

Response to reviewer

Reviewer #1 Evaluations

- Recommendation: Return to author for major revisions
- Significant: There are major errors or gaps in the paper but it could still become significant with major changes, revisions, and/or additional data.
- Supported: Mostly yes, but some further information and/or data are needed.
- Referencing: Mostly yes, but some additions are necessary.
- Quality: The organization of the manuscript and presentation of the data and results need some improvement.
- Data: Yes
- Accurate Key Points: Yes

Reviewer #1 (Formal Review for Author (shown to authors))

This paper uses simulations with the CESM2-WACCM model to investigate radiative forcing and surface temperature impacts of volcanic eruptions spanning magnitudes from Pinatubo-size to super eruptions. The paper highlights a number of non-linearities and time dependent changes in the response to different sized eruptions, and suggests that the primary difference between results from different models is tied to their chemical treatment of aerosol formation and growth.

The results are not very surprising, but there is significant value in the way the analysis is done, especially with regard to comparisons to other simulations. There are a few aspects I believe warrant further analysis before the paper is to be considered to publication.

General comments

The assumption of linearity between AOD and radiative forcing has been around for a long time, e.g. Hansen et al. (2005). What was not often stated is that this assumption is physically justified only for smaller eruptions, since by Beer's law, the reduction in radiation for a given AOD τ is $\exp(-\tau)$. For small τ , we can approximate $\exp(-\tau)$ as $1-\tau$, giving a linear relationship. If τ gets large, then clearly this linear assumption will break down. This likely is the main reason why the relationship between ERF and AOD is not the same for the different sized eruptions in this paper. I would strongly encourage the authors to revise their analysis to include consideration of the relationship between ERF and $\exp(-\tau)$. This was done e.g. by Marshall et al. (2020) which showed a more compact relationship between the two factors. I believe this will be much more important in this study since it explores an even larger range of AOD values. Note that I have not been very thorough in my critique of the results and discussion since I believe they will change a lot once this change in analysis is done.

i Reply

Letting $AOD \rightarrow 1 - \exp(-AOD)$ means that the most extreme AOD values will now be mapped to ~ 1 . Please find below the figures 1 to 4 with AOD adjusted. Since we are mostly interested in the more extreme AOD values, where $AOD > 1$, this region will be mapped to a very narrow range, and thus we found more use of the version using AOD directly. Fig3a show perhaps a more clear structure, but for consistency was kept as AOD as well. However, if you find they should be included we are open to including the adjusted figures as well.

