Comparison of methods for the estimation of inert carbon suitable for initialisation of the CANDY model

Martina Puhlmann^{1,*}, Katrin Kuka² and Uwe Franko²

¹Leibniz-Centre for Agricultural Landscape Research (ZALF) e. V., Institute of Landscape Systems Analysis, Eberswalder Straße 84, 15374, Müncheberg, Germany; ²Department of Soil Science, UFZ – Centre for Environmental Research Leipzig-Halle, Theodor-Lieser-Str. 4, 06120, Halle, Germany; *Author for correspondence (e-mail: martina.puhlmann@zalf.de; fax: +49-334-3282334)

Received 25 August 2005; accepted in revised form 23 January 2006

Key words: CANDY, Carbon and nitrogen dynamics, Inert carbon, Model initialisation, Modelling

Abstract

Almost all soil organic carbon turnover models rely on a partitioning of total organic carbon into an inert and a decomposable pool. The quantification of these pools has a large impact on modelling results. In this study several methods to estimate inert carbon in soils, based either on total soil organic matter or physical protection, were assessed with the objectives of (1) minimising errors in carbon and nitrogen dynamics and (2) ensuring usability for sites with marked differences in site conditions. CANDY simulations were carried out by varying solely the method for calculating the size of the inert carbon pool used to initialise the model. Experimental data from Bad Lauchstädt and Müncheberg were used for the simulation. The data were made available for modellers at a workshop held at Müncheberg (Germany) in 2004. The results concerning not only carbon but also nitrogen dynamics were analysed by applying selected statistical methods. It was shown that even in short-term simulations model initialisation procedure may influence the simulation results considerably. Three methods of estimating inert carbon were identified as being the most appropriate. These methods are either based on soil texture or pore-space classes and therefore account for the physical protection of soil organic matter. Thus, physical protection seems to be of major importance. By extending the scope of the investigation into nitrogen dynamics, additional support for the applicability of a selected method was obtained.

Introduction

The simulation model CANDY (Carbon and Nitrogen Dynamics, Franko et al. 1995, 1997; Franko 1996, 1997) has been developed in order to provide information about carbon stocks in soils, organic matter turnover, N uptake by crops, leaching and water quality. A comparison of soil organic carbon models in 1997 showed CANDY

to be in the group with the best performance (Smith et al. 1997).

In most long-term organic matter models (Hansen et al. 1991; Franko et al. 1995; Coleman and Jenkinson 1996; Parton 1996, etc.), soil organic matter is partitioned in several pools with different turnover rates. In CANDY the decomposable soil organic matter can be subdivided into biologically active and stabilised soil organic

matter. The decomposable and the inert organic matter contribute to the total stock of carbon in the soil. The inert organic matter is considered to be stable and, therefore, does not participate in turnover processes.

As in other models (Falloon et al. 2000; Bruun and Jensen 2002) the manner in which CANDY is initialised influences the simulation results considerably. Model initialisation may be responsible for the behaviour of a soil as a source or sink of atmospheric CO₂ (Falloon and Smith 2000). In order to initialise the CANDY model, a value for decomposable carbon (C_{DEC}) has to be specified by the user. Because it is not possible to measure the decomposable pool directly, in CANDY C_{DEC} can be calculated from the history of the plot and the site-specific turnover conditions, or from the difference between an organic carbon (C_{ORG}) measurement and the estimated inert carbon (C_I). Thus, the amount of C_{DEC} is affected by uncertainties in estimating C_I. For example, if decomposable carbon is too high, the mineralisation of carbon and nitrogen may be overestimated, leading to a misinterpretation of organic matter dynamics and a surplus of mineral nitrogen in soil.

In the literature, several methods for calculating the amount of C_I are described. They use either the relation between C_I and soil texture (Körschens 1980; Körschens et al. 1998; Rühlmann 1999) or estimate C_I as a part of the whole amount of soil organic carbon (Falloon et al. 1998, 2000). A new approach by Kuka et al. (2006), based on porespace classes, shows that soil organic matter (SOM) localised in micro-pores is stabilised over a long time. Thus, this part of organic carbon should be very similar to the inert pool of the CANDY model.

In the investigation reported here, we tested all of these methods on their applicability for initialising the inert carbon pool in the CANDY model. Our objective was to find a method which fulfils the following criteria: (1) minimises errors in carbon and nitrogen dynamics and (2) provides usability for sites with marked differences in site conditions. In order to achieve our objective CANDY simulations were carried out by varying solely the method of estimating the initial value for C_I. Simulation results for soil organic carbon and soil mineral nitrogen of the sites 'Bad Lauchstädt' and 'Müncheberg' were compared with observations. From among the huge number of statistical

methods that can be found in literature (Willmott and Wicks 1980; Fox 1981; Addiscott and Whitmore 1987; Loague and Green 1991; Smith et al. 1996, etc.) we used the mean, the root mean square error (RMSE; Fox 1981), the mean bias error (MBE; Addiscott and Whitmore 1987) and the index of agreement (IA; Willmott and Wicks 1980) for comparing predicted and observed values and assessing the C_I estimations.

