

A Conforming Finite Elements for the Pure Streamfunction Form of the Quasi-Geostrophic Equations

Erich L Foster

Department of Mathematics
Virginia Tech

09 February 2013

- In collaboration with: Traian Iliescu (VT), Zhu Wang (IMA), and Dave Wells (VT)
- National Science Foundation Grant DMS-1025314
- Institute for Critical Technology and Applied Science (ICTAS) fund 118709

- In collaboration with: Traian Iliescu (VT), Zhu Wang (IMA), and Dave Wells (VT)
- National Science Foundation Grant DMS-1025314
- Institute for Critical Technology and Applied Science (ICTAS) fund 118709

- In collaboration with: Traian Iliescu (VT), Zhu Wang (IMA), and Dave Wells (VT)
- National Science Foundation Grant DMS-1025314
- Institute for Critical Technology and Applied Science (ICTAS) fund 118709

1 Large Scale Ocean Surface Currents

2 Quasi-Geostrophic Equations

3 Argyris Finite Element

4 Error Estimates

5 Time Dependence

6 Future Work

1 Large Scale Ocean Surface Currents

2 Quasi-Geostrophic Equations

3 Argyris Finite Element

4 Error Estimates

5 Time Dependence

6 Future Work

1 Large Scale Ocean Surface Currents

2 Quasi-Geostrophic Equations

3 Argyris Finite Element

4 Error Estimates

5 Time Dependence

6 Future Work

1 Large Scale Ocean Surface Currents

2 Quasi-Geostrophic Equations

3 Argyris Finite Element

4 Error Estimates

5 Time Dependence

6 Future Work

1 Large Scale Ocean Surface Currents

2 Quasi-Geostrophic Equations

3 Argyris Finite Element

4 Error Estimates

5 Time Dependence

6 Future Work

1 Large Scale Ocean Surface Currents

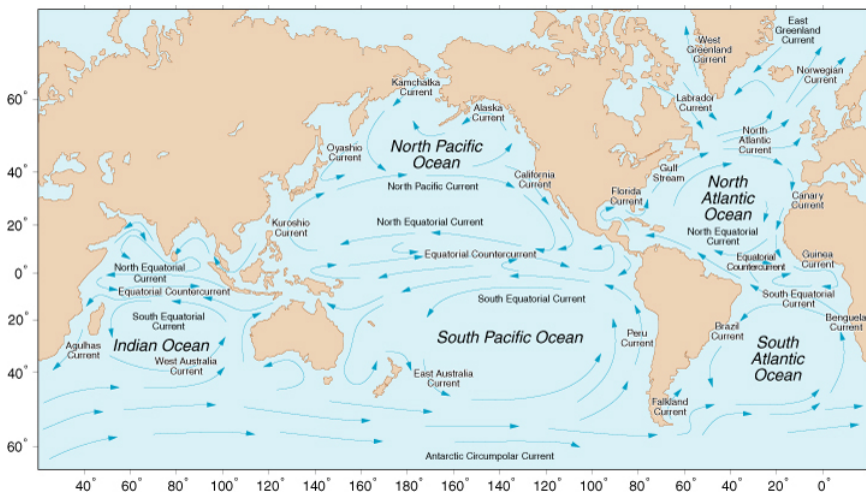
2 Quasi-Geostrophic Equations

3 Argyris Finite Element

4 Error Estimates

5 Time Dependence

6 Future Work



© 2005 American Meteorological Society

[src:http://oceanmotion.org/html/background/wind-driven-surface.htm](http://oceanmotion.org/html/background/wind-driven-surface.htm)

Large scale surface currents of the oceans

Characteristics of large scale oceanic surface currents

- Driven by forces such as
 - Wind
 - Coriolis Force, i.e. the deflection of moving objects due to the rotation of the Earth.

Quasi-Geostrophic Equations

- The QGE are a simplified model for planet-scale flows.
- Streamfunction-Vorticity Formulation

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F \quad (1)$$

$$q = -Ro \Delta \psi + y \quad (2)$$

where Ro , Re are the Rossby and Reynolds numbers, and the Jacobian

$$J(\xi, \eta) = \frac{\partial \xi}{\partial x} \frac{\partial \eta}{\partial y} - \frac{\partial \xi}{\partial y} \frac{\partial \eta}{\partial x}$$

- The Jacobian is associated with $(u \cdot \nabla) u$ in the Navier-Stokes Equations

Characteristics of large scale oceanic surface currents

- Driven by forces such as
 - Wind
 - Coriolis Force, i.e. the deflection of moving objects due to the rotation of the Earth.