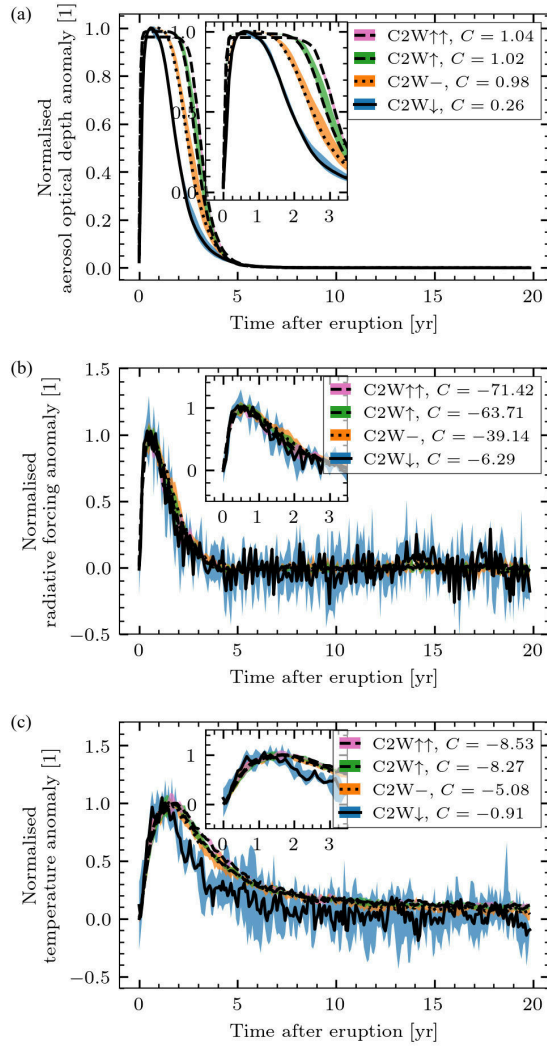


Figure 1: AOD represented as $1 - \exp(-\text{AOD})$

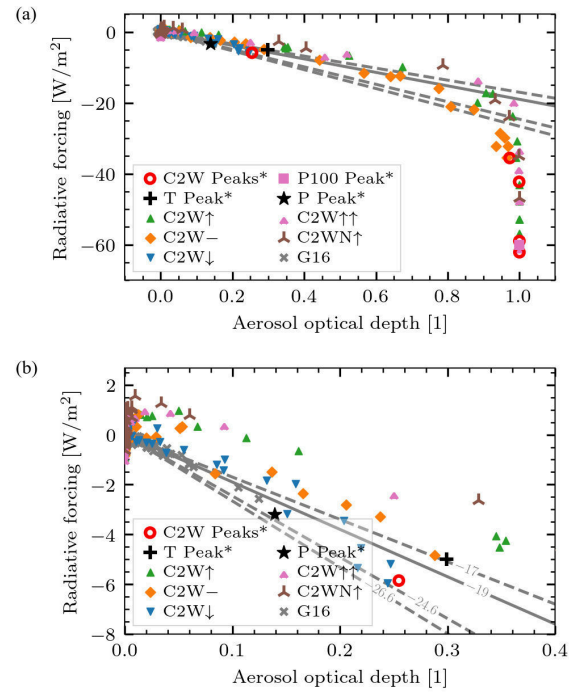


Figure 2: AOD represented as $1 - \exp(-\text{AOD})$

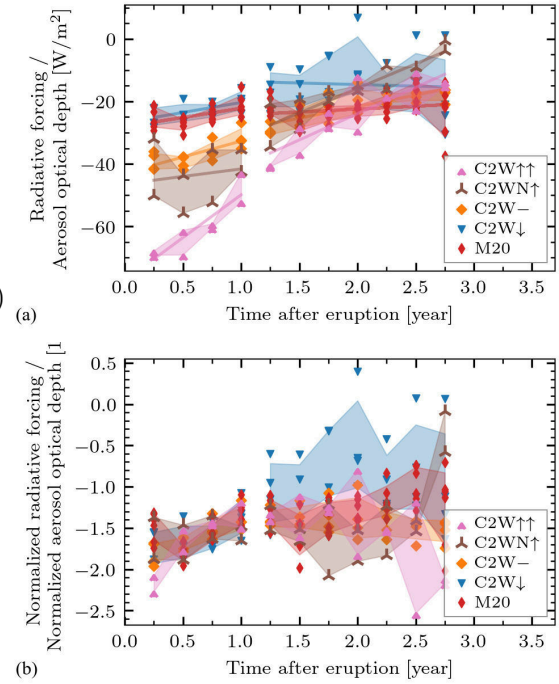


Figure 3: AOD represented as $1 - \exp(-\text{AOD})$

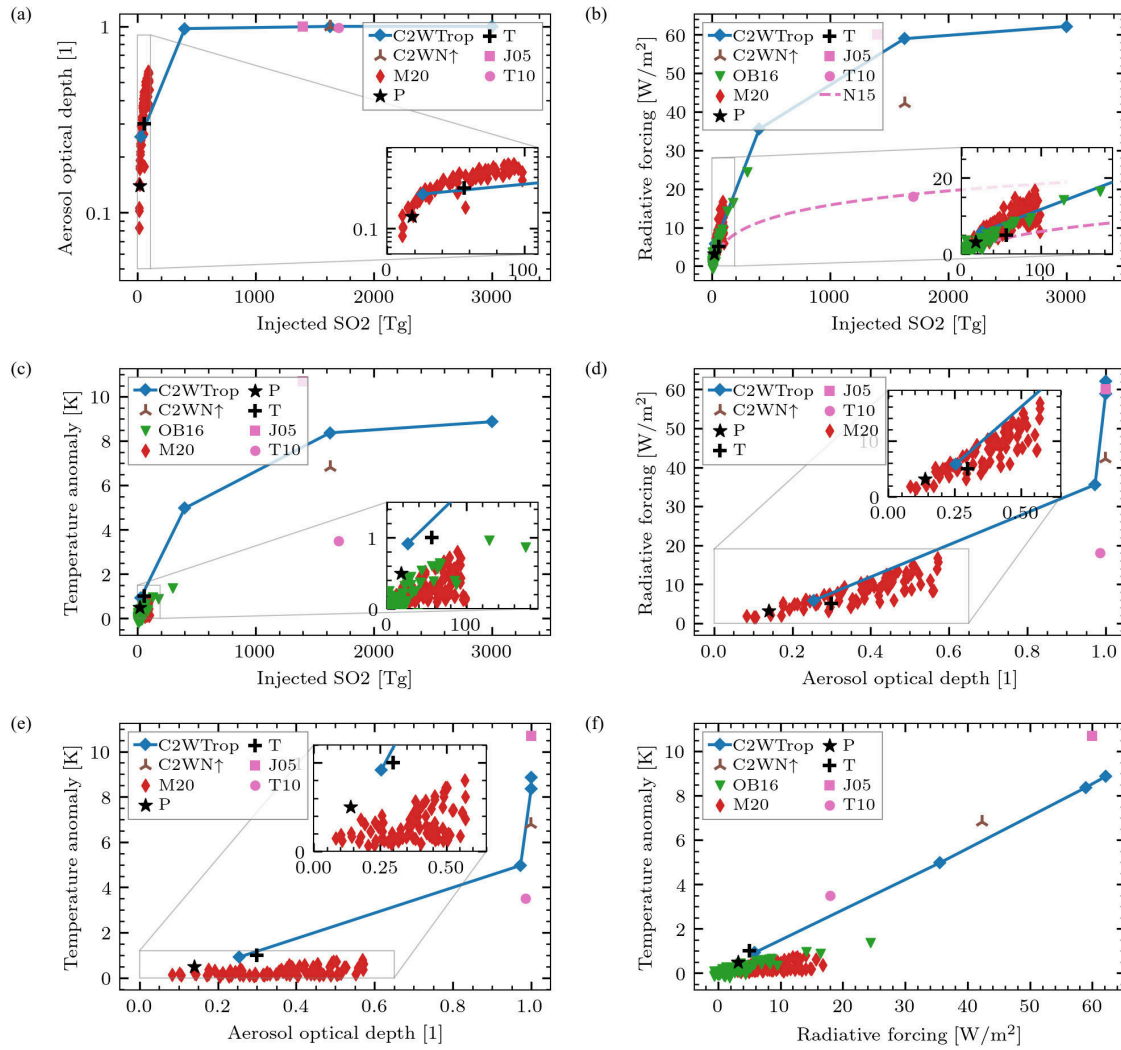


Figure 4: AOD represented as $1 - \exp(-\text{AOD})$

It is well known that as eruption magnitude increases, stratospheric aerosol size will increase, promoting aerosol loss through gravitational settling, and affecting the efficiency of radiative forcing (Pinto et al. 1989, Lacis et al., 1992, Timmreck et al., 2009, Timmreck et al., 2010). It is also clear that differences in aerosol size distribution is the primary cause of differences between models in terms of the magnitude and evolution of AOD and radiative forcing for volcanic eruptions (Clyne et al., 2020). I was surprised that these factors were not investigated in this study. It would be very valuable to include analysis of the simulated aerosol size in the present study. Failing this, at least the importance of aerosol size should be included in the introduction and discussion sections.

i Reply

