

Methods for Producing a Reliable APWP

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SUMMARY

First, new picking/weighting methods developed here and previously published picking/weighting methods are compiled together to generate 168 different paleomagnetic APWPs. Then, the APWP similarity measuring tool is used to find which methods is(are) good or bad. The final results tell us that the “Age Position Picking (APP)” method is better than the “Age Mean Picking (AMP)” method for making a reliable paleomagnetic APWP and weighting is actually unnecessary.

Key words: Moving Average – Weighting – APWP – Paleomagnetism.

1 INTRODUCTION

APWPs are generated by combining paleomagnetic poles, also known as paleopoles, for a particular rigid block over the desired age range to produce a smoothed path. See the Supplementary Material for some examples how the paleopole datasets are constrained for a particular tectonic plate during a specific time interval.

1.1 Not All Data Are Created Equal

However, uncertainties in the age and location of paleopoles can vary greatly for different poles.

1.1.1 Age Error

Although remanent magnetizations are generally assumed to be primary, many events can cause remagnetisation (in which case the derived pole is ‘younger’ than the rock). If an event that has occurred since the rock’s formation that should affect the magnetisation (e.g., folding, thermal overprinting due to intrusion) can be shown to have affected it, then it constrains the magnetisation to have been acquired before that event. Recognising or ruling out remagnetisations depends on these field tests, which are not always performed or possible. Even a passed field test may not be useful if field test shows magnetisation acquired prior to a folding event tens of millions of years after initial rock formation.

The most obvious characteristic we can observe from paleomagnetic data is that some poles have very large age ranges, e.g., more than 100 Myr. The magnetization age should be some time between the information of the rock and folding events. There are also others where we have similar position but the age constraint is much narrower, e.g. 10 Myr window or less. Obviously the latter kind of data is more valuable than the one with large age range.

1.1.2 Position Error

The errors of pole latitudes and longitudes are 95% confidence ellipses, which also vary greatly in magnitude. All paleopoles have some associated uncertainties due to measurement error and the nature of the geomagnetic field. More uncertainties can be added by too few samples, sampling spanning too short a time range to approximate a GAD field, failure to remove overprints during demagnetisation, etc.

1.1.3 Data Consistency

Paleopoles of a rigid plate or block should be continuous time series. For a rigid plate, two poles with similar ages shouldn’t be dramatically different in location. Sometimes, this is the case. Sometimes we have further separated poles with close ages.

There are a number of possible causes for these outliers, including:

Lithology

For poor consistency of data, it is potentially because of different inclinations or declinations. The first thing we should consider about is their lithology. We want to check if the sample rock are igneous or sedimentary, because sediment compaction can result in anomalously shallow inclinations (Tauxe et al. 2018). In addition, we also can check if the rock are redbeds or non-redbeds. Although whether redbeds record a detrital signal or a later Chemical Remanent Magnetization (CRM) is still somewhat controversial, both sedimentary rocks and redbeds could lead to inconsistency in direction compared to igneous rocks.

Local Rotations

Local deformation between two paleomagnetic localities invalidates the rigid plate assumption and could lead to inconsistent VGP directions. So if discordance is due to local deformation, and we would ideally want to exclude such poles from our APWP calculation.

Other Factors

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In most cases, mean pole age (centre of age error) has just been binned. If any of the poles have large age errors, they could be different ages from each other and sample entirely different parts of the APWP. Conversely, if any of the poles have too few samples, or were not sampled over enough time to average to a GAD field, a discordant pole may be due to unreduced secular variation.

1.1.4 Data Density

As we go back in time, we have lower quality and lower density (or quantity) of data, for example, the Precambrian or Early Paleozoic paleomagnetic data are relatively fewer than Middle-Late Phanerozoic ones, and most of them are not high-quality, e.g., larger errors in both age and location. The combination of lower data quality with lower data density means that a single ‘bad’ pole (with large errors in age and/or location) can much more easily distort the reconstructed APWP, because there are few or no ‘good’ poles to counteract its influence.

Data density also varies between different plates. E.g., we have a relatively high density of paleomagnetic data for North American Craton (NAC), but few poles exist for Greenland and Arabia. Based on mean age (mean of lower and upper magnetic ages), for 120–0 Ma, the **Global Paleomagnetic Database** (GPMDB) version 4.6b (McElhinny & Lock 1996; Pisarevsky 2005, updated in 2016 by the Ivar Giaever Geomagnetic Laboratory team, in collaboration with Pisarevsky) has more than 130 poles for NAC, but only 17 for Greenland and 24 for Arabia.

1.1.5 Publication Year

The time when the data was published should also be considered, because magnetism measuring methodology, technology and equipments have been improved since the early 20th century. For example, stepwise demagnetisation, which is the most reliable method of detecting and removing secondary overprints, has only been in common use since the mid 1980s.

In summary, not all paleopoles are created equal, which leads to an important question: how to best combine poles of varying quality into a coherent and accurate APWP?

1.2 Existing Solutions and General Issues

Paleomagnetists have proposed a variety of methods to filter so-called “bad” data, or give lower weights to those “bad” data before generating an APWP, e.g., two widely used methods: the V90 reliability criteria (van der Voo 1990) and the BC02 selection criteria provided by Besse & Courtillot (2002). Briefly, the V90 criteria for paleomagnetic results includes seven criteria: (1) Well determined age; (2) At least 25 samples with Fisher (Fisher 1953) precision κ greater than 10 and α_{95} less than 16° ; (3) Detailed demagnetisation results reported; (4) Passed field tests; (5) Tectonic coherence with continent and good structural control; (6) Identified antipodal reversals; (7) Lack of similarity with younger poles (Torsvik et al. 1992). The total criteria satisfied (0–7) is then used as a measure of a paleomagnetic result’s overall reliability, which is known as Q (quality) factor (Torsvik et al. 1992). Q factor is indeed a very straightforward way to get a quantified reliability score. Also it can be conveniently used in the later calculations of APWPs (Torsvik et al. 1992). But at the same time this is a fairly basic filter that lumps together criteria that may not be equally important. Compared with V90, the BC02 criteria suggests stricter filtering, e.g., using only

poles with at least 6 sampling sites and 36 samples, each site having α_{95} less than 10° in the Cenozoic and 15° in the Mesozoic. BC02 is also straightforward and convenient to use, but some useful data may be filtered out and wasted especially for a period where there are only limited number of data. In addition, there has been limited study of how effective these marking/filtering methods are at reconstructing a ‘true’ APWP, and for most studies after a basic filtering of ‘low quality’ poles, the remaining poles are, in fact, treated equally.

Above all, there haven’t been any real attempts to study how APWP fits may be improved by filtering/weighting data. This paper is presented to address these issues.

2 METHODS

2.1 General Approach

In this study, we use paleopoles extracted from the GPMDB to generate APWPs for the period 120–0 Ma. A range of possible APWP paths for North America, India and Australia can be generated from the extracted sets of paleopoles using various binning, filtering and weighting methods (Tables 1 and 2). These paths can then be compared to synthetic APWPs independently generated from an absolute plate motion model. The three plates chosen have different attributes, both in terms of the input data set and the nature of the reference APWP.

2.2 Paleomagnetic Data

Collections of poles with a minimum age (LOMAGAGE) ≤ 135 Ma for the North American (Plate ID 101), the Indian (ID 501) and Australian (ID 801) plates, were extracted from the GPMDB. In order to include valid paleopoles from blocks that moved independently prior to 120 Ma, which therefore have different assigned plate codes in the GPMDB, the spatial join technique (Jacox & Samet 2007) was used to find all poles within the geographic region that defines the rigid plate within the period of interest (see also Appendix A):

for North America, the search region was defined by the North American (ID 101), Avalon/Acadia (ID 108) and Piedmont (ID 109) blocks, as defined by the recently published plate model of (Young et al. 2018). Following extraction, 58 poles from southwestern North America that have been affected by regional rotations since 36 Ma (McQuarrie & Wernicke 2006), were removed. The final dataset consists of 135 paleopoles (Fig. 1), with 76 (about 56.3%) sampled from dominantly igneous sequences; 56 (about 41.5%) sampled from mostly sedimentary sequences, including 6 from redbeds; and 3 (about 2.2%) from metamorphic sequences. The principal features of the age distribution are a larger number of young (< 5 Ma) poles, and relatively fewer poles in the Late Cretaceous and Miocene.

for India, the Indian block (ID 501) as defined by (Young et al. 2018) was used, but following extraction 31 paleopoles associated with parts of the northern margin that have undergone regional rotations since the Jurassic (Gaina et al. 2015) were removed. The final dataset consists of 75 paleopoles (Fig. 2), with 39 (52%) sampled from dominantly igneous sequences and 36 (48%) sampled from mostly sedimentary sequences, including 3 from

redbeds). There is a high concentration of poles from the latest Cretaceous–Early Cenozoic (c. 70–60 Ma), many of which are igneous; in younger and older intervals, there are fewer, mostly sedimentary poles.

for Australia, the Australia (ID 801), Sumba (ID 675), and Timor (ID 684) blocks as defined by (Young et al. 2018) were used, in combination with data from the Tasmania block (ID 805) younger than c. 100 Ma (with a maximum age (HIMAGAGE) ≤ 100 Ma), prior to which it was not fixed with respect to Australia (Young et al. 2018). The final dataset consists of 99 paleopoles (Fig. 3), with 61 (61.6%) sampled from dominantly igneous sequences and 38 (38.4%) sampled from mostly sedimentary sequences, including 9 from redbeds). The temporal distribution of poles is relatively uniform.

Compared with North American (Fig. 1) and Australian (Fig. 3) paleopoles, Indian paleopoles are in relatively lower density in general except during the period of about 70–60 Ma (Fig. 2).

2.3 APWP Generation

Multiple APWPs were generated using the selected poles for each of the three plates as follows:

Picking/binning. A moving average or moving window technique was used: paleopoles were selected for each APWP time step (initially 5 Myr step length from 0 to 120 Ma) if their age fell within a window centered on the current step age. In this study, the width of the moving window was always twice that of time step (i.e. initially 10 Myr), such that each window half-overlaps with its neighbours.

Filtering. Poles with characteristics thought to correspond to poor data quality, or lacking characteristics thought to correspond to good data quality, were discarded (or in some cases, corrected).

Weighting. Calculation of a weighted Fisher mean (Fisher 1953) of the remaining poles within each window, using weighting functions intended to increase the influence of higher quality poles relative to lower quality ones.

28 different picking and filtering algorithms were tested (Table. 1), in combination with 6 different weighting algorithms (Table. 2), for the three plates. The effects of changing the time step length and width of the moving window, and the reference path, were also examined.

2.3.1 Moving Average

The moving average method, also called “running mean” or “moving window” (Torsvik et al. 2008) method, calculates the average of values between a certain data (age in our case) range; the average is then recalculated as the limits of the bin are repeatedly incremented upwards. In addition to this traditional moving windows averaging algorithm, a less restrictive moving average method is also used, referred to here as the “Age Position Picking (APP)” method. The difference of this new moving average method from the traditional one built in GMAP (Torsvik & Smethurst 1999; Torsvik et al. 2008) is that the whole magnetic age range is taken into account in each window, while GMAP only considers the mid-point of the low and high magnetic age of each pole, an algorithm referred to as the “Age Mean Picking (AMP)” method.

Normally each paleopole in the paleomagnetic database is treated as a point with an age that is the mid-point between the upper and lower age limits, i.e. AMP, but this is problematic for paleomagnetic data with large age ranges (especially if they turn out to be primary magnetization that should plot at old end of age range). We are trying a method, APP, that includes a paleo-pole in the moving average bin if any part of its specified age range falls within that bin. If, for example, we have a pole which is constrained to within 10 and 20 Ma of age, and we have a 2 Myr moving window with a 1 Myr age step, then it shouldn't just be in the 14–16 Ma bin (for the mid-point age of 15 Ma)—it should be in the 9–11, 10–12, 11–13, 12–14... 17–19, 18–20, and 19–21 Ma bins. So the average poles are produced from each bin, and each original pole is represented over its entire possible acquisition age. Fig. 4 shows an example of moving average with a 10 Myr window and a 5 Myr step. So, for example, for the window of 15 Ma to 5 Ma (the dashed-line bin in Fig. 4), the AMP method calculates the Fisher mean pole of only 5 poles, while the APP method calculates the mean pole of 9 poles. From comparison of mean poles of the picked poles for the dashed-line age window with the two different algorithms (the 10 Ma mean poles in Fig. 5), the mean pole from the APP method is closer to the 10 Ma pole in the FHM predicted path, which indicates more data diluting the effect of outliers.

In addition, technically we don't want to let the step length more than the window length, in which cases we would lose data between windows.

2.3.1.1 Picking The 28 picking methods include AMP (Picking No 0), APP (Picking No 1) and also those with filtering or corrections implemented onto the two (Table. 1).

Filtering through $\alpha95$ and Age Range For this specific filter (Picking No 2 and 3; Table. 1), the paleopoles are picked out through setting $\alpha95$ of $\leq 15^\circ$ (referring to the BC02's $\alpha95$ filter for the Mesozoic (Besse & Courtillot 2002) and SS05's $\alpha95$ filter (Schettino & Scotese 2005)) and age uncertainty of ≤ 20 Myr (a bit stricter than BC02's ≤ 30 Myr age uncertainty filter (Besse & Courtillot 2002) and SS05's ≤ 40 Myr filter (Schettino & Scotese 2005)). This filter considers both spatial and temporal uncertainties of the paleopoles.

Filtering Out Non-igneous Derived Poles With this filter (Picking No 4 and 5; Table. 1), the paleopoles are mainly (mostly but also containing sedimentary or metamorphic or both) or only from igneous rocks with extrusive or intrusive type. In fact, most paleopoles picked by this filter are derived from igneous-only rocks.

Filtering Out Igneous Unrelated Poles With this filter (Picking No 6 and 7; Table. 1), the poles are from rocks that contain extrusive or intrusive igneous type. In other words, the rock type could be mainly sedimentary or metamorphic.

Inclination Shallowing and Unflattening To test if unflattening possible inclination shallowing in sedimentary rocks can improve the APWP fitting outcomes, the flattening function (King 1955) $\tan I_o = f \tan I_f$ is used to unflatten assumed existing inclination shallowing in sedimentary-based or redbeds-based paleomagnetic data (No 8, 9 and 12, 13 in Table. 1), where I_o is the observed inclination, I_f is the unflattened inclination, and f is the flattening factor (or shallowing coefficient) ranging from unity (no flattening) to 0 (completely flattened). Here $f = 0.6$ is used for



Figure 1. Temporal distribution of 120–0 Ma NAC (101) paleopoles in 10 Myr window length and 5 Myr step length. For distribution a, each bin only counts in the midpoints (circles) of pole error bars (not including those right at bin edges); For distribution b, as long as the bar intersects with the bin (not including those intersecting only at one of bin edges), it is counted in. Inside the parentheses, i means igneous rocks derived (red bars; only two poles, 83–77 Ma and 80–65 Ma, from igneous and also sedimentary; only one pole, 72–40 Ma, from igneous and also metamorphic), r means sedimentary rocks with redbeds involved derived (orange bars), and m means metamorphic rocks derived (blue bars); the left are non-redbed sedimentary rocks derived (black bars; only two poles, 146–65 Ma [RESULTNO 6679] and 2–0 Ma [RESULTNO 1227], are from sedimentary and also metamorphic). The data published before 1984 are shown as circles with a dot.



Figure 2. Temporal distribution of 120–0 Ma Indian (501) paleopoles. For red bars, only one pole, 67–64 Ma (RESULTNO 8593), is from igneous and also sedimentary. See Fig. 1 for more information.



Figure 3. Temporal distribution of 120–0 Ma Australian (801) paleopoles. For black bars, only four poles, 100–80 Ma (RESULTNO 1106), 10–2 Ma (RESULTNO 1208), 4–2 Ma (RESULTNO 140) and 1–0 Ma (RESULTNO 1963), are from sedimentary and also igneous. For red bars, only one pole, 65–25 Ma (RESULTNO 1872), is from igneous and also sedimentary rocks, and only one pole, 1–0 Ma (RESULTNO 1147), is from igneous and also metamorphic rocks. See Fig. 1 for more information.

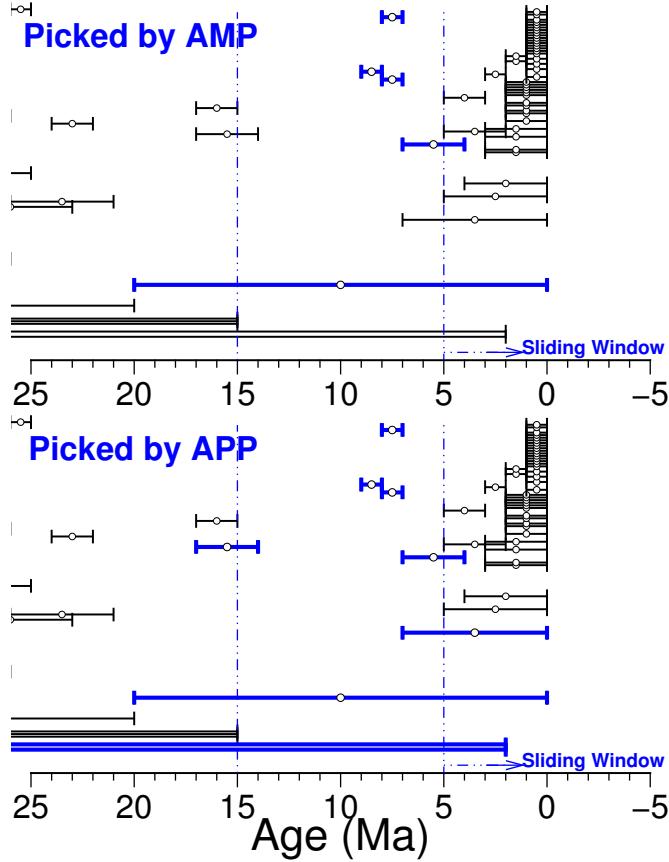


Figure 4. An example of 10 Myr moving window and 5 Myr step in the two moving average methods, AMP and APP, based on poles of the NAC. Every age window has a different color. White points are the midpoints of low and high magnetic ages. The vertical axis has no specific meaning here.

the sedimentary-dominated rocks derived paleopoles in our calculations, according to the previous experience (Torsvik et al. 2012). And $f = 0.8$ (the minimum anisotropy-of-thermal-remanence-determined f -correction) (Domeier et al. 2011a; Domeier et al. 2011b) is used for the sedimentary-contained (not dominated) rocks derived paleopoles in the calculations.

Filtering Out Poles Related to Red-beds Bias toward shallow inclinations is also observed in paleomagnetic data derived from red-beds (Tauxe & Kent 2004). The question raised here is: are the red-beds the main component that contributes the inclination shallowing? In this filter (Picking No 10–11 in Table 1), we want to simply remove all red-beds related paleopoles to see if there is any improvement from matching the hotspot model predicted path or if there is any difference from the igneous derived results (Picking No 4–7 in Table 1).

Filtering Out Poles Published Earlier or Later It is also worthy to see if recently published data is able to produce a more reliable APWP than relatively older data. Here 1983 is chosen as the division (Picking No 14, 15 and 16, 17; Table 1), because the mean of all the data publication years in GPMDB 4.6b is about 1983.

Filtering Out Poles Influenced by Local Rotations or Secondary Print Some publications of paleomagnetic data suggest the data has probably been affected by local rotations or secondary

Table 1. List of all Picking (i.e. Binning) algorithms developed in this paper. AMP, Age Mean Picking (See Section ‘Moving Average’); APP, Age Position Picking.

No.	Picking Algorithm
0	AMP
1	APP
2	AMP (“ $\alpha_{95}/\text{Age range}$ ” no more than “15/20”)
3	APP (“ $\alpha_{95}/\text{Age range}$ ” no more than “15/20”)
4	AMP (mainly or only igneous)
5	APP (mainly or only igneous)
6	AMP (contain igneous and not necessarily mainly)
7	APP (contain igneous and not necessarily mainly)
8	AMP (unflatten sedimentary)
9	APP (unflatten sedimentary)
10	AMP (nonredbeds)
11	APP (nonredbeds)
12	AMP (unflatten redbeds)
13	APP (unflatten redbeds)
14	AMP (published after 1983)
15	APP (published after 1983)
16	AMP (published before 1983)
17	APP (published before 1983)
18	AMP (exclude commented local rot or secondary print)
19	APP (exclude commented local rot or secondary print)
20	AMP (exclude local rot or correct it if suggested)
21	APP (exclude local rot or correct it if suggested)
22	AMP (filtered using SS05 palaeomagnetic reliability criteria)
23	APP (filtered using SS05 palaeomagnetic reliability criteria)
24	AMP (exclude superseded data already included in other results)
25	APP (exclude superseded data already included in other results)
26	AMP (comb of 22 and 24)
27	APP (comb of 23 and 25)

Notes: SS05, (Schettino & Scotese 2005)

overprints. So for this filter (Picking No 18 and 19; Table. 1), this type of data are removed.

Filtering Out Poles Influenced by Local Rotations or Correcting Them if Suggested Some publications suggest the data has been through local rotation and propose a solution of correction. With this filter (Picking No 20 and 21; Table. 1), if there is a correction suggestion, the data is corrected; if no, the data is simply removed.

Filtering Using SS05 Liability Criteria SS05 (Schettino & Scotese 2005) provided their criteria of picking paleomagnetic data for producing their APWPs. This filter is reproducing their criteria by setting α_{95} of $\leq 15^\circ$, age uncertainty of ≤ 40 Myr, sampling sites’ quantity of ≥ 4 , samples’ quantity of at least 4 times of the sites, and laboratory analytical procedure code of at least 2 (Picking No 22 and 23; Table. 1).

Filtering Out Superseded Data In this filter (Picking No 24 and 25; Table. 1), those superseded data already included in other newer results are excluded.

Filtering Using SS05 Liability Criteria Without Superseded Data The last two filters are also combined into one (Picking No 26 and 27; Table. 1) to see if SS05 Liability Criteria (Schettino & Scotese 2005) performs better using renewed datasets.

2.3.1.2 Weighting

Because all data is not created equal, we want to calculate a weighted mean pole for a time interval with ‘bet-

Table 2. List of all weighting algorithms developed in this paper.

No.	Weighting Algorithm
0	None (No weighting)
1	Larger Numbers of sites (B), Observations (N), greater weight
2	Lower Age uncertainty, greater weight
3	α_{95} (Normal distribution: standard deviation 10)
4	Age error Position to bin (more overlap, greater weight)
5	comb of 3 and 4

ter' (more likely to be reliable) poles counting more than 'worse'. For example, a pole with small α_{95} and very well constrained age is more likely to reflect APWP position at the selected age point than a pole with large α_{95} and very broad age range. There are many potential ways to weight this data set which can obviously greatly influence the final result, and we want to test this.

Six weighting algorithms (Table. 2) have currently been developed or reproduced according to published work to give different weights to data with different qualities.