Quasi-Geostrophic Equations

- The QGE are a simplified model for planet-scale flows.
- Streamfunction-Vorticity Formulation

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F \quad (1)$$

$$q = -Ro \Delta \psi + y \quad (2)$$

where Ro , Re are the Rossby and Reynolds numbers, and the Jacobian

$$J(\xi, \eta) = \frac{\partial \xi}{\partial x} \frac{\partial \eta}{\partial y} - \frac{\partial \xi}{\partial y} \frac{\partial \eta}{\partial x}$$

- The Jacobian is associated with $(u \cdot \nabla) u$ in the Navier-Stokes Equations

Characteristics of large scale oceanic surface currents

- Driven by forces such as
 - Wind
 - Coriolis Force, i.e. the deflection of moving objects due to the rotation of the Earth.

Quasi-Geostrophic Equations

- The QGE are a simplified model for planet-scale flows.
- Streamfunction-Vorticity Formulation

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F \quad (1)$$

$$q = -Ro \Delta \psi + y \quad (2)$$

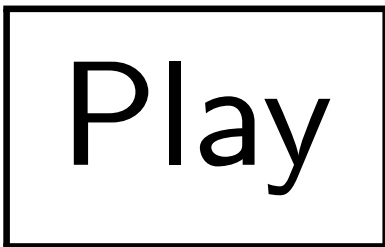
where Ro , Re are the Rossby and Reynolds numbers, and the Jacobian

$$J(\xi, \eta) = \frac{\partial \xi}{\partial x} \frac{\partial \eta}{\partial y} - \frac{\partial \xi}{\partial y} \frac{\partial \eta}{\partial x}$$

- The Jacobian is associated with $(u \cdot \nabla) u$ in the Navier-Stokes Equations

Characteristics of large scale oceanic surface currents

- Driven by forces such as
 - Wind
 - Coriolis Force, i.e. the deflection of moving objects due to the rotation of the Earth.



Demonstration of the Coriolis force.

Quasi-Geostrophic Equations

- The QGE are a simplified model for planet-scale flows.
- Streamfunction-Vorticity Formulation

$\partial \eta$

Characteristics of large scale oceanic surface currents

- Driven by forces such as
 - Wind
 - Coriolis Force, i.e. the deflection of moving objects due to the rotation of the Earth.

Quasi-Geostrophic Equations

- The QGE are a simplified model for planet-scale flows.
- Streamfunction-Vorticity Formulation

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F \quad (1)$$

$$q = -Ro \Delta \psi + y \quad (2)$$

where Ro , Re are the Rossby and Reynolds numbers, and the Jacobian

$$J(\xi, \eta) = \frac{\partial \xi}{\partial x} \frac{\partial \eta}{\partial y} - \frac{\partial \xi}{\partial y} \frac{\partial \eta}{\partial x}$$

- The Jacobian is associated with $(u \cdot \nabla) u$ in the Navier-Stokes Equations

Characteristics of large scale oceanic surface currents

- Driven by forces such as
 - Wind
 - Coriolis Force, i.e. the deflection of moving objects due to the rotation of the Earth.

Quasi-Geostrophic Equations

- The QGE are a simplified model for planet-scale flows.
- Streamfunction-Vorticity Formulation

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F \quad (1)$$

$$q = -Ro \Delta \psi + y \quad (2)$$

where Ro , Re are the Rossby and Reynolds numbers, and the Jacobian

$$J(\xi, \eta) = \frac{\partial \xi}{\partial x} \frac{\partial \eta}{\partial y} - \frac{\partial \xi}{\partial y} \frac{\partial \eta}{\partial x}$$

- The Jacobian is associated with $(u \cdot \nabla) u$ in the Navier-Stokes Equations

Characteristics of large scale oceanic surface currents

- Driven by forces such as
 - Wind
 - Coriolis Force, i.e. the deflection of moving objects due to the rotation of the Earth.

Quasi-Geostrophic Equations

- The QGE are a simplified model for planet-scale flows.
- Streamfunction-Vorticity Formulation

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F \quad (1)$$

$$q = -Ro \Delta \psi + y \quad (2)$$

where Ro , Re are the Rossby and Reynolds numbers, and the Jacobian

$$J(\xi, \eta) = \frac{\partial \xi}{\partial x} \frac{\partial \eta}{\partial y} - \frac{\partial \xi}{\partial y} \frac{\partial \eta}{\partial x}$$

- The Jacobian is associated with $(u \cdot \nabla) u$ in the Navier-Stokes Equations

Further simplification by substituting (2) into the (1).