Materials and methods

CANDY model

CANDY consists of a modular system of submodels and a data base system for model parameters, initial values, weather data, soil management data and measurement values. The submodels of CANDY are described briefly below and in detail by Franko et al. (1995).

In the soil temperature model the heat flow equation is solved based on a statistical approach for the calculation of the soil surface temperature.

The hydrological model is based on a capacity approach, and takes into consideration the draining of water by gravitation forces, interception of water by crops, potential and actual evapotranspiration, surface runoff and snow cover dynamics.

The crop model in its standard version consists of parameters describing the temporal development of crop height, soil cover and rooting depth as piecewise linear functions. Nitrate uptake by plants is calculated by a sigmoidal function for the distribution of the N demand over vegetation time. After harvest, a yield-dependent amount of organic matter from roots and plant residues is recycled to the soil as fresh organic matter with crop-specific quality parameters.

The organic matter turnover model includes the soil nitrogen model. Soil organic matter is subdivided in several compartments: up to six different pools of added organic matter, two pools of decomposable soil organic matter (one active and one stabilised) and an inert soil organic matter pool, which is independent from climate and agriculture and stays constant over time. Turnover dynamics of all degradable carbon pools follow first-order kinetics but are influenced by soil temperature, soil moisture and aeration conditions indicated by soil texture and depth of the soil

layer. Nitrogen dynamics are connected to the carbon fluxes via the C/N-ratio of the pools concerned. Mineral nitrogen is divided into a nitrate and an ammonium pool. The nitrogen-related processes are: plant uptake, mineral nitrogen input including atmospheric deposition, input of organic manure or plant residues, nitrate leaching, nitrification, denitrification and mineralisation of soil organic matter leading to nitrogen mineralisation or immobilisation.

The model inputs can be classified as parameters and scenario data. The parameters are:

- plant development characteristics (standard) crop height, vegetation time, maximum root depth, etc.;
- soil data soil density, particle density, field capacity, permanent wilting point, saturated hydraulic conductivity, amount of clay and silt;
- organic matter characteristics C/N-ratio, dry matter content, turnover time, SOM reproduction coefficient.

The scenario data include:

- initial conditions for carbon, nitrogen and soil moisture;
- agricultural management data emergence, harvesting, fertilisation, tillage, yield, N-uptake;
- weather data precipitation, air temperature and global radiation preferably on a daily basis.

Datasets from Bad Lauchstädt and Müncheberg and simulation with CANDY

The datasets from Bad Lauchstädt (crop rotation) and Müncheberg (plot 1) used for the model simulation are described in detail in Franko et al. (2006) and Mirschel et al. (2006). In addition to these data we used the organic carbon measurements presented in Table 1. For Müncheberg no value at the beginning of the simulation was available. Therefore, the first value in Table 1 was estimated with a trend line using the four measured values. The simulation for both sites has been fitted to meet the starting C_{ORG} values. For the mineral nitrogen measurements the reader is referred to Franko et al. (2006) and Mirschel et al. (2006).

The simulation period for the Bad Lauchstädt crop rotation (BL) lasted from 1 September 1996 to 31 December 2003 and at Müncheberg plot 1 (MÜ) from 1 September 1992 to 31 December

Table 1. Measured soil organic carbon values for Bad Lauchstädt and Müncheberg.

Bad Lauchstädt		Müncheberg		
Date	C_{ORG}	Date	C_{ORG}	
13 August 1997	2.13	01 September 1992	0.63 ^a	
23 September 1999	2.14	03 May 1993	0.58	
27 October 2000	2.19	21 September 1993	0.67	
23 October 2001	2.03	17 July 1995	0.51	
09 October 2002	2.07	01 October 1997	0.56	
22 September 2003	2.06	_	_	

^aEstimated with a trend line using the four measured values

1998. Input data such as soil parameters, daily meteorological data and management information were taken from the datasets. In some cases data for the nitrogen uptake by crops, the vegetation time, the carbon input via straw and the C/N-ratio of straw were not specified but deducible from the datasets. If no data were available or deducible – for example, for C and N input with roots and crop residues – CANDY standard parameters (C/N-ratio, dry matter, rate and synthesis coefficients, etc.) were used. The soil parameters for wilting point and field capacity were adapted to measurements of soil moisture.

In order to initialise the CANDY model, a value for C_{DEC} in soil has to be specified. According to the equation

$$C_{DEC} = C_{ORG} - C_{I}, \tag{1}$$

with C_{ORG} = total organic carbon in soil and C_I = inert carbon. C_{DEC} were calculated using the starting C_{ORG} values from Table 1 and C_I values were calculated with the methods mentioned below. This procedure resulted in five initial values for C_{DEC} . With each initial value a CANDY simulation was started. No other parameters or scenario data were changed. Therefore, differences in simulated carbon and nitrogen dynamics came from the different initial values used.