In order to average errors in orientation of the samples and scatter caused by secular variation, a "sufficient" number of individually oriented samples (observations) from "enough" sites must be satisfied (Tauxe et al. 2018; van der Voo 1990; Besse & Courtillot 2002). So for the "Numbers of sites (B), Observations (N)" weighting (No 1 in Table. 2), larger B and N mean stronger weighting. Through knowing the pattern of all B and N in the database, the proposed solutions are as follows. If both B and N are more than 1, weight= $(1 - \frac{1}{B}) * (1 - \frac{1}{N})$. There are data in GPMDB with only the number of sampling sites (at least greater than 1) given, but no number of samples or only one sample given, so for this case, if $B > 1$ and $N \leq 1$, weight= $(1 - \frac{1}{B}) * 0.5$. If only the number of samples (at least greater than 1) is given, and the number of sampling sites is missing or only one, i.e. $B \leq 1$ and $N > 1$, weight= $(1 - \frac{1}{N}) * 0.5$. If $B \leq 1$ and $N \leq 1$ (there are only 23 datasets for the whole GPMDB 4.6b, including 18 with both B and N informations missing), weight=0.2.

As for the "Age uncertainty" weighting (No 2 in Table. 2), a well-constrained age should be known to within a half of a geological period (e.g., Quaternary, Neogene, Triassic) for Phanerozoic data (van der Voo 1990; Tauxe et al. 2018). Generally, this work follows this principle. However, for the periods of Paleogene, Cretaceous, and Jurassic, their halves are all beyond a time span of at least 20 Myr, which is relatively large for these relatively young geologic periods. So for these three periods, a tighter age constraint is set using age uncertainties of ≤ 15 Myr. So, for example, for NAC's Neogene (23.03–2.58 Ma according to GSA Geologic Time Scale) data, if age uncertainty (the high magnetic age – the low magnetic age) ≤ 10.225 (from $0.5^*(23.03–2.58)$) Myr, its weight = 1; if age uncertainty > 10.225 Myr, its weight = $10.225 / (\text{high magnetic age} - \text{low magnetic age})$. For the periods spanning Jurassic to Paleogene, if age uncertainty ≤ 15 Myr, it get its weight of 1; if age uncertainty > 15 Myr, a weight of $15 / (\text{high magnetic age} - \text{low magnetic age})$ is assigned instead.

For the " α_{95} " weighting (No 3 in Table. 2), smaller radius of circle of 95% confidence about mean remanence direction means less error, so should get larger weight. Here, weight is from a Gaussian distribution centered on 0 with standard deviation of 10, i.e., when $\alpha_{95}>0$, weight<1. There is no such case as $\alpha_{95}=0$, while there exists a blank α_{95} in the database. If α_{95} is not given, DP (the semi-axis of the confidence ellipse along the great circle path from site to pole) is assigned as α_{95} . If DP is also not given, α_{95} was fur-

ther approximated by $\frac{140}{KD*N}$, where KD is Fisher precision parameter for mean direction if this parameter is given, or Fisher precision parameter for Normal directions (K_NORM) if only K_NORM is given when KD is missing. If N is not given, B is used as N. If even K_NORM is also missing, the lowest KD value 1.7 that can be found in GPMDB 4.6b is used as KD. It is also worthwhile to mention that if samples, where two poles are derived, are exactly from the same place and same rock, and one α_{95} is completely inside the other α_{95} , a zero is assigned as the weight of the data with the larger α_{95} . In fact, in the above described procedure A95 (circle of 95% confidence about mean pole) is a better alternative instead of α_{95} , because A95 is directly reflecting the uncertainty of the paleopoles. However, the A95s of the most paleopoles are not given in GPMDB 4.6b, so α_{95} is used instead since α_{95} is also indirectly reflecting the quality of the dataset.

For the "Age error Position to window" weighting (No 4 in Table. 2), if window intersects with young/old end of age bracket or whole window overlaps with a part of age range, weight=(overlapping part) / (age range width); if whole age range is within window, weight=(window width) / (age range width) (note that when weight > 1, it is set back to 1).

The "Age error Position to window, and α_{95} " weighting (No 5 in Table. 2), is a combination of No 3 (but here the standard deviation is 15 though) and No 4.

Some of the picking (Table. 1) and weighting (Table. 2) methods developed here are also connected with the V90 Q factors mentioned above. For example, Pt 2, 3 and Wt 2, 4, 5 are related to the V90 criteria 1; Pt 2, 3, 22, 23, 26, 27 and Wt 1, 3 are related to the V90 criteria 2; Pt 22 and 23 are related to the V90 criteria 3; The data constraining described in the Supplementary Material is related to the V90 criteria 5; and Pt 18 and 19 are related to the V90 criteria 7.

2.4 Reference Paths: The Hotspot and Seafloor Spreading Model Predicted

After about 170 Ma, multiple data sources can help constrain plate motions in more accurate ways. One of the most developed and studied plate kinematics models is the Fixed Hotspot Model (FHM) (Müller et al. 1993; Müller et al. 1999), which assumes the Atlantic and Indian hotspots are relatively fixed. Another one is the Moving Hotspot Model (MHM) (O'Neill et al. 2005), which is based on mantle convection models that indicate large motions of the Indo-Atlantic hotspots.

Such a model like FHM or MHM can predict APWPs for main continents, e.g. the North America (Plate ID 101) (Fig. 5), the India (501) (Fig. 6) and the Australia (801) (Fig. 7), with the help of global tectonic plate motion data (i.e. plate circuits) from the ocean basins (i.e. spreading sea floors) that has been reconstructed for the last c. 180–200 Myr (although the extrapolation is required; hard to constrain uncertainties). These more accurate and totally different-source derived model-predicted APWPs can be compared to data from the paleomagnetic database. For about 120 Ma to the present, India drifts much faster than North America and Australia. North America has been drifting rather slow. Australia's drifting rate is between North America and India. In addition, India drifts almost north, whereas North America east and then north, and Australia west and then north. So these three continents are representatives for three different types of plate kinematics.

The oldest pole that can be predicted from the FHM is about 120 Ma. For example, the North American 120–0 Ma APWP predicted from this rotation model and latest published spreading ridge

rotations (collected plate circuit data will be shared as a supplementary material) will be taken as a reference path (Fig. 5), which will be compared with paleomagnetic APWPs for the same plate or continent.

2.5 Path Comparison Method

Chapter 2 proposes a new method of determining the degree of similarity between two APWPs that combines three separate difference metrics that assess both spatial distance of coeval points, and similarities in the bearing and length of coeval segments using a weighted linear summation. In this chapter, the path comparison method is used to measure similarity between the hotspot and seafloor spreading model predicted APWP and the 168 different paleomagnetic APWPs. This does not only help us find the most similar paleomagnetic APWP to the other-model predicted APWP, but also help further test the validity of the measuring tool in practise.

3 RESULTS AND DISCUSSIONS

3.1 Results

Is there any common pattern of the similarities for all the three continents? First, the best and worst methods need to be determined. Here, the difference values less than the one-standard-deviation interval (containing about 68.269% of the data values) are picked out as the “best” ones (lower about 15.866% of the data values), more than the one-standard-deviation interval as the “worst” ones (upper about 15.866%). Then we will see if there is one method or several methods labeled as “best” or “worst” for all the three continents.

3.1.1 When FHM and Global Plate Tectonic Circuits Predicted APWP as Reference

First, we focus on analysing the results with FHM and global plate circuits predicted APWP (Fig. 5, Fig. 6, Fig. 7) as the reference path.

3.1.1.1 10 Myr Binning and 5 Myr Stepping

According to the results (Fig. 8 and Fig. 13), we can observe:

- (i) the groups of picking-method-no 19 (APP with commented local-rotation or secondary-print studies excluded) and 21 (APP with local rotation excluded or corrected as suggested in the original sources) (Table. 1) are among the best ones for all the three continents, while the groups of picking-method-no 2 (AMP with “ $\alpha_{95}/\text{Age range}$ ” no more than “15°/20 Myr”) and 16 (AMP with only earlier-than-1983 studies), (Table. 1) are among the worst for all them three.

- (ii) for both North America (101) and Australia (801), the groups of picking-method-no 1, 11, 13, 19, 21 and 25 (Table. 1) are the best, and 2, 14, 16, 22 and 26 the worst. For both North America (101) and India (501), the groups of picking-method-no 4, 5, 7, 19 and 21 are the best, and 2, 8, 16 and 18 the worst. For both India (501) and Australia (801), the best 19, 21 and the worst 2, 16 are the same as the ones for all the three continents as above-mentioned. These results also further indicate that APP methods generally produce better similarity than AMP methods, however, the picking-method-no 4 is special, which is one of the AMP methods but also one of the best for both North America and India.

- (iii) the results of North America (101) and Australia (801) are closer (Fig. 13).

- (iv) For North America and India, newer studies (later than 1983) give better results, whereas for Australia older studies (before 1983) give better results, mainly because the number of older studies (65; Fig. 8c) is much greater than the number of newer studies (27).

- (v) Weighting influences more to APP than AMP, especially for the cases when there are plenty of paleo-poles (e.g., for North America and Australia; Fig. 14). Wt 0 and 1 generally work well with APP.

- (vi) Lithology related picking methods (Pt 4–13) are generally improving the fits, compared with other picking methods. However, sometimes, Wt 3 does not work well with these lithology-related methods (Fig. 8 and Fig. 14).

3.1.1.2 20 Myr Binning and 10 Myr Stepping

According to the results (Fig. 15), we can observe:

- (i) the groups of picking-method-no 19 is the best, and 16 the worst, for all the three continents.

3.1.2 When MHM and Global Plate Tectonic Circuits Predicted APWP as Reference

Then, we focus on analysing the results with MHM and global plate circuits predicted APWP (Fig. 5, Fig. 6, Fig. 7) as the reference path.

3.1.2.1 10 Myr Binning and 5 Myr Stepping

According to the results (Fig. 17a, Fig. 17b and Fig. 17c), we can observe:

- (i) there is no best picking method, whereas no. 16 is the worst, for all the three continents.

3.1.2.2 20 Myr Binning and 10 Myr Stepping

According to the results (Fig. 19), we can observe:

- (i) there is no best picking method, whereas no. 16 is the worst, for all the three continents.

3.1.3 Summary of Results

According to all the above results (Fig. 8 and Fig. 15), we can observe:

- (i) generally the APP methods (adding data to a time window with overlapping age selection criterion) produce better similarities than the AMP methods (Table. 3).

- (ii) the self-explanatory topography of bands indicates that the picking methods (Table. 1) influence the similarity more than the weighting methods (Table. 2) do.

- (iii) filtering (picking no 2–7, 10, 11, and 14–27) and correcting (picking no 8, 9, 12 and 13) has limited effectiveness.

- (iv) weighting is not always making similarities better. In fact, for quite many of the methods, no weighting is the best performer (Table. 3).

- (v) the picking method no 19 is always the best when the FHM predicted APWP is the reference.

- (vi) the weighting method no 0, 1 and 5 are generally producing better similarity than 2, 3 and 4.

- (vii) North America (101) owns better similarity results than Australia (801) and India (501), because its worst and mean composite differences are always less than the other two continents’.

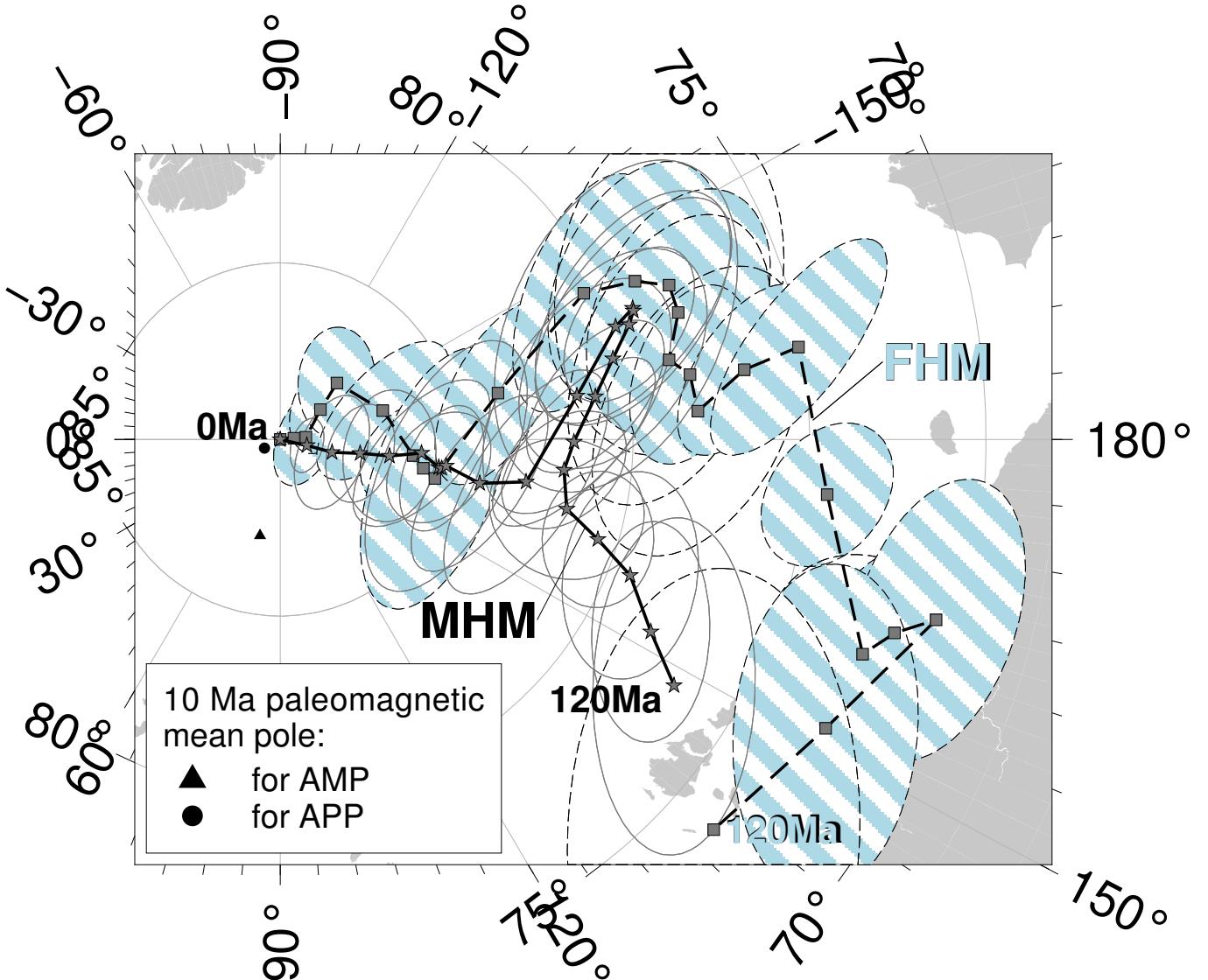


Figure 5. MHD predicted 120–0 Ma APWP (solid line) for *NAC* through the North America–Nubia–Mantle plate circuit. Its age step is 5 Myr. The dashed line is the FHM predicted path for comparison.

(viii) for both North America (101) and India (501), more recent studies generally give better results than (or results close to) older studies. However, this is not true for Australia (801) (Fig. 8c).

3.2 Discussions

The following discussions will be in Q&A style.

3.2.1 Question: Why the APP methods generally produce better similarities than AMP methods do?

Paleomagnetic (Mean) A95 represents precision (how well constrained calculated poles are), and (mean) coeval poles' GCD represents accuracy (how close calculated poles are to the reference path; Fig. 21a and Fig. 21b). Compared with AMP, APP usually improves both and generates paths with higher accuracy and also higher precision (generally increasing *N*).

The fact that APP increases the number of paleo-poles (*N*) in each sliding window would potentially average out some “bad” (i.e.

inaccurate) poles and improves the fit between the paleomagnetic APWPs and the model-predicted APWPs. The general effects that APP brings include the decreases in paleomagnetic A95s, or/and distances between compared coeval poles of paleomagnetic APWP and reference APWP (Fig. 21 and Fig. 22). However, if the added paleo-poles were all or mostly “bad”, the improvement of fit would not occur. So the improvement of fit is not only because of the increase in *N*, but also because the majority of the additional poles are “good”. AMP only regards the time uncertainty of each pole as one mid-point. Then this mid-point is treated as the most likely age of that mean pole. This is actually incorrect. The age uncertainty of paleopole is not obtained from a probability density function derived from an observed frequency distribution. As defined, the time uncertainty’s lower (older) limit is a stratigraphic age, and its upper (younger) limit could be also a stratigraphic age or be constrained by a tectonic event using the field tests (e.g. fold/tilt test and conglomerate test). So the true age of the pole could be any one that is not older than the lower limit and also not younger than the upper limit. In other words, the mid-point could be the true age of the

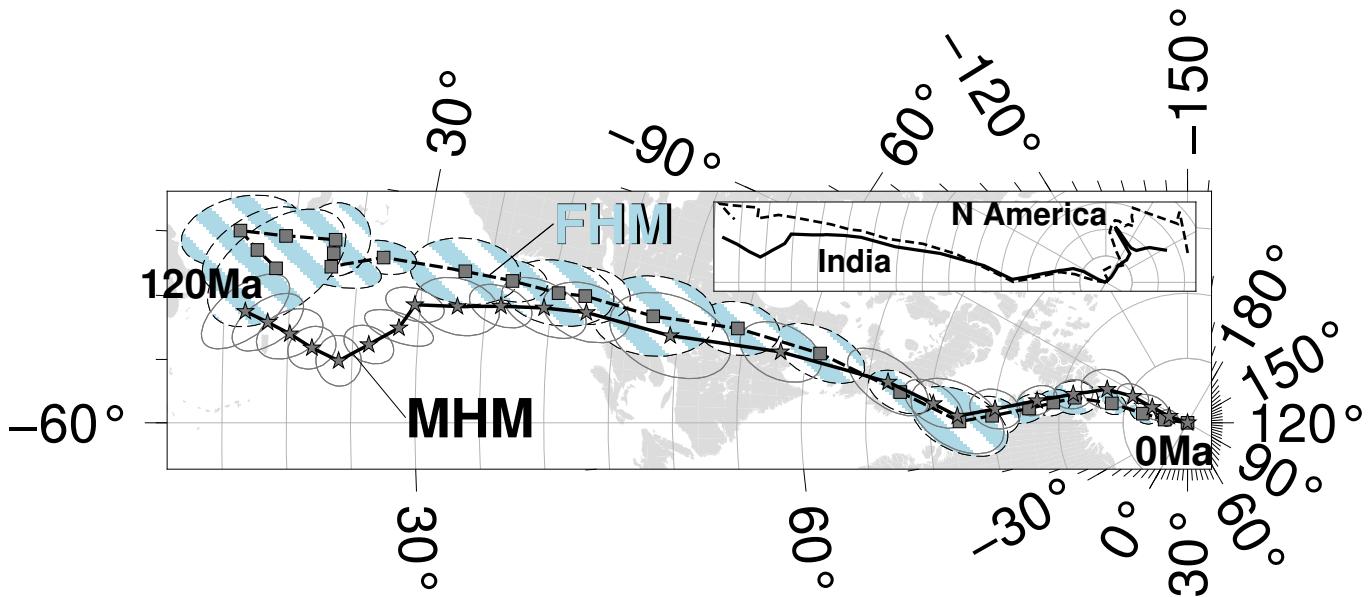


Figure 6. MHM predicted 120–0 Ma APWP (solid line) for India through the India–Somalia–Nubia–Mantle plate circuit. Its age step is 5 Myr. The dashed line is the FHM predicted path shown for comparison. The inset shows paths for fast moving India and also much slower moving North America shown in Fig. 5.

pole, but it is not known as the most likely age of that pole. If the mid-point is the most likely age of a pole, AMP should generate a path that is closer to the reference. However, mostly APP generates better similarities (See the high proportions of APP better than AMP in Table. 3). Most reasonably, the mid-point should be regarded as one possibility of all uniformly (not necessarily normally bell shaped, or U shaped, or left or right skewed) distributed ages between the two time limits.

So APP remains the effect of a paleopole borne on the mean poles during all the period of its age uncertainty, and use the increased number of paleopoles (N) to average out the negative effect of those “bad” poles, including the paleopoles that should not be included at that age for mean pole.

3.2.2 Question: Why the AMP methods sometimes unexceptionally produce better similarities than APP methods do?

Because of small number of paleo-poles (not necessarily “bad”) involved in each sliding window, the produced mean poles by AMP should be relatively far from its contemporary model-predicted pole. In other words, AMP intends to give fairly small change in accuracy. This also could potentially bring more distinguishable d_s for AMP if the corresponding A95 is not large enough. For example, for Fig. 8a, there are only three special (of 84 APP vs. AMP comparisons) cases picking/weighting 4/3, 4/5, 6/3 better than 5/3, 5/5, 7/3 respectively. Compared with the picking/weighting 4/3 APWP, although most of the mean paleopoles are closer to the FHM predicted APWP and also the number of the significant pole pairs is one less for the APP derived path (i.e. 5/3), the A95s are smaller and most importantly there are one more significant d_a orientation-change pair and two more significant d_l segment pair (Table. 4). If we observe carefully, it is because of the much smaller 15 Ma A95 for 5/3. The similar phenomenon occurs to the case of 6/3 vs 7/3, a relatively much smaller paleomagnetic A95 causes

more distinguishable d_a and d_l for the APP results, and they offset the improvement of spatial similarity d_s APP brings.

For 4/5 vs 5/5, all d_a and d_l are indistinguishable. Compared with the results from AMP, although the coeval pole GCDs are all decreased for APP, this spatial improvement is not able to offset the negative effects of also generally decreased paleomagnetic A95s, which potentially brings more statistically distinguishable coeval poles (e.g. the 15 Ma and 30 Ma poles for picking 5 and weighting 5; Table. 5). This further causes greater distinguishable mean d_s from the APP methods. The similar phenomenon occurs to Fig. 8b picking 4 vs 5 with all the six types of weightings, Fig. 8c picking/weighting 4/2–3 vs 5/2–3, 4/5 vs 5/5.

In addition, compared with AMP, APP potentially could generate more mean poles, because sometimes for some sliding window there is no paleo-pole involved at all for AMP while there is(are) paleo-pole(s) involved for APP. For APP, the mean poles at all ages should be composed of more paleo-poles than it is for AMP, which should generally decrease both coeval pole distance and paleomagnetic A95. However, sometimes a rare case (e.g. the 0 Ma comparison shown in Table. 6) happens. It is sometimes that an additional very “bad” paleo-pole gets included by APP and this increases both coeval pole distance and paleomagnetic A95 even though N increases. Such cases include Fig. 8b picking 22 vs 23 (actually exactly the same as picking 26 vs 27) with all the six types of weightings.

