"FLOW AND SEDIMENT TRANSPORT MODELING IN RIVER BASINS USING TELEMAC 2D AND 3D NUMERICAL CODES"

February 26-28, 2022





Sediment transport module - GAIA

Steering GAIA: parameters and data files

Chen Peng-An

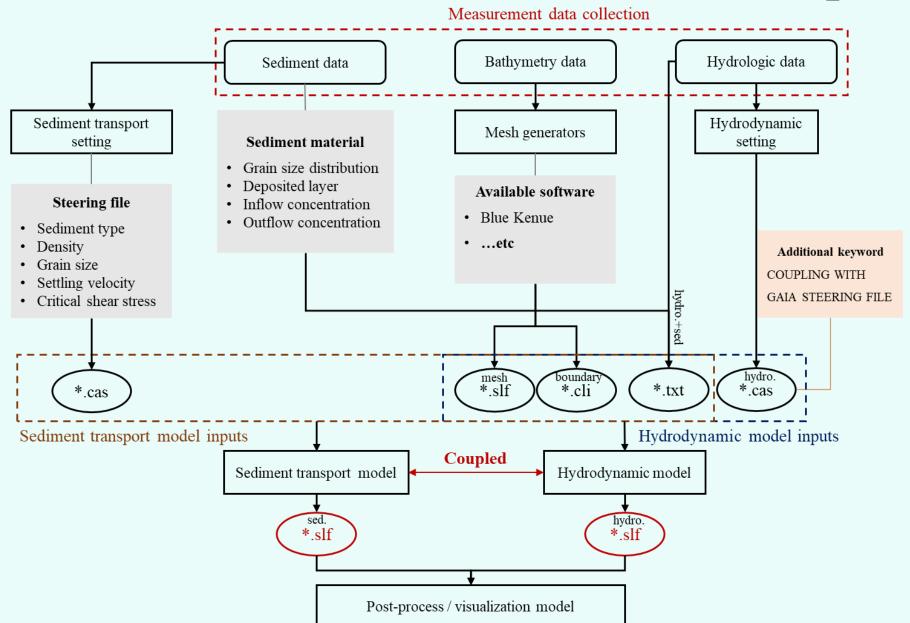


BASINS USING TELEMAC 2D AND 3D NUMERICAL CODES"

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Structure of sediment transport





"FLOW AND SEDIMENT TRANSPORT MODELING IN RIVER **BASINS USING TELEMAC 2D AND 3D NUMERICAL CODES"**

Input file: BOUNDARY **CONDITION FILE**





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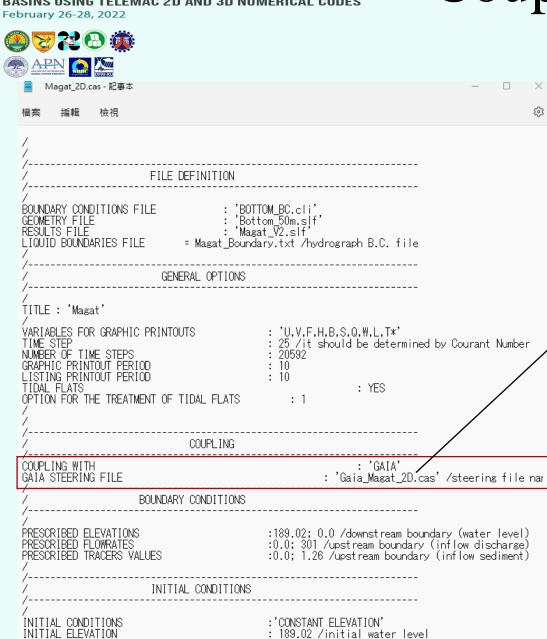