This has been valuable input which now better explains our results and conclusions, specifically in relation to Fig. 4f.

As Brenna et al. (2020) run similar simulations and go deeper into effective radius, we do not provide a deeper analysis here other than to mention the importance of effective radius. We include results from Brenna et al. (2020) and McGraw et al. (2024) in Fig. 4 for comparison which aligns well with our simulations, and instead point to their studies for a deeper analysis as this is not

our main focus. Note also that only the effective radius of $0.6\ \mu\text{m}$ from McGraw et al. (2024) is included, as this is the only effective radius that was run by McGraw et al. (2024) across the full SO_2 range.

Below is also shown a plot of effective radius from four new simulations (original model output from the initial single-eruption simulations has unfortunately been deleted to limited data storage). These simulations show equatorial double-eruptions of 26&26 and 400&400 Tg SO_2 , with separation of two and four years.

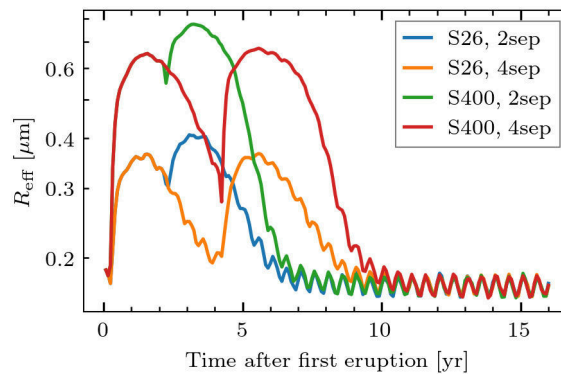


Figure 5: Effective radius of double eruption events from two identical eruptions of 26 and 400 Tg SO_2 .

