

From Precipitation to Prediction: Integrating ConvLSTM Models for Comprehensive River Runoff Forecasting

Florian Börgel¹, Sven Karsten¹

¹Leibniz-Institute for Baltic Sea Research Warnemünde,

Corresponding author: Florian Börgel, florian.boergel@io-warnemuende.de

Abstract

a

Plain Language Summary

b

1 Introduction

River runoff is an important component of the global water cycle as it comprises about one third of the precipitation over land areas (Hagemann et al., 2020). In the context of climate change studies, river runoff is usually generated in two ways. First, river runoff as input for ocean models can be created using hydrological models such as the Hydrological Discharge (HD) model (Hagemann et al., 2020). HD models calculate the water balance using hydrological processes (e.g. snow, glaciers, soil moisture, groundwater contribution). It represents a complex forecasting tool that uses underlying physical processes. A different approach would use data-based models that integrate statistical correction, using the land surface schemes of global or regional climate models.

The relatively recent rise of machine learning (ML) models has been mostly explored for river runoff forecasting, as accurate runoff forecasting, especially over extended periods, is pivotal for effective water resources management (Fang et al., 2019; Tan et al., 2018; Yang et al., 2018). Common approaches employ artificial neural networks, support vector machines, adaptive neuro-fuzzy inference systems, and notably, Long Short-Term Memory (LSTM) neural networks that have gained traction for long-term hydrological forecasting due to their excellent performance (Humphrey et al 2016, Huang et al 2014, Ashrafi et al 2017, Yuan et al 2018, Xu et al 2021).

LSTM networks, an evolution of the classical Recurrent Neural Networks (RNNs), have shown stability and efficacy in sequence-to-sequence predictions, such as using climatic indices for rainfall estimation or long-term hydrological forecasting. However, a limitation of LSTMs is their inability to effectively capture two-dimensional structures, an area where Convolutional Neural Networks (CNNs) excel. Recognizing this limitation (SHI et al., 2015) proposed a convolutional LSTM (ConvLSTM) architectures, which combines the strengths of both LSTM and CNN. The ConvLSTM network has been proven useful for precipitation nowcasting (SHI et al., 2015) or for river runoff forecasting [Ha et al. (2021); @niu2020].

In the following we will show that in absence of a fully functioning hydrological model, that also uses a rather complex parametrization, ConvLSTM also represent a robust way to predict multiple rivers at once for any given period using only atmospheric forcing. In this work we use the Baltic Sea catchment to illustrate our approach, while in principal the methodology we propose is universally applicable across various geographic regions. The Baltic Sea serves as a challenging example due to its unique hydrological characteristics, being nearly decoupled from the open ocean (see Figure). As a consequence, the salinity of the Baltic Sea is driven to a large part by freshwater supply from rivers. More generally, the freshwater input into the Baltic Sea comes either as river runoff or a positive net precipitation (precipitation minus evaporation) over the sea surface. The net precipitation accounts for 11 % and the river input for 89 % of the total freshwater input (Meier and Doescher, 2002). Modeling the Baltic Sea is therefore to a large part the result of the quality of the river input, that is used for the simulation. This makes the accurate modeling of river runoff especially critical for simulations pertaining to the Baltic Sea.

In this work we will, we present a ConvLSTM architecture that is able to predict daily river runoff for 97 rivers across the Baltic Sea catchment.

2 Methods

2.1 Runoff data used for training

The runoff data covering the period 1979 to 2011 is based on an E-HYPE hindcast simulation that was forced by a regional downscaling of ERA-Interim (Dee et al., 2011) with RCA3 (“The rossby centre regional climate model RCA3,” n.d.) and implemented into NEMO-Nordic (Hordoir et al., 2019) as a mass flux. For the periods before (1961 to 1978) and after (2012 to 2018) additional spatial temporal corrections have been applied to the runoff data, and have therefore been ignored. For more information see Gröger et al. (2022) and references therein.