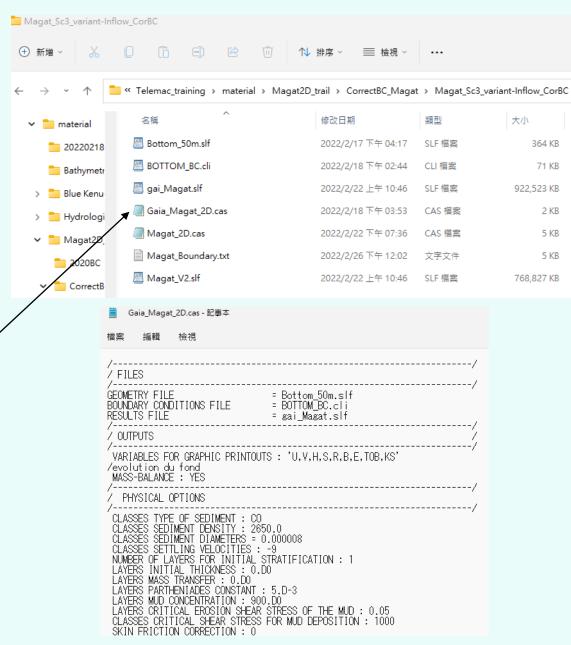


| LIHBOR | LIUBOR | LIVBOR | LITBOR | |
|--------|--------|--------|--------|--|
| 2 | 2 | 2 | 2 | Solid wall. |
| 2 | 0 | 2 | 2 | Solid wall with zero U . |
| 2 | 2 | 0 | 2 | Solid wall with zero V. |
| 2 | 0 | 0 | 2 | Solid wall with zero U and V . |
| 4 | 4 | 4 | 4 | Free H , free velocities, free T . |
| 5 | 4 | 4 | 4 | Prescribed H , free velocities, free T . |
| 5 | 4 | 0 | 4 | Prescribed H , free U , zero V , free T . |
| 5 | 0 | 4 | 4 | Prescribed H , zero U , free V , free T . |
| 1 | 1 | 1 | 4 | Incident wave, free tracer. |
| 4 | 5 | 5 | 5 | Free H , prescribed Q , prescribed T . |
| 4 | 5 | 0 | 5 | Free H , prescribed Q with zero V , prescribed T . |
| 4 | 0 | 5 | 5 | Free H , prescribed Q with zero U , prescribed T . |
| 4 | 6 | 6 | 5 | Free H , prescribed velocities, prescribed T . |
| 5 | 5 | 5 | 5 | Prescribed H and Q , prescribed T . |
| 5 | 6 | 6 | 5 | Prescribed H and velocities, prescribed T . |

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Coupled with GAIA







2 KB

5 KB

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Physical parameters: Suspended load transport



- Suspended load: with the solution of the advection-diffusion equation (ADE) plus closures for erosion and deposition fluxes, equilibrium concentration
 - two-dimensional advection-diffusion equation:

$$\frac{\partial hC}{\partial t} + \frac{\partial hUC}{\partial x} + \frac{\partial hVC}{\partial y} = \frac{\partial}{\partial x} \left(h\varepsilon_s \frac{\partial C}{\partial x} \right) + \frac{\partial}{\partial y} \left(h\varepsilon_s \frac{\partial C}{\partial y} \right) + E - D$$

where C = C(x, y, t) is the depth-averaged concentration **expressed in g/l**, (U, V) are the depth-averaged components of the velocity in the x and y directions, respectively, ε_s is the turbulent diffusivity of the sediment, often related to the eddy viscosity $\varepsilon_s = v_t/\sigma_c$, with σ_c the Schmidt number. In our case, $\sigma_c = 1.0$.

• three-dimensional advection-diffusion equation:

$$\frac{\partial C}{\partial t} + \frac{\partial uC}{\partial x} + \frac{\partial vC}{\partial y} + \frac{\partial wC}{\partial z} - \frac{\partial w_sC}{\partial z} = \frac{\partial}{\partial x} \left(\varepsilon_s \frac{\partial C}{\partial x} \right) + \frac{\partial}{\partial y} \left(\varepsilon_s \frac{\partial C}{\partial y} \right) + \frac{\partial}{\partial y} \left(\varepsilon_s \frac{\partial C}{\partial z} \right)$$

where C = C(x, y, z, t) is the concentration **expressed in g/l**, (u, v, w) are the components of the velocity in the x, y and z directions, respectively. w_s is the sediment settling velocity. ε_s is the turbulent diffusivity of the sediment, often related to the eddy viscosity $\varepsilon_s = v_t/\sigma_c$, with σ_c the Schmidt number. In our case $\sigma_c = 1.0$.



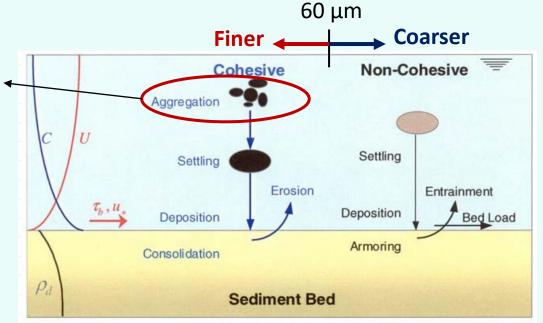


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Sediment transport: Cohesive suspended load



Aggregation of flocs can lead to the formation of macroflocs larger than 100 µm.