Specific comments

- ✓ L11: Here and throughout, be careful whether to refer to “super-volcanoes” or “super-eruptions”. A supervolcano is just a volcano that has produced a super eruption at some point in its history. Volcanoes are just mountains most of the time. An eruption is what puts sulfur and other stuff into the atmosphere.

i Reply

Text has been adjusted to use “supereruption”.

- ✓ L15: Radiative forcing (RF) is a general term, effective radiative forcing (ERF) and instantaneous radiative forcing (IRF) are two types of RF. If you are specifically dealing with ERF, I would suggest using this more specific name and acronym throughout.
- ✓ L25: The evolution of SO_2 is rather simple to describe, it is chemically converted to H_2SO_4 . The evolution of sulfate aerosol is more complicated as it involves the formation and growth of the aerosol size distribution. I assume it’s the evolution of aerosol that is most important to the radiative forcing and climate impacts in the simulations?
- ✓ L31: Unfortunately for our intro sentences, the eruptions don’t directly emit sulfate aerosols (at least not very much), they emit gases which produce aerosols. Best to be accurate even if it costs a few more words.
- ✓ L33: the ERF is not quite the change in energy balance at TOA - it is that when surface temperature is held fixed.
- ✓ L55: the origin of $\text{RF} = -25 \times \text{AOD}$ is Hansen, Sato, Ruedy, Nazarenko, Lacis, Schmidt, Russell, Aleinov, Bauer, Bauer, Bell, Cairns, Canuto, Chandler, Cheng, Genio, Faluvegi, Fleming, Friend, Hall, Jack-

man, Kelley, Kiang, Koch, Lean, Lerner, Lo, Menon, Miller, Minnis, Novakov, Oinas, Perlwitz, Perlwitz, Rind, Romanou, Shindell, Stone, Sun, Tausnev, Thresher, Wielicki, Wong, Yao, & Zhang (2005), best to cite that study here. It is also relevant that this scaling has been used in subsequent IPCC reports, but that is worth its own statement.

- ✓ L63: Robock (2000) is a review paper, better to cite direct observational studies to support this estimate of the SO₂ lifetime.
- ✓ L63: the peak in AOD (at 500 nm) and thus the RF depends also on the aerosol size distribution, which evolves on a longer time scale than the conversion of SO₂ to aerosol.
- ✓ L75: I don't remember Sigl et al. (2022) quantifying the impact of eruptions vs other sources of variability. Schurer et al. (2013) might be a more appropriate citation here.
- ✓ L94: I believe the standard format would be to use the plus/minus symbol rather than brackets on these numbers.
- ☐ L95: These are stratospheric aerosol simulations, not volcano simulations.
- ✓ L104: SO₂ injection rate would be mass per unit time, but I think you mean SO₂ amount here.
- ✓ L107: this doesn't seem right, as -18 W m^{-2} seems quite small for a 1700 TgSO₂ eruption if the scaling factor is 65 W m^{-2} per TgSO₂?

i Reply

From Niemeier and Timmreck (2015):

$$\Delta R_{\text{TOA}} = -65 \text{ W m}^{-2} \exp\left(-\left(\frac{2246 \text{ Tg(S) yr}^{-1}}{x}\right)^{0.23}\right)$$

For an annual SO₂ injection of 1700 Tg(SO₂) = 850 Tg(S) the above yields

$$\Delta R_{\text{TOA}} = -65 \text{ W m}^{-2} \exp\left(-\left(\frac{2246}{850}\right)^{0.23}\right) \approx -18.6 \text{ W m}^{-2}$$

So, from assuming an annual total injection of 850TgS, the equation is consistent with the model output from Timmreck et al. (2010) of -18 W m^{-2} .

There is of course a significant difference between continuous injection (as used in Niemeier and Timmreck (2015) to fit to the equation) and an eruption simulation as in Timmreck et al. (2010). This has hopefully been made more clear in the text now.

- ✓ L135: I don't agree with this statement-a super eruption is going to produce much different aerosol characteristics than a cluster of smaller eruptions.

i Reply

This was indeed the intended meaning, where the reference to the 100xPinatubo simulation by Jones et al. (2005) was the example used to illustrate some potential errors or biases that can come up. It is thus used as motivation for why the simulations using SO₂ input are an interesting and more physically complete comparison, but more importantly to justify our upper threshold of temperature anomaly. The section has been re-phrased to better highlight this.