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F$$

$$q = -Ro \Delta \psi + y$$

Pure Streamfunction Formulation

$$-\frac{\partial [\Delta \psi]}{\partial t} + Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} = Ro^{-1} F \quad (3)$$

Pro	Con
One Variable "Simpler"	Fourth-Order PDE Requires C^1 FE to be Conforming

Further simplification by substituting (2) into the (1).

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F$$

$$q = -Ro \Delta \psi + y$$

Pure Streamfunction Formulation

$$-\frac{\partial [\Delta \psi]}{\partial t} + Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} = Ro^{-1} F \quad (3)$$

Pro	Con
One Variable "Simpler"	Fourth-Order PDE Requires C^1 FE to be Conforming

Further simplification by substituting (2) into the (1).

$$\frac{\partial q}{\partial t} + J(\psi, q) = -Re^{-1} \Delta q + F$$

$$q = -Ro \Delta \psi + y$$

Pure Streamfunction Formulation

$$-\frac{\partial [\Delta \psi]}{\partial t} + Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} = Ro^{-1} F \quad (3)$$

Pro	Con
One Variable "Simpler"	Fourth-Order PDE Requires C^1 FE to be Conforming

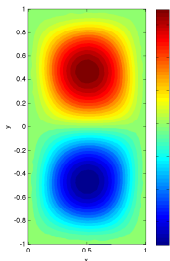
QGE is **not** Navier-Stokes

NSE

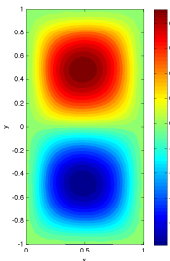
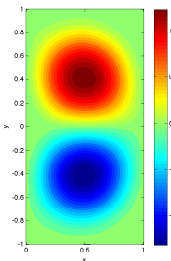
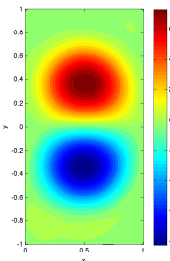
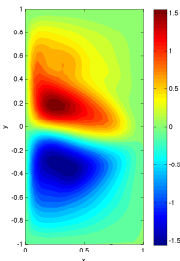
$$-\frac{\partial[\Delta\psi]}{\partial t} + Re^{-1} \Delta^2 \psi + J(\psi, \Delta\psi) = F,$$

QGE

$$-\frac{\partial[\Delta\psi]}{\partial t} + Re^{-1} \Delta^2 \psi + J(\psi, \Delta\psi) - Ro^{-1} \frac{\partial\psi}{\partial x} = Ro^{-1} F$$



NSE

 $Ro = 1$  $Ro = 0.1$  $Ro = 0.01$  $Ro = 0.001$

Time Averaged, $t = [0, 10]$, $dt = 1 \times 10^{-3}$, $Re = 200$, $F = \sin \pi y$

Stationary Quasi-Geostrophic Equations

- Pure Streamfunction Form of Stationary QGE

$$\begin{aligned} Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} &= Ro^{-1} F \\ \psi &= 0, \quad \frac{\partial \psi}{\partial \vec{n}} = 0 \text{ on } \partial \Omega \end{aligned} \tag{4}$$

- Weak Form

$$\begin{aligned} &\text{Find } \psi \in X \text{ such that } \forall \chi \in X := H_0^2(\Omega) \\ Re^{-1}(\Delta \psi, \Delta \chi) + b(\psi; \psi, \chi) - Ro^{-1}(\psi_x, \chi) &= Ro^{-1}(F, \chi), \end{aligned} \tag{5}$$

where

$$b(\psi; \psi, \chi) = [(\Delta \psi \cdot \psi_y, \chi_x) - (\Delta \psi \cdot \psi_x, \chi_y)]$$

- Conforming Finite Element Formulation

$$\begin{aligned} &\text{Find } \psi^h \in X^h \text{ such that } \forall \chi^h \in X^h \subset X \\ Re^{-1}(\Delta \psi^h, \Delta \chi^h) + b(\psi^h; \psi^h, \chi^h) - Ro^{-1}(\psi_x^h, \chi^h) &= (F, \chi^h) \end{aligned} \tag{6}$$