So generally as we discussed in the last section APP decreases the distances between paleomagnetic APWPs and the hotspot and sea-floor spreading model predicted APWP, and also the uncertainties of paleomagnetic APWPs. However, as we described in this section, special cases like decreased A95 potentially intends to make coeval poles differentiated if the coeval poles’ distance is not decreased effectively or even increased, or very “bad” paleo-poles got involved in some sliding windows, occurs. In summary, when the negative effect from these types of rare cases is beyond the positive effect the generally improved mean poles contribute,

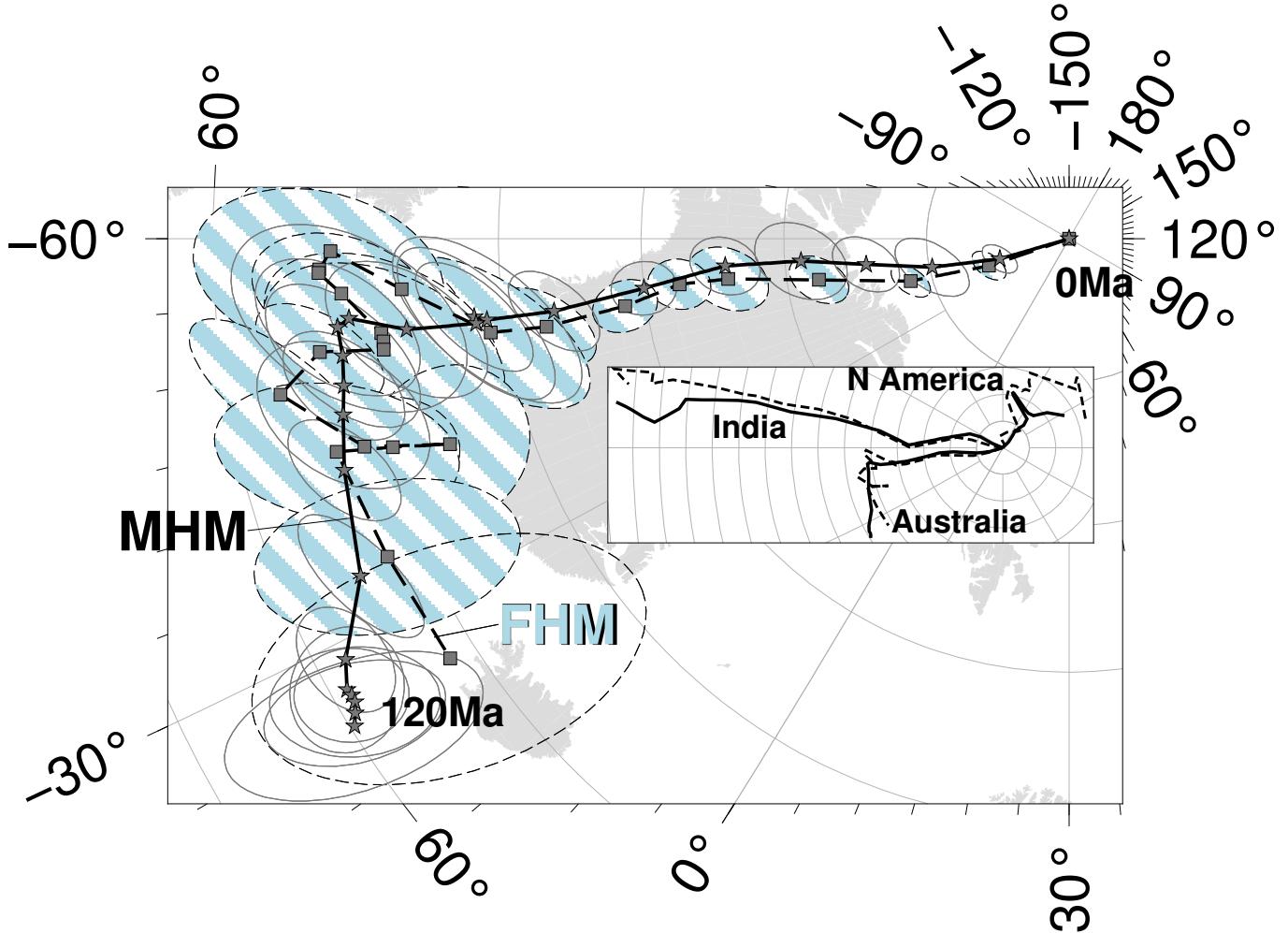


Figure 7. MHM predicted 120–0 Ma APWP (solid line) for Australia through the Australia–East Antarctica–Somalia–Nubia–Mantle plate circuit. Its age step is 5 Myr. The dashed line is the FHM predicted path shown for comparison. The inset shows paths for fast moving India shown in Fig. 6, much slower moving North America shown in Fig. 5, and also relatively intermediate moving Australia.

Table 3. Performance statistics of all the picking and weighting methods.

Grid	Best No.		Worst No.		Proportion of APP Better Than AMP	For All 28 Picking Methods, Count of Occurrences of Each Weighting No. Being Best					Picking 14/15 (Studies After 1983) Better Than 16/17 (Older)		
	Picking	Weighting	Picking	Weighting		0	1	2	3	4	5		
FHM	Fig. 8a	1, 4, 5, 7, 11, 13, 15, 19 , 21 , 25	0, 1, 5	2, 5, 7, 8, 14, 16 , 17, 18, 22, 26	0, 1, 2, 3, 4, 5	27/28	15	4	2	3	1	4	Y/Y
	Fig. 8b	4, 5, 6, 7, 9, 19 , 21	0, 1, 2, 3, 4, 5	0, 2, 8, 10, 12, 16 , 18, 20, 24	0, 1, 2, 3, 4, 5	11/14	17	1	2	7	1	0	Y/Y
	Fig. 8c	1, 11, 13, 17, 19 , 21 , 25	0, 1, 3, 5	2, 6, 14, 16 , 22, 26	0, 1, 2, 3, 4, 5	27/28	11	4	1	8	3	1	N/N
	Fig. 15a	1, 4, 5, 7, 11, 13, 15, 19 , 21 , 25	0, 1, 3, 5	2, 6, 14, 22, 26	0, 1, 2, 3, 4, 5	3/4	18	6	0	2	4	3	N(5n1y)/Y
	Fig. 15b	4, 5, 6, 7, 19	0, 1, 2, 3, 4, 5	0, 2, 10, 12, 16 , 18, 20, 23, 24, 27	0, 1, 2, 3, 4, 5	59/84	12	6	0	9	0	1	Y/Y(4y2n)
	Fig. 15c	1, 11, 13, 17, 19 , 21 , 25	0, 1, 3, 4, 5	2, 14, 16 , 22, 26	0, 1, 2, 3, 4, 5	41/42	7	9	0	11	1	2	N/N(4n2y)
MHM	Fig. 17a	1, 4, 5, 7, 11, 13, 15, 21 , 25	0, 1, 2, 3, 4, 5	0, 8, 10, 12, 14, 16 , 17, 18, 20, 22, 24	0, 1, 2, 3, 4, 5	11/12	14	6	1	1	2	4	Y/Y
	Fig. 17b	4, 5, 6, 7, 19 , 22, 26	0, 1, 2, 3, 4, 5	0, 2, 8, 10, 12, 16 , 18, 20, 24	0, 1, 2, 3, 4, 5	11/14	12	3	0	9	1	3	Y/Y
	Fig. 17c	1, 11, 13, 17, 19 , 21 , 25	0, 1, 3, 5	2, 14, 16 , 22, 26	0, 1, 2, 3, 4, 5	20/21	8	5	1	9	4	1	N/N
	Fig. 19a	4, 5, 7, 11, 15, 22, 26	0, 1, 2, 3, 4, 5	0, 8, 10, 12, 14, 16 , 18, 20, 24	0, 1, 2, 3, 4, 5	3/4	15	6	1	0	1	5	N(5n1y)/Y
	Fig. 19b	4, 5, 6, 7, 19	0, 1, 2, 3, 4, 5	0, 2, 10, 12, 16 , 18, 23, 24, 27	0, 1, 2, 3, 4, 5	29/42	8	8	2	5	3	2	Y/N(5n1y)
	Fig. 19c	1, 11, 13, 15, 17, 19 , 21 , 23, 25, 27	0, 1, 2, 3, 4, 5	2, 14, 16 , 22, 26	0, 1, 2, 3, 4, 5	41/42	9	5	0	8	5	1	N/Y(4y2n)

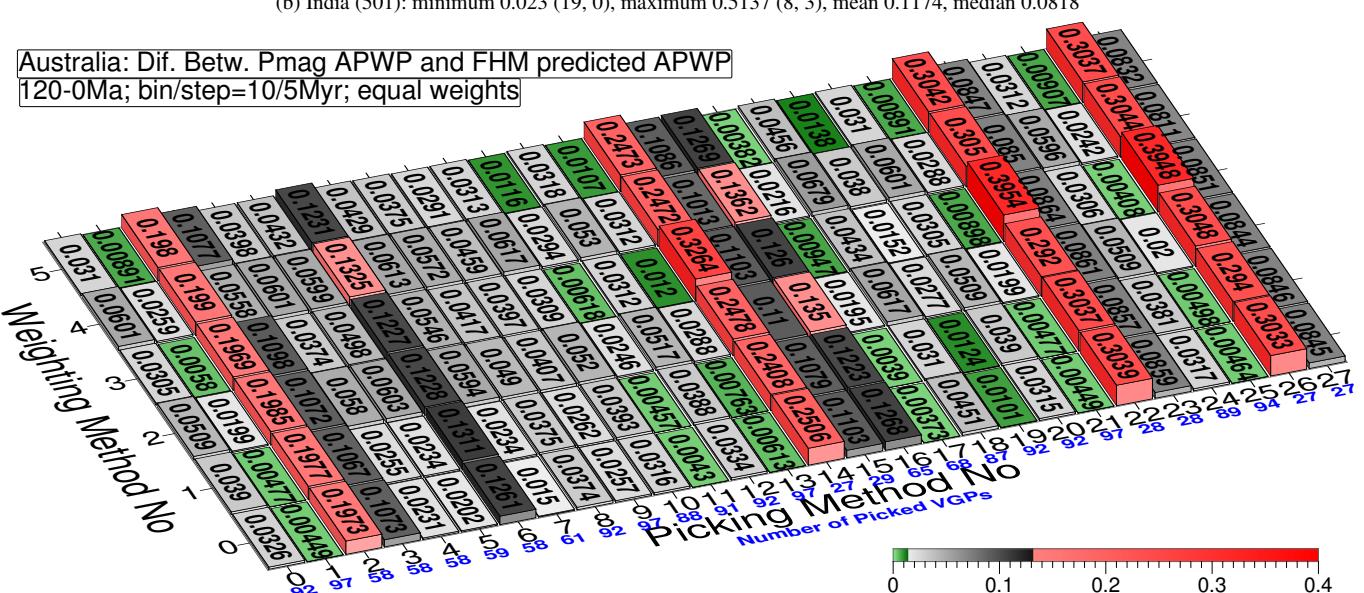
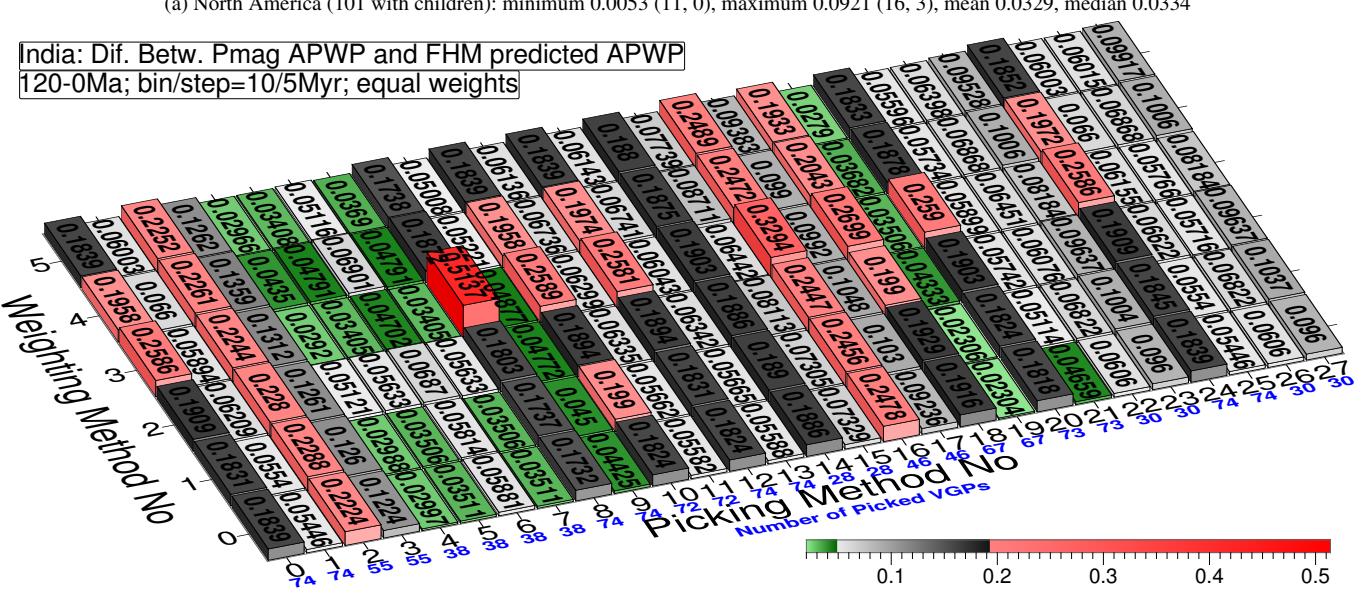
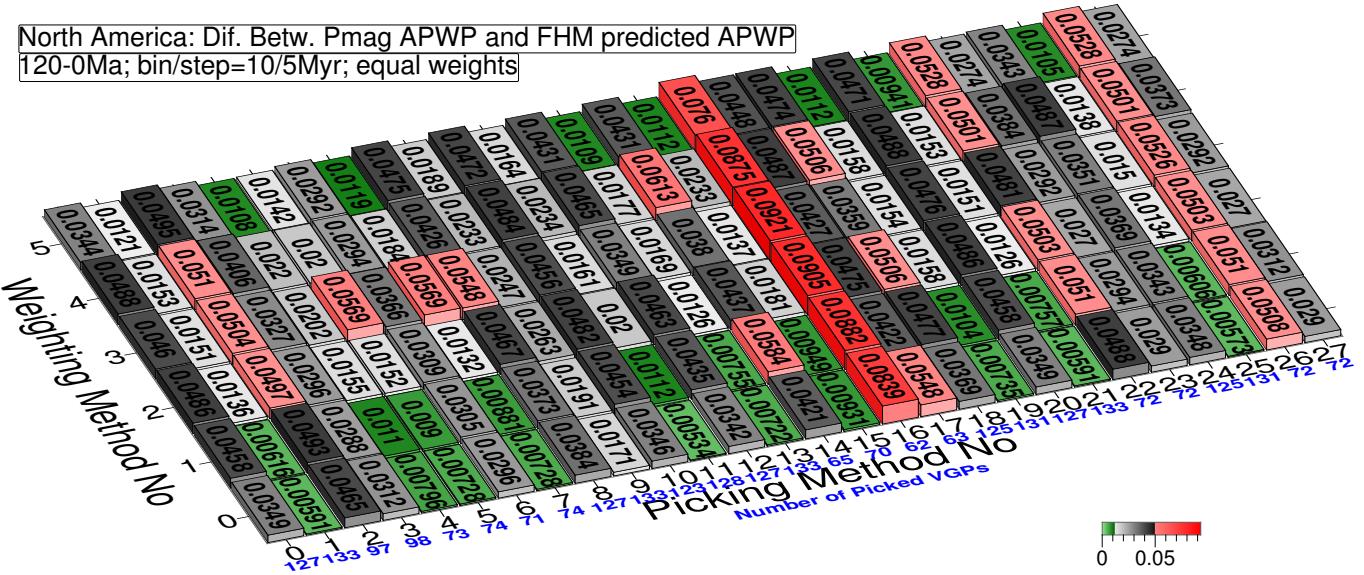


Figure 8. Equal-weight composite path difference (CPD) values between each continent's paleomagnetic APWPs and its predicted APWP from FHM and related plate circuits. The paths are in 10 Myr bin and 5 Myr step. The difference values less than one-standard-deviation interval of the whole 168 values are colored in green, more than one-standard-deviation interval colored in red.

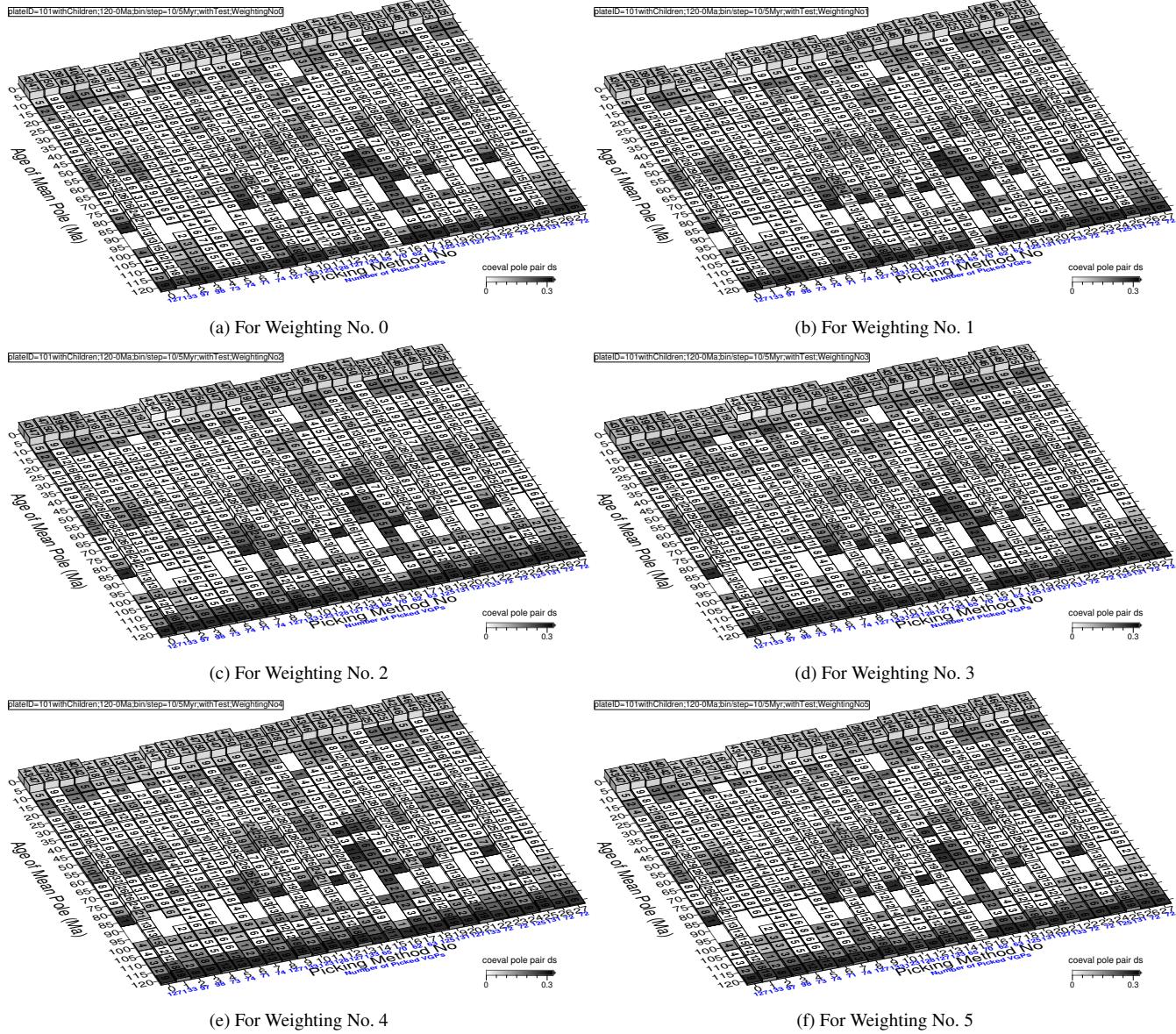


Figure 9. Tested spatial difference (ds) values (color shaded) between North American paleomagnetic APWPs and its predicted APWP from FHM and related plate circuits. The paths are in 10 Myr bin and 5 Myr step. The labeled numbers on the grids are the numbers of site mean poles that are contributing to each mean path pole.

the composite difference score would increase. However, this phenomenon seldom occur (Table. 3).

Other Type 1 (e.g. Table. 4) cases: Fig. 15a picking/weighting 22/3 vs 23/3. Fig. 17a 6/3 vs 7/3.

Other Type 2 (e.g. Table. 5) cases: Fig. 15a picking/weighting 2/0–5 vs 3/0–5, 4/0–2 vs 5/0–2, 4/4 vs 5/4, 6/5 vs 7/5, 6/2–3 vs 7/2–3, 22/4–5 vs 23/4–5, 22/1 vs 23/1, 26/4 vs 27/4, 26/2 vs 27/2. Fig. 15b 14/3 vs 15/3. Fig. 15c 4/2 vs 5/2. Fig. 17a 4/0–1 vs 5/0–1. Fig. 17b 4/0–4 vs 5/0–4, 22/0–1 vs 23/0–1, 22/3–4 vs 23/3–4, Fig. 17c 4/3 vs 5/3, 4/0 vs 5/0, 8/5 vs 9/5, 8/3 vs 9/3. Fig. 19b 14/2–3 vs 15/2–3. Fig. 19c 4/4 vs 5/4.

Combination of both Type 1 and 2 cases: Fig. 15a picking/weighting 26/0–1 vs 27/0–1. Fig. 17a 4/3 vs 5/3. Fig. 17b 22/2 vs 23/2, 22/5 vs 23/5, 26/0–5 vs 27/0–5.

Other Type 3 (e.g. Table. 6) cases: Fig. 15b picking/weighting 6/0–5 vs 7/0–5, 22/0–5 vs 23/0–5, 26/0–5 vs 27/0–5. Fig. 17a 4/2–5

vs 5/2–5. Fig. 19a 2/0–5 vs 3/0–5, 4/2 vs 5/2, 4/4 vs 5/4, 6/2 vs 7/2, 22/0–5 vs 23/0–5, 26/0–5 vs 27/0–5. Fig. 19b 6/0–5 vs 7/0–5, 22/0–5 vs 23/0–5, 26/0–5 vs 27/0–5.

3.2.3 Question: Why weighting is not affecting?

Generally, weighting does not affect the similarities dramatically, because the six results from the six weighting methods are mostly very close. There are a few special cases that one of them dramatically changes similarity (e.g. for picking method (Pk) 5 and weighting method (Wt) 3, etc.; Fig. 8a). As follows, those normal cases with close results are discussed first. Those rare cases are examined later.

For Pk 0 in the North America (101) example (Fig. 8a), the reason of why Wt 1–4 do not produce better similarity scores than Wt 0 (i.e. no weighting) does is examined here. Take the compar-

Table 4. One example of the Type 1 rare cases where AMP gives better similarity result than APP does from North America (101). Only statistically significant values are listed here.

FHM predicted		picking 4 + weighting 3				picking 5 + weighting 3			
Age (Ma)	A95 (°)	Age (Ma)	Pmag A95 (°)	Dist (°)	Age (Ma)	Pmag A95 (°)	Dist (°)	Age (Ma)	Diff (°)
0	0	0	8.3	4.29	0	7.6	3.885	10-15-20	126.59
5	1.56039/0.87367	5	7	4.853	5	6.8	4.3453		
10	2.89214/1.58743	10	12	9.91	10	8.6	5.79	dl	
15	2.575/1.63303	15	58/49	6.72	15	2.0857	11.8	Age (Ma)	Diff (°)
20	3.16077/2.20094	20	12.43	8.58				10-15	13.52
25	4.96061/2.2183	25	6.76	6.96	25	6.3358	6.873	15-20	14.68
30	3.39692/2.57114	30	6.68	6.46	30	6.68	6.4583		
115	9.27023/5.16012	115	8.6	12.2535	115	8.5	11.7		
120	14.6882/8.12086	120	7.76	15.5744	120	7.728	15.258		

Table 5. One example of the Type 2 rare cases where AMP gives better similarity result than APP does from North America (101). Only statistically significant values are listed here. Note that the number of the ages of the significant differentiated mean poles are the same for both paleomagnetic APWPs in this type of situation.