- ✓ Table 1: I find these ensemble labels very non-intuitive. Since all experiments use CESM2(WACCM) the C2W can be left out of the labels. Arrows are typically used to indicate directions, not relative

magnitudes. I strongly encourage the authors to use labels which more intuitively communicate the relative magnitude of the eruption simulations and thus the differences between the ensembles.

i Reply

Ensembles are now specified as S<TgSO₂> and an extra N for the Northern latitude eruption ensemble.

- ✓ L157: The method section needs a description of how ERF was calculated.
- ✓ L178: Well, in order of magnitude sure, but there's a factor of about 2 difference, best to be clear about this.
- ✓ L182: Is the SO₂ injected uniformly between these altitudes?

i Reply

Adjusted to now specify how it is ramped up/down over three vertical levels.

- ✓ L197: Importantly it's the Normalized RF and temperature timeseries that are indistinguishable here.
- ✓ L204: again, the normalized AOD
- ✓ Figure 3: It looks like in this plot all ensemble members are plotted, as well as a standard deviation range of the ensemble. This seems redundant, either all the members or the mean and SD would seem to be sufficient and would improve the readability of the plot.
- ✓ L269: Here is one example where I expect this result may change when you consider not AOD but exp(-AOD).
- ✓ L338: I'd argue that the temperature anomalies from a simulation with fixed SSTs are close to meaningless in comparison with coupled ocean simulations.

i Reply

This was included to illustrate the difference, but can be removed if this makes the figure more clear.

- ✓ Fig. 4: everything that is a peak value should be referred to as such on figure axis labels in in caption.

i Reply

In this figure, everything plotted are peak values. This has now been explicitly stated in the figure caption.

- ☐ L364: why is this being assumed when it has been shown this isn't very valid?
- ☐ L365: It seems as if x is referring to different quantities in the different equations. Much more attention to detail needed if this mathematical framework is to be explained in sufficient detail.
- ✓ L378: Volcanic forcing definitely lasts for more than a year!

- ✓ L398: It's quite well known that simply scaling up the forcing for Pinatubo is not a good representation of forcing from super eruptions, see e.g. Timmreck et al. (2009) and Timmreck et al. (2010). That the AOD used in Jones et al. leads to a too strong RF is therefore not very surprising.

i Reply

The aim of this section was to put more confidence into the lower bound (upper in Fig4) of the temperature perturbation. This has been slightly re-worded and an extra citation to aerosol effective radius studies has been included to hopefully make this more precise.

References

- ☐ Timmreck, C., Lorenz, S. J., Crowley, T. J., Kinne, S., Raddatz, T. J., Thomas, M. A., and Jungclaus, J. H.: Limited temperature response to the very large AD 1258 volcanic eruption, *Geophys. Res. Lett.*, 36, <https://doi.org/10.1029/2009GL040083>, 2009.
- ✓ Timmreck, C., Graf, H.-F., Lorenz, S. J., Niemeier, U., Zanchettin, D., Matei, D., Jungclaus, J. H., and Crowley, T. J.: Aerosol size confines climate response to volcanic super-eruptions, *Geophys. Res. Lett.*, 37, L24705, <https://doi.org/10.1029/2010GL045464>, 2010. (Timmreck et al., 2010)
- ☐ Schurer, A. P., Hegerl, G. C., Mann, M. E., Tett, S. F. B., and Phipps, S. J.: Separating Forced from Chaotic Climate Variability over the Past Millennium, *J. Clim.*, 26, 6954-6973, <https://doi.org/10.1175/JCLI-D-12-00826.1>, 2013.
- ☐ Pinto, J. P., Turco, R. P., and Toon, O. B.: Self-limiting physical and chemical effects in volcanic eruption clouds, *J. Geophys. Res.*, 94, 11165, <https://doi.org/10.1029/JD094iD08p11165>, 1989.
- ☐ Lacis, A., Hansen, J., and Sato, M.: Climate forcing by stratospheric aerosols, *Geophys. Res. Lett.*, 19, 1607, <https://doi.org/10.1029/92GL01620>, 1992.
- ☐ Clyne, M., Lamarque, J. F., Mills, M. J., Khodri, M., Ball, W., Bekki, S., Dhomse, S. S., Lebas, N., Mann, G., Marshall, L., Niemeier, U., Poulain, V., Robock, A., Rozanov, E., Schmidt, A., Stenke, A., Sukhodolov, T., Timmreck, C., Toohey, M., Tummon, F., Zanchettin, D., Zhu, Y., and Toon, O. B.: Model physics and chemistry causing intermodel disagreement within the VolMIP-Tambora Interactive Stratospheric Aerosol ensemble, *Atmos. Chem. Phys.*, 21, 3317-3343, <https://doi.org/10.5194/ACP-21-3317-2021>, 2021.
- ✓ Marshall, L. R., Smith, C. J., Forster, P. M., Aubry, T. J., Andrews, T., and Schmidt, A.: Large variations in volcanic aerosol forcing efficiency due to eruption source parameters and rapid adjustments, *Geophys. Res. Lett.*, 47, <https://doi.org/10.1029/2020gl090241>, 2020. (Marshall et al., 2020)
- ☐ Hansen, J., Sato, M., Ruedy, R., Nazarenko, L., Lacis, A., Schmidt, G. A., Russell, G., Aleinov, I., Bauer, M., Bauer, S., Bell, N., Cairns, B., Canuto, V., Chandler, M., Cheng, Y., Genio, A. D., Faluvegi, G., Fleming, E., Friend, A., Hall, T., Jackman, C., Kelley, M., Kiang, N., Koch, D., Lean, J., Lerner, J., Lo, K., Menon, S., Miller, R., Minnis, P., Novakov, T., Oinas, V., Perlwitz, J., Perlwitz, J., Rind, D., Romanou, A., Shindell, D., Stone, P., Sun, S., Tausnev, N., Thresher, D., Wielicki, B., Wong, T., Yao, M., & Zhang, S. (2005). Efficacy of climate forcings. *Journal of Geophysical Research: Atmospheres*, 110(D18). <https://doi.org/10.1029/2005JD005776>