2.2 Atmospheric Forcing

The UERRA-HARMONIE regional reanalysis dataset was developed as part of the FP7 UERRA project (Uncertainties in Ensembles of Regional Re-Analyses, <http://www.uerra.eu/>),. The UERRA-HARMONIE system represents a comprehensive, high-resolution reanalysis covering a wide range of essential climate variables. This dataset encompasses data on air temperature, pressure, humidity, wind speed and direction, cloud cover, precipitation, albedo, surface heat fluxes, and radiation fluxes from January 1961 to July 2019. With a horizontal resolution of 11 km and analyses conducted at 00 UTC, 06 UTC, 12 UTC, and 18 UTC, it also provides hourly resolution forecast model data. UERRA-HARMONIE is accessible through the Copernicus Climate Data Store (CDS, <https://cds.climate.copernicus.eu/#!/home>), initially produced during the UERRA project and later transitioned to the Copernicus Climate Change Service (C3S, <https://climate.copernicus.eu/copernicus-regionalreanalysis-europe>).

2.3 Ocean Model

To simulate the Baltic Sea, we use a coupled 3-dimensional ocean model Baltic Sea, called the Modular Ocean Model (MOM). This model uses a finite-difference method to solve the full set of primitive equations to calculate the motion of water and the transport of heat and salt. The K-profile parameterization (KPP) was used as turbulence closure scheme. The model’s western boundary opens into the Skagerrak and connects the Baltic Sea to the North Sea. The maximum depth was set at 264 meters. A more detailed description of the setup can be found in (Gröger et al., 2022).

2.4 LSTM network

The Long Short-Term Memory (LSTM), a specialized form of Recurrent Neural Networks (RNNs), is specifically tailored for modeling temporal sequences. Its unique design allows it to adeptly handle long-range dependencies, setting it apart from traditional RNNs in terms of accuracy (see Figure 1).

This performance in modeling long-range dependencies has been validated in various studies. The key component of LSTM’s innovation is its memory cell, c_t , which stores state information, also referred to as long-term memory. This cell is accessed, modified, and reset through several self-parameterized gates. For the input of the sequence x_t input, the forget gate f_t defines the percentage of the previous long-term memory status c_{t-1} that should be retained stored. Next the input gate i_t decides how much of the input is added to the the long-term memory, forming the updated cell state c_t . The decision to propagate the latest cell output, c_t , to the final state, h_t , is governed by the output gate, o_t , representing the updated short-term memory of the hidden state h_t . A significant advantage of this architecture is the memory cell’s ability to retain gradients. This mechanism addresses the vanishing gradient problem, where, as input sequences elongate, the influence of initial stages becomes harder to capture, causing gradients of early input points to approach zero. The