Age (Ma)	FHM predicted A95 (°)	ds			
		picking 4 + weighting 5 Pmag A95 (°)	Dist (°)	picking 5 + weighting 5 Pmag A95 (°)	Dist (°)
0	0	8.3526	3.79	7.6664	3.391
5	1.56039/0.87367	7.5215	4.17	7.29	3.74
10	2.89214/1.58743	21.351	4.76	14.598	2.87
15	2.575/1.63303			10.153	10.71
30	3.39692/2.57114			6.9664	5.7245
115	9.27023/5.16012	9.635	11.935	9.452	11.276
120	14.6882/8.12086	8.0223	15.8183	7.943	15.511

Table 6. One example of the Type 3 rare cases where AMP gives better similarity result than APP does from India (501). Only statistically significant values are listed here. Note that for the bold-number ages, there is no mean poles at all for the “picking 22 (AMP) + weighting 0” case.

FHM predicted		picking 22 + weighting 0				picking 23 + weighting 0			
Age (Ma)	A95 (°)	Pmag A95 (°)	Dist (°)	N	Pmag A95 (°)	Dist (°)	N	Age (Ma)	Diff (°)
0	0	6.85	12.973	2	23.6214	17.698	3	80-85	6.286
5	1.415/0.965	23.6214	15.941	3	23.6214	15.941	3	110-115	16.684
10	2.2425/1.34645	5.4/3.1	29.897	1	5.4/3.1	29.897	1		
15	2.2694/1.62543				5.4/3.1	28.28	1		
70	8.53016/4.97567	4.0164	4.864	16	3.246	4.436	20		
75	5.3554/3.1595				5	4.477	1		
80	8.41657/5.00588				5	3.358	1		
85	5.01489/2.49492	5	7.632	1	5	7.632	1		
90	7.77997/2.86845				5	10.884	1		
95	4.46779/3.24941				5	11.099	1		
100	5.6124/5.19639				5	11.4155	1		
105	4.64657/3.49277				5	14.908	1		
110	9.11039/4.98436				6.8/4.9	13.962	1		
115	9.27023/5.16012	10.73	10.508	5	10.73	10.508	5		

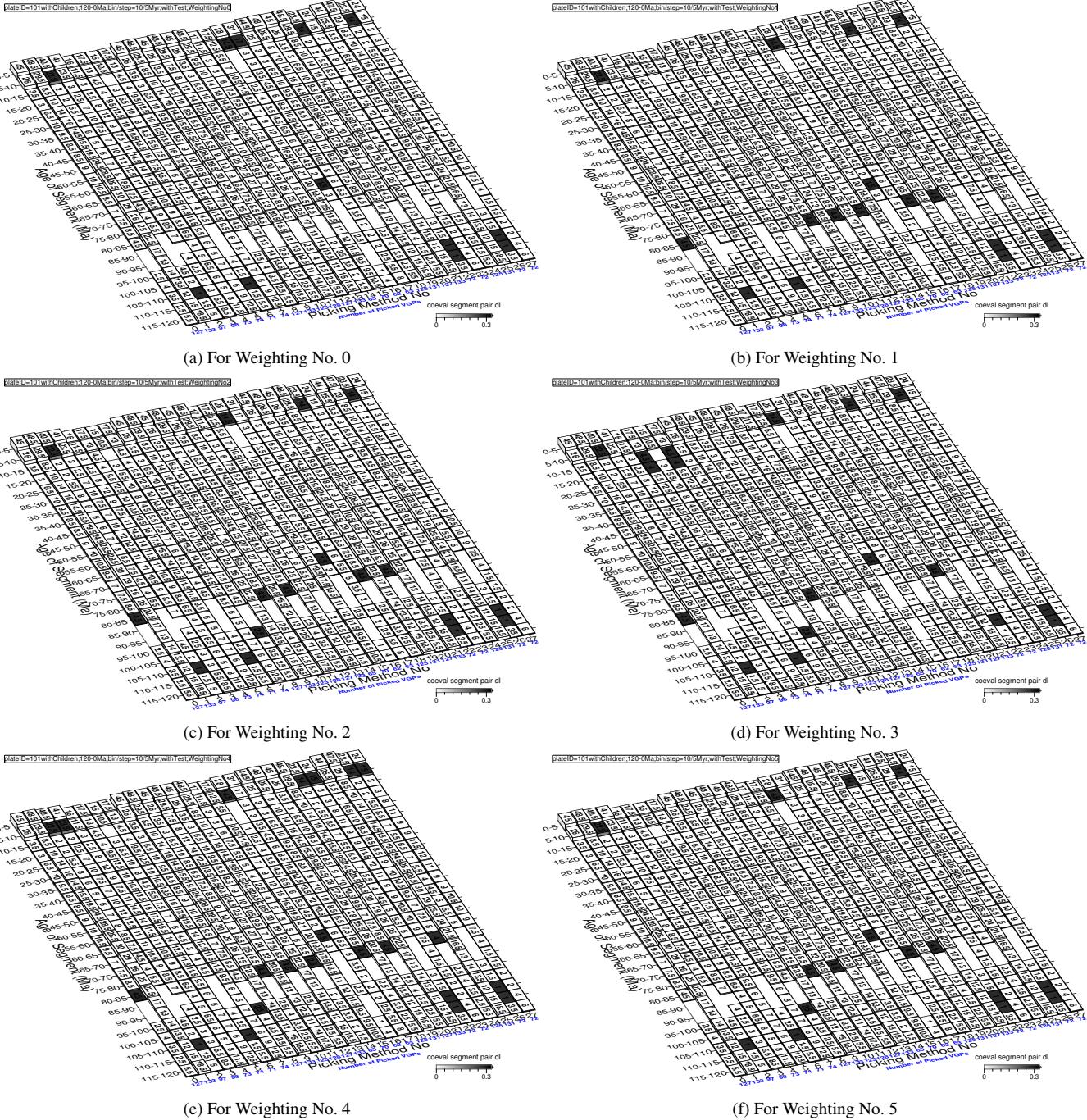


Figure 10. Tested length difference (dl) values (color shaded) between North American paleomagnetic APWPs and its predicted APWP from FHM and related plate circuits. The paths are in 10 Myr bin and 5 Myr step. The labeled numbers on the grids are the averaged numbers of site mean poles that are contributing to each segment's two mean path poles.

ison between Wt 0 and Wt 1 (Table. 7) as an example, although APP indeed decreases the values of most distances from the reference and most A95s, APP also brings one more significant differentiated segment-length difference (80–85 Ma coeval segments; Fig. 23) because of the slightly larger A95 and shorter coeval pole GCD for 80 Ma Pk 0 Wt 1. Although they are distinguishable, the overlapping of their distributions is quite small (Fig. 23b). Wt 2 and 4 cause one more pair of distinguishable mean poles (d_s) than other weighting methods. However, compared with Wt 0, Wt 5 (i.e. combination of α 95 and age error position to bin) does affect and

bring the best similarity, because Wt 5 decreases the GCDs of the 20, 80 and especially 100 Ma distinguishable pole pairs (by more than 2 degrees; Table. 7).

For Fig. 8a Pk 1, Wt 0 generates the best similarity. Wt 2–5 cause at least 3 more pairs of distinguishable mean poles (d_s) than other weighting methods. Wt 1 increases the GCDs of all the distinguishable pole pairs (0, 5 and 120 Ma).

For Fig. 8a Pk 2, Wt 0 generates the best similarity. Wt 1–5 cause at least 1 more pair of distinguishable mean poles (d_s) than

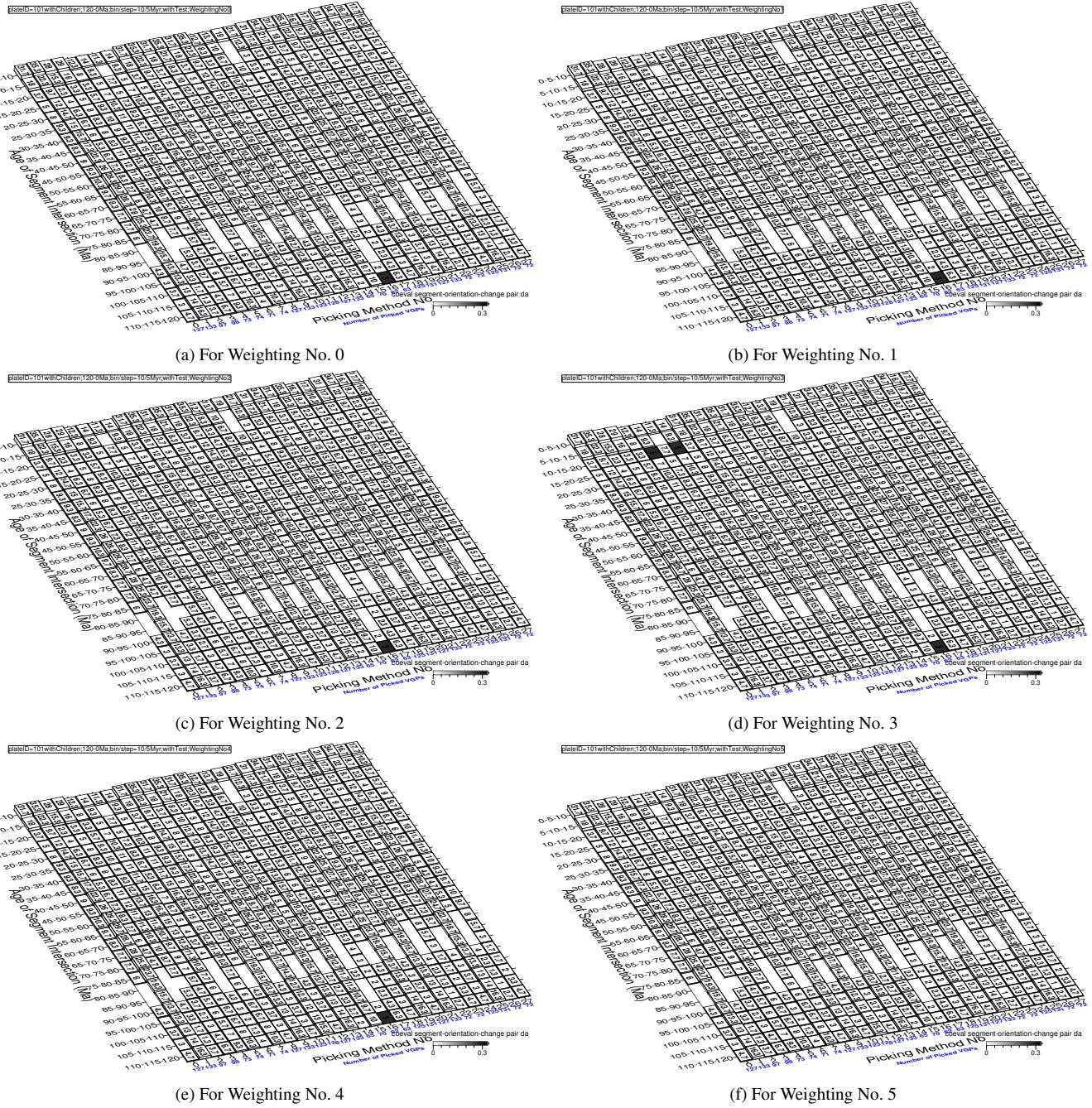


Figure 11. Tested angular difference (da) values (color shaded) between North American paleomagnetic APWPs and its predicted APWP from FHM and related plate circuits. The paths are in 10 Myr bin and 5 Myr step. The labeled numbers on the grids are the averaged numbers of site mean poles that are contributing to each segment-orientation-change's three mean path poles.

Wt 0 (60 Ma). Wt 4 also increases another pair of distinguishable mean poles (d_s) (80 Ma).

As the above examples shown, although no weighting is often better than weighting, the results are actually rather close to each other and also APP still improves precision and accuracy of most poles for the cases with larger difference scores, although APP brings extra significant differences in shape metrics which is also the main cause of their larger difference scores (Table. 7). However, APP does not always improve precision or accuracy, for example, for the Pt 5 (mainly or only igneous derived) and Wt 3 (according to α_{95}) results (Table. 8). For Pt 5, the result with no weighting is

still the best, and the results from Wt 0–2, 4, 5 are close and APP generally improves precision and accuracy, just like the above mentioned general cases. Only for Wt 3, APP worsens precision and improves accuracy of most poles at the same time. This makes the difference score dramatically larger than others. In addition, Wt 3 brings more significant differences in shape metrics, which causes the difference score even larger (Table. 8). That Wt 3 sometimes gives dramatical worse similarity could be related to the way we do the weighting according to the size of α_{95} , i.e. the smaller the α_{95} is, the larger the assigned weight is (see more details about Weighting No. 3 in the Weighting paragraph). However, small size

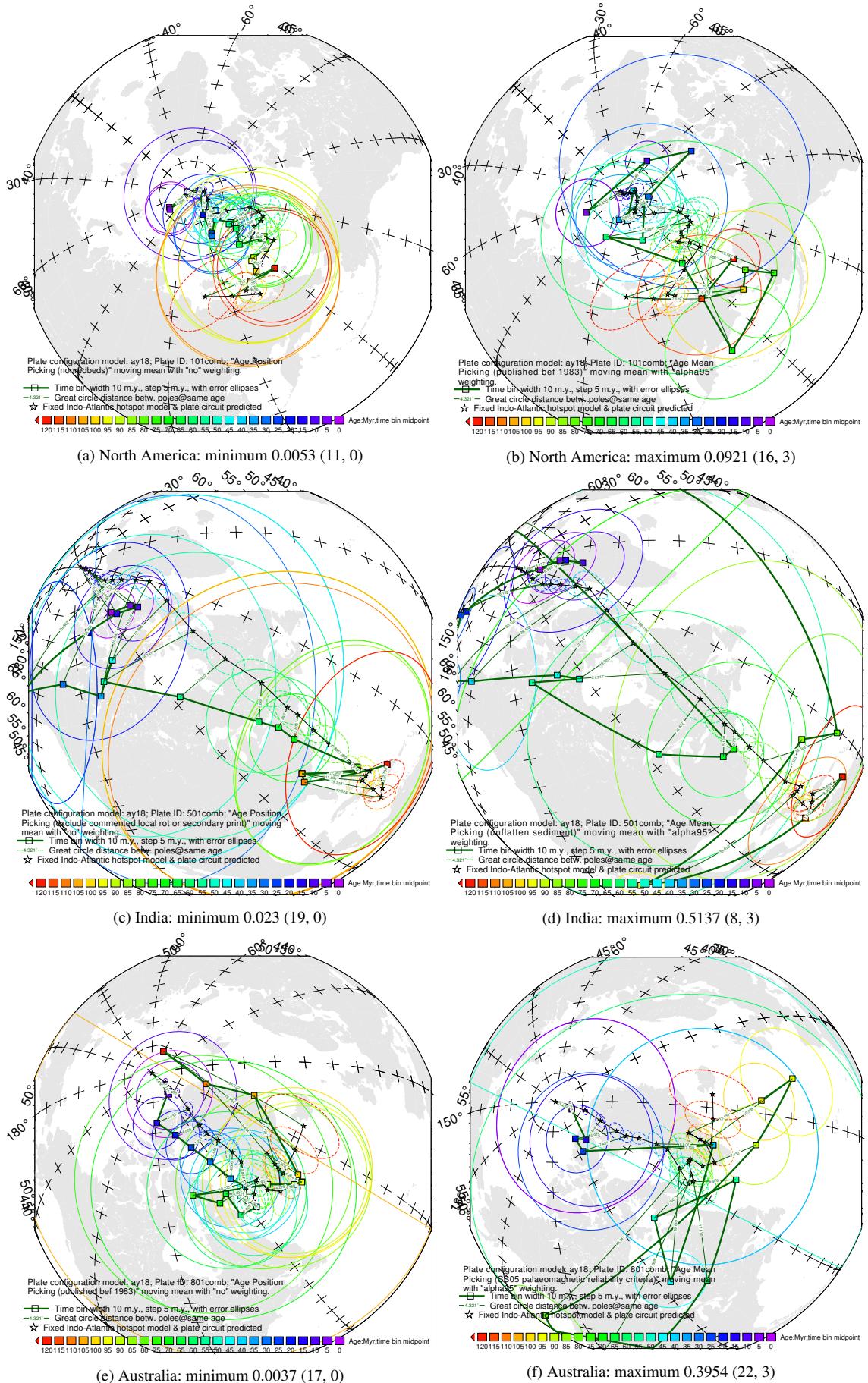


Figure 12. Path comparisons with best and worst difference values shown in Fig. 8. The parenthetical remarks are Picking No and Weighting No.

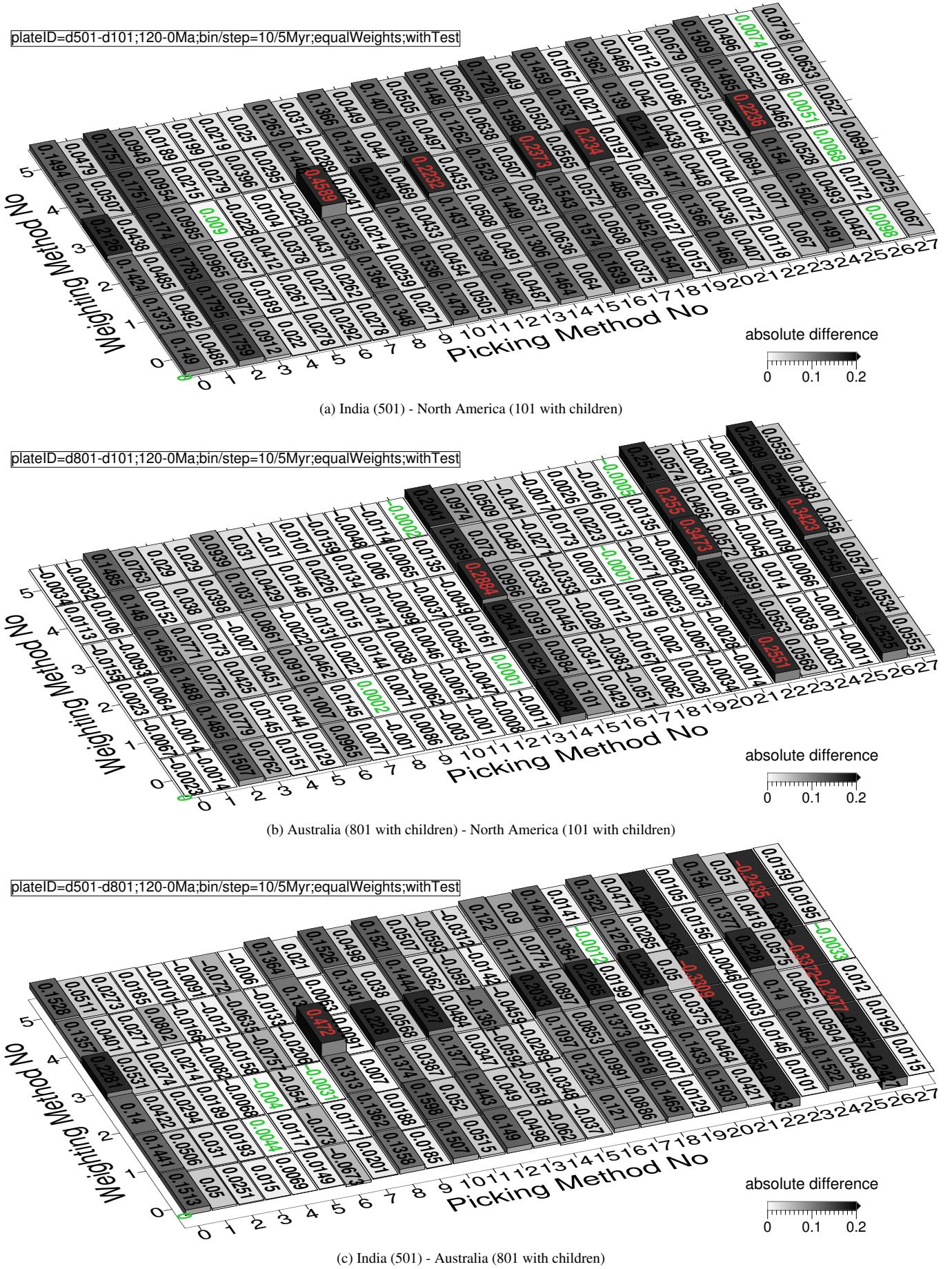


Figure 13. Differences between grids in Fig. 8. The absolute difference values less than 1.96-standard-deviation interval of the whole 168 values are labeled in green, more than 1.96-standard-deviation interval labeled in red.

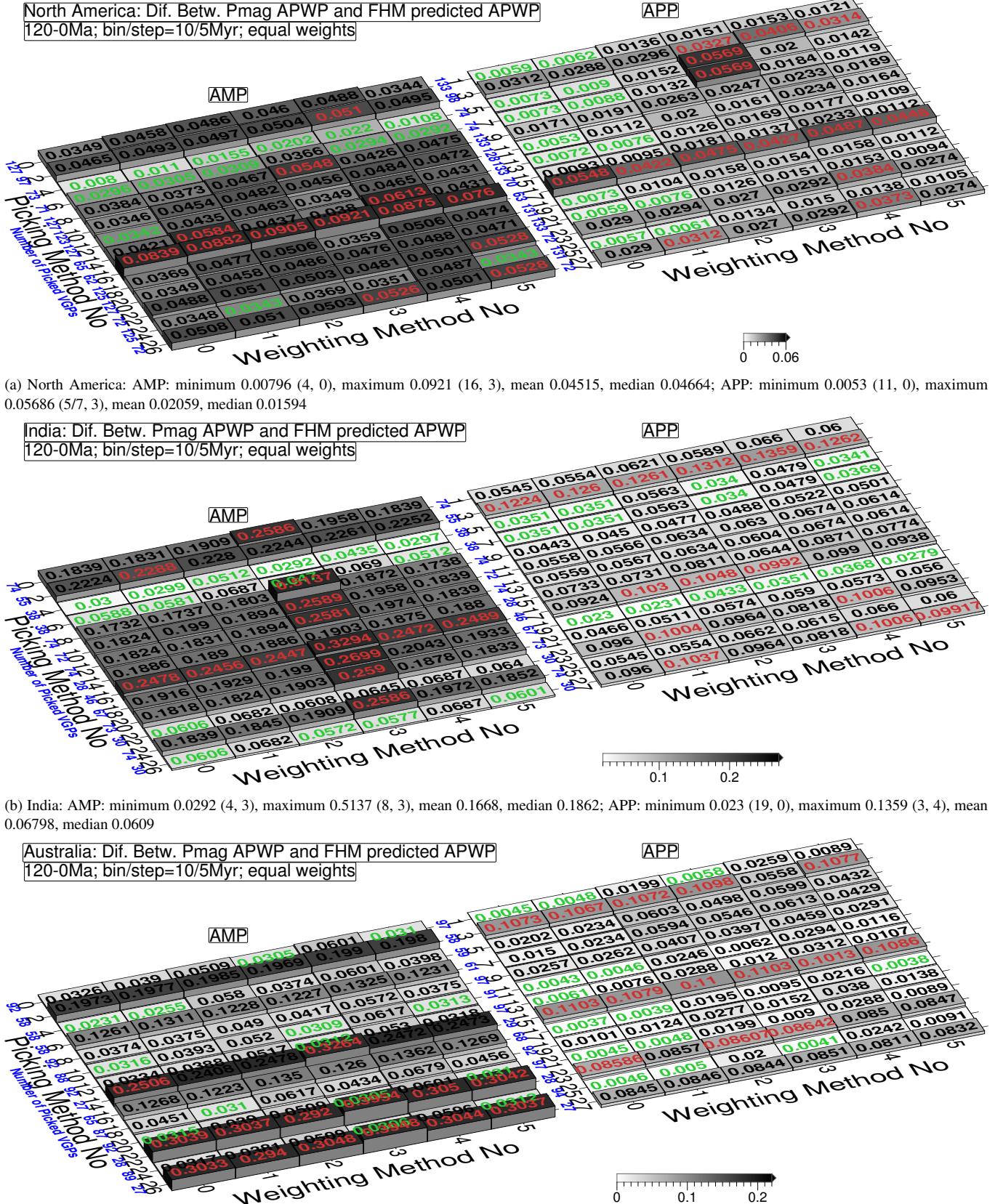


Figure 14. Separated results from AMP and APP in Fig. 8. For each grid block (left: AMP; right: APP), the difference values less than one-standard-deviation interval of the whole 84 values are labeled in green, more than one-standard-deviation interval labeled in red.