Reviewer #3 Evaluations

- Recommendation: Return to author for major revisions
- Significant: The paper has some unclear or incomplete reasoning but will likely be a significant contribution with revision and clarification.
- Supported: Mostly yes, but some further information and/or data are needed.
- Referencing: Mostly yes, but some additions are necessary.
- Quality: The organization of the manuscript and presentation of the data and results need some improvement.

- Data: Please Select
- Accurate Key Points: No

Reviewer #3 (Formal Review for Author (shown to authors))

This study investigates the atmospheric radiative forcing by super eruptions. The new aspect of this paper is on the relationship of effective radiative forcing, aerosol optical depth and surface temperature by eruptions larger than Pinatubo, which is a long-standing topic in the field of volcanic eruptions and their impacts on radiation and climate. The paper presents new model results from the CES-M2(WACCM6) model with 2 deg horizontal resolution, prescribed sea surface temperature and sea ice, 1850 conditions, and a Middle Atmosphere chemistry configuration taking different strength of sulfur-only emissions into the stratosphere, from Pinatubo strength to super-size eruption, into account. A non-linear behavior is modelled for Effective Radiative Forcing (ERF) and Aerosol Optical Depth (AOD) for eruptions larger than Pinatubo. This has been shown for Radiative Forcing (RF) estimates and model calculations before and is well accepted and taken over by the community. Thus, the authors need to substantially revise the paper to clarify the new angle in the context of the existing, large body of literature. I believe this will be feasible for the authors when addressing my below comments carefully. In that case I would be happy to review a thoroughly revised manuscript again. See my general major and minor comments, and a list of references to take care of below.

General major comments

- The background, status quo of the research field, and what is new in this study has to be carefully revised. There are many studies already published on the non-linearity of volcanic eruption strength impacts on AOD and RF. This has been taken over by the volcano climate model community. This background/status quo needs to be carefully included and revised in the ms; original references need to be cited. Following, pls reformulate what is new and different in your study. See a list of key references for an overview below.

i Reply

More on the non-linearities are now described in the introduction, as well as in the discussion. It should be more clear now that the focus regarding the non-linear relationship between AOD and ERF is on the development over time in the post-eruption period rather than across SO₂ injection magnitudes.

- Abstract and elsewhere in the ms need carefully rewording: “Unlike the traditional linear models used for smaller eruptions, our results reveal...” I would argue against that a linear model is still used (“traditionally”) for large eruptions such as Pinatubo (which are here called small).

i Reply

Re-worded slightly to specify the validity up to Mt. Pinatubo-sized eruption, and that it is known that non-linearities are found for even larger eruptiona.

- The authors need to introduce and define the concept of Effective Radiative Forcing in contrast to Radiative Forcing used in other studies.

i Reply

This is now included in the introduction.