Stationary Quasi-Geostrophic Equations

- Pure Streamfunction Form of Stationary QGE

$$\begin{aligned} Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} &= Ro^{-1} F \\ \psi &= 0, \quad \frac{\partial \psi}{\partial \vec{n}} = 0 \text{ on } \partial \Omega \end{aligned} \quad (4)$$

- Weak Form

$$\begin{aligned} &\text{Find } \psi \in X \text{ such that } \forall \chi \in X := H_0^2(\Omega) \\ Re^{-1}(\Delta \psi, \Delta \chi) + b(\psi; \psi, \chi) - Ro^{-1}(\psi_x, \chi) &= Ro^{-1}(F, \chi), \end{aligned} \quad (5)$$

where

$$b(\psi; \psi, \chi) = [(\Delta \psi \cdot \psi_y, \chi_x) - (\Delta \psi \cdot \psi_x, \chi_y)]$$

- Conforming Finite Element Formulation

$$\begin{aligned} &\text{Find } \psi^h \in X^h \text{ such that } \forall \chi^h \in X^h \subset X \\ Re^{-1}(\Delta \psi^h, \Delta \chi^h) + b(\psi^h; \psi^h, \chi^h) - Ro^{-1}(\psi_x^h, \chi^h) &= (F, \chi^h) \end{aligned} \quad (6)$$

Stationary Quasi-Geostrophic Equations

- Pure Streamfunction Form of Stationary QGE

$$\begin{aligned} Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} &= Ro^{-1} F \\ \psi &= 0, \quad \frac{\partial \psi}{\partial \vec{n}} = 0 \text{ on } \partial \Omega \end{aligned} \quad (4)$$

- Weak Form

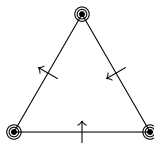
$$\begin{aligned} &\text{Find } \psi \in X \text{ such that } \forall \chi \in X := H_0^2(\Omega) \\ Re^{-1}(\Delta \psi, \Delta \chi) + b(\psi; \psi, \chi) - Ro^{-1}(\psi_x, \chi) &= Ro^{-1}(F, \chi), \end{aligned} \quad (5)$$

where

$$b(\psi; \psi, \chi) = [(\Delta \psi \cdot \psi_y, \chi_x) - (\Delta \psi \cdot \psi_x, \chi_y)]$$

- Conforming Finite Element Formulation

$$\begin{aligned} &\text{Find } \psi^h \in X^h \text{ such that } \forall \chi^h \in X^h \subset X \\ Re^{-1}(\Delta \psi^h, \Delta \chi^h) + b(\psi^h; \psi^h, \chi^h) - Ro^{-1}(\psi_x^h, \chi^h) &= (F, \chi^h) \end{aligned} \quad (6)$$



Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

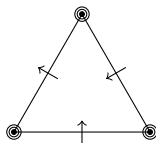
$$\|u - u_h\|_{H^2(\Omega)} \leq C \|u\|_{H^{2+\alpha}(\Omega)} h^{\alpha} \quad \text{for } \alpha = 0, 1, 2$$

11 degrees of freedom

11th-order functions at each node (5 values, 6 DOF)

- First derivatives unique at each corner (3 values total)
- Second derivatives unique at each corner (6 values total)
- Normal derivatives unique at the midpoints (3 values total)

• C^1 continuity is achieved by the 11th-order functions

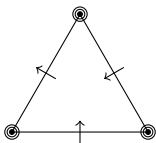


Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)
- <https://github.com/VT-ICAM/ArgyrisPack>

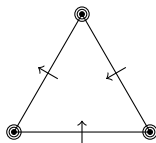


Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)
- <https://github.com/VT-ICAM/ArgyrisPack>

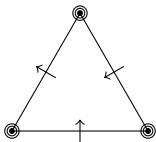


Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)
- <https://github.com/VT-ICAM/ArgyrisPack>

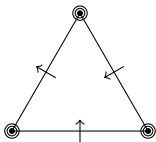


Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)
- <https://github.com/VT-ICAM/ArgyrisPack>

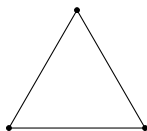


Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)
- <https://github.com/VT-ICAM/ArgyrisPack>



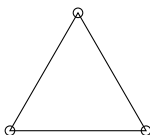
Node Values of Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)

• <https://github.com/VT-ICAM/ArgyrisPack>



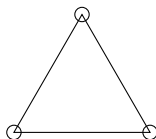
First Derivative Values of Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)