Calculation of inert carbon (C_I)

Körschens (1980) and Körschens et al. (1998) found a relationship between C_I and the content of clay and fine silt which can be described with the equation:

$$C_{I-K\ddot{O}} = a \cdot b \tag{2}$$

where b is the percentage of soil particles < 6 μ m and a is usually taken as 0.04. According to Schulz (1997) the regression coefficient a has a range from 0.04 to 0.05. In our investigation we used both 0.04 ($C_{I-K\ddot{O}I}$) and 0.05 ($C_{I-K\ddot{O}II}$) to define the possible range of C_{I} .

Rühlmann (1999) suggested an equation to describe the influence of soil texture on the carbon content of long-term bare fallow soils:

$$C_{I-RII} = 0.017 \cdot c - 0.001 \cdot \exp(0.075 \cdot c)$$
 (3)

where c is the percentage of soil particles <20 μ m and $C_{I-R\ddot{U}}$ (%) is the soil organic carbon content of long-term bare fallow soils which were assumed to be similar to the size of the inert organic carbon pool.

Based on Falloon et al. (1998, 2000) C_I (tonne C per hectare) can be estimated from the total stock of carbon using the following equation (Eq. 4) and its upper (Eq. 5) and lower (Eq. 6) confidence levels:

$$C_{I-FAL} = 0.049 \cdot C_{ORG}^{1.139}$$
 (4)

$$C_{\text{I-FAL}(+95)} = 0.1733 \cdot C_{\text{ORG}}^{1.4624}$$
 (5)

$$C_{\text{I-FAL}(-95)} = 0.01384 \cdot C_{\text{ORG}}^{0.8156} \tag{6} \label{eq:6}$$

where C_{ORG} (in tonnes C per hectare) is the total amount of organic C in the soil.

An approach to estimate organic carbon in micro-pores based on the hypothesis that the stabilisation of soil organic matter is a result of its location at places with low biological activity, which in turn is caused by a limitation of oxygen, was suggested by Kuka et al. (2006). In our investigation with the CANDY model the organic carbon in micro-pores is considered to be inert.

The total organic carbon is distributed to the different pore-space classes – micro-, meso- and macro-pores – according to their surface area. The pore-space classes used (micro, meso and macro) are related to wilting point (WP), field capacity (FC) and pore volume (PV), respectively. The inner surface area of each pore class is calculated from the volume of the considered pore-space class $V_{\rm m}$ [WP, FC–WP, PV–FC (in cubic metres)] and

the equivalent pore radius $R_{\rm m}$ (in metres). $R_{\rm m}$ for the micro-, meso- and macro-pores was set to 5×10^{-8} , 10×10^{-8} and 500×10^{-8} m, respectively.

$$A_m = 2 \cdot \frac{V_{\text{m}}}{R_{\text{m}}}, \quad m \in \{\text{micro}; \text{meso}; \text{macro}\}.$$
 (7)

Considering organic carbon in micro-pores (C_{I-MIP}) as 'inert', its amount can be calculated according to the equation:

$$C_{\text{I-MIP}} = C_{\text{ORG}} \cdot \frac{A_{\text{micro}}}{A_{\text{micro}} + A_{\text{meso}} + A_{\text{macro}}}, \quad (8)$$

with $C_{\text{I-MIP}}$ (M.%) mass of carbon in micro-pores, C_{ORG} (M.%) mass of total organic carbon in soil and $A_{\text{micro, meso, macro}}$ (in square metres) inner surface of pore class.

Statistical methods

For this investigation we selected the following statistical methods to obtain information about differences deriving from varying C_I initial values.

The root mean square error (RMSE; Fox 1981) and the mean bias error (MBE; Addiscott and Whitmore 1987) provide information about the average difference between predicted (P_i) and observed (O_i) values.

$$RMSE = \sqrt{\frac{\sum_{I=1}^{n} (P_I - O_I)^2}{n}}$$
 (9)

$$MBE = \sum_{I=1}^{n} \frac{P_I - O_I}{n}, n = \text{number of samples.} \quad (10)$$

A lower RMSE indicates a more accurate simulation. The lower limit of RMSE is 0. MBE can take positive or negative values. Its calculation does not include a square term, thus, predicted values below and above the observed values cancel out, and the result gives an indication of the bias error.

The 'index of agreement' (IA) suggested by Willmott and Wicks (1980) is intended to be descriptive and is both a relative and a bounded measure.

$$IA = 1 - \frac{\sum_{i=1}^{n} (P_I - O_I)^2}{\sum_{i=1}^{n} (|P_i - \overline{O}| + |O_I - \overline{O}|)^2},$$

$$0 \le IA \le 1,$$
(11)

with n = number of samples and $\overline{O} =$ mean of the observed data.