- There are several super-eruptions publications which are relevant for this study, but they are not included and referred to: Bekki et al 1996; Robock et al 2009; Timmreck et al 2010 and 2012; Brenna et al 2020, MacGraw & Polvani, 2024. Brenna et al 2020 shows the non-linear effect and the disagreement between models with increasing strength of eruption.

i Reply

Several new studies have now been included in figure 4, as well as referenced to and included as part of the discussing in relation to our simulation results.

- AOD-RF relationship depends on the effective radius etc (Pinto et al 1989; Bekki et al 1995&1996; Timmreck et al 2009 ff).

i Reply

This has also been included as part of the discussion and effective radius is introduced in greater depth in the introduction.

- Different eruption strength, season and latitude and their impact on AOD, RF, climate, etc is published already in Schneider et al 2009, Toohey et al 2011, 2013, 2019, Metzner et al 2014, Marshall et al 2019ff, Quaglia et al 2023, Zhuo et al 2024.

i Reply

We hope the focus in relation to ERF/AOD has now been made more clear in the direction of the ratios development over time post-eruption. The provided citations have also been included in the discussion of the results.

- Model uncertainties on the SO₂-AOD-effective radius of aerosols: Bekki et al 1995&1996, Pinto et al 1989, Timmreck et al 2009 & 2010; Toohey et al 2011/2013/2019, Marshall et al 2018; Brenna et al 2020, Quaglia et al 2023; MacGraw et al 2024, Zhuo et al 2024.

i Reply

We now include studies describing the effective radius, and compare in particular with Brenna et al. (2020) who ran the same model and did a much more in-depth analysis of the effective radius. McGraw et al. (2024) also provide a useful comparison across a wide range of eruption magnitudes that we now refer to.

- Timing of maximum AOD, RF, climate for tropical versus high-latitude eruptions see Zhuo et al 2024.

i Reply

We now also included the results from Zhuo et al. (2024) in the introduction. We hope that it is more clear from the text now that this is not a main focus of the manuscript.

- How can you address surface temperature response if the ocean-sea ice is prescribed? I suggest taking this field out or using lower tropospheric temperature instead as well as discussing the corresponding uncertainties properly.

i Reply

We try now to be more explicit about which type of simulation the different output fields come from, in the case of temperature from a fully dynamic simulation.

- The model uncertainties need to be discussed (see Zanchettin et al 2016; Marshall et al 2018; Brenna et al 2020). CESM(WACCM) tend to simulate smaller aerosols sizes and thus a longer lifetime of and large AOD, RF and climate than the other existing aerosols climate models (Fig. 3a Zanchettin et al 2016).

i Reply

This has been a valuable addition to the manuscript, and how our results compared to results from other studies are now placed in the context of how CESM(WACCM) has been found to simulate aerosol effective radius. In particular, Brenna et al. (2020) who used the same model with similar experiments provides a good comparison to our results, and includes effective radius in their analysis that we refer to.

- Figs 2-4: Need to include the other existing Pinatubo-strength and supereruption studies and data (see references).

i Reply

Fig2 for peak values is shown in Fig4d, and for Fig3 we have kept only the data we have been able to get full time series of, so Fig2 and Fig3 are therefore the same. Several new studies have been included to Fig.4 in response to your comment.

Minor comments

- ✓ “Super-eruption” but not super-volcano eruption (see f.e. Wilson et al 2021).
- ✓ Define super eruption and other strength of eruptions. Pinatubo is not a “small” eruption. (cf. Metzner et al 2014; Schmidt & Black 2022).
- ✓ Model experiments details such as fixed SST and sea ice, 1850 conditions need to go into Section 2.2 as they are important for the interpretation of the results.
- ✓ MAM3 add on changes for stratospheric aerosols need to cite Mills et al (2016) Mills, M. J., et al. (2016), Global volcanic aerosol properties derived from emissions, 1990-2014, using CESM1(WACCM), J. Geophys. Res. Atmos., 121, <https://doi.org/10.1002/2015JD024290>.
- ✓ Fig. 1 and ff Figs.: Which temperature field is plotted here; does it make sense if the ocean/sea-ice is prescribed?

References

Non-linear behavior

- Pinto et al. 1989, Hyde and Crowley 2000; Ammann et al. 2003; Gao et al. 2008; Metzner et al 2014. For large eruptions (see Methodology Section and References in Metzner et al 2014)
- Metzner, D., Kutterolf, S., Toohey, M. et al. Radiative forcing and climate impact resulting from SO₂ injections based on a 200,000-year record of Plinian eruptions along the Central American Volcanic Arc. Int J Earth Sci (Geol Rundsch) 103, 2063-2079 (2014). <https://doi.org/10.1007/s00531-012-0814-z>

Effective radius

- Pinto, J. R., R. P. Turco, and O. B. Toon (1989), Self-limiting physical and chemical effects in volcanic eruption clouds, J. Geophys. Res., 94, 11,165-11,174.