• <https://github.com/VT-ICAM/ArgyrisPack>



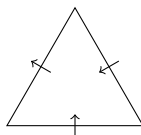
Second Derivative Values of Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)

- <https://github.com/VT-ICAM/ArgyrisPack>



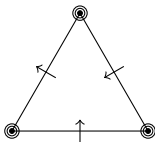
Second Derivative Values of Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)

• <https://github.com/VT-ICAM/ArgyrisPack>



Argyris Triangle

- The Argyris Finite Element is C^1
 - Conforming Finite Element for fourth-order problems
- Fifth-order basis functions
 - Interpolation Error Bounds for Argyris

$$\|u - \Pi_h u\|_{2-s} \leq C h^{4+s} |u|_6 \text{ for } s = 0, 1, 2 \quad (7)$$

- 21 degrees of freedom
 - Function values at each vertex (3 values total)
 - First derivative values at each vertex (6 values total)
 - Second derivative values at each vertex (9 values total)
 - Normal derivative values at the midpoints (3 values total)
- <https://github.com/VT-ICAM/ArgyrisPack>

Theorem (Argyris Error Estimates)

Let ψ be the solution of (5) and ψ^h be the solution of (6) and assume the small data condition

$$Re^{-2} Ro \geq \Gamma_1 \|F\|_{-2}.$$

Furthermore, assume that $\psi \in H^6(\Omega) \cap H_0^2(\Omega)$. Then there exists positive constants C_0 , C_1 , and C_2 that depend on $Re, Ro, \Gamma_1, \Gamma_2, F$ but not h such that

$$|\psi - \psi^h|_2 \leq C_2 \cdot h^4 \quad (8)$$

$$|\psi - \psi^h|_1 \leq C_1 \cdot h^5 \quad (9)$$

$$\|\psi - \psi^h\|_0 \leq C_0 \cdot h^6 \quad (10)$$

E. L. Foster, T. Iliescu, and Z. Wang, Submitted 2012, <http://arxiv.org/abs/1210.3630>

For the H^2 error estimate the proof follows in the standard way. To determine the H^1 and L^2 error estimates one uses the duality argument (Aubin-Nitsche) to bootstrap from H^2 norm into the appropriate norm.

Numerical Test

$$Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} = Ro^{-1} F$$

$$\psi = 0, \quad \frac{\partial \psi}{\partial \vec{n}} = 0 \text{ on } \partial\Omega, \quad Re = Ro = 1$$

- $\Omega = [0, 1] \times [0, 1]$
- $\psi(x, y) = \sin^2 \pi x \cdot \sin^2 2\pi y$

Numerical Test

$$Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} = Ro^{-1} F$$

$$\psi = 0, \quad \frac{\partial \psi}{\partial \vec{n}} = 0 \text{ on } \partial\Omega, \quad Re = Ro = 1$$

- $\Omega = [0, 1] \times [0, 1]$
- $\psi(x, y) = \sin^2 \pi x \cdot \sin^2 2\pi y$

Numerical Test

$$Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} = Ro^{-1} F$$

$$\psi = 0, \quad \frac{\partial \psi}{\partial \vec{n}} = 0 \text{ on } \partial\Omega, \quad Re = Ro = 1$$

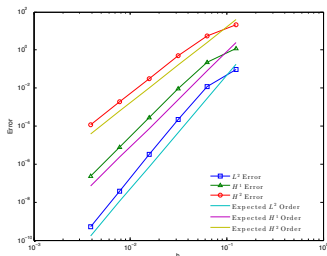
- $\Omega = [0, 1] \times [0, 1]$
- $\psi(x, y) = \sin^2 \pi x \cdot \sin^2 2\pi y$

Numerical Test

$$Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} = Ro^{-1} F$$

$$\psi = 0, \quad \frac{\partial \psi}{\partial \vec{n}} = 0 \text{ on } \partial \Omega, \quad Re = Ro = 1$$

- $\Omega = [0, 1] \times [0, 1]$
- $\psi(x, y) = \sin^2 \pi x \cdot \sin^2 2\pi y$



Numerical vs Theoretical Rates of Convergence

Time Dependence

- Recall the pure streamfunction form of QGE, i.e. Equation 3

$$\begin{aligned}
 -\frac{\partial [\Delta \psi]}{\partial t} + Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} &= Ro^{-1} F \\
 \psi(t; x, y) = \frac{\partial \psi}{\partial \vec{n}} &= 0 \text{ on } \partial \Omega, \quad \psi(0; x, y) = \psi_0(x, y)
 \end{aligned}$$