In GAIA, cohesive sediments are accounted by solving the 2D advection-diffusion equation:

$$\frac{\partial hC}{\partial t} + \frac{\partial hUC}{\partial x} + \frac{\partial hVC}{\partial y} = \frac{\partial}{\partial x} \left(h\varepsilon_s \frac{\partial C}{\partial x} \right) + \frac{\partial}{\partial y} \left(h\varepsilon_s \frac{\partial C}{\partial y} \right) + (E - D)$$

C = C(x, y, t) is the depth-averaged concentration expressed in g/I, (U, V) are the depth-averaged components of the velocity in the x and y directions, respectively, ε_s is the turbulent diffusivity of the sediment. Variables E and D are respectively the erosion and deposition fluxes.



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Advection-diffusion equation:





Numerical treatment of the diffusion terms:

The keyword OPTION FOR THE DIFFUSION OF TRACER (integer type, set = 1 by default) allows to choose the treatment of the diffusion terms in the advection-diffusion equation 2.1 for the depth-averaged suspended concentration:

- = 1: the diffusion term is solved in the form $\nabla \cdot (\varepsilon_s \nabla T)$
- = 2: the diffusion term is solved in the form $\frac{1}{h}\nabla \cdot (h\varepsilon_s \nabla T)$

This keyword must be activated in the steering file of TELEMAC-2D; user can refer to the corresponding user manual for further details.

Numerical treatment of the advection terms

The choice for the scheme for the treatment of the advection terms can be done with the keyword SCHEME FOR ADVECTION OF SUSPENDED SEDIMENTS (integer type, set = 5 by default):

1="CHARACTERISTICS"

2="SUPG"

3="CONSERVATIVE N-SCHEME"

4="CONSERVATIVE N-SCHEME"

5="CONSERVATIVE PSI-SCHEME"

13="EDGE-BASED N-SCHEME"

14="EDGE-BASED N-SCHEME"

15="ERIA SCHEME"

It is recommended to use the schemes 4 or 14 for a good compromise between accuracy and computational time (specially if tidal flats are present).

• Method of characteristics (1)

- Unconditionally stable and monotonous
- Diffusive for small time steps, Not conservative
- Method Streamline Upwind Petrov Galerkin SUPG (2)
 - Based on the Courant number criteria
 - Less diffusive for small time steps, Not conservative
- Conservative N-scheme (similar to finite volumes) (3, 4)
 - Solves the continuity equation under its conservative form
 - Recommended for correction on convection velocity
 - Courant number limitation (sub-iterations to reduce time step)
- Edge-based N-scheme NERD (13, 14)
 - Same as 3 and 4 but adapted to tidal flats
 - Based on positive-depth algorithm
- Distributive schemes PSI (5)
 - fluxes corrected according to the tracer value: relaxation of Courant number criteria, less diffusive than schemes 4, 14 but larger CPU time
 - Should not be applied for tidal flats
- Eria scheme (15)
 - Works for tidal flats

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Sediment transport: Bedload transport





In contrast to SISYPHE, in GAIA bedload fluxes are computed in terms of (dry) mass transport rate per unit width, without pores.

The dimensionless current-induced sediment transport rate (Φ_b) is expressed:

$$\Phi_b = \frac{Q_b}{\sqrt{g(s-1)D_{50}^3}}$$

where, Q_b is the bedload transport rate per unit width (m²/s); $s = \rho_s/\rho_a$ is the relative density; ρ_s and ρ_a is the density of sediment and clear ware (kg/m³), respectively; D_{50} means sediment grain size (m).

To obtain the sediment transport rate, the bedload transport formulas are commonly

computed as the function of Shield number (θ) :

$$\theta = \frac{\mu \tau_b}{(\rho_s - \rho_a)gd}$$

where, μ is the skin friction's correction factor.

Solved by bedload transport formula

Keyword BED-LOAD TRANSPORT FORMULA FOR ALL SANDS (integer type variable, set to = 1 by default) can be used to set a bedload transport formula. Available formulas in GAIA for bedload transport are:

- 1 : MEYER-PETER and MUELLER
- 2 : EINSTEIN-BROWN
- 3: ENGELUND-HANSEN + CHOLLET ET CUNGE (total sediment transport)
- 10: WILCOCK AND CROWE
- 30: ENGELUND-HANSEN (total sediment transport)
- 7 : VAN RIJN

For example, the keyword BED-LOAD TRANSPORT FORMULA FOR ALL SANDS = 7 sets the van Rijn formula. Please note that bedload transport formulas 3 and 30 account for the total sediment transport.