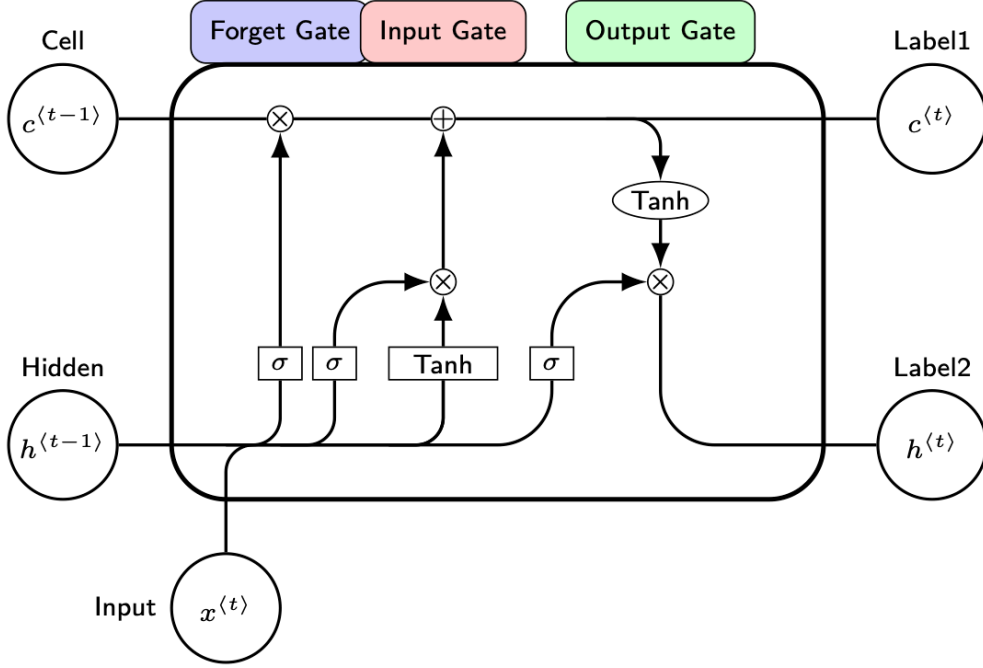


Figure 1: Inner structure of a Long Short-Term Memory Cell

LSTM's activation function, inherently recurrent, mirrors the identity function with a consistent derivative of 1.0, ensuring the gradient remains stable throughout back-propagation.

One LSTM cell hence maybe expressed as:

$$\begin{aligned}
 i_t &= \sigma(W_{xi}x_t + W_{hi}h_{t-1} + W_{ci} \circ c_{t-1} + b_i) \\
 f_t &= \sigma(W_{xf}x_t + W_{hf}h_{t-1} + W_{cf} \circ c_{t-1} + b_f) \\
 c_t &= f_t \circ c_{t-1} + i_t \circ \tanh(W_{xc}x_t + W_{hc}h_{t-1} + b_c) \\
 o_t &= \sigma(W_{xo}x_t + W_{ho}h_{t-1} + W_{co} \circ c_t + b_o) \\
 h_t &= o_t \circ \tanh(c_t)
 \end{aligned}$$

with

- x_t : Input vector at time step t .
- h_{t-1} : Hidden state from the previous time step.
- c_{t-1} : Cell state from the previous time step.
- W and b : Weight matrices and bias vectors, respectively, associated with the gates of the LSTM. The subscripts denote the specific gate or operation they are associated with (e.g., W_f and b_f are the weight matrix and bias for the forget gate).
- σ : Sigmoid activation function (0, ... 1)
- \tanh : Hyperbolic tangent activation function (-1 ... 1)

2.5 ConvLSTM network

The FC-LSTM fails to handle information when handling spatiotemporal data due to its reliance on full connections in both input-to-state and state-to-state tran-

sitions. To adress this limitation we use a convLSTM architecture. convLSTM replaces the fully connected operations in the LSTM with convolutional operations. Hence, all inputs X_1, \dots, X_t , cell outputs C_1, \dots, C_t , hidden states H_1, \dots, H_t , and gates i_t, f_t, o_t of the ConvLSTM are 3D tensors. The last two dimensions of these tensors represent spatial dimensions, specifically rows and columns. Conceptually, these inputs and states can be visualized as vectors positioned on a spatial grid.

In the ConvLSTM, the future state of a specific cell on this grid is determined by the inputs and past states of its neighboring cells. This spatial consideration is integrated by employing a convolution operator in both state-to-state and input-to-state transitions, as illustrated in Fig. 2. The foundational equations for ConvLSTM are:

$$\begin{aligned} i_t &= \sigma(W_{xi} * X_t + W_{hi} * H_{t-1} + W_{ci} \circ C_{t-1} + b_i) \\ f_t &= \sigma(W_{xf} * X_t + W_{hf} * H_{t-1} + W_{cf} \circ C_{t-1} + b_f) \\ C_t &= f_t \circ C_{t-1} + i_t \circ \tanh(W_{xc} * X_t + W_{hc} * H_{t-1} + b_c) \\ o_t &= \sigma(W_{xo} * X_t + W_{ho} * H_{t-1} + W_{co} \circ C_t + b_o) \\ H_t &= o_t \circ \tanh(C_t) \end{aligned}$$

In summary, the ConvLSTM excels at processing tasks that demand a combined understanding of spatial patterns and temporal sequences in data. It merges the image-processing capabilities of Convolutional Neural Networks (CNNs) with the time-series modeling of Long Short-Term Memory (LSTM) networks.