- Full-discretization

$$\begin{aligned}
 [K + k(L - B)] \psi_{n+1}^h + b(\psi_{n+1}^h; \psi_{n+1}^h, \chi^h) \\
 = K \psi_n^h + \ell(t_{n+1}), \quad \forall \chi^h \in X^h \subset X
 \end{aligned} \tag{11}$$

where

- K is the stiffness matrix associated with $(\nabla \psi^h, \nabla \chi^h)$
- L is the Laplace matrix associated with $Re^{-1}(\Delta \psi^h, \chi^h)$
- B is the β -plane matrix associated with $Ro^{-1}(\psi_x^h, \chi^h)$
- $\ell(t_{n+1})$ is the load vector associated with $Ro^{-1}(F(t_{n+1}), \chi^h)$
- k is the time step

Time Dependence

- Recall the pure streamfunction form of QGE, i.e. Equation 3

$$\begin{aligned}
 -\frac{\partial [\Delta \psi]}{\partial t} + Re^{-1} \Delta^2 \psi + J(\psi, \Delta \psi) - Ro^{-1} \frac{\partial \psi}{\partial x} &= Ro^{-1} F \\
 \psi(t; x, y) = \frac{\partial \psi}{\partial \vec{n}} = 0 \text{ on } \partial \Omega, \quad \psi(0; x, y) &= \psi_0(x, y)
 \end{aligned}$$

- Full-discretization

$$\begin{aligned}
 [K + k(L - B)] \psi_{n+1}^h + b(\psi_{n+1}^h; \psi_{n+1}^h, \chi^h) \\
 = K \psi_n^h + \ell(t_{n+1}), \quad \forall \chi^h \in X^h \subset X
 \end{aligned} \tag{11}$$

where

- K is the stiffness matrix associated with $(\nabla \psi^h, \nabla \chi^h)$
- L is the Laplace matrix associated with $Re^{-1}(\Delta \psi^h, \chi^h)$
- B is the β -plane matrix associated with $Ro^{-1}(\psi_x^h, \chi^h)$
- $\ell(t_{n+1})$ is the load vector associated with $Ro^{-1}(F(t_{n+1}), \chi^h)$
- k is the time step

Numerical Test

$$-\frac{\partial [\Delta\psi]}{\partial t} + Re^{-1}\Delta^2\psi + J(\psi, \Delta\psi) - Ro^{-1}\frac{\partial\psi}{\partial x} = Ro^{-1}F$$

$$\psi = 0, \quad \frac{\partial\psi}{\partial\vec{n}} = 0 \text{ on } \partial\Omega, \quad \psi(0; x, y) = \psi_0(x, y)$$

$$Re = Ro = 1$$

- $\Omega = [0, 1] \times [0, 1], t = [0, 0.25]$
- $\psi(t; x, y) = [(1 - x) (1 - e^{-0.1 x t}) \sin \pi y]^2$

Numerical Test

$$-\frac{\partial [\Delta\psi]}{\partial t} + Re^{-1}\Delta^2\psi + J(\psi, \Delta\psi) - Ro^{-1}\frac{\partial\psi}{\partial x} = Ro^{-1}F$$

$$\psi = 0, \quad \frac{\partial\psi}{\partial\vec{n}} = 0 \text{ on } \partial\Omega, \quad \psi(0; x, y) = \psi_0(x, y)$$

$$Re = Ro = 1$$

- $\Omega = [0, 1] \times [0, 1], t = [0, 0.25]$
- $\psi(t; x, y) = [(1-x)(1 - e^{-0.1xt})\sin\pi y]^2$

Numerical Test

$$-\frac{\partial [\Delta\psi]}{\partial t} + Re^{-1}\Delta^2\psi + J(\psi, \Delta\psi) - Ro^{-1}\frac{\partial\psi}{\partial x} = Ro^{-1}F$$

$$\psi = 0, \quad \frac{\partial\psi}{\partial\vec{n}} = 0 \text{ on } \partial\Omega, \quad \psi(0; x, y) = \psi_0(x, y)$$

$$Re = Ro = 1$$

- $\Omega = [0, 1] \times [0, 1], t = [0, 0.25]$
- $\psi(t; x, y) = [(1 - x) (1 - e^{-0.1 x t}) \sin \pi y]^2$