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Sediment Exchanges: Non-cohesive sediment





For non-cohesive sediments, the net sediment flux E - D is therefore determined based on the

concept of equilibrium concentration: $(E-D)_{z_{ref}} = w_s (C_{eq}) - (C_{z_{ref}})$

The equilibrium near-bed concentration determined by using an empirical formula.

Zyserman-Fredsoe

$$C_{eq} = \frac{0.331(\theta' - \theta_{cr})^{1.75}}{1 + 0.72(\theta' - \theta_{cr})^{1.75}},$$

Bijker

$$C_{eq} = \frac{Q_b}{b \, z_{ref} \, u_*}$$

van Rijn

$$C_{eq} = 0.015 d_{50} \frac{(\theta'/\theta_{cr} - 1)^{3/2}}{z_{ref} D_*^{0.3}}$$

Soulsby & van Rijn

$$C_{eq} = \begin{cases} A_{ss} \left(\sqrt{U_c^2 + \frac{0.018}{C_D} U_w^2 - U_{cr}} \right)^{2.4} & \text{if } \ge U_{cr} \\ 0.0 & \text{otherwise} \end{cases}$$

The near-bed concentration, evaluated at the interface between the bed load and the suspended load, $z = z_{ref}$.

In 2D cases, the near-bed concentration is computed assuming a Rouse profile for the vertical concentration distribution, which is theoretically valid in uniform steady flow conditions:

$$C(z) = C_{z_{ref}} \left(\frac{z - h}{z} \frac{a}{a - h} \right)^{R}, \tag{4.1}$$

where R is the Rouse number defined by

$$R = \frac{w_s}{\kappa u_*},\tag{4.2}$$

with κ the von Karman constant ($\kappa = 0.4$), u_* the friction velocity corresponding to the total bed shear stress, and a the reference elevation above the bed elevation. The distance a, defined variously by various authors, is taken to be very close to the bed.

By depth-integration of the Rouse profile (4.3), the following relation can be established between the depth-averaged concentration and the reference concentration (near-bed concentration):

$$C_{z_{ref}} = FC$$

where:

$$F^{-1} = \left(\frac{z_{ref}}{h}\right)^R \int_{z_{ref}/h}^1 \left(\frac{1-u}{u}\right)^R du. \tag{4.3}$$

In GAIA, the following expression is used to compute *F*:

$$F^{-1} = \begin{cases} \frac{1}{(1-Z)} B^R \left(1 - B^{(1-R)} \right) & \text{if } R \neq 1 \\ -B \log B & \text{if } R = 1 \end{cases}$$

with
$$B = z_{ref}/h$$
.

"FLOW AND SEDIMENT TRANSPORT MODELING IN RIVER BASINS USING TELEMAC 2D AND 3D NUMERICAL CODES"

Sediment Exchanges: Cohesive and mixed sediment





Erosion flux

The erosion flux is computed with the Partheniades formula. For uniform beds, the erosion flux is related to the excess of applied bed shear stress to the bed shear strength at the bed surface:

$$E = \begin{cases} M \left[\left(\frac{\tau_b}{\tau_{ce}} \right) - 1 \right] & \text{if } \tau_b > \tau_{ce} \\ 0 & \text{otherwise} \end{cases}$$

where M the Krone-Partheniades erosion law constant [kg/m²/s] and τ_{ce} the critical bed shear stress.

The composition of the sediment mixture in the surface (active) layer is taken into consideration when computing the critical shear stress for erosion and the erosion rate. This is achieved by combining the critical shear stresses for erosion for all the sediment classes (cohesive and non-cohesive), according to [30]:

- If the mass of cohesive sediment as a fraction of the mixture is ≥ 50%, then the erosion
 rate and critical shear stress for cohesive sediment alone is used.
- If the mass of cohesive sediment as a fraction of the mixture is ≤ 30%, then the erosion rate for non-cohesive sediment is used and the critical shear stress for non-cohesive sediment is used with a correction.
- If the mass of cohesive sediment as fraction of the mixture is ≥ 30% and ≤ 50%, then the
 values are interpolated between the previous values.

