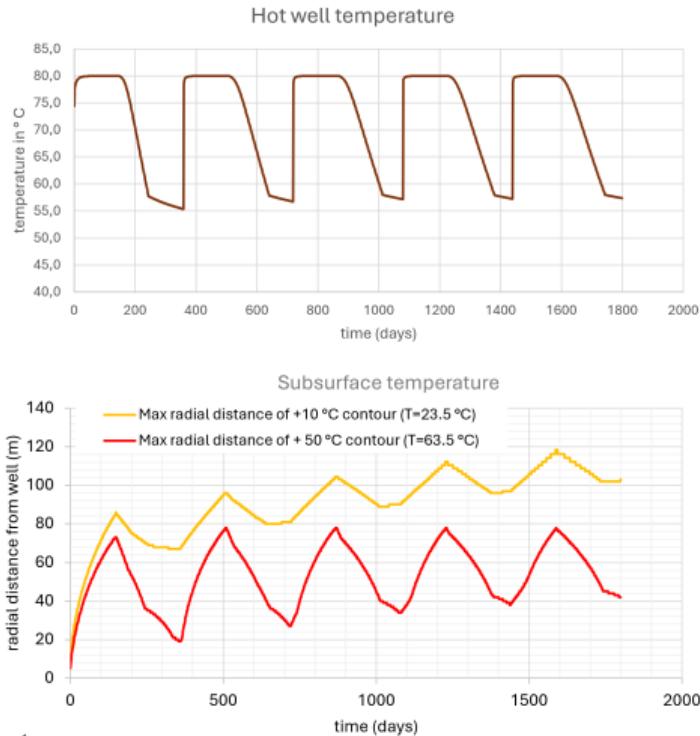


Stratigraphy with temperature log [3]



Averaged temperatures [3]

# ATES site: Heat Propagation Simulations



after 120 & 365 days, simulated with *openDARTS* [4]

# Geothermal Characterisation and Monitoring with CSEM

**Base line:** Characterising geoelectric structure before operations.

**Monitoring:** Aquifer temperature variations result in resistivity changes during operations.

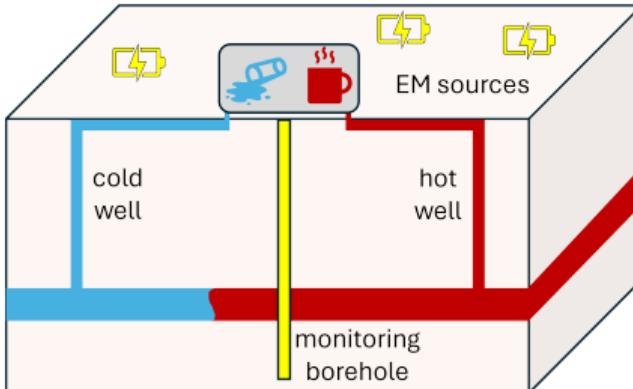
# Geothermal Characterisation and Monitoring with CSEM