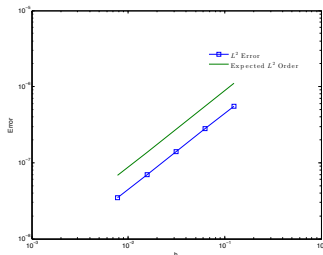
Numerical Test

$$-\frac{\partial [\Delta\psi]}{\partial t} + Re^{-1}\Delta^2\psi + J(\psi, \Delta\psi) - Ro^{-1}\frac{\partial\psi}{\partial x} = Ro^{-1}F$$

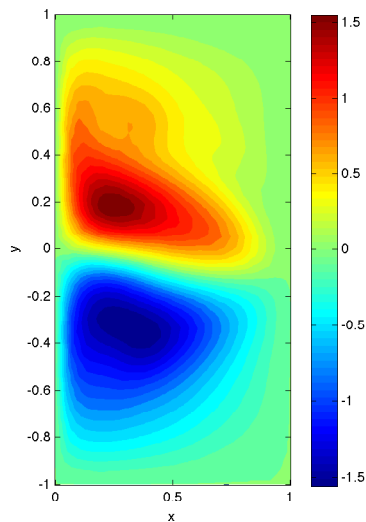
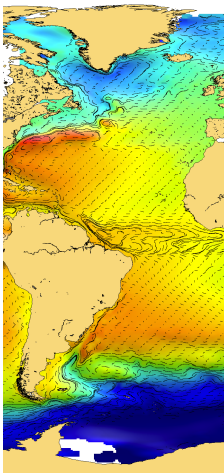
$$\psi = 0, \quad \frac{\partial\psi}{\partial\vec{n}} = 0 \text{ on } \partial\Omega, \quad \psi(0; x, y) = \psi_0(x, y)$$

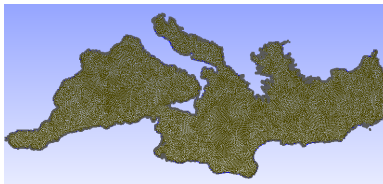
$$Re = Ro = 1$$

- $\Omega = [0, 1] \times [0, 1], t = [0, 0.25]$
- $\psi(t; x, y) = [(1 - x)(1 - e^{-0.1xt}) \sin \pi y]^2$

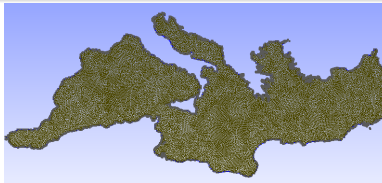


Notice anything different?





FE mesh for the Mediterranean Sea



FE mesh for the Mediterranean Sea

Play

QGE on the Mediteranean Sea

$$F = \sin(\pi y), Re = 400, Ro = 1.6E - 4, DoFs = 46480$$

Challenges and Future Work

- Realistic Parameters

Challenge	Possible Solution
Narrow Boundary Layer	Stabilization Methods
Dynamic Structures	Adaptive Mesh Refinement

- Ensemble Forecasting

Challenge	Possible Solutions
"Slow" Code	Parallel Processing Adaptive Mesh Refinement Proper Orthogonal Decomposition Large Eddy Simulation

Challenges and Future Work

- Realistic Parameters

Challenge	Possible Solution
Narrow Boundary Layer	Stabilization Methods
Dynamic Structures	Adaptive Mesh Refinement

- Ensemble Forecasting

Challenge	Possible Solutions
"Slow" Code	Parallel Processing Adaptive Mesh Refinement Proper Orthogonal Decomposition Large Eddy Simulation

Challenges and Future Work

- Realistic Parameters

Challenge	Possible Solution
Narrow Boundary Layer Dynamic Structures	Stabilization Methods Adaptive Mesh Refinement

- Ensemble Forecasting

Challenge	Possible Solutions
"Slow" Code	Parallel Processing Adaptive Mesh Refinement Proper Orthogonal Decomposition Large Eddy Simulation

Challenges and Future Work

- Realistic Parameters

Challenge	Possible Solution
Narrow Boundary Layer Dynamic Structures	Stabilization Methods Adaptive Mesh Refinement

- Ensemble Forecasting

Challenge	Possible Solutions
"Slow" Code	Parallel Processing Adaptive Mesh Refinement Proper Orthogonal Decomposition Large Eddy Simulation