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Sediment Exchanges: Cohesive and mixed sediment



Deposition flux

The deposition flux for mud is computed by the expression:

$$D = w_s C \left[1 - \left(\frac{\sqrt{\tau_b/\rho}}{u_{*mud}^{cr}} \right)^2 \right]$$

 $D = \underbrace{w_s C} \left[1 - \left(\frac{\sqrt{\tau_b/\rho}}{u_{*mud}^{cr}} \right)^2 \right] \qquad \text{where } u_{*mud}^{cr} \text{ is the critical shear velocity for mud deposition, expressed in [m/s] and computed as } \sqrt{\tau_{d,mud}/\rho} \text{ with } \tau_{d,mud} \text{ provided by the keyword CLASSES CRITICAL SHEAR STRESS FOR MUD DEPOSITION (real type, set to = 1000. N/m² by default).}$

Settling velocity, it can be specified by the user or calculated by the model as a function of grain diameter, keyword **SETTLING VELOCITIES** (real list, =-9 by default).

$$w_s = \begin{cases} \frac{(s-1)gd_{50}^2}{18\nu}, & \text{if } d_{50} \le 10^{-4} \\ \frac{10\nu}{d_{50}} \left(\sqrt{1 + 0.01 \frac{(s-1)gd_{50}^3}{\nu^2}} - 1 \right), & \text{if } 10^{-4} \le d_{50} \le 10^{-3} \\ 1.1\sqrt{(s-1)gd_{50}}, & \text{otherwise} \end{cases}$$

with $s = \rho_s/\rho_0$ is the relative density and g is the acceleration of the gravity.

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Coupled with GAIA

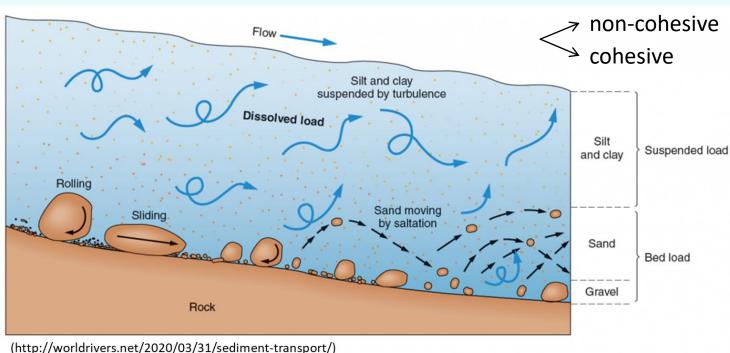




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= BOTTOM_BC.cli
BOUNDARY CONDITIONS FILE
                 = gai_Magat.slf
.
VARIABLES FOR GRAPHIC PRINTOUTS : 'U,V,H,S,R,B,E,TOB,KS'
/evolution du fond
MASS-BALANCE : YES
 CLASSES TYPE OF SEDIMENT : CO
CLASSES SEDIMENT DENSITY: 2650.0
 CLASSES SEDIMENT DIAMETERS = 0.000008
 CLASSES SETTLING VELOCITIES : -9
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 LAYERS INITIAL THICKNESS: 0.DO
 LAYERS MASS TRANSFER : 0.D0
LAYERS PARTHENIADES CONSTANT : 5.D-3
LAYERS MUD CONCENTRATION: 900.DO