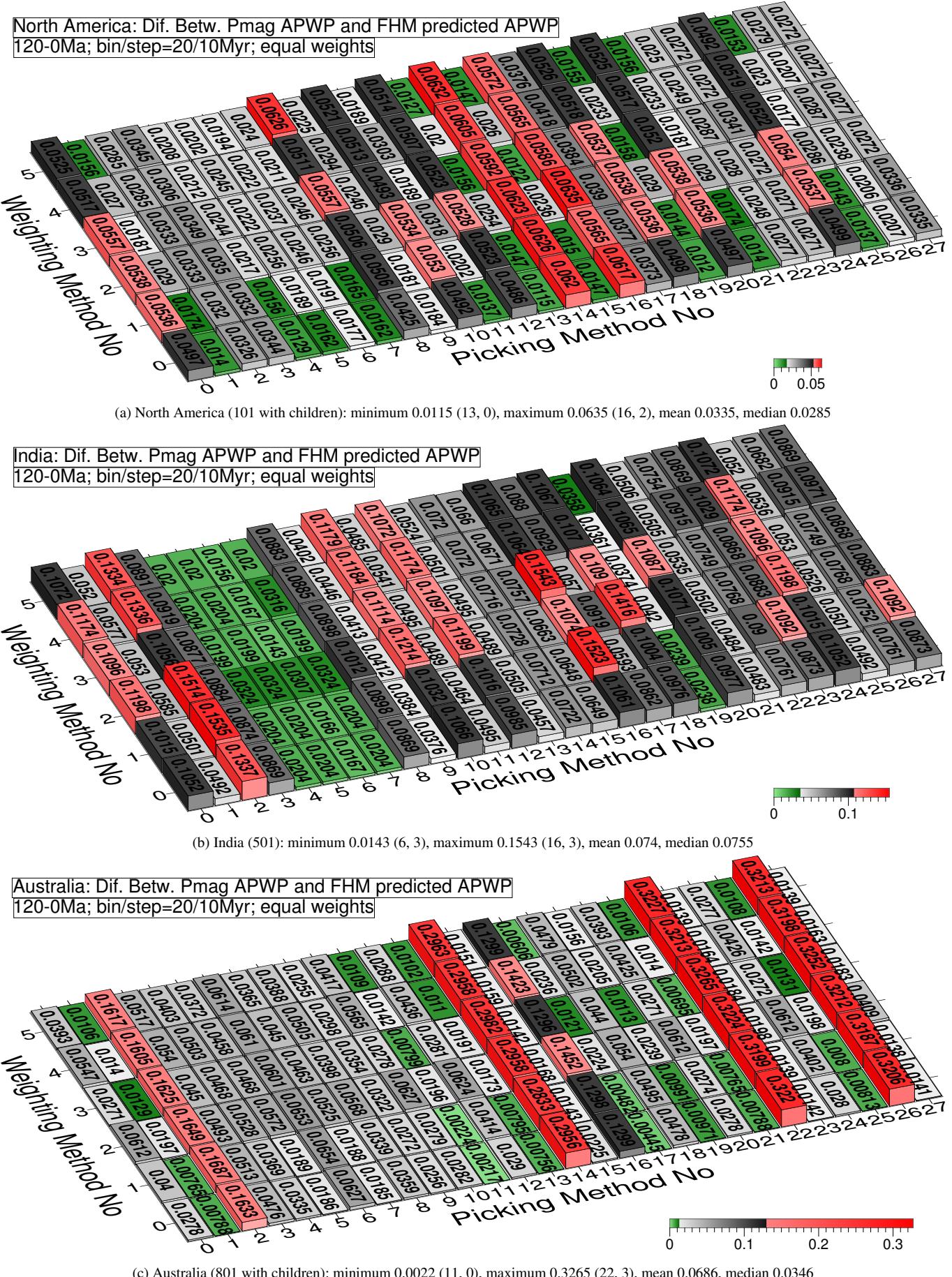


Figure 15. Same as Fig. 8. The only difference is here the paths are in 20 Myr bin and 10 Myr step. The difference values less than one-standard-deviation interval of the whole 168 values are colored in green, more than one-standard-deviation interval colored in red. See the numbers of picked paleo-poles in Fig. 8.

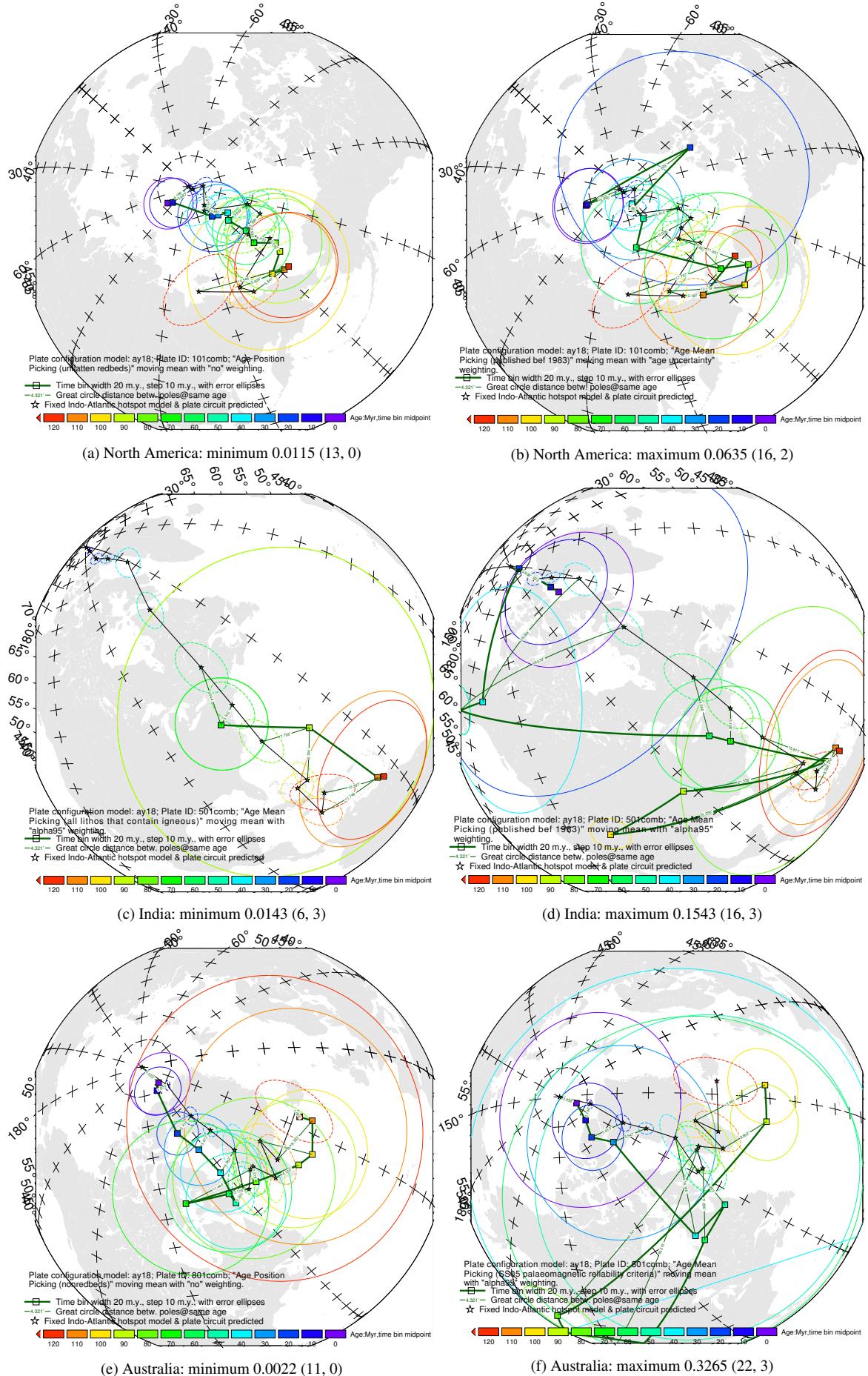
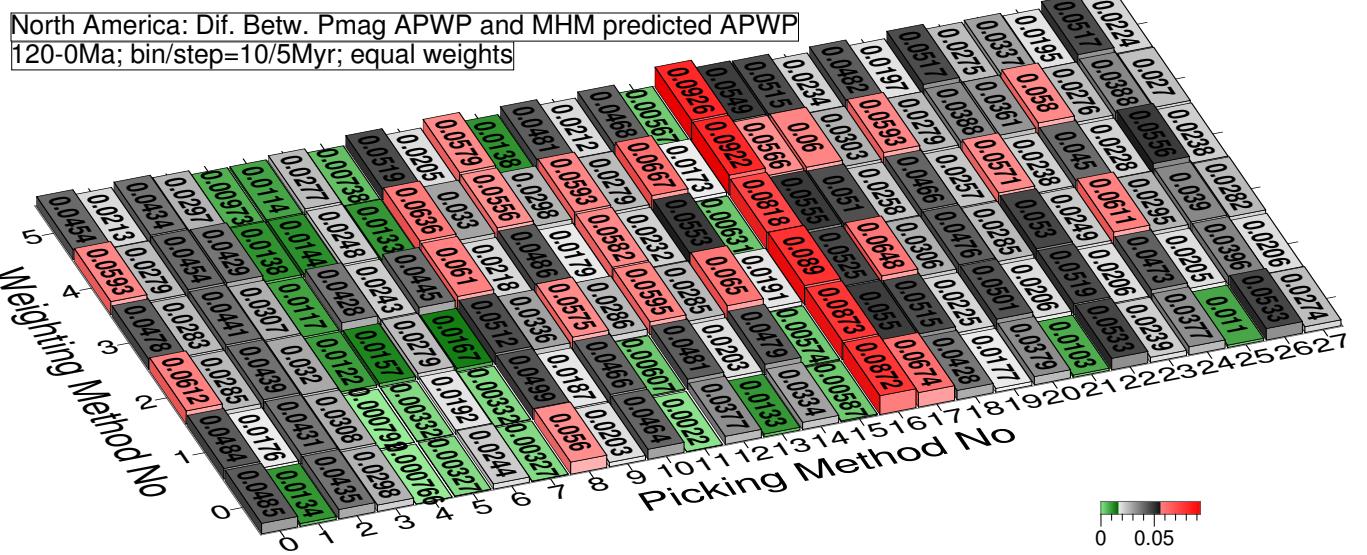
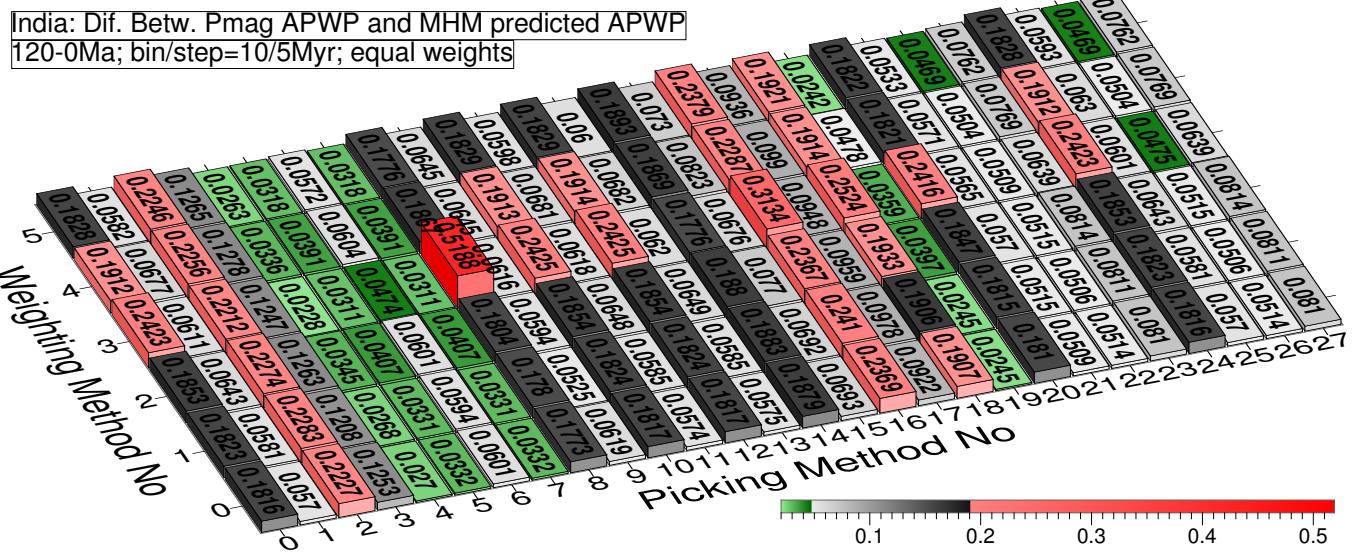


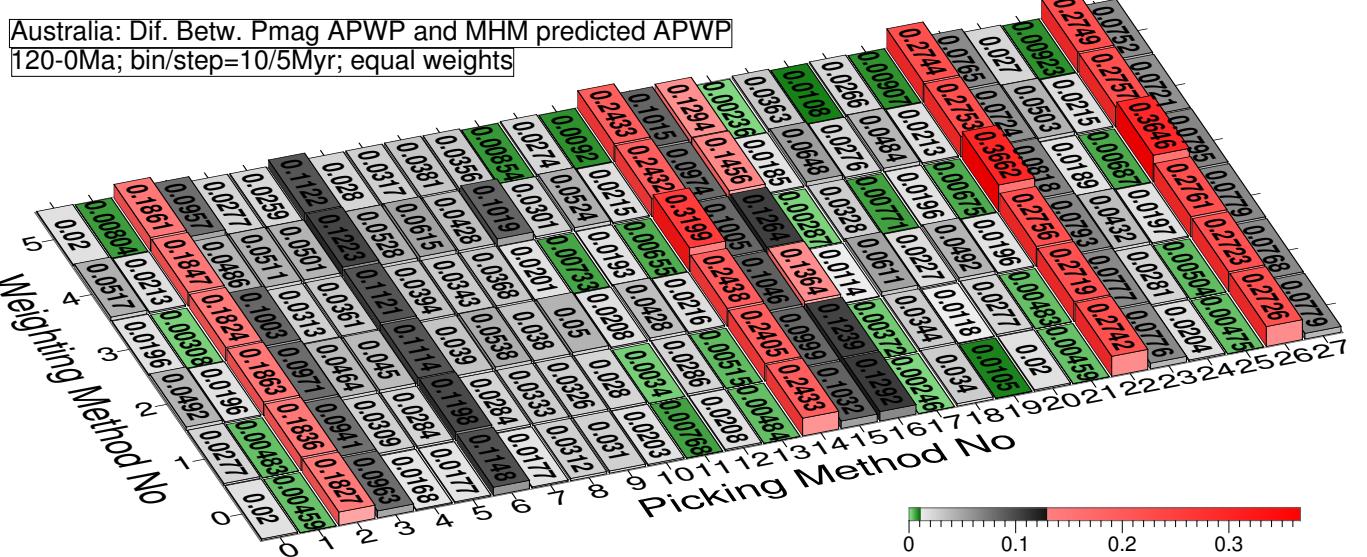
Figure 16. Path comparisons with best and worst difference values shown in Fig. 15. The parenthetical remarks are Picking No and Weighting No.



(a) North America (101 with children): minimum 0.00077 (4, 0), maximum 0.0926 (16, 5), mean 0.0361, median 0.0325



(b) India (501): minimum 0.0228 (4, 3), maximum 0.5188 (8, 3), mean 0.113, median 0.0746



(c) Australia (801 with children): minimum 0.0024 (17, 5), maximum 0.3662 (22, 3), mean 0.0744, median 0.0374

Figure 17. Same as Fig. 8 except that the reference path is predicted from MHM here. See the numbers of picked paleo-poles in Fig. 8.

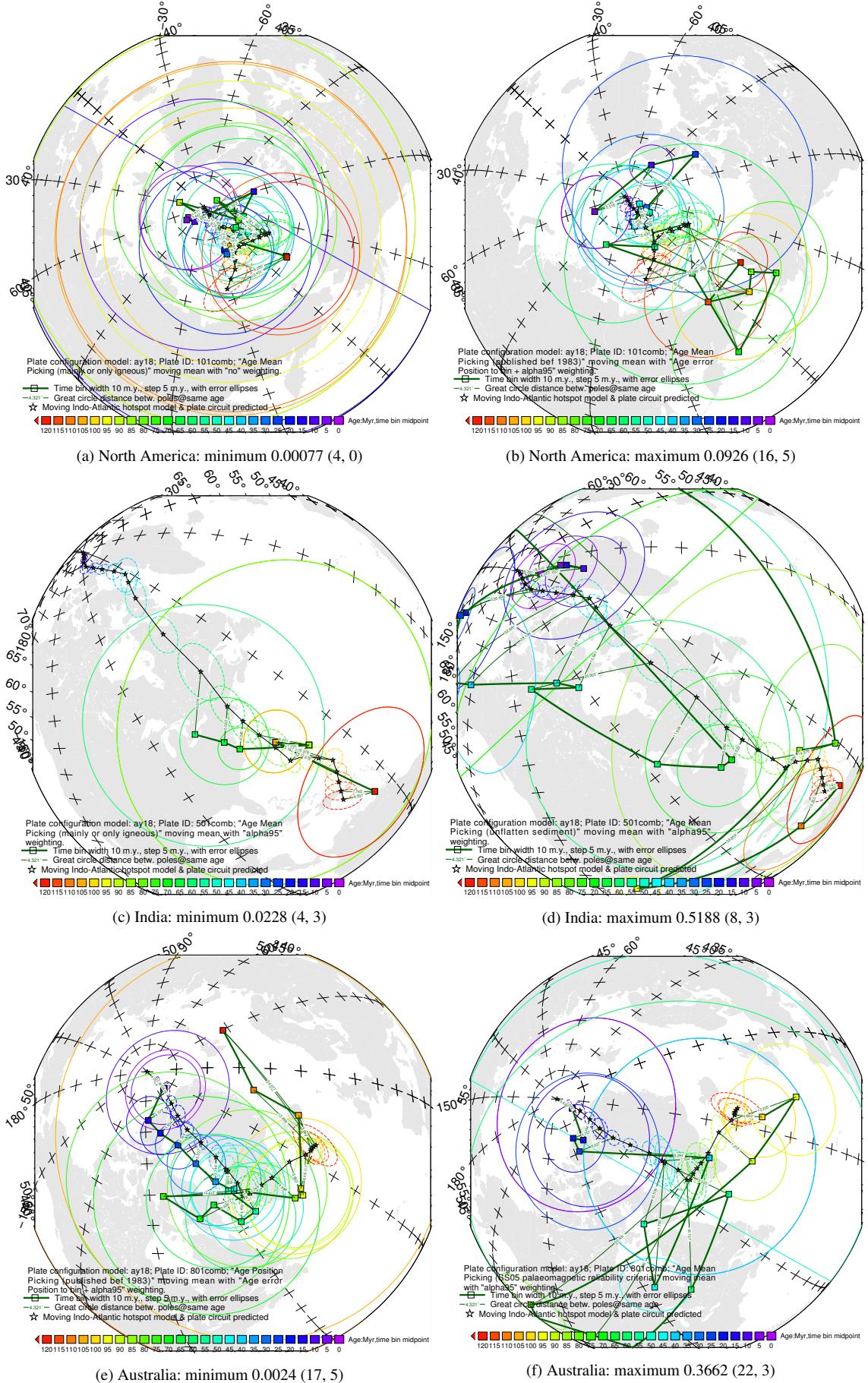
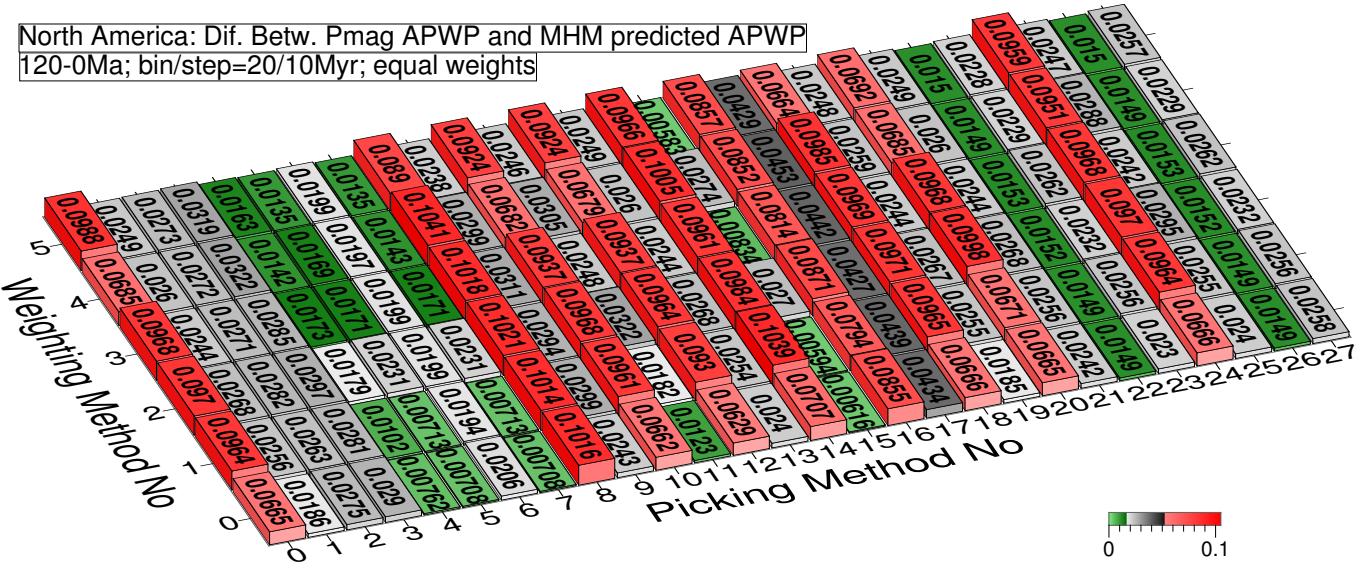
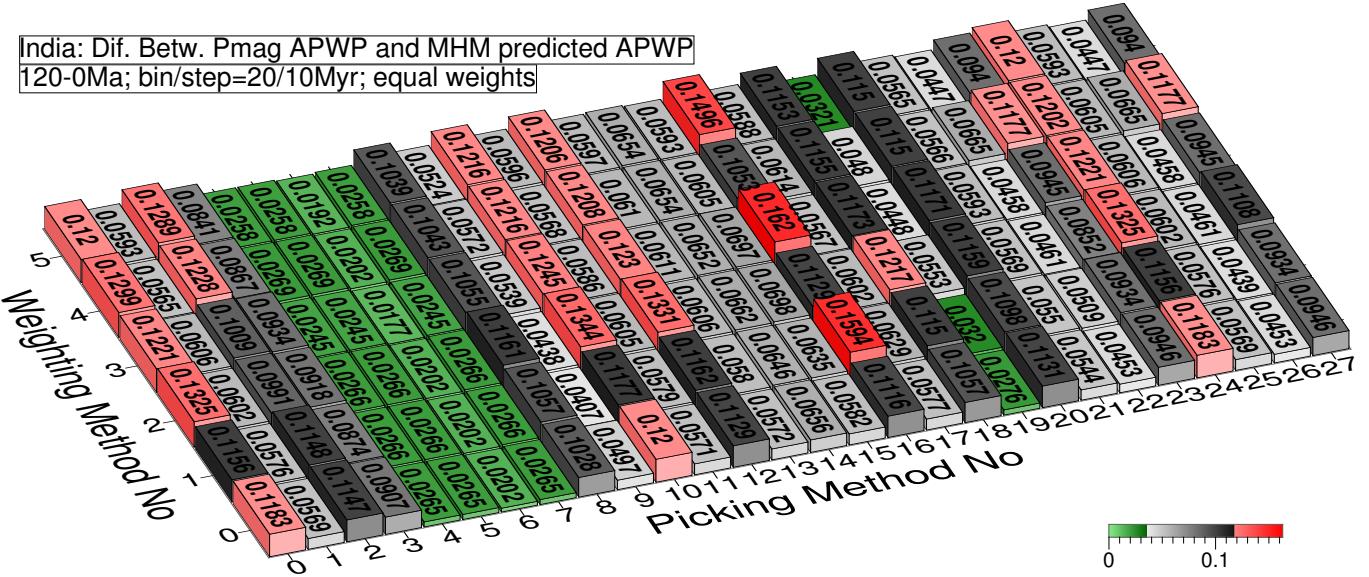