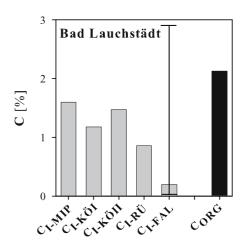
Results and discussion

Calculated inert carbon

The C_I values used to calculate initial C_{DEC} values for the simulation of the Bad Lauchstädt and Müncheberg sites and the total organic carbon content are shown in Figure 1. Apart from the remarkable differences in total organic carbon between the two sites, Figure 1 shows that with the C_{I-FAL} method, which is simply based on total C_{ORG}, the lowest C_I values are obtained. However, the lower and upper confidence levels of C_{I-FAL} cover nearly the whole range of possible inert carbon. This is in complete agreement with the statement of Falloon et al. (1998) 'that the confidence limits of the model are wide'. For the methods $C_{\text{I-K\"OI}},\,C_{\text{I-K\"OII}}$ and $C_{\text{I-R\"U}},$ which are all based on soil texture and suggest a kind of physical protection of soil organic matter, the following order for both sites was found: $C_{I-R\ddot{U}} < C_{I-K\ddot{O}I} < C_{I-K\ddot{O}II}$. Whereas, with the new C_{I-MIP} approach the highest inert carbon value of all methods was calculated in the case of the Bad Lauchstädt site. With respect to the Müncheberg site C_{I-MIP} is located between $C_{\text{I-R}\ddot{\text{U}}}$ and $C_{\text{I-K}\ddot{\text{O}}\text{I}}.$ Thus, a clear ranking of the C_{I-MIP} method with respect to the amount of calculated inert carbon is not possible. Similar to the $C_{I-K\ddot{O}I}$, $C_{I-K\ddot{O}II}$ and $C_{I-R\ddot{U}}$ approaches, the C_{I-MIP} approach is suggesting physical protection, but the latter uses pore-space classes instead of soil texture and is also dependent on the total amount of C_{ORG} . Therefore, it shows a more dynamic behaviour. For example, changes in pore-space class distribution because of a modified tillage regime result in a different C_{I-MIP} value.

Simulation results and statistics

To give the reader an impression of the model outputs produced with different initial C_I values,



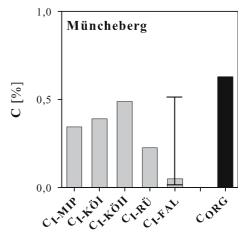


Figure 1. Calculated inert carbon (C_I) and total soil organic carbon (C_{ORG}) for Bad Lauchstädt and Müncheberg. $C_{I\text{-MIP}} = K$ uka et al. 2006; $C_{I\text{-K\bar{O}I}}$ and $C_{I\text{-K\bar{O}II}} = K$ örschens 1980 and Körschens et al. 1998 with a = 0.04 and 0.05, respectively; $C_{I\text{-R\bar{U}}} = R$ ühlmann 1999; $C_{I\text{-FAL}} = F$ alloon et al. 1998, 2000. I = Confidence interval for $C_{I\text{-FAL}}$.

the simulation results for C_{ORG} and N_{MIN} are presented in Figures 2 and 3, respectively. In the case of C_{ORG} there is a clear differentiation – increasing with time – between the model runs. The distinct influence of different initial C_I values on N_{MIN} in soil, especially for the Bad Lauchstädt site, emphasises the importance of a careful model initialisation even for short simulation periods.

For Bad Lauchstädt the carbon and nitrogen dynamics and the differences between the simulations are higher than for Müncheberg. In addition to mineralisation, this is may be due to other processes such as leaching or denitrification. In fact, nitrogen leaching is very low at the Bad Lauchstädt

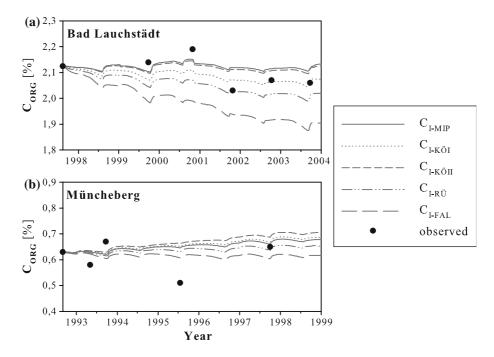


Figure 2. Simulation results for organic carbon (C_{ORG}) dynamics using different initial C_{I} values for Bad Lauchstädt (**a**) and Müncheberg (**b**). $C_{I-MIP} = K$ uka et al. 2006; $C_{I-K\ddot{0}I}$ and $C_{I-K\ddot{0}II} = K\ddot{0}$ and Körschens et al. 1998 with a = 0.04 and 0.05, respectively; $C_{I-R\ddot{0}I} = R\ddot{0}$ Rühlmann 1999; $C_{I-FAL} = F$ alloon et al. 1998, 2000.

site (see Kersebaum et al. 2006) due to a high water storage capacity and low precipitation (see Franko et al. 2006). Thus, the other nitrogen-related processes have to be adequately taken into account for an assessment of model initialisation.