LAYERS CRITICAL EROSION SHEAR STRESS OF THE MUD: 0.05

CLASSES CRITICAL SHEAR STRESS FOR MUD DEPOSITION: 1000
 SKIN FRICTION CORRECTION: 0
ADVECTION-DIFFUSION SCHEME WITH SETTLING VELOCITY: 1 / SET_DIF scheme'
SOLVER FOR DIFFUSION OF SUSPENSION = 1
```



Cohesive suspended sediment transportation

BASINS USING TELEMAC 2D AND 3D NUMERICAL CODES"

Physical parameters: settling velocity

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$$w_s = \begin{cases} \frac{(s-1)gd_{50}^2}{18\nu}, & \text{if } d_{50} \le 10^{-4} \\ \frac{10\nu}{d_{50}} \left(\sqrt{1 + 0.01 \frac{(s-1)gd_{50}^3}{\nu^2}} - 1 \right), & \text{if } 10^{-4} \le d_{50} \le 10^{-3} \\ 1.1\sqrt{(s-1)gd_{50}}, & \text{otherwise} \end{cases}$$

with $s = \rho_s/\rho_0$ is the relative density and g is the acceleration of the gravity.

Rubey:

USE DECLARATIONS_GAIA IMPLICIT NONE INTEGER :: ICLA DOUBLE PRECISION :: DENS SETTLING VELOCITY DO ICLA = 1, NSICLA IF(XWC0(ICLA).LT.-3) THEN SETTLING VELOCITY IS NOT GIVEN IN THE PARAMETER FILE AND INITIALISED IF T3D: SETTLING VELOCITY IS POTENTIALY MODIFIED IN 3D AND THEN GIVEN DENS = (XMVS0(ICLA) - XMVE) / XMVE IF (DCLA(ICLA).LT.1.D-4) THEN XWC0(ICLA) = DENS * DCLA(ICLA) * DCLA(ICLA) * GRAV / (18.D0 * VCE) ELSEIF (DCLA(ICLA).LT.1D-3) THEN $\frac{[1.636(s-1)d_{50}^3 + 9v^2]^{0.5} - 3v}{}$ XWCO(ICLA) = 10.DO * VCE / DCLA(ICLA) :(SQRT(1.D0 + 0.01D0* DENS * GRAV * DCLA(ICLA)**3.D0 / (VCE*VCE)) -1.D0) XWC0(ICLA) = 1.1D0 * SQRT(DENS * GRAV * DCLA(ICLA)) ELSEIF (XWC0(ICLA).EQ.-1) THEN DENS = (XMVS0(ICLA) - XMVE) / XMVE XWC0(ICLA) =(XMVE * (SQRT((1636/XMVE) * DENS * DCLA(ICLA)**3.D0 + 9 * VCE * VCE) - 3 * VCE)) / (500 * DCLA(ICLA))

XWC(ICLA)=XWC0(ICLA)

SUBROUTINE SETTLING_VEL

!>@brief Compute the settling velocities according to the Stokes, Zanke and Van Rijn formulation, in case of non cohesive sediments. Sets the default value, in case of cohesive

.

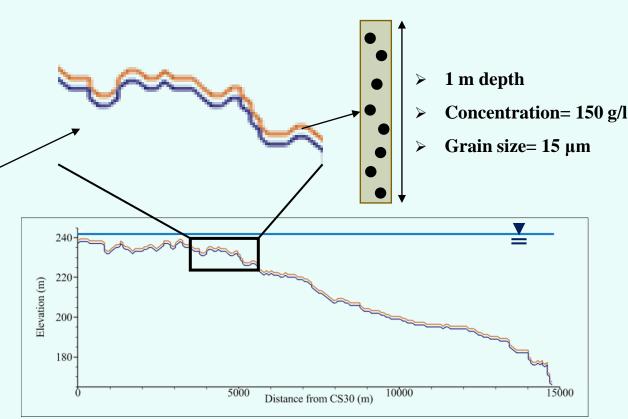


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Physical parameters: Accumulated deposition layer









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Physical parameters: Accumulated deposition layer





```
Gaia_Magat_2D.cas - 記事本
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GEOMETRY FILE = Bottom_50m.slf
BOUNDARY CONDITIONS FILE = BOTTOM_BC.cli
RESULTS FILE = gai_Magat.slf
 .
VARIABLES FOR GRAPHIC PRINTOUTS : 'U,V,H,S,R,B,E,TOB,KS'
/evolution du fond
 MASS-BALANCE : YES
/-----/
CLASSES TYPE OF SEDIMENT: CO
CLASSES SEDIMENT DENSITY: 2650.0
CLASSES SEDIMENT DIAMETERS = 0.0000008
CLASSES SETTLING VELOCITIES: -9
NUMBER OF LAYERS FOR INITIAL STRATIFICATION: 1
LAYERS INITIAL THICKNESS: 0.D0
LAYERS MASS TRANSFER: 0.D0
LAYERS MASS TRANSFER: 0.D0
LAYERS MID CONCENTRATION: 0.00.D0
  LAYERS MUD CONCENTRATION : 900 DO
 LAYERS CRITICAL EROSION SHEAR STRESS OF THE MUD: 0.05
CLASSES CRITICAL SHEAR STRESS FOR MUD DEPOSITION: 1000
  SKIN FRICTION CORRECTION : 0
 ADVECTION-DIFFUSION SCHEME WITH SETTLING VELOCITY: 1 / SET_DIF scheme SOLVER FOR DIFFUSION OF SUSPENSION = 1
```