**Base line:** Characterising geoelectric structure before operations.

**Monitoring:** Aquifer temperature variations result in resistivity changes during operations.

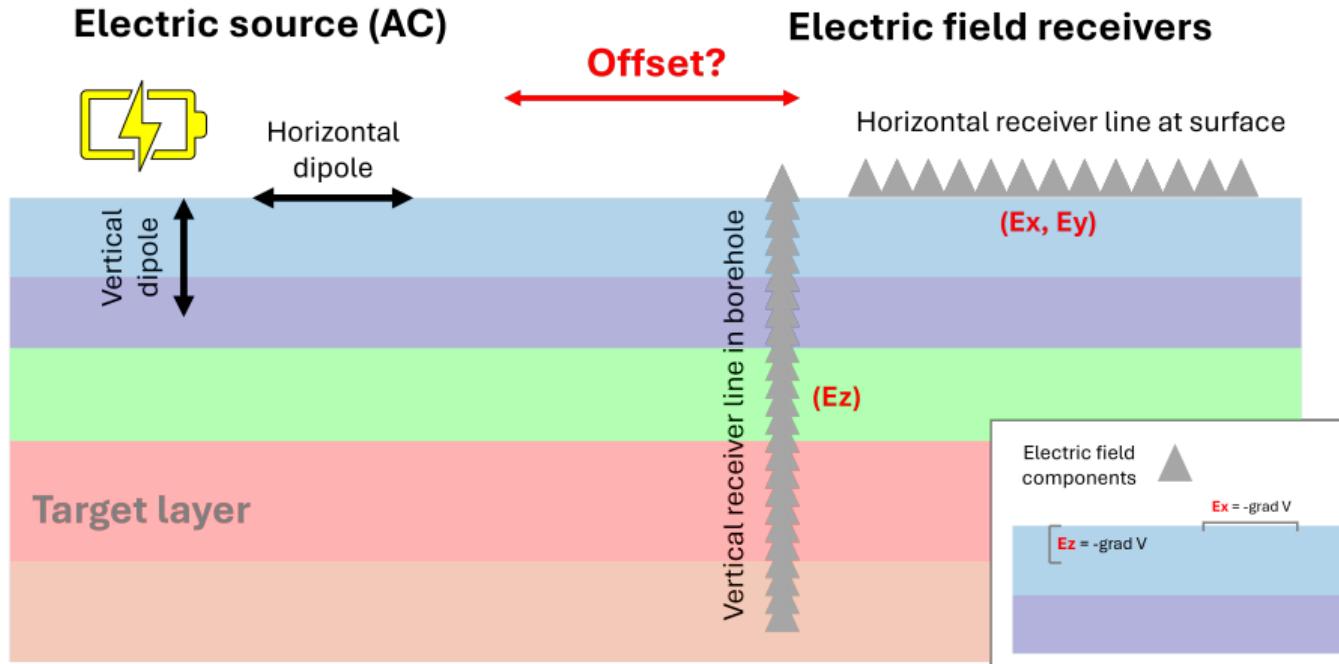
Monitoring aquifer temperature variations  $\Delta T$  through resistivity changes  $\Delta\rho$  [2] as

$$\rho(T_{hot}) = \rho(T_{cold}) [1 + 0.025(T_{hot} - T_{cold})]^{-1}.$$



For example,  $\Delta T$  from  $14^\circ\text{C}$  to  $75^\circ\text{C}$  results in  $\Delta\rho$  from  $29.5 \Omega\text{m}$  to  $11.7 \Omega\text{m}$ .

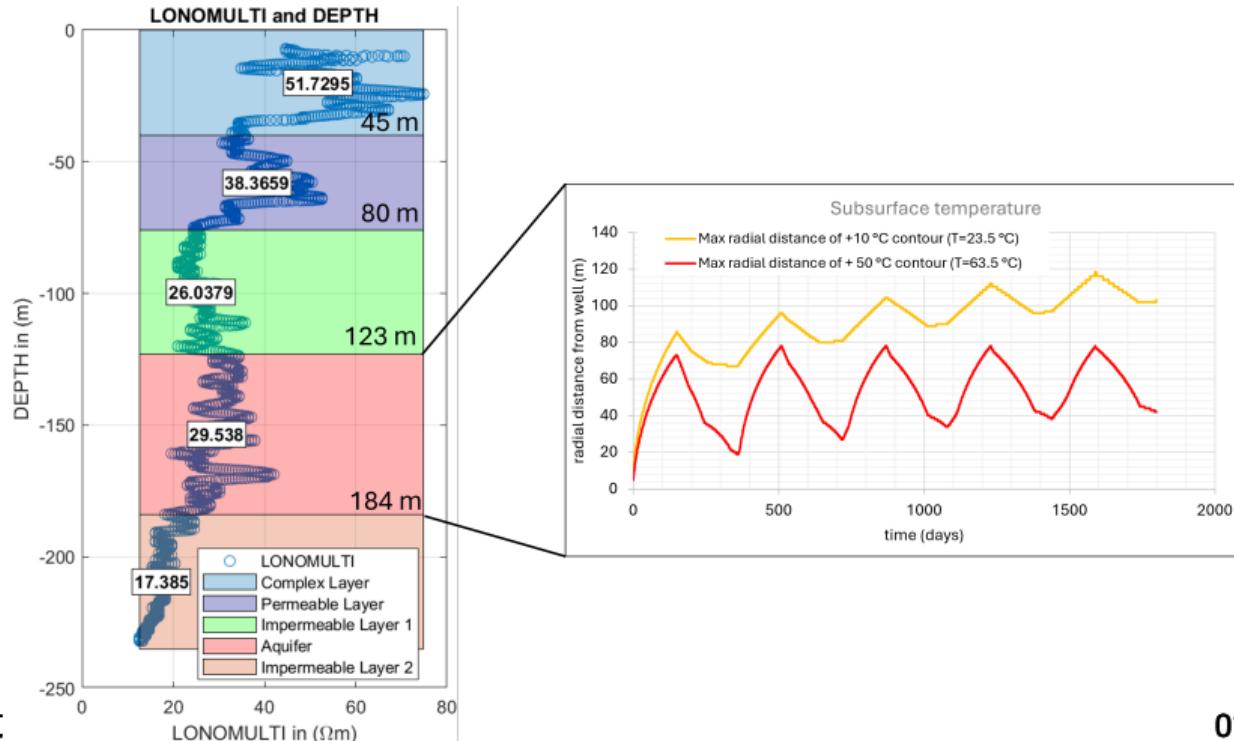
# Survey configurations and field components