2.6 Implemented model architecture

The ConvLSTM architectures uses an encoder/decoder structure as discussed in TODO. To predict all 97 rivers entering the Baltic Sea, we flatten the output and use fully connected layers to map onto the individual rivers outputs.

An overview of the model structure is given below

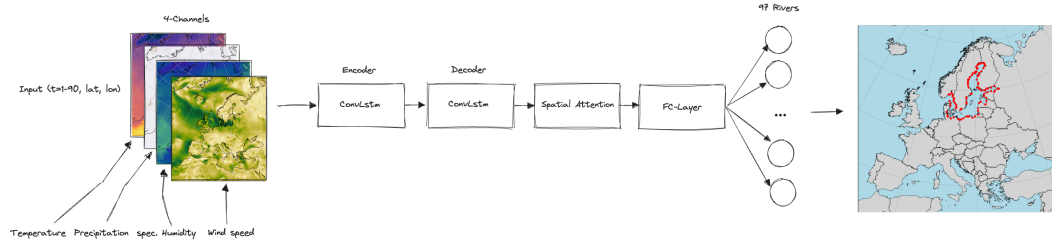


Figure 2: BaltConvLSTM

For the computation we use the following set of hyper parameters:

Table 1: Hyperparameters

Parameter name	Parameter size
Channel size	4
Num. hidden layer	10
Num. timesteps	30
Conv. Kernelsize	(9,9)
Num. ConvLSTM layers	1
Batch size	64

Parameter name	Parameter size
Learning Rate	1e-3 with CosineAnnealing

As input the model receives 30 days of atmospheric surface fields temperature T , precipitation P , specific humidity Q and wind speed W , with a daily resolution to predict the river runoff R at the time step $\Delta t + 1$, which can be summarized as

$$R_{\Delta t+1} = f(T_{t-30:t}, P_{t-30:t}, Q_{t-30:t}, W_{t-30:t})$$

with f being a function maps the 30 days of daily atmospheric surface fields data to the predicted river runoff.

The choice of atmospheric fields was based on the implemented river runoff calculation in the atmospheric model COSMO-CLM which uses these four fields to calculate an river runoff estimate.

3 Results

For the evaluation of the model performance we consider the period 1979 to 2011. For this period no bias correction was applied to the original E-HYPE dataset. We chose a split of 80% training data, 10% validation data to evaluate the performance of the model during training, and 10% training data that is finally used to evaluate the performance of the model after training.

Figure 3 shows the performance of the model using the test dataset. The predicted total river runoff for the Baltic Sea is closely matching the original data. Zooming in on the largest individual rivers (lower panels) it can be seen that that also the prediction of the individual rivers is close to the original data.

The accuracy of the model is further displayed in Figure 4. The correlation is close to 1 and the residuals show the error of the model is below 1% of the original data.

Lastly, we evaluated the performance of the runoff model by incorporating the predicted river runoff as forcing into the ocean model MOM5. This provides a robust validation of the runoff model against more complex real world conditions. This allows us to ensure that the predictions accurately reflect the impact of the river discharge on the ocean dynamics, validating the temporal and spatial variability of the the river discharge. Figure 15 shows the salinity comparison between the original E-HYPE river runoff and the predicted river runoff at BY15 - a central stations in the Baltic Sea.