- Bekki et al 1995 Oxidation of volcanic SO₂: A sink for stratospheric OH and H₂O - Geophysical Research Letters
- Bekki et al 1996 The role of microphysical and chemical processes in prolonging the climate forcing of the Toba Eruption - Geophysical Research Letters
- Timmreck et al 2009 - Limited temperature response to the very large AD 1258 volcanic eruption - Geophysical Research Letters
- Quaglia, I., Timmreck, C., Niemeier, U., Visionsi, D., Pitari, G., Brodowsky, C., Brühl, C., Dhomse, S. S., Franke, H., Laakso, A., Mann, G. W., Rozanov, E., and Sukhodolov, T.: Interactive stratospheric aerosol models' response to different amounts and altitudes of SO₂ injection during the 1991 Pinatubo eruption, *Atmos. Chem. Phys.*, 23, 921-948, <https://doi.org/10.5194/acp-23-921-2023>, 2023.

Effective Radiative Forcing

- Forster, P. M., T. Richardson, A. C. Maycock, C. J. Smith, B. H. Samset, G. Myhre, T. Andrews, R. Pincus, and M. Schulz (2016), Recommendations for diagnosing effective radiative forcing from climate models for CMIP6, *J. Geophys. Res. Atmos.*, 121, 12,460-12,475, <https://doi.org/10.1002/2016JD025320>.
- Marshall et al 2021 - Unknown Eruption Source Parameters Cause Large Uncertainty in Historical Volcanic Radiative Forcing Reconstructions - *Journal of Geophysical Research: Atmospheres* (Marshall et al., 2021)

Super eruptions

- Bekki et al 1996 The role of microphysical and chemical processes in prolonging the climate forcing of the Toba Eruption - Geophysical Research Letters
- Brenna, H., Kutterolf, S., Mills, M. J., and Krüger, K.: The potential impacts of a sulfur- and halogen-rich supereruption such as Los Chocoyos on the atmosphere and climate, *Atmos. Chem. Phys.*, 20, 6521-6539, <https://doi.org/10.5194/acp-20-6521-2020>, 2020.
- Brenna et al 2021 Decadal Disruption of the QBO by Tropical Volcanic Supereruptions - Geophysical Research Letters
- English et al 2013 - Microphysical simulations of large volcanic eruptions: Pinatubo and Toba - *Journal of Geophysical Research: Atmospheres* (English et al., 2013)
- McGraw, Z., K. DallaSanta, L. M. Polvani, K. Tsigaridis, C. Orbe, and S. E. Bauer, 2024: Severe Global Cooling After Volcanic Super-Eruptions? The Answer Hinges on Unknown Aerosol Size. *J. Climate*, 37, 1449-1464, <https://doi.org/10.1175/JCLI-D-23-0116.1>.
- Osipov, S., G. Stenchikov, K. Tsigaridis, A.N. LeGrande, S.E. Bauer, M. Fnais, and J. Lelieveld, 2021: The Toba supervolcano eruption caused severe tropical stratospheric ozone depletion. *Communications Earth Environ.*, 2, no. 1, 71, <https://doi.org/10.1038/s43247-021-00141-7>.
- Robock, A., C.M. Ammann, L. Oman, D. Shindell, S. Levis, and G. Stenchikov, 2009: Did the Toba volcanic eruption of ~74k BP produce widespread glaciation? *J. Geophys. Res.*, 114, D10107, <https://doi.org/10.1029/2008JD011652>.
- Timmreck et al 2010, Aerosol size confines climate response to volcanic super-eruptions -Geophysical Research Letters (Timmreck, Graf, Lorenz, Niemeier, Zanchettin, Matei, Jungclaus, & Crowley, 2010)
- Timmreck et al, Climate response to the Tobasuper-eruption: Regional changes, *Quaternary International*, Volume 258, <https://doi.org/10.1016/j.quaint.2011.10.008>, 2012
- Wilson, C.J.N., Cooper, G.F.,Chamberlain, K.J. et al. No single model for supersized eruptions and their magmabodies. *Nat Rev Earth Environ* 2, 610-627 (2021). <https://doi.org/10.1038/s43017-021-00191-7>

Different aerosol climate models and their uncertainties (for Tambora)

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