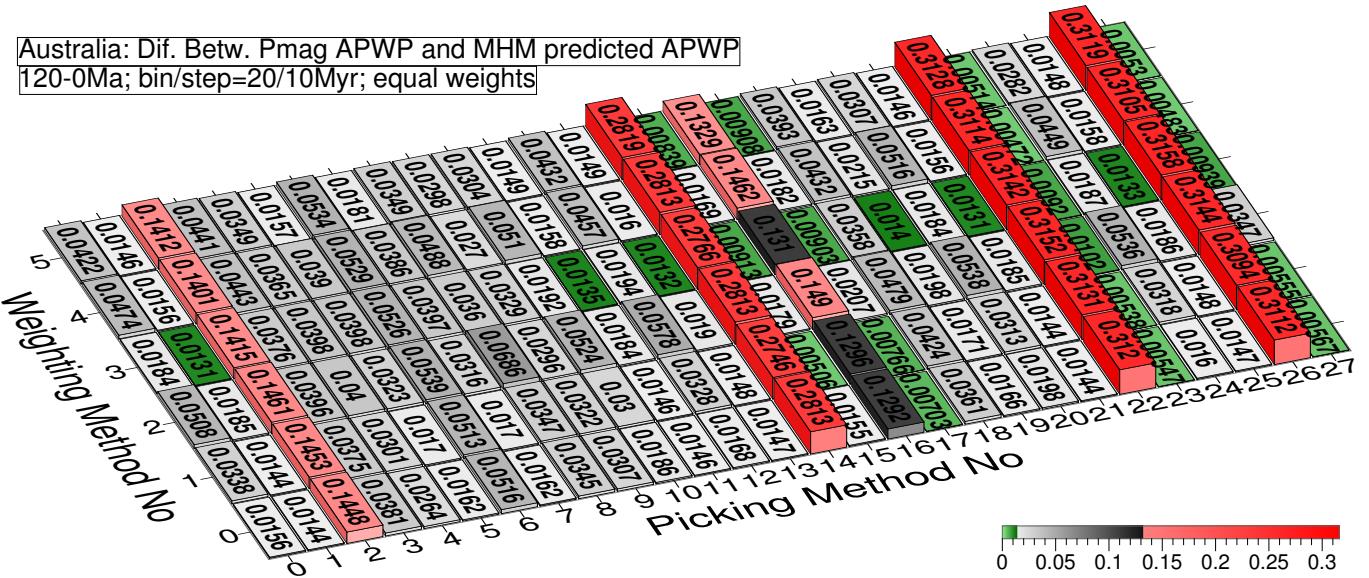
Figure 18. Path comparisons with best and worst difference values shown in Fig. 17. The parenthetical remarks are Picking No and Weighting No.



(a) North America (101 with children): minimum 0.0058 (15, 5), maximum 0.1041 (8, 4), mean 0.0437, median 0.0267



(b) India (501): minimum 0.0177 (6, 3), maximum 0.162 (16, 3), mean 0.0759, median 0.0621



(c) Australia (801 with children): minimum 0.0047 (23, 4), maximum 0.3158 (26, 3), mean 0.0638, median 0.0307

Figure 19. Same as Fig. 15 except that the reference path is predicted from MHM here. See the numbers of picked paleo-poles in Fig. 8.

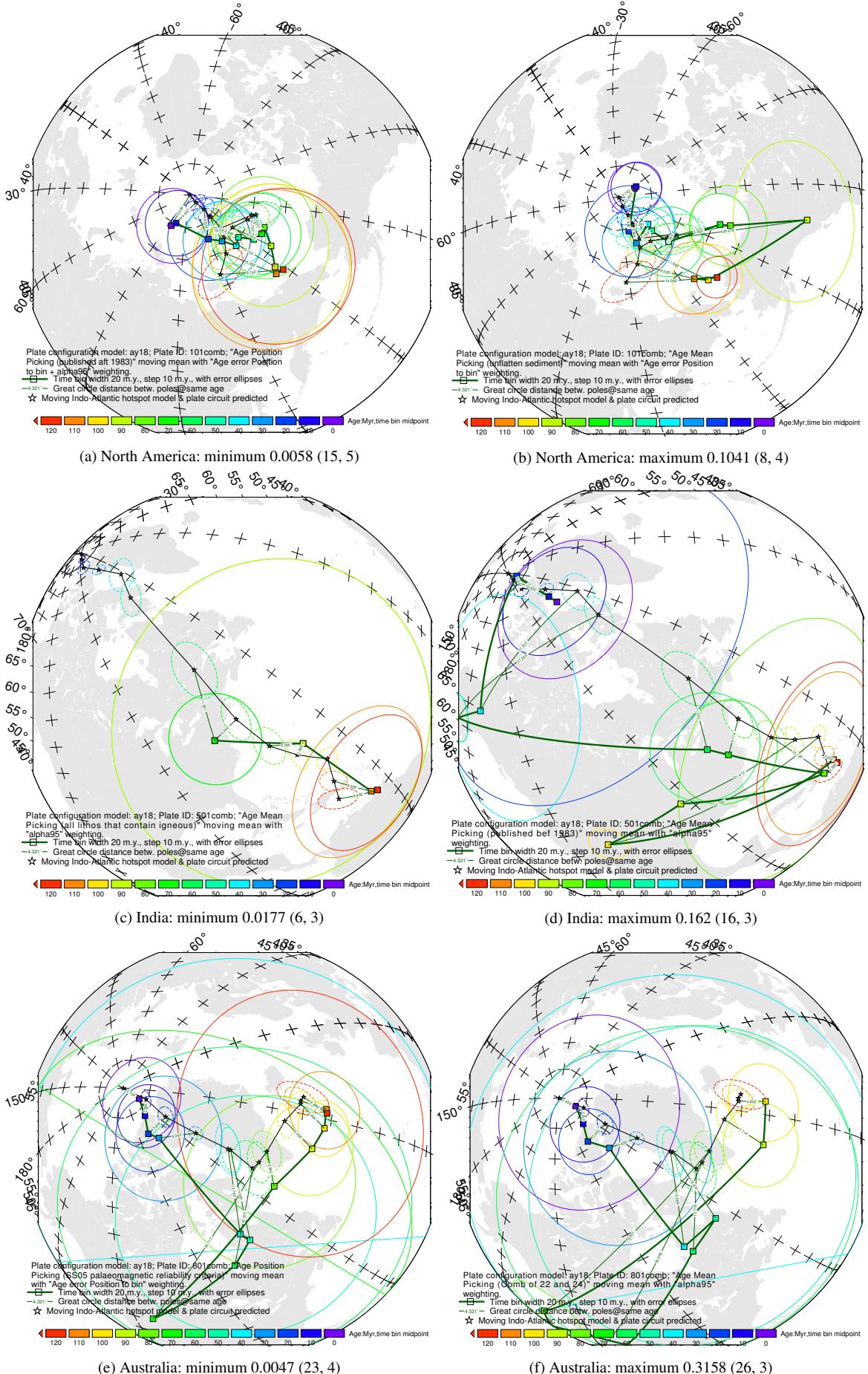


Figure 20. Path comparisons with best and worst difference values shown in Fig. 19. The parenthetical remarks are Picking No and Weighting No.

Table 7. My caption

FHM predicted for 101	Pk 0 + Wt 0				Pk 0 + Wt 1				Pk 0 + Wt 5	
	ds		ds		dl		ds			
Age (Ma)	A95 (°)	Pmag A95 (°)	Dist (°)	Pmag A95 (°)	Dist (°)	Age (Ma)	Diff (°)	Pmag A95 (°)	Dist (°)	
0	0	4.27286602	5.01	3.950661	5.05647	80-85	11.103	4.143	5.356	
5	1.56039/0.87367	4.22350537	5.146	3.936601	5.1286			4.0534	5.407	
10	2.89214/1.58743	20.9920176	3.076	19.826829	3.22624			19.868	3.36	
15	2.575/1.63303	13.8698147	10.3	13.827757	10.34			13.85	10.2753	
20	3.16077/2.20094	8.36501201	7.2162	8.463096	6.9973			8.413	6.7906	
50	7.15565/3.22656	4.3991229	6.22	4.651517	6.273			4.3326	6.3563	
55	7.17564/4.28065	5.6991191	8.647	5.670719	9.6724			5.52	8.53	
60	9.71876/6.35204	7.71537555	9.498	6.889233	8.607			7.77	9.8237	
80	8.76515/5.14459	6.29356332	9.26	6.452368	8.098			6.033	8.369	
85	5.54221/2.65419	8.7/6.7	18.995	8.7/6.7	18.995			8.7/6.7	18.995	
100	5.79659/5.36693	10.2720657	10.75658	9.286878	9.035			9.045	8.68	
115	9.27023/5.16012	19.767437	9.074	18.483547	10.054			19.652	9.3547	
120	14.6882/8.12086	3.56957955	17.3331	3.060561	17.062			3.606	17.47	

of α_{95} (high accuracy) could be becausei of those sampled directions not covering enough long period (thought to be at least about 10^4 years) to “average out” secular variation for giving a paleopole. That is to say, the smallest α_{95} s could get the greatest weights that they should not deserve.

Generally, weighting does not improve fit. In other words, Wt 0 is generaly the best. Wt 2 or 4 is not recommended, because they never have generated the best similarities (Table. 3), compared with other weighting methods. There is no general pattern about which weighting (of Wt 1–5) is better or worse. So weighting, for making a paleomagnetic APWP, is not absolutely necessary. However, there are some patterns about which weighting is better or worse for some specific continent. For example, Wt 3 prefers Australia (Table. 3). Wt 3 works fine with India. However, Wt 3 is not recommended for North America.

3.2.4 Question: Why best and worst methods are not consistent?

3.2.4.1 Question: Why the picking method 21 is not among the best for 20 Myr binning and 10 Myr stepping with both space and shape tested? As shown in Fig. 8, both the picking methods 19 and 21 are among the best. However, the picking method 21 is not one of the best any more for 20 Myr binning and 10 Myr stepping with both space and shape tested (Fig. 15). Further, in fact, in Fig. 15, we can see the picking method 21 is still one of the best for North America (101) and Australia (801), but just not for India (501). However, even for India (501), the difference values (ranging 0.0483–0.0535) produced by the picking method 21 are still closer to the left bound of the one-standard-deviation interval 0.0359–0.1072 and relatively farther from the mean 0.074, which means the picking method 21 is still a relatively better one.

3.2.5 Question: Are there particular parts of the path that are more variable? Do different methods affect different parts of the path differently?

The results may highlight the trade-off between more data diluting the effect of outliers, and fewer but ‘better’ data being more easily affected by a bad point that gets through the filters (Fig. 9, Fig. 10 and Fig. 11).

3.2.6 Question: Do time window size and step affect the results?

A balance needs to be made between having windows that are too wide and steps that are too long which will smooth the data so much we miss actual details in the APWP (e.g. those 20 Myr window 10 Myr step paleomagnetic paths in Fig. 16 and even 30/15 Myr window and step; Table. 9 and Fig. 25) and windows that are too narrow and steps that are too short which introduces noise by having too few poles in each window (e.g. 2 Myr window 1 Myr step; Table. 9 and Fig. 25). There is a dependence here on data density: higher density allows smaller windows/steps (this is one of the things we want to test with selective data removal mentioned in Chapter 4). Fitting curves by moving averaging change with different time window lengths and time increment lengths (i.e. steps) (e.g., the similarity of the pair in Fig. 16f is improved a bit compared to Fig. 12f). A variety of ways of binning the data (here 30–2 Myr window size and half of the size as step) are being tested to see which one produces the better and more appropriately smoothed fit.

For Pk 2 (AMP with “ α_{95} /Age range” no more than “ $15^\circ/20^\circ$ ”), the 20/10 Myr bin/step methods always generate better similarities than the 10/5 Myr ones (e.g. Fig. 15a versus Fig. 8a).

Interestingly, only for North America, the 10/5 Myr bin/step methods generally and unexceptionally produce better similarities than the 20/10 Myr methods do (Fig. 24), which mainly depends on the picking methods. Note that as mentioned in the Supplementary materials and also Chapter 4, there are 135, 75 and 99 paleo-poles that compose of 120–0 Ma APWPs of North America, India and Australia respectively. Does the reason could be because of the relatively larger number of paleo-poles for North America? Since theoretically for each sliding window, the more “bad” paleo-poles it contains, the worse similarity we should obtain. In the contrary, the less paleo-poles the window contains, the weaker the effect of averaging out “bad” poles’ influence would be. So is there a threshold number of paleo-poles for making an paleomagnetic APWP? For example, for making a 120–0 Ma APWP, do the results indicate the best number of paleo-poles we need should be some value between 99 and 135? We did the test in Chapter 2 on the 530–0 Ma paleomagnetic APWP using the AMP method, and we did find that larger windows and steps bring the paleomagnetic APWPs closer to the reference path. Here another test will be implemented as follows. With the results from the 10/5 and 20/10 bin/step together, 2/1, 4/2, 6/3, 8/4, 12/6, 14/7, 16/8, 20/10, 24/12 and 30/15 Myr

Table 8. Highest and lowest values for the same variable are highlighted in red and green respectively.

FHM predicted for 101	Pk 5 + Wt 0			Pk 5 + Wt 1			Pk 5 + Wt 2			Pk 5 + Wt 3			Pk 5 + Wt 4			Pk 0 + Wt 5		
	ds	ds	ds	ds	ds	ds	ds	da	ds	ds	ds	ds	ds	ds	ds	ds	ds	
Age (Ma)	A95 (°)	Pmag A95 (°)	Dist (°)	Pmag A95 (°)	Dist (°)	Pmag A95 (°)	Dist (°)	Pmag A95 (°)	Dist (°)	Age (Ma)	Diff (°)	Pmag A95 (°)	Dist (°)	Pmag A95 (°)	Dist (°)	Pmag A95 (°)	Dist (°)	
0	0	7.458	2.058	7.4575	2.387	8.027	3.539	7.598	3.885	10-15-20	126.5907	7.7351	3.624	7.67	3.3909			
5	1.56039/0.87367	7.3814	2.624	7.3814	2.995	7.887	3.876	6.8	4.3453			7.515	3.9475	7.29	3.74			
10	2.89214/1.58743					15.208	3.402	8.602	5.79	dl		16.783	4.2726	14.598	2.87			
15	2.575/1.63303	12.421	9.2	12.4213	9.077	12.384	9.4	2.0857	11.805	Age (Ma)	Diff (°)	12.3843	9.4	10.153	10.71			
20	3.16077/2.20094									10-15	13.52							
25	4.96061/2.2183			6.463	6.2097			6.336	6.873	15-20	14.68	6.435	6.68					
30	3.39692/2.37114							6.678	6.458					6.97	5.724			
50	7.15565/3.22656												3.34	4.51				
55	7.17564/4.28065												5.44	6.2034				
65	7.37969/4.60029						7.6917	7.214										
100	5.79659/5.36693												8.275	7.013				
115	9.27023/5.16012						5.1	12.92	8.5	11.704			5.92	12.355	9.452	11.276		
120	14.6882/8.12086	11.4266	13.435	11.4266	13.0664	4.7143	16.543	7.728	15.258				4.509	17.112	7.943	15.511		

bin/step will be used to generate paleomagnetic APWPs for North America, India and Australia to see which one would make the paleomagnetic APWP closest to the reference paths. Will the similarities they generate be generally worse than those the 10/5 Myr bin/step generates? Or will they be better first and then worse than those the 10/5 Myr bin/step generates when the bin/step sizes increase up to 20/10 Myr? For the best results (Table. 9), as expected, AMP needs wider sliding window and step to get close to the reference while APP does not (Fig. 25). Even the best sizes of sliding window and step are assigned for AMP, the results from APP are still much better than those from AMP. Picking methods (directly related to N) are still the key influence factor of choosing a better sliding window size and step size of moving averaging, although weighting methods are also important.

3.2.6.1 What to expect is the difference values for 20/10 window/step should be generally lower than those for 10/5 window/step, which further could result in more best methods and less worst methods.

3.2.6.2 The results are summarised in Table. 10, Table. 9 and Fig. 25.

3.2.6.3 Conclusions:

If AMP has to be used, better results can be obtained through using large sizes of sliding window and step, commonly more than 24/12 Myr. In addition, we should be cautious when Wt 3 is used with AMP.

APP is recommended, not only because the temporal uncertainty is incorporated into the algorithm but also the results from APP are not as that sensitive as AMP to the changes of sliding window and step sizes. In fact, for APP the results from different window and step sizes are much more stable than those from AMP (Fig. 25). This means we actually do not need to worry about what sizes should be chosen for the sliding window and step when we use APP method.

3.2.7 Question: Is there any difference between choosing FHM and MHM as the reference path?

The differences (Fig. 26) are all lower than 0.041, and most are less than 0.01. That indicates the uncertainties of paleomagnetic mean poles are relatively large enough to make uses of Fixed and

Moving Hotspot Models as references not quite differentiated. So far choosing Fixed or Moving model is not a priority.

However, through looking at the signs of these differences (Fig. 26), for North America (Fig. 26a), FHM and the North America–Nubia–Mantle plate circuit derived path is a better reference option in general, while for both India (Fig. 26b) and Australia (Fig. 26c), MHM and the related plate circuit derived path is a generally better choice.

4 FINAL CONCLUSIONS

From the perspective of the general similarities between those paleomagnetic APWPs and the hot spot model and sea-floor spreading model predicted APWPs, GAD hypothesis is proved valid for at least the last 120 Myr.

4.1 Universal Rules of Ways of Processing Paleomagnetic Data:

Although effects of filters (all the picking methods except Pk 1) have marginal change in reducing N (precision going down), some filters do improve the similarity score generally, for example, Pt 11 (APP without redbed-derived data) and 13 (APP with redbed-derived data corrected) are better than Pt 0 for all the three continents.

APP is better than AMP for making paleomagnetic APWPs, for both kinds of situations when there are lots of data (APP even better, e.g. for North America and Australia) and not much data (APP still a better option, e.g. for India; Table. 3). APP with filters (Pt 3, 5, 7, 9, 11, ..., 27) is generally worse than APP without any filter (Pt 1).

Weighting is actually not affecting, and in some cases makes it much worse, for example, likely worse for the combined methods of Wt 3 and AMP.

APP itself helps incorporate temporal uncertainty into the algorithm. With the bootstraps test helping incorporate spatial uncertainty into the algorithm together, both spatial and temporal uncertainties are successfully considered in APP methods.

4.2 Conditional Rules of Ways of Processing Paleomagnetic Data:

Picking method no. 16 (AMP with data from old studies) is not recommended, e.g. before 1983) for generating a paleomagnetic APWP.

Table 9. Equal-weight 120–0 Ma CPDs for the three continents' paleomagnetic APWPs compared with their FHM predicted APWPs. The best are in dark green and underlined, second best in green and third in light green.