Magnetic field components could be measured with receiver coils.

# Modelling

Design a CSEM survey for monitoring the ATES site using [empymod\\_ATES.ipynb](#)



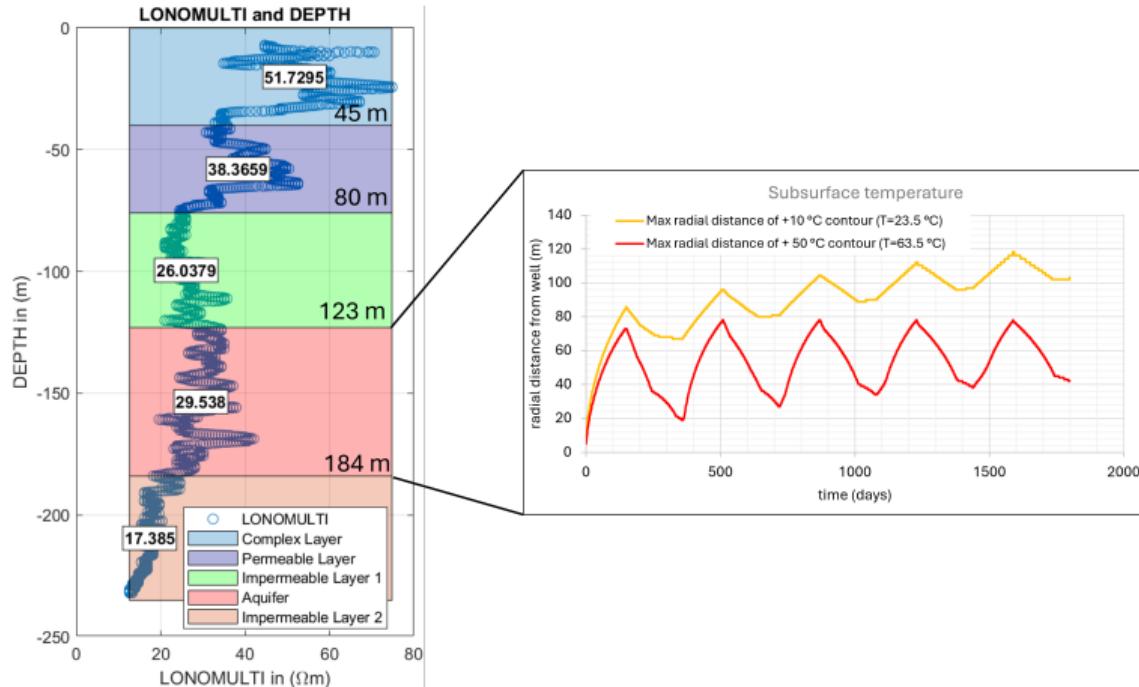
# 3D Modelling

3D modelling with `emg3d` using `emg3d.ipynb`.