The examination of the mean (m_C, m_N) and the mean bias error (MBE_C, MBE_N) for carbon and nitrogen, respectively $(Tables\ 2,\ 3)$, indicates that in most cases the observed values are exceeded by the corresponding predicted values. Only for organic carbon and $C_{I-K\ddot{O}I}$, $C_{I-R\ddot{U}}$ and C_{I-FAL} of the Bad Lauchstädt site an underestimation can be found. For the Müncheberg site, the MBE_C and MBE_N are the lowest for C_{I-FAL} and $C_{I-K\ddot{O}II}$, respectively. For Bad Lauchstädt, the lowest mean bias errors are obtained for $C_{I-K\ddot{O}I}$ and C_{I-MIP} . Consequently, no method seems to be clearly better than the others in predicting initial values for $C_{I-K\ddot{O}I}$.

Looking at the carbon root mean square error (RMSE, Figure 4) for Bad Lauchstädt, the method for estimating C_I proposed by Falloon et al. (1998, 2000) differs clearly from the other C_I methods, which account for physical protection of soil organic matter. Obviously, the $C_{I\text{-}FAL}$ method

underestimates the inert part of soil organic matter for Bad Lauchstädt (see Figure 1), leading to an overestimation of the decomposable part, followed by a fast decline in soil organic matter and a surplus of mineral nitrogen (see Figure 2). For Müncheberg there is no substantial difference in RMSE_C between any of the methods used in this investigation. In the case of mineral nitrogen, greater differences in the root mean square error can be found. C_{I-MIP} and $C_{I-K\ddot{O}II}$ are in the leading group for the Bad Lauchstädt site and, accompanied by $C_{I-K\ddot{O}I}$, in the leading group for the Müncheberg site.

The index of agreement (IA, Figure 5) more precisely suggests that for both sites the C_{I-FAL} method decreases in accuracy compared to the others with respect to predict C_I in the initialisation of the CANDY model. Falloon et al. (2000) pointed out that their C_I method was not expected to be valid for soils with a large C_I content. The 'Haplic Chernozem' (FAO 1994) of Bad Lauchstädt may belong to that category. However, Schmidt et al. (1999) detected charred organic carbon (up to 45% of the bulk organic carbon) in German chernozemic soils.

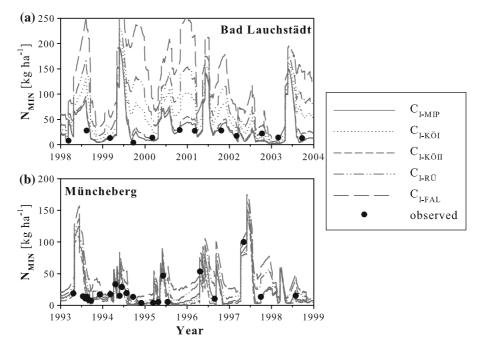


Figure 3. Simulation results for mineral nitrogen (N_{MIN}) dynamics using different initial $C_{\rm I}$ values for Bad Lauchstädt (**a**) and Müncheberg (**b**). $C_{\rm I-MIP} = {\rm Kuka}$ et al. 2006; $C_{\rm I-K\ddot{O}I}$ and $C_{\rm I-K\ddot{O}I} = {\rm K\ddot{o}}$ respectively; $C_{\rm I-R\ddot{U}} = {\rm R\ddot{u}hlmann}$ 1999; $C_{\rm I-FAL} = {\rm Falloon}$ et al. 1998, 2000.

Table 2. Quantitative statistical measures for simulations with different initial inert carbon ($C_{\rm I}$) values for the Bad Lauchstädt site. $C_{\rm I-MIP} = K$ uka et al. 2006; $C_{\rm I-K\ddot{O}I}$ and $C_{\rm I-K\ddot{O}II} = K$ örschens 1980 and Körschens et al. 1998 with a = 0.04 and 0.05, respectively; $C_{\rm I-R\ddot{U}} = R\ddot{u}hlmann$ 1999; $C_{\rm I-FAL} = Falloon$ et al. 1998, 2000

Bad Lauchstädt ^a	m_C^{b}	$MBE_C^{\ b}$	$m_N^{\ b}$	MBE_N^b
C _{I-MIP}	2.12	0.027	23.6	5.49
$C_{I-K\ddot{O}I}$	2.08	-0.018	51.6	33.5
C _{I-KÖII}	2.11	0.021	28.8	10.7
$C_{I-R\ddot{U}}$	2.03	-0.060	75.8	57.7
C_{I-FAL}	1.94	-0.149	124	106
O^c	2.09	-	18.1	_

 $^{^{}a}n_{C} = 5$; $n_{N} = 12$.