Window, Step (size in Myr)	N America Pk 0						N America Pk 1					
	Wt 0	Wt 1	Wt 2	Wt 3	Wt 4	Wt 5	Wt 0	Wt 1	Wt 2	Wt 3	Wt 4	Wt 5
2, 1	0.2864	0.26	0.2877	0.2801	0.2632	0.2619	0.00747	0.00776	0.01688	0.01283	0.0372	0.0097
4, 2	0.08205	0.08422	0.1034	0.09653	0.10295	0.09863	0.0064145	0.00711	0.0182	0.011423	0.03606	0.00909
6, 3	0.06657	0.06817	0.06788	0.06229	0.08557	0.07634	0.00627	0.007797	0.01754	0.01254	0.02113	0.008955
8, 4	0.0614	0.0772	0.0653	0.06214	0.0903	0.0646	0.006	0.01099	0.01576	0.01399	0.02027	0.01271
10, 5	0.0349	0.0458	0.0486	0.046	0.0488	0.0344	0.0059	0.0062	0.0136	0.0151	0.0153	0.0121
12, 6	0.0318	0.0316	0.0325	0.0298	0.0323	0.0299	0.0087	0.009	0.017	0.0126	0.0191	0.0145
14, 7 (119–0 Ma path)	0.0367	0.0348	0.0353	0.0369	0.0319	0.0352	0.00996	0.01198	0.023922	0.013	0.0203	0.0113
16, 8	0.0493	0.0492	0.0496	0.0486	0.0477	0.0485	0.0114	0.0117	0.023902	0.0123	0.0207	0.0129
20, 10	0.0497	0.0536	0.0538	0.0557	0.0517	0.0526	0.014	0.0174	0.029	0.0181	0.027	0.0156
24, 12	0.0274	0.0304	0.0327	0.0324	0.0315	0.0313	0.0138	0.0143	0.0221	0.0191	0.0203	0.0192
30, 15	0.0345	0.0298	0.0317	0.0307	0.03402	0.0307	0.0174	0.01797	0.0276	0.02402	0.0252	0.02414
Window, Step (size in Myr)	India Pk 0						India Pk 1					
	Wt 0	Wt 1	Wt 2	Wt 3	Wt 4	Wt 5	Wt 0	Wt 1	Wt 2	Wt 3	Wt 4	Wt 5
2, 1	0.249664	0.249951	0.249985	0.257479	0.250294	0.249702	0.0550264	0.0558274	0.0653912	0.0616122	0.0725822	0.0583844
4, 2	0.193452	0.202757	0.193438	0.241803	0.210005	0.194226	0.0570142	0.059523	0.0670142	0.0622828	0.0672158	0.0601206
6, 3	0.173961	0.174758	0.173975	0.229325	0.185139	0.175174	0.0578869	0.0588201	0.0639928	0.0627534	0.0672496	0.0624135
8, 4	0.154308	0.149658	0.151995	0.174118	0.165387	0.150164	0.0576095	0.0584774	0.0667115	0.0613883	0.0686417	0.05941
10, 5	0.1839	0.1831	0.1909	0.2586	0.1958	0.1839	0.0545	0.0554	0.0621	0.0589	0.066	0.06
12, 6	0.108924	0.105626	0.114955	0.118308	0.121497	0.105872	0.0598967	0.0629866	0.0657908	0.0644987	0.0695085	0.0629451
14, 7 (119–0 Ma path)	0.112537	0.112885	0.126554	0.116174	0.132359	0.120914	0.0547932	0.0502588	0.0579931	0.060018	0.0654112	0.0582519
16, 8	0.104461	0.1241	0.124436	0.110942	0.119599	0.118336	0.0517351	0.0528129	0.0551881	0.056389	0.0574883	0.0550421
20, 10	0.1052	0.1015	0.1198	0.1096	0.1174	0.1072	0.0492	0.0501	0.0585	0.053	0.0577	0.052
24, 12	0.0531434	0.053561	0.056986	0.0574692	0.0555799	0.0553047	0.0519257	0.0459949	0.048681	0.0557455	0.0569792	0.0545615
30, 15	0.0737324	0.0754578	0.0775947	0.0575459	0.0565421	0.056635	0.0523614	0.0519862	0.054158	0.0563985	0.0555998	0.0543501
Window, Step (size in Myr)	Australia Pk 0						Australia Pk 1					
	Wt 0	Wt 1	Wt 2	Wt 3	Wt 4	Wt 5	Wt 0	Wt 1	Wt 2	Wt 3	Wt 4	Wt 5
2, 1	0.589222	0.522048	0.523031	0.558687	0.544474	0.54495	0.00612207	0.00611382	0.025182	0.00696989	0.0355568	0.00581442
4, 2	0.268779	0.247297	0.251993	0.271535	0.353676	0.272178	0.00504377	0.00625515	0.029305	0.00753837	0.035213	0.00689346
6, 3	0.0918779	0.0956757	0.0944798	0.0852725	0.100829	0.0922681	0.00488333	0.00536227	0.0208704	0.00654597	0.0236738	0.00644947
8, 4	0.0448139	0.0972008	0.0572843	0.0859725	0.0624494	0.0455424	0.0426485	0.00419086	0.0272409	0.00670754	0.0303	0.00757145
10, 5	0.0326	0.039	0.0509	0.0305	0.0601	0.031	0.0045	0.0048	0.0199	0.0058	0.0259	0.0089
12, 6	0.0288423	0.0376182	0.0594042	0.0279421	0.0569207	0.0285109	0.00692362	0.0072455	0.0220767	0.0102027	0.0220214	0.0111216
14, 7 (119–0 Ma path)	0.0287639	0.0286588	0.0480163	0.0279289	0.0480853	0.0287173	0.00343606	0.00354029	0.0137145	0.00306601	0.0137356	0.0105205
16, 8	0.0299621	0.0398455	0.0556817	0.0293245	0.0512076	0.0299436	0.0115828	0.0057367	0.0201674	0.00643355	0.0162713	0.0123646
20, 10	0.0278	0.04	0.0612	0.0271	0.0547	0.0393	0.0079	0.0076	0.0197	0.0129	0.014	0.0106
24, 12	0.024437	0.0241335	0.0465703	0.0237562	0.0397431	0.036984	0.00584733	0.00637881	0.00599371	0.0056679	0.00547471	0.00516432
30, 15	0.0200862	0.0204561	0.0844176	0.0173471	0.0365786	0.034705	0.00412276	0.00444694	0.00448627	0.00397289	0.00369699	0.00387337

4.3 Summary

According to the results we have from the three continents, North America, India and Australia, using the similarity measuring tool developed in Chapter 2, it is recommended that APP should be used to select the input paleopoles. According to all the paleomagnetic data we have from the three continents, the results from any size for sliding window and step are interestingly and extremely close to each other when the APP method is used, compared with the results from the AMP method (Table 9 and Fig. 25). So any size for binning and stepping is ok when APP is used. Then filtering is actually not necessary. However, some filtering methods (e.g. Picking

No. 5, 7 (igneous-derived), 11, 13 (nonredbeds or corrected redbeds derived), 19, 21 (non-local-rotation/reprinted or corrected-rotation derived) and 25 (non-superseded data derived)) are fine too and will not give worst or worse results than the other filtering methods (i.e. the left Picking methods). With APP used, weighting is actually not necessary either. If a weighting has to be used, Weighting No. 1 (related to the number of paleomagnetic sampling sites and observations) is generally better than the other four given weighting methods (Fig. 27).

If AMP has to be used, relatively wide sliding window and step are needed. According to our tests, more than 20/10 Myr is recommended. In addition, AMP works relatively better with igneous-

Table 10. Consistency check on comparisons of 20/10 window/step and 10/5 window/step. Notes: E: expected; UE: unexpected.

Comparisons				Consistency of Best			Consistency of Worst			If Difference Values for 20/10 Bin/Step Are Lower (Y/N)						
10/5	20/10	Y/N	Special Case(s)	Notes	Y/N	Special Case(s)	Notes	Mean	Median	Maximum	Minimum	All	If No, Unexpected Case (s)	Notes		
FHM																
Fig. 8a	Fig. 15a	Y	–	Same: Picking no. 1, 4, 5, 7, 11, 13, 15, 19, 21 and 25	N	Picking no. 2, 5, 7, 22 and 26 for 10/5 (E); 0, 10, 12, 20 and 24 for 20/10 (UE)	Same: Picking no. 8, 14, 16 and 18	N	Y	Y	N	N	(0,1,8,10,11,12,18,19,20,21, 24,25)(0-5) (3,4,14),(0-3,5) (5,7,13,15),(0-2,4,5) 9,(0,2,4,5) 23,(2,3) 27,(0-2); account for 29/42			
Fig. 8b	Fig. 15b	Y	2 more best: Picking no. 9 and 21 only for 10/5 (UE)	Same: Picking no. 4, 5, 6, 7 and 19	N (almost Y)	Picking no. 8 for 10/5 (E); 23 and 27 for 20/10 (UE)	Same: Picking no. 0, 2, 10, 12, 16, 18, 20 and 24	Y	Y	Y	Y	N	15,3 19,(0,1,3,5) 21,0 (22,26),(0-5) 23,(1,3,4) 27,(1,3); account for 23/168			
Fig. 8c	Fig. 15c	Y	–	Same: Picking no. 1, 11, 13, 17, 19, 21 and 25	Y	1 more worst: Picking no. 6 only for 10/5 (E)	Same: Picking no. 2, 14, 16, 22 and 26	Y	Y	Y	Y	N	0,(1,2,5) (1,4,18,25),(0,1,3,5) 7,0 (8,10,20),(2,5) 9,1 11,3 (12,24),(1,2) 13,(0,1,3) (14,22,26),(0-2,4,5) (16,17),(0-5) 19,5 21,(0,1,5); account for 33/84			
MHM																
Fig. 17a	Fig. 19a	N	Picking no. 1, 13, 21 and 25 for 10/5 (UE); 22 and 26 for 20/10 (E)	Same: Picking no. 4, 5, 7, 11, and 15	Y	2 more worst: Picking no. 17, 22 only for 10/5 (E)	Same: Picking no. 0, 8, 10, 12, 14, 16, 18, 20 and 24	N	Y	N	N	N	(0,4,8,10,11,12,14,15,18, 20,24,25),(0-5) (1,19,21),(0,1,5) (5,7),(0-2,4,5) 6,1 (9,13),(0,1,3,5) 23,(1,3) 27,(1,3,5); account for 53/84			
Fig. 17b	Fig. 19b	Y	2 more best: Picking no. 22 and 26 only for 10/5 (UE)	Same: Picking no. 4, 5, 6, 7 and 19	N	Picking no. 8, 20 for 10/5 (E); 23 and 27 for 20/10 (UE)	Same: Picking no. 0, 2, 10, 12, 16, 18 and 24	Y	Y	Y	Y	N	1,5 (4,15),3 (19,23,27),(0-5) 21,(0,1,3,5) 22,(1,4) 25,(3,5) 26,4; account for 5/28			
Fig. 17c	Fig. 19c	Y	3 more best: Picking no. 15, 23 and 27 only for 20/10 (E)	Same: Picking no. 1, 11, 13, 17, 19, 21 and 25	Y	–	Same: Picking no. 2, 14, 16, 22 and 26	Y	Y	Y	N	N	(0,24),(1,2,5) (1,11,13,18,19,21,25),(0,1,3,5) 4,(0,3,5) (5,7),3 (8,17),(0-3,5) 12,(1-3,5) 20,(1,2,4,5)			

derived data (i.e. Picking No. 4 and 6), which indicates that if we have fewer data, these data need to be better in quality (Fig. 27).

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REFERENCES

- Besse, J. & Courtillot, V., 2002. Apparent and true polar wander and the geometry of the geomagnetic field over the last 200 Myr, *J. Geophys. Res.*, **107**, 2300.
- Christeson, G. L., Van Avendonk, H. J. A., Norton, I. O., Snedden, J. W., Eddy, D. R., Karner, G. D. & Johnson, C. A., 2014. Deep crustal structure in the eastern Gulf of Mexico, *J. Geophys. Res. Solid Earth*, **401**, 183–195.
- Domeier, M., van der Voo, R. & Denny, F. B., 2011. Widespread inclination shallowing in Permian and Triassic paleomagnetic data from Laurentia: Support from new paleomagnetic data from Middle Permian shallow intrusions in southern Illinois (USA) and virtual geomagnetic pole distributions, *Tectonophysics*, **511**, 38–52.
- Domeier, M., Van der Voo, R., Tomezzoli, R. N., Tohver, E., Hendriks, B. W. H., Torsvik, T. H., Vizan, H. & Dominguez, A., 2011. Support for an “A-type” Pangea reconstruction from high-fidelity Late Permian and Early to Middle Triassic paleomagnetic data from Argentina, *Journal of Geophysical Research: Solid Earth*, **116**, B12114.
- Fisher, R. A., 1953. Dispersion on a sphere, *Proc. Roy. Soc. London Ser. A*, **217**, 295–305.
- Gaina, C., van Hinsbergen D. J. J. & Spakman, W., 2015. Tectonic interactions between India and Arabia since the Jurassic reconstructed from marine geophysics, ophiolite geology, and seismic tomography, *Tectonics*, **34**, 875–906.
- Jacox, E. H. & Samet, H., 2007. Spatial Join Techniques, *ACM Trans. Database Syst.*, **32**, 7.
- King, R. F., 1955. The remanent magnetism of artificially deposited sediments, *Geophys. Suppl. Mon. Not. Roy. Astron. Soc. Geophys. Left.*, **1**, 115–134.
- Müller, R. D., Royer, J. Y. & Lawver, L. A., 1993. Revised plate motions relative to the hotspots from combined Atlantic and Indian-Ocean hotspot tracks, *Geology*, **21**, 275–278.
- McElhinny, M. W. & Lock, J., 1996. IAGA Paleomagnetic Databases with access, *Surveys in Geophysics*, **17**, 575–591.
- Müller, R. D., Royer, J. Y., Cande, S. C., Roest, W. R. & Maschenkov, S., 1999. New constraints on the Late Cretaceous/Tertiary plate tectonic evolution of the Caribbean, *Sedimentary Basins of the World*, **4**, 33–59.
- McQuarrie, N. & Wernicke, B. P., 2006. An animated tectonic reconstruction of southwestern North America since 36 Ma, *Geosphere*, **1**, 147–172.
- O’Neill, C., Müller, R. D. & Steinberger, B., 2005. On the uncertainties in hot spot reconstructions and the significance of moving hot spot reference frames, *Geochem. Geophys. Geosyst.*, **6**, Q04003.
- Pisarevsky, S. A., 2005. New edition of the Global Paleomagnetic Database, *Eos. Trans. AGU*, **86**, 170.
- Schettino, A. & Scotese, C. R., 2005. Apparent polar wander paths for the major continents (200 Ma to the present day): a palaeomagnetic reference frame for global plate tectonic reconstructions, *Geophys. J. Int.*, **163**, 727–759.
- Torsvik, T. H., Smethurst, M. A., van der Voo, R., Trench, A., Abrahamson, N. & Halvorsen, E., 1992. Baltica. A synopsis of Vendian-Permian palaeomagnetic data and their palaeotectonic implications, *Earth Sci. Rev.*, **33**, 133–152.
- Torsvik, T. H. & Smethurst, M. A., 1999. Plate tectonic modelling: virtual reality with GMAP, *Comput. Geosci.*, **25**, 395–402.
- Tauxe, L. & Kent, D. V., 2004. A simplified statistical model for the geomagnetic field and the detection of shallow bias in paleomagnetic inclinations: Was the ancient magnetic field dipolar? *Geophys. Monogr. AGU*, **145**, 101–115.
- Torsvik, T. H., Müller, R. D., van der Voo, R., Steinberger, B., & Gaina, C., 2008. Global plate motion frames: Toward a unified model, *Rev. Geophys.*, **46**, RG3004.
- Torsvik, T. H., van der Voo, R., Preeden, U., Mac Niocaill, C., Steinberger, B., Doubrovine, P. V., van Hinsbergen, D. J. J., Domeier, M., Gaina, C., Tohver, E., Meert, J. G., McCausland, P. J. A. & Cocks, L. R. M., 2012. Phanerozoic polar wander, palaeogeography and dynamics, *Earth Sci.*

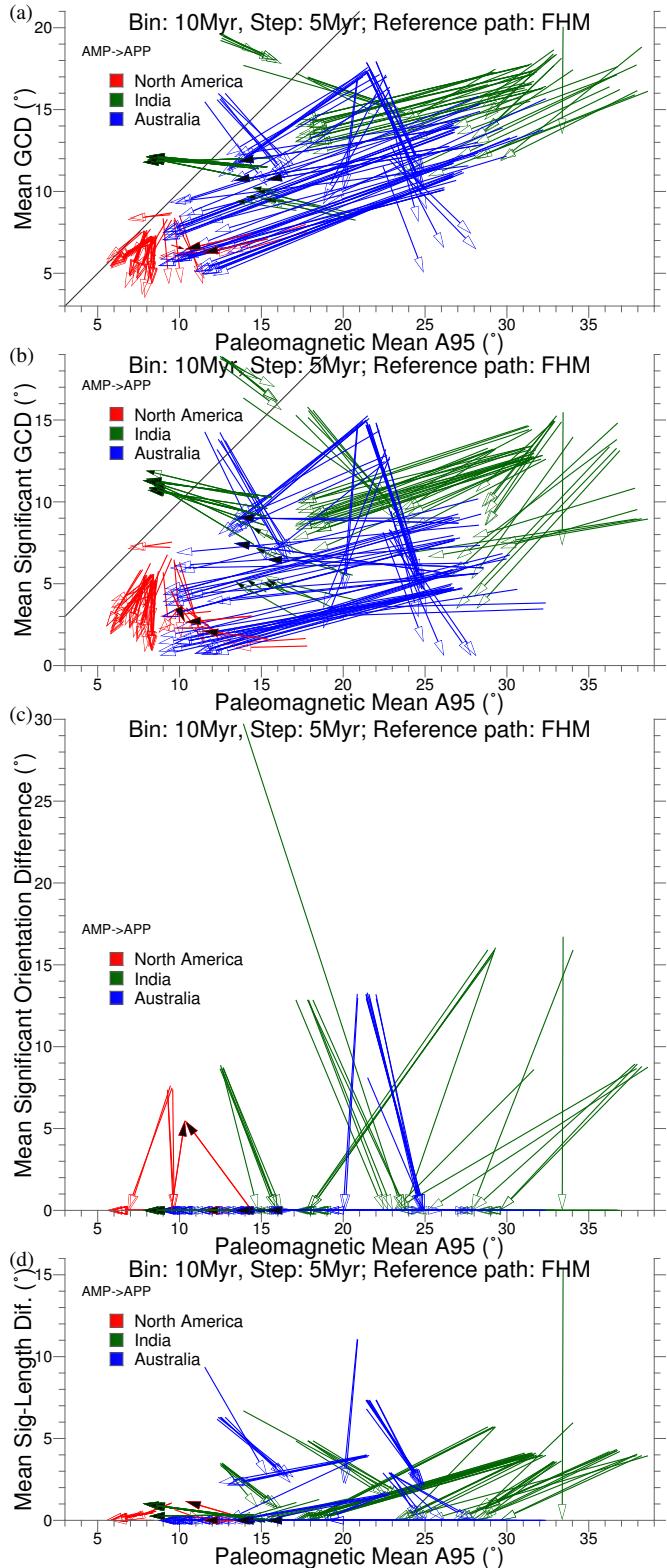


Figure 21. Paleomagnetic APWPs' mean A95 versus (a) “mean GCD”, (b) “mean significant GCD”, (c) “mean significant orientation difference”, and (d) “mean significant length difference” between paleomagnetic APWP and its corresponding FHM-and-plate-circuit predicted APWP. Starting points of the arrows are results from AMP, while ending points are from APP. Black color filled arrow heads are the small number of special cases of AMP derived equal-weight CPDs better than APP (see details in Fig. 8).

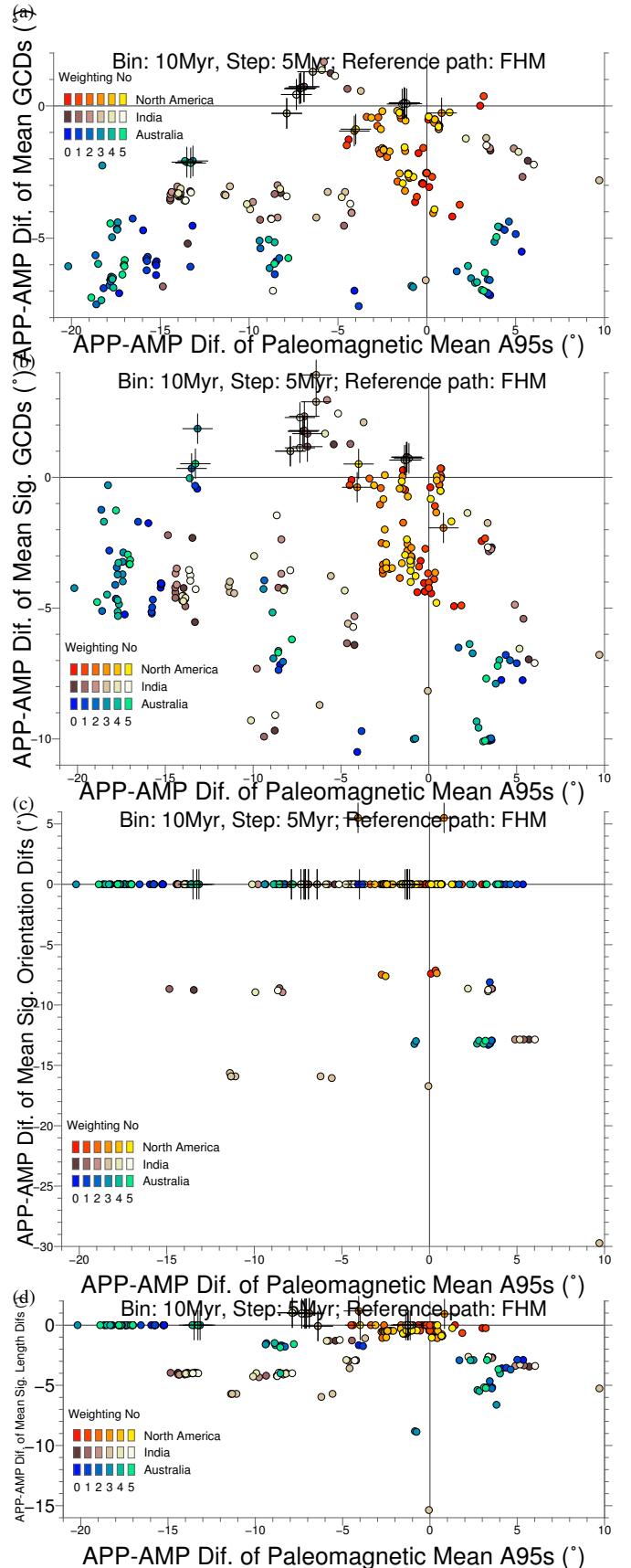


Figure 22. Differences of APP and AMP coordinates shown in Fig. 21. Crosses locates the small minority cases of AMP derived equal-weight CPDs better than APP (see details in Fig. 8).

- Rev.*, **114**, 325–368.
- Tauxe L., Banerjee S.K., Butler R.F. & van der Voo R., 2018. *Essentials of Paleomagnetism*, 5th web edn, Available on line
- van der Voo, R., 1990. The reliability of paleomagnetic data, *Tectonophysics*, **184**, 1–9.
- Wessel, P., Smith, W. H. F., Scharroo, R., Luis, J. & Wobbe, F., 2013. Generic Mapping Tools: Improved version released, *Eos. Trans. AGU*, **94**, 409–410.
- Young, A., Flament, N., Maloney, K., Williams, S., Matthews, K., Zahirovic, S. & Müller, D., 2019. Global kinematics of tectonic plates and subduction zones since the late Paleozoic Era, *Geosci. Front.*, **10**, 989–1013.

APPENDIX A: CONSTRAIN PALEOPOLES FOR A CERTAIN TECTONIC PLATE

A polygon can be drawn around a set of paleomagnetic data, whose sampling sites we believe belong to a specific plate or rigid block. Then the *Spatial Join* technique (Jacox & Samet 2007) helps join attributes from the polygon to the paleomagnetic data based on the spatial relationship allowing data within this polygon to be extracted from the whole raw large dataset without splitting a subset just for a specific plate. That allows us to quickly select subsets of the database based on geographic constraints just as easily as for age. Of course, the boundary of this polygon must be reasonably along a tectonic boundary. Regions like those close to the plate boundaries are usually tectonically active (e.g. local rotations), so we should also be careful when we deal with the paleopoles derived from this type of locations.

A1 120–0 Ma North America

The data-constraining polygons are from the recently published plate model (Young et al. 2018) (Fig. A1). Plate ID 101 polygon in the recently published Plate Model (Young et al. 2018), including its children 108 (Avalon/Acadia block) and 109 (Piedmont block) polygons for 120–0 Ma, is used to select the sampling sites of the paleopoles for North America. According to the plate model rotation data (Young et al. 2018), 108 is fixed to 101 during the geologic period from Cretaceous to the present day. 109 is also fixed to 101 since about 300 Ma (Christeson et al. 2014). Then in order to be compared with the FHM (120–0 Ma) (Müller et al. 1993; Müller et al. 1999), the paleopoles with age ranging 120–0 Ma are further selected through constraining the lower magnetic age “LOMAGAGE ≤ 135 ” (here it is not 120 but 135, because for the lower resolution case when the window length is 30 Myr, the Age Position Picking method will include those data with their lower magnetic age between 120 Ma and 135 Ma). In addition, the RESULTNO=6007 dataset should also be included according to a published plate kinematic model (McQuarrie & Wernicke 2006) with a relatively higher resolution of polygons and rotations, although the dataset is in the PlateID=178 polygon. In the end, 193 datasets in total are extracted (both white circles and red triangle-inside-circles in Fig. A1).