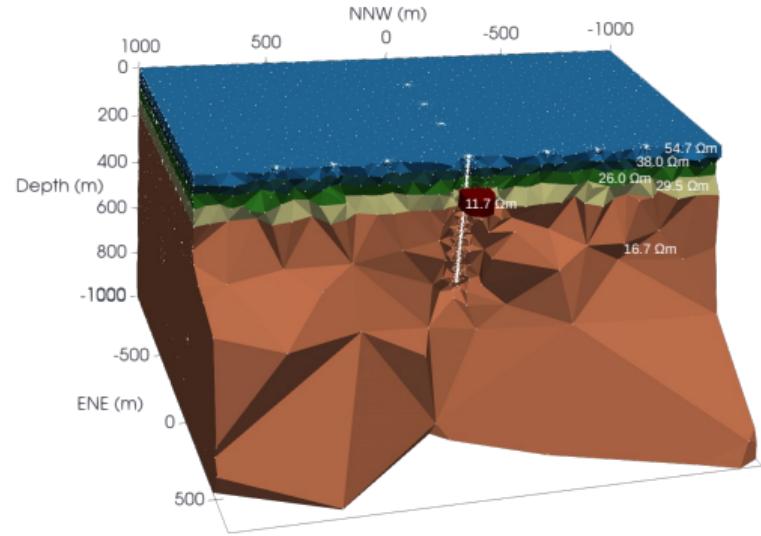
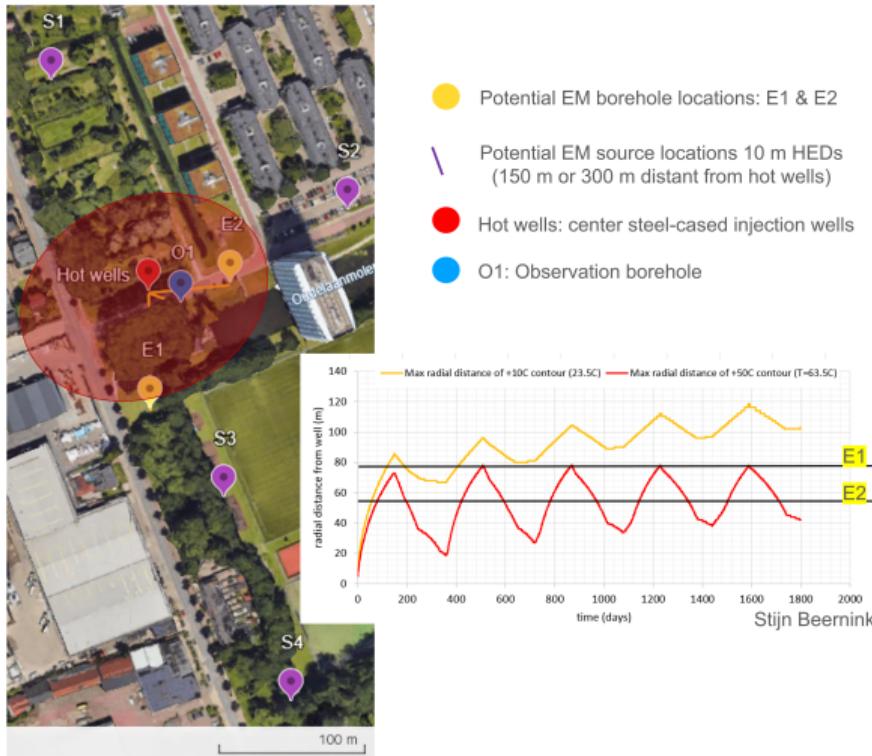


# 3D Modelling

Improve your CSEM survey design for monitoring the ATES site using [emg3d\\_ATES.ipynb](#) and including different hot plume sizes.