The new C_{I-MIP} approach proves to be as good as $C_{I-K\ddot{O}I}$ and $C_{I-K\ddot{O}II}$. $C_{I-R\ddot{U}}$ is located between the leading group and C_{I-FAL} . The usefulness of the approach by Körschens (1980) and Körschens et al. (1998) was also pointed out by Ludwig et al. (2003), however there are some indications that

Table 3. Quantitative statistical measures of simulations with different initial inert carbon ($C_{\rm I}$) values for the Müncheberg site. $C_{\rm I-MIP}=K$ uka et al. 2006; $C_{\rm I-K\bar{O}I}$ and $C_{\rm I-K\bar{O}II}=K$ örschens 1980 and Körschens et al. 1998 with a=0.04 and 0.05, respectively; $C_{\rm I-R\bar{U}}=R$ ühlmann 1999; $C_{\rm I-FAL}=F$ alloon et al. 1998, 2000

Müncheberg ^a	m_C^{b}	$MBE_C^{\ b}$	m_N^{b}	MBE_N^b
C _{I-MIP}	0.64	0.036	28.5	8.07
$C_{I-K\ddot{O}I}$	0.64	0.040	25.3	4.86
C _{I-KÖII}	0.65	0.050	20.9	0.53
$C_{I-R\ddot{U}}$	0.63	0.024	38.8	18.4
C _{I-FAL}	0.61	0.006	50.8	30.4
O^{c}	0.60	_	20.4	_

 $^{^{}a}n_{C} = 4; n_{N} = 23.$

this method is not generally applicable. Problems arise particularly for soils with a very high content of fine particles (FP = clay + fine silt) and low C_{ORG} content, a situation that can be found in the long-term experiment from Prague-Ruzyně with FP=41.2% (J. Klir, personal communication) and

^bm, Mean; MBE, mean bias error; subscript C, carbon (%); subscript N, mineral nitrogen (kg ha⁻¹).

^cO, Observed.

^bm, Mean; MBE, mean bias error; subscript C, carbon (%); subscript N, mineral nitrogen (kg ha⁻¹).

^cO, Observed.

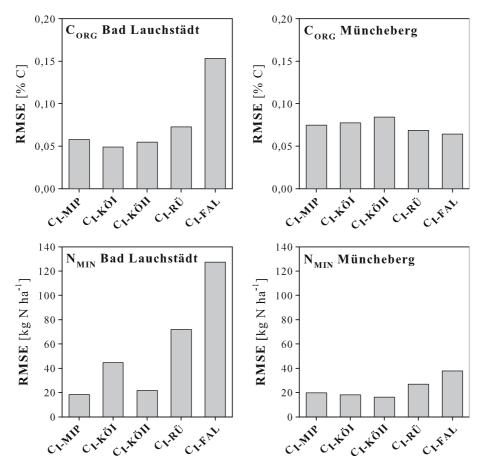


Figure 4. Root mean square error (RMSE) for organic carbon (%) and mineral nitrogen (kg ha⁻¹) of simulations with different initial C_I values for the sites Bad Lauchstädt and Müncheberg. $C_{I-MIP} = Kuka$ et al. 2006; $C_{I-K\bar{O}I}$ and $C_{I-K\bar{O}II} = K\ddot{o}$ rschens 1980 and Körschens et al. 1998 with a = 0.04 and 0.05, respectively; $C_{I-R\bar{U}} = R\ddot{u}$ hlmann 1999; $C_{I-FAL} = Falloon$ et al. 1998, 2000).

 $C_{\rm ORG}$ values from 1.1 to 1.4% (Kubat et al. 2003). Other examples are the long-term experiments of Giessen with $C_{\rm ORG} = 0.99\%$ and FP = 25% (Boguslawski and Debruck 1976) and Gembloux with $C_{\rm ORG} = 0.92\%$ and FAT = 24% (Droeven et al. 1982). In all these cases the approach by Körschens (1980) and Körschens et al. (1998) leads to an overestimation of $C_{\rm I}$.

Conclusion

Our results show that an important factor in the CANDY model initialisation procedure is the method chosen for the calculation of C_I as this will have a strong influence on modelling results, even in short-term simulations. The coupled investigation of carbon and nitrogen dynamics leads to a more secure decision on the applicability of a C_I

method than simply looking at carbon dynamics because an underestimation of inert carbon causes an overestimation of decomposable carbon and a surplus of mineral nitrogen in the soil (or vice versa).

Of the C_I methods considered, the $C_{I\text{-}FAL}$ method, based simply on total soil organic carbon, is shown to be the least appropriate for a soil with a high amount of inert carbon, such as the Bad Lauchstädt 'Haplic Chernozem', whereas the $C_{I\text{-}K\ddot{O}I}$ and $C_{I\text{-}K\ddot{O}II}$ methods as well as the new $C_{I\text{-}MIP}$ approach, all of which allow for the effects of physical protection of soil organic matter, proved to be successful in estimating the inert organic carbon pool in soils of both of the sites investigated. Hence, physical protection seemed to be of major importance for simulating carbon and nitrogen dynamics.