Also based on this model of southwestern North America since 36 Ma (McQuarrie & Wernicke 2006), part of the paleopoles constrained by the four small western terranes whose Plate IDs are also 101 (white circles in Fig. A1) in fact had gone through regional rotations and here are removed. However, the poles with age younger than 10 Ma located within the largest western 101 terrane (on the south of the smallest western 101 terrane; corresponding to the RANGE_ID=74 polygon in the model (McQuarrie & Wernicke

2006)) should be included. So finally 135 of the 193 datasets remain (Fig. A1). Spatially North American paleomagnetic data are mainly from the western and eastern margins of the plate.

A2 120–0 Ma India

Plate ID 501 polygons in the recently published Plate Model (Young et al. 2018) also include the two small polygons of the northern “Lesser Himalayan passive margin of Greater Indian Basin” and “Tethyan Himalayan microcontinent of Greater India” (Fig. A2). The polygons are used to select the sampling sites of the paleopoles for India (Fig. A2).

Based on the model of the tectonic interactions between India, Arabia and Asia since the Jurassic (Gaina et al. 2015) (Fig. A2), part of the paleopoles constrained by the north two small terranes whose Plate IDs are also 501 in fact had gone through regional rotations and here are removed. So finally 75 datasets are left (Fig. A2). Spatially Indian paleomagnetic data are more evenly distributed on the India plate than North American and Australian poles.

A3 120–0 Ma Australia

Plate ID 801 polygon in the recently published Plate Model (Young et al. 2018), including its children 675 (Sumba block) and 684 (Timor block) polygons for 120–0 Ma (Fig. A3), is used to select the sampling sites of the paleopoles for Australia. According to the plate model rotation data (Young et al. 2018), 675 and 684 are fixed to 801 during the geologic period from c.145 Ma to the present.

On the southeast of the main Australia plate (the blue polygon in Fig. A3), there is a triangle-shaped small polygon 850 (Tasmania block) which is fixed to 801 since c.100 Ma according to the (Young et al. 2018) rotation data. With that attribute, 805 contributes more data younger than c.100Ma for the later analysis. Ultimately the final 99 extracted datasets is shown in Fig. A3.

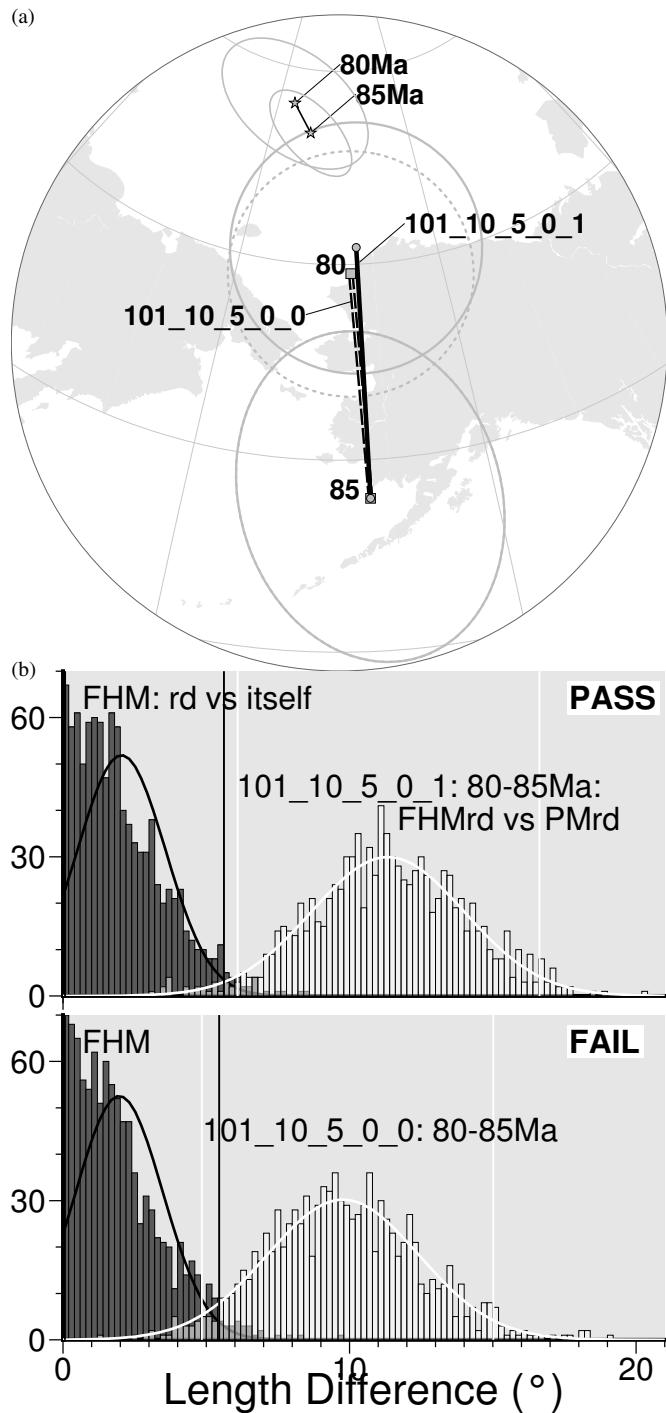


Figure 23. Significance testing on the 80–85 segment length difference between North American paleomagnetic APWPs derived from picking 0 and weightings 0 and 1 and the FHM predicted APWP (Fig. 5; Table 7). (a) The thin segment through stars is from FHM predicted path, and the bold solid and dashed are from paleomagnetic paths. (b) The results from Wt 1 are differentiated (Fig. 10a and Fig. 10b).

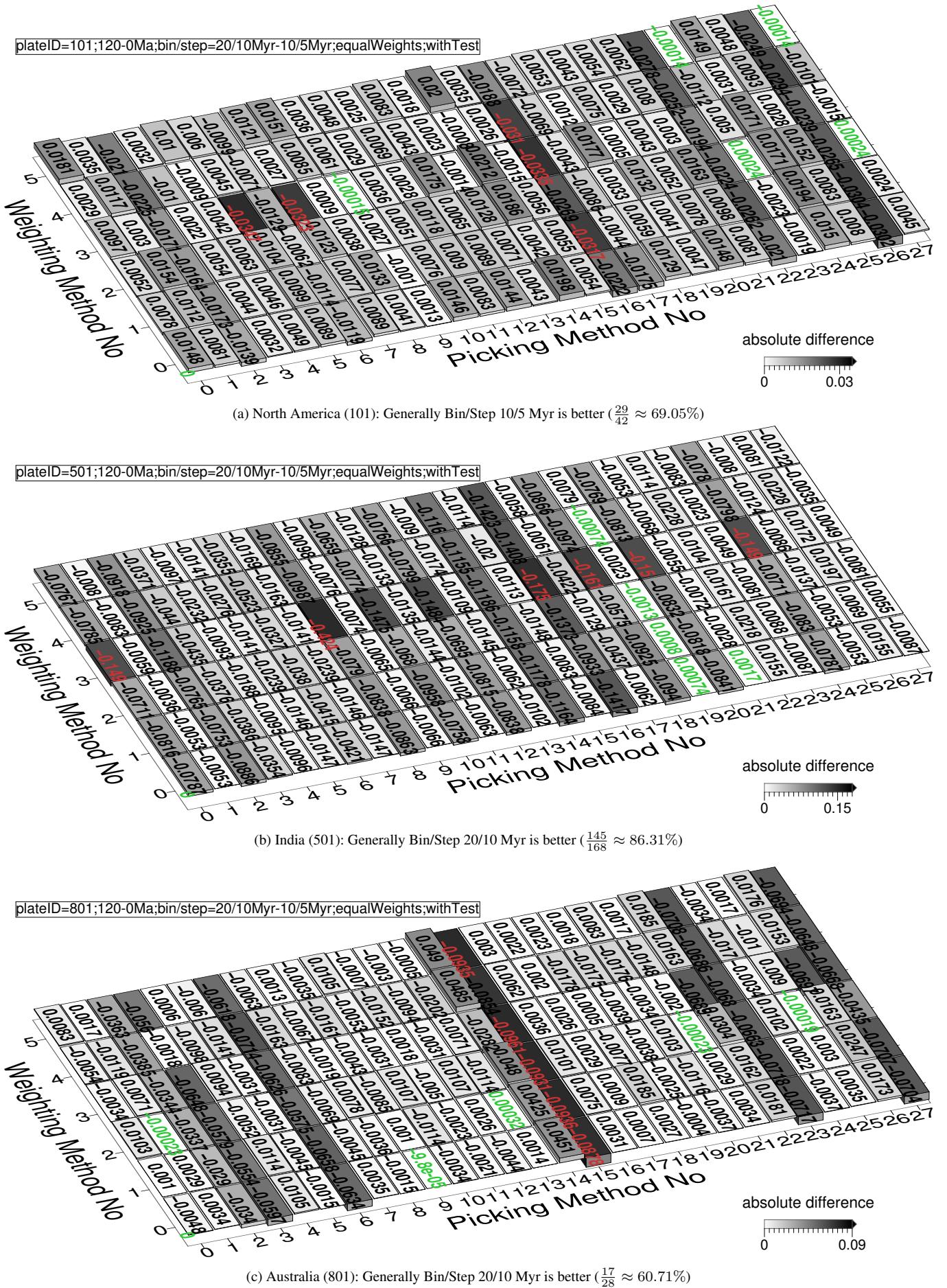


Figure 24. Differences between grids in Fig. 8 (10 Myr bin, 5 Myr step) and Fig. 15 (20 Myr bin, 10 Myr step). The absolute difference values less than 1.96-standard-deviation interval of the whole 168 values are labeled in green, more than 1.96-standard-deviation interval labeled in red.

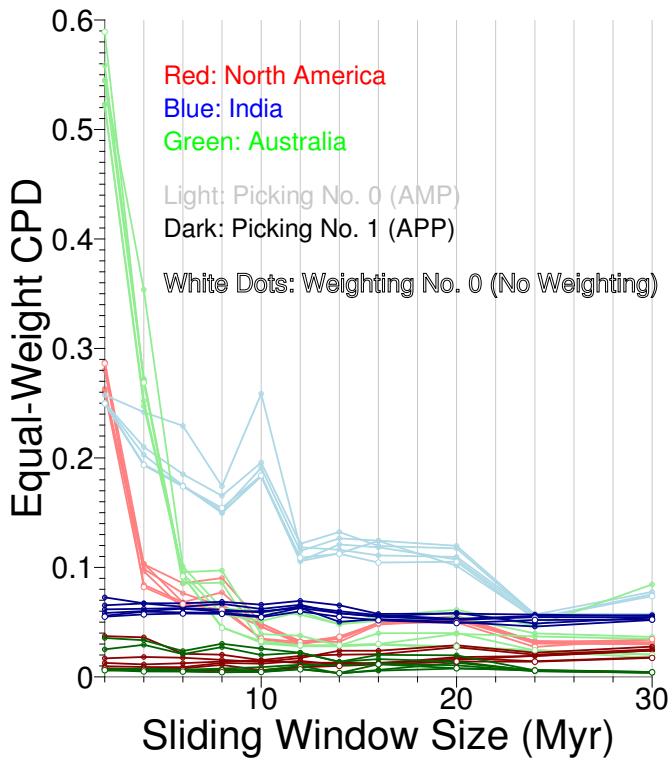


Figure 25. Plot of the data shown in Table. 9. Note that here the step size is always half of the sliding window size and the reference path is the FHM derived.

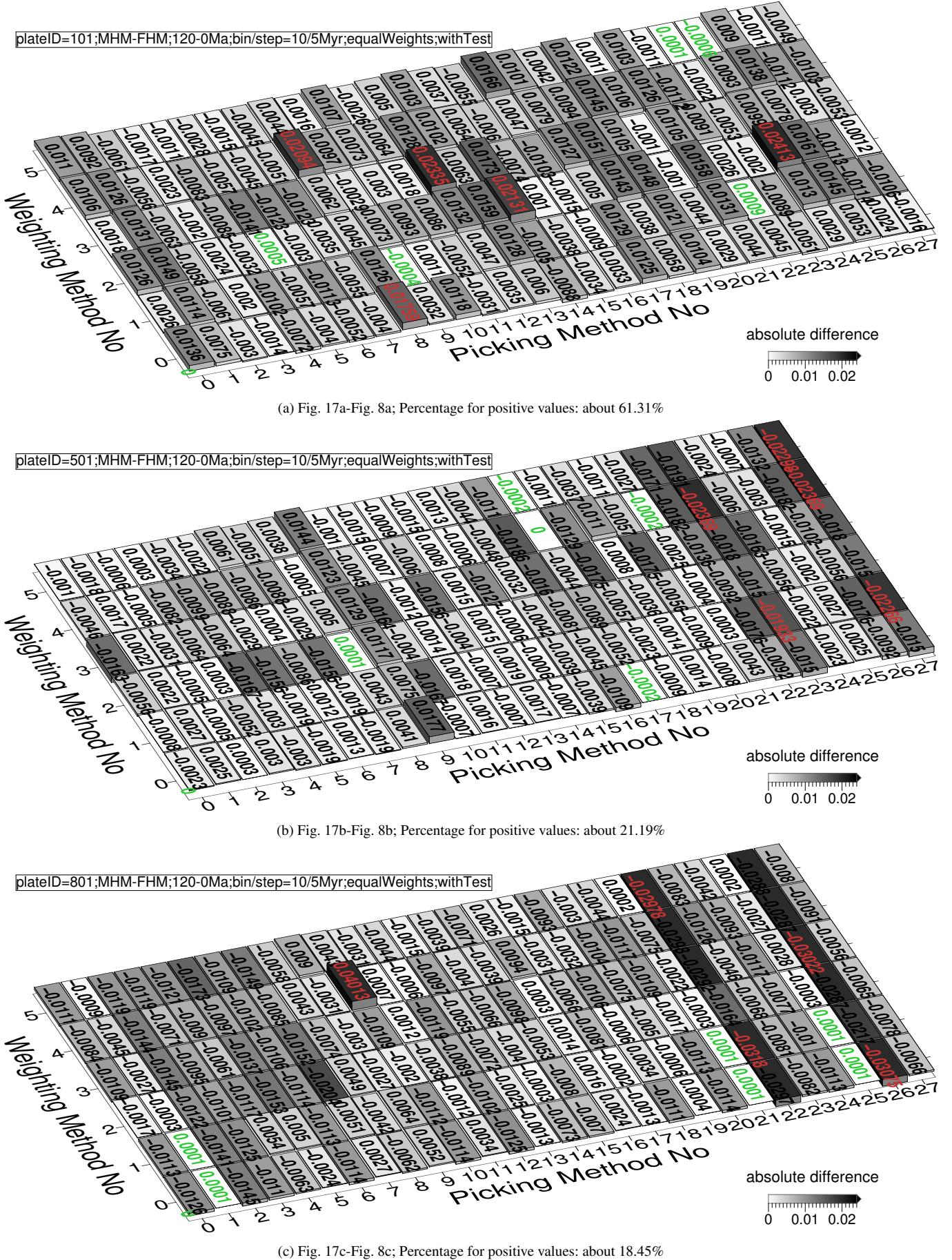


Figure 26. Differences between results from FHM (Fig. 8) and MHM (Fig. 17) as reference paths.

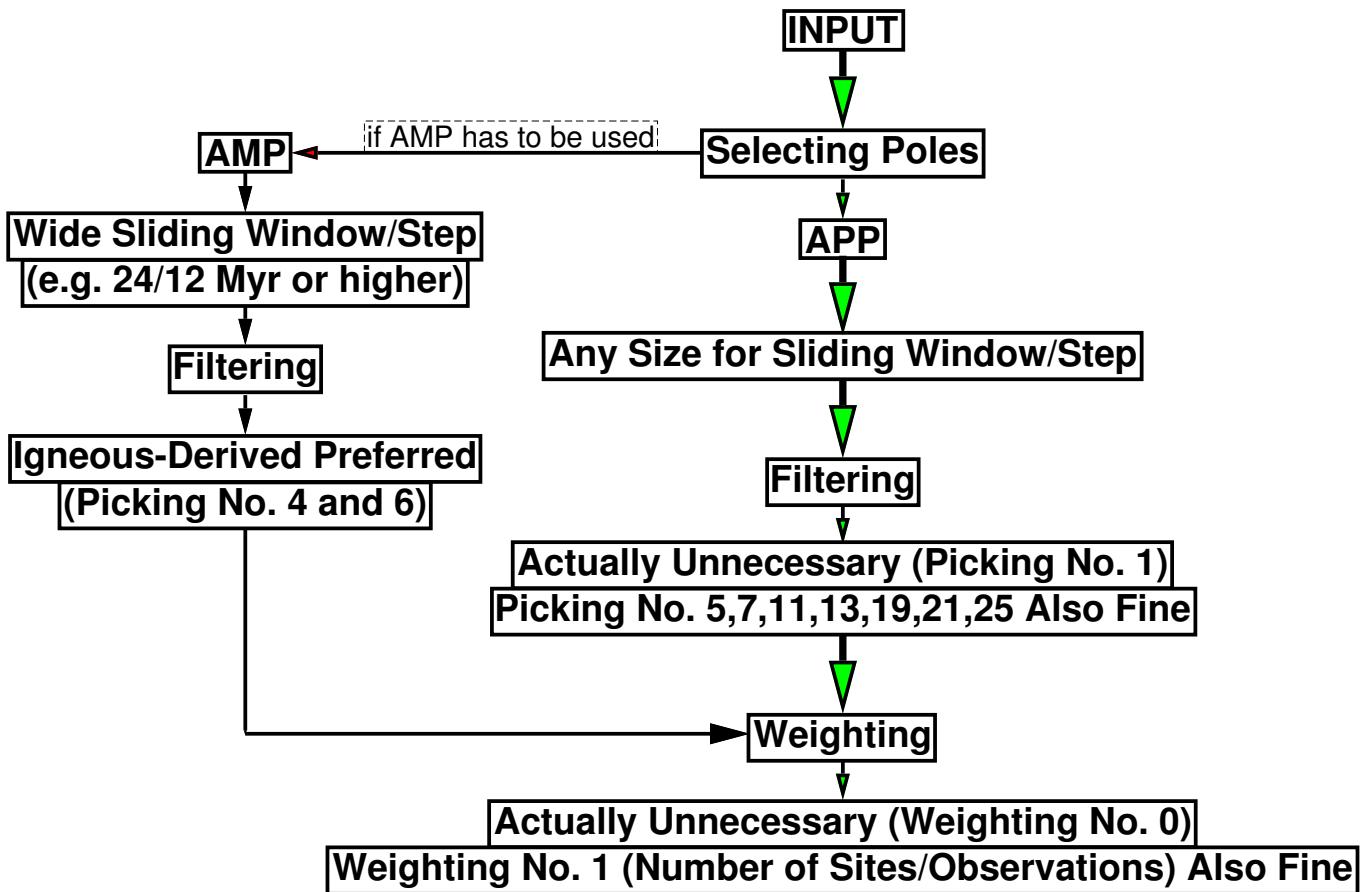


Figure 27. Flowchart for recommended procedure of processing paleomagnetic data.

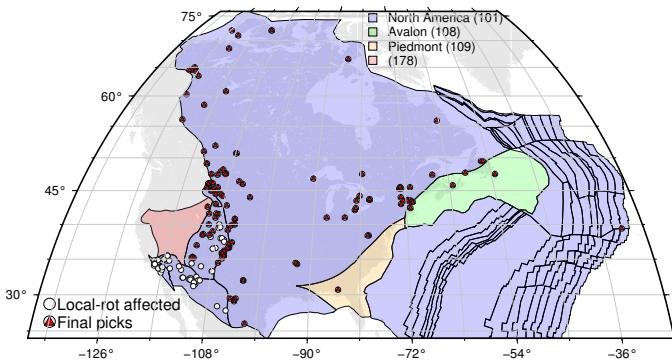


Figure A1. The final filtered datasets (red triangle-inside-circles) for later analysis on 120–0 Ma North America. Those poles that had been influenced by local tectonic rotations are shown as white circles.

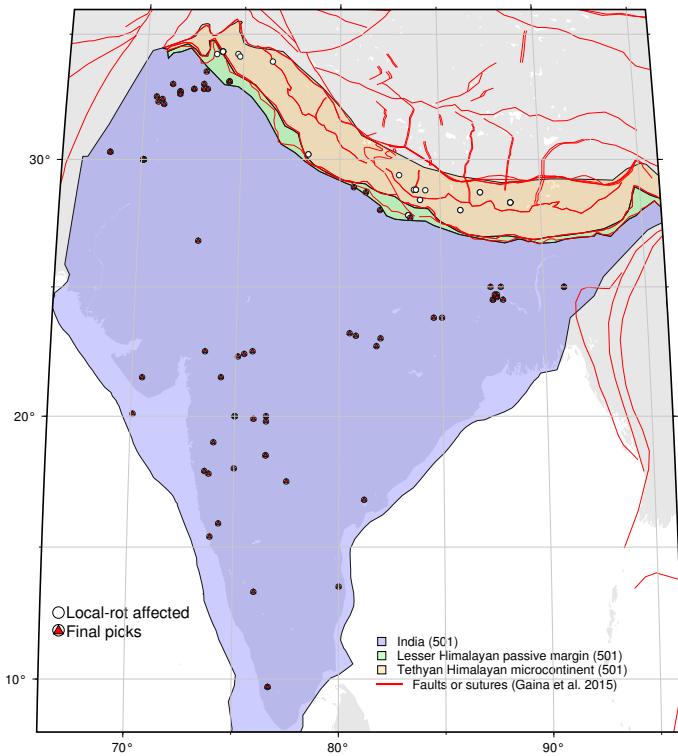


Figure A2. The final filtered datasets (red triangle-inside-circles) for later analysis on 120–0 Ma India. Those poles that had been influenced by local tectonic rotations are shown as white circles. The rifts, faults and detachments (red lines) around India are used to filter out those data that are influenced by local tectonic rotations.

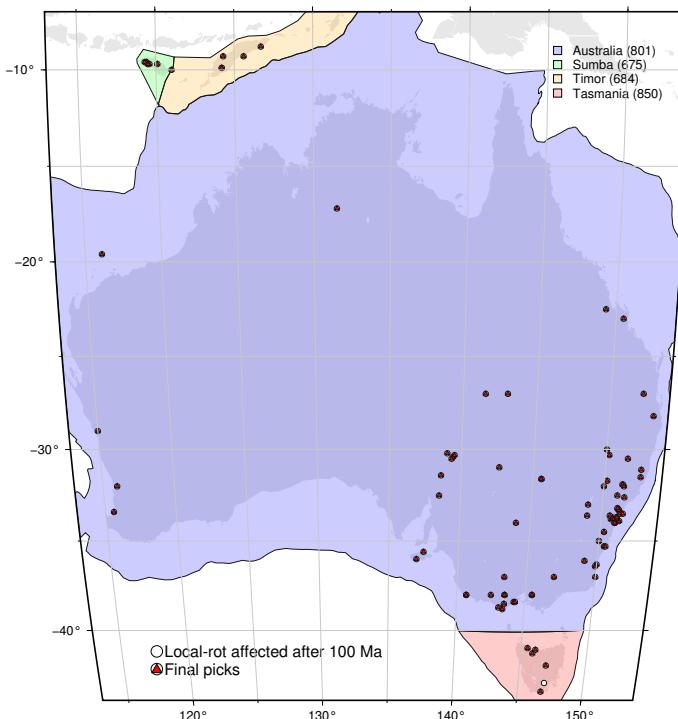


Figure A3. The final filtered datasets (red triangle-inside-circles) for later analysis on 120–0 Ma Australia. Those poles that had been influenced by local tectonic rotations are shown as white circles. The Plate ID 850 helps increase the amount of qualified datasets for 100–0 Ma.