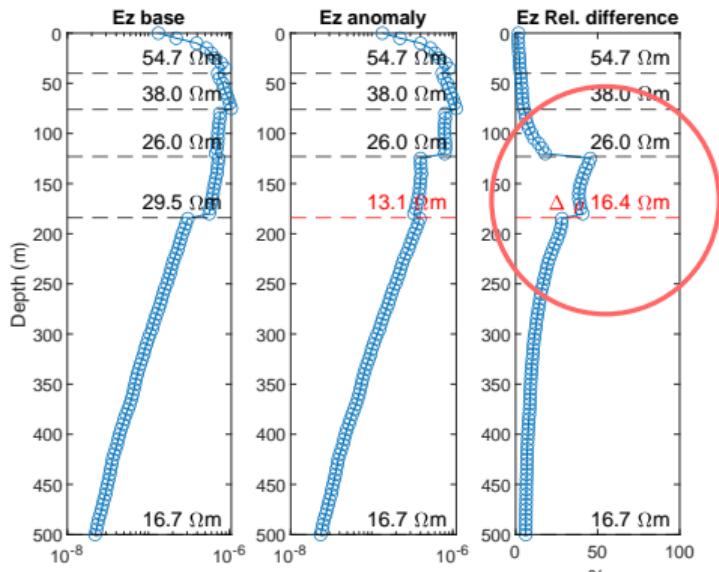
# ATES 3D EM Simulations



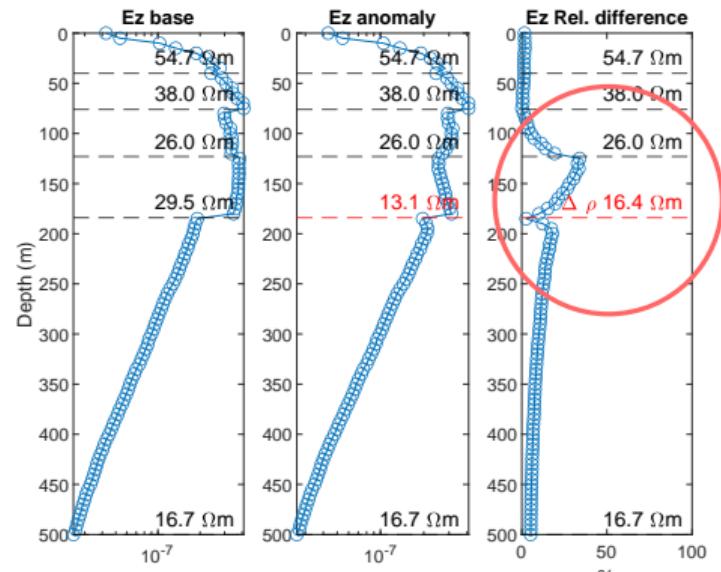
3D finite-element model: averaged layer resistivities, oval conductive heat-plume, sensor locations and radially distributed electric dipole sources

# ATES 3D EM Simulations

Densely (5 m) sampled  $E_z$  field components along borehole simulated with elfe3D [5].  
Electric dipole source at the surface. Aquifer at 123 - 184 m depth.