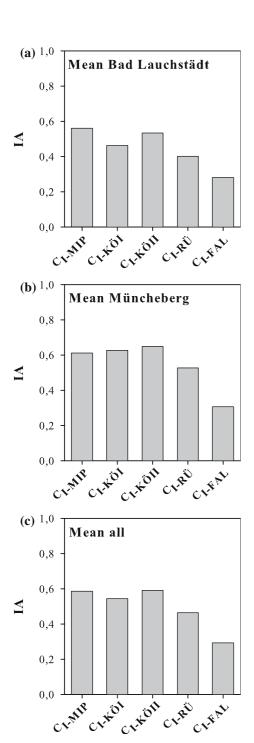


Figure 5. Mean index of agreement (IA) of simulations with different initial C_I values for the sites Bad Lauchstädt (a) and Müncheberg (b) (in each case: mean of organic carbon and mineral nitrogen) and for both sites (c) (mean of organic carbon and mineral nitrogen and sites). $C_{I-MIP} = Kuka$ et al. 2006; $C_{I-K\bar{O}I}$ and $C_{I-K\bar{O}II} = K\ddot{o}$ rschens 1980 and Körschens et al. 1998 with a = 0.04 and 0.05, respectively; $C_{I-R\bar{U}} = R\ddot{u}$ hlmann 1999; $C_{I-FAL} = F$ alloon et al. 1998, 2000.

In contrast to the methods which primarily consider soil texture ($C_{I\text{-}K\ddot{O}I},\,C_{I\text{-}K\ddot{O}II}$ and $C_{I\text{-}R\ddot{U}})$ the $C_{I\text{-}MIP}$ approach uses pore-space classes and also includes the total amount of $C_{ORG}.$ Changes in pore-space class distribution of one soil, because of a modified tillage or a changed organic manure regime, would result in a different $C_{I\text{-}MIP}$ value for the soil. Conseuently, this $C_{I\text{-}MIP}$ method is able to express differences in soil structure and organic matter supply at one site. Further studies, with a wider range of soil types and soils with different tillage regimes and soil organic matter levels, are required to see whether the new $C_{I\text{-}MIP}$ approach is generally applicable.

Acknowledgements

We thank Drs. W. Mirschel and K.C. Kersebaum for providing additional C_{ORG} data of the Müncheberg site, Mr. A.D. Liston for improving the English and two anonymous reviewers for helpful comments on an earlier draft of this manuscript. We are also grateful to the organisers of the workshop.

References

Addiscott T.M. and Whitmore A.P. 1987. Computer simulation of changes in soil mineral nitrogen and crop nitrogen during autumn, winter and spring. J. Agric. Sci. Cambridge 109: 141–157.

von Boguslawski E. and Debruck J. 1976. Ergebnisse aus einem langjährigen Stallmist-Schafpferchversuch in Rauisch-Holzhausen. Z. Acker-u. Pflanzenbau 143: 223–242.

Bruun S. and Jensen L.S. 2002. Initialisation of soil organic matter pools of the Daisy model. Ecol. Model. 153: 291–295.

Coleman K. and Jenkinson D.S. 1996. Roth C-26.3-A model for the turnover od carbon in soil. In: Powlson D.S., Smith P. and Smith J.U. (eds), Evaluation of Soil Organic Matter Models Using Existing Long-term Datasets, Vol. 38. NATO ASI series, Series I, Springer-Verlag, Heidelberg, pp. 237–246.

Droeven G., Rixhon L., Crohain A. and Raimond Y. (1982). Long term effects of different systems of organic matter supply on the humus content and on the structural stability of soils regard to the crop yields in loamy soils. Soil degradation. In:Balkema A.A. (ed.), Proc Land Use Sem Soil Degradation, PUDOC, Wageningen, pp. 203–222.

Falloon P. and Smith P. 2000. Modelling refractory soil organic matter. Biol. Fertil. Soils 30: 388–398.

Falloon P., Smith P., Coleman K. and Marshall S. 1998. Estimating the size of the inert organic matter pool from total soil organic carbon content for use in the Rothamsted carbon model. Soil Biol. Biochem. 30: 1207–1211.