Challenges and Future Work

- Realistic Parameters

Challenge	Possible Solution
Narrow Boundary Layer Dynamic Structures	Stabilization Methods Adaptive Mesh Refinement

- Ensemble Forecasting

Challenge	Possible Solutions
"Slow" Code	Parallel Processing Adaptive Mesh Refinement Proper Orthogonal Decomposition Large Eddy Simulation

Challenges and Future Work

- Realistic Parameters

Challenge	Possible Solution
Narrow Boundary Layer Dynamic Structures	Stabilization Methods Adaptive Mesh Refinement

- Ensemble Forecasting

Challenge	Possible Solutions
"Slow" Code	Parallel Processing Adaptive Mesh Refinement Proper Orthogonal Decomposition Large Eddy Simulation

Summary

- Large Scale Ocean Currents
- Quasi-Geostrophic Equation
- Pure Stream Function form of QGE
- Optimal Error estimates
- Numerical Experiments
- Preliminary and Future Work

We would like to acknowledge the NSF and ICTAS for their generous sponsorship of this work.

We would also like to acknowledge GMSH for their excellent meshing software.

Summary

- Large Scale Ocean Currents
- Quasi-Geostrophic Equation
- Pure Stream Function form of QGE
- Optimal Error estimates
- Numerical Experiments
- Preliminary and Future Work

We would like to acknowledge the NSF and ICTAS for their generous sponsorship of this work.

We would also like to acknowledge GMSH for their excellent meshing software.

Summary

- Large Scale Ocean Currents
- Quasi-Geostrophic Equation
- Pure Stream Function form of QGE
- Optimal Error estimates
- Numerical Experiments
- Preliminary and Future Work

We would like to acknowledge the NSF and ICTAS for their generous sponsorship of this work.

We would also like to acknowledge GMSH for their excellent meshing software.

Summary

- Large Scale Ocean Currents
- Quasi-Geostrophic Equation
- Pure Stream Function form of QGE
- Optimal Error estimates
- Numerical Experiments
- Preliminary and Future Work

We would like to acknowledge the NSF and ICTAS for their generous sponsorship of this work.

We would also like to acknowledge GMSH for their excellent meshing software.

Summary

- Large Scale Ocean Currents
- Quasi-Geostrophic Equation
- Pure Stream Function form of QGE
- Optimal Error estimates
- Numerical Experiments
- Preliminary and Future Work

We would like to acknowledge the NSF and ICTAS for their generous sponsorship of this work.

We would also like to acknowledge GMSH for their excellent meshing software.

Summary

- Large Scale Ocean Currents
- Quasi-Geostrophic Equation
- Pure Stream Function form of QGE
- Optimal Error estimates
- Numerical Experiments
- Preliminary and Future Work

We would like to acknowledge the NSF and ICTAS for their generous sponsorship of this work.

We would also like to acknowledge GMSH for their excellent meshing software.

Summary

- Large Scale Ocean Currents
- Quasi-Geostrophic Equation
- Pure Stream Function form of QGE
- Optimal Error estimates
- Numerical Experiments
- Preliminary and Future Work

We would like to acknowledge the NSF and ICTAS for their generous sponsorship of this work.

We would also like to acknowledge GMSH for their excellent meshing software.

Summary

- Large Scale Ocean Currents
- Quasi-Geostrophic Equation
- Pure Stream Function form of QGE
- Optimal Error estimates
- Numerical Experiments
- Preliminary and Future Work

We would like to acknowledge the NSF and ICTAS for their generous sponsorship of this work.

We would also like to acknowledge GMSH for their excellent meshing software.

Thank you!

References

- J. H. Argyris, I. Fried, and D. W. Scharpf, 1968.
- P. G. Ciarlet, 2002.
- M. E. Cayco and R. A. Nicolaides, 1986.
- C. Johnson, 2009.
- V. Dominguez and F. J. Sayas, 2006.
- F. Fairag, 1998.
- F. Fairag and N. Almulla, 2004.
- E. L. Foster, T. Iliescu, and Z. Wang, 2012.
- E. L. Foster, T. Iliescu, and D. R. Wells, 2012.
- V. Girault and P. A. Raviart, 1979.
- M. D. Gunzburger, 1989.
- W. Layton, 1993.
- A. Logg, K-A Mardal, and G. Wells, 2012.
- A. J. Majda and X. Wang, 2006.
- G. K. Vallis, 2007.
- J. Xu, 1994.