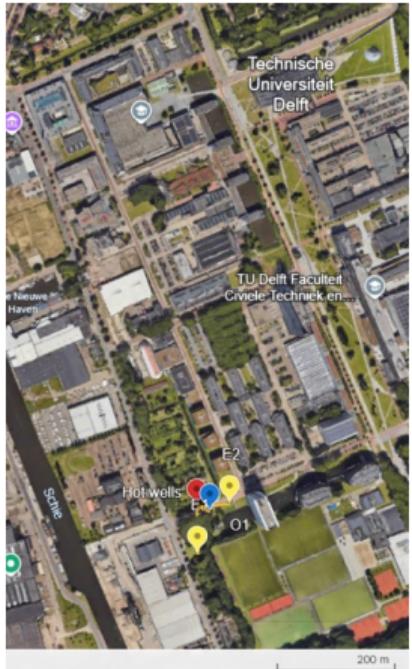


Source: 150 m North, borehole location E2



Source: 150 m North, borehole location E1

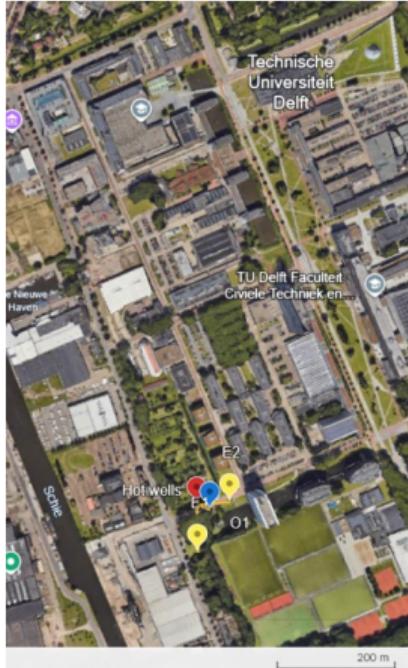
# ATES 3D EM Simulations: Infrastructure



ATES location on TU Delft campus

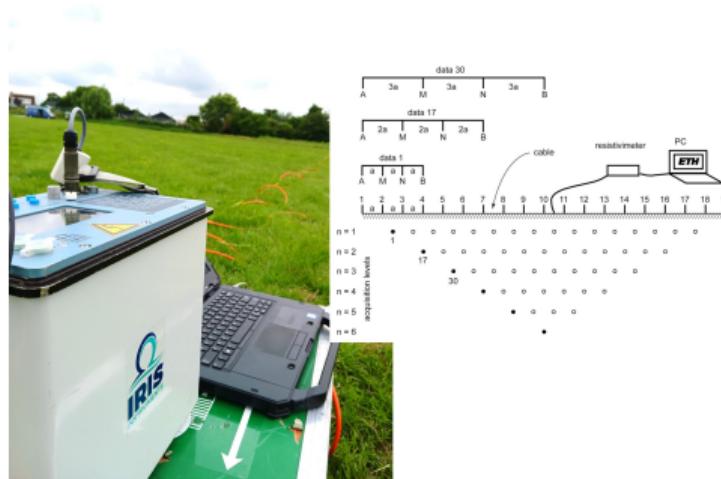
Infrastructure, particularly power lines and metallic objects, represents a significant source of "cultural noise" in electromagnetic geophysics, interfering with the measurement of subsurface electrical properties.

# ATES 3D EM Simulations: Infrastructure



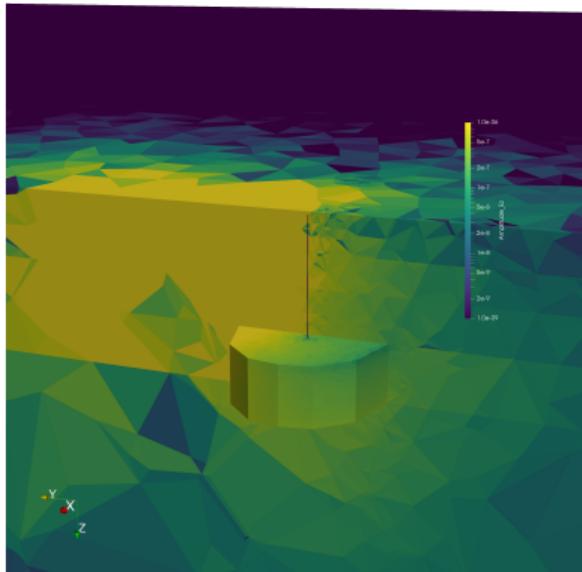
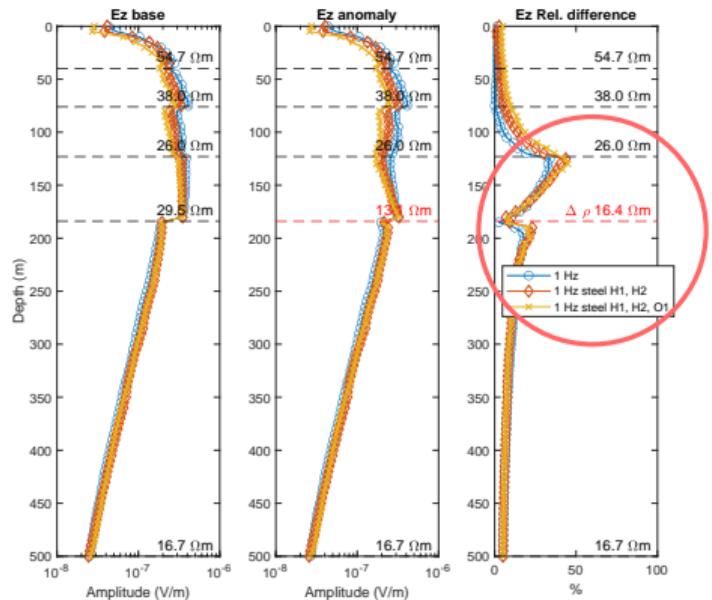
ATES location on TU Delft campus