- Falloon P., Smith P., Coleman K. and Marshall S. 2000. How important is inert organic matter for predictive soil carbon modelling using the Rothamsted carbon model?. Soil Biol. Biochem. 32: 433–436.
- FAO 1994. FAO-UNESCO soil map of the world. Revised legend. FAO, Rome.
- Fox D.G. 1981. Judging air quality model performance: a summary of the AMS Workshop on dispersion model performance. Bull. Am. Meteorol. Soc. 62: 599–609.
- Franko U. 1996. Modelling approaches of soil organic matter within the CANDY system. In: Powlson D.S., Smith P. and Smith J.U. (eds), Evaluation of Soil Organic Matter Models Using Existing Long-term Datasets, Vol. 38, NATO ASI Series, Series I, Springer, Berlin Heidelberg New York, pp. 247–254.
- Franko U. 1997. Modellierung des Umsatzes der organischen Bodensubstanz. Arch. Acker-Pfl. Boden 41: 527–547.
- Franko U., Oelschlägel B. and Schenk S. 1995. Simulation of temperature-, water- and nitrogen dynamics using the model CANDY. Ecol. Model. 81: 213–222.
- Franko U., Crocker G.J., Grace P.R., Klir J., Körschens M., Poulton P.R. and Richter D.D. 1997. Simulating trends in soil organic carbon in long-term experiments using the CANDY model. Geoderma 81: 109–120.
- Franko U., Puhlmann M., Kuka K., Böhme F. and Merbach I. 2006. Dynamics of water, carbon and nitrogen in an agricultural used Chernozem soil in Central Germany. In: Kersebaum K.C., Hecker J.-M. and Mirschel W. (eds), Modelling Water and Nutrient Dynamics in Soil Crop Systems. Springer, Stuttgart (in press).
- Hansen S., Jensen H.E., Nielsen N.E. and Svendsen H. 1991. Simulation of nitrogen dynamics and biomass production in winter wheat using the Danish simulation model DAISY. Fert. Res. 27: 245–259.
- Kersebaum K.C. (2006). Water and nutrient dynamics in soil-crop systems with HERMES. In: Kersebaum K.C., Hecker J.-M. and Mirschel W. (eds), Modelling Water and Nutrient Dynamics in Soil Crop Systems. Springer, Stuttgart (in press). Klir J. 2005. Personal communication.
- Körschens M. 1980. Beziehungen zwischen Feinanteil, Ct- und Nt-Gehalt des Bodens. Arch. Acker- Pfl. Boden 24: 585–592.
- Körschens M., Weigel A. and Schulz E. 1998. Turnover of soil organic matter (SOM) and long-term balances – tools for evaluating sustainable productivity of soils. Z. Pflanzenernähr. Bodenk. 161: 409–424.
- Kubat J., Klir J. and Pova D. 2003. The dry matter yields, nitrogen uptake and the efficacy of nitrogen fertilisation in long-term field experiments in Prague. Plant Soil Environ. 49: 337–345.

- Kuka K., Franko U., Rühlmann J., Martens R., Vogt M. and Kalbitz K. 2006. Modelling the impact of pore space distribution on carbon turnover. Ecol. Model. (submitted).
- Loague K. and Green R.E. 1991. Statistical and graphical methods for evaluating solute transport models: overview and application. J. Contam. Hydrol. 7: 51–73.
- Ludwig B., John B., Ellerbrock R., Kaiser M. and Flessa H. 2003. Stabilization of carbon from maize in sandy soil in a long-term experiment. Eur. J. Soil Sci. 54: 117–126.
- Mirschel W., Wenkel K.-O., Wegehenkel M., Kersebaum K.C., Schindler U. and Hecker J.-M. (2006). Müncheberg field trial data set for agro-ecosystem model validation. In: Kersebaum K.C., Hecker J.-M. and Mirschel W. (eds), Modelling Water and Nutrient Dynamics in Soil Crop Systems. Springer, Stuttgart (in press).
- Parton W.J. (1996). The CENTURY model. In: Powlson D.S., Smith P. and Smith J.U. (eds), Evaluation of Soil Organic Matter Models Using Existing Long-term Datasets, Vol. 38, NATO ASI Series, Series I, Springer, Berlin Heidelberg New York, pp. 283–293.
- Rühlmann J. 1999. A new approach to estimating the pool of stable organic matter in soil using data from long-term field experiments. Plant Soil 213: 149–160.
- Schulz E. 1997. Charakterisierung der organischen Bodensubstanz (OBS) nach dem Grad ihrer Umsetzbarkeit und ihre Bedeutung für Transformationsprozesse für Nähr- und Schadstoffe. Arch. Acker-Pfl. Boden 41: 465–483.
- Schmidt M.W.I., Skjemstad J.O., Gehrt E. and Kögel-Knabner I. 1999. Charred organic carbon in German chernozemic soils. Eur. J. Soil Sci. 50: 351–365.
- Smith J.U., Smith P. and Addiscott T. 1996. Quantitative methods to evaluate and compare soil organic matter (SOM) models. In: Powlson D.S., Smith P. and Smith J.U. (eds), Evaluation of Soil Organic Matter Models Using Existing Long-term Datasets, Vol. 38, NATO ASI Series, Series I, Springer, Berlin Heidelberg New York, pp. 181–199.
- Smith P., Smith J.U., Powlson D.S., McGill W.B., Arah J.R.M., Chertov O.G., Coleman K., Franko U., Frolking S., Jenkinson D.S., Jensen L.S., Kelly R.H., Klein-Gunnewiek H., Komarov A.S., Li C., Molina J.A.E., Mueller T., Parton W.J., Thonley J.H.M. and Whitmore A.P. (1997). A comparison of the performance of nine soil organic matter models using datasets from seven long-term experiments. Geoderma 81: 153–225.
- Willmott C.J. and Wicks D.E. 1980. An empirical method for the spatial interpolation of monthly precipitation within California. Phys. Geogr. 1: 59–73.