BY15

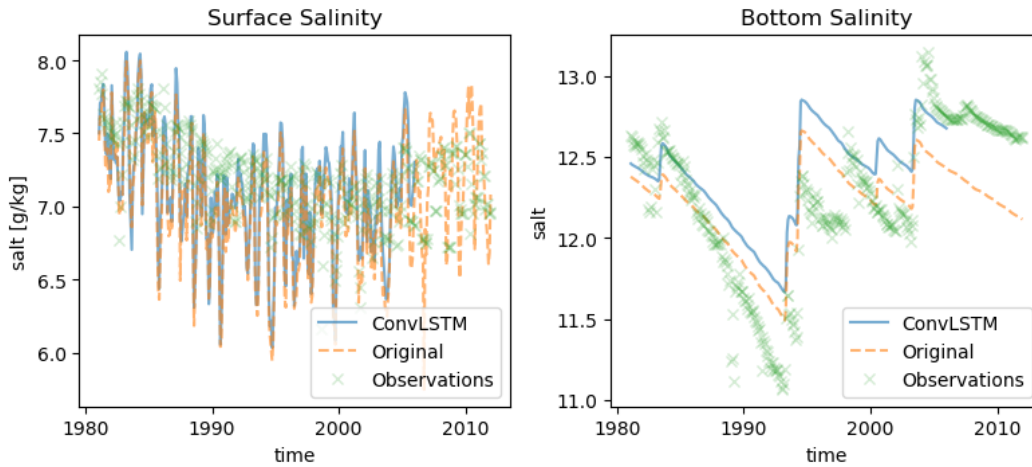




Figure 3: Figure ..

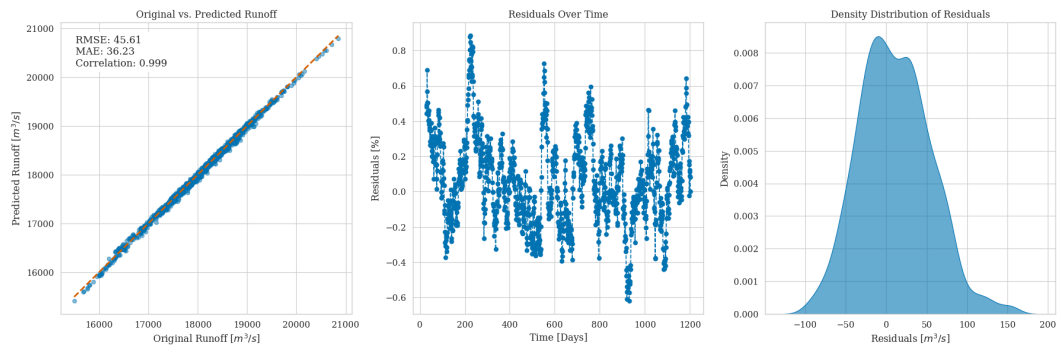
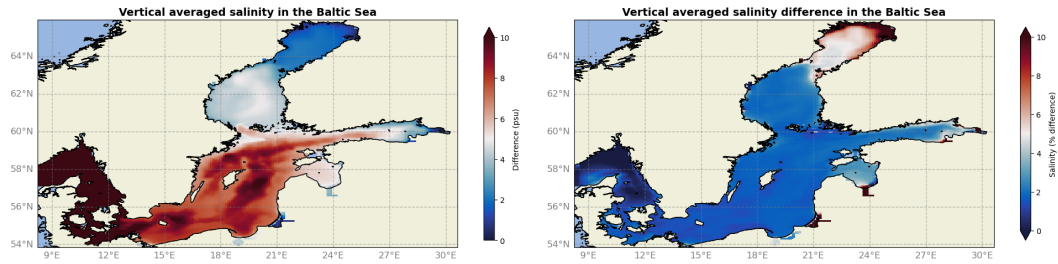


Figure 4: Figure ...



4 Acknowledgments

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5 Open research

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