Infrastructure, particularly power lines and metallic objects, represents a significant source of "cultural noise" in electromagnetic geophysics, interfering with the measurement of subsurface electrical properties.



# ATES 3D EM Simulations: Steel casings

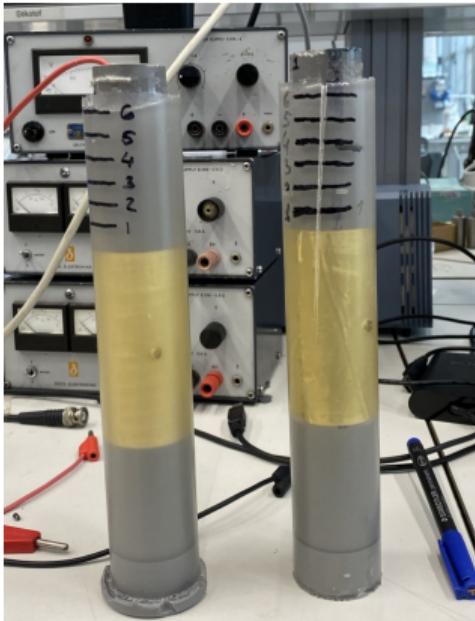
Densely (5 m) sampled  $E_z$  field components along borehole simulated with elfe3D [5].  
Electric dipole source at the surface. Aquifer at 123 - 184 m depth, 73 m thermal radius.



Ez field amplitudes in the domain

Source: 150 m North, borehole location E1

# Borehole Electrodes



Lab-test to develop zero-resistance ammeter with Integrator.



Casings with capacitive electrodes.

© Marat Ravilov

Capacitively coupled copper electrodes and non-corrosive casings [6]

- Combination makes long-term downhole monitoring possible.
- Crucial to minimise the capacitive effect of the cables to make SNR levels of galvanic and capacitive electrodes comparable.

# Q&A



# Questions & Feedback

p.rulff@tudelft.nl - dieter.werthmuller@eaps.ethz.ch

## References

- 1 Dentith, M., & Mudge, S. T. (2014). Geophysics for the Mineral Exploration Geoscientist. Cambridge: Cambridge University Press.
- 2 Keller, G. V., & Frischknecht, F. C. (1966). Electrical Methods in Geophysical Prospecting. Oxford: Pergamon Press Inc.
- 3 Bortolotti, M. (2024). Optimal Placement of Capacitive Electrodes in a Monitoring Borehole for the Campus Geothermal Project at TU Delft. BSc thesis. TU Delft.
- 4 Voskov et al. (2024). open Delft Advanced Research Terra Simulator (open-DARTS). Journal of Open Source Software, 9(99), 6737, <https://doi.org/10.21105/joss.06737>
- 5 Rulff, P. (2025). elfe3D v1.0.1: Modelling with the total electric field approach using finite elements in 3D. Journal of Open Source Software, 10(110), 7949.
- 6 Drijkoningen, G. et al. (2024). Capacitively coupled EM sensors integrated in non-corrosive casings for long-term CSEM monitoring. 85th EAGE Annual Conference & Exhibition.